# Contributions to Stratigraphy 1965

GEOLOGICAL SURVEY BULLETIN 1224

This volume was published as separate chapters A-J

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# UNITED STATES DEPARTMENT OF THE INTERIOR

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# STEWART L. UDALL, Secretary

#### **GEOLOGICAL SURVEY**

William T. Pecora, Director

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# Changes in Stratigraphic Nomenclature by the U.S. Geological Survey 1964

By GEORGE V. COHEE and WALTER S. WEST

CONTRIBUTIONS TO STRATIGRAPHY

GEOLOGICAL SURVEY BULLETIN 1224-A



UNITED STATES GOVERNMENT PRINTING OFFICE, WASHINGTON: 1965

## UNITED STATES DEPARTMENT OF THE INTERIOR

STEWART L. UDALL, Secretary

#### GEOLOGICAL SURVEY

William T. Pecora, Director

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### CONTRIBUTIONS TO STRATIGRAPHY

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE BY THE U.S. GEOLOGICAL SURVEY, 1964

By GEORGE V. COHEE and WALTER S. WEST

#### LISTINGS OF NOMENCLATURAL CHANGES

In the following listings, the changes in stratigraphic nomenclature are grouped together in the categories of (1) new names adopted, (2) previously used names adopted, (3) names revised, (4) changes in age designation, (5) names reinstated, and (6) names abandoned. The stratigraphic names involved in change are listed alphabetically under each category. The age of the unit, the area in which the name is employed, the title of the report, and the publication in which the change is described are given.

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Year of	publication				1964 [1065]	1964		1964	1965	1964	1964		1965	1964	1965
s adopted	Publication (U.S. Geol. Sur- vey except as indicated)	This report, p. A30	This report, p. A50	This report, p. A33	Bull. 1179.	Map MF-282.	This report, p. A25	Map GQ-345	Bull. 1194-M	Am. Assoc. Petroleum Geol- ogists Bull., v. 48, no. 9.	do	This report, p. A23	Prof. Paper 477	Prof. Paper 475-D	Bull. 1194-K
Report in which new name i	Title and <b>author</b> ship	Nomenclature and age of formations in the An- sonia quadrangle, Fairfield and New Haven	Counties, Connecticut, by C. E. Fritts. Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Belted Range Tuff of Nye and Lincoln Coun- ties, Nevada, by K. A. Sargent, D. C. Noble.	and B. B. Ekren. Geology of the Frenchie Creek quadrangle,	no un contact verveus, by L. J. F. Muller. Recommissance bedrock geology of the Wabas- sus Lake quadrangle, Maine, by D. M. Lar-	rabee. Demry Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke Granodiorite Monrif Flarrord anoferation	Geologic map of the Prospect Peak quad- Geologic map of the Prospect Peak quad-	Lower Mesozoic extrusion of the Volumerul Lower Mesozoic extrusion and the Canalo Hills Volumerules by	F. T. Hayes, F. S. Simons, and R. S. Raup, New Permian stratigraphic units in the Glass Mountains, west Taxas, by G. A. Cooper and R. E. Grant	op	Denny Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke	Granodiorich, Mount Harvard quadrangle, Golorado, by Fred Barker and M. R. Brock, Lake Bonneville: Quakernary stratigraphy of esstern Jordan Valley south of Salt Lake	City, Utan, by K. B. Morrison. Thirsty Caryon Tuff of Nye and Esmeralda Counties, Nevada, by D. C. Noble, R. E.	Anderson, E. E. Erren, and J. T. O. Connor. The Foote Creek and Duton Creek Forma- tions, two new formations in the north part of the Laramie basin, Wyoming, by H J.
	Location	Connecticut	Nevada	op	do	Maine	Colorado	California	Arizona.	Texas.	do	Colorado	Utah	Nevada	Wyoming
	Age	Ordovician(?)	Pliocene	Miocene or Pliocene	Mesozoic	Devonian	Precambrian	Early Pleistocene	Triassic and Jurassic	Early Permian (Leonard).	do	Precambrian.	Pleistocene	Pliocene.	Paleocene
	Name	Allingtown Metadiabase	Ammonia Tanks Member (of Tim- ber Mountain Tuff) (of Piapi	Canyon Group). Belted Range Tuff	Big Pole Formation (of Pony Trail	Group). Bottle Lake Quartz Monzonite	Browns Pass Quartz Monzonite	Burney Basalt.	Canelo Hills Volcanics	Cathedral Mountain Formation	Decie Ranch Member (of Skinner	Ranch Formstion). Denny Creek Granodiorite Gneiss.	Draper Formation (of Lake Bonne- ville Group).	Dry Lake Member (of Thirsty Canyon Tuff).	Dutton Creek Formation

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do 	Bull. 1179	Prof. Paper 475-	Man GO-303	This second as	TILS Report, p.	Bull. 1205	Bull. 1181-R.	This report, p.		Prof. Paper 475-	Bull. 1194-0		Prof. Paper 477.	Bull. 1194-N		Map 44-358	This report, p. /	Bull. 1194-B	This report, p. /	mbio montat
Hyden, H. McAndrews, and R. H Tschudy.	Geology of the Frenchie Creek quadrangle,	north-central Nevada, by L. J. P. Muffler. Thirsty Canyon Tuff of Nye and Esmeralda	Counties, Nevada, by D.C. Noble, R. E. Anderson, E. B. Ekren, and J. T. O'Connor. Geology of the Twrone curedranels Kentucky	by E. R. Cressman.	Grante of northeastern Wisconsin, by W Dring	Geology of the Garns Mountain quadrangle, Bonneville, Madison and Teton Counties,	Iteanto, DY M. H. STARZS BUT H. F. ALDGE. Reconnaissance geology of Admiratty Island, Alaska, by E. H. Lathram, J. S. Pomeroy,	H. C. Berg, and R. A. Loney. Denny Creek Granodiorite Gneiss, Browns	Fass Quartz Monzonite, and Kroenke Granodiorite, Mount Harvard quadrangle, Colorada by Freed Barkar and M R Brook	Thirsty Canyon Tuff of Nye and Esmeralda Counties, Neveda, by D. C. Noble, R. E.	Anderson, E. B. Ekren, and J. T. O'Connor. Stratigraphy and chronology of late intergla- cial and early Vashon glacial time in Seattle	H. Washington, by D. R. Mullineaux, H. Waldron and Mever Rubin.	Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake	City, Utah, by R. B. Morrison. Nomenclature and correlation of lithologic subdivisions of the Jefferson and Three	Forks Formations of southern Montana and northern Wyoming, by C. A. Sandberg.	bedrock geologic map of the Big Lake quad- drangle, Maine, by D. M. Larrabee.	Nomenclature and age of formations in the Ansonia quadrangle, Fairfield and New Haven Counties, Connecticut, by C. E.	Fritts. Hovey Group, a redefined stratigraphic name for the Hovey Formation of northeast Maine.	by Louis Pavlides. Marinette Quartz Diorite and Hoskin Lake	Graute of northesserth wiscousin, by W.
đo	Nevada.	do	Kentucky.	Wisconsin		Idaho.	Alaska.	Colorado		Nevada	Washington		Utah	Montana and Wyo- ming.		Multiple.	Connecticut	Maine	Wisconsin	
Late Cretaceous and	Paleocene. Mesozoic	Pliocene	Middle Ordovician	Procembrian		middle Pliocene or younger.	Paleocene through Miocene.	Precambrian		Pliocene	Pleistocene		do	Late Devonian		Levonan	Ordovician(?)	Silurian	Precambrian	
Foote Creek Formation	Frenchie Creek Rhyolite (of Pony	Trail Group). Gold Flat Member (of Thirsty	Canyon Tutt). Grier Limestone Member (of	Lexington Limestone).	AND	Kirkham Hollow Volcanics	Kootznahoo Formation	Kroenke Granodiorite		Labyrinth Canyon Member (of Thirsty Canyon Tuff).	Lawton Clay Member (of Vashon Drift).		Little Cottonwood Formation (of Lake Bonneville Group).	Logan Gulch Member (of the Three Forks Formation).		TOVE HIDE BURNER TRUNCING BURNER	Maltby Lakes Volcanics	Maple Mountain Formation (of Hover Group).	Marinette Quartz Diorite	

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE

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<b>URVEY REPORTS-Continued</b>
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				vey except as indicated)	
Matagamon Sandstone	Early Devonian	Maine	The Matagamon Sandstone: a new Devonian formation in north-central Maine, by D. W.	Bull. 1194-F	1965
Nine Lake Formation (of Hovey Group).	Ordovician or Silurian.	dodo	Rankin. Hovey Group, a redefined stratigraphic name for the Hovey Formation of northeast Maine,	Bull. 1194–B	1964
Jronoque Member (of Derby Hill Schist).	Ordovician(?)	Connecticut.	by Louis Favlides. Nomenclature and age of formations in the Ansonia quadrangle, Fairfield and New Haven Counties, Connecticut, by C. E.	This report, p. A30	•.
bah Canyon Member (of Paint- brush Tuff) (of Plapi Canyon	Miocene(?) and Pliocene.	Nevada	Fritts. Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	This report, p. A49.	
Paintupy.	dododo	do	do	This report, p. A45	
Pocamoonshine Gabbro-Diorite	Devonian	Maine	Bedrock geologic map of the Big Lake quad-	Map GQ-358	1964.
Pony Trail Group	Mesozoic	Nevada	Geology of the Frenchie Creek quadrangle,	Bull. 1179	1964
Poplar Tank Member (of Skinner Ranch Formation).	Early Permian (Leonard).	Texas	north-central Nevada, by L. J. F. Mumer. New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper-	Am. Assoc. Petroleum Geol- ogists Bull., v. 48, no. 9.	[1965] 1964
Pungo River Formation	middle Miocene	North Carolina	and R. E. Grant. The Pungo River Formation, a new name for middle Miocene phosphorites in Beaufort	Southeastern Geology, v. 5, no. 4.	1964
Road Canyon Member (of Word Formation).	Early and Late Per- mian (Guadalupe).	Texas.	County, North Carolina, by J. O. Kimrey. New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper	Am. Assoc. Petroleum Geologists Bull., v. 48,	1964
Salyer Formation	late Miocene	Nevada.	and K. E. Grant. Salyer and Wahmonie Formations of south- eastern Nye County, Nevada, by F. G.	no. 9. This report, p. A40	
Skinner Ranch Formation	Early Permian (Leonard).	Texas	Poole, W. J. Carr, and D. P. Elston. New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper	Am. Assoc. Petroleum Geologists Bull., v. 48,	1964
Sod House Tuff (of Pony Trail	Mesozoic	Nevada	Geology of the Frenchie Creek quadrangle,	Bull. 1179.	1964
South Pass Formation	late Miocene to middle Pliocene.	Wyoming	north-central Nevada, by L. J. F. Mumer. South Pass Formation on the southwest flank of Wind River Mountains, Wyoming, by	This report p. A27	[1965]
			N. M. Denson, H. D. Zeller, and E. V. Stenhens.		

ate Jurassic and Early Cretaceous.	Alaska	Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. O. Porr, and D. A. Tanay,	Bull. 1181-R	1965
larly Permian (Leonard).	Texas	New Permian stratigraphic units in the Glass Mountains, west Texas, by G. A. Cooper	Am. Assoc. Petroleum Geologists Bull., v. 48,	1964
liocene	Nevada	and K. E. Urant. Thirsty Canyon Tuff of Nye and Esmeralda Counties, Nevada, by D. C. Noble, R. E. Anderson, E. B. Ekren, and J. T.	по. <sup>и</sup> о. <sup>у</sup> . Prof. Paper 475-D	1964
	dodo	O'Comnor. Paintbursh Tuff and Timber Mountain Tuff of Nyes County Neveda hy P. P. Orbild	This report, p. A49	
do	do	Thirsty Canyon Tuff of Nye and Esmeralda Counties, Nevada, by D. C. Noble, R. E.	Prof. Paper 475-D	1964
ate Devonian	Montana and Wyoming.	Anderson, B. B. Ekren, and J. T. O'Connor. Nomenclature and correlation of lithologic subdivisions of the Jefferson and Three Forks Formations of southern Montana and	Bull. 1194-N	1965
evonian	Maine	northern Wyoming, by C. A. Sandberg. Bedrock geologic map of the Big Lake quad-	Map GQ-358	1964
ocene	Wyoming.	Terugue, Maure, by D. M. Larrabee. Tertiary geology of the Beaver Divide area, Fremont and Natrona Counties, Wyoming,	Bull. 1164	1964
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leistocene	Wyoming	Foole, W. J. Carr, and D. F. Elston. Three pre-Bull Lake tills in the Wind River Mountains, Wyoming, reinterpreted, by	Prof. Paper 501-D	196 <b>4</b>
iarly and Middle(?) Devonian.	Virginia.	G. M. Richmond. Wildcat Valley Sandstone (Devonian) of wouthwest Virginia, by R. L. Miller, L. D.	Prof. Paper 501-B	1964
arly Pennsylvanian (Morrow).	Arkansas	The Witts, Burds, 20, 20, 20, 20 The Witts, Springs, Formation of Morrow age in the Snowball quadrangle, north-central Arkansas, by E. E. Qlick, S. E. Frezon, and	Bull. 1194-D	1964 [1965]
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				Report in which name	is adopted	
Name	Age	Location	Original authorship	Title and authorship	Publication (U.S. Geol. Survey except as indicated)	Year of publication
Argentine Limestone Member (of Wyandotte Limestone) (of Kansas	Late Pennsylvanian	Nebraska, Iowa, and Kansas.	Newell, 1935	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	Prof. Paper 472	1964 [1965]
City Group) (of Missouri Series). Beirdneau Sandstone Member (of Jefferson Formation).	Late Devonian	Utah	Williams, 1948	R. D. Miller. Geology of the Paradise quad- rangle, Cache County, Utah, by T. B. Mullens and G. A.	Bull. 1181-8	1964
Bisher Limestone	Middle Silurian	Kentucky	Foerste, 1917	Lett. Geology of the Charters quad- rangle, Kentucky, by R. H.	Map GQ-293	1965
Block Limestone Member (of Cherry- vale Formation) (of Kansas City	Late Pennsylvanian	Nebraska, Iowa, and Kansas.	Newell, 1935	Morris. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by	Prof. Paper 472	1964 [1965]
Group) (of Missouri Series). Boyle Limestone	Middle Devonian	Kentucky	Foerste, 1906	R. D. Muler. Geology of the Dunville quad- rangle, Kentucky, by C. H.	Map GQ-367	1965
Canville Limestone Member (of Dennis Limestone) (of Kansas	Late Pennsylvanian	Nebraska, Iowa, and Kansas.	Jewett, 1932	Maxwell. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	Prof. Paper 472	1964 [1965]
City droup) of (Missouri Scries). Chorry Valley Limestone Member (of Marcellus Shale).	Middle Devonian	New York	Clarke, 1903	Dy K. D. MILLET. A new Species of the rugose coral genus Natinkinella from the Middle Devonian of eastern Pennsylvania, by W. A. Oliver,	Jour. Palenotology, v. 38, no. 5.	1964
Crab Orchard Formation	Early and Middle Silurian.	Kentucky	Linney, 1882	Jr. Geology of the Charters quad- rangle, Kentucky, by R. H.	Map GQ-293	1965
Devils Hollow Member (of Cynthi- ana Formation).	Middle Ordovician	do	McFarlan and White, 1948.	Geology of the Tyrone quad- rangle, Kentucky, by E. R.	Map GQ-303	1964
Dockendorff Group	Early Devonian	Maine	Boucot, Field, Fletcher, Forbes, Naylor, and Pavildes,	Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Pavior, and Mencher, R. S. Navior, and	Prof. Paper 501-C	1964
Edmunds Hill Andesite (of Dock-	do	do	1964. Gregory, 1899	A. J. Boucot.	do	1964
endorff Group). Fontana Shale Member (of Cherry- vale Formation) (of Kansas City Group) (of Missouri Serles).	Late Pennsylvanian	Nebraska, Iowa, and Kansas.	Newell, 1935	Geology of the Omaha-Council Bluffs aree, Nebraska-Iowa, by R. D. Miller.	Prof. Paper 472	1964 [1965]

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do	Map I 426	Map GQ-379	Bull. 1179	Prof. Paper 501-C	Bull. 1181-S	Prof. Paper 472	Map GQ-306	Bull. 1181-8	Bull. 1194-J	do	Map I-426	Bull. 1194-E	
dodo	Geologic map of the Lucerne Valley quadrangle, San Bernar- dino, County, California, by	T. W. Diouee, Jr. Geology of the Mount Aire quad- rangle, Utah, by M. D. Critten-	Geology of the Frenchie Creek quadrangle, north-central	Nevada, by L. J. P. Muffler. Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Pavlides, Ely Mercher, R. S. Naylor, and	A. J. Boucot. Geology of the Paradise quad- rangle, Cache County, Utah, by T. E. Mullens and G. A.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by	K. D. Muldr. Geologic map of the Santa Rita quadrangle, New Mexico, by R. M. Hernon, W. R. Jones,	and S. L. Moore. Geology of the Paradise quad- rangle, Cache County, Utah, by T. E. Mullens and G. A.	Lett. Members and Loyd Sandstone Members of the Mancos Shale, Moffat and Rio Blanco Coun- ties, Colorado, by J. R. Dyni	and H. L. Cullins. do-	Geologic map of the Lucerne Valley quadrangle, San Bernardino County, California,	by T. W. Dibbles, Jr. pper Miocene and Pliocene marine stratigraphy in southern Salinas Valley, California, by D. L. Durham and W. O. Addicott.	•
do	Vaughan, 1922	Thomas and Krueger, 1946.	Regnier, 1960	Gregory, 1900	Williams, 1948	Neweil, 1935	Kuellmer and others, 1953.	Holland, 1952	Konishi, 1959	do	Shreve in Rich- mond, 1960.	Reed, 1925	
do	California	Utah	Nevada	Maine.	Utah	Nebraska, Iowa, and Kansas.	New Mexico	Utah	Colorado	do	California	do	
do	Paleozoic	Late Triassic	middle Pliocene to middle Pleistocene.	Early Devonian	Late Devonian	Late Pennsylvanian	Miocene(?)	Early Mississippian	Late Cretaceous	do	Tertiary	early Pliocene	·
Frisbie Limestone Member (of Wyandotte Limestone) (of Kansas	City Group) (of Missouri Series). Furnace Limestone	Gartra Grit Member (of Ankareh Formation).	Hay Ranch Formation	Hedgehog Formation (of Docken- dorf Group).	Hyrum Dolomite Member (of Jeffer- son Formation).	Wyandotte Limestone) (of Kansas	Kneeling Nun Rhyolite Tuff	Leatham Formation	Loyd Sandstone Member (of Man- cos Shale).	Meeker Sandstone Member (of	Mancos suale). Old Woman Sandstone.	Pancho Rico Formation	

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ontinued		Year of publication	1964	1964 [1965] 1964	[1965] 1964 [1965]	1964	1964	1964	1964	1964 [1965]	1964	1964	
IY REPORTS-C	is adopted	Publication (U.S. Geol. Survey except as indicated)	Prof. Paper 501-C	Prof. Paper 472dodo	Map I-412	do	Map GQ-306.	Map I-426	Map I-412	Prof. Paper 472	Map GQ-306	Prof. Paper 501-C	
S. GEOLOGICAL SURVE	Report in which name	Title and authorship	Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Pavlides, Ely Mencher, R. S. Naylor, and	A. J. Boutoot. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Geologic map and sections, Deep Creek area, Stevens and Pend Orelile Counties, Washington,	by R. G. Yates.	Geologic map of the Santa Rita quadrangle, New Mexico, by R. M. Hernon, W. R. Jones,	Gend S. L. Moore. Gends in the Lucerne Valley quadrangle, San Bernar- dino County, California, by T.	Geologic map and sections, Deep Geologic map and sections, Deep Creek area, Stevens and Pend Orelle Counties, Washington,	by K. G. Yates. Geology of the Omaha-Council Buffs area. Nebraska-Iowa, by	Geologic map of the Santa Rita quadrangle, New Mexico, by R. M. Hernon, W. R. Jones,	and s. L. Moore. Outline of stratigraphic and tectonic features of northeast Maine, by Louis Pavides, Bly Memher P. S. Navhr. and	A. J. Boucot.
LL USE IN U.S		Original authorship	Boucot, Field, Fletcher, Forbes, Naylor, and Pavlides,	1994. Newell, 1935	Fyles and Hew- lett, 1959.	Frebold and	Little, 1962. Hernon, Jones, and Moore, 1953.	Vaughan, 1922	Daly, 1912	Jewett, 1932	Kuellmer and others, 1953.	Fletcher, 1960	
FOR OFFICIA		Location	Maine	Nebraska, Iowa, and Kansas. do	Washington	do	New Mexico	California	Washington	Nebraska, Iowa, and Kansas.	New Mexico	Maine	
MES ADOPTED		Age	Late Silurian	Late Pennsylvanian do	Cambrian	Jurassic.	Miocene (?)	Paleozoic.	Tertiary	Late Pennsylvanian	Miocene (?)	Early Devonian	:
PREVIOUSLY USED NAI		Маше	Perham Formation	Juindaro Shale Member (of Wyan- dotte Limestone) (of Kansas City Group) (of Missouri Sates). Juivira Shale Member (of Cherry-	vale Formation) (of Kansas City Group) (of Missouri Series). Reeves Limestone Member (of Mait- len Phyllite).	Rossland Group.	Rubio Peak Formation	saragossa Quartzite	sheppard Granite	stark Shale Member (of Dennis Limestone) (of Kansas City	Group) (of Missouri Series). Sugarlump Tuffs	wanback Formation (of Docken- dorff Group).	

