

Geologic Map of the Death Valley ground-water model area, Nevada and California

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Introduction

The purpose of this map is to provide the surface expression of the geology in the Death Valley ground-water model area to be incorporated initially into a 3-D geologic framework model and eventually into a transient ground-water flow model by the U.S. Geological Survey (D'Agnese, 2000; D'Agnese and Faunt, 1999; Faunt and others, 1999; and O'Brien and others, 1999). This work has been conducted in collaboration with the U.S. Department of Energy in order to assess regional ground water flow near the Nevada Test Site (NTS) and the potential radioactive waste repository at Yucca Mountain. The map is centered on the NTS and its perimeter encircles the entire boundary of the numerical flow model area, covering a total area of 57,000 km². The physiography, geology, and tectonics of the model area are extremely complex (Hunt and Mabey, 1966; Stewart, 1980; Jennings, 1994; Slate and others, 2000; Wright and others, 1999b). The northern and eastern part of the area includes typical Basin and Range topography consisting of north-trending block-faulted ranges and intervening valleys. The central part contains diverse ranges, plateaus, basins, and alluvial flats (for example, the NTS volcanic highlands and Amargosa Valley). The rugged ranges and deep basins of the Death Valley region in eastern California are characteristic of the topography of the southern and western parts of the map area. The map spans numerous tectonic subdivisions of the Great Basin. Deformation includes several generations of upper Paleozoic to Mesozoic thrust faulting that have been dismembered by extensive regional Tertiary to Quaternary normal and strike-slip faults. Much of this extensional and translational deformation is active today, with rates and amounts that vary from low to moderate in the central, eastern, and northern parts of the study area in southern Nevada, to very high in the southwestern and western parts in eastern California. For detailed discussion of the tectonic framework of the map area, the reader is referred to Workman and others (2002).

Methodology

This geologic map of the Death Valley ground-water model area involves the compilation of geologic data taken from many sources (figure 1). Cenozoic volcanic and sedimentary rocks are compiled and divided into major regional units based upon their isotopically determined time ranges combined with general petrographic characteristics evident in thin-section analyses. These units follow the mapping strategy of regional ash-flow tuff sheets in Ekren and others (1977), except in the Nevada Test Site area where more detailed Cenozoic stratigraphy was maintained

according to Slate and others (2000). Ash-flow stratigraphy is from Mackin (1960), Cook (1965), Williams (1967), Ekren and others (1971), Anderson and Rowley (1975), Best and others (1989a, b; 1993), Page and Dixon (1994), Sawyer and others (1994), Rowley and others (1994, 1995), Scott and others (1995a, b), and Warren and others (1998). Mesozoic, Paleozoic, and Proterozoic units and their thicknesses were compiled mainly from Slate and others (2000), Albers and Stewart (1972), Burchfiel and others (1983a, b), Page and others (in press), Cornwall (1972), Jennings (1961), Jennings and others (1962), Kleinhampl and

Ziony (1985), A. Jayko (unpublished mapping, 1990), Snow (1992), Strand (1967), Streitz and Stinson (1974), Wright and others (in press), and P. Guth and J. Yount (unpublished mapping, 1994). Unit thicknesses are not given for volcanic rocks because of the extreme lateral variability in thicknesses that are commonly observed in volcanic units within the map area.

In contrast to the bedrock units, the compilation of upper Tertiary to Quaternary surficial units is based on new interpretive mapping supported by reconnaissance field observations (figure 2). Surficial units were identified and compiled on rectified Landsat 5 Thematic Mapping (TM) image maps at 1:100,000-scale, and incorporated into the map and database. The TM data in these satellite image maps were specially processed to maximize the spectral discrimination of surficial deposits in arid to semiarid basins using procedures outlined in Menges and others (1999, 2000). Unit identification on the image maps was supplemented by (a) stereographic photo-interpretation of black and white high-altitude aerial photographs (nominal scale of 1:80,000), (b) calibration using field data collected at over 600 GPS-located stations in the map area (figure 2), and (c) variations in geomorphic characteristics derived from 30-m-DEMs using the techniques of Jayko and Pritchett (1999) and Jayko and Menges (2001). Surficial units on the map were provisionally correlated with upper Tertiary to Quaternary units identified in other studies conducted at varying scales and levels of detail in the map area. Although most of the surficial units are derived from new mapping, some existing regional-scale maps were adapted and incorporated (table 1). A number of more localized and detailed maps and reports were used for reference during map compilation in specific localities and (or) as general age control for the surficial units used in this map (Menges and others, 2001). Surficial mapping in the southern and western sections of the map were compiled at greater detail because the major Death Valley ground-water basin discharge occurs in these areas, and this part of the map contains a greater density of field control and more pre-existing detailed geologic maps and studies. Refer to Menges and others (2001) for more information on methodologies and correlations of surficial map units used in this study.

SECONDARY COMPILATION SOURCES

1.
McKee (1985)
Reheis (1992)
Ross (1967)
2.
Ekren and others (1971; 1973; 1976)
Gardner and others (1980)
Klinger and Piety (1996)
Klinger and Sarna-Wojcicki (2001)
Moring (1986)
3.
E. Ekren and P. Rowley
(unpublished mapping, 2000)
4.
Ekren and others (1977)
Guth (1980)
Longwell and others (1965)
Tschanz and Pampeyan (1970)
5.
Ekren and others (1977)
Tschanz and Pampeyan (1970)
6.
B. Burchfiel and G. Davis
(unpublished mapping, 1992)
8.
Ekren and others (1971; 1973; 1976)
9.
B. Burchfiel (unpublished mapping, 1971)
Carr and Pinkston (1987)
10.
Abolins (1999)
B. Burchfiel (unpublished mapping, 1971)
Burchfiel and others (1974; 1983a, b)
Lundstrom and others (1998; in press)
Maldonado and Schmidt (1990)
11.
Barnes and others (1982)
Burchfiel (1965)
Hay and others (1986)
Machette and others (1999)
Machette, Klinger, and Knott (2001)
Potter and others (in press)
Sarna-Wojcicki and others (2001)
12.
Klinger and Sarna-Wojcicki (2001)
McKee (1985)
Moring (1986)
Reheis (1992)
Ross (1967)
Sarna-Wojcicki and others (2001)
13.
Burchfiel (1969)
Hall (1971)
Klinger and Peity (1996)
Klinger and Sarna-Wojcicki (2001)
McAllister (1956)
Ross (1967)
Sarna-Wojcicki and others (2001)
Snow and Lux (1999)
Stone (1984)
14.
Abolins (1999)
Beratan and others (1999)
Brady (1986)
Brady and Troxel (1999)
B. Burchfiel and G. Davis
(unpublished mapping, 1992)
Burchfiel and others (1983a, b)
Butler and others (1988)
Calzia and others (2000)
Cemen and others (1999)
Denny (1965)
Denny and Drewes (1965)
Drewes (1963)
Hall and Stephens (1963)
Hay and others (1986)
Hillhouse (1987)
Hunt and Mabey (1966)
Klinger and Peity (1996)
Klinger and Sarna-Wojcicki (2001)
Knott (1998; 1999)
Machette, Klinger, and Knott (2001)
Machette, Menges, and others (2001)
McAllister (1970)
McKittrick (1988)
D. Miller and J. Yount
(unpublished mapping, 2000)
Morrison (1999)
Prave and McMackin (1999)
Sarna-Wojcicki and others (2001)
Snow and Lux (1999)
Stone (1984)
Taylor (1993)
Troxel (1986)
Troxel and Butler (1986a, b)
Wagner and Hsu (1987)
Wright and Troxel (1984; 1993)
Wright and others (1999a)

Table 1. Correlation table of upper Tertiary and Quaternary surficial units in the geologic map of the Death Valley ground-water model area and other stratigraphic sequences in and adjoining the Death Valley area. Only correlations with regional-scale mapping projects which were directly incorporated into the geologic map of the Death Valley ground-water model area are included.

[First column from left includes surficial units for Death Valley ground-water model area at regional (1:250,000) scale. The next five columns to right are Quaternary surficial units from five mapping projects both published and in progress that were incorporated into the Death Valley ground-water model area mapping program. Correlations on chart are based primarily on corresponding age ranges of units and only secondarily on genetic association (for example, alluvium vs. discharge deposits) to minimize effects of contrasting level of detail in various types of units among studies. Numbers in parentheses refer to minimum and maximum age estimates, in ka]

Geologic map of the Death Valley ground-water model area— 1:250,000 scale ¹			Geologic map of the Yucca Mountain region— 1:50,000 scale ²	Geologic map of the Nevada Test Site— 1:100,000 scale ³	Geologic map of the Indian Springs quadrangle— 1:100,000 scale ⁴	Geologic map of the Pahranaagat quadrangle— 1:100,000 scale ⁵	Geologic map of the Las Vegas quadrangle— 1:100,000 scale ⁶		
QTau, QTd, QTsf, QTls (0->758)	Qau, Qe	Qc, Qp (0-1)	Qay, Qayf (Qayfe) (0-18)	Qar (0-2)	Qay, Qey, Qed, Qp, Qps	Qya, Qyf, Qae, Qve, Qyl	Qs, Qsc	Qay, Qayy, Qayo, Qfy, Qfo, Qpy, Qd, Qsy, Qsyy, Qse	
		Qayo (5-18)		Qay, Qey (5-18)					
	Qao, Qlc (18-758)		Qam, Qem, Qsd (30-250)	QTm, QTc, QTu	Qai, Qeo	Qua, Qp, Qal, Qse, Qls	Qia, Qof, Qol	Qe, Qls, Qp, QTol	Qog, QTos?, QTs?
QTa (500->758)		Qao (500->758)		QTa		Qoa, Qof QToa	QTog, QTos, QTs,	QTa, Qb, Qbo, Qfw	

¹ Units are summarized in unit description text and in Table 2.

² Potter and others (in press): Geologic Map of Yucca Mountain region (Surficial units mapped by E. Taylor).

³ Slate and others (2000): Digital geologic map of the Nevada Test Site (revision 4, at 1:100,000 scale). Surficial units mapped by J. Slate, M. Berry, and V. Williams, based primarily on photo-interpretation on Beatty 1:100,000-scale quadrangle and the westernmost part of the Pahranaagat Range and Indian Springs 1:100,000-scale quadrangles. Surficial deposits on the Pahute Mesa 1:100,000-scale quadrangle adapted from mapping of W. Swadley.

⁴ P.L. Guth and J.C. Yount, unpublished mapping: Geologic map of Indian Springs 1:100,000-scale quadrangle (Quaternary and upper Tertiary surficial units mapped by J. Yount).

⁵ A.S. Jayko, unpublished mapping: Geologic map of Pahranaagat 1:100,000-scale quadrangle (Quaternary and upper Tertiary surficial units included in mapping).

⁶ Page and others (in press): Geologic map of the Las Vegas 1:100,000-scale quadrangle (Quaternary and upper Tertiary surficial units mapped by S.C. Lundstrom and J.B. Workman).

Description of Map Units

Surficial Deposits

The surficial units in this map include some modifications from the standard types of surficial-deposit units encountered in most geologic maps. These modifications are designed to emphasize factors of potential importance to the hydrogeologic framework model for the Death Valley ground-water model, as well as infiltration-recharge and discharge elements of the model. This includes not only differentiation of specific units, such as modern channels, discharge deposits, and playas, that directly impact these factors, but also identification of generalized textural variations in deposits of potential hydrologic significance. That is, variation in grain size was approximated by differentiating where possible (a) fine-grained deposits on distal alluvial fans and alluvial plains in basin interiors from coarse-grained alluvial deposits in proximal or medial piedmont positions, and (b) fine-grained Quaternary to Tertiary basin sediments from coarse-grained alluvium of similar age range. The latter fine-grained basin sediments typically include Pleistocene to upper Tertiary basin-fill deposits of mixed lacustrine, alluvial, and (or) paludal origin exposed in some dissected basins (Hillhouse, 1987; Morrison, 1999), but locally may delineate primarily lacustrine deposits in a few closed basins (Klinger and Sarna-Wojcicki, 2001). Alluvial units are further subdivided according to more standard age categories that reflect other potentially important hydrologic characteristics such as degree of soil and pavement development, cementation, and amount of internal dissection. Colluvial deposits are not included on this map because (a) they are difficult to systematically identify on the image maps and aerial photography, (b) these deposits are commonly thin and discontinuous, especially at map scale, and (c) they typically occur in range blocks and upland areas where they would tend to obscure the hydrogeologically more significant lithology and structure of subjacent bedrock.

