gallon per minute (gal/min)

square mile (mi²)

ground-water flow systems in the Great Basin region of Nevada, Utah, and adjacent States

CONVERSION FACTORS Multiply By To obtain barrel (bbl) 0.1590 cubic meter barrel (bbl) 43.23 gallon gallon meters

0.06309

liter per second

square kilometer

INTRODUCTION

This map publication is one of several in a series concerning various aspects of the ground-water hydrology of the Great Basin in Nevada, Utah, and adjacent States. One report in the series describes the hydrogeologic framework of the Great Basin (Plume and Carlton, 1988). Another shows the ground-water levels for the aquifer systems of the Great Basin (Thomas and others, 1986). A third report in the series describes the regional ground-water flow patterns in the Great Basin (Harrill and others, 1988).

The Great Basin as discussed here includes about 140,000 mi², largely in Nevada and Utah but with smaller components in Arizona, California, Idaho, and Oregon (fig. 1). The study area, identical to the area described by Harrill and others (1988), is characterized by generally north-trending mountain ranges more than 50 mi long, separated by alluvial and fluviolacustrine basins. Most mountain ranges are 5 to 15 mi wide and rise 1,000 to 5,000 ft above adjoining basin floors. Many of the structural basins are topographically closed; however, others are interconnected by several river systems that terminate in major lakes or sinks. The area has had a complex geologic history that includes major episodes of sedimentation, igneous activity, orogenic deformation, and continental rifting.

The area includes several oil-producing fields in Nevada and Utah. Oil production in Nevada was approximately 3.4 million bbl in 1991 (Nevada Bureau of Mines and Geology, 1993). Several studies in recent years have shown a relation between ground-water movement and the occurrence of hydrocarbons. An early report by Hubbert (1953) and more recent reports by Tóth (1980, 1988) and by Garven (1989) indicate that ground-water flow may act as the driving force behind oil migration in certain areas. Garven describes conditions where regional flow in Upper Devonian carbonate rock and Lower Cretaceous sandstone in a western Canada sedimentary basin may have been responsible for the formation of an oil field containing more than 2,700 billion bbl of oil. In this conceptual model, ground water descends vertically from the recharge area. Soluble hydrocarbons are leached near deeply buried source rocks and transported by ground-water flow through regional aquifers. The hydrocarbons then accumulate in the discharge area of the flow system. Several key factors, such as the presence of heat sources and of forms of hydrocarbons capable of migration, must be in effect for this sequence of events to take place (Tóth, 1988, p. 485). Knowing that the occurrence and movement of hydrocarbons in some areas of the Great Basin can be due to regional ground-water flow could prove useful in future hydrocarbon explo-

The objective of the study described herein is to illustrate the possible relation between ground-water flow and the location of oil and natural gas within the regional flow systems of the Great Basin. The study area for this project is the Great Basin of Nevada and western Utah; however, much of the data collection was concentrated in Railroad Valley, in east-central Nevada. This valley, the most productive oil area of the State, yielded about 2.86 million bbl of oil in 1991 (Nevada Bureau of Mines and Geology, 1993).

This study included compiling all available data on hydrocarbon exploration and production wells drilled in the Great Basin and on hydrocarbon shows detected while drilling other wells. Garside and others (1988) compiled an extensive collection of hydrocarbon-related data for the State of Nevada. Other data were obtained in 1992 from State agencies (Nevada Bureau of Mines and Geology and Nevada Department of Minerals). Well data for Nevada were current as of April 1992. Well-location and well-status data for Utah were collected by Brady (1984) and were current as of 1982. Also collected was information on the locations of oil seeps in the study area and water wells that had an indication of hydrocarbons. The hydrocarbon data were integrated with hydrologic and geologic information to construct a map that shows the spatial relation of hydrocarbon exploration and production wells to regional and local ground-water flow systems. The question of whether ground-water flow is the motive force behind oil and gas migration in the Great Basin is not addressed further herein, but future directions of research to test this theory may become apparent.

