

U.S. Department of the Interior
U.S. Geological Survey

Ground-Water Discharge Determined from Estimates of Evapotranspiration, Death Valley Regional Flow System, Nevada and California

Water-Resources Investigations Report 01-4195

Prepared in cooperation with the
OFFICE OF ENVIRONMENTAL RESTORATION AND WASTE MANAGEMENT,
U.S. DEPARTMENT OF ENERGY
National Nuclear Security Administration
Nevada Operations Office, under
Interagency Agreement DE-AI08-96NV11967



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By Randell J. Laczniak, J. LaRue Smith, Peggy E. Elliott, Guy A. DeMeo,
and Melissa A. Chatigny, U.S. Geological Survey; *and* Gaius J. Roemer,
GeoTrans

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Carson City, Nevada
2001

U.S. DEPARTMENT OF THE INTERIOR
GALE A. NORTON, Secretary

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CONVERSION FACTORS AND VERTICAL DATUM

	Multiply	By	To obtain
	acre	0.4047	square hectometer
	acre-foot (acre-ft)	0.001233	cubic hectometer
	acre-foot per year (acre-ft/yr)	0.001233	cubic hectometer per year
	cubic foot per day (ft ³ /d)	0.02832	cubic meter per day
	cubic foot per second (ft ³ /s)	0.02832	cubic meter per second
	foot (ft)	0.3048	meter
	foot per second (ft/s)	0.3048	meter per second
	foot per year (ft/yr)	0.3048	meter per year
	gallons per minute (gal/min)	0.0631	liter per second
	inch (in.)	25.4	millimeter
	mile (mi)	1.609	kilometer
	square mile (mi ²)	2.590	square kilometer

Temperature: Degrees Celsius (°C) can be converted to degrees Fahrenheit (°F) by using the formula °F = [1.8(°C)] + 32. Degrees Fahrenheit can be converted to degrees Celsius by using the formula °C = 0.556(°F-32).

Sea level: In this report, “sea level” refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929, formerly called “Sea-Level Datum of 1929”), which is derived from a general adjustment of the first-order leveling networks of the United States and Canada.

ACRONYMS

DVRFS	Death Valley regional flow system
ET	Evapotranspiration
MSAVI	Modified soil-adjusted vegetation index
NTS	Nevada Test Site
TM	Thematic Mapper

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ABSTRACT

The Death Valley regional flow system (DVRFS) is one of the larger ground-water flow systems in the southwestern United States and includes much of southern Nevada and the Death Valley region of eastern California. Centrally located within the ground-water flow system is the Nevada Test Site (NTS). The NTS, a large tract covering about 1,375 square miles, historically has been used for testing nuclear devices and currently is being studied as a potential repository for the long-term storage of high-level nuclear waste generated in the United States. The U.S. Department of Energy, as mandated by Federal and State regulators, is evaluating the risk associated with contaminants that have been or may be introduced into the subsurface as a consequence of any past or future activities at the NTS. Because subsurface contaminants can be transported away from the NTS by ground water, components of the ground-water budget are of great interest. One such component is regional ground-water discharge.

Most of the ground water leaving the DVRFS is limited to local areas where geologic and hydrologic conditions force ground water upward toward the surface to discharge at springs and seeps. Available estimates of ground-water discharge are based primarily on early work done as part of regional reconnaissance studies. These early efforts covered large, geologically complex areas and often applied substantially different techniques to estimate ground-water discharge. This report

describes the results of a study that provides more consistent, accurate, and scientifically defensible measures of regional ground-water losses from each of the major discharge areas of the DVRFS.

Estimates of ground-water discharge presented in this report are based on a rigorous quantification of local evapotranspiration (ET). The study identifies areas of ongoing ground-water ET, delineates different ET areas based on similarities in vegetation and soil-moisture conditions, and determines an ET rate for each delineated area. Each area, referred to as an ET unit, generally consists of one or more assemblages of local phreatophytes or a unique moist soil environment. Ten ET units are identified throughout the DVRFS based on differences in spectral-reflectance characteristics. Spectral differences are determined from satellite imagery acquired June 21, 1989, and June 13, 1992. The units identified include areas of open playa, moist bare soils, sparse to dense vegetation, and open water. ET rates estimated for each ET unit range from a few tenths of a foot per year for open playa to nearly 9 feet per year for open water.

Mean annual ET estimates are computed for each discharge area by summing estimates of annual ET from each ET unit within a discharge area. The estimate of annual ET from each ET unit is computed as the product of an ET unit's acreage and estimated ET rate. Estimates of mean annual ET range from 450 acre-feet in the Franklin Well area to 30,000 acre-feet in Sarcobatus Flat. Ground-water discharge is estimated as annual ET minus that part of ET attributed to local precipitation

Mean annual ground-water discharge estimates range from 350 acre-feet in the Franklin Well area to 18,000 acre-feet in Ash Meadows. Generally, these estimates are greater for the northern discharge areas (Sarcobatus Flat and Oasis Valley) and less for the southern discharge areas (Franklin Lake, Shoshone area, and Tecopa/ California Valley area) than those previously reported.

INTRODUCTION

The Death Valley regional flow system (DVRFS) as delineated by Harrill and others (1988, sheet 1) covers an area of 15,800 mi² of southern Nevada and southeastern California. This ground-water flow system is west of Las Vegas, Nevada, and includes parts of Clark, Esmeralda, Lincoln, and Nye Counties, Nevada and Inyo and San Bernardino Counties, California. Death Valley, the largest valley within the flow system, forms the southwestern boundary of the flow system.

The DVRFS is centered about the Nevada Test Site (NTS) and Yucca Mountain. The NTS is fairly extensive at approximately 1,375 mi² and serves as the primary continental location for testing nuclear devices. Yucca Mountain, which borders the NTS on the west, is the site for a potential high-level nuclear waste disposal facility. The U.S. Department of Energy is currently evaluating the risk to the general public associated with past testing activities and the future storage of high-level radioactive waste as required by Federal and State mandates. As part of this evaluation, studies are being done to develop ground-water flow and transport models of the region. The accuracy of these models and confidence given to their results are dependent on an understanding of the region's geology and hydrology. Ground-water discharge is a major component of the hydrology controlling ground-water flow throughout the region.

Accurate estimates of ground-water discharge from the DVRFS are crucial to developing realistic simulations of ground-water flow and contaminant transport away from the NTS and Yucca Mountain. Although early reconnaissance studies completed in the 1960s and 1970s (Malmberg and Eakin, 1962; Walker and Eakin, 1963; Pistrang and Kunkel, 1964; Malmberg, 1967; Cardinalli and others, 1968; Rush, 1968; Winograd and Thordarson, 1975) provide general estimates of regional discharge from many

of the major valleys in the region, the methodologies and techniques applied often differed among the various investigations. More accurate and defensible estimates of regional ground-water discharge across the DVRFS require that a more thorough and consistent approach be applied to all major discharge areas within its boundaries.

The U.S. Geological Survey, in cooperation with the U.S. Department of Energy, began long-term studies to refine and improve estimates of mean annual ground-water discharge from the Ash Meadows (Laczniak and others, 1999) and Oasis Valley discharge areas (Reiner and others, 2002). In 1998, the U.S. Geological Survey, also in cooperation with the U.S. Department of Energy, initiated a 2-year effort to extend the knowledge gained and techniques developed during these investigations to estimate ground-water discharge from other major discharge areas throughout the DVRFS. The latter effort is the focus of this report.

Purpose and Scope

The purpose of the study was to reduce uncertainty in estimates of ground-water discharge from the DVRFS by refining existing estimates for most major discharge areas in the flow system (fig. 1). Discharge areas evaluated specifically for this study include Chicago Valley, the Franklin Well area, Franklin Lake, the Shoshone area, Stewart Valley, and the Tecopa/ California Valley area in California and Sarcobatus Flat in Nevada (fig. 1). Estimates for Ash Meadows and Oasis Valley in Nevada, although studied and documented independently as part of more detailed efforts (Laczniak and others, 1999; and Reiner and others, 2002, respectively), are included for completeness. For reasons given in a later section, estimates for other major discharge areas, namely Death Valley in California and Pahrup and Lida Valleys in Nevada, are not given.

The report presents estimates of mean annual ground-water discharge and evapotranspiration, and describes the general procedure used to make the estimates. This effort resulted in many digital products that are available from the USGS node of the National Spatial Data Infrastructure (NSDI) at URL: <<http://nsdi.usgs.gov/>>. These revised estimates of ground-water discharge can be used to improve general concepts of regional flow and as input to models simulating ground-water flow and transport, and should increase confidence in model-based risk assessments of past,

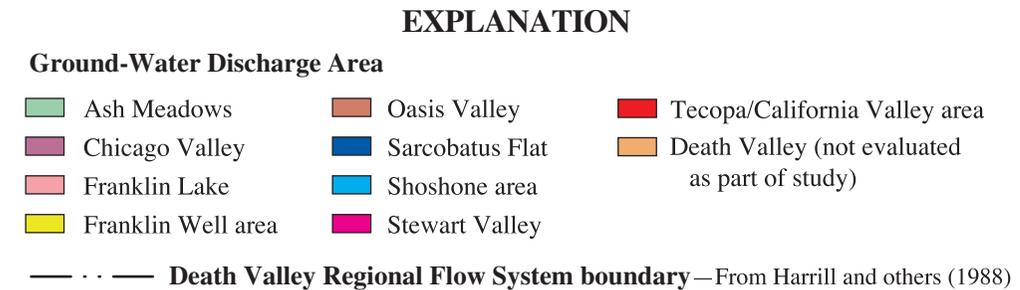
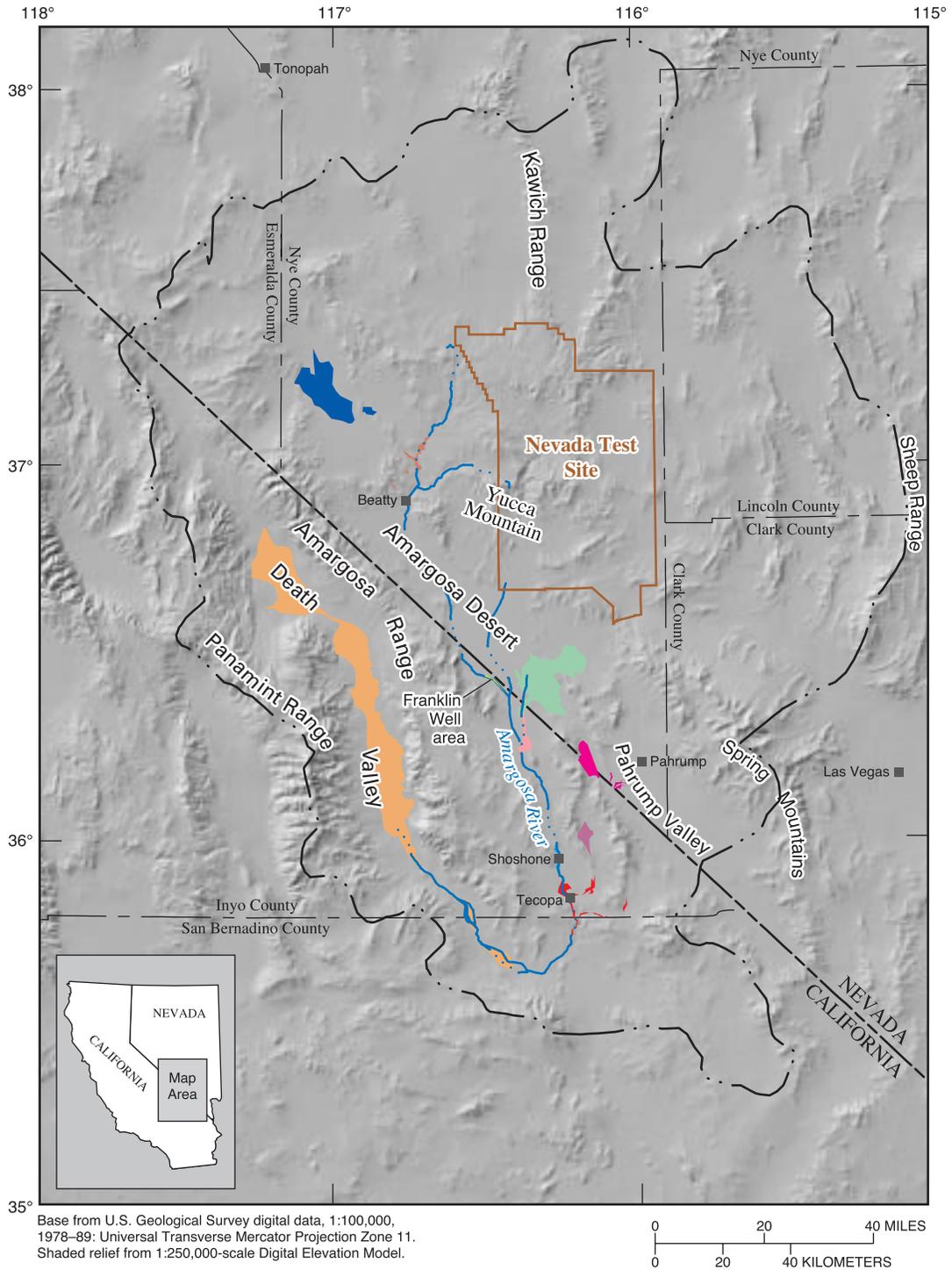


Figure 1. Major areas of natural ground-water discharge in Death Valley regional flow system.

ongoing, or future activities at the NTS and Yucca Mountain site. To provide readers, and flow and transport modelers, with a quantification of the error associated with these estimates, an uncertainty analysis is included as an appendix.

Acknowledgments

The authors express their appreciation to the U.S. Fish and Wildlife Service, the National Park Service, the Bureau of Land Management, and the many private landowners in the area that openly provided access to their property. The authors wish to thank S.R. Reiner for providing meteorological data and imagery from the Oasis Valley area. Author appreciation also is due to the U.S. Department of Energy for continued support given to this and other investigations of ground-water discharge throughout the Death Valley region.

Stewardship

Much of the land in the Death Valley region is owned and administered by the U.S. Government (fig. 2). The primary stewards of these Federal lands are the National Park Service, U.S. Forest Service, Bureau of Land Management, U.S. Air Force, U.S. Department of Energy, and U.S. Fish and Wildlife Service. Much smaller parcels of privately owned land are concentrated in a few agricultural areas, mining centers, and recreational gateways including Pahrump, Amargosa Farms, Amargosa Valley, and Beatty, Nevada, and Tecopa and Shoshone, California.

GENERAL DESCRIPTION AND SETTING

Physiography and Geology

The DVRFS lies within the southern Great Basin region, an internally drained subdivision of the Basin and Range physiographic province. The region is characterized by low rainfall, intermittent streams, large internal surface drainages, and sparsely distributed spring-fed oases. The dominant physiographic features are linear, north trending mountain ranges separating broad, elongated valleys, formed in response to a long and ongoing period of crustal extension (Stewart, 1980, p. 110). Large vertical displacements along faults offset bedrock blocks separating north-trending mountain ranges from similar trending alluvial-filled valleys. The

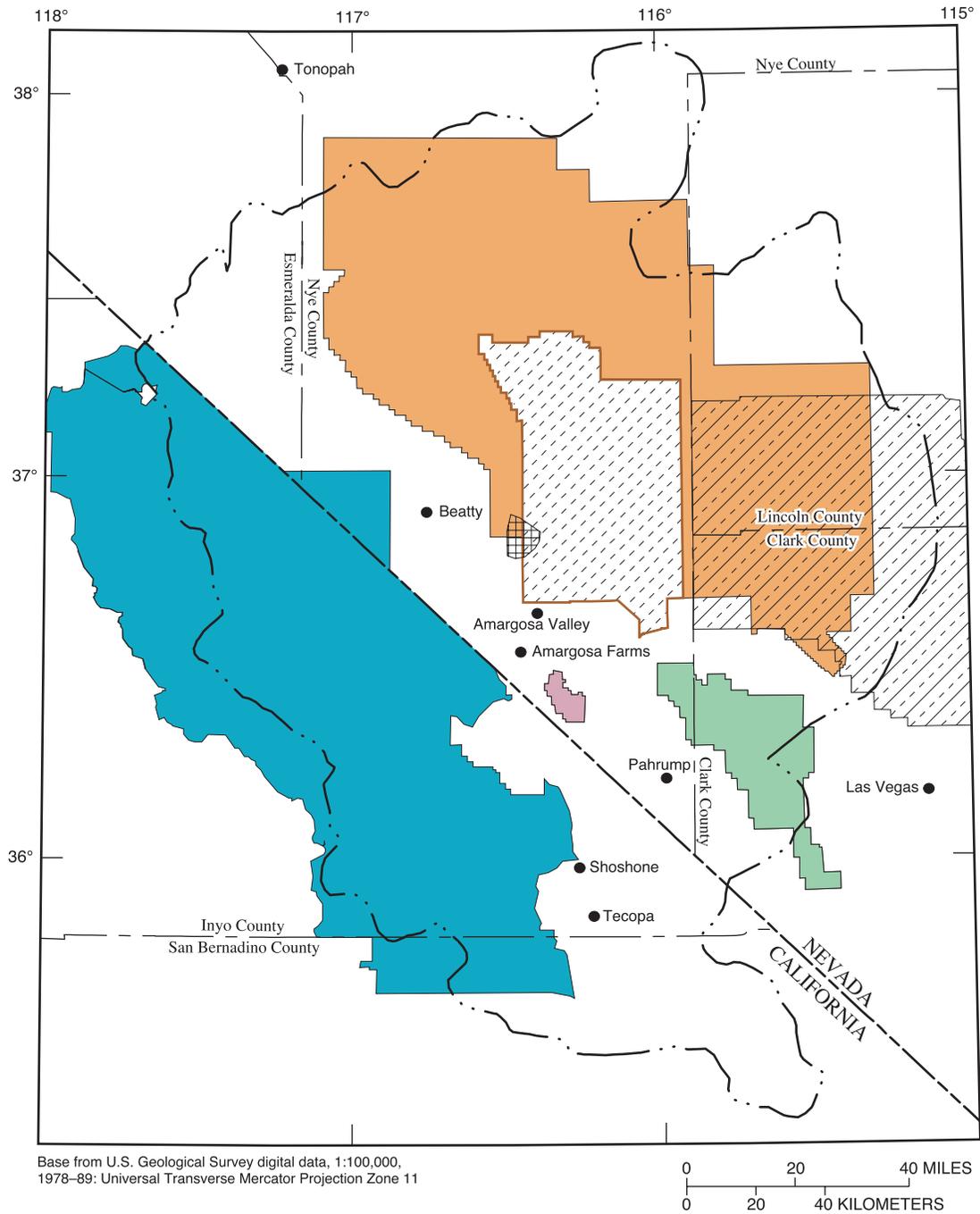
region includes several large, prominent valleys (Amargosa Desert, and Pahrump and Death Valleys; fig. 1) and major mountain ranges (Amargosa, Panamint, Kawich, and Sheep Ranges, and Spring Mountains; fig. 1).

Major mountain ranges rimming the larger valleys typically rise more than 5,000 ft above valley floors. Higher peaks typically crest at more than 8,000 ft above sea level. The highest peaks, reaching more than 11,000 ft above sea level, are found in the Spring Mountains and Panamint Range. One of the more striking elevation contrasts occurs in Death Valley where peaks in the Panamint Range rise as much as 11,000 ft above the valley floor. Valley floors are typified by a gently sloping terrain and range in altitude from about 200 ft below sea level in Death Valley to about 5,000 ft above sea level in the northernmost valleys.

Most of the major mountain ranges in the general region are composed of pre-Cenozoic rocks of diverse age and lithology (Stewart, 1980). Paleozoic carbonate rock and Paleozoic and Proterozoic siliciclastic rock, and Tertiary volcanic rock constitute the primary rock types of the hills, ridges, and mountain ranges. The intervening valleys are filled primarily with alluvium, lacustrine and palustrine deposits, and volcanic rock.

Climate

The climate of the DVRFS is arid in the south to semiarid in the north. The southern part of the flow system lies within the Mojave Desert, an area characterized by short mild winters, long hot summers, and low annual rainfall and humidity. The northern part of the flow system lies within the Great Basin Desert, an area characterized by cold winters, warm summers, and low-to-moderate annual rainfall and humidity. The wide ranges in altitude and latitude over the DVRFS contribute to climatic conditions that vary dramatically on both seasonal and daily time scales. Temperatures, higher at lower altitudes and more southern latitudes of the DVRFS, range from winter lows below zero in mountainous areas to summer highs that exceed 120°F in Death Valley. In many valleys, the daily temperature range exceeds 30°F. Precipitation also ranges widely over the DVRFS and is dependent on altitude, latitude, and location relative to surrounding mountain peaks. Mean annual precipitation is lowest on the valley floors



EXPLANATION

- | | |
|--------------------------------------|--------------------------------|
| Ash Meadows National Wildlife Refuge | Desert National Wildlife Range |
| Death Valley National Park | Nevada Test Site |
| Nellis Air Force Range | Yucca Mountain Control Area |
| Toiyabe National Forest | |
- Death Valley Regional Flow System boundary—From Harrill and others (1988)

Figure 2. Major areas of Federally administered land in Death Valley regional flow system.

at the more southern latitudes and ranges from less than 2 in. in Death Valley to more than 25 in. in the Spring Mountains (Daly and others, 1994).

Vegetation and Wildlife

A great diversity of plants, fish, and local wildlife are found throughout the area. Much of the biota relies on water originating from mountain runoff and from mountain and valley springs scattered about the area for survival. Spring pools and their associated drainages and wetlands provide habitat for numerous species of endemic and rare fish, aquatic insects, and plants. Ash Meadows, a major ground-water discharge area (fig. 1), has the highest concentration of endemic species in the continental United States (Chaplin and others, 2000, p. 198). A notable example of one of these endemic species is the endangered Devils Hole pupfish (*Cyprinodon diabolis*) whose only natural habitat is limited to a small ground-water pool known as Devils Hole in Ash Meadows. Other prominent species prevalent throughout the area are the Desert Bighorn Sheep (*Ovis canadensis*) and the Desert Tortoise (*Gopherus agassizii*).

