


**THE 1992 ERUPTIONS OF
CRATER PEAK VENT,
MOUNT SPURR VOLCANO,
ALASKA**



U.S. GEOLOGICAL SURVEY BULLETIN 2139

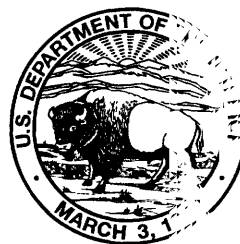


Cover—Ash cloud rising from Crater Peak vent of Mount Spurr volcano, Alaska, during the eruption on August 18, 1992. Crater Peak is not visible in image. A pyroclastic flow that descended the volcano's flank produced the smaller steam and ash cloud shown at lower right of large cloud. Photograph taken by Robert G. McGimsey.

THE 1992 ERUPTIONS OF CRATER PEAK VENT, MOUNT SPURR VOLCANO, ALASKA

Terry E.C. Keith, Editor

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The 1992 Eruptions of Crater Peak Vent, Mount Spurr Volcano, Alaska: Chronology and Summary

By John C. Eichelberger, Terry E.C. Keith, Thomas P. Miller, and Christopher J. Nye

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ABSTRACT

Following 39 years of inactivity, Crater Peak vent on the south flank of Mount Spurr volcano burst into eruption at 7:04 a.m. Alaska daylight time (ADT) on June 27, 1992. This and subsequent eruptions on August 18 and September 16–17 were subplinian, had a volcano explosivity index (VEI) of 3, and lasted about 4 hours. The June, August, and September eruptions released 44, 52, and 56 million m³ (12, 14, and 15 million m³ dense rock equivalent or DRE) of andesitic (57 weight percent SiO₂) tephra, respectively and drove tephra columns 14,000 to 15,000 m above sea level. The August eruption had the most far-reaching

effects—it deposited 3 mm of ash on Anchorage, 125 km east of the volcano. Proximal pyroclastic flows accompanied the August and September eruptions. In all eruptions, lahars and debris flows descended Crater Peak's south flank, and some reached the Chakachatna River.

Real-time seismic monitoring tracked the 10-month crescendo of precursory earthquakes and allowed timely warning of the increasing unrest to State and Federal government officials, the military, air carriers, and local citizens. This monitoring was augmented with other types of observations and provided the basis for accurate eruption advisories that minimized economic losses. In particular, because of an efficient ash-cloud warning system in Alaska and new awareness of the problem of ash clouds within the aviation community, no jets were damaged. Unavoidable losses of \$5 to 8 million were sustained from the August ashfall on Anchorage and the Matanuska-Susitna Valley, both of which make up the State of Alaska's center of population and economic activity. Additional but unevaluated costs were incurred from flight delays in large North American airports to the south and east as ash clouds of the August and September eruptions passed overhead.

INTRODUCTION

The Crater Peak vent of Mount Spurr volcano, located 125 km west of Anchorage (figs. 1, 2) and last active in 1953 (Juehle and Coulter, 1955; Wilcox, 1959), erupted three times in 1992 following 10 months of heightened seismicity. The first eruption occurred on June 27, 1992, and was followed by similar events on August 18 and September 16–17, 1992. This paper provides an overview of these eruptions and the Alaska Volcano Observatory's response to them.