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1964 [1965] 1964	[1965] 196 <del>4</del>	
Prof. Paper 472	Map GQ-335	
Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller. -do	Bedrock geology of the Willi- mantic quadrangle, Connecti- cut, by G. L. Snyder.	
Newell, 1935  Bain, 1898	Gregory, 1906	
Nebraska, Iowa, and Kansas. do	Connecticut	
Late Pennsylvanian	pre-Pennsylvanian	
Wea Shala Member (of Cherryvale Formation) (of Kansas City Group) (of Missouri Series). Westervilla Limestone Member (of _	Cherryvale Formation) (of Kansas City Group) (of Missouri Series). Willimantic Gneiss	

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#### CHANGES IN STRATIGRAPHIC NOMENCLATURE

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CONTRIBUTIONS TO STRATIGRAPHY

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	Year of publication	1965	1965	1965	1965	1965	1965	1962	1964
ge is revised	Publication (U.S. Geol. Survey except as indicated)	Prof. Paper 477	Bull. 1188.	Мар ФQ-381.	Prof. Paper 477	Prof. Paper 525-B	Bull. 1181-R.	Map GQ-362	Bull. 1194–D
Report in which usa	Title and authorship	Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R. B.	Morrison. Geology of northwestern North Park, Colorado, by	W. J. Hall, Jr. Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R. Mudge.	Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R. B. Morrison.	Correlation of Cretaceous and lower Tertiary rocks near Livingston, Montana, by	R. D. KOUEUS. Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. C. Berg, and	K. A. Loney. Geology of the Rome quad- rangle, Kentucky, by M. D. Crittenden, Jr., and R. K.	The Witts Springs Formation of Morrow age in the Browball quadrangle, north-central Arkansas, by E. B. Gibles, S. Fregori, and Mackenzie Gordon, Jr.
	Revision	Formerly Alpine Formation of Lake Bonneville Group in Jordan Valley. Alpine Formation in good usage	elsewhere in Utah. Mowry Shale Member in- cluded in Benton Shale in	In report area. Birdbear made member of Jefferson Formation in report area. Birdbear For- mation in good usage else-	Water Formary Bonneville Forma- tion of Lake Bonneville Group in Jordan Valley. Bonneville Formation in good usage elsewhere in	Utan. Reassignment from member of Colorado Shale to mem- ber of Frontier Formation.	Brothers Volcanics assigned to Stephens Passage Group.	Vienna Limestone made a member of the Buffalo Wallow Formation in	Kentucky. Formerly Cane Hill Member of the Hale Formation in report area. Age changed from Fennsylvanian to Missispipian and Early Pennsylvanian.
	Location	Utah	Colorado	Montana.	Utah	Montana.	Alaska	Kentucky	Arkansas
	Age	Pleistocene.	Early and Late Cre- taceous.	Late Devonian	Pleistocene	Late Cretaceous	Late Jurassic and Early Cretaceous.	Late Mississippian	Mississippian and Early Pennsylvanian.
	Name	Alpine Member (of Little Cotton wood Formation) (of Lake Bonneville Group).	Benton Shale	Birdbear Member (of Jeffer- son Formation).	Bonneville Member (of Little Cottonwood For- mation) (of Lake Bonne- ville Group).	Boulder River Sandstone Member (of Frontier For- mation).	Brothers Volcanics (of Stephens Passage Group).	Buffalo Wallow Formation	Cane Hill Formation

STRATIGRAPHIC NAMES REVISED

CHANGES	IN	STRATIGRAPHIC	NOMENCLATURE

F961	1964	1964	1964	1964	1964		1964 [1965]		1965	1964	1965
Prof. Paper 501-C	Prof. Paper 475-D	do	Am. Jour. Science, v. 262, no. 5.	Map GQ-303	Basic Data Report 15		Prof. Paper 472	This report, p. A30	Water Supply Paper 1750-A.	Bull. 1194-B.	Map GQ-381
Outline of the stratigraphic and tectonic features of northeast Maine, by Louis Pavidies, Ely Mencher, R. S. Naylor and A. J. Pavivor	Surface and subsurface surface and subsurface southeastern Mississippi, by D. H. Parrio	do	Reaction between mafic magma and pelific schist, Cortlandt, New York, by Wood Rather	Geology of the Tyrone quad- rangle, Kentucky, by E. R.	Records of wells and test holes, water analyses, and physical properties of water-bearing materials for the Denver bssin, Colorado, by J. A.	A.J. Boettcher, and T.J. Mainr	Geology of the Omaha-Coun- cil Bluffs area, Nebraska- rowa hy R D Miller	Nomenclature and age of for- mations in the Ansonia quadrangle, Fairfield and New Haven Counties, Con-	necticut, by C. E. Fritts. Ground and surfaces water in the Messbi and Vermilion Iron Range area, north- eastern Minnesota, by R. D. Cotter, H. L. Young, L. R. Petri, and C. H.	Fror. Hovey Group, a redefined stratigraphic name for the Hovey Formation of north- east Maine, by Louis	Favrages Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R. Mudge.
Chapman Sandstone included in Dockendorff Group.	Formerly Clayton Formation, which remains in good usage outside of report area.	Formerly Cook Mountain Formation, which remains in good usage outside of remort area	Formerly Cortlandt Series	Formerly Curdsville Lime- stone of Lexington Group.	Dawson Formation in report area. Dawson Arkose in good usage elsewhere.		Winterset Limestone Member placed in Dennis Limestone.	Oronoque Member made a member of the Derby Hill Schist.	Formerly Duluth Gabbro	Formerly Dunn Brook Mem- ber (Early Silurian age) of Hovey Formation.	Formerly Flood Member of Blackleaf Formation.
Maine	Mississippi	do	New York	Kentucky	Colorado		Nebraska and Iowa.	Connecticut	Minnesota.	Maine	Montana
Early Devonian	Paleocene.	middle Eocene.	unknown	Middle Ordovician	Late Cretaceous and Paleocene.		Late Pennsylvanian	Ordovician(?)	Late Precambrian	Ordovician or Sil- urian.	Early Cretaceous
Chapman Sandstone (of Dockendorff Group).	Clayton Limestone (of Midway Group).	Cook Mountain Limestone (of Claiborne Group).	Cortlandt Complex	Curdsville Limestone Member (of Lexington	Dawson Formation		Dennis Limestone (of Kan- sas City Group) (of Mis-	Derby Hill Schist	Duluth Gabbro Complex	Dunn <b>]</b> Brook Fc <b>rm</b> 2tion (of Hovey Group).	Flood Shale Member (of Blackleaf Formation) (of Colorado Group).

		Year of publication	:	1964	1964	1964	1965	1965	1964	1964	1964
	ge is revised	Publication (U.S. Geol. Survey except as indicated)	This report, p.A36	Am. Assoc. Petro- leum Geologists Bull., v. 48, no. 9.	Bull, 1194–B	Bull. 181-S	Map GQ-381	Prof. Paper 477	Map GQ-303	do	Map I-412
-Continued	Report in which us	Title and authorship	Belted Range Tuff of Nye and Lincoln Counties, Ne- vada, by K. A. Sargent, D. C. Noble, and E. B. Ekren.	New Permian stratigraphic units in the Glass Moun- tains, west Texas, by G. A.	Hovey Group, a redefined stratigraphic name for the Hovey Formation of north- east Maine, by Louis	Favilies. Geology of the Paradise quadrangle, Cache County, Utah, by T. E. Mullens and G. A. Izett.	Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R.	Mudge. Lake Bonneville: Quaternary stratigraphy of eastern Jordan Valley south of Salt Lake City, Utah, by R.	B. Morrison. Geology of the Tyrone quad- rangle, Kentucky, by E.	к. Стезящан. do	Geologic map and sections, Deep Creek area, Stevens and Pend Oreille Counties,
IIC NAMES REVISED-		Revision	Member of Belted Range Tuff except in eastern part of Neveda Test Site where it remains a member of the Indian Trail Formation of Oak Spring Group. Re- unner nart of the original	unit throughout its extent. Formerly Hess Limestone Member of Leonard Formation.	Formerly Hovey Formation of Early Silurian (?) and Early Silurian age.	Jefferson Formation divided into the Hyrum Dolomite Member (at botrom) and the Beirdneau Sandstone Member (at top) in report	area. Formation includes Birdbear Member in report area.	Lake Bonneville Group redefined to include Draper Formation (top) and Little Cottonwood Formation	(bottom). Formerly Lexington Group	Formerly Logana Formation	of Lexington Group. Reeves Limestone Member made a member of Maitlen Phyllite.
<b>TRATIGRAPH</b>		Location	Nevada	Teras	Maine	Utah	Montana	Utah	Kentucky		Washington
02		Age	Miocene or Pliocene.	Early Permian (Leonard).	Ordovician or Silurian.	Late Devonian.	do	Pleistocene	Middle Ordovician	do	Early or Middle Cambrian.
		Name	Grouse Canyon Member (of Belted Range Tuff).	Hess Formation	Hovey Group.	Jefferson Formation	Do	Lake Bonneville Group	Lexington Limestone	Lorana Member (of	Lexington Limestone). Maitlen Phyllite

	1965		1965	1964		1965	1965	1964		1964	1964	1965
	Bull. 11 <del>91-</del> J		Bull. 1188.	Map GQ-282	This report, p. A44	Prof. Paper 477	Prof. Paper 525-B	Map I-417	This report, p. A50	Map GQ-282	do	Bull. 1104-N
Washington, by R. G.	Meaker and Loyd Sandstone Members of the Mancos Shale, Moffat and Rio	bianco Countries, Cotorado, by J. R. Dyni and M. L. Cullins.	Geology of northwestern North Park, Colorado, by	Geology of the Shopville quadrangle, Kentucky, by N. L. Hatch, Jr.	Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P.	Lake Borneville: Quater- nary stratigraphy of eastern Jordan Valley south of Salt Lake Oity, Utah, by R. B.	Correlation of Cretaceous and lower Tertiary rocks mear Livingston, Montana, by A F Poborte	Preliminary geologic map of the Tepee Creek quad- rangle, Montana-Wyoming,	Puy troy wukumu Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Geology of the Shopville quadrangle, Kentucky, by N. L. Hatch, Jr.	dodo	Nomenclature and correlation of ithologic subdyrsions of the Jefferson and Three Forks Formations of south- ern Montana and northern Wyoming, by C. A. Sandberg,
	Meeker Sandstone Member and Loyd Sandstone Mem- ber made members of Man-	COS SUIME IN LEDOIL MEN.	Mowry Shale made a mem- ber of Benton Shale in	Ste. Genevieve and St. Louis Limestones made members of Newman Limestone in	Feport area. Formerly Piapi Canyon Formation of Oak Spring Group. Age was early Dilocons or requires	Formerly Prove Formation of Lake Bonneville Group in Jordan Valley. Provo Formation in good usage	Reassignment - is a member ber of the Kootenai Forma- tion as well as of the Clov-	Quadrant Sandstone used in report area. Quadrant Formation of Quadrant	Formerly Rainier Mesa ber of Plapi Canyon For- mation of Oak Spring Oroup. Age was Pllocene	Made member of Newman Limestone in report area. Ste. Genevieve Limestone	Made member of Newman. Limestone in report area. St. Louis Limestone in coord area.	Formerly Saphington Sand- stone Member of Three Forks Formation.
	Colorado	:	do	Kentucky	Nevada	Utah	Montana	Montana and Wyoming.	Nevada	Kentucky	do	Montana and Wyoming.
	Late Cretaceous		Early Creataceous	Late Mississippian	Miocene(?) and Pliocene.	Pleistocene	Early Cretaceous	Pennsylvanian	Pllocene	Late Mississippian	do	Late Devonian and Barly Mississip- pian.
	Mancos Shale	2-62	či Mowry Shale Member (of Benton Shale).	ce Newman Limestone	Piapi Canyon Group	Provo Member (of Little Cottonwood Formation) (of Lake Bonneville Group).	Pryor Conglomerate Mem- ber (of Kootenai Forma- tion).	Quadrant Sandstone	Rainier Mesa Member (of Timber Mountain Tuff) (of Piapi Canyon Group).	Ste. Genevieve Limestone Member (of Newman Limestone).	St. Louis Limestone Mem- ber (of Newman Limestone).	Supplugton Member (of Three Forks Formation).

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# CHANGES IN STRATIGRAPHIC NOMENCLATURE A13

<b>REVISED</b> -Continued	
NAMES	
STRATIGRAPHIC	

	<b>4</b> 40	Location	Darrieion	Report in which usa	ge is revised	
Age Locat	Locat	fo	Kevision	Title and authorship	Publication (U.S. Geol. Survey except as indicated)	Year publics
Late Jurassic and Early Cretaceous.	Alaska		Seymour Canal Formation assigned to Stephens Pas- sage Group.	Reconnaissance geology of Admiratty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. C. Berg, and	Bull 1181-R.	1961
PlioceneNevada	Nevada.		Formerly Spearhead Rhyo- lite of Pllocene(?) age.	Thirsty Canyon Tuff of Nye and Esmeralda Counties, Nevada, by D. O. Noble, R. E. Anderson, E. B.	Prof. Paper 475-D	1964
Miocene(?) anddo	op		Formerly Stockade Wash Member of Piapi Canyon Formation of Oak Spring Droup. Former age was	EKTEN, and J. J. V. Connor. Paintburah Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	This report, p. A45	
Early Cretaceous Montana	Montana		Formerly Taft Hill Glau- contitic Member of Black- leaf Formation.	Bedrock geology of the Saw- tooth Ridge quadrangle, Montana, by M. R. Mindea	Map GQ-381	1965
middle Eocene Mississippi	Mississippi		Formerly Tallahatta For- mation. Tallahatta For- mation in good usage ex-	Surface and subsurface stratigraphic sequence in southeastern Mississippi,	Prof. Paper 475-D	1964
Late Devonian and Montana a Barly Mississippian. Wyomin	Montana <sup>g</sup> Wyomin	g.	Theory are so the point. Theory are so that formation di- vided into (ascending order): Logan Guich, Tridant, and Sappington Members in report area.	Norver, I	Bull. 11 <del>94</del> –N	1965
Miocene(?) and Nevada	Nevada		Formerly Tiva Canyon Member of Piapi Canyon Formation of Oak Spring Group. Age formerly early	C. A. Sautuber, Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	This report, p. A45	×
dodo	do		Pluce of younger. Formerly Topopah Spring Member of Plapi Canyon Formation of Oak Spring Group. Age was Pliocene or younger.	op	This report, p. A49	

	1965	1965	1965	1964	1964 [1965]	1964	
This report, p. A34	Bull. 1194-0	Map GQ-381	Map GQ-362	Prof. Paper 475-D	Prof. Paper 472	Am. Assoc. Petro- leum Geologists Bull., v. 48, no. 9.	This report, p. A49
Belted Range Tuff of Nye and Lincoln Counties, Nevada, by K. A. Sargent, D.C. Noble, and E. B. Ekren.	Stratigraphy and chronology of tote interguation and early Vashon glacial titme in Seattle area, Washington, by D. R. Mullineaux, H. H. Waldron, and Meyer	Redrocting Bedrock Redogy of the Saw- tooth Ridge quadrangle, Montana, by M. R.	Geology of the Rome quad- rangle, Kentucky, by M. D. Crittenden, Jr., and	Surface and subsurface stratigraphic sequence in southeastern Mississippi,	Geology of the Omaha- Council Bluffs area, Ne- braska-Iowa, by R. D. Miller.	New Permian stratigraphic units in the Glass Moun- tains, west Texas, by G. A. Cooper and R. E.	Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.
Member of Belted Range Tuff except in eastern part of Nevada Test Site where it remains a member of the Indian Trail Formation of	Vashon Drift divided into Lawton Clay Member (at base) and Esperance Sand Member (at top).	Formerly Vaughn Bento- nitic Member of Blackleaf Formation.	Formerly Vienna Limestone	Formerly Winona Sand or Formation and remains as such outside of report area.	Formerly Winterset Lime- stone of Pennsylvanian age.	Road Canyon Member replaces the "First Lime- stone" member of P. B. King (1931).	Formerly Yucca Mountain Member of Plapi Canyon Formation of Oka Spring Group. Age was early Pliocene or younger.
qo	Washington	Montana	Kentucky	Mississippi	Nebraska and Iowa.	Texas	Nevada
Miocene or Pliocene	Pleistocene	Early Cretaceous	Late Mississippian	middle Eocene	Late Pennsylvanian	Early and Late Permian (Guada- lupe).	Miocene (?) and Pliocene.
Tub Spring Member (of Belted Range Tuff).	Vashon Drift	Vaughn Member (of Blackleaf Formation) (of Colorado Group).	Vienna Limestone Member (of Buffalo Wallow Formation).	Winona Marl (of Clai- borne Group).	Winterset Limestone Member (of Dennis Limestone) (of Kansas City Group) (of Mis-	Word Formation	Yucca Mountain Member (of Paintbrush Tuff) (of Plapi Canyon Group).