Vegetation throughout the area is diverse and exhibits a strong dependence on climate, water availability, and soil chemistry. More dense growths are concentrated near springs, on poorly drained bottomland, and at higher altitudes in the major mountain ranges. This vegetation provides food and shelter to numerous birds, insects, reptiles, and small mammals. Although the area is characterized as desert, plant assemblages and species are numerous and include many varieties of grasses, reeds, shrubs, and trees. Vegetation supported by local springflow includes groves of ash, cottonwood, willow, and mesquite; thick stands of saltcedar; expansive meadows of salt, wire, and bunch grasses; and open marshland of cattails, reeds, and bulrush. Extensive stands of phreatophytic shrubs, including greasewood, rabbitbrush, seepweed, seep willow, and wolfberry, are found in many valley lowlands. Upland areas beyond the influence of local drainages support more classic desert flora including sparse covers of healthy creosote bush, saltbush, and desert holly. Mountainous areas and highlands are dominated by sparse to dense woodlands that primarily include juniper, piñon pine, and various species of conifer.

Drainage

The area occupied by the DVRFS is drained internally with Death Valley serving as the terminal drainage of the flow system. Drainage features consist primarily of intermittent streams fed by spring snow-melt or infrequent, major storms. Only a few short reaches, located downgradient of major springs, flow year round. Perennial flows are greatest in winter, when cooler temperatures and dormant vegetation result in reduced ET rates. Many of the individual intermountain basins are themselves internally drained with only a few having any surface outlet. The Amargosa River, the largest and most continuous drainage within the flow system, drains about 5,800 mi² and is the only major drainage into Death Valley originating outside the valley proper (fig. 1).

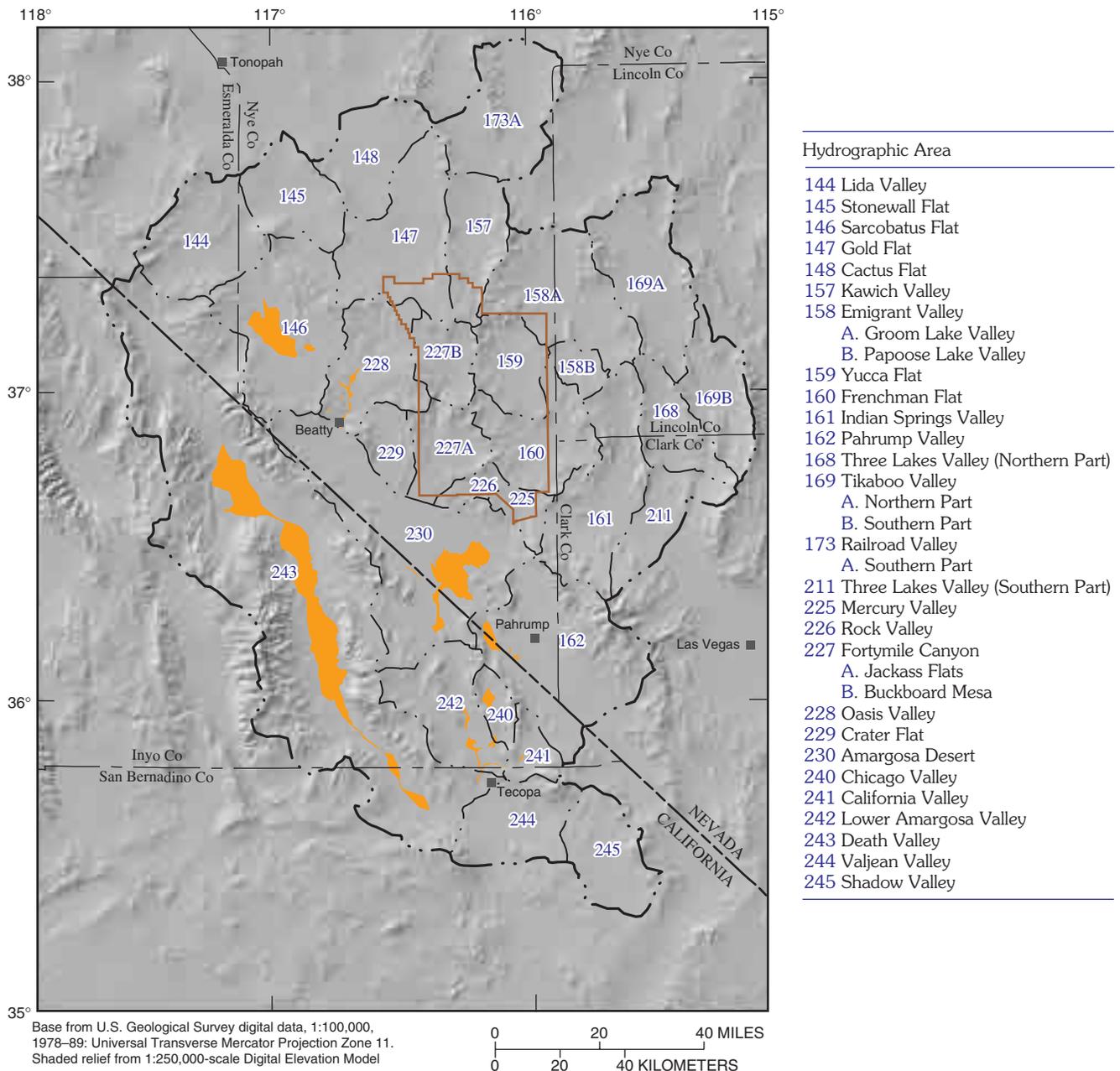
A total of 27 hydrographic areas¹ (fig. 3) are recognized within the DVRFS. Hydrographic areas are delineated primarily from topography and geologic structures, and generally correspond to major surface drainages. These areas serve as the basic units used by State and local agencies for water-resources planning, and in combination, form the regional flow systems delineated throughout the area (Harrill and others, 1988).

Hydrogeology

Ground water in the DVRFS passes through a diverse assemblage of rocks. These rocks differ substantially in terms of age, composition, and water-transmitting properties. The northwestern part of the flow system is predominantly volcanic rock, the eastern and southern parts are predominantly carbonate and siliciclastic rocks, and most major valleys are filled with unconsolidated deposits derived from the surrounding highlands (Laczniaik and others, 1996, fig. 4).

Ground water originates from precipitation falling on the higher mountain ranges and mesas in the northern and eastern parts of the region (Winograd

¹ Formal hydrographic areas in Nevada were delineated systematically by the U.S. Geological Survey and Nevada Division of Water Resources in the late 1960's (Rush, 1968; Cardinalli and others, 1968) for scientific and administrative purposes. The official hydrographic-area names, numbers, and geographic boundaries continue to be used in U.S. Geological Survey scientific reports and Nevada Division of Water Resources administrative activities.



EXPLANATION

- Ground-water discharge area**
- Death Valley Regional Flow System Boundary**—From Harrill and others (1988)
- Hydrographic area boundary**—Label shows hydrographic area number. See table on right for hydrographic area name. Modified from Cardinalli and others (1968), Rush (1968), and Harrill and others (1988)
- Nevada Test Site boundary**

Figure 3. Hydrographic areas of Death Valley regional flow system.

Hydrographic Area

- 144 Lida Valley
- 145 Stonewall Flat
- 146 Sarcobatus Flat
- 147 Gold Flat
- 148 Cactus Flat
- 157 Kawich Valley
- 158 Emigrant Valley
 - A. Groom Lake Valley
 - B. Papoose Lake Valley
- 159 Yucca Flat
- 160 Frenchman Flat
- 161 Indian Springs Valley
- 162 Pahrump Valley
- 168 Three Lakes Valley (Northern Part)
- 169 Tikaboo Valley
 - A. Northern Part
 - B. Southern Part
- 173 Railroad Valley
 - A. Southern Part
- 211 Three Lakes Valley (Southern Part)
- 225 Mercury Valley
- 226 Rock Valley
- 227 Fortymile Canyon
 - A. Jackass Flats
 - B. Buckboard Mesa
- 228 Oasis Valley
- 229 Crater Flat
- 230 Amargosa Desert
- 240 Chicago Valley
- 241 California Valley
- 242 Lower Amargosa Valley
- 243 Death Valley
- 244 Valjean Valley
- 245 Shadow Valley

and Thordarson, 1975; Waddell and others, 1984; Laczniaik and others, 1996; Harrill and Prudic, 1998). Regional flow generally is southward and westward away from the major recharge areas toward Death Valley and other intermediate areas of ground-water discharge (fig. 1). Ground water flows primarily through faults, fractures, and joints in consolidated rock, and through the coarser grained sand and gravel deposits in unconsolidated sediment. Beneath the major valleys, ground water flows primarily through unconsolidated valley-fill deposits and, where present, through underlying fractured and faulted volcanic and carbonate rocks. Beneath highland areas, ground-water flows through fractured and faulted volcanic and carbonate rock. Low permeability siliciclastic (quartzite and siltstone) and crystalline rocks, fine-grained valley-fill deposits (clay), and non-fractured volcanic rock form the major confining units of the flow system. Total ground-water flow is estimated at about 70,000 acre-ft/yr with the largest proportion moving through highly transmissive, fractured and faulted carbonate rock in the eastern and southern parts of the flow system (Winograd and Pearson, 1976; Harrill and others, 1988, sheet 2; Dettinger and others, 1995).

Ground-water flow rates and directions are influenced by abrupt changes in the water-transmitting properties of the rocks. Tectonic movements along major faults often juxtapose rocks of differing permeability and can intensely fracture the nearby host rock. Permeability contrasts caused by these faults and other lithologic controls create local pathways by which some ground water exits the regional flow system and discharges to land surface. Ground water, some of which originates at or flows beneath the NTS and Yucca Mountain area, emerges as seeps and springs in the alluvium, and as springs in fractured and faulted volcanic and carbonate rock. Most local springflow evaporates or is transpired by the local vegetation. Springflow, while relatively constant throughout recent history (Tim Mayer, U.S. Fish and Wildlife Service, written commun., 1997), varies substantially across the flow system. Measured springflow ranges from less than 1 gal/min at numerous small springs scattered about the area to nearly 3,000 gal/min at the largest spring in Ash Meadows (fig. 1). Ground water not discharged to the surface, exits the discharge area and continues flowing down gradient toward the next area of discharge. This underflow differs from area to area and can be a major component of the water budget.

The few communities, Federal facilities, and agricultural, mining, ranching, and recreational interests within the area of the DVRFS rely almost entirely on ground water for their water supply. A large part of this need is fulfilled by spring-flow diversions, but the demands of larger users and those more distant from springs are met by pumping ground water. Most wells produce water from valley-fill deposits, but some larger supply wells produce from fractured and faulted carbonate and volcanic rock. The withdrawal of ground water from much of the area is limited by Federal mandates instituted to protect rare and endangered species dependent for their survival on regional springflow in Ash Meadows and Death Valley.

ESTIMATION OF ANNUAL EVAPOTRANSPIRATION AND GROUND-WATER DISCHARGE

Historically, estimates of evapotranspiration in the DVRFS were computed as part of regional assessments of the ground-water resource (Malmberg and Eakin, 1962; Walker and Eakin, 1963; Pistrang and Kunkel, 1964; Malmberg, 1967; Rush, 1968). These regional assessments estimated annual ET losses as the product of the acreage of phreatophytes within a discharge area and an annual ET rate representative of the vegetation and soil conditions of the discharge area. Acreages were estimated by delineating vegetation from aerial photographs and field mapping. Annual ET rates typically were estimated from values reported for similar plant assemblages found throughout the western United States (Lee, 1912; White, 1932; Young and Blaney, 1942; Gatewood and others, 1950; Robinson, 1958). Most of these rates were determined from tank experiments as described by Gatewood and others (1950, p. 105–111). Although the methodology is technically sound, its overall accuracy depends on the accuracy of the individual components. Inexact estimates of the acreage, the annual ET rate, or both could introduce substantial error into the calculation. Recent studies of ET rates for vegetation and soil conditions in and near the DVRFS (Johnson, 1993; Czarnecki, 1997; Laczniaik and others, 1999; Nichols, 2001) indicate rates somewhat different than those presented in earlier reconnaissance studies. These more recent estimates of ET rates were determined primarily from energy-balance techniques (Brutsaert, 1982) that require rigorous and long-term measurements of micrometeorological

data. Noted differences between early and more recent measurements of ET rates are expected considering differences in the local climatic, vegetation, and soil conditions between sites and differences in the methods used to measure ET.

Most previous attempts to quantify ground-water discharge from the DVRFS have been based on measurements of springflow or on estimates of ET from the major discharge areas (Malmberg and Eakin, 1962; Walker and Eakin, 1963; Pistrang and Kunkel, 1964; Malmberg, 1967; Rush, 1968, Dudley and Larson, 1976). In discharge areas where seeps are dominant or where springs are difficult to measure, estimates of ground-water discharge account only for measurable outflow. Alternatively, ground-water discharge has been estimated by efforts to quantify local ET. Assuming that all spring and seep flow is evaporated or transpired by the local vegetation, ET-based estimates also include water upwardly diffused into the shallow flow system from the underlying regional flow system.

The need for accurate estimates of ground-water discharge and the differences in ET rates estimated by past and more recent studies prompted studies to re-evaluate and more rigorously and consistently quantify mean annual ET and ground-water discharge throughout the DVRFS. The approach used in this study is similar to that used in the early reconnaissance studies in that ET is calculated as the product of an area and a rate. Although the approach is similar, the specific techniques used to delineate ET areas and to determine an appropriate ET rate differ in that they take advantage of modern technologies and are applied consistently over the DVRFS. Because these techniques are expected to reduce errors, the revised ET estimates should be more reliable. Early reconnaissance studies delineated ET areas based on generalized vegetation and soil mapping—this study utilizes satellite imagery in combination with remote sensing and a geographic information system. The reconnaissance studies estimated ET rates based on measurements made for similar phreatophytes found at locations outside the study area—this study relies primarily on recent estimates determined from micrometeorological measurements made in areas within or adjacent to the study area (Johnson, 1993; Czarnecki, 1997; Laczniaak and others, 1999, and Nichols, 2001; and Reiner and others, 2002). Another difference between the two approaches is in the calculation of ground-water discharge. The reconnaissance studies assume that ground-water discharge is equivalent to ET; this study, however, estimates discharge by

reducing the ET rate to account for any local precipitation falling on the area. This reduction to the ET rate generally is small considering the aridity of the area.

ET and ground-water discharge estimates computed in this study account only for ground water lost to the atmosphere. Estimates do not account for ground water pumped from the flow system or for ground water exiting the discharge area as underflow. Presently, ground water consumed to support agricultural, domestic, and operational needs is minimal and probably does not exceed more than a few hundred acre-feet per year in any major discharge area other than in Pahrump and Lida Valleys. In Pahrump Valley, about 30,000 acre-ft of ground water was pumped for municipal and domestic supply and irrigation in 1997 (Division of Water Resources, State of Nevada, written commun., 1999). In Lida Valley ground water used for irrigation is less than a few hundred acre-feet per year but accounts for a significant portion of the ground water discharged from that area. Because the focus of this study is on natural ground-water discharge, and the difficulty associated with separating transpiration by natural vegetation from that of irrigated crops, ET and ground-water discharge estimates for Pahrump and Lida Valleys are not included in this report.

Annual ET is estimated for Chicago Valley, the Franklin Well area, Franklin Lake, Sarcobatus Flat, the Shoshone area, Stewart Valley, and the Tecopa/California Valley area (fig. 1). Also included in this report are estimates for Ash Meadows and Oasis Valley, which were studied primarily as part of previous efforts (Laczniaak and others, 1999; and Reiner and others, 2002, respectively). Although Death Valley is an area of substantial natural ground-water discharge, it is not included because evapotranspiration rates for salt-encrusted playas are unavailable. Evapotranspiration rates in Death Valley are being studied by the U.S. Geological Survey in cooperation with the National Park Service as part of an ongoing study. Ground water removed by processes other than ET from the major discharge areas included in this report is assumed to account for less than a few percent of the total discharge.

Evapotranspiration

ET is a process by which water from the Earth's (or a plant's) surface is transferred to the atmosphere. Estimates of mean annual ET were determined using the procedure established during recent studies of

Ash Meadows (Laczniaik and others, 1999) and Oasis Valley (Reiner and others, 2002). The underlying assumption of this procedure is that ET rates vary with the health, density, and type of the vegetation and the wetness of the soil. The procedure assumes ET rates within a discharge area can be generalized on the basis of similarities in vegetation and soil conditions. Areas of similar vegetation and soil conditions are referred to in this report as ET units. The procedure computes annual ET from each ET unit within a discharge area and sums these values to estimate the total annual ET from the discharge area. Annual ET estimates for each ET unit are computed by multiplying the acreage of the unit by an appropriate ET rate for the unit's vegetation and soil conditions. Acreage was determined through a combination of field mapping and satellite imagery analyzed using remote-sensing techniques. ET rates were estimated primarily from data collected at micrometeorological stations established in different vegetation and soil assemblages found throughout Ash Meadows (Laczniaik and others, 1999) and Oasis Valley (Reiner and others, 2002), and supplemented with values reported in Young and Blaney (1942), Walker and Eakin (1963), Weeks and others (1987), Johnson (1993), Nichols (1993), Czarniecki (1997), DeMeo and others (1999), and Nichols (2001).

Evapotranspiration Units

Past studies have shown that ET rates throughout the Great Basin region vary with vegetation and soil conditions. In general, the more dense and healthy the vegetation and the wetter the soil, the greater the rate of ET (Ustin, 1992; Nichols, 2001). Vegetation, water, and soil covers reflect incoming solar radiation differently. These differences have been used to determine type, density, and health of vegetation, and type and moisture content of soil from aerial and satellite imagery (Anderson and others, 1976, p. 2; American Society of Photogrammetry, 1983, p. 23–25; Goetz and others, 1983, p. 576–581). This process is referred to as a land-cover classification. The procedure used in this study takes advantage of (1) the relation between ET, and vegetation and soil conditions, and (2) differences in the spectral reflectance of the different vegetation and soil covers to map distinct ET units within the major discharge areas of the DVRFS, and hereafter is referred to as an ET-unit classification.

Thematic Mapper (TM) imagery was the source of the spectral data used to classify ET units. Earth-orbiting satellites equipped with sensors to detect solar-reflected and earth-emitted radiation acquire TM imagery. TM satellites collect spectral information across seven wavelength bands referred to as TM channels (fig. 4). Each channel spans a discrete part of the electromagnetic spectrum. Six of the channels (1, 2, 3, 4, 5, and 7) measure reflected radiation in the visible, near infrared, and short-wave infrared regions. These six channels were used in this study to define the spectral character or response of an area imaged by the satellite. A seventh band, TM channel 6, measures thermal energy radiated from the Earth and was not used in this study. Spectral reflectance, as acquired by TM sensors, represents an average value over an area measuring about 100 ft by 100 ft. Each square area is referred to as a pixel (picture element) and defines the spatial resolution of the imagery. Differences in the spectral response of land covers having different vegetation, soil, and moisture conditions are shown in figure 4A.

Figure 4B shows these same spectral responses as would be inferred from TM imagery. The TM imagery used to classify all discharge areas, except Sarcobatus Flat, was acquired June 13, 1992. The decision to use June 1992 imagery was based on (1) June typically being a period of high vegetation vigor, and (2) 1992 having slightly above normal in terms of precipitation. Two TM scenes (entity-id number LT5040034009216510 and LT5040035009216510) were required for complete coverage of the area of interest. A pseudo-color infrared composite of the combined TM imagery is shown in figure 5. Excessive cloud cover in the June 13, 1992, imagery precluded accurate classification of ET units in the Sarcobatus Flat area. Instead, the Sarcobatus Flat area was classified from imagery acquired June 21, 1989 (LT5040034008917210). One major difference between the procedure used here and that used by Laczniaik and others (1999) for the Ash Meadows area was in the number of dates used for the classification process. Here only June imagery was used to classify ET units, whereas, Laczniaik and others used both imagery from June and September 1999. The June imagery represented conditions of near maximum plant vigor and high moisture, whereas the September imagery represented conditions of high plant stress (dormancy) and low moisture. Based on the Ash