HYDROGEOLOGIC UNITS

The Great Basin has had a complex geologic history resulting in a variety of rock types and structural features (Jennings, 1977; Walker, 1977; Bond, 1978; Stewart and Carlson, 1978; Hintze, 1980). Surficial geology in the area comprises three categories: basin fill characterized by mostly moderate to high permeability, consolidated rocks characterized by generally low permeability, and consolidated rocks characterized by moderate and locally high permeability.

The basin fill consists primarily of unconsolidated and semiconsolidated deposits of gravel, sand, silt, and clay derived from adjacent mountain ranges. Thickness of the basin fill generally ranges from 2,000 to 5,000 ft but exceeds 10,000 ft in the deeper basins. The basin fill has mostly moderate to high permeability. Information on the hydrologic properties of the deeper basin-fill deposits is limited, although the deeper deposits generally are assumed to have lower permeabilities and poorer water quality than the shallower deposits. In most basins, the uppermost 1,500 ft of basin fill forms the most productive aquifer.

The consolidated materials of generally low permeability include clastic sedimentary rocks, intrusive and extrusive igneous rocks, and metamorphic rocks, generally none of which transmit water readily unless extensively fractured. Basalt can have localized zones characterized by high permeability. Carbonate rocks, such as limestone and dolomite, generally have low permeability where they are unfractured or where fractures enlarged by solution are uncommon or not intercon-

Bedrock characterized by moderate to locally high permeability includes carbonate rocks within which secondary permeability has been developed, probably by solution enlargement of fractures. These rocks underlie a large area in western Utah and eastern Nevada known as the carbonate-rock province, where thick sequences of carbonate sediments were deposited in Paleozoic and early Mesozoic time. In this province, the Paleozoic and early Mesozoic sequence ranges in aggregate thickness from 20,000 to 30,000 ft and consists mostly of carbonate ocks. The carbonate-rock province is typified by complex interbasin systems of regional ground-water flow that include both basin-fill and carbonate-rock aquifers. Regional structural features that may account for significant local fracturing and resultant secondary permeability include overthrust zones, strike-slip shear zones, and low-angle fault zones (detachment surfaces). Locally, high-angle normal or listric faults may also cause significant fracturing and resultant permeability. The overall distribution of hydrogeologic units throughout the Great Basin is shown on Plume and Carlton's (1988) map, and the hydrogeology also was described by Dettinger and others (1990, 1995) and by Plume (1996).

GROUND-WATER FLOW SYSTEMS Large ground-water flow systems are driven by hydraulic gradients that are

continuous over long distances. These large systems are of either the regional or intermediate type (Harrill and others, 1988). A major flow system is considered to be local, intermediate, or regional depending upon what percentage of an area's ground water it conveys. Where consolidated rocks are permeable enough to allow hydraulic continuity on a regional scale, the local and intermediate flow systems are considered to be subsystems of the regional flow system. Where consolidated rocks have a low permeability and where no interbasin flow is evident from flow-system budgets, local and intermediate ground-water flow systems are thought to predominate.

Harrill and others (1988) delineated 39 flow systems in the Great Basin that range in area from 30 to $18,000~\text{mi}^2$. Of these, 16~are single-basin systems and 23 are multibasin systems. One of the multibasin systems includes as many as 34 hydrographic areas and subareas. Large, multibasin flow systems outside the carbonate-rock province are generally coincident with major river systems.

A good example of a regional ground-water flow system in the Great Basin is the White River system (figs. 1 and 2), described by Thomas (1988) and Dettinger (1989). Much of the ground water in the White River system is recharged in the Ely area of eastern Nevada (fig. 2). Additional recharge feeds the system near the center and also in the southern part. Ground water discharges at various locations throughout the flow system but primarily at several regional springs in White River and Pahranagat Valleys; some recharge is consumed by evapotranspiration in many parts of the White River system. Ground water ultimately discharges in Moapa Valley through springs along the Muddy River, in the southern part of the

In the Great Basin, deep ground water flows through the carbonate-rock aquifer and discharges through regional springs (Mifflin, 1968). These springs result from the combined influence of topography and geologic factors such as rock type and structure. Several regional springs in the White River Valley apparently are related to structural constriction of the carbonate-rock aquifer, which forces ground-water flow to the surface. Plume (1988) used aeromagnetic modeling to support this interpretation.