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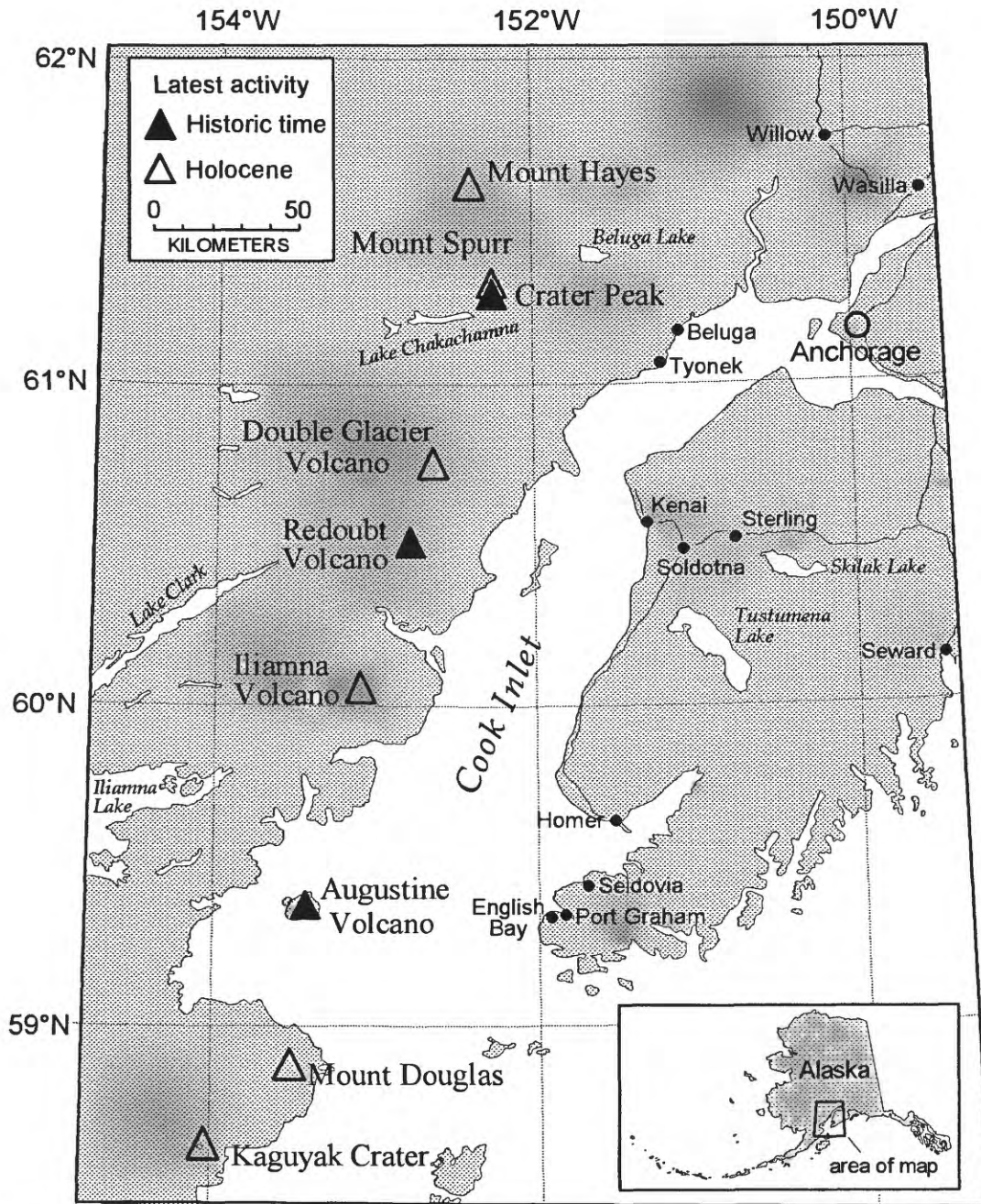


Figure 1. Index map showing volcanoes and population centers of the Cook Inlet region, Alaska.

GEOLOGIC SETTING AND HISTORY OF MOUNT SPURR VOLCANO

Mount Spurr sits 100 km above the Wadati-Benioff zone and 500 km inboard of the Aleutian Trench (Jacob and others, 1977) on a basement of mainly intrusive rocks (Nye and Turner, 1990). Its large stratovolcano edifice, built of two-pyroxene andesite (58–60 weight percent SiO_2), began to form before 255,000 yr ago (Nye and Turner, 1990). Pyroclastic rocks dominate the lower part of the volcano, and lava flows dominate the upper part. Within the past 58,000 yr

but before Holocene time, the volcano collapsed southward in a manner similar to Bezymianny Volcano in Kamchatka, Russia (Gorshkov, 1959). The debris avalanche generated at Mount Spurr had a minimum runout of 25 km by 6 km and left the volcano with a horse-shoe-shaped caldera rim. Collapse was followed immediately by eruption of more silicic (60–63 weight percent SiO_2) andesitic ash-flow tuffs. A large dome complex, which now forms Mount Spurr's 3,374-m summit, grew in the center of the caldera probably no later than 5,000 yr ago (Riehle, 1985). The complex consists of two-pyroxene silicic andesite, chemi-