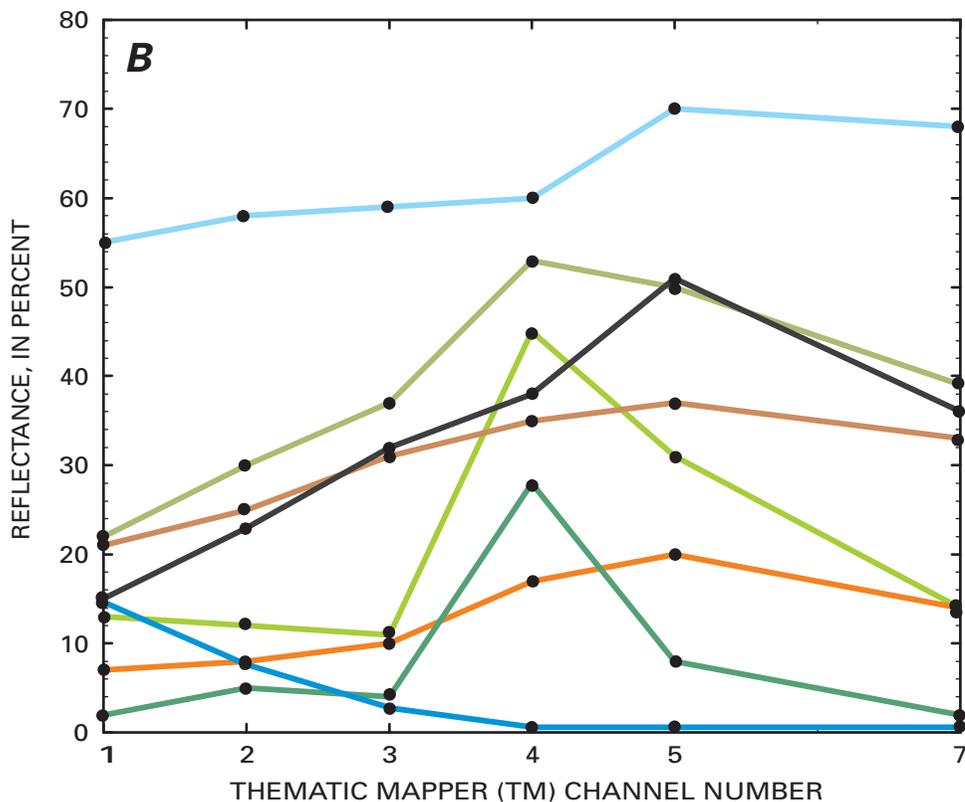
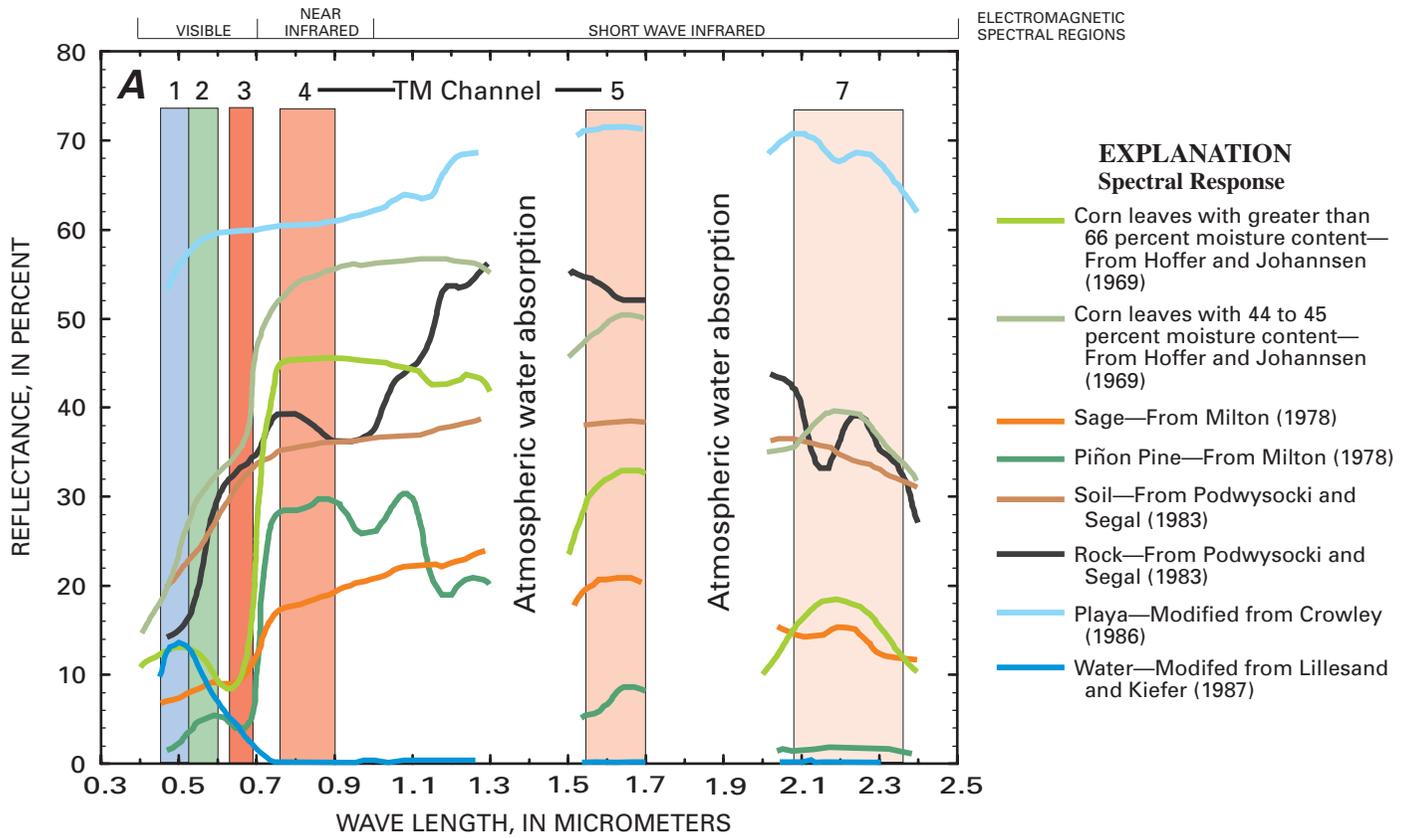


Figure 4. Spectral response of land covers having different vegetation, soil, and moisture conditions.

Meadows results, a single-date classification was assumed adequate for the discharge areas being evaluated.

The procedure used to classify ET units applied an unsupervised approach (Avery and Berlin, 1992, p. 452) to identify the unique spectral responses present within the TM imagery. Uniqueness of the spectral response was based on statistical differences between reflectance values in TM channels 1, 2, 3, 4, 5, and 7 of the imaged pixels falling within the boundaries of the major discharge areas (fig. 1). Boundaries initially were located using a raster image representing a modified soil adjusted vegetation index (Qi and others, 1994) developed from June 13, 1992, TM imagery. Boundaries were refined continually on the basis of information acquired during numerous field visits. The use of boundaries reduced the area within the TM image that needed to be classified to that of the major discharge areas. Final boundary locations are available in digital format from the USGS node of NSDI at URL: <http://water.usgs.gov/lookup/getspatial?darea>.

Each pixel was assigned a unique spectral response using the maximum likelihood classification technique (Lillesand and Kiefer, 1987, p. 685–689). This technique compares reflectance values of each pixel with the statistics defining each unique spectral response, calculates the statistical probability of a pixel being associated with each unique spectral response, and assigns each pixel to the unique response having the highest probability.

The classification process then reduces the number of unique spectral responses by grouping them on the basis of similarities in their reflectance statistics. Each group typically represents a unique land cover. Groupings were further lumped into clusters on the basis of similarities in vegetation and soil conditions. The consolidation process was an interactive procedure that involved comparing similarities in the spectral response with information gathered in the field. The process continued until a manageable number of clusters remained whereby each represented a distinct ET unit. This process ultimately resulted in 10 clusters, each representing a unique area of significant ground-water ET typified by a land cover dominated by open water, phreatophytes, or moist bare soil (table 1; fig. 6). An eleventh cluster included all spectral responses representing areas of insignificant ground-water ET typified by land covers dominated by sparse upland desert vegetation and xeric landscape. Each cluster was given a number from 0 to 10 (table 1; fig. 6).

Most spectral responses included within a cluster exhibit a characteristic shape (fig. 6). The two exceptions are clusters 4 (dense meadow and forested vegetation) and 7 (moist bare soil). Each of these clusters is defined by spectral responses exhibiting multiple characteristic shapes. The multiple patterns within each of these clusters are explained by the inclusion of more than one vegetation or soil type within the cluster. For example, cluster 4 includes dense meadow and forested vegetation, each of which exhibits a different spectral response. The two spectral responses are combined into one cluster because ET rates between the vegetation types are assumed to be similar.

The final step in the classification procedure is to assign an ET unit to each pixel. Pixels outside the boundary of a major discharge area were assigned a value of zero (an area of insignificant ground-water ET). Pixels within discharge areas were assigned a value of 1 through 10 to represent their associated cluster. This process created a raster image whereby each pixel within the image was associated with a single ET unit. The image was resampled at a finer resolution (60 ft by 60 ft) to allow for direct comparison with results presented for Ash Meadows (Laczniaik and others, 1999) and Oasis Valley (Reiner and others, 2002). The resampled image was then smoothed using a filter, which replaced spuriously classified pixels (areas defined by less than three adjacent pixels) in the image and filled single-pixel gaps by assigning the pixel to the ET unit most representative of its neighbors.

The spatial distribution of ET units as classified by the above procedure was mapped for each major discharge area (figs. 7–15). The classification shown for Ash Meadows (fig. 7) is that given by Laczniaik and others (1999) with the addition of one ET unit to represent open playa. Open playa was added to maintain consistency with other mapped discharge areas. Raster data sets representing the final classification of each discharge area are available from the USGS node of NSDI at URL: <http://water.usgs.gov/lookup/getspatial?etunit>. The acreage of each ET unit and the total ET-unit acreage are given for each discharge area in table 2.

The discharge area having the largest ET unit acreage is Sarcobatus Flat at 34,250 acres (table 2; fig. 12). Of this total, the majority is classified as sparse to moderately dense shrubs at 19,372 acres or open playa at 10,817 acres (tables 1 and 2). The dominant phreatophyte found in the shrubland area of Sarcobatus Flat is greasewood (*Sarcobatus vermiculatus*). The

Table 1. Evapotranspiration (ET) units identified and classified in major discharge areas of Death Valley regional flow system, Nevada and California

[Sources for depth to water information given in description: U.S. Geological Survey National Water Information System (accessed June 1999) and Laczniak and others (1999)]

ET-unit number ¹	ET-unit description ²
0	Area of no significant ET from ground-water source (unclassified); water table typically greater than 50 feet below land surface.
1	Area of open water , primarily reservoir or large spring pool.
2	Area of submerged aquatic vegetation ; includes sparse emergent vegetation and shallow part of open water areas; perennially flooded; water at surface.
3	Area dominated by dense wetland vegetation , primarily tall reedy and rushy marsh plants, typically tule, cattail, or giant reed; perennially flooded; water at surface.
4	Area dominated by dense meadow and forested vegetation , primarily trees, meadow grasses, or mixed trees, shrubs, and grasses; trees include saltcedar, mesquite, or desert willow; water table typically ranges from a few feet to about 20 feet below land surface; soil moist to dry.
5	Area dominated by dense to moderately dense grassland vegetation , primarily saltgrass, and/or short rushes with an occasional tree or shrub; intermittently flooded; water table typically less than 5 feet below land surface; soil wet to moist.
6	Area dominated by sparse grassland vegetation , primarily salt and bunch grasses but also includes areas of very low density shrubs (mesquite); water table typically ranges from a few feet to about 12 feet below land surface; soil dry.
7	Area dominated by moist bare soil ; vegetation very sparse, primarily grasses; intermittently flooded, water table typically near land surface throughout most of the year but in some areas declines to a maximum depth of about 5 feet below land surface during late summer and early fall; soil typically moist.
8	Area dominated by sparse to moderately dense shrubland vegetation , primarily greasewood, rabbitbrush, wolfberry, and seepweed; water table typically ranges from about 5 feet to about 20 feet below land surface; soil dry.
9	Area dominated by sparse woodland vegetation , primarily mesquite; water table typically ranges from about 10 to 40 feet below land surface; soil dry.
10	Area dominated by open playa , primarily bare soil, often encrusted with salts; water table ranges from about 5 to 40 feet below land surface; soil typically dry but can be moist for short periods after intermittent flooding.

¹ Integer value assigned to each grid cell to identify ET unit. Numbers are color coded to match ET unit as mapped in [figures 7–15](#).

² Vegetation cover descriptors: very sparse, less than 5 percent; sparse, 5 to 20 percent; moderate, 20 to 75 percent; and dense, greater than 75 percent.

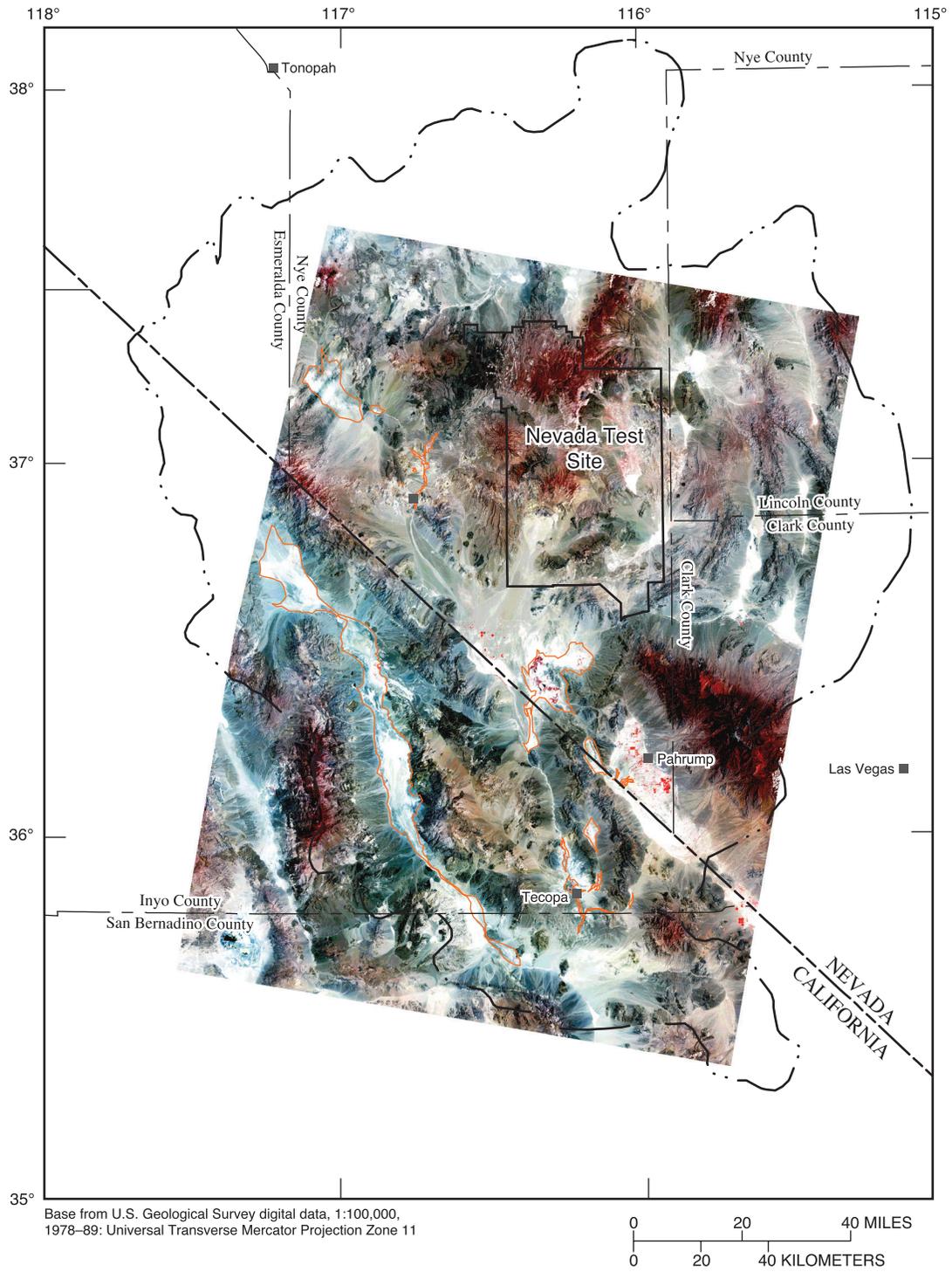


Figure 5. Pseudo-color infrared false color composite of Death Valley regional flow system.

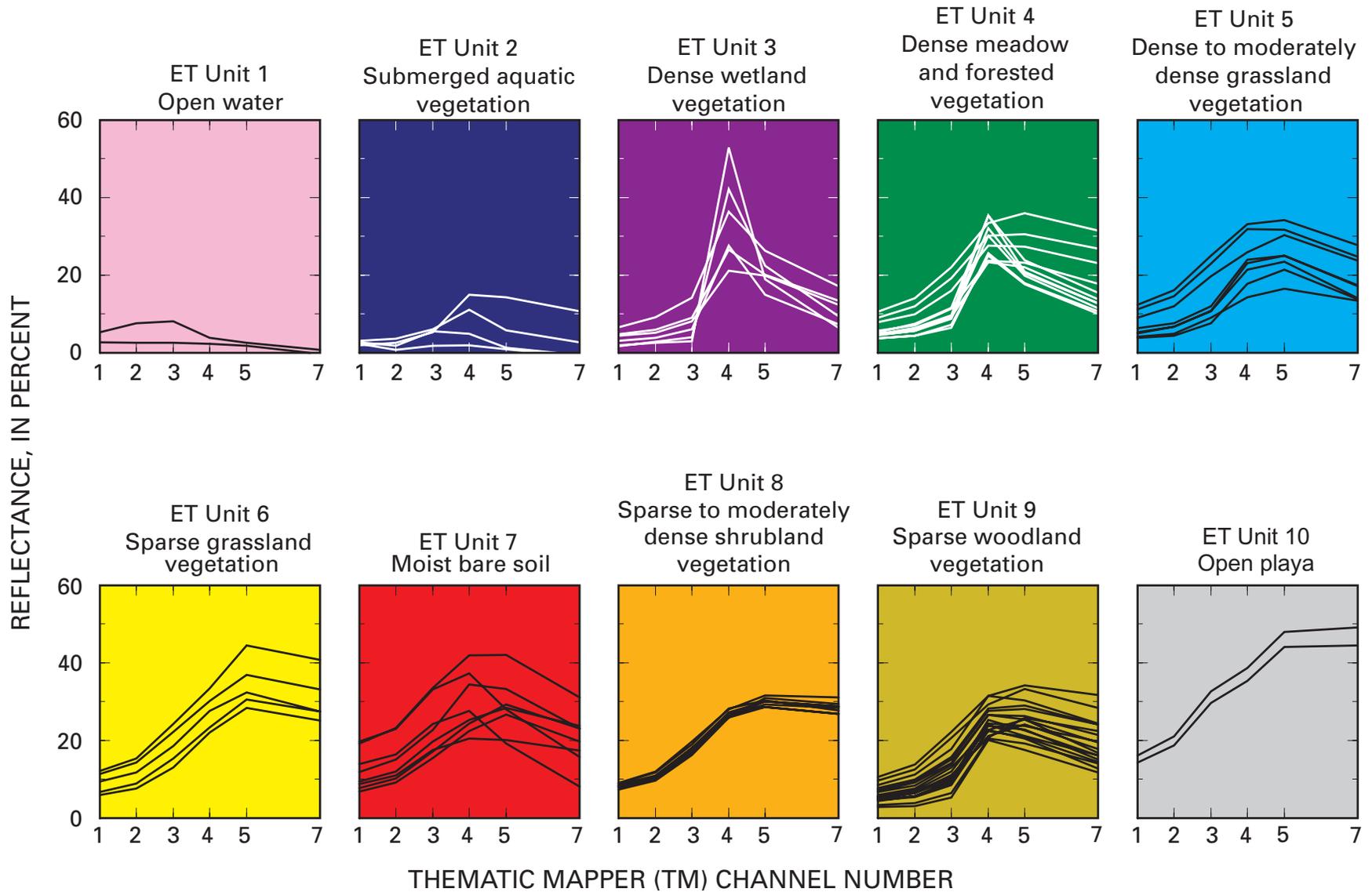


Figure 6. Spectral clusters used to classify evapotranspiration units in discharge areas of Death Valley regional flow system.

discharge area having the next largest acreage is Ash Meadows at 12,467 acres (table 2; fig. 7). This area is about 2,100 acres greater than that given by Lacznia and others (1999, table 10). This difference results from a slightly different interpretation of the Ash Meadows boundary and the inclusion of 2,241 acres of open playa as an ET unit. The largest ET unit classified within Ash Meadows is sparse grassland vegetation at 7,059 acres (tables 1 and 2). This grassland vegetation is dominated by expansive meadows of saltgrass (*Distichlis spicata* var. *stricata*). The area for sparse grassland vegetation and moist bare soil is about 150 acres less than that reported in Lacznia and others (1999, table 10). This area was not overlooked but is instead included in the adjacent Franklin Lake discharge area (figs. 1 and 10). Together these two disparities account for the difference in total acreage between the two reports. The discharge area having the smallest total acreage is the Franklin Well area at 297 acres (table 2; fig. 9). This discharge area is a narrow crescent shape that spans only about a 5-mi section of the Amargosa River along the Nevada–California border (fig. 1).

Evapotranspiration Rate and Volume Estimates

The rate at which water is evaporated and transpired is defined as the ET rate. The ET rate is a function of the vegetation, soil, soil moisture conditions, and micrometeorological factors. Because project timelines did not allow the opportunity to quantify

ET rates for each of the identified ET units, ET rates instead were estimated from those given in other studies of the general area.

The primary sources of ET rates were from recent studies of Ash Meadows (Lacznia and others, 1999, table 7) and Oasis Valley (Reiner and others, 2002). As part of these two studies, 15 sites were instrumented to collect micrometeorological data (table 3; figs. 7 and 11). These data were used to quantify ET rates for most of the ET units found within the discharge areas of interest. Daily ET rates were computed using the Bowen ratio solution of the energy-budget equation (Bowen, 1926). ET sites were instrumented from 1 to 3 years, and annual ET rates were determined from the data collected.

ET rates for each ET unit are presented as ranges in table 4. The range given for each ET unit is inclusive of all ET rates computed for Ash Meadows (Lacznia and others, 1999) and Oasis Valley (Reiner and others, 2002), and of rates estimated in other selected studies of ET throughout the general area. Data sources and the relative significance of each source used to construct the range are listed in table 4. The range includes ET rates computed for different sites during different years. The highest annual rate is near 9 ft for open water and submerged aquatic vegetation. The lowest annual rate is less than 1 ft for open playa. The difference between the minimum and maximum values varies among ET units, ranging from 0.4 ft/yr for open water and submerged aquatic vegetation to 1.8 ft/yr for sparse to

Table 2. Acreage of evapotranspiration (ET) units by major discharge area in Death Valley regional flow system, Nevada and California

Major discharge area ¹	ET-unit number ²										Total
	1	2	3	4	5	6	7	8	9	10	
Ash Meadows	158	81	385	490	1,499	7,059	554	0	0	2,241	12,467
Chicago Valley	0	0	1	0	11	0	3	0	596	1	612
Franklin Lake	0	0	0	0	9	550	148	136	0	1,488	2,331
Franklin Well area	0	0	0	13	15	0	1	0	268	0	297
Oasis Valley	1	5	40	931	645	962	81	807	0	1	3,473
Sarcobatus Flat	0	0	0	0	0	2,611	1,450	19,372	0	10,817	34,250
Shoshone area	0	1	12	62	344	659	152	22	136	0	1,388
Stewart Valley	0	1	0	22	18	49	87	9	382	2,443	3,011
Tecopa/California Valley area	2	12	16	655	379	396	894	45	1,106	8	3,513

¹ Discharge areas are shown in figure 1.

² ET unit is described in table 1 and mapped in figures 7–15.