OIL AND NATURAL-GAS WELLS AND GROUND-WATER FLOW

In the White River flow system, several of the hydrocarbon test wells are just upgradient from some of the regional springs. Compared to the overall distribution of test wells in the White River flow system, hydrocarbon shows tend to be concentrated at these locations (fig. 1). This distribution may result from a restriction in the ground-water flow system that not only aids in the formation of the regional springs but may also act as a barrier to hydrocarbon migration. Plume's (1988) aeromagnetic modeling indicated that the ground water can be affected by geologic barriers, but no definitive proof exists for the possible constriction of hydrocarbon migration in this area.

1). At depth the Grant Range is composed primarily of crystalline basement rock, which acts as a barrier to easterly ground-water flow and possibly to easterly oil migration. Railroad Valley itself is the location of numerous springs that form a terminal discharge zone for the Railroad Valley regional flow system.

Several large, producing oil fields are on the east edge of Railroad Valley (fig.

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BOX ELDER CARSON CITY Table 1.—Hydrographic areas [Numbers and names correspond to those used by Harrill and others (1988) and are used to label the areas in figure 1] Continental Lake Valley Cactus Flat Gridley Lake Valley Stone Cabin Valley Virgin Valley Little Fish Lake Valley Swan Lake Valley Massacre Lake Valley Antelope Valley (Eureka and Nye) Stevens Basin Long Valley Mosquito Valley Diamond Valley Newark Valley Boulder Valley Little Smoky Valley Duck Lake Valley Hot Creek Valley Kawich Valley Smoke Creek Deser **Emigrant Valley** San Emidio Desert Frenchman Flat Granite Basin Indian Springs Valley Hualapai Flat High Rock Lake Valley Pahrump Valley Mud Meadow Mesquite Valley Summit Lake Valley Ivanpah Valley¹ Jean Lake Valley Black Rock Desert Pine Forest Valley Hidden Valley (South) Eldorado Valley Kings River Valley¹ Three Lakes Valley (Northern Part) Desert Valley Silver State Valley Tikapoo Valley Penoyer Valley Starr Valley Area Garden Valley Jakes Valley Lamoille Valley South Fork Area Long Valley **Huntington Valley** Ruby Valley Tenmile Creek Area Clover Valley Butte Valley¹ Elko Segment Susie Creek Area Steptoe Valley Cave Valley Maggie Creek Area Marys Creek Area Dry Lake Valley Pine Valley Delamar Valley Crescent Valley Lake Valley Spring Valley Carico Lake Valley Upper Reese River Valley Antelope Valley Antelope Valley Middle Reese River Valle Goshute Valley Lower Reese River Valley Independence Valley Whirlwind Valley Thousand Springs Valley¹ Pilot Creek Valley Rock Creek Valley Dry Valley Rose Valley Willow Creek Valley Eagle Valley Clovers Area Pumpernickel Valley Spring Valley Kelly Creek Area Patterson Valley Little Humboldt Valley Panaca Valley Clover Valley Hardscrabble Area Lower Meadow Valley Wash Paradise Valley Kane Springs Valley Winnemucca Segment **EXPLANATION** Grass Valley White River Valley Imlay Area Pahroc Valley Flow systems Lovelock Valley¹ Pahranagat Valley Coyote Spring Valley White Plains Humboldt River Three Lakes Valley (Southern Part) Bradys Hot Springs Area Las Vegas Valley Fernley Area Black Mountains Area Dixie Valley area Fireball Valley **EXPLANATION** Granite Springs Valley Garnet Valley Kumiva Valley Hidden Valley (North) Smith Creek Valley California Wash Winnemucca Lake Valley Basin fill (Quaternary to Tertiary)—Unconsolidated to consolidated Muddy River Springs Area Pyramid Lake Valley sedimentary deposits. Includes some tuff and interbedded lava flows Lower Moapa Valley South-central marsh area Dodge Flat Tule Desert Tracy Segment Consolidated rock (Tertiary and older) Warm Springs Area Virgin River