Specific characteristics of all units, including physiographic position, dissection, surface and weathering characteristics, and selected soil and sedimentologic properties, are summarized in table 2. General ranges of unit thickness are estimated below, but are highly variable and poorly constrained due to a lack of basal exposures of most surficial deposits in the area. The typical range of unit thickness given first applies to most surficial deposits located in basins within the central and eastern parts of the map with low to moderate rates of tectonic activity and basin subsidence in localized areas. A larger maximum range is given for some deposits to reflect their likely greater thicknesses adjacent to the fault-bounded margins of basins such as central Death Valley in the western part of the map with high rates of tectonic activity and basin subsidence.

Qc Channel alluvium (Holocene)—Typically unconsolidated gravel, sand, and (or) silt; generally associated with zones of active to very recent surface flow and deposition. Identified on primary map only for selected major drainages that are differentiable at

map scale. No significant weathering or modification of the original depositional surface. May locally include braided channel systems with small inclusions of **Qay**, **Qayf**, **Qayo**, and (or) **Qao** alluvium. Thickness varies from 0 to 30 m, commonly from 1 to 10 m

Qp Playa and (or) salt pan deposits (Holocene)—Variably consolidated mixtures of silt, clay, and fine sand. May contain varying proportions of secondary carbonate, gypsum, and (or) salts in very arid basins. Generally restricted to sites of active deposition in the interior of closed to poorly drained basins. No significant surface modification or soil development, although salt crusts of variable hardness and thickness may develop on surface. Deposit thickness is poorly constrained, but may vary from 1 to 10 m for deposits of this age; commonly underlain by poorly constrained and potentially thicker accumulations of older playa or lacustrine deposits of middle to early Holocene and Pleistocene age

Qay Young alluvium (Holocene to latest Pleistocene?)—Mostly medium- to coarse-grained alluvium consisting of boulder to pebble gravel and sand, but may locally include fine-grained deposits of sand and silt. Typically unconsolidated to poorly consolidated. Commonly located in the proximal and distal positions of alluvial fans flanking basins, but also may include gravel- and sand-dominated axial drainages and fluvial terraces. Typically has no or weak soil development, no to moderate varnish development on clasts, and no to moderately modified surfaces characterized by prominent to indistinct relict depositional landforms (for example, bar and swale) with at most localized and incipient surface pavement development. Typically low surface relief related mainly to primary depositional features with at most minor dissection, although latter may increase adjacent to incised major drainages. Variable thickness commonly ranging from 0 to 20 m; possibly as much as 30 m adjacent to tectonically active mountain fronts

Qayf Young fine-grained alluvium (Holocene to latest Pleistocene?)—Mostly sand and (or) silt, locally with minor pebble gravel. Unconsolidated to poorly consolidated. Mostly associated with alluvial plains along the axial drainages of basin interiors, but may include fine-grained distal fan margins where clearly distinguishable from coarser-grained facies on aerial photography or image maps. Typically gentle surface gradients and minimal local relief. Has soil and surface characteristics similar to unit **Qay**, with some variation related to topographic position and finer-grained character (for example, minimal bar and swale development and fine-textured soil characteristics). Variable thickness from 0 to 10 m; possibly as much as 30 m in the interiors of tectonically active subsiding basins

- Qayfe** **Evaporite surface crust of salts and (or) carbonate (Holocene to latest Pleistocene?)**—Localized areas of thin evaporative surface crust of salts and (or) carbonate overlying unit **Qayf**
- Qayo** **Intermediate-age alluvium (middle Holocene to latest Pleistocene?)**—Mostly medium- to coarse-grained gravel (boulder to pebble) and sand deposits in proximal and medial basin positions. Poorly consolidated to unconsolidated. Generally has weakly developed soils, weak to moderate varnish development on clasts, and weak to moderate surface modification characterized by no to incipient pavement development with strong to weak relict bar and swale landforms. Local relief similar to that of unit **Qay**. Unit mapped where sufficiently large older elements of unit **Qay** may be differentiated on images or aerial photography and (or) are identified in field or on existing maps; otherwise unit is included within **Qay**. Variable thickness ranges from 0 to 10 m, with possibly as much as 20 m adjacent to tectonically active mountain fronts
- Qe** **Eolian deposits (Holocene to middle Pleistocene)**
Generally loose, but locally consolidated sand and silt, with local concentrations of gravel on stabilized surfaces commonly associated with buried paleosols. Typically consists of Holocene sand sheets or dune fields sufficiently thick and continuous to be mapped and (or) large individual dunes. May also locally include large relict late to mid-Pleistocene sand ramp deposits that are banked along the flanks of some ranges and may contain one or more variably developed calcic to petrocalcic relict or buried soils. Generally restricted to areas with sufficient sand accumulation to form discrete eolian landforms. Thickness of most sand sheets and dune fields vary from 0 to 10 m, but deposits in individual large dunes and sand ramps may increase locally from 30 to 200 m (for example, northern Amargosa Desert and southeastern Eureka Valley)
- Qau** **Undifferentiated alluvium (Holocene to Pleistocene)**—Mainly used to delineate areas with intermixed **Qay**, **Qayf**, and (or) **Qao** deposits, with at most minor inclusions of **QTa** that cannot be depicted individually at map scale. Generally consists of various alluvial units of recent to middle Pleistocene age that individually comprise from 20% to 80% of area, with at most 5-10% inclusions of unit **QTa**. Commonly used on piedmonts with complex suites of alluvial deposits, or along drainages with narrow suites of terraces or inset surfaces, within that age range
- QTau** **Undifferentiated older alluvium (Holocene to latest Tertiary)**—Mainly used to delineate areas of mixed **Qay**, **Qayf**, **Qayo**, **Qao**, **QTa** or **QTsf** units that are too small or too complexly intermingled to depict individually at map scale. Commonly identified where complex suites of deposits of widely mixed ages are located in proximal piedmonts or range-front embayments, interiors of dissected basins, or small narrow intramontane valleys or basins. In contrast to unit **Qau**, older **QTa** and **QTsf** comprise a significant component of unit **QTau** (commonly >10-20% of unit area). Also used to indicate areas where one or more of the Quaternary or Quaternary and Tertiary deposits listed above are complexly intermingled with dissected tracts of unconsolidated to poorly consolidated Tertiary sediments in patterns too complex or small to discriminate individually at map scale
- QTd** **Deposits associated with modern or past groundwater discharge (Holocene to late Tertiary)**—Typically fine-grained deposits of fine sand, silt, and mud, but may locally include dense crystalline deposits of limestone or travertine. Variably indurated, from unconsolidated to cemented. Generally associated with variable amounts of organic material, phreatophytic or wetland fauna, and (or) secondary precipitates of carbonate or silica commonly derived from ground-water. Commonly consists of fossil deposits (latest Pleistocene or older in age) that formed in paludal to lacustrine environments associated with formerly expanded sites of groundwater paleodischarge that are entirely or largely inactive under current climatic conditions, although unit locally is associated with areas of active seepage or spring discharge. May include older fine-grained (inactive) playa or lacustrine deposits locally. Thickness of unit variable, depending on age range of exposed units. Commonly from 0 to 10 m for exposed deposits, but possibly as much as 100 m or more in subsurface or in some dissected basin interiors (for example, southwestern Pahrump Valley; Page and others, in press)
- QTls** **Landslide block (Holocene to late Tertiary)**—Generally brecciated but coherent (commonly re cemented) bedrock outcrop, typically emplaced above or interstratified within surficial materials or upper basin-fill sediments in basin piedmonts or along range fronts. Thickness variable and poorly determined, as base of slide blocks generally are not exposed and thus basal depths must be estimated mainly from geophysical data. Exposed thicknesses commonly from 10 to 100 m, but may range to more than 500 m in subsurface
- QTsf** **Old alluvial, paludal, or lacustrine sediments (Pleistocene to late Tertiary)**—Typically fine-grained fluvial, paludal, or lacustrine basin-fill sediments consisting of mixtures of poorly to well stratified silt, clay and fine sand. Variable induration, from unconsolidated to well consolidated. May contain local gravel-dominated channels and interbeds, paleodischarge deposits, interbedded tuffs, and (or) buried soils. Deposits are typically exposed in bluffs, ridges, or badland areas in dissected basin interiors (for example, beds of Lake Tecopa Allotment of Morrison, 1999) or in uplifted fault blocks

Table 2. Summary of selected surficial deposit units mapped in Death Valley ground-water model area.

[Table summarizes characteristics of the most common and distinctive surficial-deposit units in the Death Valley groundwater model mapping program subdivided into primarily alluvial units and other types of deposits. Mixed undifferentiated units (Qau and QTau) are not included because they consist of various proportions of other units listed below and thus lack a consistent set of characteristics. QTls, the bedrock landslide unit, is not included, because it resembles bedrock on the images and aerial photography and in the field consists of brecciated bedrock outcrop. Qayfe is not listed separately, as it is a variant of Qayf distinguished only by a thin crust of carbonate or salt on the surface of the deposit (see unit descriptions). The first three columns list map unit, general character, and age range estimated from correlations with specific local- to regional-scale studies with independent age control in and adjacent to the map area (listed in Menges and others, 2001). The next two columns list characteristics of units observed in the satellite image maps and black and white aerial photography used to interpret and compile units for map. The remainder of the columns list characteristics of deposits and related geomorphic surfaces that are primarily based on observations from field traverses. Characteristics are summarized from full area of Death Valley ground-water model, and thus vary significantly, reflecting the wide range of geologic, tectonic, and local climatic conditions across the region. N/A indicates not applicable.]

Unit ¹	Type ²	Age ³	Image/Photo Characteristics ⁴		Physiographic Position ⁵	Internal Dissection ⁶	Surface Characteristics ⁷		Varnish ⁸	General Soil Characteristics ⁹	Typical Grain Size ¹⁰	Cementation ¹¹
			Image Maps	Black & White Aerial Photography			Pavement	Bar and Swale				
Qc	Channel alluvium	Recent	Variable, light colored to white; light blue where salt encrusted	White to light gray	Intrarange valleys, piedmonts, and axial basin drainages	None to minor thalweg incision	None	Unmodified primary	None (only inherited)	None	Variable, from mixed gravel to fine sand	None (except local gully bed)
Qay	Medium - to coarse-grained alluvium	Recent to Holocene, locally latest Pleistocene	Light to dark colors, including browns; rough to medium textures	Light to medium gray	Proximal to distal piedmonts, terraces, coarse axial drainages	None to minor (mostly primary)	None to weak, locally moderate	Unmodified to faint (but discernable)	None to moderate	None to weak (cambic and Stage I calcic horizons)	Coarse to fine gravel and sand	None
Qayo	Medium - to coarse-grained alluvium	Middle to early Holocene, locally latest Pleistocene	Medium to dark colors, including browns; medium textures	Medium to dark grays	Proximal to distal piedmonts, terraces	None to minor	Weak to locally moderate	Faint, but discernable	Weak to moderate	Weak (cambic and Stage I calcic horizons)	Coarse to fine gravel and sand	None
Qayf	Fine-grained alluvium	Recent to Holocene, locally latest Pleistocene	Light to medium colors, including brown; light blue where salt encrusted; smooth to medium textures	White and light to medium grays	Mostly fine-grained axial drainages and interior alluvial plains, some distal piedmonts	None to very minor	None to weak, locally moderate	None to faint	None to moderate	None to weak (cambic and Stage I carbonate horizons)	Sand, silt, minor fine gravel	None
Qao	Medium - to coarse-grained alluvium	Pleistocene	Light to dark colors, including brown to reddish brown; smooth to medium textures	Light to medium grays	Proximal to medial, and locally distal, piedmonts; terraces in intrarange valleys	Commonly moderate to strong; locally minor	Strong to moderate; present on interfluvial where dissected	None	Strong to moderate	Strong to moderate; includes argillic and (or) Stage II-IV calcic to petrocalcic horizons	Coarse to fine gravel and sand	Poor to moderate
QTa	Medium - to coarse-grained alluvium	Middle Pleistocene to Pliocene	Light to medium colors, possible light-brown tint; smooth to medium textures	Light to medium grays	Proximal to locally medial piedmonts; range-front embayments; intrarange valleys	Commonly strong and deep (balena topography)	None to weak, commonly with carbonate chips	None	None to weakly preserved	Strong to moderate; variably eroded Stage III-V petrocalcic where preserved	Coarse to fine gravel and sand	Moderate to strong

Table 2—Continued. Summary of selected surficial deposit units mapped in Death Valley ground-water model area.