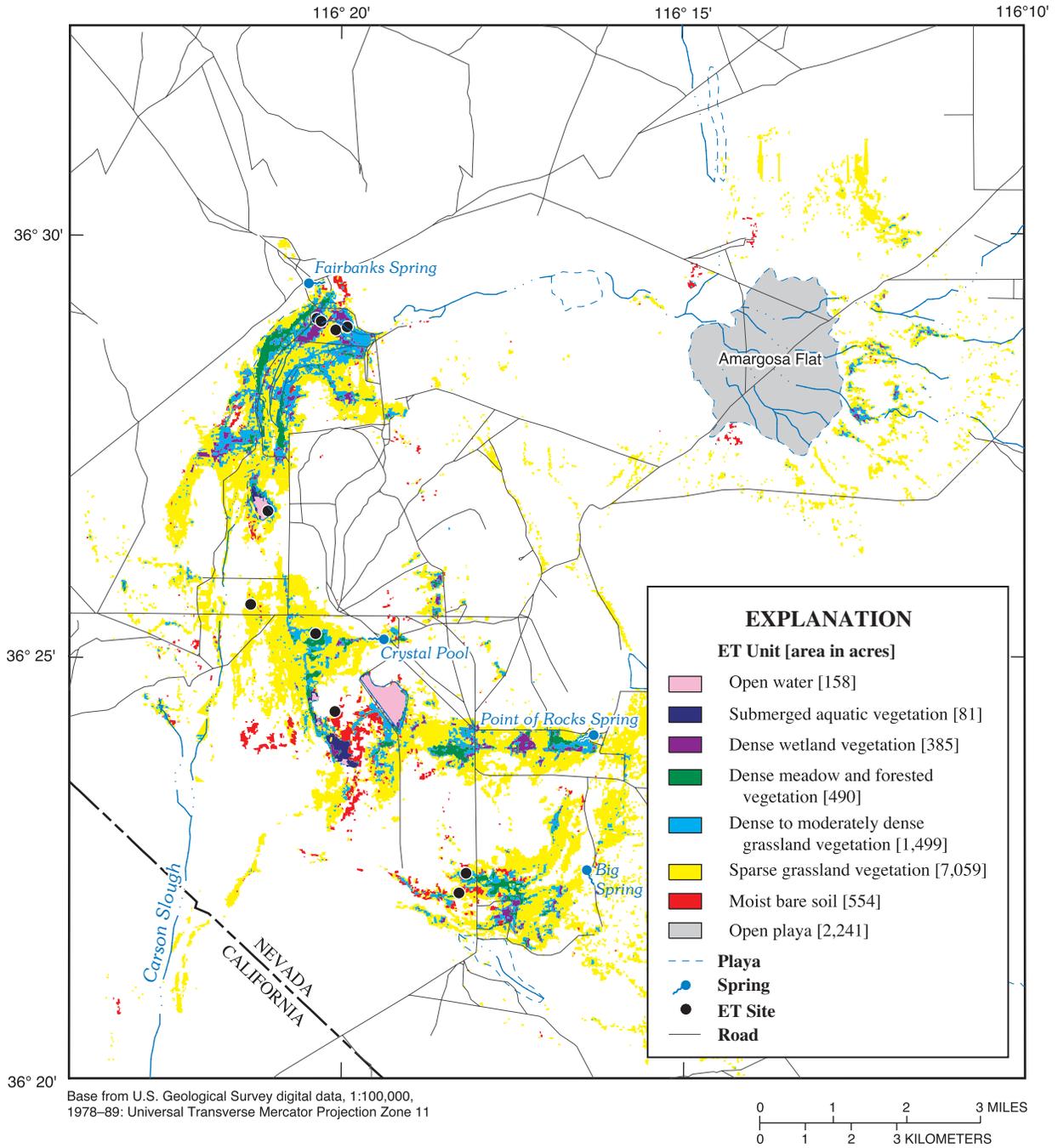


Figure 7. Classification of evapotranspiration (ET) units in Ash Meadows, Nevada and California.

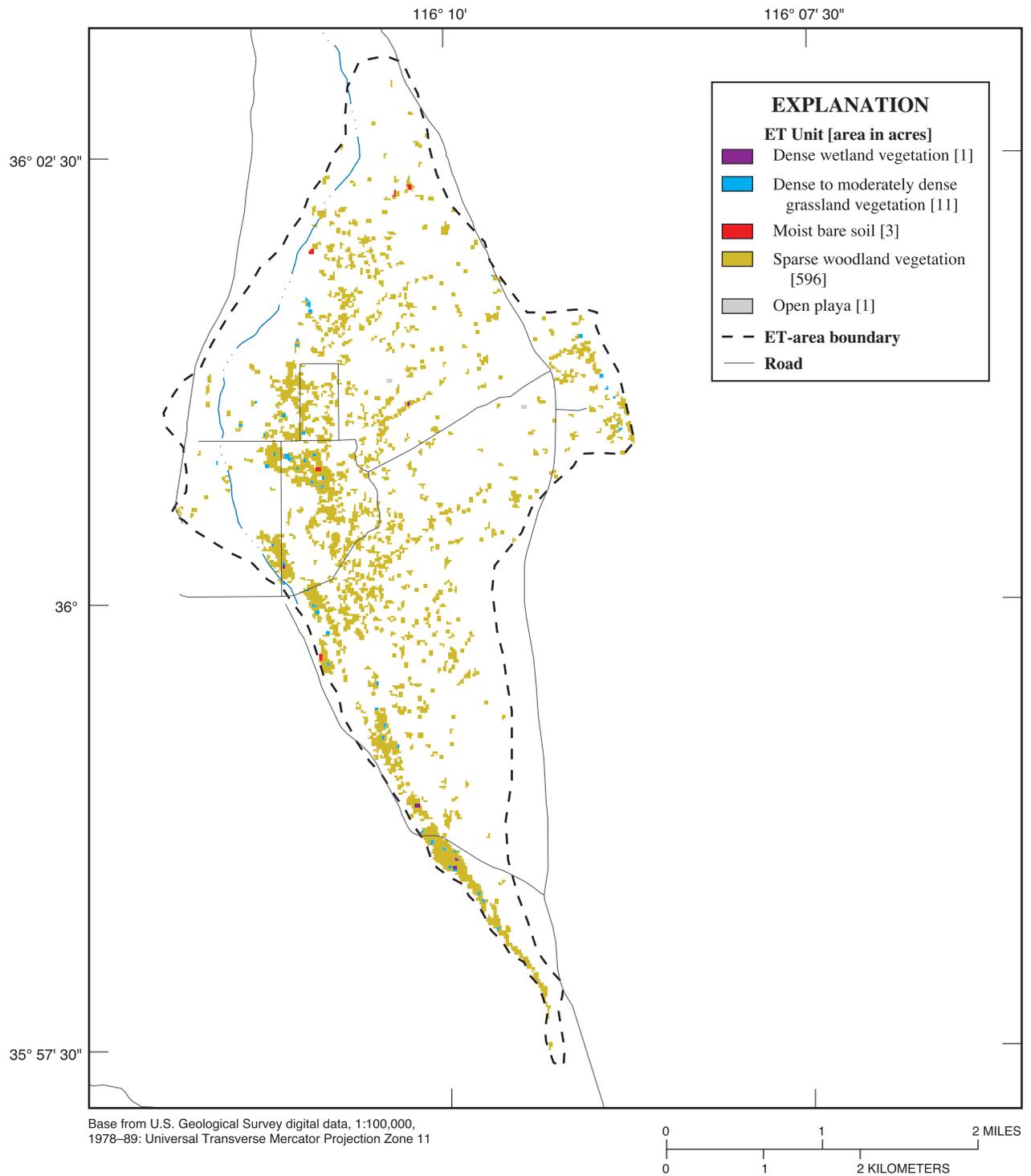


Figure 8. Classification of evapotranspiration (ET) units in Chicago Valley, California.

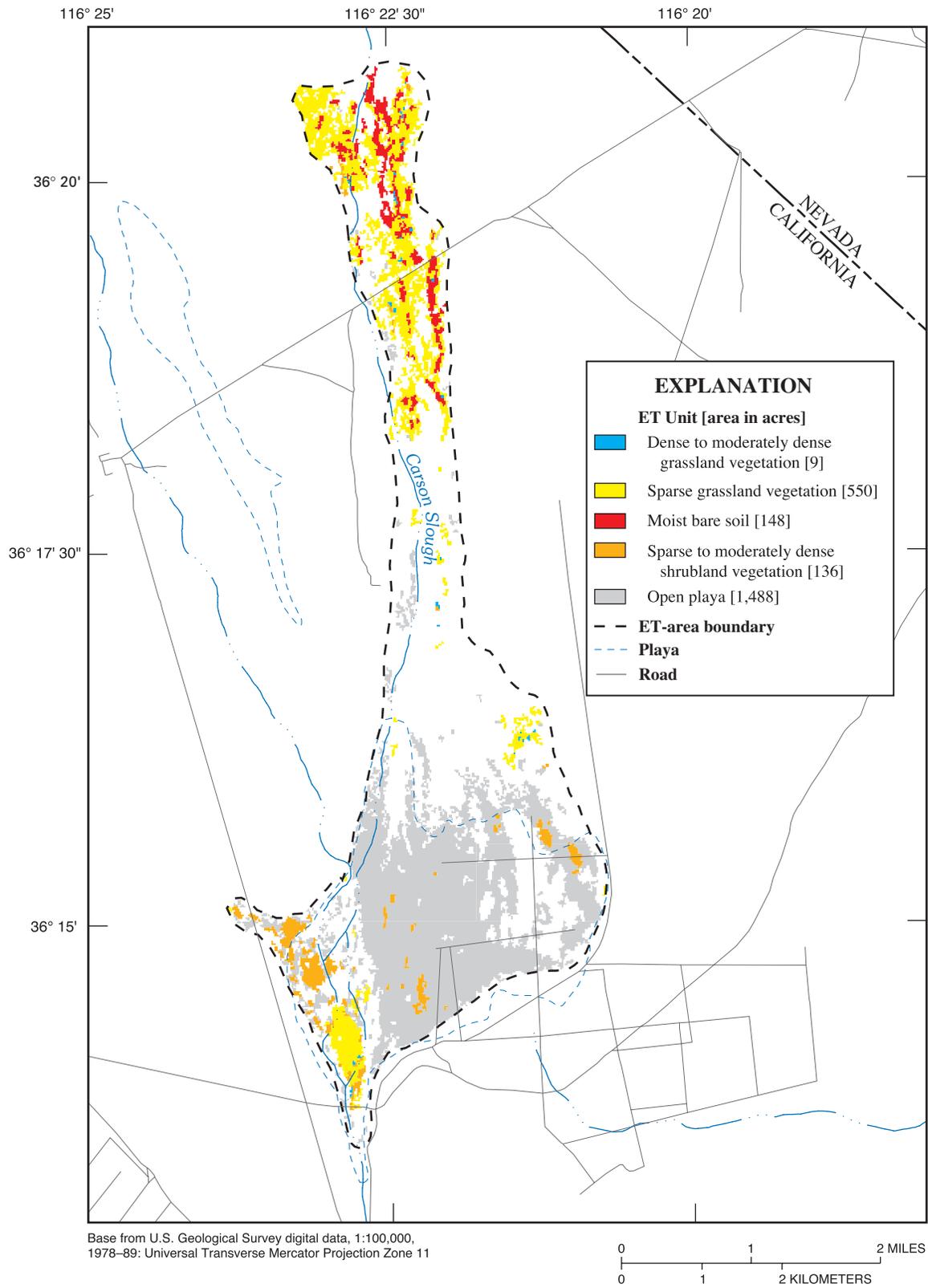


Figure 9. Classification of evapotranspiration (ET) units in Franklin Lake, California.

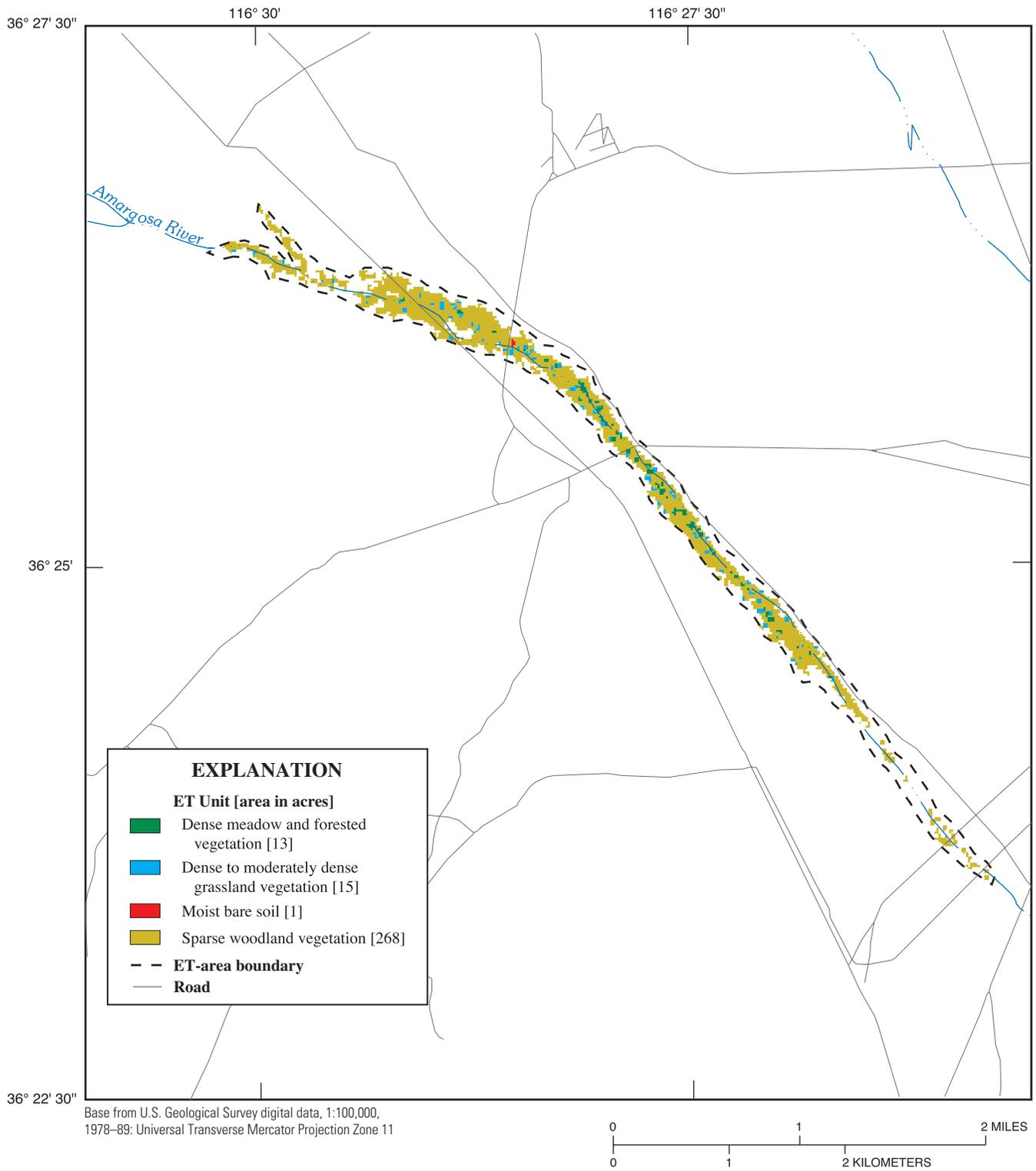


Figure 10. Classification of evapotranspiration (ET) units in Franklin Well area, California.

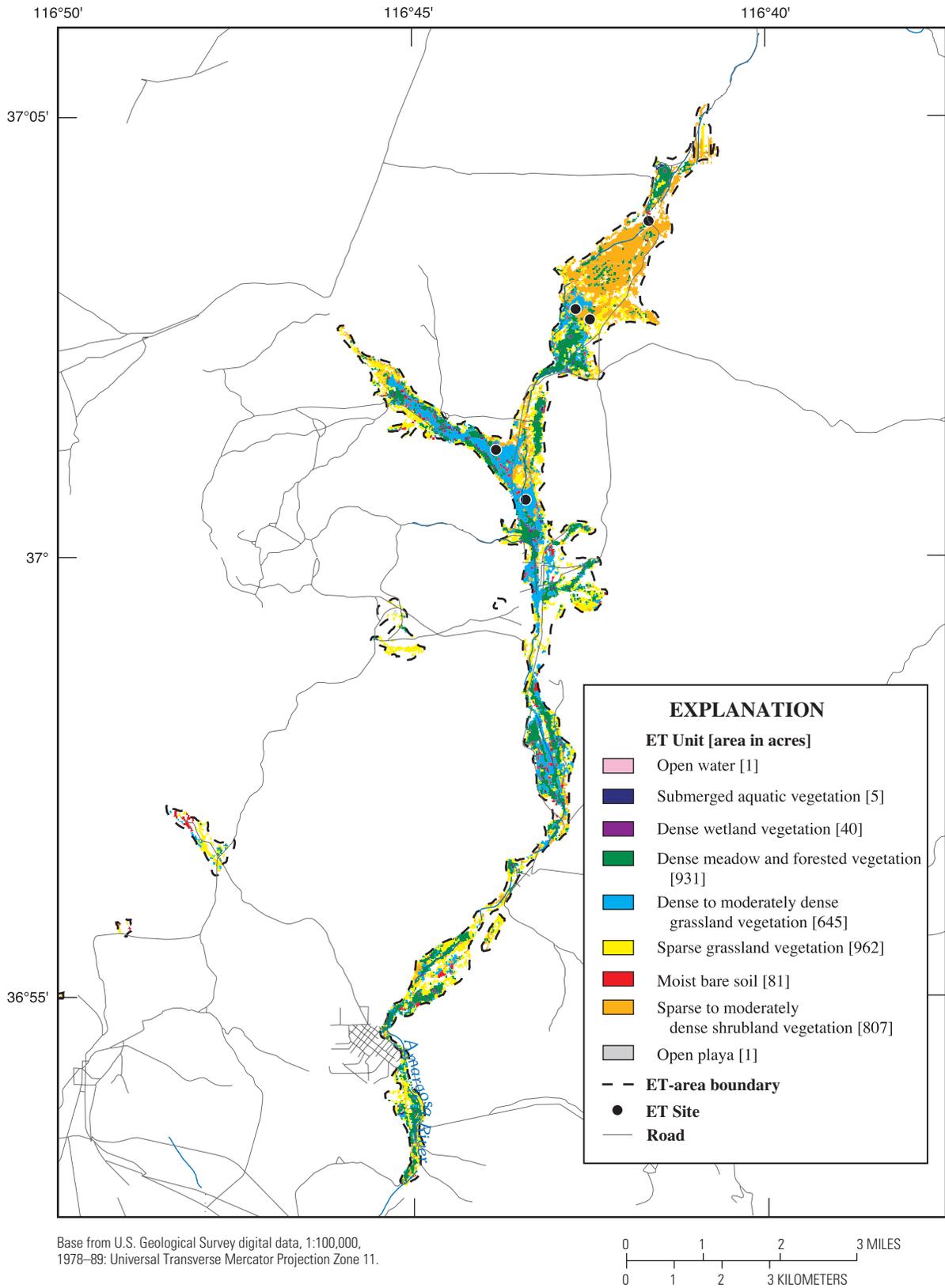


Figure 11. Classification of evapotranspiration (ET) units in Oasis Valley, Nevada.

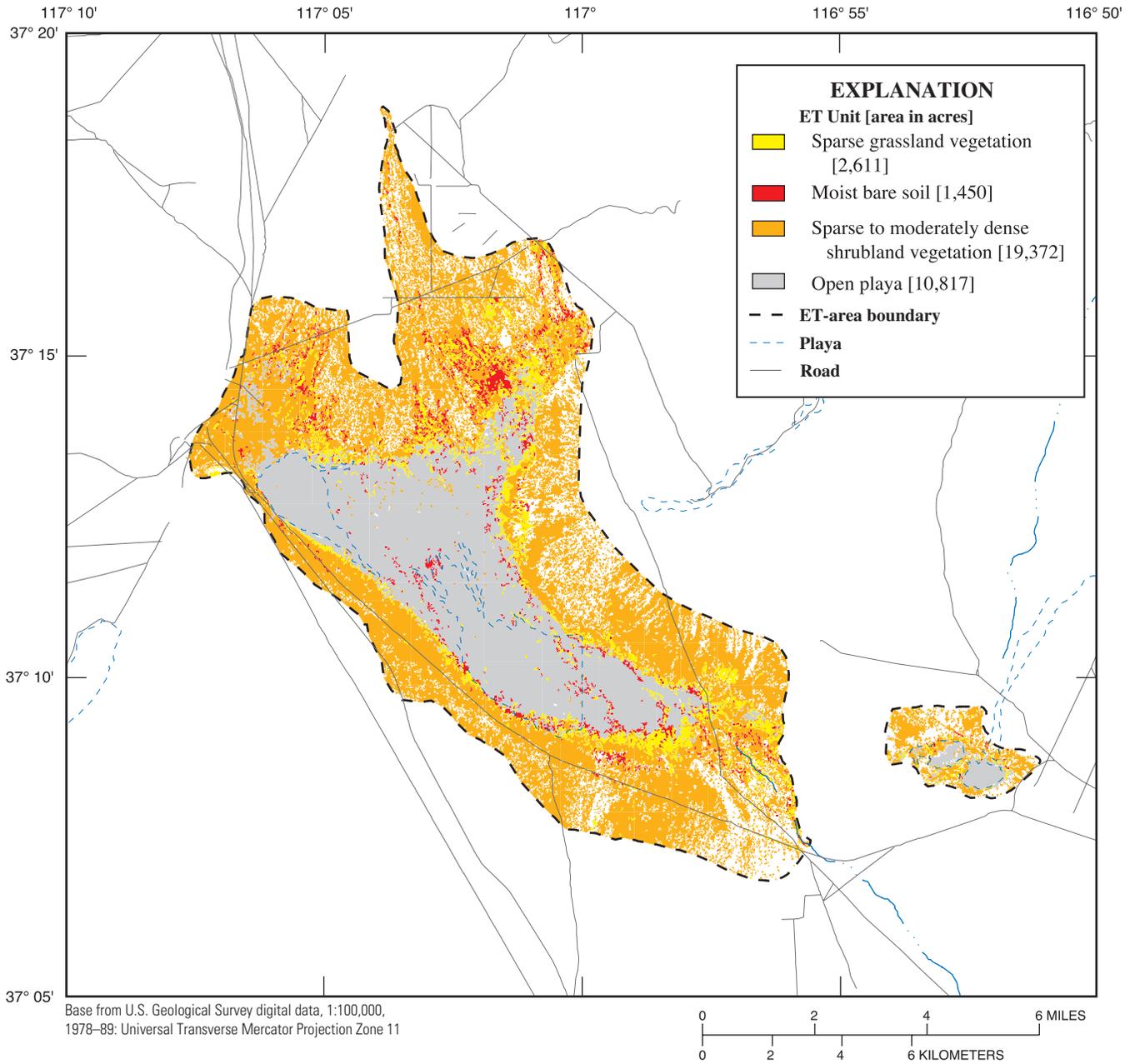


Figure 12. Classification of evapotranspiration (ET) units in Sarcobatus Flat, Nevada.