Valley White River Spanish Springs Valley Mercury Valley Rock Valley Sun Valley Dil field—Includes many wells not shown individually Fortymile Canyon Truckee Meadows Railroad Valley Oasis Valley 36° — Pleasant Valley Crater Flat Washoe Valley ----- Boundary of study area—Coincides Boundary of carbonate-rock province Lemmon Valley¹ Amargosa Desert with Great Basin boundary Antelope Valley Chicago Valley Cold Spring Valley California Valley ————— Boundary of flow system—Dashed Lower Amargosa Valley Carson Desert¹ with hydrographic-area or -subarea boundary where uncertain Death Valley Churchill Valley Valiean Valley Dayton Valley Eagle Valley Shadow Valley Grouse Creek Valley Antelope Valley Pilot Valley table 1 and figure 2 Deep Creek Valley Smith Valley Snake Valley Mason Valley General direction of ground-water flow in basin-fill deposits East Walker Area Pine Valley Wah Wah Valley Walker Lake Valley Huntoon Valley Tule Valley Flow across hydrographic-area or -subarea boundary—Solid Teels Marsh Valley Fish Springs Flat Fish Lake Valley arrow indicates flow primarily through basin fill; dashed arrow Columbus Salt Marsh Valley indicates flow primarily through permeable consolidated rock Great Salt Lake Desert¹ Rhodes Salt Marsh Valley Tooele Valley Garfield Flat Soda Spring Valley¹ Rush Valley **Boundary of study area**—Coincides with Great Basin boundary Gabbs Valley Cedar Valley Rawhide Flats Utah Valley Area¹ Figure 2.—Index map, showing examples of ground-water flow systems in study area (modified from Harrill and others, 1983, Northern Juab Valley Fairview Valley Large spring—In Utah and most of Nevada, springs having discharge fig. 3): single-basin system (Smith Creek Valley flow system), multibasin system linked by saturated basin fill (Dixie Valley area and Stingaree Valley Salt Lake Valley generally greater than 1,000 gallons per minute; in more arid East Shore Area Cowkick Valley Humboldt River flow system), multibasin system linked by fracture zones (south-central marsh area flow system), and deep regional parts of Nevada and in California, regionally significant springs West Shore Area Eastgate Valley Area system with ground water flowing through both basin fill and adjacent and underlying permeable rock (White River and having discharges greater than 200 gallons per minute Skull Valley Railroad Valley flow systems). Buena Vista Valley Sink Valley Cache Valley Pleasant Valley Malad-Lower Bear River Area Buffalo Valley Producing oil well—Approximately located Pocatello Valley Jersey Valley SCALE 1:1,000,000 Blue Creek Valley Edwards Creek Valley Smith Creek Valley Hansel and Northern Rozel Flat Test wells 100 MILES Promontory Mountains Area **Curlew Valley** Monte Cristo Valley No oil or gas shows Big Smoky Valley1 Great Salt Lake 100 KILOMETERS Beryl-Enterprise Area Grass Valley Oil show Parowan Valley Kobeh Valley Monitor Valley Cedar City Valley Gas show Ralston Valley Beaver Valley Alkali Spring Valley Milford Area Oil and gas shows Leamington Canyon Clayton Valley Pavant Valley Lida Valley Sevier Desert Stonewall Flat Oil show in water well or spring—Approximately located Sarcobatus Flat **Figure 1.**—Distribution of oil and gas wells in relation to ground-water flow systems. Oil seep or hydrocarbon occurrence—Approximately located ¹Includes two or more subareas, which may be in different major flow systems.

Base from U.S. Geological Survey digital data, 1:100,000, 1987 Lambert Conformal Conic projection Standard parallels 33° and 45°, central meridian –117°

DISTRIBUTION OF OIL AND NATURAL-GAS WELLS IN RELATION TO GROUND-WATER FLOW SYSTEMS IN THE GREAT BASIN REGION OF NEVADA, UTAH, AND ADJACENT STATES

1996



Geology modified from Plume and Carlton (1988)

INTERIOR—GEOLOGICAL SURVEY, RESTON, VIRGINIA—1996