[Table summarizes characteristics of the most common and distinctive surficial-deposit units in the Death Valley groundwater model mapping program subdivided into primarily alluvial units and other types of deposits. Mixed undifferentiated units (Qau and QTau) are not included because they consist of various proportions of other units listed below and thus lack a consistent set of characteristics. QTIs, the bedrock landslide unit, is not included, because it resembles bedrock on the images and aerial photography and in the field consists of brecciated bedrock outcrop. Qayfe is not listed separately, as it is a variant of Qayf distinguished only by a thin crust of carbonate or salt on the surface of the deposit (see unit descriptions). The first three columns list map unit, general character, and age range estimated from correlations with specific local- to regional-scale studies with independent age control in and adjacent to the map area (listed in Menges and others, 2001). The next two columns list characteristics of units observed in the satellite image maps and black and white aerial photography used to interpret and compile units for map. The remainder of the columns list characteristics of deposits and related geomorphic surfaces that are primarily based on observations from field traverses. Characteristics are summarized from full area of Death Valley ground-water model, and thus vary significantly, reflecting the wide range of geologic, tectonic, and local climatic conditions across the region. N/A indicates not applicable.]

Qlc	Medium to coarse-grained lacustrine deposits	Late to middle Pleistocene	Light to dark colors, including brown to reddish brown; smooth to medium textures	Light to medium grays	Mostly medial to distal piedmonts	Commonly moderate to strong; locally minor	Strong to moderate, but present on interfluvial where dissected	None	Strong to moderate	Strong to moderate; includes argillic and (or) Stage II-IV calcic to petrocalcic horizons	Medium- to fine gravel and sand (moderately to well sorted)	Poor to moderate
QTsf	Fine-grained lacustrine, inactive playa, and alluvial deposits	Pleistocene to Pliocene	White to light colors; medium to smooth textures	White to light grays	Variable, including proximal to distal fans and basin interiors	Moderate to strong, in dissected basins; locally none	N/A	N/A	N/A	Local buried paleosols, variable weak to strong development	Silt, clay, fine sand; local gravel interbeds	Poor to moderate
Qp	Active playa and salt flat deposits	Recent to late Holocene	White to variegated light and pastel colors; light blue where moist; smooth textures	White to light grays	Basin interiors	None	None	None	None	None; local near-surface encrustations on salt pans	Clay and silt; local fine sand	None to moderate (on salt pans)
QTd	Inactive and active discharge deposits	Recent to Pliocene	White to variegated light and pastel colors; light blue where moist; smooth to medium textures; local green vegetation	White to light grays	Mostly distal piedmonts and basin interiors; locally proximal to medial piedmonts	None to moderate	None; locally weak to moderate on alluvial caps	N/A; none to weak (on alluvial caps)	None to weak (on alluvial caps)	None to moderate (may include St I-II calcic horizons)	Clay, silt, fine sand; local fine gravels and organic mats	None to moderate; locally strong where crystalline precipitate
Qe	Eolian deposits	Holocene to middle Pleistocene	Light brown, pinkish gray to pink; medium textures	Light grays	Mostly basin interiors; locally variable piedmont	None	None	N/A	None	None	Fine to medium sand; silt	None

¹ Labels used to identify surficial units for Death Valley ground-water model area geologic map. Units are subdivided into alluvial and other (that is, lacustrine, playa/salt flat, eolian, ground-water discharge). Excluded from table are cone point or mixed age-range alluvial units (Qau and QTau) and the bedrock landslide unit (QTIs) for reasons explained in table header.

² General type or classification of deposit, including general size range of coarse fraction, such that medium- to coarse-grained refers to gravel and sand and fine-grained refers to sand, silt and clay.

³ Age range of deposits, based on correlations to local or regional studies with geochronologic control listed in Menges and others (2001).

⁴ Typical characteristics of unit observed in satellite image maps and aerial photography used to interpret and compile maps. Satellite image refers to the specially processed Landsat 5 Thematic Mapper image projected onto 1:100,000-scale quadrangles for compilation of units and Black & White aerial photography refers to high-altitude black and white aerial photography that were used as aid to interpretive mapping and for field observations, as described in text.

⁵ Typical physiographic positions of unit within landscape, commonly given in reference to alluvial basins and (or) bedrock ranges.

⁶ Relative amount of local dissection or incision of drainages within or along margins of given mapped area of deposit. Common depth ranges include 0-2 m (minor), 2-5 m (moderate), and > 5 m (strong).

⁷ Degree of desert varnish observed at surface, given qualitatively in reference to darkening of tops and (or) reddening of undersides of surface clasts.

⁸ Degree of desert varnish observed at surface, given qualitatively in reference to darkening of tops and (or) reddening of undersides of surface clasts.

⁹ General strength of maximum soil development associated with upper surface of deposits, typically as observed in natural outcrops (gullies) or road cuts, or as exposed in shallow pits. Soil pits exposing the full depth range of deep profiles could not be excavated during the reconnaissance field traverses required for this project. Observations focused on relative degree of development of horizons of secondary accumulation of carbonate, after Machette (1985), and locally salt or silica, and presence or absence of cambic or argillic horizons. In some cases, generalized soil characteristics were inferred less reliably where soil was not exposed on undissected surfaces from features on surface, including carbonate chips, carbonate rinds on clasts, and materials brought up from burrowing insects or rodents.

¹⁰ Typical textural size range of deposits, given for gravels in terms of coarse (cobble to boulder) or fine (pebble to cobble) grained.

¹¹ Degree of internal cementation or induration throughout deposits, below surface zone of soil development. Bonding agent varies from fines (silt, clay) in uncemented to poorly cemented deposits to secondary carbonate, salt, or silica cement in moderately to strongly cemented materials.

(beds of Confidence Hills Formation of Beratan and others, 1999), although unit locally includes fine-grained primarily lacustrine deposits as young as late to middle Pleistocene related to pluvial lakes within enclosed basins (for example, northern Death Valley; Klinger and Sarna-Wojcicki, 2001). Variable thickness from 0 to greater than 100 m in exposed sections. Total thickness in basin subsurface may exceed 300 m

Qao Old alluvium (late to middle Pleistocene)—Typically medium- to coarse-grained gravel (boulder to pebble gravel dominated) and sand deposits in proximal and medial basin positions, but may locally include regions with fine-grained sand and silt at surface in some closed-basin settings. Variable induration but generally moderately consolidated to unconsolidated below potentially cemented well-developed soils. Has moderately to well-developed soils, moderate to strong varnish development on clasts, and moderate- to well-developed surface pavements with no detectable bar-and-swale landforms. Variable local relief, ranging from undissected to well incised, depending on local base-level conditions, but typically with at least some local incision; unit typically contains at least some varnished pavement surfaces preserved on flat to broadly rounded interfluves even where strongly dissected. Thickness of individually mapped units at surface typically from 0 to 30 m, but may increase to 50-100 m along tectonically active range-fronts. Aggregate thickness of all Pleistocene-age alluvial units adjacent to tectonically active range fronts may exceed 500 m in subsurface

Qlc Old lacustrine deposits (late to middle Pleistocene)—Generally fine- to medium-grained gravel (mostly pebble to cobble, some small boulder) and sand, related primarily to pluvial lakes formed in closed basins such as central Death Valley (Machette, Klinger, and Knott, 2001). May locally include fine-grained sand and silt beds and (or) variably cemented tufa deposits. Generally moderately consolidated to unconsolidated beneath surface soils, except where associated with cemented densely crystalline tufa. Has moderate- to well-developed soils, moderate to strong varnish development on surface clasts, and moderate- to well-developed surface pavements with no detectable bar and swale landforms. Variable local relief, ranging from undissected to deeply dissected, depending on local base-level conditions, but typically with at least some local incision; unit typically contains at least some varnished pavement surfaces preserved on flat to broadly rounded interfluves even where strongly dissected. Unit similar in many respects to unit **Qao**, but is differentiable based on generally finer gravel texture, better sorting and rounding of coarse fraction, and common association with lacustrine depositional features and landforms (for example,

shorelines; spits and bars; foreset, backset, or deltaic stratification). Thickness variable, but typically from 0 to 20 m; locally as much as 50 m

QTa Oldest alluvium (middle Pleistocene to late Tertiary)—Generally medium- to coarse-grained gravel (boulder to pebble) and sand. Variable induration, from poorly consolidated beneath cemented surface soils to well cemented throughout. Mostly associated with old dissected fans and (or) gravel-sand dominated basin fill deposits, commonly located at basin margins. Typically has strongly developed or eroded relict soils (usually dominated by variably degraded cemented petrocalcic to petrosalic horizons). Generally strongly incised into ballena (ridge-ravine) topography with little or no preservation of pavements or original surfaces on narrow interfluves. Thickness generally from 0 to 100 m, but may increase to 300 m locally. Generally not preserved at surface on basin side of tectonically active range-front faults

Bedrock Units

Within the Tertiary section, we have followed the mapping strategy of Ekren and others (1977). Sedimentary and volcanic units are subdivided based upon age range. Within each age range, the rocks are further subdivided into lava flows, tuffs, and associated sedimentary rocks. These units represent four major stages in the Tertiary tectonic evolution of the region. Unit 1 consists of sedimentary rocks that predate the oldest Tertiary volcanic units associated with extension. These rocks fall into regionally extensive formations which have been significantly deformed by more recent extension and which everywhere lack volcanic material. Unit 2 consists of the oldest Tertiary volcanic sequence ranging in age from 27 to 32 Ma. These rocks are commonly difficult to differentiate in the field and in thin section and consequently source associations are poorly constrained. No sedimentary rocks associated with unit 2 are exposed within the map area. Unit 3 consists of the younger sequence of calc-alkaline volcanic rocks and associated sedimentary units ranging in age from 18 to 27 Ma. Unit 4 consists of the entire bimodal sequence of volcanic rocks and associated sedimentary rocks ranging in age from 5 to 17 Ma. Within the Nevada Test Site and Death Valley regions, these units are subdivided into a more detailed stratigraphic sequence representing both the Death Valley volcanic field and southwestern Nevada volcanic field. Sedimentary rocks of unit 4 are not subdivided due to a lack of consistent, regional stratigraphic correlations. Felsic-composition lava flows and intrusive rocks are not subdivided due to lack of age data.