CHANGES IN STRATIGRAPHIC NOMENCLATURE A15

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	Year of publication	- - -		1965	1965	1965	1965	196 <del>4</del> [1965]	1965	1964	1964 [1965]
ation is changed	Publication (U.S. Geol. Survey except as indicated)	This report, p. A63	This report, p. A32	Bull, 1183.	Map GQ-293	do	Bull, 1183	Prof. Paper 472	Bull. 1183.	Science, v. 145, no. 3633.	Prof. Paper 472
Report in which age design	Title and authorship	Precambrian and Lower Cam- brian formations in the Last Channes Range area, Inyo Country, California, by J. H.	Stewart. Nomenclature and age of forma- tions in the Ansonia quadran- gie, Fairfield and New Haven Counties, Connecticut, by	C. E. Fritts. Provisional geologic map of the crystalline rocks of South Car- olina, by W. C. Overstreet and	Henry Bell, 3d. Geology of the Charters quad- rangle, Kentucky, by R. H. Morris.	do	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet	and Henry Bell, 3d. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet	and Henry Bell, 3d. Potassium-argen and lead-alpha ages of some plutonic rocks, Bokan Mountain area, south-	about Anakak, Dy M. A. Lau- phere, E. M. MacKevett, Jr., and T. W. Stern. Bology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.
	Location	California.	Connecticut	North Carolina and South Carolina.	Kentucky, Ohio Michigan, Pennsyl- vania, and West	virginia.	North Carolina and South Carolina.	Nebraska and Iowa	North Carolina and South Carolina.	Alaska.	Nebraska and Iowa
e	Former	Preeambrian(?) and Early Cambrian,	Ordovician or Devonian.	Precambrian(?) or Paleozoic(?).	Mississippian	do	Precambrian	Pennsylvanian	Cambrian(?)	Cretaceous(?) and Tertiary(?).	Pennsylvanian
Ag	New	Precambrian and Early Cambrian.	Devonian	Ordovician to Mis- sissippian,	Devonian or Mississippian.	do	Ordovician to Mississippian.	Late Pennsylvanian	Ordovícían to Mississippian.	Late Triassic or Early Jurassic.	Late Pennsylvanian
	Name	Andrews Mountain Mem- ber (of Campito For- mation).	Ansonia Greiss	Battleground Schist	Formation.	Berea Sandstone or	r ormation. Bessemer Granite	Bethany Falls Limestone Member (of Swope Limestone) (of Kansas City Group) (of Missouri	Series). Blacksburg Schist	Bokan Mountain Granite	Bonner Springs Shale (of Kansas City Group) (of Missouri Serles).

	1965	1964 [1965]	1964	1964 [1965]	1965	1964 [1965]	1965		1965	1965	1964 [1965]	
This report p. A63	Bull. 1181-R	Prof. Paper 472	Prof. Paper 501-C	Prof. Paper 472	Bull. 1201-A	Prof. Paper 472	Bull. 1183	This report, p. A62	Bull. 1181-R.	Bull. 1183	Prof. Paper 472	This report, p. Aöl
Precambrian and Lower Cam- brian formations in the Last Ohance Range area, Inyo County, California, by J. H.	Stewart. Reconnaissance geology of A dmiraity Island, Alaska, by E. H. Lathram, J. S. Pomeroy,	H. C. Berg, and R. A. Loney. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Chemistry of greenstones of the Catootin Formation in the Blue Ridge of central Virginia, by	J. C. Reed, Jr. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	by A. D. Muler. Petrography of evaporities from the Weilington Formation near Hutchinson, Kansas, by C. L.	Jones. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	Provisional geologic map of the crystalline rocks of South Caro- lina, by W. C. Overstreet and	Heury Bell, 3d. Precembrian and Lower Cam- brian formations in the Last Chance Range area, Inyo County, California, by J. H.	Reconnaissance geology of Ad- mirality Island, Alaska, by E. H. Lathram, J. S. Pomeroy,	H. O. Berg, and H. A. Loury. Provisional geologic map of the crystalline rooks of South Caro- lina, by W. C. Overstreet and	Henry Bell, 3d. Geology of the Omaha-Council Bluffs area, by R. D. Miller.	Age of the Eleana Formation (Devonian and Mississipian) in the Nevada Test Site, by F. G. Poole, P. P. Orkild, Mackenzie Gordon, Jr. and Helen Duncen.
California	Alaska	Nebraska and Iowa	Virginia	Nebraska and Iowa	Kansas	Nebraska and Iowa	North Carolina and South Carolina.	California	Alaska	North Carolina and South Carolina.	Nebraska and Iowa	Nevada
Precambrian(?) and Early Cambrian.	Permian.	Pennsylvanian	late Precambrian	Pennsylvanian	Permian	Pennsylvanian	Devonian(?)	Precambrian(?)	Jurassic(?) to Early Cretaceous (?).	Cambrian	Pennsylvanian	Mississippian to Early Pennsyl- vanian.
Precambrian and Early Cambrian.	Early Permian	Late Pennsylvanian	late Precambrian(?)	Late Pennsylvanian	Early Permian	Late Pennsylvanian	Mississippian(?) to Permian(?).	Precambrian	Late Jurassic and Early Cretaceous.	Ordovician to Missis- sippian.	Late Pennsylvanian	Devonian and Miss- issippian.
Campito Formation	Cannery Formation	Captain Creek Limestone Member (of Stanton Limestone) (of Lansing Group) (of Missouri	Catoctin Formation	Chanute Shale (of Kansas City Group) (of Missouri	Chase Group	Cherryvale Formation (of Kansas City Group) (of	Cherryville Quartz Mon- zonite.	Deep Spring Formation	Douglas Island Volcanics (of Stephens Passage Group).	Draytonville Conglomer- ate Member (of Kings Mountain Quartzite).	Drum Limestone (of Kansas City Group) (of Missouri	Eleans Formation

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# CHANGES IN STRATIGRAPHIC NOMENCLATURE A17

CHANGES IN AGE DESIGNATION-Continued

	Year of publication	1964	1964	1964 [1965]	1964 [1965]	1965	1964 [1965]	1965	1965	1964 [1965]	
tion is changed	Publication (U.S. Geol. Survey except as indicated)	Bull. 1180-C	Prof. Paper 483-B	Prof. Paper 472	do	Bull. 1183	Prof. Paper 472.	Bull. 1183	Prof. Paper 474-B	Prof. Paper 472.	This report, p. A62
Report in which age designs	Title and authorship	Middle and lower Ordovician for- mations in southernmost Ne- yada and adjacent California,	by R. J. Ross. The Devonian colonial coral genus. <i>Billingsustraca</i> and its earliest known species, by W. A. Oliver,	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	do	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet	and Henry Bell, 3d. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Over-	street and Henry Bell, 3d. Geology and uranium deposits of Filk Ridge and vicinity, San Juan County, Utah, by R. Q. Lewis, Sr. and R. H.	Campbell. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Precambrian and Lower Cam- brian formations in the Last Chance Range area, Inyo Country, California, by J. H. Stewart
	Location	Nevada and Califor- nia.	New York, Pennsyl- vania and New Jer- sey.	Nebraska and Iowa	do	North Carolina and South Carolina.	Nébraska and Iowa	North Carolina and South Carolina.	Utah	Nebraska and Iowa	California
65	Former	Late Ordovician	Early or Middle De- vonian.	Pennsylvanian	do	Cambrian(?)	Pennsylvanian	Precambrian or early Paleozoic.	Middle Penn- sylvanian.	Pennsylvanian	Precambrian (?)
ĀĘ	New	Middle and Late Ordovician.	Early Dévonian	Late Pennsylvanian	op	Mississippian	Late Pennsylvanian	Ordovician to Devonian.	Middle and Late Pennsylvanian.	Late Pennsyl- vanian.	Precambrian
	Name	3ly Springs Dolomite	Ssopus Shale, Grit, Silt- stone.	Eudora Shale Member (of Stanton Limestone) (of Lansing Group) (of Mis-	souri Series). Parley Limestone Member (of Wyandotte Lime- stone) (of Kansas City Group) (of Missouri Se-	ries). Aaffney Marble	Jalesburg Shale (of Kansas City Group) (of Mis-	sourt Series). Henderson Gneiss	Hermosa Formation	Hickory Creek Shale Member (of Plattsburg Limestone) (of Lansing Grount) (of Missonri	Series). Hines Tongue (of Reed Dolomite).

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1965	1965		1964 [1965]	1964 [1965]	1965	1964 [1965]	1964	[1965] 1964		1964	1964 [1965]	1965	1965		
Bull. 1201-A	Prof. Paper 483-G	This report, p. A45	Prof. Paper 472	do	Bull. 1183	Prof. Paper 472	do	Prof. Paper 458-B		Prof. Paper 501-C	Prof. Paper 472	Prof. Paper 483-G	Bull. 1201–A	This report, p. A29	-
Petrography of evaporites from the Weilington Formation near Hutchinson, Kansas, by C. L.	Jones. Evolution and distribution of the genus Mys and Tertiary migrations of Mollusca, by F.	Paintbrush Tuff and Timber Mountain Tuff of Nye County,	Nevada, by F. F. UKUM. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	do	Provisional geologic map of the crystalline rocks of South Caro- lina, by W. C. Overstreet and	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by	R. D. Miller. do	Rocks, structure, and geologic history of Steamboat Springs thermal area, Washoe County,	G. A. Thompson, and C. H.	Ouline of the stratigraphic and tectonic features of northeast Maine, by Louis Pavlides, Ely Mencher, R. S. Naylor, and A.	J. Bouccor. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Evolution and distribution of the genus Mya and Tertiary migra- tions of Mollusca, by F. S.	Machen. Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C.L.	Jones. South Pass Formation on the southwest flank of Wind River Mountains, Wyoming, by N. M. Denson, H. D. Zeller, and	L. V. ORPHIELS.
Kansas	California	Nevada	Nebraska and Iowa	op	North Carolina and South Carolina.	Nebraska and Iowa	do	Nevada		Maine	Nebraska and Iowa	Alaska	Kansas	Eastern Wyoming and Nebraska	•
Permian	early Miocene	late Miocene or early Pliocene.	Pennsylvanian	do	Cambrian	Pennsylvanian	do	late Pliocene or early Pleistocene.		Late Devonian	Pennsylvanian	Miocene	Permian	Pliocene	
Early Permian (Leonard).	late Miocene or early Pliocene.	Miocene and Plio- cene (?).	Late Pennsylvanian.	do	Oruovician to Missis- sippian.	Late Pennsylvanian.	do	Pliocene (?) and Pleistocene.		early Middle Devo- nian.	Late Pennsylvanian	Oligocene or Miocene.	Early Permian (Leonard).	late Miocene and Pliocene.	
Hutchinson Salt Member (of Wellington For- mation) (of Sumner	Group). Imperial Formation	Indian Trail Formation.	Iola Limestone (of Kansas City Group) (of Missouri Series).	Kansas City Group (of Missouri Series).	Kings Mountain Quartzite.	Lane Shale (of Kansas City Group) (of Missouri	Series). Lansing Group (of Missouri	Series). Lousefown Formation		Mapleton Sandstone	Merriam Limestone Mem- ber (of Plattsburg Lime- stone) (of Lansing Group)	Meshik Formation	Ninnescah Shale (of Sunner Group).	Ogallala Formation	

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE

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	Υ.			Report in which age designa	ition is changed	
Name	New	Former	Location	Title and authorship	Publication (U.S. Geol. Survey except as indicated)	Year of publication
Prospect Gneiss	Devonian	Ordovician or Devonian.	Connecticut	Nomenclature and age of formations in the Ansonia guadrangle, Fairfield and New	This report, p. A32	Ţ
Redwall Limestone	Early and Late Mississippian.	Mississippian	Arizona	Haves Counties, Connecticut, by C. E. Fritts. Stratgraphic importance of corals in the Redwall Linestone, northern Arizona, by W. J.	Prof. Paper 501-C	1964
Reed Dolomite	Precambrian	Precambrian (?)	California	Sando. Precambrian and Lower Cam- brian formations in the Last Chance Range area, Inyo County, California, by J. H.	This report, p. A62	
Retreat Group.	Middle(?) Devonlan	Silurian and Devon- ian.	Alaska	Stewart. Reconnaissance geology of Ad- miralty Island, Alaska, by E. H. Lathram, J. S. Pomeror	Bull. 1181-R.	1965
Rico Formation	Permian. Permian.	Pennsylvanian and Permian(?).	Utah, Arizona, and New Mexico.	H. C. Berg, and R. A. LOUP. Geology and uranium deposits of Blk Ridge and vicinity, San Juan County, Utah, by K. Qamp- Lewis, Sr. and R. H. Camp-	Prof. Paper 474-B	1965
Rock Lake Shale Member (of Stanton Limestone) (of Lansing Group) (of	Late Pennsylvanian	Pennsylvanian	Nebraska and Iowa	Dell. Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Prof. Paper 472	1964 [1965]
Missouri Series). Schoharie Grit	Early Devonian	Early or Middle Devonian.	New York	The Devonian colonial coral genus <i>Billingaastraca</i> and its earliest known species, by	Prof. Paper 483-B	1964
South Bend Limestone Mem- ber (of Stanton Limestone) of Lansing Group) (of	Late Pennsylvanian	Pennsylvanian	Nebraska and Iowa	W. A. Oliver, Jr. Geology of the Ornaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Prof. Paper 472	1964 [1965]
Missouri Series). Spring Hill Limestone Mem- ber (of Plattsburg Lime- stone) (of Lansing Groun)	do	do	do	dodo	do	1964 [1965]
(of Missouri Series). Stanton Limestone (of Lansing Group) (of Mis- souri Series).	do	do	do	op	do	196 <del>4</del> [1965]

1964	1964 [1965]	1965	1964 [1965]	1965	1964	1964 [1965]	1965	1964	1965	1964 [1965]	1965
Prof. Paper 458-B	Prof. Paper 472.	Bull. 1201-A	Prof. Paper 472	Map GQ-379	Prof. Paper 437-A	Prof. Paper 472	Bull. 1201-A.	Bull. 1144-F	Bull. 1183	Prof. Paper 472	Bull. 1183
Rocks, structure, and geologic history of Steamboat Springs thermal area, Washoe County, Nevada, by D. E. White, G. A. Phompson, and C. H. Sand-	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa, by R. D. Miller.	Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by C. L.	Geology of the Omaha-Council Bluffs area, Nebraska-Iowa,	Geology of the Mount Aire quad- rangle, Utah, by M. D. Crit-	Alluvial fans and near-surface subsidence in western Fresno County, California, by W. B.	Geology of the Omaha-Council Beology of the Omaha-Council Buffs area, Nebraska-Iowa, by D. D. Millar	Petrography of evaporites from the Wellington Formation near Hutchinson, Kansas, by O. L.	Distribution of thorium and uranium in three early Paleo- zoic plutonic series of New	Hampshire, by J. B. Lyons. Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet	Build nearly Deal, ou. Geology of the Omaha-Council Fluffs area, Nebraska-Iowa, by D. D. Millor	Provisional geologic map of the crystalline rocks of South Carolina, by W. C. Overstreet and Henry Bell, 3d.
Nevada	Nebraska and Iowa	Kansas	Nebraska and Iowa	Utah	California	Nebraska and Iowa	Kansas and Oklahoma.	New Hampshire	North Carolina and South Carolina.	Nebraska and Iowa	North Carolina and South Carolina.
Pliocene or Pleis- tocene.	Pennsylvanian	Permian.	Pennsylvanian	Early Cambrian	Pliocene and Pleis- tocene(?).	Pennsylvanian	Permian.	do	late Carboniferous(?)	Pennsylvanian	Early Mississippian (?).
Pliocene(?) and Pleis- tocene.	Late Pennsylvanian.	Early Permian (Leonard).	Late Pennsylvanian	Early and Middle Cambrian.	Pliocene and Pleis- tocene.	Late Pennsyl- vanian.	Early Permian (Leonard).	Early Jurassic or Late Triassic.	Ordovician to Devonian.	Late Pennsyl- vanian.	Permian
Steamboat Hills Rhyolite 22	V Stoner Limestone Member (of Stanton Limestone) ev (of Lansing Group) (of Missouri Series).	Summer Group	A Swope Limestone (of Kan- sas City Group) (of Missouri Saries).	Tintic Quartzite	Tulare Formation	Vilas Shale (of Lansing Group) (of Missouri Series)	Weilington Formation (of Sumner Group).	White Mountain Plutonic Volcanic Series.	Whiteside Granite	Wyandotte Limestone (of Kansas City Group) (of Miscouri Series)	Yorkville Quartz Monzonite.

# CHANGES IN STRATIGRAPHIC NOMENCLATURE A21

	s reinstated	Publication (U.S. Geol. Sur- vev except as indicated)
VAMES REINSTATED	Report in which name i	Title and authorship
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	1965		Year of	publication	1964	1964	1964		1965	1964			1965
Publication (U.S. Geol. Sur- vey except as indicated)	Prof. Paper 525-B		oandoned	Publication (U.S. Geol. Sur- vey except as indicated)	Map GQ-303	Prof. Paper 501-D	Map GQ-303	This report, p. A44	Bull. 1194–N	Bull. 1194–B	This report, p. A70	This report, p. A49	Bull. 1181-R
Title and authorship	Suggestions for prospecting for evaporite de- posits in southwest Virginia, by C. F. With- ington.	NAMES ABANDONED	Report in which name is at	Title and authorship	Geology of the Tyrone quadrangle, Kentucky,	Three pre-Bull Lake tills in the Wind River Mountains, Wyoming, reinterpreted, by G.	deology of the Tyrone quadrangle, Kentucky,	Paintbrush Tuff and Timber Mountain Tuff	OLYS COULTY, Neverals, DY F. F. UKIUL. Normenclature and correlation of lithologics ub- divisions of the Jefferson and Three Forks Formations of southern Montana and north-	Wyoming, by C. A. Saudberg. Hovey Group, a redefined stratigraphic name for the Hovey Formation of northeast	Abule, by Louis Faviues. Application and use of the name Arikaree For- mation for lower and middle Miocene rocks in Granite Mountains area of central Wyo-	Paintbrush Tuff and Timber Mountain Tuff of Nye County, Nevada, by P. P. Orkild.	Reconnaissance geology of Admiralty Island, Alaska, by E. H. Lathram, J. S. Pomeroy, H. C. Berg, and R. A. Loney.
	Virginia	<b>FRATIGRAPHIC</b>		Location	Kentucky	Wyoming	Kentucky	Nevada	Montana, Wyoming, North Dakota, and South Dakota.	Maine	Wyoming	Nevada	Alaska
20	Mississippian	18		Age	Middle Ordovician	Pleistocene	Middle Ordovician	Eocene to Pliocene or	younger. Late Devonian	Early Silurian (?)	early and middle. Miocene.	Pliocene or younger	Early Cretaceous (?)
	Little Valley Limestone.			Nane	Benson Limestone (of Lexington	Dinwoody Lake Till and Glaciation.	Jessamine Limestone (of	Oak Spring Group	Potlatch Member (of Three Forks Formation).	Saddleback Mountain Mem- ber (of Hovey Formation).	Split Rock Formation	Survey Butte Member (of Piapi Canyon Formation)	Symonds Formation

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Year of publication

#### DENNY CREEK GRANODIORITE GNEISS, BROWNS PASS QUARTZ MONZONITE, AND KROENKE GRANODIORITE, MOUNT HARVARD QUADRANGLE, COLORADO

#### By FRED BARKER and M. R. BROCK

Three plutonic rocks of Precambrian age in the Mount Harvard quadrangle, Colorado, are here named Denny Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke Granodiorite. Part of the rock assigned to the Denny Creek was previously called Pikes Peak Granite by Stark (1935) and Stark and Barnes (1935), a designation now known to be erroneous; another part and the other two intrusive rocks and all the enclosing schists and gneisses, were assigned by them to the "Sawatch schist and migmatite". The Denny Creek Granodiorite Gneiss, the oldest of the three plutonic rocks, and the Kroenke Granodiorite, the youngest, form batholiths that extend beyond the boundaries of the Mount Harvard quadrangle (fig. 1). The Browns Pass Quartz Monzonite forms two stocks in the central part of the quadrangle but thus far has not been recognized outside the quadrangle.