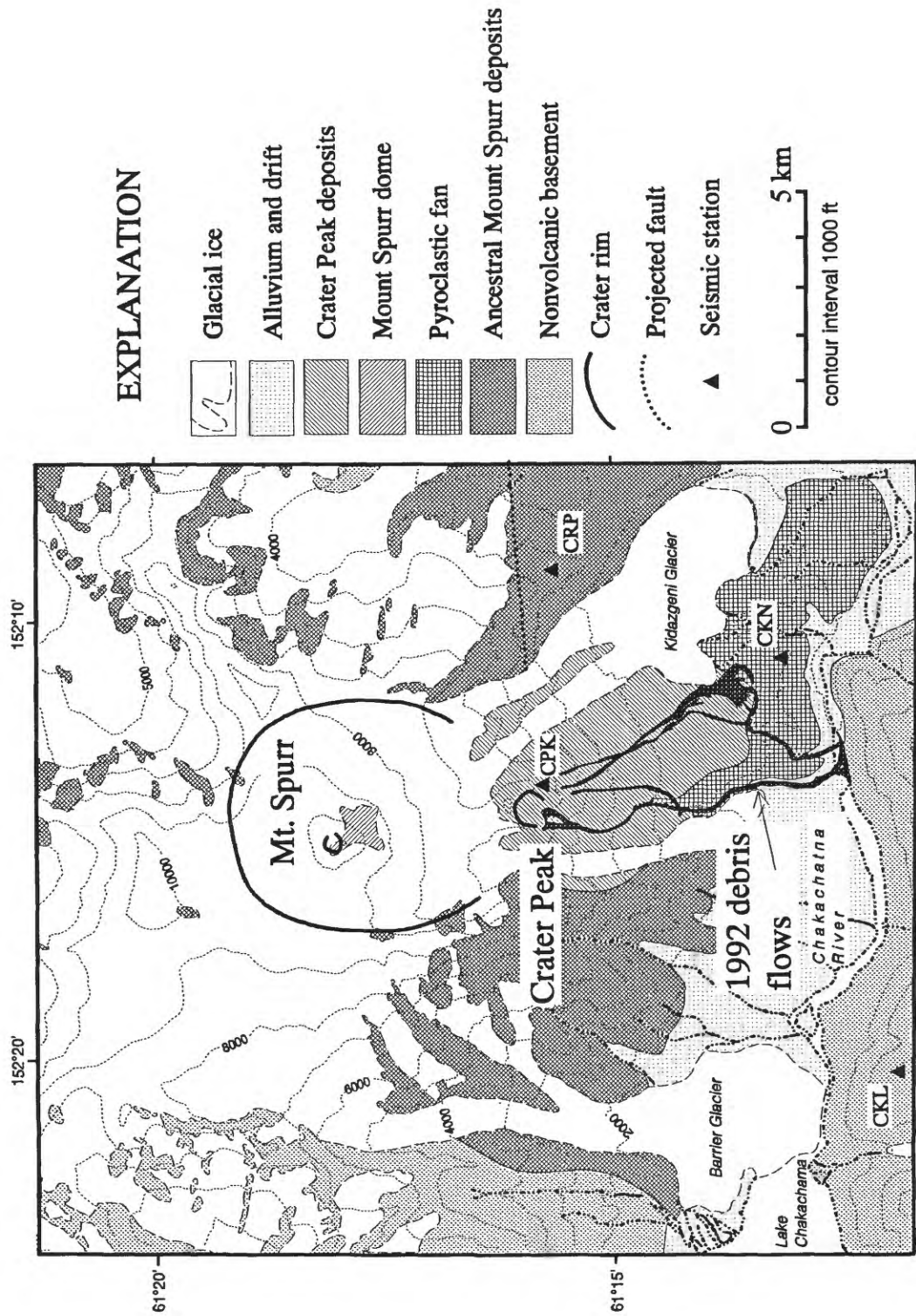


Figure 2. Generalized geologic map of Mount Spurr area, southern Alaska.