Silicic to Intermediate-composition Volcanic Rocks

Qa Young intermediate-composition lava flows, undivided (Pleistocene)—Partially reworked, intermediate-composition lava flows, pyroclastic rocks and

interbedded sedimentary rocks exposed in the eastern-most Saline Range and the northern tip of the Cottonwood Mountains (Strand, 1967; Burchfiel, 1969; Streitz and Stinson, 1974)

Tv Volcanic rocks, undivided (Pliocene to Oligocene)—Volcanic rocks whose age and stratigraphic position are poorly constrained

Tr Felsic-composition lava flows, undivided (Pliocene to Oligocene)—Mostly resistant, mostly gray, white, and tan, rhyolitic lava flows and volcanic domes, and subordinate ash-flow and air-fall tuff. Includes many rhyolite and lesser trachyte and comendite masses, many of them associated with calderas. Includes the Oddie Rhyolite of Esmeralda County (Albers and Stewart, 1972)

Ta4 Intermediate-composition lava flows, unit 4 (Pliocene and Miocene)—Mostly resistant, intermediate-composition lava flows, flow breccia, and subordinate volcanic mudflow breccia. Includes the andesite of Sarcobatus Flat west of the Nevada Test Site (Slate and others, 2000), 9-12 Ma dacitic rocks in Death Valley and the Grapevine Mountains (Wright and others, in press), and 5.9-Ma porphyritic latite or trachyandesite and rhyolitic to andesitic volcanic breccia in Esmeralda County (Albers and Stewart, 1972). Age range: 5-17 Ma

Tt4 Ash-flow tuffs and interbedded air-fall tuffs, unit 4 (Pliocene and Miocene)—Mostly resistant, poorly to highly welded, mostly silicic ash-flow tuffs and minor air-fall tuff. Includes the peralkaline Kane Wash Tuff (14.4-14.7 Ma) and older (15-16 Ma) peralkaline tuffs formerly included in the Kane Wash Tuff (Scott and others, 1995a). Where not mapped separately, includes tuffs of the Timber Mountain Group. Includes the upper nonwelded ash-flow tuff (6 Ma) and the upper welded ash-flow tuff (5.9-6.1 Ma) of Albers and Stewart (1972) in Esmeralda County. Age range: 5-17 Ma

Rocks of the southern Death Valley region

(Miocene)—Lava flows and tuffs exposed in southern Death Valley and areas to the south and east including the Black Mountains, Greenwater Valley, Greenwater Range, southern Amargosa Desert, southern Panamint Range, Owlshead Mountains, Sperry Hills, and northern Kingston Range. These rocks are described in detail by Wright and others (in press). All dates taken directly from Wright and others (in press). Divided into:

Tvg Greenwater Volcanics—Rhyodacite flows and air-fall tuffs exposed in the northern Greenwater Range and Black Mountains. Age Range: 5.4-6.0 Ma

Tar Intermediate- to felsic-composition lava flows of the central Death Valley volcanic field—Andesitic to rhyolitic lava flows and interbedded tuffs in the Black Mountains, Greenwater Range, and Panamint Mountains where unit includes Rhodes Tuff,

Sheephead Andesite, and Shoshone Volcanics (8-11.5 Ma). Includes intermediate- to felsic-composition hypabyssal intrusive rocks in Black Mountains and Greenwater Range. Also includes dacitic lava flows in the southern Black Mountains and southern Resting Spring Range. Age range: 8-12 Ma

Tas Intermediate- to felsic-composition lava flows of the southern Death Valley-Kingston Range volcanic belt—Basaltic andesite to andesite lava flows and associated breccias and pyroclastic deposits in the southern Panamint Range, Owlshead Mountains, Avawatz Range, and north end of California Valley and north and east of the Saddle Peak Hills. Age Range: 11-14 Ma

Rocks of the Nevada Test Site and vicinity

(Miocene)—Ash-flow tuffs and interbedded air-fall tuffs exposed in and near the Nevada Test Site that are described in detail by Slate and others (2000). All dates taken directly from Slate and others (2000). Divided into:

Tst Stonewall Flat Tuff and tuffs of the Thirsty Canyon Group, undivided—Peralkaline ash-flow tuffs and associated rocks. The Stonewall Flat Tuff was derived from the Stonewall Mountain caldera at about 7.5-7.6 Ma. The Thirsty Canyon Group was derived from the Black Mountain caldera at about 9.15-9.4 Ma

Tmt Tuffs of the Timber Mountain Group—Metaluminous ash-flow tuffs and associated rocks derived from the Timber Mountain caldera complex at 11.4-11.6 Ma

Tpt Tuffs of the Paintbrush Group—Metaluminous ash-flow tuffs and associated rocks derived from the Claim Canyon caldera at 12.7-12.8 Ma. Locally includes the Kane Wash Tuff east of the Nevada Test Site

Tw Wahmonie and Salyer Formations, undivided—Intermediate-composition lava flows, flow breccia, volcanic mudflow breccia, and source intrusions of the Wahmonie and Salyer Formations

Tct Tuffs of the Crater Flat Group—Calc-alkaline ash-flow tuffs and associated rocks of 13.1-13.5 Ma exposed in, and derived from an unknown caldera, perhaps the Silent Canyon caldera complex

Tbt Tuffs of the Belted Range Group—Peralkaline ash-flow tuffs and associated rocks derived from the Silent Canyon caldera complex at 13.5-13.85 Ma

Tlt Lithic Ridge Tuff, rhyolite of Picture Rock, Tunnel Formation, and comendite of Quartet Dome, undivided—Rhyodacitic, rhyolitic, and comendite ash-flow tuff, lava flows, and associated rocks of 14.0-14.3 Ma. Sources known or inferred to be in or near the Nevada Test Site

Tqv Volcanic rocks of Quartz Mountain—Rhyolite to quartz trachyte lava flows, ash-flow tuff, and air-fall tuff of 14.2-14.9 Ma derived probably from the Mount Helen caldera

- Tuv Older ash-flow tuffs and lava sequences, undivided**—Peralkaline and rhyolite ash-flow tuff, lava flows, and associated rocks of about 14.9-16 Ma. Consists of the comendite of Ocher Ridge, Tub Spring Tuff, comendite of Emigrant Valley, older tunnel beds, tuff of Yucca Flat, Redrock Valley Tuff, tuff of Twin Peaks, and volcanic rocks of Gold Flat. Comendite units are from local sources but the sources of the tuffs are unknown, although some could be from the Nevada Test Site
- Tkv Lava flows of the Belted Range**—Rhyolite and latite lava flow sequences. Consists of the rhyolite of Belted Peak, rhyolite of Wheelbarrow Peak, and latite of Kawich Valley
- Ta3 Intermediate-composition lava flows, unit 3 (Miocene and Oligocene)**—Mostly resistant, intermediate-composition lava flows, flow breccia, and subordinate volcanic mudflow breccia. Includes the Mizpah Trachyte (20.4 Ma), the Milltown Andesite (21.5 Ma), and 20.8-21.6 Ma dacitic rocks and the Chispa Andesite in the northern part of the area (Albers and Stewart, 1972; Cornwall, 1972; Kleinhampl and Ziony, 1985). Age range: 19-26 Ma
- Tt3 Ash-flow tuffs and interbedded air-fall tuffs, unit 3 (Miocene and Oligocene)**—Mostly resistant, poorly to highly welded, mostly silicic ash-flow tuffs and minor air-fall tuff. Includes the Hiko Tuff (18.2-18.6 Ma; Rowley and others, 1995), Fraction Tuff (18.7-20.5 Ma; Best and others, 1989b), tuff of Saulsbury Wash (21.6 Ma; Kleinhampl and Ziony, 1985), Harmony Hills Tuff (22.0-22.5 Ma; Rowley and others, 1995), Pahrnagat Formation (22.39 Ma; Best and others, 1995), Condor Canyon Formation (22.8-23.7 Ma; Rowley and others, 1995), Leach Canyon Formation (23.8 Ma; Best and others, 1989b; Rowley and others, 1994), Bates Mountain Tuff (23-24 Ma; Best and others, 1989b), tuff of White Blotch Spring (24-25 Ma; Ekren and others, 1977), tuffs of Antelope Springs (Cornwall, 1972), rocks of the Reveille Range (24-25 Ma; Kleinhampl and Ziony, 1985), tuff of Kiln Canyon (24-25 Ma (Kleinhampl and Ziony, 1985), tuff of the Quinn Canyon Range, Kendall Tuff (Albers and Stewart, 1972), Vindicator Rhyolite (Albers and Stewart, 1972), tuff of Murphy Gap, tuff of Bald Mountain (about 25 Ma; Best and others, 1989b), Shingle Pass Tuff (26.0-26.7 Ma; Best and others, 1989b), tuff of Hancock Summit (about 26 Ma; Best and others, 1989b), and Isom Formation (25.7-27.0 Ma; Best and others, 1989b; Rowley and others, 1994). Also includes poorly dated (but less than 25 Ma) young tuffs of the Pancake caldera complex, and poorly dated tuffs of the Hannapah and Ellendale mining districts, including the tuff of McKinney Tanks and the tuff of Hannapah (Kleinhampl and Ziony, 1985). Age range: 18-27 Ma
- Ta2 Intermediate-composition lava flows, unit 2 (Oligocene)**—Mostly resistant, intermediate-composition lava flows, flow breccia, and subordinate volcanic mudflow breccia. Age range: 27-32 Ma
- Tt2 Ash-flow tuffs and interbedded air-fall tuffs, unit 2 (Oligocene)**—Mostly resistant, poorly to highly welded, mostly silicic ash-flow tuffs and minor air-fall tuff. Includes the Monotony Tuff (27.3 Ma; Best and others, 1989b), tuff of the Seaman Range (Ekren and others, 1977), Lund Formation (27.9 Ma), Wah Wah Springs Formation (29.5 Ma), and Cottonwood Wash Tuff (30.6 Ma) of the Needles Range Group (Best and others, 1989a), tuff of Williams Ridge and Morey Peak (31.3 Ma; Best and others, 1989b), Windous Butte Formation (31.4 Ma; Best and others, 1989b), and the Stone Cabin Tuff (35.4 Ma; Best and others, 1989b). Age range: 27-32 Ma

Basaltic Volcanic Rocks

- Qb Young basaltic lava flows (Holocene to Pleistocene)**—Mostly resistant, basaltic lava flows, flow breccia, scoria, and cinder cones
- QTb Basaltic lava flows, undivided (Pleistocene to Miocene)**—Mapped where age is less certain. Includes Malpais Basalt and Rabbit Spring Formation, undivided, of Esmeralda County (Albers and Stewart, 1972) and Pliocene? basaltic rocks in the Reveille Range (Cornwall, 1972; Kleinhampl and Ziony, 1985)
- Tb Old basaltic lava flows (Pliocene to Miocene)**—Mostly resistant, basaltic lava flows, flow breccia, scoria, and cinder cones. Includes the Mira Basalt (Albers and Stewart, 1972)

Intrusive Rocks

- Tgy Young granitic intrusive rocks (Miocene)**—Mostly resistant silicic intrusions. Includes intrusive rocks on Timber Mountain associated with the tuffs of the Timber Mountain Group (11.4-11.6 Ma; Slate and others, 2000) and rocks in the Death Valley region (8.5-12.5 Ma; Wright and others, in press)
- Tai Younger intermediate-composition intrusive rocks (Miocene to Oligocene)**—Mostly resistant plugs and dikes in the Panamint Range and Owlshhead Mountains, which are correlative with a 12.5 Ma sill in the Kingston Range (Wright and others, in press). In the central and northern sections of the map area, this unit includes exposures associated with caldera magmatism and shallow plutonism ranging in age from 11-32 Ma (Ekren and others, 1971; Cornwall, 1972; Ekren and others, 1977; Kleinhampl and Ziony, 1985; Slate and others, 2000). Locally includes associated extrusive rocks