The Browns Pass Quartz Monzonite discordantly intrudes the Denny Creek Granodiorite Gneiss both in the Denny Creek drainage basin and at contacts of the small stock lying about 2 miles north of Cottonwood Pass (fig. 1). Xenoliths of the granodiorite gneiss lie in the quartz monzonite along parts of the contacts of these two intrusives. The Kroenke Granodiorite intrudes the Denny Creek Granodiorite Gneiss along many miles of contact (fig. 1) and sharply transects the foliation of the granodiorite gneiss. Although the Browns Pass Quartz Monzonite and the Kroenke Granodiorite are nowhere in contact, the Browns Pass is partly syntectonic whereas the Kroenke is posttectonic; moreover, on the south side of Texas Creek, the Kroenke batholith crosscuts and brecciates gneisses and schists that 1 mile to the south contain syntectonic well-foliated sills and dikes of the Browns Pass unit. The Browns Pass Quartz Monzonite evidently is the older of the two.

#### DENNY CREEK GRANODIORITE GNEISS

The Denny Creek Granodiorite Gneiss is named for the exposures along Denny Creek at altitudes of 10,800–11,000 feet. This area is here designated as the type area. This formation consists of gneissic biotite granodiorite and biotite-quartz diorite and underlies much of the west-central and north-central parts of the quadrangle (fig. 1). It is medium to dark gray, is well foliated, and consists of lenticular single grains and aggregates of medium- to coarse-grained plagioclase and quartz and, also, augen of pale-pink microcline perthite; the grains, aggregates, and augen are set in a continuous schistose matrix CONTRIBUTIONS TO STRATIGRAPHY



FIGURE 1.—Generalized distribution of major Precambrian intrusive rocks, Mount Harvard quadrangle, Colorado. D, Denny Creek Granodiorite Gneiss; B, Browns Pass Quartz Monzonite; K, Kroenke Granodiorite.

of fine-grained biotite that, in addition, commonly contains either hornblende or muscovite. Proportions of minerals, especially microcline, differ from one hand specimen to the next in most parts of the batholith so that quartz diorite and granodiorite are closely intermixed in many outcrops. The plagioclase commonly is andesine, rimmed with oligoclase, and is slightly to markedly sericitized. The biotite is pleochroic, X being straw yellow and Y and Z dark brown to dark green. The hornblende is pleochroic, X being golden yellow brown, Y green, and Z bluish green. Chemical and modal analyses

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE

TABLE 1.—Chemical and modal analyses of Denny Creek Granodiorite Gneiss, Browns Pass Quartz Monzonite, and Kroenke Granodiorite, Mount Harvard quadrangle, Colorado

	Denny ri	Creek Gr te Gneis	anodio- ss	Brown M	s Pass Aonzonit	Quartz œ	Kroenke Granodiorite			
	Bf-1	Bf-17	Bf-18	Bf-20	Bf-76	Bf-141	Bf-111	Bf-176	Bf-196	
		Cl [Analyse	nemical A	Analyses S. Dan	iels]					
$\begin{array}{c} {\rm SiO}_{3} \\ {\rm Al}_{3}{\rm O}_{3} \\ {\rm Feo}_{0} \\ {\rm Feo}_{0} \\ {\rm CaO}_{-} \\ {\rm CaO}_{-} \\ {\rm NajO}_{-} \\ {\rm K}_{2}{\rm O}_{-} \\ {\rm H}_{3} \\ {\rm +}_{-} \\ {\rm H}_{3}{\rm O}_{-} \\ {\rm TiO}_{3} \\ {\rm F}_{2}{\rm O}_{3} \\ {\rm MnO}_{-} \\ {\rm CO}_{3} \\ {\rm Cl}_{-} \\ {\rm F}_{-} \\ \\ {\rm Subtotal}_{-} \end{array}$	58. 11 16. 29 3. 45 5. 85 2. 14 5. 17 2. 97 2. 97 1. 29 1. 29 1. 29 1. 29 . 19 . 19 . 02 . 02 . 02 . 02 . 02 . 02 . 02 . 99. 95	59.14 14.65 3.06 7.13 2.54 4.23 2.58 2.57 1.46 .08 1.67 .45 .17 .02 .03 .10 99.88	61. 73 15. 66 2. 88 4. 12 1. 48 2. 91 3. 00 4. 38 1. 83 1. 83 . 10 1. 05 . 35 . 11 . 01 . 03 . 09 99, 73	$\begin{array}{c} 69.\ 71\\ 14.\ 02\\ 1.\ 78\\ 2.\ 36\\ 80\\ 1.\ 91\\ 2.\ 82\\ 4.\ 32\\ 97\\ .14\\ .66\\ 6.\ 25\\ .08\\ .06\\ .02\\ .10\\ \hline \end{array}$	65. 04 15. 14 1. 96 3. 60 1. 65 3. 34 3. 07 3. 40 - 93 . 21 . 95 . 35 . 03 . 02 . 02 . 24 100. 00	74.17 14.12 .41 .63 .96 3.76 5.01 .44 .18 .11 .04 .05 .02 .02 .01 100.11	70. 80 15. 71 . 70 1. 28 . 62 2. 40 5. 05 2. 48 . 50 . 26 . 03 . 01 . 01 . 06 100. 07	68. 32 16. 60 .85 1.98 .97 2.74 5.67 1.51 .66 .05 .38 .11 .04 .03 .02 .06 .99.99	$\begin{array}{c} 66.\ 88\\ 16.\ 88\\ 1.\ 07\\ 2.\ 20\\ 1.\ 20\\ 3.\ 01\\ 5.\ 45\\ 1.\ 98\\ 50\\ 0.\ 52\\ 1.\ 52\\ 0.\ 52\\ 0.\ 52\\ 0.\ 03\\ 0.\ 03\\ 0.\ 08\\ \hline 100.\ 08\\ \end{array}$	
Less O	99.90	99.83	99.68	. 05	. 10	100.11	. 03 100. 04	. 03 99. 96	. 04	

Mode

[Analyses by R. C. Bucknam]

Plagioclase Potassic feldspar	50	49	47 15	31 29, 5	32.5 17.5	36. 5 28	53 10	63. 5 1	63. 5 1. 5
Quartz	25.5	22	21	29.5	29.5	32	30	25	22
Biotite	21	18	14	5.5	16.5	3.5	4.5	9	12
Hornblende		8	. 5						
Muscovite				1.5		Trace	1		
Sphene	1.5	Trace	Trace	Trace	1.5		.5	Trace	Trace
Magnetite and ilmenite	1.5	2.5	1	1.5	1		. 5	Trace	.5
Epidote				.5	Trace	Trace	Trace	1	
Allanite			Trace		Trace		Trace	Trace	
Apatite	Trace	Trace	Trace	.5	.5	Trace	Trace	Trace	. 5
Zircon	Trace	Trace	Trace	Trace	Trace		Trace	Trace	Trace
					•	1	· · · · · · · · · · · · · · · · · · ·		

Bf- 1. Gneissic biotite-quartz diorite, 0.1 mile north of Pass Creek, 11,040 ft altitude.
Bf- 17. Gneissic biotite-hornblende-quartz diorite, 100 ft east of north fork of Denny Creek, 10,970 ft altitude.
Bf- 18. Gneissic biotite-hornblende granodiorite, 150 ft east of north fork of Denny Creek, 10,970 fet altitude.
Bf- 20. Leucoquartz monzonite, 0.25 mile southeast of Browns Pass, 11,960 ft altitude.
Bf- 76. Biotite-quartz monzonite, 0.18 mile,north of Peak 12,955; 12,760 ft altitude.
Bf- 141. Alaskite, ridge between north fork of Denny Creek and Delaney Gulch, 0.1 mile north of 12,320,ft closed contour, 12,300 ft altitude.
Bf-111. Leucogranodiorite, northwest side of Peak 12,778; 12,220 ft altitude; 0.4 mile west-northwest of Kroenke Lake.
Bf-1616. Leucograne diorite, 0.19 mile southeast of Jako Babasea, 10,070 ft altitude;

Bf-176. Leucoquartz diorite, 0.12 mile southeast of Lake Rebecca, 12,270 ft altitude. Bf-196. Blottle-quartz monzonite, north side of canyon of Pine Creek, 0.11 mile northeast of Peak 11,391; 11,150 ft altitude.

of three typical varieties of the granodiorite gneiss are presented in table 1.

#### BROWNS PASS QUARTZ MONZONITE

The Browns Pass Quartz Monzonite, named for exposures west and southeast of Browns Pass (fig. 1), here designated the type area, consists of three rock types that, from oldest to youngest, are (1)

buff medium- to coarse-grained foliated homogeneous biotite-quartz monzonite (sample Bf-76, table 1), found northeast of Browns Pass; (2) pink to buff coarse-grained massive to foliated quartz monzonite and granite (sample Bf-20, table 1), found southeast to west of Browns Pass and also in the small stock north of Cottonwood Pass; and (3) buff fine- to medium-grained massive alaskite (sample Bf-141, table 1), found as dikes southeast of Browns Pass. The plagioclase of the different varieties ranges from median andesine to calcic oligoclase, and much of it has been altered to sericite. The potassic feldspar is well-twinned slightly perthitic microcline. The biotite is typically pleochroic, X being straw yellow and Y and Z dark olive brown. Chemical and modal analyses of the Browns Pass Quartz Monzonite are shown in table 1.

#### **KROENKE GRANODIORITE**

The Kroenke Granodiorite, named for excellent exposures west of Kroenke Lake, the type area, consists of quartz diorite, granodiorite, and quartz monzonite in a sharply discordant body that underlies much of the drainage areas of Pine Creek, North Cottonwood Creek, and the uppermost reaches of Texas Creek, and a small pluton in the Ptarmigan Creek drainage area in the southern part of the quadrangle. It ranges in composition from quartz monzonite to quartz diorite, but these types are so similar in outcrop and hand specimen that one cannot be distinguished from the other in the field. The rocks are mostly quartz diorite in the Pine Creek drainage area and granodiorite and quartz monzonite from Horn Fork Creek to Texas Creek. Thev are white, light gray, or buff and fine to medium grained; the rocks are commonly foliated and contain single grains and wisplike clusters of biotite but in many areas are massive. They range from homogeneous forms to banded forms consisting of alternating biotite-rich and quartz feldspar-rich laminae. The granodiorite contains swarms of amphibolite inclusions in some places.

Plagioclase in the Kroenke Granodiorite is calcic oligoclase or sodic andesine. Microcline is well twinned and slightly perthitic to nonperthitic. The biotite is pleochroic, X being straw yellow or pale brownish green and Y and Z dark olive green or greenish brown. The biotite commonly is partly altered to chlorite and typically is intergrown with apatite, epidote, sphene, magnetite, and allanite. In the vicinity of amphibolite inclusions the rocks contain hornblende. Three chemical and modal analyses are shown in table 1.

#### SOUTH PASS FORMATION ON THE SOUTHWEST FLANK OF WIND RIVER MOUNTAINS, WYOMING

By NORMAN M. DENSON, HOWARD D. ZELLER, and E. VERNON STEPHENS

The name South Pass Formation is here used for a generally conglomeratic sequence of rocks that includes beds of sandstone, limestone, and volcanic ash; the rocks have a combined average thickness of about 350 feet. The formation unconformably overlies rocks of early and middle Miocene, early and middle Eocene, and Precambrian ages in the vicinity of South Pass, Fremont County, Wyo. (fig. 2).

The rocks in the South Pass Formation were first described by Comstock (1874) as beds of conglomerate, sandstone, and marl near South Pass, Wyo.; however, his names South Pass Group and South Pass Beds (Comstock, 1874, chart opposite p. 102 and p. 130, respectively) were never formally defined, and these names have not been used by other geologists working in the area (Wilmarth, 1938, p. 2032). A section of the rock sequence showing stratigraphic and structural relations is exposed along the Continental Divide near South Pass (alt 7,550 ft); therefore, the authors believe that the name South Pass, given to these beds by Comstock, should be retained as a formation name. The area around South Pass in Tps. 28 and 29 N. and Rs. 98 through 103 W. is herein designated as the type area.

The South Pass Formation is composed dominantly of a basal pinkish-gray pebble-to-boulder conglomerate having a matrix of fine-grained tuffaceous sandstone and siltstone that locally contains pebbles of moss agate. Locally the conglomerate is interbedded with very coarse grained arkosic sandstone and is overlain by fine- to coarse-grained laminated strongly fluorescent sandstone as much as 200 feet thick and fresh-water limestone and volcanic ash as much as The thickness of the formation ranges from 0 to more 15 feet thick. than 500 feet. More than 100 feet of this formation is exposed along many north-facing escarpments on the south side of the Sweetwater River near the center of the map area. There the South Pass Formation is heterogeneous and includes boulders and cobbles of Precambrian metamorphic and igneous rocks set in a matrix consisting of both fine-grained tuffaceous siltstone, reworked from the White River Formation (Oligocene), and tuffaceous sandstone, which is from the lower and middle Miocene sequence.

The South Pass Formation fills preexisting valleys and forms pediment and coalescing alluvial fanlike deposits along both flanks at the southern end of the Wind River Mountains. The heterogeneity of the rocks and the abrupt changes in lithology along the outcrop are



distinguishing characteristics of the unit and probably represent torrential, short-duration deposition following major uplift of the Wind River Mountains at the end of middle Miocene time. The formation has been mapped as irregular, discontinuous patches in an area of about 350 square miles which extends from the mouth of Rock Creek northwestward to a point near the head of Big Sandy Creek in T. 30 N., 104 W.—a distance of about 40 miles. The rocks herein referred to as the South Pass Formation have been designated as post-Bridgerian (McGrew and others, 1959, fig. 1), as basal part of the White River (Nace, 1939, pl. 1), and as middle or upper Eocene (Love and others, 1955).

Although the exact age of the South Pass Formation is unknown, it is no older than the lower and middle Miocene rocks which it overlies near the west-central edge of the map area. At various places in the area, the South Pass also overlies the Bridger (middle Eocene) and Wasatch (early Eocene) Formations and in addition Precambrian granite and metagraywacke. Northwest of the area in figure 2, in T. 30 N., R. 104 W., and along the northeast flank of the Wind River Mountains in Tps. 29 and 30 N., Rs. 98 and 99 W., the formation rests on the beveled edge of as much as 800 feet of the White River Formation (Oligocene). The South Pass Formation is overlain by pediment gravel and alluvium of Quaternary age.

Four miles west of the map area, near the McCann Ranch (SW<sup>1</sup>/<sub>4</sub> T. 28 N., R. 103 W., and NW¼ T. 27 N., R. 103 W.), and in the SW cor. T. 28 N., R. 102 W. (fig. 2), the formation is offset by movement along the Continental fault and the beds are tilted as much as 45°. The cementation of the basal conglomerate is noticeably greater along the Continental fault whereas elsewhere the conglomerate is generally semiconsolidated. If the latest movement along the Continental fault is post-middle Pliocene, which regional evidence indicates (Love, 1954, p. 1312), the South Pass Formation may be considered as late Miocene to middle Pliocene in age. Because the South Pass Formation can be traced eastward discontinuously into fossiliferous rocks assigned to the Ogallala Formation (upper Miocene and Pliocene) in southeastern Wyoming and because it has lithologic, stratigraphic, and structural relations similar to those of the Ogallala, a correlation between the two formations is suggested.
## NOMENCLATURE AND AGE OF FORMATIONS IN THE ANSONIA QUADRANGLE, FAIRFIELD AND NEW HAVEN COUNTIES, CONNECTICUT

## By CRAWFORD E. FRITTS

The names and geologic ages of most of the formations in this area are discussed by Fritts (1962a, b). Additions to the nomenclature and revisions of geologic ages are discussed briefly here.

The Oronogue Member of the Derby Hill Schist is named here for the unincorporated community of Oronoque (pronounced Or-o-noke') in the northwestern part of the adjacent Milford quadrangle. The type locality of this member is a long roadcut at the intersection of the Merritt Parkway and State Route 110 on the west bank of the Housatonic River about 1.300 feet south of the Ansonia guadrangle. The rocks exposed there are in the kyanite zone of regional metamorphism and are similar to rocks mapped as part of the Oronogue Member just north of the Far Mill River in the western part of Pine Rock Park in the Ansonia quadrangle. Abundant quartz-rich paragneiss, or siliceous metatuff, distinguishes this member from the predominant schist of the Derby Hill Schist, which is also well exposed just north of the Far Mill River. Farther north near Shelton and Derby the entire formation is highly sheared; the Oronoque Member there is difficult to recognize but presumably is in the upper part of the formation, and grades into the predominant schist. Rocks of the Oronoque Member in zones of low-grade metamorphism in the eastern part of the Ansonia and Milford quadrangles are also difficult to recognize, but in at least some places there the rocks appear to contain more primary (nonintroduced) silica than the underlying phyllites, which probably are equivalent to the main part of the Derby Hill The Oronoque Member, especially the upper part, also con-Schist. tains subordinate impure limestone of metasedimentary origin and amphibolite and greenschist, which probably represent mainly metavolcanic rocks similar to those found in the overlying Maltby Lakes Volcanics.

The names Maltby Lakes Volcanics and Allingtown Metadiabase were chosen in consultation with John Rodgers and J. E. Sanders of Yale University, whose students, Holdaway<sup>1</sup> and Burger,<sup>2</sup> mapped rocks of pre-Triassic age in the western part of the New Haven quadrangle during the period 1957-62 when work by the present writer was in progress in the Mount Carmel, Southington, Ansonia, and Milford quadrangles. The rocks mapped by Holdaway and

<sup>&</sup>lt;sup>1</sup>Holdway, M. J., 1958, The bedrock geology of the Maltby Lakes area [Conn.]: Yale Univ. unpub. senior thesis.

<sup>&</sup>lt;sup>3</sup> Burger, H. R., 1962, Stratigraphy and structure of the Milford Group, New Haven quadrangle, Connecticut: Yale Univ. unpub. senior thesis.

Burger formerly were included in the Milford Chlorite Schist (Rice and Gregory, 1906), a name recently abandoned (Fritts, 1962a).