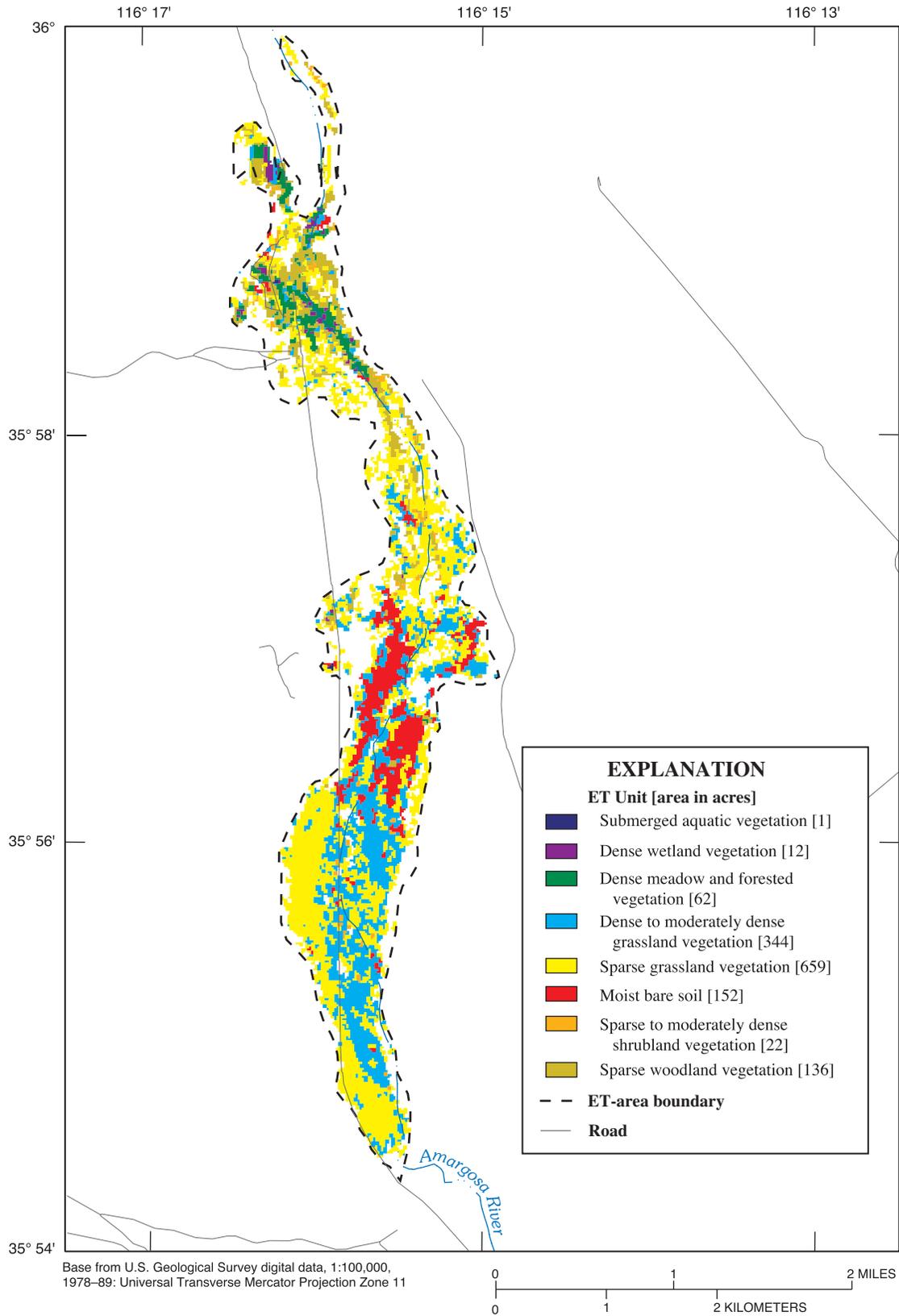


Figure 13. Classification of evapotranspiration (ET) units in Shoshone area, California.

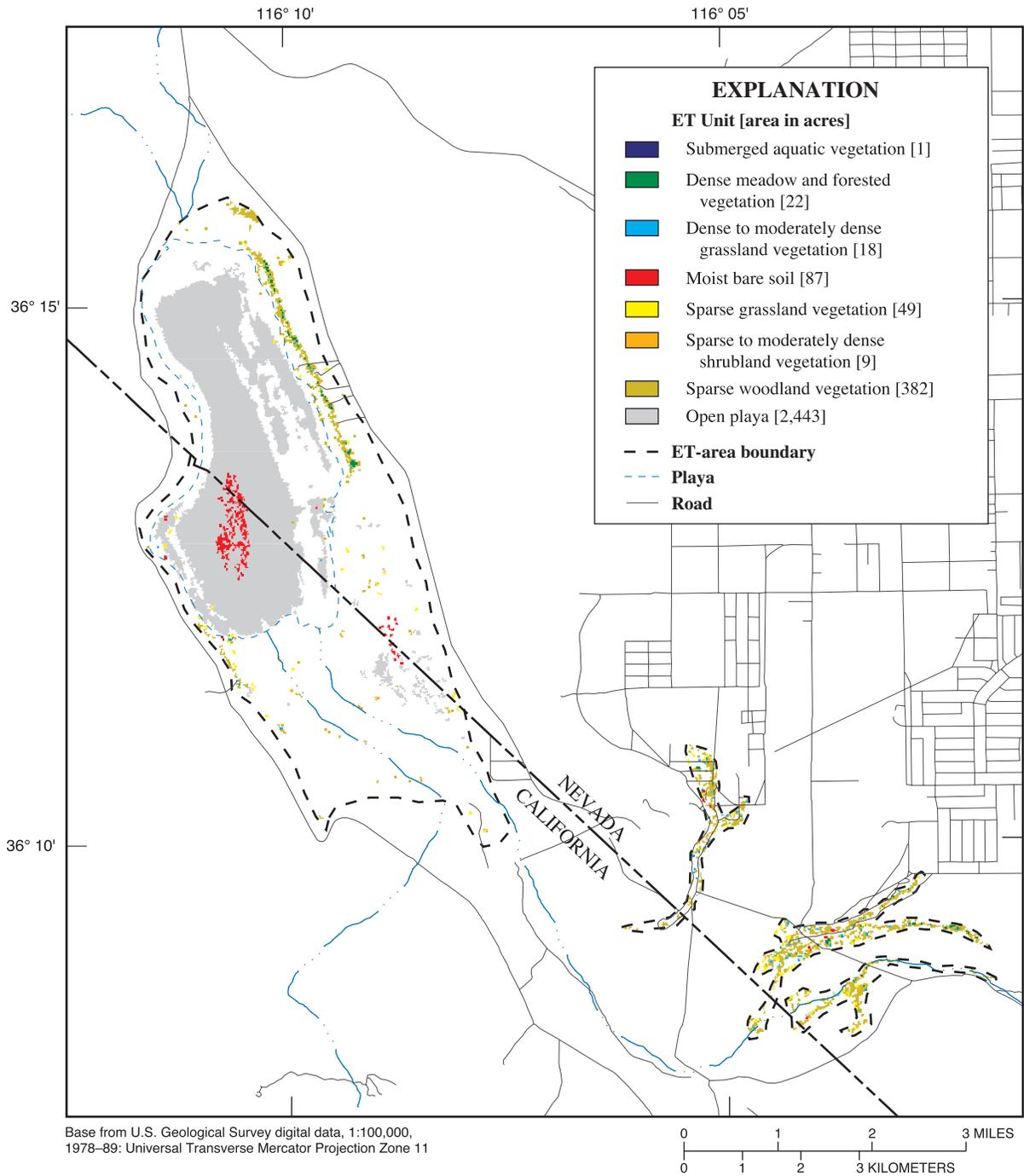


Figure 14. Classification of evapotranspiration (ET) units in Stewart Valley, California.

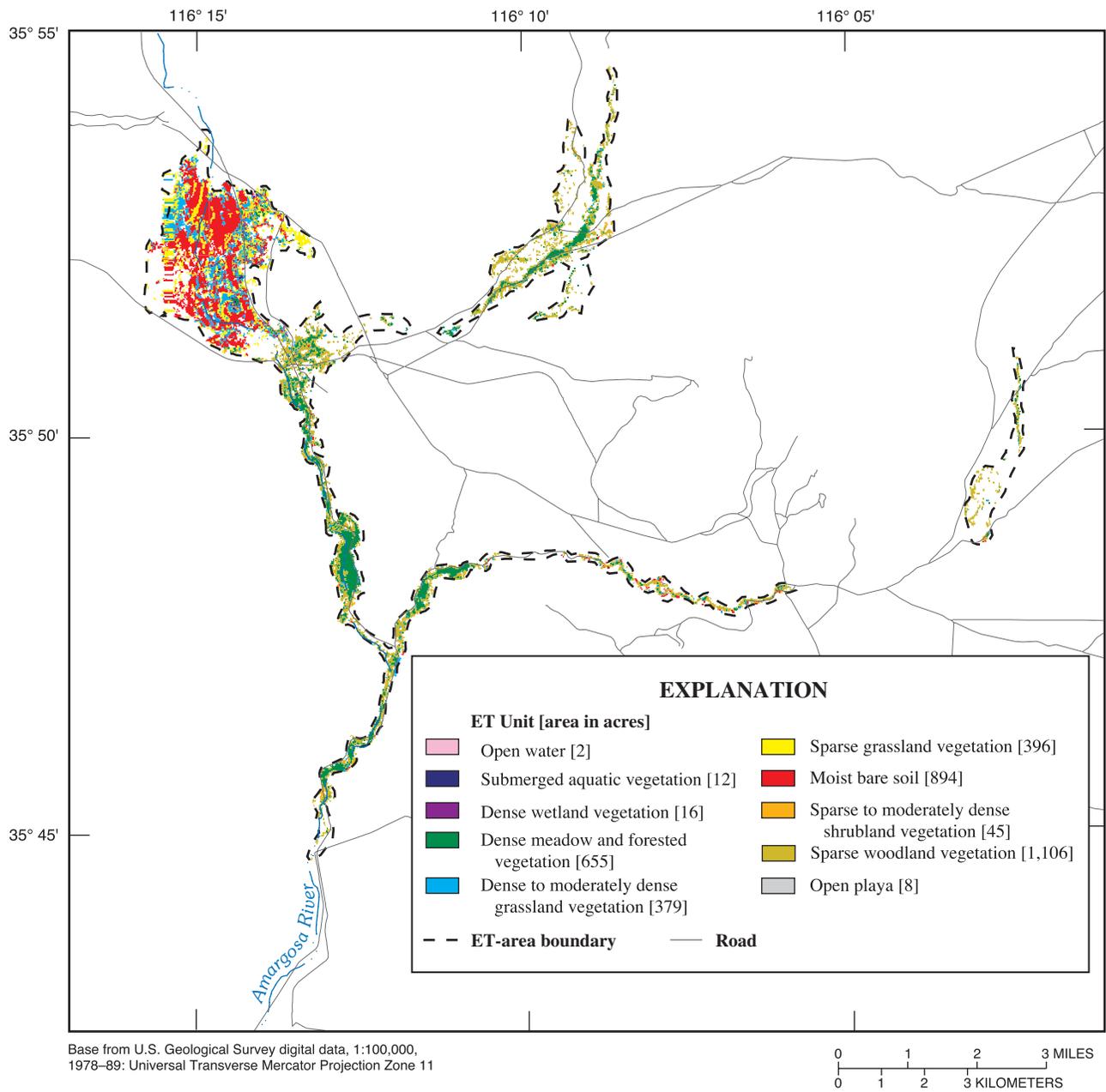


Figure 15. Classification of evapotranspiration (ET) units in Tecopa/California Valley area, California.

The estimated ET rate for each ET unit within a discharge area is listed in [table 5](#). Estimates represent the mean annual ET rate. Excluding open playa, ET rates given for Ash Meadows are those estimated by Lacznik and others (1999). Rates given for ET units in other discharge areas reflect differences between the different discharge areas in vegetation density. The density adjustment was made based on Ash Meadows values for ET units within Ash Meadows, and for other discharge areas on the mean value of the range.

Density differences between ET units were determined from a modified soil adjusted vegetation index (MSAVI; Qi and others, 1994). The MSAVI uses TM channels 3 and 4 to compute a vegetation index for each pixel in the imagery. The MSAVI increases the dynamic range of the vegetation signal by minimizing background influences from the soil. Index values range from 0 to 1, where 1 represents the strongest vegetation signal. A MSAVI was computed using the June 13, 1992 ([fig. 16](#)), and June 21, 1989, TM imagery. Raster data sets representing these indices are available from the USGS node of NSDI at <http://water.usgs.gov/lookup/getspatial?msavi89> and <http://water.usgs.gov/lookup/getspatial?msavi92>. The average MSAVI computed for each ET unit is listed in [table 6](#). Averages are based on an index computed from the imagery used to classify the discharge area, and range from 0.045 for submerged aquatic vegetation to nearly 0.395 for wetland vegetation. In general, lower values are associated with units having the sparsest vegetation and higher values with units having the densest vegetation. ET units having a greater average MSAVI were assigned higher ET rates, and those having a lesser average were assigned lower rates.

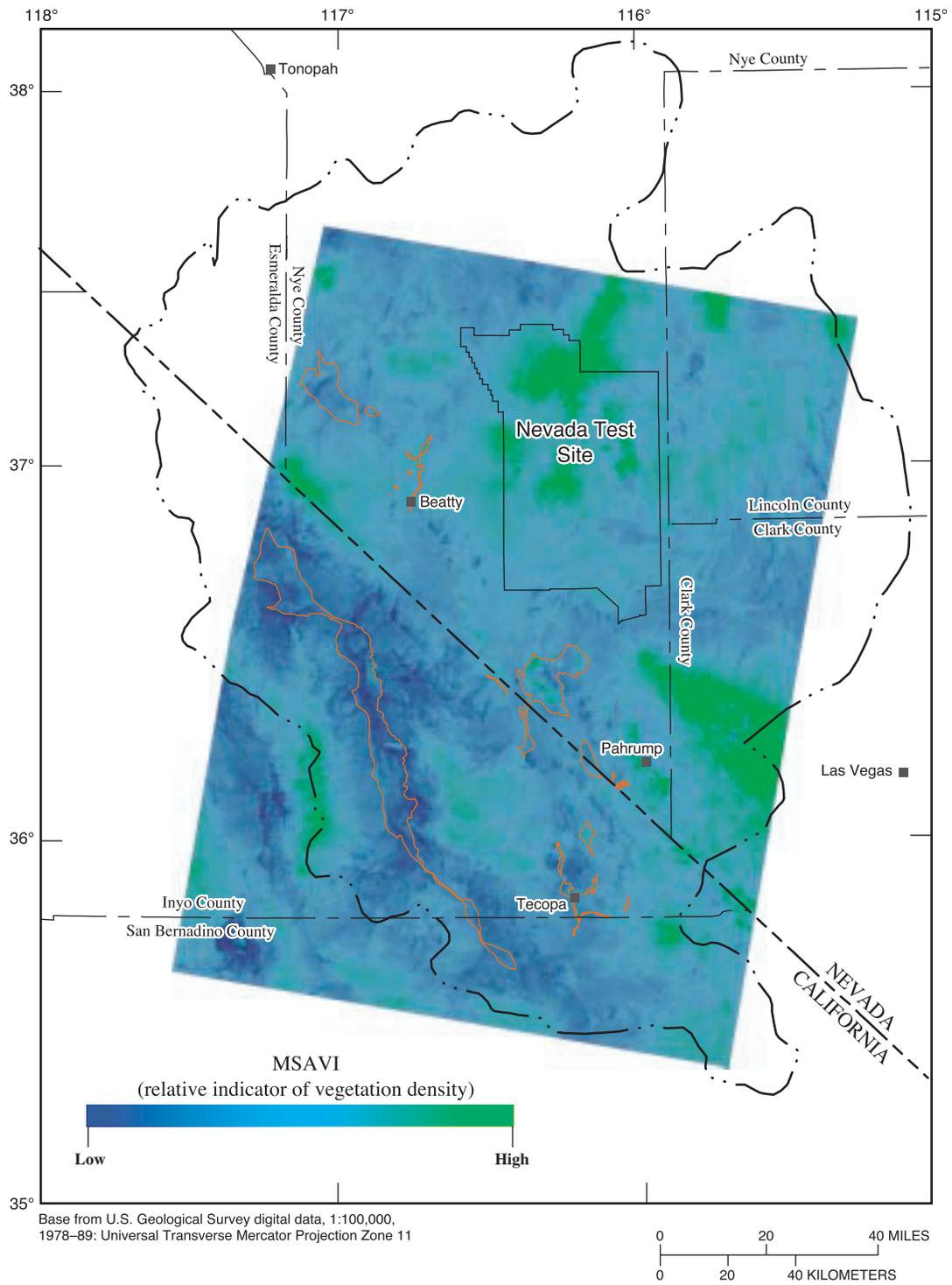
Estimates of the mean annual volume of ET from each ET unit and discharge area are listed in [table 5](#). Volume estimates for each ET unit are computed as the product of a unit's acreage and rate, and the total for each discharge area is computed by summing the individual ET-unit estimates. Mean annual ET ranges from about 450 to 30,000 acre-ft. Discharge areas having the largest mean annual ET are Ash Meadows and Sarcobatus Flat at 22,000 and 30,000 acre-ft, respectively, and those having the smallest are the Franklin Well area and Chicago Valley at 450 and 650 acre-ft, respectively.

Ground-Water Discharge

Sources contributing water to ET include direct precipitation, surface-water inflow from the surrounding drainage area, and regional ground-water inflow. Regional ground-water inflow includes recycled spring and seep flow and any diffuse upward flow from the underlying regional flow system. Assuming that all ground water discharged from the regional flow system is locally evaporated or transpired, ground-water discharge can be estimated from ET knowing the precipitation and surface-water inflow components. In this study ground-water discharge is computed assuming that (1) the surface-water contribution is negligible, and (2) the precipitation component of ET is approximately equal to precipitation falling on the area—assumptions considered reasonable for these arid valleys. Spring and seep flow are not directly accounted for in ET-based estimates, but are indirectly included by the assumption that all local surface discharge (spring and seep flow) is evaporated or recycled back into the shallow ground-water flow system, where eventually, it is evaporated or transpired from within the area.

The average annual precipitation to each discharge area was estimated from precipitation measurements made at ET sites in Ash Meadows, Oasis Valley, and Death Valley ([table 7](#)); long-term continuous data collected at nearby National Weather Service (NWS) climate stations ([tables 8 and 9](#)); and values interpreted from different published maps of precipitation (Hardman, 1965; Houghton and others, 1975, [fig. 40](#); Winograd and Thordarson, 1975, [fig. 3](#)) and a map generated by parameter-elevation regressions on independent slopes model (PRISM; Daly and others, 1994). Estimates range from 3.5 in. in the Shoshone and Tecopa/California Valley discharge areas to 6 in. in the Sarcobatus Flat and Oasis Valley discharge areas.

An estimate of mean annual ground-water discharge from each discharge area was computed by summing individual estimates of mean annual ground-water discharge computed for each ET unit within the discharge area ([table 10](#)). Mean annual ground-water discharge from each ET unit was computed as the product of an adjusted mean annual ET rate and the acreage of the ET unit. Adjustments were made to remove any water from the ET estimate that was contributed by precipitation. The remainder of the water consumed by ET is assumed to have originated from ground water. Mean annual ET rates ([table 5](#)) were adjusted for precipitation



EXPLANATION

- Ground-water discharge area
- - - Death Valley Regional Flow System boundary—From Harrill and others (1988)

Figure 16. Modified soil-adjusted vegetation index (MSAVI) of Death Valley regional flow system.

Table 3. Location and general description of sites equipped with micrometeorological instruments and used to determine evapotranspiration (ET) in Ash Meadows and Oasis Valley, Nevada

[Geographic coordinates given in degrees, minutes, seconds and referenced to North American Datum of 1927. Sites in Ash Meadows are shown as ET sites in [figure 7](#) and in Oasis Valley in [figure 11](#). Abbreviations: L, lower; M, middle; and U, upper]

Site name	Latitude	Longitude	Altitude (feet above sea level)	Period of data acquisition	Description of dominant vegetation cover ¹ and soil moisture conditions ²	ET-unit number ³
Ash Meadows						
Rogers Spring 1	36°28'53"	116°20'05"	2,256	Dec. 1993–Jan. 1996	Sparse to moderate saltgrass; soil moist in winter; very dry in summer.	6
Rogers Spring 2	36°28'55"	116°19'55"	2,253	Dec. 1993–Jan. 1996	Moderate wire and saltgrass; soil wet in winter; moist in summer.	5
Fairbanks Swamp	36°29'01"	116°20'22"	2,248	Mar. 1995–Sept. 1997	Dense cattails and reeds; surface flooded throughout year.	3
Carson Meadow	36°25'17"	116°20'23"	2,171	Mar. 1995–Mar. 1997	Dense mixed grasses, clover, and scattered shrubs; soil moist in winter, dry in summer.	4
Spring Meadow	36°25'38"	116°21'20"	2,139	June 1995–Sept. 1997	Sparse to very sparse saltgrass; soil dry in winter, very dry in summer.	6
Lower Crystal Flat	36°24'22"	116°20'06"	2,148	Nov. 1995–June 1997	Very sparse bunch grass; soil wet in winter, moist in summer.	7
Bole Spring North	36°22'27"	116°18'11"	2,180	Jan. 1996–Sept. 1997	Sparse to very sparse bunch grass; soil moist in winter, dry in summer.	7
Bole Spring South	36°22'13"	116°18'17"	2,175	Jan. 1996–Sept. 1997	Sparse saltgrass; surface periodically floods during late winter and early spring, otherwise soil moist in winter, very dry in summer.	6
Peterson Reservoir	36°26'44"	116°21'05"	2,169	March 1996–Sept. 1997	Open-water body.	1
Fairbanks Meadow ⁴	36°28'59"	116°20'18"	2,249	March 1997–June 1999	Dense saltgrass; surface periodically floods during late winter, otherwise soil wet in winter, dry in summer.	5
Oasis Valley						
Middle Oasis Valley ⁴	37°00'39"	116°43'24"	3,691	April 1997–June 1999	Moderate saltgrass, periodically flooded in early spring; soil wet in winter, dry in summer.	5
Springdale ⁴	37°01'13"	116°43'49"	3,714	July 1996–Dec. 1998	Dense marsh grass, soil moist year-round.	4
Upper Oasis Valley L ⁴	37°02'42"	116°42'29"	3,861	July 1998–June 1999	Moderate shrubs, primarily greasewood, soil dry year round.	
Upper Oasis Valley U ⁴	37°03'49"	116°41'39"	3,930	Jan. 1998–June 1999	Sparse shrubs, primarily rabbit brush and wolfberry; soil moist in winter, dry in summer.	8
Upper Oasis Valley M ⁴	37°02'49"	116°42'41"	3,856	Dec. 1998–June 1999	Sparse saltgrass; soil moist in winter, dry in summer.	6

¹ Vegetation cover descriptors: very sparse, less than 5 percent; sparse, 5 to 20 percent; moderate, 20 to 75 percent; and dense, greater than 75 percent.