- Tws Gabbro to diorite intrusive rocks of Willow Springs (Miocene)**—Mostly resistant and mafic intrusion of about 11.6 Ma in the Black Mountains associated with the Willow Springs pluton (Wright and others, in press)
- Tgo Older granitic intrusive rocks (Miocene? to Oligocene)**—Mostly resistant silicic intrusions. Includes intrusive quartz monzonitic rocks in the northern Quinn Canyon Range (26.8 Ma; Kleinhampl and Ziony, 1985) and coarse-grained rhyolite porphyry in the Cactus Range which intrudes the tuff of White Blotch Spring (Cornwall, 1972)
- TKi Older intermediate-composition intrusive rocks (Oligocene to Cretaceous)**—Mostly resistant intrusions and dikes whose age is poorly constrained. Includes “dacite to gabbro dikes, many too small to show at map scale, along the west side of Bare Mountain (Slate and others, 2000)”
- TKd Mafic intrusive rocks (Oligocene to Cretaceous)**—Mostly resistant mafic intrusions and dikes. Exposed in the Palmetto and Montezuma ranges of Esmeralda County (Albers and Stewart, 1972)
- Trg Oldest granitic intrusive rocks (Tertiary? to Triassic)**—Mostly resistant, mostly silicic intrusions and dikes. Ages generally decrease eastward across the map area from predominantly Jurassic in the western ranges to predominantly Cretaceous in the eastern ranges. Includes exposures along the eastern Cottonwood Mountains of known Triassic age (McAllister, 1956)
- JFqm Avawatz Mountains quartz monzodiorite, undivided (Early Jurassic to Late Triassic)**—Coarse-grained quartz monzodioritic plutonic rocks underlying most of the Avawatz Range, and exposed as fault slices near the Garlock fault and in the Slate Range and southern Panamint Range. K/Ar and Rb/Sr ages fall within the range of 180 to 220 Ma (Wright and others, in press)

Sedimentary and Metamorphic Rocks

- Ts Sedimentary rocks, undivided (Pliocene to Eocene)**—Poorly to well consolidated fluvial and lacustrine sedimentary rocks whose age is poorly constrained. Includes the Siebert Tuff in the Tonopah area (Albers and Stewart, 1972). Locally divided into:
- Ts4 Sedimentary rocks, unit 4 (Pliocene and Miocene)**—Poorly to moderately consolidated, mostly fluvial sandstone, conglomerate, mudstone, limestone, and siltstone. The upper part of the unit is composed of deposits in local basins that reflect the present topography formed during basin-range extension. These deposits include unnamed basin-fill sediments, the Funeral Formation, the Furnace Creek Formation, caldera moat-filling sediments,

bedded tuff and sedimentary rocks, and sedimentary rocks and landslide deposits exposed in and near the Nevada Test Site (Slate and others, 2000; Wright and others, in press). Includes the sedimentary deposits (7-14 Ma) of the Shadow Valley basin (Wright and others, in press). In Esmeralda County, the upper part of the unit includes the Pozo Formation, deposits that overlie tuff dated at 5.9 Ma, and other (4-17 Ma) deposits (Albers and Stewart, 1972). The lower part of the unit includes pre-basin-range deposits filling local basins that predate the present topography. On and near the Nevada Test Site, these rocks, previously called the rocks of Pavits Spring, are of about 11.7-18.3(?) Ma (Slate and others, 2000). In Esmeralda County, the lower part of the unit contains the Esmeralda Formation. Age range: 5-17 Ma

- Tls Landslide and mega-breccia deposits (Miocene to Oligocene)**—Large, angular blocks of older bedrock ranging in size up to hundreds of meters, commonly retaining original sedimentary bedding, contained within a recemented matrix of identical lithology. Generally underlain by mappable, gravity-induced faults related to Tertiary extension (Slate and others, 2000). Includes masses in the Timpanahute Range inferred to post-date 24 Ma (Ekren and others, 1977), in the Sheep Range where deposition was synchronous with deposition of the Horse Springs Formation (Guth, 1981; P. Guth and J. Yount, unpublished mapping, 1994), in the Reveille Range dated at 18.5-26 Ma (Ekren and other, 1973), in the Beatty region dated at 10-12.7 Ma (Slate and others, 2000), in the Nevada Test Site region overlying volcanic rocks dated at 13.7 to 15.3 Ma and overlain by 11.6 Ma volcanic rocks (Slate and others, 2000), at the northern end of the Spring Mountains dated at 12.7-13.7 Ma (Slate and others, 2000), and in the Blue Diamond area where the age is unknown (Page and others, in press). Unit is everywhere assumed to be older than 10 Ma

- Tso Sedimentary rocks older than the rocks of the Nevada Test Site and vicinity (Miocene to Eocene?)**—Moderately consolidated, fluvial and lacustrine sandstone, conglomerate, limestone, mudstone, and siltstone whose age is poorly constrained. Includes conglomerate with interbedded lacustrine limestone, siltstone, sandstone and a 30.2 Ma air-fall tuff (Barnes and others, 1982), formerly called the rocks of Winapi Wash on the northern side of the Spotted Range. Includes the upper tuffaceous sandstone and lower conglomerate of the Titus Canyon Formation along the north side of the Funeral and Grapevine Mountains (Slate and others, 2000). Also includes a small exposure north of Yucca Flat composed of locally derived clasts of the Eleana Formation which separates the Paleozoic basement rock from overlying Tertiary volcanic units beneath

much of the northern end of Yucca Flat as interpreted from drill-hole data (J. Cole, oral communication, 2001)

- Ts3** **Sedimentary rocks, unit 3 (Miocene and Oligocene)**—Moderately to well consolidated, fluvial and lacustrine sandstone, limestone, conglomerate, and mudstone. Deposited in basins that formed before the present topography was developed. Includes the Sandstorm Formation of Esmeralda County (Albers and Stewart, 1972), the Horse Spring Formation in the southern Las Vegas Range (Page and others, in press), and deposits younger than the Titus Canyon Formation in Death Valley National Park (Wright and others, in press). Age range: 18-26 Ma
- Ts1** **Sedimentary rocks, unit 1 (Eocene and Paleocene)**—Coarse, predominantly conglomeratic sedimentary rocks which pre-date Tertiary volcanic activity based on a lack of volcanic clasts and stratigraphic relationships. Includes the Sheep Pass Formation in Lincoln County (Tschanz and Pampeyan, 1970; Ekren and others, 1977) and the conglomerate of Ekren and others (1971) in the Mount Helen area
- Mzsv** **Sedimentary and volcanic rocks, undivided (Mesozoic)**—Exposed throughout the southwestern part of the map area including the southern Inyo Mountains, Slate Range, southern Panamint Range, and Avawatz Mountains. This unit is locally metamorphosed and includes interbedded sedimentary and volcanic rocks (Wright and others, in press). Includes the Butte Valley Formation in the southern Panamint Mountains which is age-equivalent to the Moenkopi Formation in the eastern part of the map area (Johnson, 1957)
- Klw** **Lavinia Wash sequence (Upper and Lower Cretaceous)**—Heterogeneous synorogenic sequence of interbedded conglomerate, sandstone, and mudstone defined by Carr (1980) in exposure in the Goodsprings mining district. Underlain by an inferred unconformity (Carr and Pinkston, 1987). Unit originally assigned Late Jurassic to Early Cretaceous(?) age based on conventional K-Ar dates of volcanic rocks within unit, but new laser fusion $^{40}\text{Ar}/^{39}\text{Ar}$ methods yielded a late Early Cretaceous age (99.0±0.4 Ma) (Fleck and Carr, 1990; Fleck and others, 1994); this date constrains the maximum age of movement on the Keystone thrust in the eastern Spring Mountains. Unit from 5 to 20 m thick
- Ja** **Aztec Sandstone (Jurassic)**—Exposed in eastern Spring Mountains. Tan and red quartz sandstone. Conspicuous 5-10-meter scale cross-bedding. Unit is truncated by the Keystone thrust fault. Thickness ranges from 650 to 800 meters (Longwell and others, 1965; Page and others, in press)
- Tc** **Chinle Formation (Upper Triassic)**—Exposed in eastern Spring Mountains. Interbedded sandstone, siltstone, shale, and minor conglomerate and limestone beds. Gypsum locally in discontinuous beds and irregular lenses. Base of formation is Shinarump Member, composed of conglomerate with distinctive reddish-brown and yellowish-brown chert clasts, and minor carbonate clasts. Unit conformably underlies Aztec Sandstone (Ja) and unconformably overlies Moenkopi Formation (Longwell and others, 1965; Stewart, 1980; Page and others, in press). Unit is 180 m thick
- Tm** **Moenkopi Formation (Middle? and Lower Triassic)**—Exposed in eastern Spring Mountains. In descending order, unit includes upper red member, Virgin Limestone Member, lower red member, and Timpoweap Member. Upper red member (Burchfiel and others, 1974) consists of claystone, mudstone, dolomite, limestone, siltstone, shale, and some gypsum. Virgin Limestone Member consists of bioclastic limestone, sandy limestone, dolomite, calcareous sandstone, and siltstone; fossils include brachiopods, gastropods, pelecypods, and pelmatozoan fragments; limestone locally contains ooids. Lower red member consists of siltstone, sandstone, shale, and minor dolomitic limestone, silty limestone, and gypsiferous siltstone beds. Timpoweap Member marks base of formation and is limestone and chert-clast conglomerate with minor pebbly carbonate and calcareous sandstone beds; grades upward into lower redbed sequence. Unit is 585 m thick (Page and others, in press)
- Pzu** **Sedimentary rocks, undivided (Paleozoic)**—Unit mapped in Cottonwood Mountains, Hunter Mountain area, Grapevine Mountains north of Scotty's Castle, and the Reveille Range. Consists of undivided marine sedimentary rocks (Jennings, 1961; Jennings and others, 1962; Strand, 1967; Cornwall, 1972; Streitz and Stinson, 1974)
- Pkt** **Kaibab and Toroweap Formations, undivided (Lower Permian)**—Exposed in eastern Spring Mountains. Harrisburg and Fossil Mountain Members of the Kaibab Formation are recognized (Carr, 1992; Carr and McDonnell-Canan, 1992). Harrisburg Member is micritic limestone and dolomite, bedded gypsum, and some red shale and claystone; pectinoid valves and pelmatozoan ossicles observed; common chert in nodules and layers. Fossil Mountain Member is thick-bedded fossiliferous limestone and dolomitic limestone with brown-weathering chert layers and nodules; includes brachiopods, bryozoans, and pelmatozoan columnals. Woods Ranch, Brady Canyon, and Seligman Members of the Toroweap are recognized (Carr, 1992; Carr and McDonnell-Canan, 1992); Woods Ranch Member is shale, dolomite, siltstone, and some gypsum. Brady Canyon Member is limestone and dolomitic limestone with common chert nodules and layers. Seligman Member is siltstone, sandstone, and gypsum; some dolomite and dolomitic sandstone in upper part. Unit forms ledgy to massive

cliffs and is overlain unconformably by Timpoweap Member of the Moenkopi Formation; in places, unit is completely removed along this unconformity, and Timpoweap Member rests directly on Permian redbeds. Combined thickness of the two formations is 0-300 m in the eastern Spring Mountains (Page and others, in press)

Pr Redbeds (Lower Permian)—Exposed in eastern Spring Mountains. Consists of sandstone and some interbedded shale and siltstone; locally gypsiferous. Conformably overlies the Bird Spring Formation; contact with overlying Toroweap Formation is sharp and disconformable. Unit forms ledgy slope and is about 580 m thick (Longwell and others, 1965; Carr and Pinkston, 1987; Page and others, in press)

Pov Owens Valley Group (Lower Permian)—Exposures in the Dry Mountain area consist of the Lone Pine Formation, composed of thin-bedded calcareous mudstone, shale, and siltstone. Exposures in the Cottonwood Mountains and Panamint Range consist of the Darwin Canyon Formation, composed of well-bedded very fine grained sandstone, calcarenite sandstone, and subordinate calcarenitic, bioclastic limestone, and limestone conglomerate underlain by the Osborne Canyon Formation, composed of calcareous mudstone, bioclastic limestone, limestone conglomerate, and subordinate calcareous siltstone (Burchfiel, 1969; Stone, 1984; Stone and Stevens, 1987; Stone and others, 1987). The Osborne Canyon Formation and underlying Tihvipah Limestone have been incorrectly mapped as the Keeler Canyon Formation by Streit and Stinson (1974) within the map area (Stone and others, 1987). Hall (1971) reported 730 m of unit exposed in Panamint Range area but Stone (1984) remapped much of this unit as Bird Spring Formation, and Burchfiel (1969) reported 180 m of the unit in the Dry Mountain area.