A threefold division of the Milford Chlorite Schist was made by Burger. A southeastern unit, which Burger called the Savin Schist, consists mainly of chloritic phyllite similar to that exposed near Savin Rock in the New Haven quadrangle. This unit extends southwestward across the northwest corner of the Woodmont guadrangle and into the Milford quadrangle, where it is interpreted as low-grade Derby Hill Schist (Fritts, 1965). A central unit, which Burger called the Allingtown Formation, consists of abundant metadiabase intruded into phyllitic metasedimentary rocks. This unit extends southwestward into the southeast corner of the Ansonia guadrangle and into the adjacent Milford quadrangle, where the metasedimentary rocks are mapped separately as low-grade Oronoque Member of the Derby Hill Schist. Burger's northwestern unit, which includes rocks mapped previously by Holdaway near the Maltby Lakes Reservoirs in the New Haven quadrangle, consists mainly of metavolcanic rocks but contains subordinate metasedimentary rocks and minor intrusive metadiabase similar to that characteristic of Burger's Allingtown This northwestern unit of predominantly metavolcanic Formation. rocks also extends southwestward into the Ansonia and Milford quadrangles, where it lies above the Derby Hill Schist and unconformably below the Wepawaug Schist. The present writer uses the name Maltby Lakes Volcanics for the metavolcanic rocks and the name Allingtown Metadiabase only for intrusive metadiabase or metabasalt which is probably younger than both the Derby Hill Schist and the Maltby Lakes Volcanics.

The unit of rocks mapped as Maltby Lakes Volcanics is named here for the Maltby Lakes Reservoirs near State Route 34 in the New Haven quadrangle: this locality is also designated the type locality. The unit contains a distinctive basal pyroclastic schist, characterized by numerous lapilli and bomblike masses of metabasalt, but consists mainly of a thick sequence of greenschists and amphibolites. The rocks probably represent metamorphosed marine tuffs and(or) lava The formation also contains minor metasedimentary schists flows. and impure limestones. The stratigraphic position occupied by the Maltby Lakes Volcanics is comparable to that of the Barnard Volcanic Member of the Mississquio Formation of Doll, Cady, Thompson, and Billings (1961) in Vermont. The age of the Maltby Lakes Volcanics, therefore, probably is Ordovician(?). The basal pyroclastic schist is well exposed at the type locality and was mapped by Holdaway and Burger (see footnotes 1 and 2, p. A30) as a separate unit, although they did not identify it as a metamorphosed pyroclastic rock. They also mapped several thin but apparently distinctive greenschist and amphibolite units now known to be stratigraphically above the pyroclastic schist and below the Wepawaug Schist. The thin stratigraphic units beneath the Wepawaug, however, have not been mapped separately in the Ansonia and Milford quadrangles mainly because of a scarcity of outcrops in these areas and in the southwestern part of the New Haven quadrangle. In the Ansonia and Milford quadrangles, therefore, the name Maltby Lakes Volcanics is used for all metavolcanic (extrusive) rocks and subordinate metasedimentary rocks that are underlain by the Derby Hill Schist and are overlain unconformably by the Wepawaug Schist.

The Allingtown Metadiabase is named here for the community of Allingtown (spelled Allington on recent topographic maps of the New Haven quadrangle) just southwest of the city of New Haven on U.S. Highway 1. This is also the type locality of the formation. The rock is a metamorphosed porphyritic diabase or basalt, which forms numerous closely spaced sills and gently inclined dikes in the metasedimentary rocks interpreted by the writer as low-grade Oronoque Member of the Derby Hill Schist. Dikes of similar porphyritic rocks also intruded the Maltby Lakes Volcanics sometime before metamorphism but are not known to have intruded the Wepawaug Schist. The age of the metadiabase therefore is probably Ordovician.

The age of the Prospect and Ansonia Gneisses formerly was reported as Ordovician or Devonian (Fritts, 1962a); however, the age of the Ansonia now is believed to be Devonian, because in this quadrangle the rock appears to be a gneissic equivalent of the Woodbridge Granite. The Woodbridge intruded the Wepawaug Schist of Silurian and Devonian age before the climax of progressive regional metamorphism in Middle to Late Devonian time (Fritts, 1962b). The regional distribution of the Prospect Gneiss (Fritts, 1962a, fig. 128.1) suggests that this intrusive rock was emplaced mainly along the eastern flank of a north-trending line of domes, such as the Waterbury dome, after initial folding of the Wepawaug Schist and underlying metasedimentary rocks but before emplacement of the Ansonia Gneiss, which crosscuts the Prospect. Thus the age of the Prospect Gneiss also is reported here as Devonian.

# BELTED RANGE TUFF OF NYE AND LINCOLN COUNTIES, NEVADA

## By K. A. SARGENT, D. C. NOBLE, and E. B. EKREN

As the result of mapping about 2,000 square miles north and west of the Nevada Test Site during the past two years, problems in naming the volcanic rock units have developed that cannot be solved by use

of the previous nomenclature applied to rocks within the Test Site. Volcanic rocks, more than 10,000 feet thick and equivalent in age to the Indian Trail Formation (Poole and McKeown, 1962), have been divided into more than 20 mappable units, some of which are several thousand feet thick. In contrast, the Indian Trail Formation in the Test Site is only about 1,000 feet thick and is subdivided into an informal lower member and the Tub Spring and Grouse Canyon Members (Hinrichs and Orkild, 1961; Poole and McKeown, 1962). Though some age equivalence is certain, use of the name Indian Trail Formation for the thick sequence of volcanic rocks north of the Test Site would imply lithologic and genetic relations that do not exist and would preclude the development of a reasonable nomenclature for the rocks north of the Test Site. For these reasons, the name Indian Trail Formation is here restricted geographically to the eastern part of Nevada Test Site, that is, east of long 116°15′ (fig. 3).

The recent geologic mapping has shown that the Tub Spring and Grouse Canyon Members underlie a large area north and west of the Test Site. As the Indian Trail Formation is restricted to the Test Site, the lateral extensions of the rocks of the Tub Spring and Grouse Canyon Members are no longer parts of a formal stratigraphic unit outside the Test Site; therefore, north and west of the Nevada Test Site, the Tub Spring and Grouse Canyon Members are placed in a new formation herein named the Belted Range Tuff (fig. 4).

## BELTED RANGE TUFF

The Belted Range Tuff crops out over a broad area in the southern parts of the Belted and Kawich Ranges and in the vicinity of Pahute Mesa (fig. 3). In addition, the Grouse Canyon Member is present over an extensive area west of that shown on figure 3. The formation is best exposed in the southern part of the Belted Range, which is designated the type area. At most places the Belted Range Tuff unconformably overlies local informal units of tuff and lava and is unconformably overlian by tuff, lava, the Paintbrush Tuff, the Timber Mountain Tuff, or the Thirsty Canyon Tuff (p. A44; Noble and others, 1964). Many of the lava flows that underlie or overlie the tuff in the southern part of the Belted and Kawich Ranges have a soda-rhyolite composition and are genetically related to the tuff.

Both the Tub Spring and Grouse Canyon Members are compound cooling units of comenditic (peralkaline soda-ryholite) ash-flow tuff. Rocks of both members are peralkaline. Soda-rich sanidine is the dominant phenocryst mineral in rocks of both members; fayalite, apatite, zircon, and sodic iron-rich clinopyroxene are minor ubiquitous phenocrysts. The Grouse Canyon Member is virtually quartz



FIGURE 3—Distribution of the Tub Spring and Grouse Canyon Members of the Belted Range Tuff in the northeastern part of the Nevada Test Site and the southeastern part of the Las Vagas Bombing and Gunnery Range.

free, whereas the Tub Spring Member contains abundant quartz phenocrysts.

#### TUB SPRING MEMBER

At the type locality, Tub Spring (fig. 3), the Tub Spring Member is approximately 250 feet thick; north of Oak Spring Butte the member is locally 300 feet thick. Rocks of the member are typically buff or bluish gray but locally are brick red. The degree of welding

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE



<sup>1</sup> See p. A44-A51.



ranges from poor to dense. At most outcrops the member is devitrified, but poorly welded vitric tuff, in many places partly to completely zeolitized, commonly occurs at the base and top of the member; vitrophyre is locally present where the member is thick and densely welded. Phenocrysts compose approximately 20-25 percent of the member. Acicular crystals of acmite and sodic amphibole of vapor-phase origin are typical of poorly welded parts. The member locally contains large fragments of reddish-brown or buff pumice and lithic fragments of rhyolite, welded tuff, and Paleozoic sedimentary rocks.

## CONTRIBUTIONS TO STRATIGRAPHY

#### GROUSE CANYON MEMBER

At Grouse Canyon, the type locality (fig. 3), the Grouse Canyon Member is 80 feet thick; north of Oak Spring Butte it has a maximum thickness of approximately 300 feet. The member, as originally described by Hinrichs and Orkild (1961), is composed of an upper part of ash-flow tuff and a lower part of ash-fall and reworked tuff. Although the upper part of the member was described as a simple cooling unit (Poole and McKeown, 1962), further work has demonstrated that it is a compound cooling unit. In most places the unit is densely welded and, with the exception of a thin basal vitrophyre, is almost everywhere devitrified. Gas cavities containing vaporphase crystals of sodic amphibole are common. Rocks of the cooling unit range from greenish to bluish gray to gray buff and brown and locally are brick red. The various ash flows composing the compound cooling unit contain from less than 0.1 to approximately 20 percent phenocrysts, and average 5 percent.

The ash-fall and reworked tuffs composing the lower part of the Grouse Canyon Member are indistinguishable from ash-fall and reworked tuffs locally intercalated with lavas both above and below the member. Correlation of such tuffs over distances of several miles is uncertain; the lower part of the member is, therefore, no longer retained anywhere as a part of the member; these tuffs are treated as informal units of the formation.

On the basis of the data given on page A51, the Belted Range Tuff is of Miocene or Pliocene age.

## SALYER AND WAHMONIE FORMATIONS OF SOUTHEASTERN NYE COUNTY, NEVADA

## By F. G. POOLE, W. J. CARR, and D. P. ELSTON

A Tertiary sequence of lava flows, volcanic breccia, tuff, and sandstone is exposed in the southern part of the Nevada Test Site. Part of this sequence of rocks is divided into two new formations here named the Salyer and Wahmonie Formations. The Salyer Formation, which crops out in an area exceeding 300 square miles that is centered near Mount Salyer and Cane Spring (fig. 5), is best exposed in the Cane Spring quadrangle. Its maximum exposed thickness is nearly 2,000 feet at its type area in the vicinity of Mount Salyer and Cane Spring (fig. 6). The original volume of the Salyer was at least 20 cubic miles.

Overlying the Salyer is the Wahmonie Formation, which crops out in an area exceeding 500 square miles centered near its type area at Wahmonie Flat (fig. 5). The Wahmonie is best exposed in the Cane Spring and Skull Mountain quadrangles. Its maximum exposed thickness is 3,500 feet near Wahmonie Flat (fig. 6). The original volume of the Wahmonie was at least 25 cubic miles.



FIGURE 5.—Map of Nevada Test Site showing topographic quadrangles and approximate distribution of Salyer and Wahmonie Formations.

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## CONTRIBUTIONS TO STRATIGRAPHY



the Cane Spring quadrangle showing relations between volcanic units. f symbols given below:	<ul> <li>Salyer Formation (ruled pattern slanted to right):</li> <li>Unit 10 (Tsh<sub>1</sub>): breccia flow</li> <li>9 (Tsh<sub>2</sub>): interbedded lithic breccia, breccia flow,</li> <li>8 (Tst): lava flow</li> <li>8 (Tsh): breccia flow</li> <li>6 (Tsh<sub>2</sub>): breccia flow</li> <li>6 (Tsh): briotite tuff</li> <li>5 (Tsh<sub>2</sub>): interbedded tuff, sandstone, siltstone, and</li> <li>6 (Tsh<sub>2</sub>): interbedded tuff, sandstone, siltstone, and</li> <li>7 (Tsh<sub>2</sub>): breccia flow</li> <li>8 (Tsh<sub>2</sub>): breccia flow</li> <li>9 (Tsh<sub>2</sub>): breccia flow</li> <li>1 (Tsh<sub>1</sub>), tuffaceons clavatone</li> </ul>	T (I off). ULL account and account
FIGURE 6.—Paruly restored generalized north-south section in Explanation c	<ul> <li>Wahmonie Formation (ruled pattern slanted to left): Unit 5 (Twf): many lava flows and thin interflow tuffs 4 (Twc): conglomerate and lithic tuff-breccia 3 (Twb): lithic breccia</li> <li>2 (Twtb): interbedded tuff, sandstone, lithic tuff- breccia, and pumice agglomerate</li> <li>1 (Twfa): older lava flows</li> </ul>	

# CHANGES IN STRATIGRAPHIC NOMENCLATURE

The type section of the Salyer Formation is a composite of partial sections in the area from Cane Spring to Mount Salyer to Hampel Hill (figs. 5, 6). The type localities are as follows: Units 1 to 3, a northtrending ridge near the head of Mara Wash about 1¼ miles south of Cane Spring; units 4, 5, 9, 10, a northeast-trending ridge about half a mile southeast of Cane Spring; unit 6, about 1 mile northeast of Hampel Hill; units 7 and 8, about three-fourths of a mile west of Mount Salyer. As only part of unit 9 is present on the ridge half a mile southeast of Cane Spring, a well-exposed section extending for about half a mile northeast from Mount Salyer also is considered the type locality for this unit.

The type section of the Wahmonie Formation is also a composite of partial sections found on the east and south sides of Wahmonie Flat, at Mount Salyer, and in the area between Cane Spring and Hampel Hill (figs. 5, 6). The type localities are as follows: Unit 1, a small exposure on a low ridge about 1¼ miles northwest of Cane Spring; unit 2, on a northeast-trending ridge about half a mile southeast of Cane Spring and around Hampel Hill; unit 3, at Mount Salyer; unit 4, on the northwest side of Hampel Hill; unit 5, incomplete sections at the east end of Wahmonie Flat and on the north side of Skull Mountain.

#### STRATIGRAPHIC RELATIONS

In general the Salyer and Wahmonie Formations occur in the middle part of the Tertiary volcanic sequence at the Nevada Test Site (Poole and McKeown, 1962, fig. 80.2). In the type area the Salyer Formation unconformably overlies yellowish zeolitized rhyolitic tuffs of Mara Wash and underlies the Wahmonie Formation (fig. 6). The Wahmonie is older than most of the Paintbrush Tuff (see p. A44) of the Piapi Canyon Group. On Skull Mountain the Wahmonie is directly overlain by the Topopah Spring Member of the Paintbrush. Northwest of Wahmonie Flat a late flow of the Wahmonie occurs between the Topopah Spring and Tiva Canyon Members of the Paintbrush.

Although the Salyer Formation is in general overlain by the Wahmonie, one unit of the Salyer (Tsb<sub>4</sub>, fig. 6) intertongues with the tuffs of the Wahmonie south of Cane Spring.

## LITHOLOGIC DESCRIPTION

In some areas the Salyer and Wahmonie are not easily separated, but generally rocks of the Wahmonie are more mafic, contain primary hornblende, are less altered, and consist mainly of lava flows and related tuffs, whereas the rocks of the Salyer Formation are more acidic, contain primary pyroxene and secondary hornblende, are more altered, and consist of breccia flows and interstratified tuff, sandstone, and volcanic breccia.

The Salyer Formation is divided into 10 informal units and the Wahmonie into 5 informal units, as described and illustrated on figure 6. Unless otherwise stated, contacts between units are sharp or abruptly gradational.

The following definitions are for terms used in the stratigraphic sections of the Wahmonie and Salver Formations.

- Lithic tuff-breccia. An extrusive multilithic breccia that is composed of fragments of angular to rounded older volcanic rock and sparse pre-Cenozoic sedimentary rock contained in a matrix of lapilli and ash tuff.
- Lithic breccia. A breccia similar to lithic tuff-breccia except that it has a matrix of finely fragmented previously formed volcanic rock. In addition, lithic breccia includes some monolithic breccia.
- **Pumice agglomerate.** An extrusive monolithic rock that is composed mainly of angular to rounded pumice blocks and lapilli in a matrix of ash tuff.
- **Breccia flow.** Lava flows that are almost entirely composed of angular to subrounded fragments of altered previously solidified lava contained in a less altered lava matrix of similar composition. Most fragments probably formed prior to, rather than during, emplacement of unit. Fragments occur throughout entire flow body.
- Flow breccia. An extrusive breccia on the margins of lava flows, formed by the breaking up of lava that was solidifying as it moved. The matrix is solidified lava and comminuted flow rock.

Composite sections of	' Wahmonie	and Salyer	Formations
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[Maximum measured thickness is given]

Wahmonie Formation:	(feet)
5. Black, gray, red, purple, and olive rhyodacitic to dacitic lava flows	
and related thin internow tuffs and preceiss. Lavas are flow	
banded, sheeted, porphyritic, glassy to stony, and contain	
about 40 percent phenocrysts of labradorite, orthopyroxene,	
clinopyroxene, hornblende, biotite, and magnetite. Fine-	
grained mafic inclusions are common. Intense hydrothermal	
alteration is common locally in Wahmonie Flat and Pluto	
Valley. Individual flows are as much as 800 ft thick. Con-	
tacts with older units are poorly exposed in Wahmonie Flat	
area. Unit 5 intertongues with unit 2 and contains at least	
one major erosional unconformity	3,500+
4. Conglomerate or lithic tuff-breccia consisting of porphyritic stony	
and glassy flow-banded rhyodacite and dacite cobbles and	
boulders in a gray tuffaceous sandy matrix. Unit 4 is present	
only in the southwestern part of the Cane Spring quadrangle	50
3. Gray to pinkish-gray lithic breccia consisting of dense to vesicu-	• •
lated, glassy to devitrified blocks crudely stratified in 1- to	
2 ft-thick zones. This unit has a few thin local interbeds of tuff	
and sandstone	270
	~10

Thickness Wahmonie Formation—Continued

- 2. Interbedded tuff, sandstone, lithic tuff-breccia, and pumice agglomerate. The tuff is white, gray, pink, green, and brown; locally contains abundant lithic inclusions of volcanic rock and is partly zeolitized and is generally well bedded. The sandstone is white, gray, or red, very fine to coarse grained, thinly laminated to thick bedded. Lithic tuff-breccia consists of red, purple, and gray blocks of volcanic rock in gray to green tuff matrix. Many breccia zones are crudely stratified, and some are glassy and similar to some of the breccia of unit 3. The pumice agglomerate consists of gray to tan pumice in a lapilli and ash matrix. The pumice contains 25 percent phenocrysts of plagioclase, biotite, and some hornblende and quartz. Silicified conifer wood is common at several places in unit 2. Unit 2 contains a lens of unit 10 of the Slaver, and its basal contact is locally gradational with unit 9 of the Salyer. Relations with unit 5 of the Wahmonie north of the Cane Spring fault are not clear\_\_\_\_\_ 1, 700 1. Red, purple, and brown rhyodacitic lava flows, commonly hydrothermally altered. This unit probably does not occur south of the Cane Spring fault Salyer Formation:
  - 10. Purple and red breccia flow, petrographically like other breccia flows of the Salyer. The breccia contains 30 percent phenocrysts of andesine-labradorite, partly altered biotite, altered pyroxene, and secondary hornblende, hematite, quartz. and plagioclase. Crudely sorted and locally layered. The unit occurs as a tongue within the Wahmonie Formation in the Cane Spring area\_\_\_\_\_
  - 9. Interbedded lithic breccia, breccia flow, lithic tuff-breccia, tuff. and sandstone. Top of unit is a gray to green glassy breccia flow as thick as 320 ft. The rocks of this unit contain about 30 percent phenocrysts of labradorite, hypersthene, and biotite. Middle of unit 9 is a lithic breccia as thick as 280 ft that contains large blocks of volcanic rock and a few pieces of Paleozoic carbonate rock. Lower part of unit is gray, pink, and yellow lithic tuff-breccia, crudely bedded locally. At the base of the unit is a widespread grav and green interbedded tuff and sandstone as thick as 90 ft. A major erosional unconformity is at the base. Upper contact is gradational in Cane Spring area
  - 8. Purplish lava flow and minor flow breccia. Flow is holocrystalline-porphyritic rock that contains 25 percent phenocrysts of labradorite and altered biotite. Lower contact is locally gradational with unit 7\_\_\_\_\_ 50
  - 7. Purplish- to reddish-brown breccia flow, similar to unit 10 but thicker. Major erosional unconformity at base transects some earlier structures\_\_\_\_\_ 450 6. Gray, pink, and orange biotite tuff that contains volcanic rock
  - fragments. Unit is probably two ash-flow tuffs separated by a persistent thin stratified tuff zone. Widespread redbrown sandstone layer is at top. Unit is not present in the Salyer type section near Cane Spring 200