² Soil moisture descriptors are presented as relative terms.

³ ET unit identified and described in [table 1](#). Numbers are color coded to match ET units as mapped in [figures 7–15](#).

⁴ Operation of site continued beyond end date specified by period of record (Reiner and others, 2002).

Table 4. Estimated evapotranspiration (ET) rates for ET units classified in major discharge areas of Death Valley regional flow system, Nevada and California

[NA, not applicable]

ET-unit number ¹	Evapotranspiration unit ²	Estimated ET rate (feet per year) ³	Source ⁴
0	Area of no significant ground-water ET (unclassified)	0.0	NA
1	Area of open water	8.4–8.8	1
2	Area of submerged aquatic vegetation	8.1–8.5	1
3	Area dominated by dense wetland vegetation	3.7–4.3	1
4	Area dominated by dense meadow and forested vegetation	3.0–4.0	1, 10, 5, 6, 7, 8
5	Area dominated by dense to moderately dense grassland vegetation	2.5–3.7	1, 10, 8
6	Area dominated by sparse grassland vegetation	.6–2.3	1, 10, 3, 11
7	Area dominated by moist bare soil	2.2–3.0	1
8	Area dominated by sparse to moderately dense shrubland vegetation	.7–2.5	10, 4, 11, 8
9	Area dominated by sparse woodland vegetation	.7–1.8	10, 8, 9
10	Area dominated by open playa	.1–.7	2, 11, 3

¹ Integer value assigned to represent evapotranspiration (ET) unit. Numbers are color coded to match ET units as mapped in [figures 7–15](#).

² See [table 1](#) for a more detailed description of ET unit.

³ Estimate given as range. Range is inclusive of rates determined at different sites, for different years, and from different sources.

⁴ Primary sources used to develop range are listed by priority of significance. Number identifies source(s): 1, Lacznia and others (1999); 2, DeMeo and others (1999); 3, Czarnecki (1997); 4, Nichols (1993); 5, Johnson (1993); 6, Weeks and others (1987); 7, Gay and Fritschen (1979); 8, Walker and Eakin (1963); 9, Young and Blaney (1942); 10, Reiner and others (2002); 11, Nichols (2001).

by subtracting an estimate of annual precipitation. Adjusted rates are listed in [table 10](#) for each discharge area.

The estimate of mean annual ground-water discharge for each discharge area, determined from adjusted ET rates, is listed in [table 10](#). Volumes range from 350 acre-ft for the Franklin Well area to 18,000 acre-ft for Ash Meadows. Differences between mean annual ET ([table 5](#)) and mean annual ground-water discharge ([table 10](#)) range from 100 acre-ft in the Franklin Well area to 17,000 acre-ft in Sarcobatus Flat. Differences on a percentage basis are largest in Sarcobatus Flat, Stewart Valley, and Franklin Lake at 57, 48, and 44 percent, respectively. The smallest difference is 14 percent in the Tecopa/California Valley area. The larger percent differences generally are associated with discharge areas dominated in part by open playa.

Estimates of mean annual ground-water discharge listed in [table 10](#) differ some from previous estimates reported in the literature (Malmberg and Eakin, 1962, p. 16–17, 25; Walker and Eakin, 1963, p. 24; Malmberg, 1967, p. 29; Blankennagel and Weir, 1973, p. 21; Winograd and Thordarson, 1975, p. 84; Czarnecki and Waddell, 1984; Czarnecki, 1997, p. 58; D’Agnese and others, 1997, p. 46). For a particular discharge area, previous estimates often show a wide range between low and high values and may vary by a factor of two or more. This wide range makes it difficult to determine differences between estimates presented here and previous estimates. The determination is further complicated because previous estimates often are reported for large areas that include multiple discharge areas and the actual extent of individual discharge areas often differs from that defined in this report. In general, ground-water discharge estimates reported here are greater than those reported in the literature for the northern discharge areas (Sarcobatus Flat and Oasis

Table 5. Mean annual evapotranspiration (ET) from major areas of ground-water discharge in Death Valley regional flow system, Nevada and California

[ft/yr, feet per year; acre-ft, acre-feet; —, no data]

Discharge area	ET component	ET-unit number ^a										Total
		1	2	3	4	5	6	7	8	9	10	
Ash Meadows	Acreage (acre)	158	81	385	490	1,499	7,059	554	0	0	2,241	12,467
	ET Rate (ft/yr)	8.6	8.5	4.0	3.4	3.5	1.3	2.6	—	—	0.5	^b 1.8
	Annual ET (acre-ft) ^c	1,400	690	1,500	1,700	5,200	9,200	1,400	—	—	1,100	22,000
Chicago Valley	Acreage (acre)	0	0	1	0	11	0	3	0	596	1	612
	ET Rate (ft/yr)	—	—	4.1	—	3.3	—	2.8	—	1.0	0.5	² 1.1
	Annual ET (acre-ft) ³	—	—	4	—	36	—	8	—	600	1	650
Franklin Lake	Acreage (acre)	0	0	0	0	9	550	148	136	0	1,488	2,331
	ET Rate (ft/yr)	—	—	—	—	2.9	1.0	2.4	1.0	—	0.5	² 0.77
	Annual ET (acre-ft) ³	—	—	—	—	26	550	360	140	—	740	1,800
Franklin Well area	Acreage (acre)	0	0	0	13	15	0	1	0	268	0	297
	ET Rate (ft/yr)	—	—	—	3.4	3.3	—	2.8	—	1.3	—	² 1.5
	Annual ET (acre-ft) ³	—	—	—	44	50	—	3	—	350	—	450
Oasis Valley	Acreage (acre)	1	5	40	931	645	962	81	807	0	1	3,473
	ET Rate (ft/yr)	8.6	8.5	4.2	3.1	3.1	1.2	2.7	1.9	—	0.5	² 2.3
	Annual ET (acre-ft) ³	9	42	170	2,900	2,000	1,200	220	1,500	—	1	8,000
Sarcobatus Flat	Acreage (acre)	0	0	0	0	0	2,611	1,450	19,372	0	10,817	34,250
	ET Rate (ft/yr)	—	—	—	—	—	0.8	2.4	1.0	—	0.5	² 0.88
	Annual ET (acre-ft) ³	—	—	—	—	—	2,100	3,500	19,000	—	5,400	30,000
Shoshone area	Acreage (acre)	0	1	12	62	344	659	152	22	136	0	1,388
	ET Rate (ft/yr)	—	8.4	4.3	3.6	2.8	1.0	2.4	1.5	1.3	—	² 1.8
	Annual ET (acre-ft) ³	—	8	52	220	960	660	360	33	180	—	2,500
Stewart Valley	Acreage (acre)	0	1	0	22	18	49	87	9	382	2,443	3,011
	ET Rate (ft/yr)	—	8.4	—	3.1	3.1	1.2	2.4	1.9	1.1	0.5	² 0.66
	Annual ET (acre-ft) ³	—	8	—	68	56	59	210	17	420	1,200	2,000
Tecopa/California Valley area	Acreage (acre)	2	12	16	655	379	396	894	45	1,106	8	3,513
	ET Rate (ft/yr)	8.6	8.5	3.9	3.3	2.9	0.9	2.4	1.8	1.2	0.5	² 2.1
	Annual ET (acre-ft) ³	17	100	62	2,200	1,100	360	2,100	81	1,300	4	7,300

^a ET unit is described in table 1 and shown in figures 7–15.^b Number is not a total but the area-weighted average of ET rates for ET units in the discharge area.^c Annual ET is rounded to two significant digits. Total is computed as the sum of the rounded values listed for each ET unit and may not be consistent with that computed as the product of the area-weighted average and the total acreage.

Table 6. Mean modified soil-adjusted vegetation index (MSAVI) value by evapotranspiration (ET) unit for major discharge areas in the Death Valley regional flow system, Nevada and California

[Mean MSAVI value derived from TM imagery acquired June 13, 1992, except in Sarcobatus Flat which was derived from TM imagery acquired June 21, 1989; MSAVI values range from 0 to 1, where 1 indicates the densest vegetation. —, not present or not applicable]

Discharge area ²	ET Unit ¹									
	1	2	3	4	5	6	7	8	9	10
Ash Meadows	—	0.065	0.195	0.290	0.180	0.125	0.105	—	—	0.065
Chicago Valley	—	—	.260	—	.165	—	.145	—	0.140	.065
Franklin Lake	—	—	—	—	.095	.080	.085	0.075	—	.070
Franklin Well area	—	—	—	.285	.170	—	.160	—	.180	—
Oasis Valley	—	.075	.305	.240	.140	.110	.120	.120	—	.075
Sarcobatus Flat	—	—	—	—	—	.070	.075	.075	—	.060
Shoshone area	—	.045	.395	.315	.080	.080	.080	.105	.175	—
Stewart Valley	—	—	—	.250	.135	.110	.070	.120	.150	.060
Tecopa/California Valley area	—	.060	.185	.275	.100	.075	.065	.115	.155	.070

¹ Discharge areas are shown in figure 1.

² ET unit is described in table 1 and mapped in figures 7–15.

Valley) and less than those reported in the literature for the southern discharge areas (Chicago Valley, Franklin Lake, Shoshone, and Tecopa/California Valley). These discrepancies relate to differences in estimates of ET acreage, ET rates, or both.

Limitations of Methodology

The accuracy of the ground-water discharge estimates is limited primarily by the assumptions inherent in the classification procedure used to define ET units and in the energy-budget methods (primarily Bowen ratio) used to compute ET rates (Laczniak and others, 1999, p. 21-22). Other factors potentially affecting the accuracy of the ground-water discharge estimates include (1) the assumption that all springflow ultimately is evaporated or transpired from within the bounds of defined discharge areas; (2) the assumption that no external surface inflow contributes water to the local ET process; (3) the short-term nature of the data used to compute mean annual ET rates; (4) the limited number of sites used to estimate mean annual ET rates; (5) uncertainty in estimates of ET rates based on computed relative density differences; and (6) the uncertainty in the adjustment applied to remove precipitation from ET estimates. Uncertainty associated with ground-water discharge estimates is determined using a Monte Carlo analysis. Results of the analysis are discussed in the appendix.

With one exception, ET units and their associated acreage were determined from TM imagery acquired on a single date (June 13, 1992). Cloud cover over Sarcobatus Flat on that date necessitated the use of TM imagery acquired on June 21, 1989, for that area. June represents a period of the year when vegetation is relatively robust, and 1992 was a year of slightly above normal precipitation. Vegetation on this date is assumed to be healthy and vigorous, and thus should be easily classified through spectral methods. Although this date is assumed to be a reasonable representation of average vegetation and soil conditions throughout the DVRFS, conditions are known to change from one year to the next (Laczniak and others, 1999, p. 31-33). Most of these changes are a consequence of changes in precipitation, which affects local vegetation, soil-moisture conditions, and the depth to the water table—all of which affect ET rates. Thus, a classification technique based on multiple years of TM imagery or some other type of spectral data would provide more confidence in the classification of ET units and acreage calculations in terms of long-term averages.

Estimates of the ET rate for each ET unit were determined using rates computed from micrometeorological data collected at 15 sites instrumented throughout Ash Meadows and Oasis Valley, and from ET rates given in other selected studies of the general area. Together these rates established a range of values that generally define the ET rate for each ET unit. In some

Table 7. Annual precipitation estimated from bulk precipitation measurements taken at evapotranspiration (ET) sites in Ash Meadows, Oasis Valley, and Death Valley areas, Nevada and California

[Geographic coordinates in degrees, minutes, and seconds and referenced to North American Datum of 1927. Abbreviation of locality: AM, Ash Meadows, Nevada; OV, Oasis Valley, Nevada; DV, Death Valley, California; U, upper; L, lower]

Site name	Latitude	Longitude	Annual precipitation (inches)	Locality
1995				
Fairbanks Meadow ¹	36°28'59"	116°20'18"	4.26	AM
1996				
Fairbanks Meadow ^{1, 2}	36°28'59"	116°20'18"	2.36	AM
1997				
Fairbanks Meadow ^{1, 2}	36°28'59"	116°20'18"	4.22	AM
Springdale ³	37°01'13"	116°43'49"	5.92	OV
1998				
Fairbanks Meadow ³	36°28'59"	116°20'18"	7.18	AM
Springdale ³	37°01'13"	116°43'49"	⁴ 12.58	OV
Bare Soil Playa ⁵	36°13'40"	116°47'07"	4.26	DV
1999				
Fairbanks Meadow ³	36°28'59"	116°20'18"	2.34	AM
Upper Oasis Valley U ³	37°03'49"	116°41'39"	4.69	OV
Upper Oasis Valley L ³	37°02'42"	116°42'29"	4.09	OV
Bare Soil Playa ⁵	36°13'40"	116°47'07"	.75	DV
Salt Playa ⁵	36°12'52"	116°46'29"	.70	DV
Pickleweed ⁵	36°17'12"	116°53'12"	1.04	DV

¹ Laczniaik and others (1999).

² Precipitation gage moved during year.

³ Reiner and others (2002).

⁴ Total based on 11 months of data.

⁵ G.A. DeMeo (U.S. Geological Survey, Las Vegas, written commun., 1999).

units, the range includes rates computed at multiple instrument sites (up to five), while in others, it includes only values reported in the literature. Some sites were instrumented for periods of up to 3 years, while others only for a single year. No sites were instrumented in the discharge areas of Chicago Valley, Franklin Lake, Franklin Well area, Sarcobatus Flat, Shoshone area, Stewart Valley, and Tecopa/California Valley area. ET rates determined over longer time periods and from additional sites in other discharge areas would help refine, improve, and provide more confidence in estimates of mean annual ground-water discharge. This is especially true of more sensitive units, such as open playa, where estimates of annual ET rates are based

on limited local data and on values reported in the literature for vegetation and soil conditions outside the study area.

The ET rate assigned to each ET unit within a discharge area was selected from a range of values based on relative differences in vegetation density. The range is intended to be a general indicator of likely ET rates expected for a particular ET unit and its width defines the variability in the vegetation and moisture conditions within a unit. For example, a larger range suggests greater changes in vegetation density. Relative density differences between ET units from different discharge areas were determined from differences in the average MSAVI value computed from TM imagery. Although

Table 8. Annual precipitation measurements from National Weather Service climate stations near the major discharge areas of Death Valley regional flow system, Nevada and California

[All precipitation data are in inches. Data are available from the National Climatic Data Center. Number given in station name is National Weather Service identifier (table 10). Symbol: —, missing data]

Calendar year	Precipitation station											
	Beatty (260714)	Beatty 8N (260718)	Sarcobatus (267319)	Goldfield (263285)	Amargosa Farms (260150)	Desert Rock WSMO (262251)	Pahrump (265890)	Indian Springs (263980)	Desert Game Range (262243)	Pahrnagat Wildlife Refuge (265880)	Penoyer (266130)	Death Valley (042319)
1948	^a 0.44	—	^a 1.06	^a 0.53	—	—	^a 0.45	^a 0.33	^a 0.70	—	—	—
1949	5.43	—	4.07	2.76	—	—	5.73	7.04	^a 6.28	—	—	—
1950	2.05	—	1.30	^a 1.53	—	—	1.88	.66	1.41	—	—	—
1951	2.60	—	3.57	^a 3.46	—	—	^a 1.28	1.84	4.21	—	—	—
1952	8.36	—	7.98	^a 7.27	—	—	^a 2.59	3.26	6.54	—	—	—
1953	.69	—	2.52	^a 1.39	—	—	^a .00	1.41	1.00	—	—	—
1954	^a 6.44	—	4.06	^a 4.20	—	—	^a .00	2.92	3.05	—	—	—
1955	4.31	—	2.29	^a .00	—	—	^a .00	1.73	3.64	—	—	—
1956	1.67	—	1.51	^a .00	—	—	^a .00	1.86	.85	—	—	—
1957	5.91	—	4.77	^a .00	—	—	^a .00	4.70	5.87	—	—	—
1958	^a 4.23	—	3.39	^a .07	—	—	^a 2.07	4.38	5.86	—	—	—
1959	3.39	—	4.41	2.87	—	—	2.79	3.50	3.94	—	—	—
1960	5.12	—	2.09	6.55	—	—	4.77	5.14	4.20	—	—	—
1961	3.29	—	^a .45	4.47	—	—	^a 2.30	2.75	5.76	—	—	^a 1.27
1962	2.67	—	—	6.23	—	—	^a 2.46	2.02	1.79	—	—	.77
1964	2.11	—	—	^a 1.69	—	—	^a 1.21	^a .00	1.11	^a 2.38	—	1.05
1965	7.33	—	—	^a 7.35	^a 1.93	—	9.12	—	7.41	7.80	—	3.32
1966	3.22	—	—	4.71	1.28	—	1.58	—	2.34	3.93	—	.79
1967	4.88	—	—	10.88	^a 1.42	—	^a 1.13	—	4.93	8.83	^a 4.86	1.39
1968	4.45	—	—	5.77	^a 3.00	—	^a .00	—	1.47	6.60	^a 1.77	^a 1.62
1969	^a 4.99	—	—	7.43	6.95	—	^a 5.41	—	7.53	8.35	7.62	3.42
1970	^a 3.00	—	—	4.32	^a 2.65	—	2.93	—	2.60	4.41	5.77	2.29
1971	^a 1.98	—	—	2.45	^a .00	—	2.04	—	1.50	3.93	^a 4.26	.96
1972	^a 1.95	^a 0.00	—	^a 3.36	^a .00	—	3.99	—	6.57	5.36	^a .00	2.25
1973	—	6.72	—	^a 4.75	^a .00	—	4.21	—	4.58	7.14	^a .00	^a 2.27
1974	—	5.80	—	^a 4.87	^a .00	—	3.35	—	4.77	5.45	^a .00	3.47
1975	—	4.84	—	5.68	^a .00	—	2.04	—	^a 5.15	^a 5.01	^a .00	1.51
1976	—	7.82	—	12.00	^a .00	—	6.73	—	6.35	8.16	^a .00	4.21
1977	—	6.44	—	7.33	^a .00	—	4.10	—	3.03	^a 4.03	^a .00	2.80
1978	—	10.80	—	13.19	^a 1.90	—	8.78	—	6.88	^a 9.46	^a .00	4.03

Table 8. Annual precipitation measurements from National Weather Service climate stations near the major discharge areas of Death Valley regional flow system, Nevada and California—Continued

Calendar year	Precipitation station											
	Beatty (260714)	Beatty 8N (260718)	Sarcobatus (267319)	Goldfield (263285)	Amargosa Farms (260150)	Desert Rock WSMO (262251)	Pahrump (265890)	Indian Springs (263980)	Desert Game Range (262243)	Pahranagat Wildlife Refuge (265880)	Penoyer (266130)	Death Valley (042319)
1979	—	4.71	—	4.68	2.16	—	7.17	—	^a 1.95	^a 0.00	^a 0.00	1.76
1980	—	6.01	—	6.84	6.24	—	^a 7.83	—	4.54	^a 5.58	^a .00	3.19
1981	—	3.73	—	^a 7.14	3.20	—	^a 3.62	—	3.26	6.58	^a .00	1.77
1982	—	6.07	—	7.75	4.30	—	5.11	—	4.87	9.87	^a .00	2.62
1983	—	11.49	—	^a 5.60	10.37	—	8.55	—	6.34	11.54	^a .00	4.54
1984	—	^a 5.45	—	^a .27	8.80	^a 10.76	8.41	—	10.38	8.86	^a .00	^a 2.17
1985	—	2.79	—	^a 2.18	1.00	4.77	2.45	—	2.58	4.84	^a .00	.87
1986	—	5.22	—	^a 4.11	3.75	6.92	3.23	—	4.66	4.37	^a 1.28	1.00
1987	—	7.38	—	9.23	8.18	8.56	7.41	—	7.21	7.71	^a 7.40	3.74
1988	—	6.21	—	8.17	5.59	5.84	^a 4.03	—	3.12	5.32	^a 4.36	3.84
1989	—	2.43	—	^a 5.00	.72	1.25	2.11	—	1.51	2.23	^a .00	.45
1990	—	4.92	—	5.07	2.58	4.88	5.75	—	2.82	^a 5.48	^a 3.14	1.28
1991	—	5.15	—	5.74	3.40	5.55	5.10	—	4.07	5.52	^a .83	^a 1.33
1992	—	7.37	—	^a 6.66	6.09	6.35	8.34	—	8.85	9.66	^a .00	3.36
1993	—	5.71	—	^a 6.08	5.68	7.55	7.71	—	4.91	6.82	^a 1.74	2.58
1994	—	3.44	—	5.86	2.27	3.28	3.18	—	^a 3.38	^a 4.36	^a 2.75	^a .45
1995	—	8.45	—	9.50	5.27	^a 7.90	^a 5.43	—	4.53	7.53	^a 5.91	3.27
1996	—	5.60	—	5.16	2.83	^a 2.02	^a 1.90	—	2.74	3.33	8.00	1.15
1997	—	6.64	—	6.38	3.33	^a .00	3.77	—	4.40	5.55	3.53	2.81
1998	—	12.62	—	^a 10.88	8.71	^a .00	8.49	—	9.23	^a 7.71	^a 5.45	4.26
1999	—	^a 4.69	—	^a 4.91	2.35	^a 4.44	3.64	—	3.66	^a 3.97	3.15	1.18
Annual mean	4.03	6.33	3.50	6.50	4.57	5.49	4.94	3.07	4.41	6.53	5.61	2.37

^a Value represents incomplete year.