PSu Sedimentary rocks, undivided (Permian to Silurian)—Exposed in southwestern part of map area. In Grapevine Mountains, Awawatz Mountains, and Slate Range consists of rocks correlative with parts of the Owens Valley Group, Keeler Canyon Formation, Perdido Group, Tin Mountain Limestone, Lost Burro Formation, and Hidden Valley Dolomite from top to bottom. In Funeral Mountains and northern Panamint Range consists of rocks correlative with parts of the Owens Valley Group, Bird Spring Formation, Indian Springs Formation, Santa Rosa Hills Limestone, Stone Canyon Limestone, Tin Mountain Limestone, Lost Burro Formation, and Hidden Valley Dolomite from top to bottom. About 2,370 m thick

PIPmb Bird Spring (Lower Permian and Pennsylvanian) and Indian Springs Formations (Upper Mississippian)—Exposed in eastern part of the map area in the Spring Mountains, Las Vegas Range, Sheep Range, Pahrangat Range, and Timpahute Range

and in the western part of the map area in the Cottonwood Mountains and Panamint Range (included in **PSu** in this area). Bird Spring Formation consists of limestone, silty limestone, dolomite, chert, quartzite, siltstone, and shale. Common layers and nodules of brown-weathering chert. Contains abundant macrofossils including brachiopods, ostracods, colonial and solitary corals, bryozoans, sponges, pelmatozoan fragments, and fusulinids. In Las Vegas Range, Sheep Range, and part of the Spring Mountains, unit includes a distinctive carbonate slope to basin marker unit with submarine debris-flow conglomerates—marker unit is about 1,000 m above the base of the formation in the region. The underlying Indian Springs Formation of Webster and Lane (1967) consists of interbedded limestone, shale, and quartzite (Page and others, in press). Unit is from 1,300 to 2,355 m thick in the Spring Mountains; maximum thickness of 2,500 m in the Las Vegas Range; top not exposed (Page and others, in press)

PIPt Tippipah Limestone (Lower Permian and Pennsylvanian)—Exposed in Nevada Test Site area. Consists of thin-bedded, fossiliferous limestone with chert nodules, calcareous siltstone to sandstone, and minor chert-pebble conglomerate, sandy limestone, and quartz sandstone. The clastic component increases up section. Top of unit unconformably overlain by younger Mesozoic and Cenozoic units. Unit unconformably overlies the Scotty Wash Quartzite (mapped as **Msc**) which is completely removed in some locations. Correlative with part of the Bird Spring Formation (**PIPmb**). Maximum thickness about 1,250 m (Tschanz and Pampeyan, 1970; Kleinhampl and Ziony, 1985; Slate and others, 2000)

PIPkc Keeler Canyon Formation (Lower Permian to Middle Pennsylvanian)—Exposed in the Dry Mountain area. Composed of silty, bioclastic limestone and siltstone and a basal micritic limestone with abundant chert nodules; fossils include fusulinids, pelmatozoans, bryozoans, and brachiopods. Unit is differentiated from the partly age-equivalent Bird Spring Formation by an abundance of turbidite facies. Conformably overlain by the Lone Pine Formation and conformably underlain by the Rest Spring Shale. (Ross, 1967; Streit and Stinson, 1974; Stone and Stevens, 1987). Unit 1,164 m thick in Dry Mountain area (Burchfiel, 1969)

Mu Sedimentary rocks, undivided (Upper and Lower Mississippian)—Mapped mostly in Cottonwood Mountains and Last Chance Range, Inyo County, California, and also in the Spotted Range, Lincoln and Clark Counties, Nevada. In Spotted Range, includes, from top to base, Indian Springs Formation of Webster and Lane (1967), Limestone of Timpi Canyon, Mercury Limestone, and Narrow Canyon Limestone (Poole and others, 1961; Stevens

and others, 1996). Lower Mississippian rocks are correlative with parts of the Monte Cristo Group and Joana Limestone. Indian Springs Formation is mostly of variegated shale, but also has bioclastic limestone and quartzite beds; fossils include the biostratigraphically diagnostic Chesterian brachiopod *Rhipidomella nevadensis*. Limestone of Timpi Canyon, and the Mercury and Narrow Canyon Limestones, consist of limestone and dolomite, silty limestone, sandstone and quartzite, and shale; fossils include gastropods, pelmatozoan fragments, and brachiopods. Unit about 230 m thick (Longwell and others, 1965; Stevens and others, 1996).

Exposures in the Cottonwood Mountains are equivalent to the Monte Cristo Group as described below (Stone, 1984). In the Last Chance Range, unit includes Rest Spring Shale, Kearsarge Formation, Perdido Group, and Tin Mountain Limestone. Rest Spring Shale is thin-bedded, micritic limestone and black fissile shale; limestone is bioclastic and phosphatic (Stevens and others, 1996). Mexican Spring Formation of Perdido Group is mostly siltstone and conglomerate; siltstone is bioturbated and contains turbidite beds with graded bioclasts (brachiopods, gastropods, foraminifera); the underlying Leaning Rock Formation is thin-bedded, micritic, argillaceous limestone, and minor mudstone and conglomerate. Tin Mountain Limestone is limestone with distinctive pale-reddish-purple partings. Entire assemblage is 400 m thick in the Dry Mountain area of the Cottonwood Mountains (Burchfiel, 1969), about 140 m thick in the Panamint Range (Hall, 1971), and as great as 420 m thick in the Quartz Spring area, Inyo County (Stevens and others, 1996).

Msc Scotty Wash Quartzite (Upper Mississippian) and Chainman Shale (Upper and Lower Mississippian), undivided—Exposed in the Timpahute Range, Pahrnagat Range, and eastern Nevada Test Site area (Syncline Ridge, CP Hills, and Calico Hills). Scotty Wash Quartzite consists of tabular to lensoidal quartzite beds and fossiliferous limestone. Chainman Shale is black pyritic shale with sparse siltstone partings interbedded with thicker siltstone, sandstone, and bioclastic limestone beds. Forms strike valleys. East of the Nevada Test Site area, unit is poorly exposed, and thickness is estimated at more than 1,200 m (Trexler and others, 1996); unit is about 300 m thick in the Spotted and Pahrnagat Ranges (Tschanz and Pampeyan, 1970).

Mm Monte Cristo Group (Upper and Lower Mississippian)—Exposed in the Nopah Range, Montgomery Mountains, Spring Mountains, southern Las Vegas Range, southern Sheep Range, and in the Indian Springs area, south of U.S. Highway 95. Includes following formations, from top to base: Yellowpine Limestone, Bullion Limestone, Anchor Limestone, and Dawn Limestone. Formations consist of

bioclastic limestone with pelmatozoan fragments, solitary and colonial corals, and brachiopods, and minor dolomite and chert beds. Yellowpine contains unusually large solitary rugose corals; also contains *Lithostrotionella* colonial corals, and pelmatozoan ossicles. Encrinitic limestone (grainstone to wackestone) typifies the Bullion Limestone. Anchor Limestone consists of alternating thin-bedded limestone and chert beds. Dawn Limestone is abundantly fossiliferous and has common dark-brown-weathering chert beds. Unit is from 300 to 480 m thick (Burchfiel and others, 1983a, b; Page and others, in press).

MDe Eleana Formation (Mississippian and Upper Devonian)—Exposed in Nevada Test Site area.

Described by Poole and others (1961); reinterpreted by Trexler and others (1996). Sandstone, pebble conglomerate, siliceous siltstone, and minor bioclastic limestone and bedded chert. Unit varies laterally and vertically and was deposited in deep-water turbiditic trough and submarine fan depositional environments. Upper 150 m consists of interbedded bioclastic limestone and chert-litharenite sandstone. Middle 1,700 m consists of chert-litharenite sandstone, conglomerate, and siltstone. Lower 170 m consists of debris-flow conglomerate deposits, silty and sandy limestone, and quartzite. Top of unit is everywhere truncated by the Belted Range thrust (Cole and Cashman, 1999). Unit maximum thickness is 2,040 m.

Mj Joana Limestone (Lower Mississippian)—Exposed in northeastern part of map area in the Pahrnagat and Timpahute Ranges. Limestone with interbedded silty limestone, shale, and chert; upper part weathers to lighter shades of gray and contains more chert and shale beds; lower part generally forms massive limestone cliffs; abundant pelmatozoan fragments and corals (Tschanz and Pampeyan, 1970). Correlative with lower part of Monte Cristo Group (**Mm**), and lower part of Mississippian undivided unit (**Mu**). About 200 to 300 m thick.

MDu Sedimentary rocks, undivided (Lower Mississippian and Upper and Middle Devonian)—Exposed in eastern part of the map area in Lincoln and Clark Counties and parts of eastern Nye County. In Keystone thrust plate (eastern Spring Mountains) includes the Sultan Limestone. In western thrust plates of Spring Mountains and majority of Lincoln and Clark Counties, includes Guilmette Formation and either the Crystal Pass Member of the Sultan Limestone or the Pilot Shale to the south or north of the Las Vegas Valley shear zone, respectively. Unit is partly correlative with the Devils Gate Limestone as mapped by Burchfiel (1965) and Burchfiel and others (1974; 1983a, b) in the Specter Range, Spring Mountains, Montgomery Mountains, and Nopah and Resting Spring Ranges. Guilmette Formation is dolomite, limestone, and dolomitic

quartzite. Limestone and dolomite in upper part of the formation are generally micritic, and beds are cyclic, laminated, and burrow mottled. Dolomitic quartzite beds are in upper part of formation. Middle part of formation is mostly cliff-forming dolomite with abundant stromatoporoid biostromes and patch reefs; also contains *Amphipora* sp., gastropods, and brachiopods. Lower part of formation forms ledgy slope beneath cliff-forming dolomite and contains thin-bedded limestone with stromatoporoids alternating with yellowish-gray weathering dolomicrite. Lower part is “yellow bed” of Guilmette (Tschanz and Pampeyan, 1970) which marks the base of the formation in much of eastern Nevada. Map unit is 380-520 m thick in the Spring Mountains, 630 m thick in the southern Sheep Range, 730-990 m thick in the Pahrnagat Range, 670-750 m thick in the Spotted and Desert Ranges, and 350 m thick in the Nevada Test Site area (Slate and others, 2000). Sultan Limestone (Hewett, 1931) consists, from top to base, of the Crystal Pass, Valentine, and Ironside Members. Crystal Pass Member is partly Mississippian and is light-gray-weathering micritic limestone (Page and others, in press). Valentine and Ironside Members are thin- to thick-bedded limestone and dolomite with silicified stromatoporoids. Sultan Limestone is 175-200 m thick (Page and others, in press)

D1b Lost Burro Formation (Upper and Middle Devonian)—Exposed in Cottonwood Mountains and Panamint Range. Uppermost 10 m of unit includes calcareous, laminated to thin-bedded sandstone and sandy dolomite (Quartz Spring Sandstone Member). Remainder of unit consists of dolomite and limestone, sandy dolomite, quartzite and sandstone beds; dolomite and limestone beds contain *Amphipora* sp. and silicified stromatoporoids. Includes quartzitic basal Lippincott Member. Included within map unit **PSu** in Panamint Range and southeastern Funeral Mountains. Correlative with parts of the Guilmette Formation and Sultan Limestone. Unit about 500 m thick in the Dry Mountain area (Burchfiel, 1969) and 457 m thick in the Panamint Range (Hall, 1971)

D1c Dolomite and limestone, undifferentiated (Devonian to Upper Cambrian)—Exposed only in northern Nye County, in the Reveille and Quinn Canyon Ranges. Includes, from top to base, rocks correlative with parts of the Guilmette Formation, Simonson Dolomite, Sevy Dolomite, Laketown Dolomite, Ely Springs Dolomite, Eureka Quartzite, and Pogonip Group; thickness unknown