Thicks ness (feet)

100 +

50

700

#### Salyer Formation—Continued

- 5. Interbedded tuff, sandstone, siltstone, and claystone. The tuff is mainly olive, green, yellow, and red porphyritic ash and pumice. Reddish-brown and green tuff, containing conspicuous white altered pumice lumps, marks base of unit. Sandstone and siltstone are gray, green, red, and brown, tuffaceous, pebbly, very fine to very coarse grained. Claystone is pinkish gray, silty, sandy, and pebbly, and contains abundant small fragments of volcanic rocks, and a few small pieces of pre-Cenozoic sedimentary rocks in the lower part\_\_\_\_\_\_
- 4. Breccia flows, purple and red in lower part, grading upward through pale red, pale purple, and pale yellow zones into yellow and olive. Lower part similar to units 7 and 10. Reddish blocks as large as 20 ft across occur in upper part of this unit. Several red-brown breccia zones are in upper part of unit\_\_\_\_ 1,000
- Pink, brown, gray, and red interbedded tuff, sandstone, lithic tuff-breccia, and lithic breccia. A blood-red sandstone 1-3 ft thick marks the top of unit 3. The lithic tuff-breccia and lithic breccia contain blocks of volcanic rock and some pre-Cenozoic sedimentary rocks. A widespread breccia zone as much as 15 ft thick in the upper part of unit 3 contains pre-Cenozoic rock fragments\_\_\_\_\_\_\_\_250
   Mottled purplish, reddish, and brownish breccia flow similar to units 7, 10, and lower part of unit 4\_\_\_\_\_\_\_\_150
   Reddish-orange to purplish tuffaceous claystone that locally contains pebbles of pre-Cenozoic sedimentary rocks; weathers

to soft clayey slope\_\_\_\_\_

## CHEMICAL COMPOSITION

Rocks of the Salyer and Wahmonie sequence grade upward chemically and mineralogically from acidic to intermediate composition. Plots of the major elements against silica for rocks of the Wahmonie through the Salyer generally form smooth curves along a differentiation trend common to many volcanic sequences. Chemical analyses and norms of several breccia flows and a lava flow indicate that the Salyer Formation is dellenitic to rhyodacitic in composition. See Nockolds (1954). Lavas of the Wahmonie are modally dacite and andesite, but chemical analyses and norms of several lava flows indicate that the lower part of the Wahmonie is rhyodacite and that the upper part is dominantly dacite. The chemical and mineralogical similarities of the Salyer and Wahmonie strongly suggest that the two formations are comagmatic.

#### AGE

The Salyer and Wahmonie Formations are dated as late Miocene and as late Miocene and early Pliocene(?), respectively, primarily on potassium-argon age determinations. The Topopah Spring Member of the Paintbrush Tuff (see p.A45-A47), which directly overlies the Wah-

Thickness (feet)

150

35

monie, has been dated at 13.3 million years  $\pm 0.5$  million years (R. W. Kistler, written commun., 1963). The lower and upper part of the Wahmonie, however, have been dated by the same method by Kistler (written commun., 1963) at 12.9 million years  $\pm 0.5$  million years and 12.5 million years  $\pm 0.5$  million years, respectively. The precision error in these dates may account for the apparent age inversion, as the dates of the Topopah and Wahmonie are only 0.4–0.8 million years apart, less than the possible error. Present data, therefore, place the Wahmonie very close to the Miocene-Pliocene Epoch boundary (Kulp, 1961), and as the Salyer underlies the Wahmonie its age is tentatively considered to be late Miocene.

## PAINTBRUSH TUFF AND TIMBER MOUNTAIN TUFF OF NYE COUNTY, NEVADA

## By PAUL P. ORKILD

Various names have been applied to the Tertiary rocks in the Nevada Test Site since the turn of the century. In 1907 Ball (p. 31-34) described some of the Tertiary rocks as rhyolite and latite flows and correlated other with the "Siebert tuff (lake beds)" (Spurr, 1905, p. 51-55). Fifty years later the Tertiary volcanic rocks in the eastern part of the Nevada Test Site were named the Oak Spring Formation by Johnson and Hibbard (1957, p. 367) (fig. 7). In 1957 Hansen, Lemke, Cattermole and Gibbons (1963, p. A7) divided the Oak Spring Formation at Rainier Mesa into numbered units (fig. 8). Hinrichs and Orkild (1961, p. D96) divided the Oak Spring Formation in the Yucca Flat area into seven named members and one informal member. In 1962 Poole and McKeown (p. C60, C61) raised the Oak Spring in the Nevada Test Site to the rank of a group and included two formations in it (fig. 8).

Additional detailed mapping of the Test Site and vicinity since 1962 has resulted in new stratigraphic information that requires another revision of the stratigraphic nomenclature of the Tertiary rocks. The use of Oak Spring Group for all Tertiary rocks in the Test Site is no longer practicable because the Tertiary rocks have been divided into a large number of units which have complex relations among the units. This group, therefore, is herein abandoned. Rockstratigraphic units that have similar chemical and mineralogical composition and mode of deposition constitute the major genetic units and are given formational rank. New data have shown that the Piapi Canyon Formation includes two lithologically dissimilar sequences of volcanic rocks that have different source areas. As each sequence is comprised of several units, the Piapi Canyon Formation herein is redefined and raised to the rank of group. It includes two new forma-



FIGURE 7.—Nevada Test Site and vicinity showing localities referred to in text. Area mapped by Johnson and Hibbard (1957) shown by stipple. The surface and subsurface distribution of Pah Canyon Member of the Paintbrush Tuff is enclosed by dashed line.

tions here named the Paintbrush Tuff and the overlying Timber Mountain Tuff (fig. 8). The rank of the older Indian Trail Formation is unchanged. The names Piapi Canyon Group and Indian Trail Formation, however, are restricted to the areas where they were originally used: namely, the Yucca Flat and Frenchman Flat areas (figs. 7, 9). This is the area (eastern part of Nevada Test Site) mapped by Johnson and Hibbard (1957).

## PAINTBRUSH TUFF

The Paintbrush Tuff includes the Stockade Wash, Topopah Spring, and Tiva Canyon Members of Hinrichs and Orkild (1961), the Yucca



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CHANGES IN STRATIGRAPHIC NOMENCLATURE



FIGURE 9.---Schematic diagram of the Tertiary volcanic rocks.

Mountain Member of Lipman and Christiansen (1964), and a new member herein defined as the Pah Canyon Member. These rocks are all similar in mineralogy, chemistry, and areal distribution. The type locality of this formation is in Paintbrush Canyon, which is located 2½ miles northeast of Yucca Mountain (fig. 7).

The Survey Butte Member has been a useful term for certain rocks in the eastern part of the Nevada Test Site. As additional data have become available, however, positive correlation of similar rocks in the same stratigraphic interval has become increasingly difficult as rocks are found farther from the type locality of the member. This difficulty results partly from the fact that the rocks of this member are derived from many sources. To avoid the connotation of positive correlation and implies genetic relations, therefore, the name Survey Butte Member is here abandoned. Rocks in this stratigraphic interval will hereafter be designated as bedded tuff of Paintbrush Tuff.

## PAH CANYON MEMBER

The Pah Canyon Member of the Paintbrush Tuff is here named for Pah Canyon (fig. 7). The member is a simple cooling unit of rhyolitic ash-flow tuff and is 280 feet thick at the type section at Pah Canyon. It has a maximum thickness of 300 feet at Yucca Mountain, but generally ranges from 40 to 180 feet in thickness. In most outcrops the member includes a basal zone of ash-fall tuff; this tuff is generally less than 3 feet thick but is locally as much as 15 feet thick. In much of the area of outcrop, the member consists of light-gray to light-brown partly to densely welded devitrified ash-flow tuff containing 5–15 percent phenocrysts of biotite, alkali feldspar, and plagioclase; quartz and clinopyroxene are very rare. The tuff of this member is characterized by abundant small pumice and felsic lithic inclusions and by abundant biotite. The member and the overlying Yucca Mountain Member.

## TIMBER MOUNTAIN TUFF

The well-exposed Timber Mountain Tuff attains its maximum thickness in the Timber Mountain region (fig. 7), which is the type area for the formation. It is unconformable on the Paintbrush Tuff. The formation includes the Rainier Mesa Member formerly of the Oak Spring Formation described by Hinrichs and Orkild (1961); this member is overlain by two informal members and one formal member, which are, in ascending order: tuff of Cat Canyon, tuff of Transvaal, and Ammonia Tanks Member.

The Rainier Mesa Member and Ammonia Tanks Member have greater areal extent than any other members of the Timber Mountain Tuff. They are nearly coextensive and crop out in approximately a 330° arc around Timber Mountain. The tuff of Cat Canyon is limited to Timber Mountain, and the tuff of Transvaal is restricted to an area southwest of Timber Mountain near Camp Transvaal (fig. 7).

Preliminary study indicates that the four members of the Timber Mountain Tuff have similar chemical and mineralogical compositions. These similarities suggest that the members are comagmatic. Phenocrysts, which comprise 10-40 percent of the rock, include quartz, alkali feldspar, plagioclase, and biotite.

## RAINIER MESA MEMBER

The Rainier Mesa Member is well exposed in Piapi Canyon and on Pahute Mesa (fig. 7). The member is a compound cooling unit of rhyolitic to quartz latitic ash-flow tuff containing abundant phenocrysts. Mafic minerals are generally most abundant in the upper or quartz latite part of the cooling unit.

#### TUFF OF CAT CANYON

The tuff of Cat Canyon is best exposed in Cat Canyon on the east flank of Timber Mountain (fig. 7). The tuff consists of a lower composite ash-flow sheet as much as 2,850 feet thick and an upper compound ash-flow cooling unit 0-150 feet thick (terminology of Smith, 1960). Despite its considerable thickness, tuff of Cat Canyon is confined to the collapsed area of the Timber Mountain caldera (Byers and others, 1963).

## TUFF OF TRANSVAAL

The tuff of Transvaal crops out south and southwest of Camp Transvaal, an abandoned mining camp 5 miles southwest of Timber Mountain (fig. 7). The tuff consists of a lower nonwelded rhyolitic ash-flow tuff 0-200 feet thick and an upper compound cooling unit of rhyolitic ash-flow tuff which locally is as much as 300 feet thick.

## AMMONIA TANKS MEMBER

The newly named Ammonia Tanks Member is well exposed both a quarter of a mile north of Ammonia Tanks (fig. 7), the type locality for the member, and for several miles along the south rim of Pahute Mesa (fig. 7). The member is a composite cooling unit of rhyolitic to quartz latitic ash-flow tuff which locally reaches a thickness of almost 300 feet in the vicinity of Pahute Mesa but in most places ranges in thickness from 50 to 200 feet. At the type locality the member is 250 feet thick. The lower 5–20 feet of the member is composed of light-gray or pink vitric ash-fall tuff containing about 15 percent white pumice fragments. The basal vitric zone is overlain in much of the area by pink to reddish-purple moderately to densely welded generally devitrified ash-flow tuff containing abundant phenocrysts.

The Ammonia Tanks Member is characterized by wedge-shaped sphene crystals as large as 1 mm and by locally abundant reddishbrown porphyritic lithic fragments derived mainly from the underlying Rainier Mesa Member.

## AGE

The age of the Oak Spring Group of Poole and McKeown (1962, p. C61), as indicated by plant, invertebrate, and vertebrate fossils, ranges from Eocene to Pliocene.

Potassium-argon age determinations of biotite and sanidine from various units in the Piapi Canvon Group have been made by R. W. Kistler (written commun., 1963), but the data are incomplete; adequate discussion of the results is beyond the scope of this paper. The data do provide, however, an interim and tentative basis for dating the formations. The base of the Indian Trail Formation in the northern part of the Test Site is about 16 million years old; the age of its top is unknown but is older than 13.5 million years. The Paintbrush Tuff ranges from about 13.5 to 12.5 million years in age, and the Timber Mountain Tuff has an age of about 10.5 million years. Translation of these potassium-argon dates to epochs in the geologic scale must be done somewhat arbitrarily because of the inconsistent placement of epochs recorded in the literature (Holmes, 1947, p. 145; Kulp, 1961; and Evernden and others, 1964). Using Kulp's geologic time scale, in which the Miocene-Pliocene boundary is between 12 to 14 million years before the present, the Indian Trail Formation is Miocene and Pliocene(?) rather than Miocene or Pliocene as given in Poole and McKeown (1962, p. C62). the Paintbrush Tuff is Miocene(?) and Pliocene, and the Timber Mountain Tuff is Pliocene.

## AGE OF THE ELEANA FORMATION (DEVONIAN AND MIS-SISSIPPIAN) IN THE NEVADA TEST SITE

By F. G. POOLE, P. P. ORKILD, MACKENZIE GORDON, JR., and HELEN DUNCAN

Field investigations in the Nevada Test Site by the authors and colleagues in August 1961 provided considerably more extensive fossil collections from the Eleana Formation than were available when an earlier paper on the formation was prepared (Poole and others, 1961). These collections provided a more reliable basis for dating than the previous collections from which the Eleana was dated as Mississippian and Early Pennsylvanian.

The Eleana was divided into ten informal units lettered A (at the base) through J (at the top). As mapped, unit A contains faunal assemblages that cannot be younger than early Late Devonian (Frasnian). The Devonian corals originally collected were fragmentary specimens in the matrix of a limestone conglomerate. Laboratory study did not establish that these specimens were not derived from older rocks and redeposited in beds of Mississippian age; hence, the age of unit A was given as Mississippian, even though it was suggested that unit A might be as old as Lake Devonian. Field study later revealed, however, that layers containing *Atrypa* occur between the conglomeratic coral beds. Inasmuch as *Atrypa* is not known in rocks younger than the lower Upper Devonian and as no fossils characteristic of the Mississippian have been found in unit A, it seems safe to conclude that unit A is of Devonian age.

Units B to G of the Eleana Formation have not yielded fossil collections that provide evidence for precise dating. The authors have recognized no undoubtedly Early Mississippian assemblages in the Eleana; however, several collections which contain fossils that cannot be older than Late Mississippian also contain some fossils that probably were derived from beds of Early Mississippian age.

In 1964 Gordon and Duncan completed a detailed study of the fossils collected from the upper part of the Eleana and the lower beds of the overlying Tippipah Limestone. Information gained from their work indicates that units H, I, and J are of Late Mississippian age and that no rocks of Pennsylvanian age are included in the formation. The basal beds of the Tippipah Limestone, at a locality in the C. P. Hills northwest of Frenchman Flat, 2.2. miles southwest of Yucca Pass in the Yucca Lake quadrangle, contain goniatites (*Branneroceras, Bisatoceras*, and *Boesites?*) indicative of Early Pennsylvanian age.

One collection from unit H or I in the Eleana in the foothills of Quartzite Ridge, in the Oak Spring quadrangle, was originally considered to be of Pennsylvanian age because the only identifiable fossils belonged to a species of the horn coral Lophophyllidium, a genus generally considered to be diagnostic of Pennsylvanian and Permian age. The author's later collection from the same locality yielded, among other fossils, the lithostrotionoid coral Siphonodendron (Cionodendron) sp. and the brachiopod Reticulariina cf. R. spinosa (Meek and Worthen), which are characteristic Late Mississippian fossils. Moreover. at another locality specimens of Lophophyllidium were associated with an extensive assemblage of corals, bryozoans, and brachiopods among which the critical elements for dating purposes are forms of early Late Mississippian age (Zaphrentites aff. Z. californicus (Tischler) Cyathaxonia n. sp., Antiquatonia sp., Auloprotonia? sp., and Semicostella n. sp.).

Several collections made from conglomeratic limestone beds in unit I of the Eleana indicate that at least some of the fossils were derived from erosion of considerably older rocks. A stromatoporoid and

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corals of Devonian aspect associated with corals and brachiopods indicative of both Early and Late Mississippian age provide evidence of such reworking; however, the conglomeratic layers containing mixed associations occur well below unit J, which carries go liatites (*Cravenoceras hesperium* Miller and Furnish and *C. merriami* Youngquist, which are typical of the *Eumorphoceras bisulcatum* zone) and other fossils characteristically found in late Late Mississippian faunas in the Great Basin. The faunal assemblages in which these goniatites occur seem to be natural associations that are not contaminated by elements diagnostic of the Pennsylvanian or of the pre-Late Mississippian.

## MARINETTE QUARTZ DIORITE AND HOSKIN LAKE GRANITE OF NORTHEASTERN WISCONSIN

## By WILLIAM C. PRINZ

Much of northern Wisconsin is underlain by granitic rocks of Precambrian age; these rocks compose the so-called Wisconsin batholith. The U.S. Geological Survey, as part of its restudy of the Menominee district, has mapped a small part of the northern edge of the batholith in northern Marinette County and Florence County, Wis. Two units recognized within the batholith—Marinette Quartz Diorite and Hoskin Lake Granite (Prinz, 1959)—are here adopted for use by the Geological Survey. Cain (1963) has mapped the extension of these two units to the south and has distinguished several other granitic units.

Marinette Quartz Diorite forms a pluton of about 12 square miles, mainly in Marinette County. The diorite is here named for that county. It had previously been referred to as biotite-hornblendequartz diorite by Lyons (Emmons and others, 1953, pl. 16 and p. 108). Good exposures of Marinette Quartz Diorite occur south of Niagara, Wis., in the NW¼ sec. 22 and SW¼ sec. 15, T. 38 N., R. 20 E., and along the Chicago, Milwaukee, St. Paul and Pacific Railroad in the  $S_{\frac{1}{2}}NE_{\frac{1}{4}}$  sec. 18, T. 38 N., R. 20 E.

Typical Marinette Quartz Diorite is medium grained, is massive or slightly foliated, and consists of 40-50 percent oligoclase, 0-20 percent potassium feldspar, 10-30 percent quartz, and 20-30 percent prochlorite, biotite, or hornblende. The texture of the rock is hypidiomorphic to allotriomorphic granular consisting of large subhedral to anhedral plates of plagioclase and flakes of biotite or chlorite and consisting of finer grained interstitial grains of quartz and potassium feldspar. Commonly, the oligoclase shows normal zoning and has cores as calcic as andesine; most plagioclase is partly altered to sericite or epidote. Most potassium feldspar is unaltered and shows microcline grid twinning. The major mafic constituent of the quartz diorite is either prochlorite or biotite, and some biotitic samples also contain hornblende. Ubiquitous accessory minerals are magnetite, zircon, and apatite: pyrite, calcite, and tourmaline are found locally.