cally similar to the ash-flow tuffs (Nye and Turner, 1990). A basaltic-andesite satellitic cone developed at the same time in the caldera breach. The present Crater Peak represents a second period of satellite cone-building at the same location; it reaches an elevation of 2,309 m, 3.5 km south of Mount Spurr's summit. Crater Peak consists of lavas and pyroclastic flows of mafic andesite (53–57 weight percent SiO_2 ; Nye and Turner, 1990). The 40 tephra units correlated to this vent range in age from 5,000 yr to present (Riehle, 1985). Eruption products that are apparent mixtures of Crater Peak and summit dome complex magmas indicate that the two systems were partly coeval (Nye and Turner, 1990).

Prior to 1992, the only witnessed eruption of Mount Spurr occurred on July 9, 1953. This eruption also issued from the Crater Peak vent and consisted of two main explosive pulses, followed by steam and minor ash emission. The eruption cloud rose to over 20,000 m above sea level and deposited 3 to 6 mm of andesitic (55 weight percent SiO_2) ash on Anchorage (Juehle and Coulter, 1955; Wilcox, 1959). A debris flow dammed the Chakachatna River downslope from Crater Peak, and the dam created an 8-km-long lake. The dam was overtopped and breached within a week.

ROLE OF THE ALASKA VOLCANO OBSERVATORY

The Alaska Volcano Observatory (AVO) was established in 1988, following the 1986 eruption of Augustine Volcano; the observatory conducted its first full eruption response with the Redoubt Volcano episode of 1989–90 (Miller and Chouet, 1994). AVO is a cooperative program of the U.S. Geological Survey (USGS), University of Alaska Fairbanks Geophysical Institute (UAFGI), and Alaska Division of Geological and Geophysical Surveys (ADGGS). AVO is supported by both the Department of Interior through the USGS Volcano Hazards and Geothermal Studies Program and by the State of Alaska. AVO's primary mission is to communicate timely warnings of unrest and potential eruption of Alaska's volcanoes and to investigate the fundamental processes of hazardous volcanism, with a view toward improved eruption warnings and hazard assessments.

Since World War II, Anchorage has become the population and economic center of Alaska; over half the population of the State is currently concentrated in the Anchorage area. Three of the Cook Inlet volcanoes—Augustine Volcano, Redoubt Volcano, and Mount Spurr—are frequently active. Within the past decade, their eruptions have produced ash clouds that were hazardous to aircraft as well as to the general

populace. Accordingly, the Cook Inlet region presents the greatest array of volcano hazards in Alaska and is the focus of AVO's efforts.

AVO distributes updates (fig. 3) each Friday to summarize the weekly status of the monitored Cook Inlet volcanoes and any reported activity of unmonitored volcanoes on the Alaska Peninsula and Aleutian Islands. During periods of volcanic unrest and eruption crises, additional updates are issued when significant changes in activity should be made known to the public (Brantley, 1990). Updates were issued frequently to a prioritized list of well over 100 recipients during the Crater Peak eruptions. AVO at Fairbanks faxes to the local Federal Aviation Administration (FAA), the local National Weather Service (NWS), Alaska Department of Emergency Services (ADES), local military bases, Governor's and State offices, television and radio stations, and airlines. AVO at Anchorage faxes to the local FAA, the local NWS, local military bases, USGS offices, other federal agencies, television and radio stations, news wire services, and airlines. Updates are also distributed by electronic mail to the volcano information networks.