Dsf Slope-facies carbonate rocks (Upper to Lower Devonian)—Exposed at Bare Mountain, Calico Hills, Ranier Mesa, and Carbonate Wash areas of Nevada Test Site. Unit consists of limestone, dolomite, and silty carbonate rocks. Contains submarine debris-flow conglomerate; clasts composed of quartzite,

limestone, dolomite, and chert. Maximum thickness is 300 m (Slate and others, 2000)

Ds Simonson Dolomite (Middle Devonian)—Exposed in eastern part of the map area in Lincoln, Clark, and parts of Nye Counties. Includes mostly finely crystalline dolomite; contains abundant stromatoporoids and *Amphipora* sp.; in some locations (northwestern Spring Mountains, Pahrnagat Range), upper beds of the Simonson contain the Givetian brachiopod *Strigocephalus* sp. In much of Lincoln and Clark Counties, base of formation is marked by distinctive quartzite unit which unconformably overlies the Sevy Dolomite. Unit is partly correlative with the Nevada Formation as mapped by Burchfiel and others (1974) in the Spring Mountains, and Burchfiel and others (1983a, b) in the Montgomery Mountains and Nopah Range. Unit is 120 m thick in northwest Spring Mountains, 190 m thick in the southern Sheep Range, 365 m thick in the Pahrnagat Range, and 300 m thick in the Nevada Test Site area (Slate and others, 2000)

DSu Sedimentary rocks, undivided (Devonian and Silurian)—Mapped in Lee Canyon and Deer Creek thrust plates, central Spring Mountains (Page and others, in press). Unit includes strata below basal yellow-bed sequence of Guilmette Formation (Devonian) and above Ely Springs Dolomite (Upper Ordovician). These rocks are equivalent with parts of the Simonson Dolomite, Sevy Dolomite, and Laketown Dolomite. Unit is 60-80 m thick

D1cm Mountain Springs Formation (Middle Devonian, Lower Ordovician, and Upper Cambrian)—Mapped only in eastern Spring Mountains. Defined by Gans (1974) and subsequently modified by Miller and Zilinsky (1981) and Page and others (in press). Subdivided into three informal members from top to base (Page and others, in press): upper member restricted to Middle Devonian strata equivalent to part of the Simonson Dolomite, and consists of cycles of finely to coarsely saccharoidal dolomite and dolomudstone with scattered rip-up clasts. Contains pentamerid brachiopods, stromatoporoids, and more rarely, bryozoan, coral, and pelmatozoan fragments. Middle member is early Late Ordovician age, is equivalent to part of the Ely Springs Dolomite, and is burrow-mottled, bioclastic dolopackstone. Lower member is Early Ordovician in age and equivalent to part of the Pogonip Group. Member is burrowed, dolomitized wackestone, packstone, and mudstone, and contains gastropods (*Palliseria* sp. and *Maclurites* sp.) and sponges (*Receptaculites* sp.). Unit is 295 m thick

D1cu Mountain Springs (Middle Devonian, Lower Ordovician, and Upper Cambrian) and Nopah Formations (Upper Cambrian), undivided—Mapped only in eastern Spring Mountains in southeastern part of area. Unit is about 600 m thick

DSsl Sevy (Middle and Lower Devonian) and Laketown (Upper and Middle Silurian) Dolomites, undivided—In most parts of Lincoln and Clark Counties (Desert, Pintwater, Sheep, Spotted, Pahrnagat and Timpahute ranges, and Worthington Mountains), Sevy Dolomite is cliff-forming, light-gray, aphanitic dolomite containing thin beds of cross-bedded sandy dolomite and dolomitic quartzite; lacks macrofossils; top of Sevy marked by cherty argillaceous unit composed of siltstone and cherty aphanitic dolomite; this unit was originally defined as an informal member of the Sevy Dolomite (Osmond, 1962), but was later redefined as a separate map unit (Johnson and others, 1989); overlain unconformably by Simonson Dolomite. Combined Sevy Dolomite and cherty argillaceous unit is 275 m thick in the Nevada Test Site area (Slate and others, 2000), and its maximum thickness is nearly 500 m in the Pahrnagat Range (Reso, 1963).

Laketown Dolomite underlies Sevy Dolomite in Lincoln and Clark Counties where it displays typical tri-part character widely recognized in the eastern Great Basin—upper dark, middle light, and lower dark dolomite parts. Upper dark part is mostly finely saccharoidal dolomite containing brachiopods, halysitid corals, and pelmatozoan ossicles. Middle part is massive dolomite that contains silicified pentamerid brachiopods. Lower part is saccharoidal dolomite with brachiopods, corals, pelmatozoan ossicles, and stromatoporoids; also contains beds and nodules of dark-brown-weathering chert. Laketown Dolomite unconformably overlies Ely Springs Dolomite. The Laketown Dolomite is 250 m thick in the Nevada Test Site area, and reaches a maximum thickness of 285 m in the Pahrnagat Range

DSlm Lone Mountain Dolomite (Lower Devonian and Silurian)—Exposed only at Bare Mountain. Fine- to medium-crystalline dolomite; pelmatozoan fragments are common. Contact with underlying Roberts Mountain Formation gradational. Unit is about 490 m thick (Cornwall and Kleinhampl, 1961)

DShv Hidden Valley Dolomite (Lower Devonian and Silurian)—Mapped in the Cottonwood and Funeral Mountains; included with unit **PSu** in the Nopah and Panamint Ranges and Montgomery Mountains. Consists mostly of massive dolomite, but includes minor quartzite beds and brown-weathering chert at base in the Panamint Range area (Hall, 1971). Generally lacks macrofossils but age equivalent with parts of the Sevy and Laketown Dolomites based on fossils collected by McAllister (1952). Contact with underlying Ely Springs Dolomite is gradational. Unit is 440-460 m thick in southeastern Funeral Mountains (McAllister, 1974) and in the Dry Mountain area (Burchfiel, 1969), about 120 m thick in Panamint Range, and about 100 m thick in the Nopah Range (Burchfiel and others, 1983a, b)

Sr Roberts Mountains Formation (Silurian)—Exposed only at Bare Mountain. Dolomite and limestone, and minor interbedded silty and sandy dolomite, dolomite pebble-conglomerate, and dark-brown-weathering chert nodules and layers. Contains brachiopods, corals, graptolites, and conodonts of Late to Early Silurian age (Slate and others, 2000). Contact with underlying Ely Springs Dolomite is disconformity. Unit is 200 m thick (Monsen and others, 1992)

Opa Palmetto Formation (Upper, Middle, and Lower Ordovician)—Exposed in western part of map area in Palmetto Mountains, Magruder Mountain, and Silver Peak Range. Mostly shale and siltstone, but also includes thin-bedded chert, limestone, and quartzite; quartzite beds are more common in lower part. Rocks are commonly metamorphosed to hornfels. Shale contains Middle Ordovician graptolites (Albers and Stewart, 1972; McKee, 1985). Collections by Albers and Stewart (1972) from the southern Silver Peak Range yielded Late Ordovician fossils and from the Goldfield area yielded Early Ordovician fossils. Unit is deep-water facies (eugeosynclinal) and interpreted as allochthonous with respect to the Roberts Mountain thrust (Albers and Stewart, 1972; McKee, 1985); upper and lower unit contacts are nearly everywhere represented by faults; maximum thickness estimated at about 600 m

O€u Sedimentary rocks, undivided (Ordovician and Upper Cambrian)—Mapped in southern part of area in Nopah Range, Montgomery Mountains, Panamint Range, Grapevine and Funeral Mountains, Cottonwood Mountains, and Last Chance Range. Includes parts of the Ely Springs Dolomite, Eureka Quartzite, and Pogonip Group. Unit is about 720 m thick in the Dry Mountain area (Burchfiel, 1969) and in the Panamint Range area (Hall, 1971), and about 610 m thick in the Nopah Range (Burchfiel and others, 1983a, b)

Oee Ely Springs Dolomite (Upper Ordovician) and Eureka Quartzite (Upper and Middle Ordovician), undivided

Oes Ely Springs Dolomite (Upper Ordovician)—Widely exposed in map area. Upper 20-25 m consists of slope-forming, finely crystalline dolomite with sparse pelmatozoan fragments. Most of the Ely Springs (lower two thirds to three quarters) is cliff-forming, burrow-mottled, irregularly thin- to thick-bedded dolomite with common planar laminations; includes nodules and lenses of dark-gray to dark-brown chert. Fossils include brachiopods, colonial (*Halysites sp.* and favositids) and solitary corals, gastropods, pelmatozoan ossicles, and stromatoporoids. Lower contact with Eureka Quartzite is disconformity; upper contact with Laketown Dolomite (Silurian) in eastern part of map area is disconformable; in western part of map area, upper contact

- with Hidden Valley Dolomite is gradational and conformable (Hall, 1971). Unit is 50-200 m thick
- Oe Eureka Quartzite (Upper and Middle Ordovician)**—Widely exposed in map area. Consists of fine- to medium-grained orthoquartzite and sandstone. Beds contain tabular-planar and trough crossbed sets and *Skolithos* burrows. At some locations, as in the Spring Mountains, has sandy carbonate beds and colophane nodules (Ross, 1964), and in Panamint Range area, lower part contains silty and sandy dolomite, argillaceous quartzite, and minor shale. Unit is 6-150 m thick
- O€pn Pogonip Group (Middle and Lower Ordovician and Upper Cambrian) and Nopah Formation (Upper Cambrian), undivided**—Mapped in parts of the Nopah Range, Spring Mountains, Quinn Canyon Range, and southernmost Monitor Range
- O€p Pogonip Group (Middle and Lower Ordovician and Upper Cambrian)**—Widely exposed across map area. Consists of (from top to base) the Antelope Valley Limestone, Ninemile Formation, and the Goodwin Limestone (Byers and Barnes, 1967; Ross, 1964). Antelope Valley Limestone is finely to coarsely crystalline limestone and silty limestone; beds are commonly burrowed and have yellowish-orange to pale-red tracks and trails on bedding planes; contains orthocone cephalopods, gastropods (*Maclurites* sp. and *Palliseria* sp.), pelmatozoans, *Receptaculites* sp., and trilobites; also contains ooids and oncoids. Ninemile Formation is shaly siltstone, silty limestone, and dolomite. Goodwin Limestone is limestone and dolomite with dark-brown-weathering chert layers and distinctive intraclastic conglomerate; contains trilobite and brachiopod fragments and sparse ooids. Base of Goodwin contains Late Cambrian conodonts in the Spring Mountains and southern Sheep Range (Page and others, in press). Ninemile Formation is absent in most locations in the eastern part of the map area (Lincoln and Clark Counties). Group is about 410 m thick in the Dry Mountain area (Burchfiel, 1969), 384 m thick in the Panamint Range (Hall, 1971), 320 m thick in Nopah Range (Burchfiel and others, 1983a, b), 360-700 m thick in the Spring Mountains and southern Sheep Range, (Page and others, in press), 957 m thick in the Pahrnagat Range (Tschanz and Pampeyan, 1970), and 1,050 m thick in Nevada Test Site area (Slate and others, 2000)
- €e Emigrant Formation (Upper and Middle Cambrian)**—Exposed in Esmeralda County, Goldfield Hills, Palmetto Mountains, Clayton Valley, and Montezuma Range areas. Consists of an upper limestone and siltstone member, a middle shale member, and lower limestone and chert member. Upper member is interbedded limestone and siltstone; contains sponge spicules and phosphatic brachiopod shells. Middle shale member is olive-gray to greenish-gray shale. Lower member is interbedded limestone and laminated chert, and some intraclastic conglomerate; trilobites common in Magruder Mountain area (McKee, 1985). Unit is 1,524-1,828 m thick (Albers and Stewart, 1972)
- €u Sedimentary rocks, undivided (Upper to Lower Cambrian)**—Mapped in the Last Chance Range and Cottonwood Mountains. In Last Chance Range, northwest of the Cucomongo fault, includes rocks equivalent with parts of the Emigrant, Mule Spring, Harkless, Poleta, and Campito Formations. Southeast of the Cucomongo fault in the Last Chance Range and Cottonwood Mountains, includes rocks equivalent with parts of the Nopah, Bonanza King, Carrara, Zabriskie, and Wood Canyon Formations (McKee, 1985)
- €nb Nopah (Upper Cambrian) and Bonanza King Formations (Upper and Middle Cambrian), undivided**—Mapped in the west-central Pintwater Range; thickness unknown
- €n Nopah Formation (Upper Cambrian)**—Widely exposed in map area. In central part of the map area unit includes, from top to base, Smoky, Halfpint, and Dunderberg Shale Members. Smoky and Halfpint Members consist of alternating light- to dark-gray cliff-forming dolomite that gives these rocks a banded appearance; members contain common stromatolites, oncoids, and brachiopod fragments, and less common pelmatozoan fragments. Dunderberg Shale Member consists of thin-bedded silty limestone, shale, and some siltstone; *Dunderbergia* sp. trilobite fragments are common. Formation is about 345-380 m thick in the Spring Mountains and southern Sheep Range (Page and others, in press), 365 m thick in the northern Last Chance Range (McKee, 1985), 457 m thick in the Panamint Range area (Hall, 1971), about 530 m thick in the Nopah Range and in the Cottonwood Mountains (Burchfiel and others, 1983a, b; Burchfiel, 1969), 766 m thick in the Pahrnagat Range (Reso, 1963), and 720 m thick in the Nevada Test Site area (Slate and others, 2000)
- €nbc Nopah (Upper Cambrian), Bonanza King (Upper and Middle Cambrian), and Carrara Formations (Middle and Lower Cambrian), undivided**—Mapped in the Quinn Canyon, Nopah, Kingston, Resting Springs, Montgomery, Greenwater, Panamint, Funeral, and Grapevine ranges
- €b Bonanza King Formation (Upper and Middle Cambrian)**—Banded Mountain and Papoose Lake Members recognizable in most parts of the map area. Partly correlative with the Highland Peak Formation in Lincoln County (Tschanz and Pampeyan, 1970). Banded Mountain Member is fine- to medium crystalline dolomite and limestone; alternating light- to dark-gray colors of beds give the member a distinctive banded appearance. Orange burrow mottling is prevalent, and several layers have dark-brown-weathering chert beds and nodules. Base of the member is marked by the regionally