Table 9. National Weather Service (NWS) climate stations near major discharge areas of Death Valley regional flow system, Nevada and California

[Geographic coordinates (latitude and longitude) are given in degrees and minutes and referenced to North American Datum of 1927]

Station name	NWS number	Latitude	Longitude	Altitude (feet above sea level)
Beatty	260714	36°55'	116°45'	3,300
Beatty 8N	260718	37°00'	116°43'	3,550
Sarcobatus	267319	37°16'	117°01'	4,020
Goldfield	263285	37°42'	117°14'	5,690
Amargosa Farms	260150	36°34'	116°28'	2,450
Desert Rock WSMO	262251	36°37'	116°01'	3,300
Pahrump	265890	36°12'	115°59'	2,670
Indian Springs	263980	36°35'	115°41'	3,120
Desert Game Range	262243	36°26'	115°22'	2,920
Pahrangat WL Refuge	265880	37°16'	115°07'	3,400
Penoyer	266130	37°37'	115°47'	4,800
Death Valley	042319	36°28'	116°52'	-190

this technique is likely to provide a reasonable estimate of the ET rate, more accuracy could be achieved either by classifying more ET units or by improving the relation between density and ET rate. Either approach requires the installation of additional ET sites to measure local ET rates.

ET estimates were adjusted to remove contributions of local precipitation. Some uncertainty is inherent in the precipitation adjustment. This uncertainty is attributed to errors in estimating the average annual precipitation and to the uncertainty in the actual amount of local precipitation included in estimated ET rates. No adjustment was made to remove any surface-water inflow contribution from the estimate of the ET rate. The aridity of the area is likely to produce minimal surface-water inflows, but in discharge areas where precipitation is relatively high and the surrounding drainage area is dominated by highlands, such as Oasis Valley and Sarcobatus Flat, surface-water inflow may be a more substantial component of the estimated ET rate. The decision not to adjust ET rates for local surface-water inflow was based on the scarcity of available data and may be partly responsible for the larger ground-water discharge estimates given for the northern discharge areas than reported in the literature. Additional data defining the amount of water contributed by the many sources to local ET would greatly improve estimates of ground-water discharge.

Unclassified areas are assumed to be zones of no substantial ground-water discharge. This assumption, although strongly supported by the lack of vegetation, dryness of soil, and greater depths to the water table (generally exceeding 50 ft), could result in some error when estimating ground-water discharge. Even over vast areas, volumetric losses of ground water are likely to be negligible considering the very low ET rates (less than 1.0×10^{-4} ft/yr) associated with these areas (Andraski, 1997, table 2).

Ground-water discharge estimates as reported include only water lost through evaporation and transpiration, and do not include any water that may be leaving discharge areas through subsurface flow. Until additional data become available from which to define water-level distributions and spatial variations in the hydraulic properties controlling ground-water movement throughout the shallow local and deeper regional flow systems, reliable estimates of subsurface outflow are not possible. Considering the high potential for subsurface outflow from many of the major discharge areas in the DVRFS, the estimate of ground-water discharge presented should be considered a minimum value for the total amount of ground water exiting a discharge area.

Table 10. Mean annual ground-water discharge from major areas of ground-water discharge in Death Valley regional flow system, Nevada and California

[ET, evapotranspiration; ft/yr, feet per year; acre-ft, acre/feet; —, no data]

Discharge area	Precipitation adjustment (ft/yr)	ET component	ET-unit number ¹										Total
			1	2	3	4	5	6	7	8	9	10	
Ash Meadows	0.35	Area (acre)	158	81	385	490	1,499	7,059	554	0	0	2,241	12,467
		Adjusted ET rate (ft/yr)	8.25	8.15	3.65	3.05	3.15	0.95	2.25	—	—	0.15	² 1.4
		Ground-water discharge (acre-ft) ³	1,300	660	1,400	1,500	4,700	6,700	1,200	—	—	340	18,000
Chicago Valley	0.35	Area (acre)	0	0	1	0	11	0	3	0	596	1	612
		Adjusted ET rate (ft/yr)	—	—	3.75	—	2.95	—	2.45	—	0.65	0.15	² 0.70
		Ground-water discharge (acre-ft) ³	—	—	4	—	32	—	7	—	390	0	430
Franklin Lake	0.33	Area (acre)	0	0	0	0	9	550	148	136	0	1,488	2,331
		Adjusted ET rate (ft/yr)	—	—	—	—	2.57	0.67	2.07	0.67	—	0.17	² 0.43
		Ground-water discharge (acre-ft) ³	—	—	—	—	23	370	310	91	—	250	1,000
Franklin Well area	0.35	Area (acre)	0	0	0	13	15	0	1	0	268	0	297
		Adjusted ET rate (ft/yr)	—	—	—	3.05	2.95	—	2.45	—	0.95	—	² 1.2
		Ground-water discharge (acre-ft) ³	—	—	—	40	44	—	2	—	260	—	350
Oasis Valley	0.50	Area (acre)	1	5	40	931	645	962	81	807	0	1	3,473
		Adjusted ET rate (ft/yr)	8.10	8.00	3.70	2.60	2.60	0.70	2.20	1.40	—	0.01	² 1.8
		Ground-water discharge (acre-ft) ³	8	40	150	2,400	1,700	670	180	1,100	—	0	6,200
Sarcobatus Flat	0.50	Area (acre)	0	0	0	0	0	2,611	1,450	19,372	0	10,817	34,250
		Adjusted ET rate (ft/yr)	—	—	—	—	—	0.30	1.90	0.50	—	0.01	² 0.38
		Ground-water discharge (acre-ft) ³	—	—	—	—	—	780	2,800	9,700	—	110	13,000
Shoshone area	0.29	Area (acre)	0	1	12	62	344	659	152	22	136	0	1,388
		Adjusted ET rate (ft/yr)	—	8.11	4.01	3.31	2.51	0.71	2.11	1.21	1.01	—	² 1.5
		Ground-water discharge (acre-ft) ³	—	8	48	200	860	470	320	27	140	—	2,100
Stewart Valley	0.35	Area (acre)	0	1	0	22	18	49	87	9	382	2,443	3,011
		Adjusted ET rate (ft/yr)	—	8.05	—	2.75	2.75	0.85	2.05	1.55	0.75	0.15	² 0.33
		Ground-water discharge (acre-ft) ³	—	8	—	60	50	42	180	14	290	370	1,000
Tecopa/ California Valley area	0.29	Area (acre)	2	12	16	655	379	396	894	45	1,106	8	3,513
		Adjusted ET rate (ft/yr)	8.31	8.21	3.61	3.01	2.61	0.61	2.11	1.51	0.91	0.21	² 1.8
		Ground-water discharge (acre-ft) ³	17	99	58	2,000	990	240	1,900	68	1,000	2	6,400

¹ ET unit is described in table 1 and shown in figures 7–15.² This number is not a total ET rate but the area-weighted average of the adjusted ET rates for the ET units in the discharge area.³ Ground-water discharge is rounded to two significant digits. Total is computed as the sum of the rounded values listed for each ET unit and may not be consistent with that computed as the product of the area-weighted average and the total acreage.

SUMMARY

The Death Valley regional flow system (DVRFS) is one of the larger ground-water flow systems in the southwest United States and includes much of southern Nevada and the Death Valley region of eastern California. Although situated in an arid region, large quantities of ground water discharge from a few localized areas where geologic and hydrologic conditions are such that ground water is pushed upward to the surface and is discharged from springs and seeps. The water emerging from these sources supports a great diversity of vegetation and wildlife and provides habitat for a variety of endangered plant and animal species. Some water flowing from these sources evaporates shortly after emerging, some water flows to pools and reservoirs where it too evaporates, and the remainder of the water infiltrates downward from drainage channels to recharge the underlying shallow flow system. Moisture held in the local soils and water contained in the shallow flow system sustains thriving populations of phreatophytes year round. Together these water sources and local plant communities create distinct oases within an expansive, generally barren desert region.

The Nevada Test Site (NTS), with an area of about 1,375 mi², is centrally located within the DVRFS. Some of the ground water emerging from the major discharge areas of the DVRFS originates or flows beneath the NTS. The NTS historically has been used for testing nuclear devices and currently is a potential location for the permanent disposal of high-level nuclear waste in United States. The U.S. Department of Energy, as mandated by Federal and State regulators, is evaluating the risk associated with contaminants that have been or may be introduced into the subsurface as a consequence of any past or future activities at the NTS. To assess risk, the potential for contaminant transport away from the site must be determined. Because subsurface contaminants can be transported away from the NTS by ground water, components of the ground-water budget that provide information on the flow of ground water are of extreme interest. One such component influencing ground-water flow is regional ground-water discharge. Because some uncertainty exists as to the amount of ground water that discharges from the DVRFS, studies were initiated to re-evaluate and better quantify estimates of ground-water discharge. This report documents the result of a previous study that applies a technique to estimate ground-water discharge at Ash Meadows to other major

discharge areas of the DVRFS. The discharge areas evaluated as part of this effort include Chicago Valley, the Franklin Well area, Franklin Lake, Sarcobatus Flat, the Shoshone area, Stewart Valley, and the Tecopa/California Valley area. Although Ash Meadows and Oasis Valley were studied as part of independent efforts, estimates for these two areas are included in this report for purposes of completeness. This effort does not include Death Valley, which currently is being evaluated as part of a separate study by the U.S. Geological Survey in cooperation with the National Park Service.

Ground-water discharge is estimated throughout much of the DVRFS from a rigorous quantification of evapotranspiration. This approach assumes that all ground-water discharge is evaporated or transpired from within the discharge area. Although the approach does not account for springflow directly, it assumes that all springflow is evaporated or recycled back into the shallow flow system where later it is transpired or evaporated. Any recycled water not locally evaporated or transpired is not accounted for in the discharge estimate. Mean annual evapotranspiration (ET) from each discharge area is calculated as the sum of mean annual ET estimates determined for each of the ET units (an area of similar vegetation and moisture conditions) present within the discharge area. Mean annual ET from an ET unit is computed as the product of the unit's acreage and its estimated annual ET rate.

ET units are defined on the basis of differences in spectral-reflectance characteristics. Spectral differences are determined from thematic mapper (TM) imagery acquired June 21, 1989, and June 13, 1992. Except for Sarcobatus Flat, all discharge areas are classified from the June 13, 1992, imagery. TM imagery acquired June 21, 1989, is used to classify ET units in Sarcobatus Flat because excessive cloud cover over this area diminished the overall quality of the 1992 imagery. Ten unique ET units are identified that include areas of open playa, sparse to dense vegetation, moist bare soil, and open water. Sarcobatus Flat has the largest ET unit acreage of any discharge area at 34,250 acres. Of this total, 19,372 acres are classified as sparse to moderately dense shrubs and 10,817 acres as open playa. The dominant phreatophyte found in Sarcobatus Flat is greasewood (*Sarcobatus vermiculatus*). The discharge area having the next largest acreage is Ash Meadows at 12,467 acres. The largest ET unit classified within Ash Meadows is sparse grassland vegetation at 7,059 acres

and is dominated by expansive meadows of saltgrass (*Distichlis spicata* var. *stricata*). The smallest discharge area is the Franklin Well area with 297 acres.

Mean annual ET rates for seven ET units are estimated from daily ET rates computed primarily from micrometeorological data collected at 15 sites instrumented in Ash Meadows and Oasis Valley. Sites were instrumented for periods of 1 to 3 years as part of separate studies of Ash Meadows and Oasis Valley and are located within 7 of the 10 ET units identified. In the three ET units not instrumented, annual ET rates were estimated from rates reported in the literature. Together, these values established a range of ET rates for each ET unit. Estimated mean annual ET rates range from 0.5 ft/yr for open playa to nearly 9 ft/yr for open water. Mean annual ET estimates for discharge areas range from 450 acre-ft in the Franklin Well area to 30,000 acre-ft in Sarcobatus Flat.

Ground-water discharge for each ET unit is estimated as the product of an adjusted ET rate and the ET unit acreage. The ET rate is adjusted to remove local precipitation components from the ET estimate. Precipitation ranges from 3.5 in. in the Shoshone and Tecopa discharge areas to 6 in. in the Sarcobatus Flat and Oasis Valley discharge areas. Estimates of mean annual ground-water discharge range from 350 acre-ft in the Franklin Well area to 18,000 acre-ft in Ash Meadows. Ground-water discharge estimates generally are greater for the northern discharge areas (Sarcobatus Flat and Oasis Valley) than previous estimates reported in the literature and less for the southern discharge areas (Franklin Lake, Shoshone, and Tecopa/California Valley).

The accuracy of ground-water discharge estimates is limited primarily by the assumptions inherent in the classification procedure and in the energy-budget methods (primarily Bowen ratio) used to compute daily ET. Other factors potentially affecting the accuracy of the ground-water discharge estimates include (1) the assumption that all springflow ultimately is evaporated or transpired from within the discharge area; (2) the assumption that no external surface-water inflow contributes to local ET measurements; (3) the short period of record used to compute mean annual ET rates; (4) the limited number of local sites used to estimate mean annual ET rates; (5) the uncertainty associated with computing local ET estimates on computed relative density differences; and (6) the uncertainty in the adjustment applied to remove precipitation from ET estimates. Multi-year classifications, longer-term data

acquisition, and a greater number of local ET-site installations would help refine, improve, and provide more confidence in estimates of mean annual ground-water discharge. Additional measurements of ET rates in areas such as open playa where estimates of annual ET rates are based on values reported in the literature for vegetation and soil conditions outside the study area or on limited data, and a better understanding of the contribution of precipitation and surface-water inflow to measured ET rates are vital to establishing a higher degree of confidence in estimates of ET and ground-water discharge. The stated conjecture that these data would improve confidence is supported by results of the uncertainty analysis presented in the appendix. Results indicate that the largest uncertainties are most often coincident with discharge areas dominated by open playa and that the adjustment made to remove the nonground-water component (precipitation and surface-water inflow) from the measurements of ET rates is an extremely sensitive parameter.

The estimate of ground-water discharge presented in this report includes water lost through evaporation and transpiration but does not include water that may exit discharge areas as subsurface flow. Absent better estimates of subsurface flow, annual estimates of ground-water discharge presented here should be considered a minimum value of total outflow from a discharge area.

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APPENDIX

Uncertainty and Sensitivity Analysis of Annual Estimates of Ground-Water Discharge for Death Valley Regional Flow System

By Gaius J. Roemer, GeoTrans

The objective of this analysis is to quantify the uncertainty associated with estimates of annual ground-water discharge from the nine discharge areas addressed in the main body of this report (table 2). Because discharge estimates are expected to be used as targets for calibrating ground-water flow models, it was considered beneficial to quantify, at least in a general sense, the uncertainty associated with these estimates. The results of this evaluation can be used to better understand the uncertainty associated with each discharge estimate and to appropriately weight estimates to best constrain model results. For this effort, uncertainty is evaluated through Monte Carlo simulations performed using Crystal Ball (Decisioneering, 1996, Crystal Ball Version 4.0), a Microsoft Excel add-in. The input parameters required for Monte Carlo analysis include an estimate of the annual precipitation rate for each discharge area, and estimates of the acreage and ET rate for each of the ET units within a discharge area. Although 10 ET units are identified throughout the study area (table 1), not all are present within each discharge area. In this analysis, a total of 141 input parameters were used to evaluate the uncertainty in estimates of discharge: 61 acreages, 61 ET rates, and 9 precipitation rates (tables 5 and 7).

For this basic analysis, each input parameter is assumed to be characterized by a normal distribution centered about a mean value. The mean of each input parameter is the value of the parameter as estimated in tables 5 and 7. The spread about the mean is described by the coefficient of variability (CV), which is defined as the standard deviation divided by the mean. The CV used for the acreage of each ET unit listed in table 7 is assumed to be 10 percent. A CV value of 10 percent is considered reasonable based on accuracy assessments of about 90 percent for ET-unit classifications of the Ash Meadows and Oasis Valley discharge areas (Laczniaak and others, 1998; and Reiner and others, 2002, respectively). The CV for each ET rate was determined from ranges listed in table 4, and for each precipitation rate from measurements given in tables 8 and 9.

CV values for the ET rate and precipitation parameters are computed on the assumption that ranges represent ± 2 standard deviations of a normal population (95 percent of the measurements are contained in the range; table 11).

The general procedure used to quantify the uncertainty in estimates of ground-water discharge consists of four basic steps:

1. Randomly selecting a value from the normal distribution of each input parameter. For example, 17 random values were selected for the 8 ET units in Ash Meadows discharge area (tables 5 and 7): 8 for acreage, 8 for ET rate, and 1 for the precipitation rate;
2. Adjusting the selected ET rate of each ET unit by subtracting the selected precipitation rate;
3. Multiplying the adjusted ET rate by its selected acreage to calculate the annual rate of ground-water discharge from each ET unit; and
4. Summing ET-unit discharges and acreages within each discharge area to compute totals.

This process of randomly selecting values for each input parameter and calculating the annual discharge from each discharge area is termed a realization. A test performed to determine the number of realizations needed to produce stable estimates of the standard deviation generated frequency distributions and statistics of annual discharge from sample sizes of 500; 1,000; 2,000; and 3,000 realizations. Test results indicated that a sample size of 1,000 realizations was sufficient.

To exemplify this general procedure, Monte Carlo results of the Ash Meadows discharge area (table 12, figs. 17 and 18A) are discussed in some detail. Table 12 gives the mean acreage, ET rate and precipitation rate calculated from 1,000 realizations. The table also gives the adjusted ET rate (calculated by subtracting the mean precipitation rate from the mean ET rate), annual ET,

and annual ground-water discharge computed for each ET unit and the totals. Summary tables for the eight other discharge areas are given in [tables 13–20](#).

The eight input parameters to which annual ground-water discharge in Ash Meadows is most sensitive are shown in [figure 17](#). The sensitivity of each parameter is measured by rank correlation (correlation based on ranks rather than on values). Because ET units 6 and 5, respectively, are the largest contributors to ground-water discharge at Ash Meadows, their parameters have the greatest effect on the estimate. The sensitivity of the precipitation rate always is negative because it is subtracted from the ET rate to calculate ground-water discharge.

The five input parameters from each of the nine discharge areas having the greatest effect on annual ground-water discharge are shown in [figure 17](#). Because the precipitation rate is an essential component in calculating ground-water discharge from every ET unit, it always is one of the more sensitive input parameters. Typically the precipitation rate and the ET rate associated with the largest ET unit are the two most sensitive parameters. The lone exception is Oasis Valley, where ET units 4 and 6 have the largest acreage but the ET rate associated with ET unit 8 is the most sensitive parameter. This anomaly can be explained in part by (1) the low CV of the ET rate for ET unit 4 (0.07) relative to that of ET units 6 and 8 (0.29 and 0.28, respectively; [table 11](#)), and (2) the high ET rate of ET unit 8 relative to ET unit 6 (1.92 and 1.19 ft/yr, respectively, [table 16](#)).

Descriptive statistics and a frequency chart generated from 1,000 realizations of annual ground-water discharge for Ash Meadows are listed in [table 21](#) and shown in [figure 18A](#). The mean, standard deviation, and the coefficient of variability (CV) of the simulated ground-water discharge for all nine discharge areas are compared in [table 22](#). Assuming that CV is a reasonable estimator of the relative uncertainty (larger values represent a greater uncertainty), the discharge estimates for Oasis Valley and the Tecopa/California Valley area are most certain (0.12 and 0.11, respectively) and those for Stewart Valley and Sarcobatus Flat are least certain (0.42 and 0.48, respectively). The largest CV values (those greater than 0.20) are associated with discharge areas made up of five or fewer ET units and/or with discharge areas dominated by open playa (ET unit 10). Sarcobatus Flat, which has only four ET units and 10,771 acres of open playa, has a CV of 0.48.

Whereas, Oasis Valley, which has nine ET units and only one acre of open playa, has a CV of 0.12. In general, the smaller the number of ET units the greater the uncertainty. Where fewer units are present, the greater uncertainty can be explained by the fact that deviations from the mean are less likely to compensate each other. In discharge areas dominated by open playa, the CV of the discharge estimate is high ([table 22](#)) because (1) the CV of the unit itself is high (0.38, [table 11](#)), and (2) the ET rate is low and usually not much greater than the estimated precipitation rate. Because ET and precipitation rates are similar in magnitude, the CV of their difference is larger than the CV of either parameter. [Figure 18 \(B–I\)](#) presents descriptive statistics and frequency charts generated from 1,000 realizations of ground-water discharge in the eight other discharge areas.

An analysis of the Ash Meadows discharge area was used to examine the uncertainty associated with the classification procedure. In this analysis, classified ET units similar in terms of their spectral response ([fig. 6](#)) were correlated using a correlation coefficient of -1.0. Differences in the standard deviations generated from correlating different combinations of three similar ET units in Ash Meadows are listed in [table 22](#). The three correlated ET units are ET unit 3 (dense wetland vegetation), ET unit 4 (dense meadow and forested vegetation), and ET unit 5 (dense to moderately dense grassland vegetation). The small differences (less than 2 percent, [table 22](#)) in the standard deviation of the first three combinations are attributed to (1) the decreased likelihood that the sampled area will be much greater than or less than their respective means, and (2) similarities in their ET rates. The program does not allow correlation of ET units 3, 4, and 5 by a coefficient of -1.0. Instead, the program determines the maximum allowable correlation coefficient at -0.5. Because effects of correlation are shown by this example to be minimal, any further efforts to correlate ET units were considered unnecessary.

An additional analysis was performed to evaluate uncertainty related to the assumption of a 10 percent CV for ET-unit acreage. Five Monte Carlo simulations of 1,000 realizations using CV values for acreage of 10, 20, 30, 40, and 50 percent resulted in similar mean values having standard deviations that varied nearly proportionally with changes in the input CV. These results indicate that the predicted uncertainty in the estimate is nearly proportional to the CV of the acreage.

Table 11. Values of coefficient of variability (CV) used in Monte Carlo analysis to simulate ground-water discharge: (A) evapotranspiration (ET) rate; and (B) precipitation rate

[Min., minimum; Max., maximum; R. Width, range width; SD, standard deviation; and CV, coefficient of variability]

A.

ET Unit	Evapotranspiration rate (feet per year)					
	Min.	Max.	R. Width	SD ¹	Mean ²	CV
1	8.4	8.8	0.4	0.1	8.6	0.01
2	8.1	8.5	.4	.1	8.3	.01
3	3.7	4.3	.6	.15	4.0	.04
4	3.0	4.0	1.0	.25	3.5	.07
5	2.5	3.7	1.2	.3	3.1	.10
6	.6	2.3	1.7	.425	1.45	.29
7	2.2	3.0	.8	.2	2.6	.08
8	.7	2.5	1.8	.45	1.6	.28
9	.7	1.8	1.1	.275	1.25	.22
10	.1	.7	.6	.15	.4	.38

B.