A level of concern color code was established during the Redoubt Volcano eruptions to quickly and simply convey AVO's evaluation of eruption potential and eruption severity (Brantley, 1990; Miller and Chouet, 1994). This code was modified slightly for use with the Crater Peak eruptions (fig. 4). When a volcanic crisis begins, telephone calls are made to Federal, State, and local agencies with critical public safety responsibilities to give the level of concern color code as the update is being prepared. A calldown procedure, likewise established for the Redoubt Volcano eruptions (Brantley, 1990; Miller and Chouet, 1994), was used for the Crater Peak eruptions. Responsibility for calls are shared between AVO at Anchorage and AVO at Fairbanks, so all concerned agencies are contacted within minutes.

1992 ERUPTIONS—MONITORING AND RESPONSE CHRONOLOGY

PRE-ERUPTION PERIOD

Early indication of unrest at Mount Spurr was a conspicuous swarm of volcano-tectonic (VT) earthquakes during the last half of August 1991 (table 1; Power and others, this volume). Unlike swarms that had occurred in 1982 and 1989, this one was directly under Crater Peak within a volume that had been aseismic for at least a decade (Jolly and others, 1994), and it was followed by caldera-wide seismicity. By February of 1992, seismic activity had increased to such a level that discussions of Mount Spurr's unrest were included in AVO's weekly updates. Through April

ALASKA VOLCANO OBSERVATORY

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MT. SPURR UPDATE

Thursday, September 17, 1992 12:30 AM ADT

CURRENT LEVEL OF CONCERN COLOR CODE IS **RED**

LAST LEVEL OF CONCERN WAS YELLOW
For definitions, see update of 8/24/92

At approximately 00:04 ADT, AVO seismometers indicated a second eruption at Mt. Spurr which is ongoing at this time. An ash plume reaching an estimated altitude of at least 35,000 feet has been confirmed by NWS radar and pilot observations.

Based on forecast winds, the plume should travel in an easterly direction from Mount Spurr and could effect the Mat-Su and greater Anchorage area region with the southern edge of the fallout impacting the Anchorage area within the next two hours.

The Alaska Volcano Observatory is watching the situation closely and will remain on 24-hour staffing until further notice.

PLEASE CONTACT AVO IF YOU HAVE ANY QUESTIONS OR COMMENTS.

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Figure 3. Example of update issued by Alaska Volcano Observatory during the eruption of Mount Spurr volcano on September 17, 1992.

LEVEL OF CONCERN COLOR CODE

To more concisely describe our level of concern about possible eruptive activity at Mt. Spurr, the Alaska Volcano Observatory has developed a color-coded classification system. The definitions of the various colors have been modified slightly since the 1989-90 eruption of Redoubt Volcano. General descriptions of the type of activity associated with each of the four color codes are included below. The various colors indicating our level of concern about potential eruptive behavior are as follows:

GREEN : Volcano is in its normal "dormant" state.

YELLOW: Volcano is restless.

Seismic activity is elevated. Potential for eruptive activity is increased. A plume of gas and steam may rise several thousand feet above the volcano which may contain minor amounts of ash.

ORANGE: Small ash eruption expected or confirmed. Plume(s) not likely to rise above 25,000 feet above sea level.

Seismic disturbance recorded on local seismic stations, but not recorded at more distant locations.

RED : Large ash eruptions expected or confirmed. Plume likely to rise above 25,000 feet above sea level.

Strong seismic signal recorded on all local and commonly on more distant stations.

Figure 4. Alaska Volcano Observatory's level of concern color code, as used for the 1992 eruptions of Crater Peak, Mount Spurr, Alaska. This code was distributed to government agencies, the media, airlines, and the public with the announcement of level Yellow for Mount Spurr on June 26, the day prior to the first eruption.

Table 1. Chronology of major events and Alaska Volcano Observatory notifications pertaining to the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska.

[Compiled by C.A. Neal, T. Mattox, I. Ellersieck, and T. Keith; see also Power and others and McNutt and others, this volume; Universal Time (UT) = Alaska daylight time (ADT) + 8 hours, Alaska standard time (AST) + 9 hours; * = next day in UT; altitudes are above sea level; PIREP, pilot report. Beginning and end of each eruption are in bold type.]