Hoskin Lake Granite forms an arucate-shaped intrusion of at least 20 square miles that in most places lies between metabasalts of the Quinnesec Formation to the north and Marinette Quartz Diorite to the south. It is here named for Hoskin Lake which lies within the granite pluton in the NE¼ sec. 23, T. 38 N., R. 19 E. Previous workers referred to this unit as porphyritic granite or granite porphyry. Good exposures are found almost everywhere within the mapped limits of the unit; those exposures near the county secondary roads in the S½ sec. 12 and N½ sec. 13, T. 38 N., R. 19 E., and the NE¼ sec. 17, T. 38 N., R. 20 E. are readily accessible.

Most Hoskin Lake Granite is coarse grained and porphyritic, having large subhedral, white, gray, or flesh-colored microcline phenocrysts as long as 2½ inches; the phenocrysts are set in a matrix of partially sericitized oligoclase, quartz, microcline, and biotite. In many places the phenocrysts are more or less oriented to form a crude foliation. Chlorite and some hornblende occur with the biotite, and epidote is a sparse alteration product of oligoclase. Zircon, apatite, and subhedral to euhedral sphene, in places altered to leucoxene, are ubiquitous accessory minerals; allanite, monazite, tourmaline, calcite, magnetite, and pyrite are also present in some of the granite. Commonly, half the granite consists of microcline phenocrysts; locally, however, the number of phenocrysts diminishes, and the granite grades into almost nonporphyritic coarse-grained quartz monzonite identical with the matrix of the porphyritic granite.

Hoskin Lake Granite intrudes Marinette Quartz Diorite, and both units intrude metabasalts of the Quinnesec Formation. The contacts between metabasalt and granite or quartz diorite are sharp. The contacts between granite and quartz diorite are generally gradational, because near most of them the granite contains many quartz diorite inclusions and the quartz diorite is cut by many stringers and dikes of granite.

The age of neither the Marinette Quartz Diorite nor the Hoskin Lake Granite is known with certainty. Intrusive relations show that these rocks are younger then the Quinnesec Formation, which is considered to be of early Precambrian age, and Hoskin Lake Granite is intruded by unmetamorphosed diabase dikes of Keweenawan age (late Precambrian). Their age relation to the Animikie metasedimentary rocks (middle Precambrian) is controversial, Prinz (1959) considered them both to be pre-Animikie, whereas R. W. Bayley (written commun., 1960), who has restudied the eastern part of the Menominee district in Dickinson County, Mich., concluded that they are post-Animikie.

## MASHEL FORMATION OF SOUTHWESTERN PIERCE COUNTY, WASHINGTON

#### By KENNETH L. WALTERS

The name Mashel Formation is given herein to a sequence of unconsolidated fluvial and lacustrine deposits of Miocene age that unconformably underlie Pleistocene deposits and overlie consolidated rocks in southwestern Pierce County, Wash.

J. E. Sceva proposed in 1955 (written commun.) that these deposits be named after the Mashel River (fig. 10), in whose valley walls they are exposed.

## CHARACTER AND THICKNESS

The Mashel Formation consists of a predominantly fine-grained upper part and a coarse-grained lower part. The upper part is composed mostly of clay and sand. The clay is predominantly





light colored and commonly contains plant material ranging in degree of preservation from unidentifiable fragments to whole leaves and sections of logs. The sandy phases of the upper part of the formation contain tuffaceous material, pumice, and volcanic ash.

The lower part of the formation is composed mostly of iron-stained medium- to coarse-grained poorly cemented gravel of predominantly dark volcanic rock types. Most of the granitic pebbles, where present, are badly decomposed and crumble if struck with a hammer.

The greatest thickness of the Mashel observed is about 225 feet. The total thickness may be in excess of 500 feet. Locally, the finegrained upper part may be more than 400 feet thick.

The following section, measured by the author and Grant E. Kimmel, is designated the type section. An excellent but rather inaccessible section of the formation was measured by J. E. Sceva and B. A. Liesch in the north bluff of the Mashel River about 1,350 feet east of a power transmission line, near the center of sec. 20, T. 16 N., R. 4 E.

Type section of Mashel Formation measured along Weyerhaeuser Road descending from Mashel Prairie to Mashel River, SE4SW4 sec. 20 and NE4NW4 sec. 29, T. 16 N., R. 4 E.

Vashon 1	Drift:
----------	--------

Gravel, coarse-grained.	m talan ara
Mashel Formation:	(feet)
Clay, massive, rusty; weathers white to cream; contains organic ma	-
terial. Lower few inches contains streaks of fine sand	_ 7
Sand, medium, angular to subangular, tan; composed of clear quartz biotite, muscovite, and miscellaneous dark rock fragments; include	s, s
thin beds of silty clay that contain a small amount of organic ma	-
terial and lenses of partly decomposed pebbles having a purple hue.	_ 14
clay, massive, cream to white; has rusty streaks; contains bits of or ganic material; becomes silty and sandy near base	- - 7
Sand, medium, angular to subangular; contains bits of organic ma terial	- 2
Clay, massive, cream to white; contains organic material	. 1.5
Sand	. 1.9
Clay, cream to white; contains wood and leaves near base (fossil loc. 4) _ Sand, gray if wet; contains silt and clay; grades laterally into coars	. 3 e
sand and gravel	- 3.8
Lignite	- 0.7
Clay, massive; appears waxy on fractures; contains scattered organi- material and fragments of tuff; has same thin sand beds	c - 14
Covered	- 22
Gravel, coarse, and boulders, rusty: contains interstratified lenses o sand and blocks of clay that are as much as 5 ft across; has some granitic boulders, badly decomposed; base not exposed	f e _ 75

Section of Mashel Formation measured in north bluff of Mashel River near the center of sec. 20, T. 16 N., R. 4 E.

[Measured September 12, 1953 by J. E. Sceva and B. A. Liesch]

ashel Formation:
Sand, silty, stratified, light-brown
Sand and fine gravel, stratified
Sand, fine, cemented, light-brown
Conglomerate; composed of small white pumice pebbles
Sand, stratified, rusty-red
Clay and silt, gray
Sand, crossbedded, gray; contains many streaks of white pumice pebbles
Sand, silty, brown
Sand and gravel, rusty; composed predominately of pink and gray andesite
Sand, silt, and ash
Ash, hard, brittle, brown
Tuff, hard, green
Sand, rusty, brown
Clay, and fine sand, silty, laminated
Clay, gray, and fine sand
Silt, sand, and clay
Lignite
Sand, silt, and clay, stratified, gray
Gravel, basaltic, cemented, rusty; base not exposed

#### AGE AND CLIMATIC IMPLICATIONS-

Floras collected by the author and Grant E. Kimmel from the Mashel at the localities indicated in the following table were identified by Jack A. Wolfe of the U.S. Geological Survey:

		Locality			
Flora	1	2	3	4	
Pinus ponderosa Lawson Populus trichocarpa Torrey and Gray Salix hesperia (Knowl.) Condit sp., n. sp Carya simulata (Knowl.) Brown		 × ×	× × ×	×	
Pterocarya mixta (Knowl.) Brown sp., n. sp Alnus relata (Knowl.) Brown sp., n. sp sp., cones	×	× × ×	 	×	
Betula lacustris MacGinitie Fagus sanctieugeniensis Hollick Quercus chrysolepis Liebmann Ulmus speciosa Newberry Zelknova oregoniana (Knowl.) Brown		×××× ××××	×		
Mahonia reticulata (MacG.) Brown Cinnamomum sp., n. sp Persea lanceolata (Berry) Brown Platanus dissecta Lesquereux Sp., n. sp	 × 	× × ×	×××		
Acer sp., n. sp macrophyllum Pursh Paulownia columbiana Smiley Cornus sp., n. sp Fraxinus sp		× ×	× ×	× 	

Tanwax Lake quadrangle, SW¼SE¼ sec. 22, T. 17 N., R. 4 E., on east shoulder of Clear Lake highway, 478 ft north of Golden Road junction. Altitude about 675 ft.
 Eatonville quadrangle, SE¼NE¼ sec. 25, T. 16 N., R. 3 E., on southeast side of Weyerhaeuser logging road, in bluff southeast of Ohop Creek, 260 ft northeast of bend in road at creek level. Altitude about 525 ft.
 Eatonville quadrangle, NW¼SE¼ sec. 18 T. 16 N., R. 4 E., on west side of State highway 7, in bluff west of Ohop Valley.
 Eatonville quadrangle, NW¼NW¼ sec. 29, T. 16 N., R. 4 E., on northwest shoulder of Weyerhaeuser logging road, in bluff west of Mashel River. Altitude about 525 ft.

Wolfe (written commun., 1961) assigned the flora from locality 2, and also that from locality 1, a probable late middle Miocene age. He considered the flora from locality 3 to be no older than late Miocene and that from locality 4 to be either late middle or late Miocene.

The climate at the time of deposition of the Mashel Formation, according to Wolfe, was probably warmer than at present, and there may have been at least as much precipitation. The flora from locality 3 represents a climate somewhat cooler than that represented by the flora from locality 2.

## DISTRIBUTION

The Mashel Formation is well exposed in the bluffs along the lower Mashel River, the Ohop Valley, the south valley wall of Tanwax Creek, and along the Nisqually River valley from near LaGrande to the mouth of Tanwax Creek. The formation is typically exposed in sec. 20 and 29, T. 16 N., R. 4 E., along a logging road descending from Mashel Prairie to the Mashel River. Near McKenna, in the Nisqually River valley, the Mashel disappears under younger formations.

Similar unconsolidated Miocene deposits of fluvial and lacustrine origin containing pumice and volcanic ash are reported in the vicinity of Voight Creek southeast of Orting (Mullineaux and others, 1959) and in the valley walls of the Green River in King County, Wash. (Glover, 1941, p. 138).

## MODE OF DEPOSITION

The Mashel Formation probably was deposited in a piedmont environment during and after the uplift of the ancestral Cascade Range. The areal distribution of the lower part of the formation is not known, but the deposition of much of the gravel was probably limited to areas marginal to the upland, except along northwesttrending drainage lines that may have been in existence on the lowland before the uplift took place. Early in the development of the piedmont plain underlain by the Mashel, sand was deposited in the area midway between the upland and the center of the lowland; silt and clay were deposited near the center of the lowland. As the upland area was eroded and the lowland was aggraded, the area in which gravel was being deposited was reduced in size, and finegrained material was deposited closer to the upland front.

## CONTRIBUTIONS TO STRATIGRAPHY

#### PRECAMBRIAN AND LOWER CAMBRIAN FORMATIONS IN THE LAST CHANCE RANGE AREA, INYO COUNTY, CALIFORNIA

## By John H. Stewart

About 8,000 feet of conformable strata of Precambrian and Early Cambrian age are exposed in the Last Chance Range area (fig. 11). The strata in the northern part of the range and nearby areas are assigned to the Deep Spring Formation of Precambrian age, the Campito Formation of Precambrian and Early Cambrian age, the Poleta, Harkless, Saline Valley, Mule Spring Formations of Early Cambrian age, and the Emigrant(?) Formation of Middle and Late Cambrian age (fig. 12). The strata exposed farther south in the range are assigned to the Wood Canyon Formation of Precambrian and Early Cambrian age, the Zabriskie Quartzite of Early Cambrian age, and the Carrara Formation of Early and Middle Cambrian age (fig. 12).

The nomenclatural differences between the northern and southern parts of the area reflect a change in the lithologic character of the



FIGURE 11.—Map showing location of measured sections in the Last Chance Range area, Inyo County, Calif. Numbers and letters refer to units in figures 13 and 14.

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE



FIGURE 12.—Precambrian and Lower Cambrian formations identified in the Last Chance Range area, Inyo County, Calif.

strata. The strata in the northern part of the area contain more siltstone and limestone, and less quartzite, than comparable strata in the southern part of the area. The terms "Deep Spring" "Campito," "Poleta," "Harkless," "Saline Valley," and "Mule Spring," indicate the lithologic similarity of the strata in the northern part of the Last Chance Range area to the rocks in the stratigraphic sections to the west and north of the Last Chance Range in the Inyo and White Mountains in California (Nelson, 1962) and in Esmeralda County, Nev. (Albers and Stewart, 1962; McKee and Moiola, 1962) where these terms are used. The terms "Wood Canyon," "Zabriskie," and "Carrara" indicate the close lithologic similarity of the strata in the southern part of the Last Chance Range area to the rocks in the stratigraphic sections to the south and east of the Last Chance Range, in southern Nevada and adjoining parts of California (Nolan, 1929; Hazzard, 1937; Cornwall and Kleinhampl, 1961; Barnes and Palmer, 1961; Barnes and others, 1962) where these terms are generally used.

The author's study is based on reconnaissance geologic mapping of much of the Last Chance Range area and on measurement of detailed stratigraphic sections in the southern and northern parts of the area. Parts of the Last Chance Range area have been mapped by Whetten (1959) and McKee (1962), and their studies were helpful in the present geologic study in the area. McKee and Moiola (1962) mentioned briefly some of the exposures of Precambrian and Cambrian strata in the range directly northwest of the Last Chance Range. The present study is part of a continuing investigation of the correlation of the Precambrian and Lower Cambrian strata in the southern Great Basin.

Strata assigned to the Deep Spring Formation (fig. 13), including at the base some strata considered to be equivalent to the Reed Dolomite, are exposed in badly faulted sections in the northern part of the Last Chance Range. These are the oldest rocks exposed in the range, and the total section exposed amounts to about 1,560 feet. Unit 1A (fig. 13) consists of yellowish-brown fine- to medium-grained quartzite, minor siltstone, and, in the upper part, some limestone. Units 1B through 1L consist of gray and grayish-orange limestone, sandy limestone containing scattered fine to medium quartz grains, and minor amounts of dolomite, quartzite, and siltstone. Unit 1M consists of yellowish-gray quartzite and greenish-gray siltstone; unit 1N, of limestone and minor amounts of dolomite and siltstone; unit 1O, mostly of dark siltstone and quartzite similar to that in the Campito Formation; and unit 1P, of limestone and dolomite.

Correlation of units 1A to 1L with described units (Nelson, 1962) in the Inyo and White Mountains is uncertain. Some or all of units 1A to 1L must correlate with Nelson's (1962, p. 141; 1963) lower member of the Deep Spring Formation. Units 1A through 1L total 913 feet in the Last Chance Range as compared with only 496 feet (C. A. Nelson, written commun., 1963) for the lower member near Andrews Mountain (fig. 11) in the Inyo Mountains; thus the section exposed in the Last Chance Range probably extends down into rocks laterally equivalent to the Reed Dolomite, the formation underlying the Deep Spring Formation in the Inyo and White Mountains. This correlation is supported by the lithologic similarity of unit 1A with the Hines Tongue of the Reed Dolomite as described by Nelson (1962, p. 141). The dolomite that forms the top of the Reed Dolomite in the Inyo and White Mountains may be represented by limestone in the section in the Last Chance Range.

Units 1M and 1N are considered to correlate with Nelson's (1962, p. 141, 1963) middle member of the Deep Spring Formation, and units 1O and 1P with his upper member of that formation.

The Campito Formation consists of olive-gray, greenish-gray, darkgreenish-gray, and medium-gray very fine grained quartzite and siltstone and conformably overlies the Deep Spring Formation. The formation is extensively exposed in the northern part of the area. The thickness is estimated to be about 2,500 feet, although little significance can be placed on this figure because faulting in the area makes accurate measurements of the thickness impossible. Both the quartzitic Andrews Mountain Member and the overlying thinner and silty Montenegro Member (Nelson, 1962, p. 141) were recognized in the northern part of the area.

The Poleta Formation conformably overlies the Campito Formation and is divided into a lower member (unit 3A, fig. 14) of archeocyathidbearing limestone; a middle member (units 3B-3E, fig. 14) of greenishgray and olive-gray siltstone, minor limestone, and, in the upper part, some quartzite; and an upper member (unit 3F, fig. 14) of limestone. Division of the formation into three members follows the practice of McKee and Moiola (1962, p. 534-535) although Nelson (1962, p. 142) originally divided the formation into only two members; the upper two members of McKee and Moiola correspond to the upper member of Nelson. The middle member contains the trilobite Nevadella (identified by A. P. Palmer, written commun., 1963) in the basal 181 feet of unit 3C. Rare Scolithus (worm borings) occur in the upper part of the middle member. The Poleta Formation is about 850 feet thick.

The Harkless Formation is not completely exposed in any one section in the Last Chance Range area. The lower 580 feet (unit 4A. fig. 14) of the formation is exposed 1 mile northwest of Cucomungo Canvon (fig. 11). There the strata which lie conformably on the Poleta Formation consist of greenish-gray and grayish-olive siltstone containing layers of quartzite in a few places. Trilobites that probably can be designated Fremontia(?) and Paedeumias(?) (identified by C. A. Nelson, written commun., 1964) occur in this section about 280 feet above the base of the Harkless Formation. Along Cucomungo Canyon, about 1,300 feet of the Harkless Formation are exposed conformably below the Saline Valley Formation. Parts of this section are highly faulted. The strata consist mostly of gravishpurple, pale-red, yellowish-gray, and greenish-gray fine- to mediumgrained quartzite containing numerous siltstone layers from onequarter inch to several feet thick. Unit 4C is a particularly conspicuous siltstone unit. Unit 4D is composed of dolomite and lime-

#### Thick Unit Description ness (feet) 2A Olive-gray to medium-dark-gray siltstone 1P 61 Yellowish-gray and light-gray dolomite in basal 35 ft; overlain by very pale orange limestone JPPER MEMBER FEET 0 10 223 Greenish-gray, olive-gray, and medium-gray siltstone, yellowish-gray, greenish-gray, and medium-gray very fine to fine-grained quartzite; silty limestone, siltstone and dolomite from 69 to 85 ft above base - 100 FORMATION 1N 55 Gray limestone, very pale orange dolomite in top 6 ft, some siltstone in middle SPRING 200 DEEP WIDDLE MEMBER 300 1M 310 Yellowish-gray very fine to fine-grained quartzite and greenish-gray siltstone. Amount of quartzite decreases upward - 400 500

UNITS 1A-1N MEASURED IN SEC. 31 AND 32 (UNSURVEYE D), T. 7 S., R. 39 E. UNITS 10-2A MEASURED IN SEC. 24, T. 7 S., R. 38 E.

FIGURE 13.-Columnar section of Deep Spring Formation and related

#### CHANGES IN STRATIGRAPHIC NOMENCLATURE



strata in northern Last Chance Range area, Inyo County, Calif.
NORTHERN PART OF THE LAST CHANCE RANGE AREA, UNITS 2B-4A MEASURED IN SEC. 15, T. 7 S., R. 38 E., UNITS 4B-6A MEASURED IN SECS. 14 AND 23, T. 7 S., R. 38 E.



FIGURE 14.—Columnar sections and correlations of Lower Cambrian

SOUTHERN PART OF THE LAST CHANCE RANGE AREA, UNITS 7A-8A MEASURED IN SECS. 7 AND 8 (UNSURVEYED), T. 10 S., R. 40 E., UNITS 9A-10A MEASURED IN SEC. 29, 30, 31, AND 32 (UNSURVEYED), T. 9 S., R. 40 E.