Discharge area	Precipitation rate (feet per year)					
	Mean	Min.	Max.	R. Width	SD ¹	CV
Ash Meadows	0.35	0.29	0.41	0.12	0.03	0.09
Chicago Valley	.35	.27	.43	.16	.04	.11
Franklin Wells area	.35	.29	.41	.12	.03	.09
Franklin Lake	.33	.26	.41	.16	.04	.11
Oasis Valley	.5	.42	.58	.16	.04	.08
Sarcobatus Flat	.5	.33	.67	.34	.085	.17
Shoshone area	.29	.21	.37	.16	.04	.14
Stewart Valley	.35	.27	.43	.16	.04	.11
Tecopa/California Valley area	.29	.21	.37	.16	.04	.14

¹ Standard deviation assumes that the range represents 95 percent of the sample population.

² Mean used in analysis is the value given for input parameter in tables 5 and 7.

The accuracy of the uncertainty and sensitivity analyses can be improved by better quantifying the errors associated with the parameters used to compute ground-water discharge. The errors most significant to the procedure used are those errors inherent in the methods used to measure ET-unit acreage and to calculate ET rates. Of particular significance are errors associated with using spectral data from Thematic Mapper imagery to classify ET units and the Bowen ratio solu-

tion to compute ET rates. Having a more thorough and rigorous analysis of these errors would result in more realistic formulations of the distributions describing the acreage and ET rate of the individual ET units. Although these more accurate error estimates would improve estimates of uncertainty, they probably would not have a substantial effect on estimates of annual ground-water discharge.

Table 12. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Ash Meadows

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	159.2	81.0	385.2	486.2	1,497.8	7,075.3	553.1	0.0	0.0	2,238.2	12,477
ET rate (feet per year)	8.60	8.50	4.00	3.41	3.49	1.29	2.59	—	—	.50	—
Annual ET (acre-feet)	1,369.1	688.0	1,540.8	1,657.9	5,227.3	9,127.1	1,432.5	—	—	1,119.1	22,162
Adjusted ET rate (feet per year)	8.25	8.15	3.64	3.06	3.14	0.93	2.24	—	—	0.15	—
Annual ground-water discharge (acre-feet)	1,313.4	660.1	1,402.1	1,487.8	4,703.1	6,580.0	1,238.9	—	—	335.7	17,721

Table 13. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Chicago Valley

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	0.0	0.0	1.4	0.0	11.3	0.0	2.7	0.0	595.1	1.0	611
ET rate (feet per year)	—	—	4.10	—	3.30	—	2.79	—	1.00	.50	—
Annual ET (acre-feet)	—	—	5.7	—	37.3	—	7.5	—	595.1	.5	646
Adjusted ET rate (feet per year)	—	—	3.75	—	2.95	—	2.44	—	0.66	0.16	—
Annual ground-water discharge (acre-feet)	—	—	5.2	—	33.4	—	6.6	—	390.1	.2	435

Table 14. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Franklin Lake

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	0.0	0.0	0.0	0.0	9.0	548.3	147.6	135.7	0.0	1,494.0	2,335
ET rate (feet per year)	—	—	—	—	2.89	1.00	2.39	1.01	—	.50	—
Annual ET (acre-feet)	—	—	—	—	26.0	548.3	352.8	137.1	—	747.0	1,811
Adjusted ET rate (feet per year)	—	—	—	—	2.56	0.67	2.06	0.68	—	0.18	—
Annual ground-water discharge (acre-feet)	—	—	—	—	23.0	367.4	304.1	92.3	—	269.0	1,056

Table 15. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Franklin Well area

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	0.0	0.0	0.0	12.9	15.2	0.0	0.5	0.0	268.0	0.0	297
ET rate (feet per year)	—	—	—	3.40	3.29	—	2.80	—	1.32	—	—
Annual ET (acre-feet)	—	—	—	43.9	50.0	—	1.4	—	353.8	—	449
Adjusted ET rate (feet per year)	—	—	—	3.05	2.94	—	2.45	—	0.97	—	—
Annual ground-water discharge (acre-feet)	—	—	—	39.3	44.7	—	1.2	—	260.0	—	345

Table 16. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Oasis Valley

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	1.0	5.0	39.7	928.6	647.2	961.8	80.6	808.4	0.0	1.0	3,473
ET rate (feet per year)	8.60	8.50	4.20	3.10	3.09	1.19	2.69	1.92	—	.50	
Annual ET (acre-feet)	8.6	42.5	166.7	2,878.7	1,999.8	1,144.5	216.8	1,552.1	—	.50	8,010
Adjusted ET rate (feet per year)	8.10	8.00	3.71	2.60	2.59	0.69	2.20	1.42	—	0.00	
Annual ground-water discharge (acre-feet)	8.1	40.0	147.3	2,414.4	1,676.2	663.6	177.3	1,147.9	—	.00	6,275

Table 17. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Sarcobatus Flat

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	0.0	0.0	0.0	0.0	0.0	2,603.7	1,446.3	19,368.4	0.0	10,771.7	34,190
ET rate (feet per year)	—	—	—	—	—	.80	2.39	1.02	—	.50	—
Annual ET (acre-feet)	—	—	—	—	—	2,083.0	3,456.7	19,755.8	—	5,385.8	30,681
Adjusted ET rate (feet per year)	—	—	—	—	—	0.30	1.90	0.52	—	0.00	—
Annual ground-water discharge (acre-feet)	—	—	—	—	—	781.1	2,748.0	10,071.6	—	0.00	13,601

Table 18. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Shoshone area

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	0.0	1.0	12.0	62.0	341.7	658.5	152.4	21.9	135.9	0.0	1,385
ET rate (feet per year)	—	8.40	4.30	3.60	2.81	.99	2.39	1.49	1.31	—	—
Annual ET (acre-feet)	—	8.4	51.6	223.2	960.2	651.9	364.2	32.6	178.0	—	2,470
Adjusted ET rate (feet per year)	—	8.11	4.01	3.31	2.53	0.70	2.10	1.21	1.02	—	—
Annual ground-water discharge (acre-feet)	—	8.1	48.1	205.2	864.5	461.0	320.0	26.5	138.6	—	2,072

Table 19. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Stewart Valley

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	0.0	1.0	0.0	21.9	18.0	48.7	86.9	9.0	381.7	2,448.8	3,016
ET rate (feet per year)	—	8.40	—	3.10	3.09	1.22	2.39	1.89	1.11	.49	—
Annual ET (acre-feet)	—	8.4	—	67.9	55.6	59.4	207.7	17.0	423.7	1,199.9	2,040
Adjusted ET rate (feet per year)	—	8.05	—	2.75	2.75	0.87	2.05	1.54	0.76	0.14	—
Annual ground-water discharge (acre-feet)	—	8.0	—	60.2	49.5	42.4	178.1	13.9	290.0	342.8	985

Table 20. Simulated mean annual evapotranspiration (ET) and ground-water discharge from Tecopa/California Valley area

[Data are simulated means of 1,000 realizations. Mean annual precipitation is 0.35 foot; —, no data or not applicable]

	Evapotranspiration unit										Total
	1	2	3	4	5	6	7	8	9	10	
Area (acres)	2.0	11.9	16.0	656.1	379.3	394.5	896.3	45.2	1,105.4	8.0	3,515
ET rate (feet per year)	8.60	8.50	3.90	3.30	2.89	.89	2.39	1.81	1.20	.49	—
Annual ET (acre-feet)	17.2	101.2	62.4	2,165.1	1,096.2	351.1	2,142.2	81.8	1,326.5	3.9	7,348
Adjusted ET rate (feet per year)	8.31	8.21	3.60	3.01	2.60	0.60	2.10	1.52	0.91	0.20	—
Annual ground-water discharge (acre-feet)	16.6	97.7	57.6	1,974.9	986.2	236.7	1,882.2	68.7	1,005.9	1.6	6,328

Table 21. Summary statistics data and simulated means of 1,000 realizations by discharge area for the Monte Carlo analysis used to simulate annual ground-water discharge

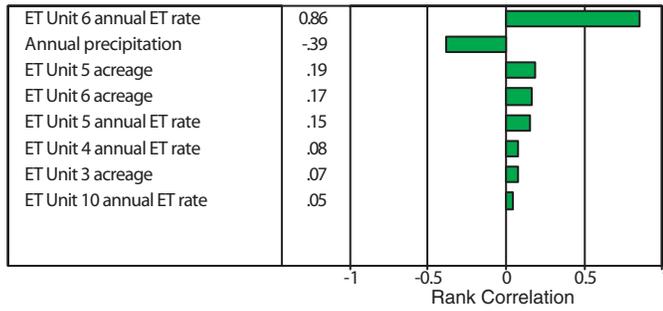
Discharge Area	Simulated annual ground-water discharge (acre-feet)					
	Mean	Median	Range width		Standard deviation	CV ^a
			Minimum	Maximum		
Ash Meadows	17,811	17,761	9,908	28,267	3,091	0.17
Chicago Valley	436	433	40	959	159	.36
Franklin Lake	1,096	1,057	237	2,372	352	.32
Franklin Well area	345	346	70	646	88	.26
Oasis Valley	6,287	6,290	4,169	8,922	773	.12
Sarcobatus Flat	15,005	14,655	2,014	37,791	7,187	.48
Shoshone area	2,075	2,071	1,167	2,639	316	.15
Stewart Valley	1,080	990	357	2,990	457	.42
Tecopa/California Valley area	6,319	6,313	4,448	8,342	673	.11

^a Coefficient of variability (CV) is standard deviation divided by mean.**Table 22.** Differences in the standard deviation of ground-water discharge estimates resulting from correlating similar ET units in Ash Meadows

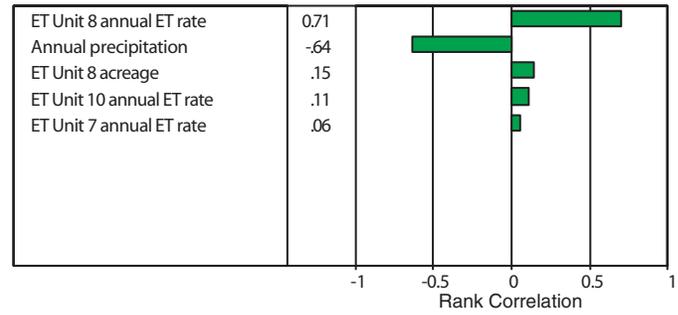
[Abbreviation and symbol: ET, evapotranspiration; —, not applicable or no data]

Unit correlation	Standard deviation	Percent difference
None	3,091	—
ET units 3 and 4	3,139	1.55
ET units 4 and 5	3,095	.13
ET units 3 and 5	3,092	.03
ET units 3, 4, and 5	3,026	-2.10

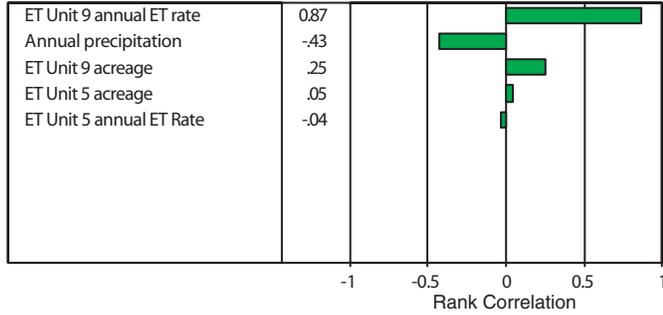
Ash Meadows



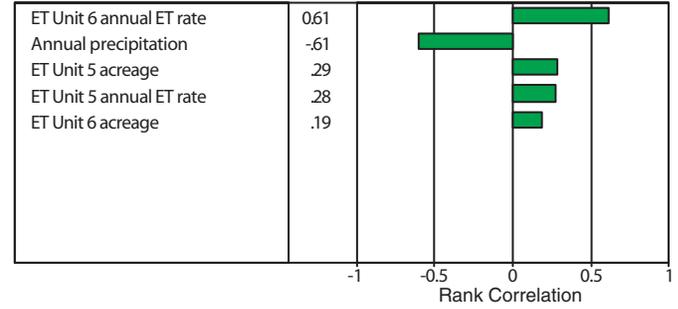
Sarcobatus Flat



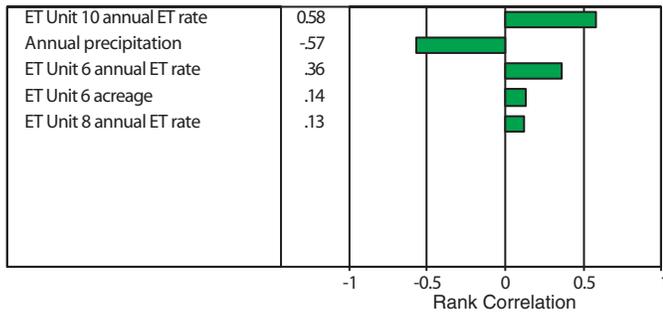
Chicago Valley



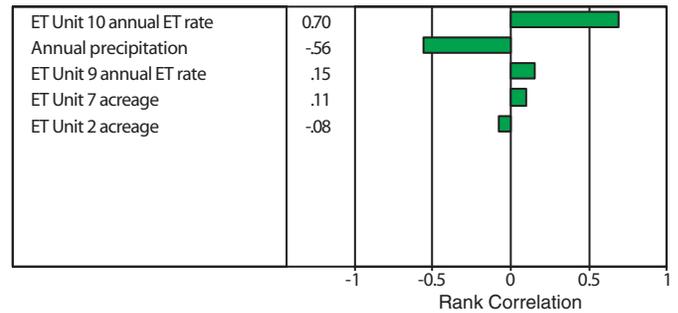
Shoshone area



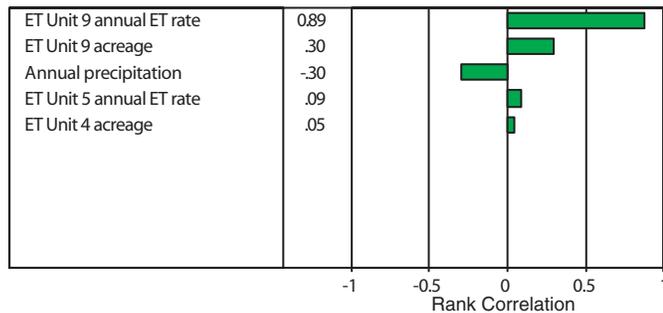
Franklin Lake



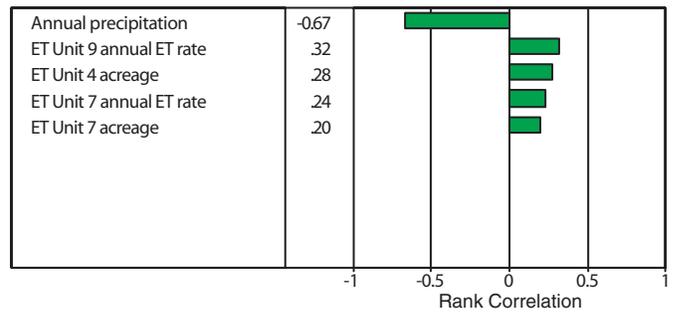
Stewart Valley



Franklin Well area



Tecopa/California Valley area



Oasis Valley

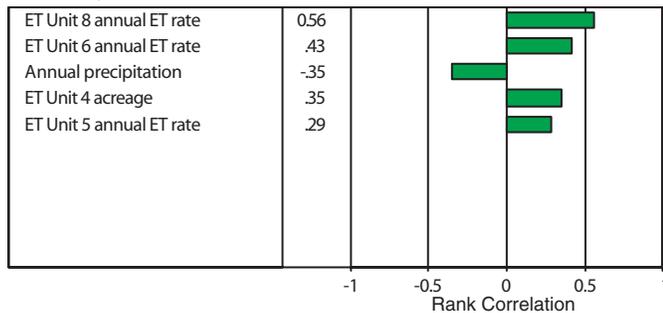


Figure 17. Parameters having the greatest effect on simulated annual ground-water discharge measured by Rank Correlation.

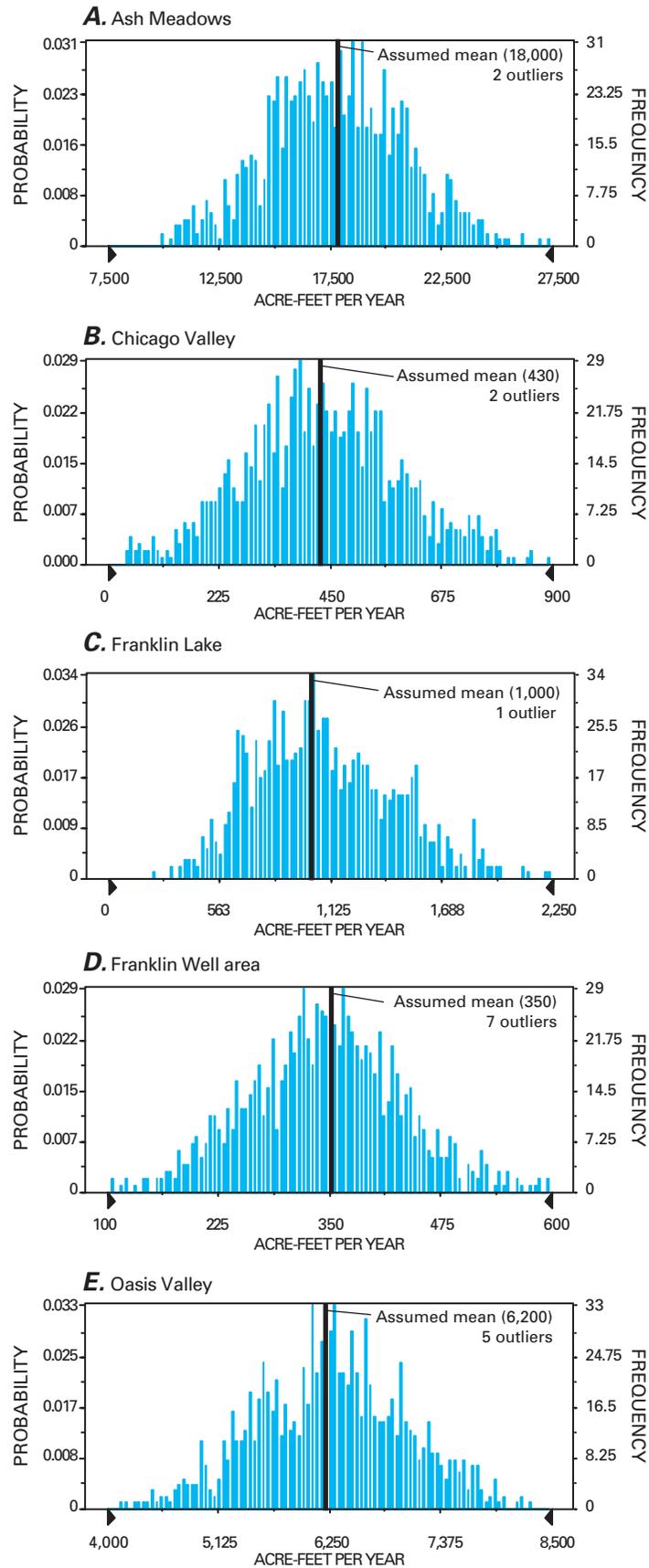


Figure 18. Frequency charts generated from 1,000 realizations of simulated annual ground-water discharge from major discharge areas.

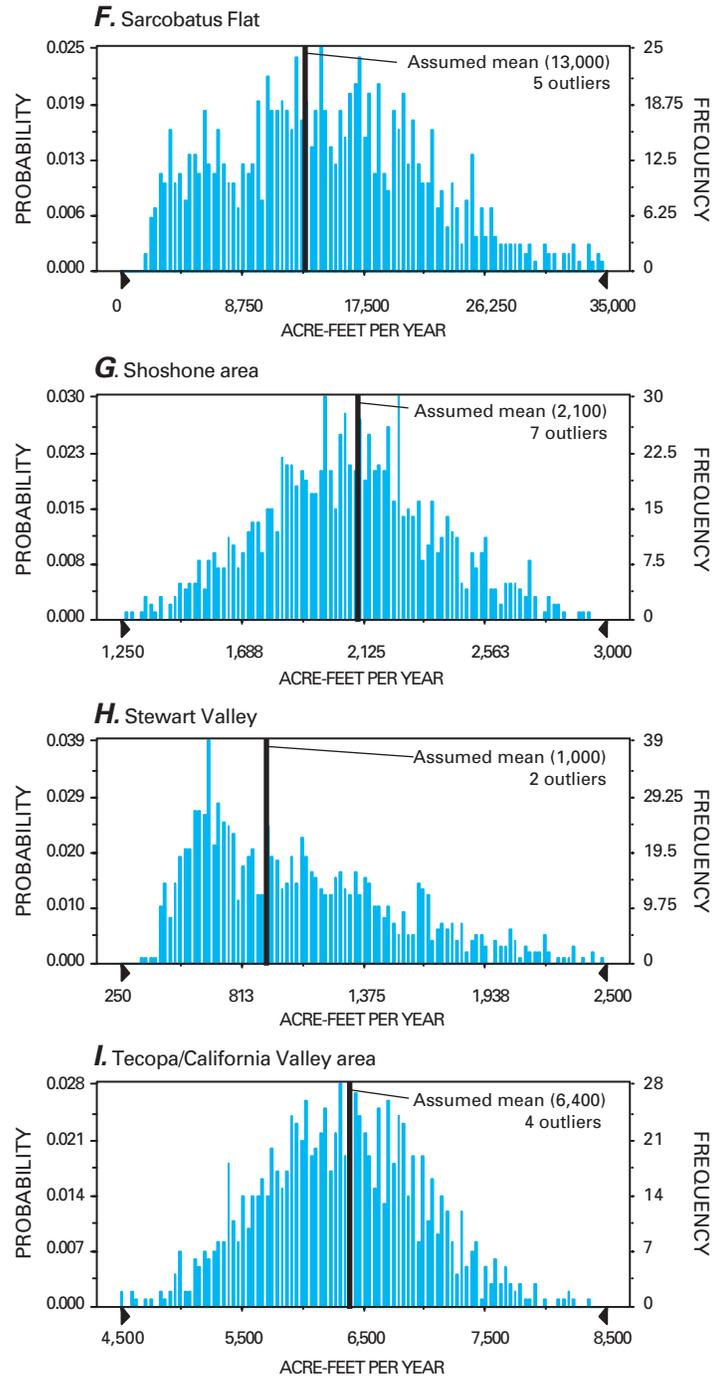


Figure 18. Frequency charts generated from 1,000 realizations of simulated annual ground-water discharge from major discharge areas — Continued.