DATE	TIME		EVENT	COLOR CODE	EVENT DESCRIPTION OR NOTIFICATION SUMMARY ISSUED
	Local	Universal			
8/15–30/91	- -	- -	Seismicity	- -	Conspicuous swarm of shallow volcano-tectonic earthquakes beneath Crater Peak accompanied by increased seismicity in a previously aseismic zone between 10 and 40 km depth.
8/30/91	- -	- -	Update	- -	First announcement of seismic swarm of past two weeks at Mount Spurr; similar to swarms in 1982 and 1989; no further notifications of increased seismicity through end of year.
2/7/92	- -	- -	Update	- -	Increased seismicity beneath Mount Spurr volcano.
3/13/92	- -	- -	Update	- -	Began weekly reporting of increased seismicity at Mount Spurr; noted seismicity had been above background since mid-August 1991.
6/5/92	- -	- -	Seismicity	- -	Highest number of earthquakes per day (28) recorded to date.
6/6/92	- -	- -	Seismicity	- -	First small tremor bursts detected beneath Crater Peak.
6/8/92	- -	- -	Update	- -	Discussion of increased seismicity at Mount Spurr volcano localized beneath Crater Peak vent; highest seismicity in 10 years; possibility of a coming eruption.
6/26/92	10:30 a.m.	18:30	Update	- -	Increasing number and duration, as long as 10 minutes, of tremor bursts beneath Crater Peak; two tremor bursts of 154 and 142 minutes, respectively, on June 24 and 25.
	12:04 p.m.	20:04	Seismicity	- -	Continuous tremor began.
	3:00 p.m.	23:00	Seismicity	- -	Decision made to announce level of concern color code Yellow.
	4:30 p.m.	00:30*	Update	Yellow	Announcement of color code Yellow. Continued increase in seismicity, appearance of continuous tremor, crater lake disappeared.
6/27/92	3:00 a.m.	11:00	Seismicity	Yellow	Vigorous swarm of volcano-tectonic earthquakes began, and it included some long-period earthquakes.
	7:04 a.m.	15:04	Eruption	Yellow	Eruption tremor began; poor weather precluded visual observations.
	7:16 a.m.	15:16	Calldown	Orange	Color code Orange announced by calldown.
	7:16 a.m.	15:16	PIREP	Orange	Eruption column rose to 5,200 m.
	7:40 a.m.	15:40	PIREP	Orange	Plume rose to 5,600 m.
	8:00 a.m.	16:00	Update	Orange	Color code Orange announced by the update. Increased eruption column height; potential for ashfall north and northwest of Mount Spurr.
	8:10 a.m.	16:10	PIREP	Orange	Plume rose to 8,530 m.
	8:22 a.m.	16:22	PIREP	Orange	Plume rose to more than 8,500 m; strong sulfur smell 25 km south of Mount Spurr.
	8:50 a.m.	16:50	Eruption	Orange	Seismic station CPK on crater rim destroyed.
	9:10 a.m.	17:10	Calldown	Red	Color code Red announced.
	9:10 a.m.	17:10	PIREP	Red	Plume height rose to more than 9,000 m.
	9:15 a.m.	17:15	Update	Red	Announcement of color code Red. Plume height rose to 9,100 m (from PIREPS.)
	9:35 – 10:25 a.m.	17:35– 18:25	Seismicity	Red	Tremor amplitude peaked.

Table 1. Chronology of major events and Alaska Volcano Observatory notifications pertaining to the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska—Continued.