	NZA G		Unit	Thickne (feet)	ess Description
	BONAI		10A	unmeas- ured	Gray limestone
	Ľ		9N	115	Gray and yellowish-brown limestone and minor silty lime- stone and limy siltstone
		室室	9M	176	Gray limestone; conspicuous white band at base
			9L	163	Yellowish-gray and medium-gray limestone, minor silty limestone and siltstone
	TION		9K	250	Gray limestone; white band at top, locally oolitic. Common Girvanella in lower half
	RMA 0 FI		16	23	Gray limy siltstone
	UT 164		91	44	Gray limestone
	CARR		9н	240	Greenish-gray phyllitic siltstone, scarce thin yellowish-brown or gray limestone layers
	72		9G	200	Gray limestone. Common Girvanella throughout; forms prominent cliff
	~~ `		]  9F	82	Gray limestone; mostly phyllitic siltstone in lower half
			9E	112	Gray limestone and minor phyllitic siltstone in lower part. Some Girvanella
		^ · · · ^	9D	71	Gray limestone with minor siltstone in lower part. Girvanella at top
			90	47	Greenish-gray phyllitic siltstone, minor limestone
	IZITE	· · · · · ·	9B	32	Gray limestone and dolomitic limestone; sandy limestone at top
	QUAR1	···· .	9A	88	Yellowish-brown very fine to fine-grained quartzite, minor phyllitic siltstone
	ZABRISKIE 1360±	×	8A	1360±	Medium - dark-gray to grayish-purple fine- to medium- grained rarely coarse-grained vitreous quartzite; lami- nated and minor small-scale cross-strata. Minor siltstone and conspicuous cross-strata in lower 88 ft. Cut by fault about 950 ft above base
		· · · · ·	7F	124	Dark-greenish-gray siltstone. Quartzite layers minor in lower part and dominant at top. Contains Paedeumios?
	l		/7E	28	Gray limestone
			70	194	Yellow-brown and yellow-gray very fine grained quartzite. Abundant Scolithus tubes
	ON FORMATION 0 FEET		70	424	Greenish-gray to medium-gray phyllitic siltstone, and minor yellowish-brown and medium-gray very fine sandstone and yellowish-brown and medium-gray limestone. Lime- stone contains pelmatozoa debris and, in basal 140 ft, archeocyathids. A few Scolithus tubes. Nevodella? occurs in basal 10 ft
·	0 CAN	<u>E</u>	7B	123	Very pale orange and yellowish gray limestone. Archeo- cyathids and pelmatozoa debris
	DOW		7A	407	Yellowish-brown very fine grained quartzite, minor greenish- gray and gray phyllitic siltstone. Two thin limestone beds in upper half, the lower limestone bed contains archeocyathids and pelmatozoa debris

strata in the Last Chance Range area, Inyo County, Calif.

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stone and contains *Salterella* in places; the unit is considered to correlate with a fairly persistent limestone found at the top of the Harkless Formation directly below the Saline Valley Formation in the Waucoba Spring area of the Inyo Mountains.

Unfortunately the strata of the Harkless Formation exposed 1 mile northwest of Cucomungo Canyon cannot be matched with any of the strata of that formation exposed along Cucomungo Canyon, and a small amount of strata may be missing between the two stratigraphic sections. The two sections may possibly overlap, however, and the difference in lithologic character may be due to abrupt facies changes coupled with thrust faulting; such faulting would bring unrelated facies into close juxtaposition. Thrust faults have been recognized in the vicinity of Cucomungo Canyon and one, at least, may occur between the section in Cucomungo and the section a mile to the northwest; critical outcrops are, however, covered by younger strata.

The Saline Valley Formation as exposed in Cucomungo Canyon is about 750 feet thick and consists, in the lower part, of pinkish-gray and grayish-red-purple quartzite (unit 4E) and, in the upper part (units 4F-4I), of gray limestone and greenish-gray siltstone, a few layers of sandy limestone near the base, and a layer of quartzite near the top. The quartzite in the lower part of the Saline Valley Formation is very similar to that in the Harkless Formation, and the two formations apparently can be separated only if the intervening carbonate layer (unit 4D) is between them. Salterella and the trilobites Wanneria sp. and Paedeumias? sp. have been collected and identified by A. R. Palmer (written commun., 1963) from the lower part of the u it 4F of the Saline Valley Formation in Cucomungo Canyon.

The Mule Spring Limestone (units 5A-5C, fig. 14) conformably overlies the Saline Valley Formation. It is about 510 feet thick and consists of gray limestone commonly containing algal structures (*Girvanella*). The trilobites *Bonnia* sp., *Peachella iddingsi* (Walcott), *Paedeumias* sp., and *Bristolia* n. sp. have been collected and identified by A. R. Palmer (written commun., 1963) from the upper part of unit 5B of the Mule Spring Limestone in Cucomungo Canyon.

The Mule Spring Limestone is overlain by, and is partly or entirely in thrust contact with a thick and badly faulted sequence of generally thin-bedded limestone and cherty limestone containing some dolomite and siltstone units. A siltstone unit may occur at the base of this sequence directly above the Mule Spring Limestone (fig. 14), although the outcrop showing this relationship is small and faulted. This thick sequence above the Mule Spring Limestone has not been studied critically but is probably assignable to the Emigrant Formation of

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Middle and Late Cambrian age, a formation overlying the Mule Spring Limestone in Esmeralda County, Nev. (Albers and Stewart, 1962, p. D27; McKee and Moiola, 1962, p. 536-537). Trilobites collected in this limestone sequence near Cucomungo Canyon by A. R. Palmer and the author have been identified as Middle Cambrian in age (A. R. Palmer, written commun., 1963).

In the southern Last Chance Range, the upper 1,300 feet of the Wood Canyon Formation (units 7A-7F, fig. 14) is exposed. The exposed section consists of very fine grained quartzite, greenish-gray and gray siltstone, and yellowish-brown and gray limestone. Archeocyathids and pelmatozoan debris occur in a thin limestone bed 177 feet below the top of unit 7A, throughout much of unit 7B, and in the basal part of unit 7C. One specimen of a poorly preserved trilobite, identified as possibly *Nevadella* (A. R. Palmer, written commun., 1963), was found in the basal 10 feet of unit 7C. Trilobite remains, possibly in part *Paedeumias* (A. R. Palmer, written commun., 1963), are locally common in unit 7F. *Scolithus* (worm borings) occur in the upper half of unit 7C and are abundant in unit 7D. Units 7B through 7E are correlative with the Poleta Formation of the northern Last Chance Range. This correlation is substantiated by similar thicknesses and sequences of units, as well as by similar fossil occurrences.

The Zabriskie Quartzite (unit 8A, fig. 14) in the southern Last Chance Range consists of medium-dark-gray to grayish-purple fine-to medium-grained vitreous quartzite that is coarse grained in a few places. The quartzite is mostly laminated but locally contains smallscale cross strata. The lower 88 feet contains a minor amount of siltstone, and the quartzite is conspicuously cross stratified. The measured thickness of the Zabriskie Quartzite is about 1,360 feet, although a prominent fault cuts the unit about 950 feet above its base and makes the exact thickness uncertain.

The Zabriskie Quartzite correlates clearly with the quartzite in the Harkless Formation and with that in the lower part of the Saline Valley Formation of the northern part of the Last Chance Range area. Much of the quartzite in the northern part of the area, however, contains thin layers of siltstone, which are not present in the Zabriskie Quartzite. These siltstone layers represent a change to a more silty facies within this part of the sequence to the north. This facies change is most noticeable in the lower part of the Harkless Formation. The basal 580 feet (unit 4A) of the Harkless Formation 1 mile northwest of Cucomungo Canyon is dominantly siltstone and contains only a few thin layers of quartzite. These quartzite layers are tongues of the Zabriskie Quartizite, and the siltstone, except in the basal 100-200 feet, is probably a lateral equivalent of the Zabriskie Quartzite, but of a different facies. The Carrara Formation (units 9A-9N, fig. 14) is 1,640 feet thick in a measured section in the southern part of the Last Chance Range area. It consists of gray limestone, commonly containing *Girvanella*, greenish-gray phyllitic siltstone, and yellowish-brown silty limestone. Units 9G and 9K form persistent cliffs, whose tops mark vertical lithologic changes that divide the Carrara Formation into three mappable units or members. Unit 9A is intermediate in lithologic type from the Zabriske Quartzite below to the Carrara Formation above. It is included with the Carrara Formation because a similar, and probably correlative unit, has been included in the type Carrara Formation (Cornwall and Kleinhampl, 1961) on Bare Mountain. The Carrara Formation is overlain by the Bonanza King Formation, a thick formation of Middle and Late Cambrian age composed of dolomite and limestone.

The Carrara Formation of the southern part of the Last Chance Range area contains correlatives of the upper part of the Saline Valley Formation, all the Mule Spring Limestone, and the lower part of the Emigrant(?) Formation of the northern part of the Last Chance Range area (fig. 14). Some of the quartzite in unit 9A of the Carrara Formation in the southern part of the Last Chance Range area may correlate with the quartzite at the top of unit 4H of the Saline Valley Formation in Cucomungo Canyon in the northern part of the area. Part of unit 9E and all of units 9F and 9G of the Carrara Formation probably correlate with the Mule Spring Limestone in Cucomungo Canyon. The remainder of the Carrara Formation is correlative with the lower part of the Emigrant(?) Formation, of the northern part of the area.

## MIOCENE AND PLIOCENE ROCKS OF CENTRAL WYOMING

## By NORMAN M. DENSON

A widespread succession of light-gray fossiliferous tuffaceous siltstone and fine-grained sandstone extends from the vicinity of the Alcova Dam westward along the north side of the Seminoe, Ferris, and Green Mountains to the southern terminus of the Wind River Rangea distance of about 85 miles. Love (1961) applied the term Split Rock Formation to this succession and some associated rocks because fossiliferous Miocene rocks in eastern Wyoming are difficult to correlate with named formations and groups in Nebraska and because the lower and middle Miocene rocks of the Granite Mountains area can not be mapped continuously into the type area of the Arikaree in northwestern Nebraska.

Recent studies and regional mapping by the author from northwestern Nebraska into central Wyoming and local studies by Harshman (1964) in the Shirley Basin, Stephens (1964) at Crooks Gap, Rich (1962) at Clarkson Hill, and Zeller, Soister, and Hyden (1956) in the Gas Hills indicate that most of the rocks originally assigned to the Split Rock Formation in the Granite Mountains area are laterally equivalent and remarkably similar lithologically and chemically to the rocks in eastern Wyoming and northwestern Nebraska described and first assigned by Darton (1899) to the Arikaree Formation. Furthermore, the upper part of the Split Rock Formation at its type locality includes rocks lithologically very similar to the lower part of the Ogallala Formation and, for this reason, is here assigned to that part of the Ogallala. In other areas the lower part of the Split Rock includes the upper part of the White River Formation (Oligocene). A correlation chart of Miocene and Pliocene rocks of central Wyoming follows:

	L	ove (1961)	This report			
Series	Age	Formation	Subdivision	Formation	Age	Series
Pliocene	Early or middle Pliocene	Moonstone			Pliocene and late Miocene	Pliocene and Miocene
Miocene	Middle Miocene	Split Rock	Upper porous	Ogallala		
			Local fauna1		Middle and early Miocene	Miocene
			sandstone			
			sequence			
			Silty sandstone sequence	Arikaree		
			Clayey sandstone sequence			
			Vertebrate for	sils <sup>2</sup>		
	Early Miocene		Lower porous sandstone sequence	White River (upper part)	Late Oligocene	Oligocene

<sup>1</sup>"Split Rock local fauna " of middle Miocene age (Love, 1961, p. 19)

<sup>2</sup>Merycoides cursor Douglass of early Miocene (Gering) age (Rich, 1962, p. 506)

Rocks of widely different lithologies and ages were originally assigned to the Split Rock Formation in two measured sections at the type locality. They include middle Miocene rocks (sec. 36, T. 29 N., R. 90 W.), assigned in this report to the Arikaree, as well as Pliocene rocks (secs. 25 and 36, T. 29 N., R. 89 W.) that are here referred to as the Ogallala Formation. The lower of these two distinct stratigraphic units contains many middle Miocene vertebrate fossils ("Split Rock local fauna" of Love and others) and was assigned by Love (1961, p. 19) to the upper porous sandstone sequence of the Split Rock Formation. Rocks below the local fauna are composed predominantly of wind-blown buff and tan fine- to medium-grained poorly bedded sandstone having abundant tiny rounded grains of bluish-gray magnetite. Lateral persistence in lithology and the general absence of coarse detritus and locally derived debris from the surrounding highland are outstanding characteristics of rocks below the local fauna at the type locality of the Split Rock and elsewhere over the whole region.

Conformably overlying the middle Miocene windblown sandstone at the type locality are Pliocene rocks which Love (1961, p. 17, 18) also assigned to the Split Rock. The rocks comprising this stratigraphic unit consist mostly of thin beds of relatively pure white pumicite, pumiceous limestone, sandstone, claystone, and tuff which grade mountainward into fanlike deposits of coarse-grained sandstone, conglomerate, and gravel. The basal contact becomes an unconformity near the mountains. Most of the rocks contain a preponderance of volcanic ash. These volcanic-rich rocks have vielded many species of diatoms and spores (Love, 1961, p. 17, 20, 21) and have been traced eastward to the vicinity of the Pathfinder Reservoir where they have yielded vertebrate fossils determined by P. O. Mc-Grew to be of Pliocene age (J. D. Love, oral commun., 1964). These rocks range from a few feet to at least 300 feet in thickness and are assigned on the basis of lithologic similarity and age to the Ogallala Formation; some vertebrate paleontologists (Wood and others, 1941, p. 27) consider the oldest of several Ogallala local faunas to be of latest Miocene age, whereas others (for example, Schultz and Falkenbach, 1949, p. 80, 83) believe that all the Ogallala is Pliocene.

The Oligocene-Miocene (White River-Arikaree) contact along the eastern and northeastern margins of the Granite Mountains area was drawn by Rich (1962, pl. 7, p. 503-506) at the base of a persistent and wide-spread conglomerate that is 150-600 feet thick. Love (1961, p. 9-12) referred to it as the lower porous sandstone sequence. Areal mapping by the writer from the Granite Mountains eastward into the Shirley Basin now indicates that this widespread conglomeratic succession is a lateral equivalent of a conglomerate, 300-350 feet thick, that some previous workers (Love and others, 1955) assigned to the Miocene and Pliocene, but that Harshman (1964) correctly mapped as an upper coarse-grained member of the Oligocene White

River Formation. At several localities in the Shirley Basin, this conglomeratic sequence has vielded Oligocene (Brule) vertebrate fossils from its lower and middle parts (Harshman, oral commun. 1965; Whitmore, F. C., Jr., and Lewis, G. E., written commun., 1962, 1963). Rocks identical to the Arikaree in the Clarkson Hill area overlie the conglomerate and were mapped by Rich (1962, pl. 7) as Miocene. In the eastern and northeastern parts of the Granite Mountains area, these thick conglomerates directly overlie fossiliferous lower Oligocene (Chadron) rocks and are assigned herein for the first time to the upper part of the White River. At many places in the Shirley Basin, light-gray calcareous sandstone at the base of the Arikaree directly overlies Harshman's upper coarse-grained member of the White River Formation. In the eastern part of the Granite Mountains and in the Shirley Basin, conglomerates occur at the base of the Arikaree but are not widespread; these conglomerates are largely channellike lenticular deposits generally less than 60 feet thick. At many localities the conglomerates are interbedded with fine- to medium-grained light-gray calcareous sandstone lithologically similar to that in the lower part of the Arikaree, from which early Miocene fossils have been reported in the Granite Mountains (Rich, 1962, p. 506).

The following analyses are presented to show the striking similarity in chemical composition of the Miocene rocks on the west and east sides of the Laramie Range.

Rapid rock analyses of very fine grained tuffaceous sandstone of Miocene age

	Lower and stone in o age thick	middle Mic central Wyon ness 1,000 ft)	ocene sand- ming (aver-	Arikaree Formation in southeast Wyoming and northwest Nebras- ka (average thickness 700 ft)		
	Range (5	analyses)	Average	Range (10 analyses)		A verage
· · · ·	From-	То		From-	То—	
$\begin{array}{c} {\rm SiO}_2 \\ {\rm Al}_2 {\rm O}_3 \\ {\rm K}_2 {\rm O} \\ {\rm Fe}_2 {\rm O}_3 \\ {\rm CaO} \\ {\rm CaO} \\ {\rm Ma}_2 {\rm O} \\ {\rm MgO} \\ {\rm TiO}_2 \\ {\rm P}_2 {\rm O}_5 \\ {\rm MnO} \\ {\rm WeO} \end{array}$	65. 1 10. 7 2. 0 2. 2 1. 6 1. 3 1. 3 . 21	74. 3 12. 5 4. 2 3. 1 2. 6 2. 0 2. 6 . 32	69. 7 11. 7 3. 1 2. 5 2. 3 1. 6 1. 9 . 26	$\begin{array}{c} 66.8\\ 10.5\\ 2.3\\ 2.2\\ 1.2\\ 1.4\\ .9\\ .02\\ .04\\ \end{array}$	75. 8 13. 2 3. 8 3. 4 3. 1 2. 4 1. 7 . 46 . 12 . 08	$\begin{array}{c} 71.9\\ 11.9\\ 2.9\\ 2.7\\ 2.3\\ 2.0\\ 1.3\\ .07\\ .07\\ .06\\ \end{array}$
H <sub>2</sub> U	5.2	11.0	7.3	2.7	7.3	4.50
			100. 99			100.00

[Method described by Shapiro and Brannock (1956)]

The Moonstone Formation defined by Love (1961, p. 25-35) for some lower or middle Pliocene rocks in the Granite Mountains area is a sufficiently distinctive lithologic unit to warrant a separate rockstratigraphic designation; however, the white uranium-bearing shale, green tuff, bedded chalcedony, and finely-laminated tuffaceous arenites, algal reefs, and lenticular beds of conglomerate to which Love applied the term Moonstone are present only locally in the Granite Mountains. These rocks occur principally in an area of about 50 square miles in and adjacent to T. 30 N., R. 89 W. Elsewhere in the region a succession of rocks lithologically similar to those assigned to the Moonstone is not present. The term Moonstone, therefore, has limited use as a rock-stratigraphic designation and is referred to in this report as part of the Ogallala.

In summary, the lower and middle Miocene rocks of the Granite Mountains area of central Wyoming are composed largely of windblown fine- to medium-grained tuffaceous sandstone having thin and relatively unimportant interbeds of limestone, tuff, and conglomerate. These rocks average about 1,000 feet in thickness and constitute an easily recognized lithogenetic unit. This widespread unit has been mapped discontinuously from the vicinity of Oregon Buttes along the southwest flank of the Wind River Mountains through central and southeastern Wyoming into northwestern Nebraska. Because this unit is strikingly similar lithologically to the Arikaree Formation as defined by Darton and because the rocks that constitute it are unconformably overlain and underlain at most places by rocks that can properly be assigned to the Ogallala and White River Formations, respectively, the term Arikaree is applied here in central Wyoming with the same meaning given it by Darton in 1899. Since 1899, Arikaree has been used for Miocene (unrestricted) rocks in Wyoming by the U.S. Geological Survey, although the stratigraphic and chronologic range elsewhere has been restricted. As indicated in this report, the term "Split Rock Formation" is not regionally useful or meaningful and, therefore, is abandoned.

The Moonstone, originally defined to include some lower or middle Pliocene rocks in the Granite Mountains area, is referred to in this report as part of the Ogallala because of its limited areal extent.

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