extensive “silty unit” (Barnes and Palmer, 1961; Guth, 1980; Gans, 1974). The underlying Papoose Lake Member is dolomite with distinctive orange burrow mottles; sparse limestone and silty dolomite beds. Formation is 580-1,700 m thick in the Nevada Test Site area (Slate and others, 2000), about 1,000 m thick in the Dry Mountain area (Burchfiel, 1969), 1,150-1,350 m thick in the Montgomery Mountains and Nopah Range, respectively (Burchfiel and others, 1983a, b), about 610 m thick in the northern Last Chance Range (McKee, 1985), 445-900 m thick in the Spring Mountains and southern Sheep Range (Page and others, in press), and reaches a maximum thickness of 1,633 m in the Groom Range area, Lincoln County (Tschanz and Pampeyan, 1970)

€c Carrara Formation (Middle and Lower Cambrian)—Interbedded limestone, silty limestone, siltstone, sandstone, and shale. Upper part is mostly cliff-forming silty limestone, and lower part contains mostly clastic rocks. Limestone beds contain ooids, oncoids, and stromatolites. Contact with overlying Bonanza King Formation is commonly marked by a change from slope-forming limestone of the Carrara to more massive cliff-forming limestone and dolomite of the Bonanza King Formation. Unit is 457 m thick in the northwestern Spring Mountains (Vincelette, 1964), 500 m thick in the Striped Hills, 350 m thick at Bare Mountain (Slate and others, 2000), and 420 m thick in the Montgomery Mountains (Burchfiel and others, 1983a, b)

€ms Mule Spring Limestone (Lower Cambrian)—Exposed in Palmetto Mountains, Clayton Ridge, Goldfield Hills, and Montezuma Range. Thin- to medium-bedded, aphanitic to finely-crystalline limestone with trilobite fragments; thicker beds have oncoids; calcareous siltstone and silty limestone are in lower part. Contact with overlying Emigrant Formation is sharp and disconformable. Total thickness estimated at 120-150 m (Albers and Stewart, 1972)

€h Harkless Formation (Lower Cambrian)—Exposed mostly in Esmeralda County, Nevada. Upper member is siliceous siltstone and shale. Lower member is interbedded sandstone, siltstone and thin-bedded limestone. Unit contains *Olenellid* trilobites and archaeocyathids, especially in lower member. Conformable with overlying Mule Spring Formation and underlying Poleta Formation; about 760-1,066 m thick (Albers and Stewart, 1972; McKee, 1985)

€z Zabriskie Quartzite (Lower Cambrian)—Laminated to thick-bedded orthoquartzite, commonly cross-bedded and well-cemented; has *skolithos*. Unit is 30-350 m thick in the Nevada Test Site area (Slate and others, 2000), 70 m thick in the Montgomery Mountains (Burchfiel and others, 1983a, b), and 106 m thick in the Dry Mountain area

€Zws Wood Canyon Formation (Lower Cambrian and Late Proterozoic) and Stirling Quartzite (Late

Proterozoic), undivided—Mapped in the Resting Springs, Nopah, Kingston, and Panamint ranges, and Montgomery Mountains

€Zw Wood Canyon Formation (Lower Cambrian and Late Proterozoic)—Quartzite, sandstone, siltstone, shale, and dolomite. Shaly and silty units increase in abundance in the upper part of the formation. Upper and lower thirds of unit consist of interbedded quartzite, micaceous siltstone, and minor silty dolomite. Middle third is mostly quartzite and siltstone, and minor distinctive arkosic conglomerate (Slate and others, 2000). Locally metamorphosed to subgreenschist and lower greenschist facies south of Nevada Test Site in Funeral Mountains and Specter Range (Slate and others, 2000). Unit is from 700 to 1,150 m thick in the Nevada Test Site area (Slate and others, 2000), about 600 m thick in the Montgomery Mountains (Burchfiel and others, 1983a, b), and 640-820 m thick in the northwest Spring Mountains (Vincelette, 1964)

€po Poleta Formation (Lower Cambrian)—Exposed in the Montezuma Range, Magruder Mountain, and Palmetto Mountains in Esmeralda County, Nevada. Upper, middle, and lower informal members recognized. Upper member is thin-bedded limestone with some poorly preserved archaeocyathids. Middle member is phyllitic siltstone, limestone, and sandstone or quartzite. *Olenellid* trilobites common in siltstone, and archaeocyathids in limestone. Contact with overlying Harkless Formation is sharp. Unit is greater than 600 m thick (McKee, 1985)

Campito Formation (Lower Cambrian and Late Proterozoic)—Exposed in Esmeralda County, Nevada. Divided into:

€cam Montenegro Member (Lower Cambrian)—Siltstone, shale, and sandstone. *Olenellid* trilobites and archaeocyathids are rare in the shaly units. Unit is 315 m thick (Albers and Stewart, 1972)

€Zcaa Andrews Mountain Member (Lower Cambrian and Late Proterozoic)—Thin-bedded quartzite and sparse beds of greenish shale. Sandstone is cross-bedded, and has ripple marks and worm trails. About 910 m thick (McKee, 1985)

Zd Deep Spring Formation (Late Proterozoic)—Exposed in Esmeralda County, Nevada, where upper, middle, and lower members are recognized. Upper member is siltstone, quartzite, and minor dolomite. Middle member is oolitic-limestone, quartzite, and some dolomite. Lower member is limestone and dolomite, and some siltstone. Unit apparently conformable with overlying Campito Formation. About 487-548 m thick (Albers and Stewart, 1972; McKee, 1985)

Zs Stirling Quartzite and Reed Dolomite, undivided (Late Proterozoic)—The Stirling Quartzite and Reed Dolomite are time-equivalent units with the

Stirling Quartzite representing the eastern facies and the Reed Dolomite the western facies within the map area. These units are mapped as undivided due to similar hydrologic properties. The Stirling Quartzite contains quartzite and sandstone, conglomeratic quartzite, and minor beds of micaceous siltstone. Quartz grains are usually fine- to medium-grained. A 20 m thick white quartzite marks the base of the unit in the Specter Range and northwestern Spring Mountains. Locally metamorphosed to subgreenschist and lower greenschist facies at Bare Mountain, Funeral Mountains (Slate and others, 2000), and parts of the northwestern Spring Mountains (Vincellette, 1964). Unit forms massive cliffs and is 975 m thick in northwestern Spring Mountains (Nolan, 1924; Vincellette, 1964), 1,100 m thick in the Montgomery Mountains (Burchfiel and others, 1983a, b), 2,100 m thick in the Funeral Mountains (Wright and Troxel, 1993), 1,500 m thick in the Belted Range, and 700 m thick at Bare Mountain (Slate and others, 2000; Stewart, 1974). The Reed Dolomite is exposed in Slate Ridge, Gold Mountain, Palmetto Mountain, and Magruder Mountain areas, Esmeralda County, Nevada, and parts of Inyo County, California. The Reed Dolomite contains medium to coarsely crystalline dolomite, sandy dolomite, and minor amounts of limestone, siltstone, and quartzite; contains ooids and locally is pelletal. Reed Dolomite is greater than 457 m thick (Albers and Stewart, 1972)

Zj Johnnie and Wyman Formations, undivided (Late Proterozoic)—The Johnnie and Wyman Formations are time-equivalent units with the Johnnie representing the eastern facies and the Wyman, the western facies within the map area. These units are mapped as undivided due to similar hydrologic properties. The Johnnie Formation is exposed in the Nopah, Panamint, and Kingston Ranges and Spring and Funeral Mountains. Unit contains quartzite, conglomeratic quartzite, siltstone, shale, and minor limestone and dolomite beds. Upper part contains Rainstorm Member (Stewart, 1974), which can be traced regionally across parts of southern Nevada and southeastern California. Rainstorm Member contains mixed clastic and carbonate units and the “Johnnie oolite,” a distinctive ooid-bearing dolomite unit in lower part of member. Formation exceeds 900 m thick in the Halfpint Range (Barnes and Christiansen, 1967), and is estimated at 2,000 m thick in the Funeral Mountains (Wright and Troxel, 1993); only 600 m of unit exposed in Spring Mountains (Vincellette, 1964). The Wyman Formation is exposed in the Slate Ridge and Gold Mountain areas of Esmeralda County and western parts of Inyo County, California. Unit contains phyllitic siltstone, limestone, sandy limestone, and calcareous sandstone. Maximum incomplete thickness of the Wyman Formation is 400 m (Albers and Stewart, 1972)

ZYp Pahump Group (Late and Middle Proterozoic)—Exposed in Black and Funeral Mountains and Panamint and Kingston Ranges of Death Valley region (Wright and others, 1974). Includes, from top to base, Kingston Peak, Beck Spring, and Crystal Spring Formations. Kingston Peak is staurolite- and biotite-bearing pelitic schist, metaconglomerate, and calcitic marble; contains some basalt flows. Beck Spring is calcitic marble and minor dolomite. Crystal Spring is staurolite-biotite-bearing pelitic schist, micaceous quartzite, calcitic marble, and amphibolite; has diabase sills that were dated at 1.08 Ga (Heaman and Grotzinger, 1992). Group is 1,570 m thick in Funeral Mountains (Wright and Troxel, 1993)

Xmi Metamorphic and igneous rocks, undifferentiated (Early Proterozoic)—Exposed mainly in the southern part of the map, in Greenwater, Funeral, Nopah, and Panamint Ranges, where it consists of biotite schist, biotite-hornblende-schist, and biotite-epidote schist. Schist is intruded by gneissic monzogranite in Trappman Hills, and north of western Gold Flat, schist and granite intruded by aplite and pegmatite dikes. Includes metasedimentary rocks in Funeral Mountains that are locally intruded by silicic plutons dated at 1.7 Ga (Wright and Troxel, 1993)

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