DATE	TIME		EVENT	COLOR CODE	EVENT DESCRIPTION OR NOTIFICATION SUMMARY ISSUED
	Local	Universal			
	10:23 a.m.	18:23	Eruption	Red	Plume height reached maximum of 14,500 m, as determined from NWS radar.
	11:07 a.m.	19:07	Eruption	Red	Eruption tremor ended; 24 hours of long-period events followed
	9:00 p.m.	05:00*	Update	Red	Based on morning eruption, as well as character and elevated level of continued seismicity, color code remained at Red.
6/28/92	9:00 a.m.	17:00	Update	Yellow	Announcement of return to color code Yellow. Seismicity less than previous day but still elevated.
6/29/92	1:00 p.m.	21:00	Update	Yellow	Seismicity elevated but it steadily declined.
6/30/92	3:30 p.m.	23:30	Update	Yellow	Seismicity elevated but it steadily declined.
7/1/92	4:00 p.m.	00:00*	Update	Yellow	Seismicity elevated but it steadily declined. Ash cloud reported over Yakutat, Juneau, and Sitka on 6/30/92.
7/3/92	3:30 p.m.	23:30	Update	Yellow	Seismicity elevated but it steadily declined. Two-page summary of eruption issued.
7/7/92	4:00 p.m.	00:00*	Update	Yellow	Continued high level of microearthquakes.
7/8/92	3:00 p.m.	23:00	Update	Green	Announcement of downgrade to color code Green. Microseismicity returned to lowest level of the past few months.
7/10–8/14/92	--	--	Updates	Green	Return to weekly update schedule. Continued low-level seismicity and steam plumes.
8/18/92	3:37 p.m.	23:37	Seismicity	Green	Weak tremor and long-period events detected on proximal Mount Spurr seismic stations (post-eruption analysis).
	3:41 p.m.	23:41	Seismicity	Green	Start of 12-minute period of tremor indicated small premonitory eruption (post-eruption analysis).
	3:48 p.m.	23:48	PIREP	Green	Ash cloud rose 300 to 600 m above Crater Peak.
	4:25 p.m.	00:25*	PIREP	Green	Additional PIREPs confirmed ash plume.
	4:25 p.m.	00:25*	Calldown	Yellow	Announcement of color code Yellow; still no strong seismicity.
	4:42 p.m.	00:42*	Eruption	Yellow	Strong tremor began; NWS radar saw no evidence of eruption.
	4:47 p.m.	00:47*	Calldown	Orange	Announcement of color code Orange.
	4:55 p.m.	00:55*	Eruption	Orange	Plume reached maximum height of 13,700 m, as determined from NWS C-band radar.
	4:58 p.m.	00:58*	Calldown	Red	Announcement of color code Red.
	4:58 p.m.	00:58*	Update	Red	Announcement of color code Red; PIREPS indicated plume rose to 10,700 m; warning issued that winds were toward the east and that ash could begin falling on Anchorage within several hours.
	8:10 p.m.	04:10*	Eruption	Red	Tremor ended.
	8:20 p.m.	04:20*	Eruption	Red	Ash began falling in Anchorage, lasted about 4 hours.
	8:30 p.m.	04:30*	PIREP	Red	Ash-free steam rose to about 5,000 m above Crater Peak vent.
	9:30 p.m.	05:30*	Update	Orange	Announcement of color code Orange based on decreasing seismicity.
8/19/92	5:15 a.m.	13:15	Update	Orange	Decreasing but still elevated seismicity. Ash fell on the northern Kenai Peninsula, Hope, and Whittier. Ash began falling about 1:45 a.m. ADT at Valdez and Cordova.
	8:30 a.m.	16:30	Update	Yellow	Color code downgraded to Yellow. Update included three-page summary of the eruption.
	11:45 a.m.	19:45	PIREP	Yellow	Steam plume and minor amounts of ash emitted from vent.

Table 1. Chronology of major events and Alaska Volcano Observatory notifications pertaining to the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska—Continued.

DATE	TIME		EVENT	COLOR CODE	EVENT DESCRIPTION OR NOTIFICATION SUMMARY ISSUED
	Local	Universal			
	5:00 p.m.	01:00*	Update	Yellow	Decreasing seismicity; steam plumes and minor amounts of ash rose above Crater Peak. Update dispelled rumors of a second eruption that day.
8/20–9/3/92	--	--	Updates	Yellow	Issued every several days as necessary to dispel rumors of eruptions and report investigations at Mount Spurr; color code remained at Yellow; seismicity gradually declined but remained above normal background levels.
9/4/92	3:30 p.m.	23:30	Update	Yellow	Increased long-period seismicity at depths of 20 to 40 km beneath Mount Spurr suggested that magma was again being replenished within the volcano. Update warned of possible eruption within days or weeks; color code continued at Yellow.
9/11/92	2:00 p.m.	23:00	Update	Yellow	Weekly update reported continued high level of seismicity; no additional seismicity in 20- to 40-km depth range, but eruption within days to weeks possible.
9/16/92	7:30 p.m.	3:30*	Seismicity	Yellow	Weak tremor began and was accompanied by an increase in deep earthquakes.
	10:25 p.m.	6:25*	Seismicity	Yellow	Tremor amplitude increased.
	10:33 p.m.	6:33*	Calldown	Red	Announcement of color code Red.
	10:36 p.m.	6:36*	Eruption	Red	Tremor signified an 11-minute preliminary eruptive burst.
	10:37 p.m.	6:37*	PIREP	Red	Ash column rose to 4,600 m.
	10:45 p.m.	6:45*	Update	Red	Announcement of color code Red: eruption began at 10:36 p.m. ADT and NWS wind data showed that the ash cloud would move easterly with fallout expected over the Matanuska-Susitna Valley region and perhaps as far south as Anchorage.
	11:11 p.m.	7:11*	PIREP	Red	Glow observed over Crater Peak.
9/17/92	12:03 a.m.	8:03	Eruption	Red	Seismicity indicated start of main eruption.
	12:08 a.m.	8:08	Eruption	Red	Telemetry lost from seismic station CPK on Crater Peak rim.
	12:30 a.m.	8:30	Update	Red	Plume rose to 10,700 m, reported by PIREPS and NWS radar. Warning issued that southern edge of plume could produce ashfall on Anchorage.
	12:48 a.m.	8:48	Eruption	Red	Plume rose to 12,200 m, as determined from NWS radar.
	12:50 a.m.	8:50	Eruption	Red	Beginning of intense incandescence, as recorded on the slow-scan television camera at Kasilof and on the video camera at AVO in Anchorage.
	2:21 a.m.	10:21	Eruption	Red	Maximum plume height of 13,900 m measured from NWS radar.
	3:39 a.m.	11:39	Eruption	Red	Tremor ended.
	5:45 a.m.	13:45	Update	Red	Eruption ended and seismicity declined although tremor continued; NWS radar showed no plume at Crater Peak; ashfall affected the Matanuska-Susitna Valley and stayed north of Anchorage.
	4:00 p.m.	00:00*	Update	Orange	Color code downgraded to Orange due to decreased seismicity, low-level tremor continued.
9/18/92	3:00 p.m.	23:00	Update	Orange	Summary of eruption issued; color code remained at Orange because of low-level tremor.
9/19/92	4:00 p.m.	00:00*	Update	Orange	Steam plumes rose 3,000 to 6,000 m above Crater Peak; seismicity continued but declined.
9/20/92	4:00 p.m.	00:00*	Update	Orange	Elevated seismicity indicated the possibility of a small eruption with little warning. Therefore, color code remained at Orange.

Table 1. Chronology of major events and Alaska Volcano Observatory notifications pertaining to the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska—Continued.

DATE	TIME		EVENT	COLOR CODE	EVENT DESCRIPTION OR NOTIFICATION SUMMARY ISSUED
	Local	Universal			
9/21/92	4:00 p.m.	00:00*	Update	Orange	Elevated seismicity indicated the possibility of a small eruption with little warning. Therefore, color code remained at Orange.
9/22/92	4:00 p.m.	00:00*	Update	Orange	Elevated seismicity indicated the possibility of a small eruption with little warning. Therefore, color code remained at Orange.
9/23/92	4:00 p.m.	00:00*	Update	Orange	Elevated seismicity indicated the possibility of a small eruption with little warning. Therefore, color code remained at Orange.
9/24/92	4:00 p.m.	00:00*	Update	Yellow	Announcement of downgrade to color code Yellow. Tremor amplitude decreased and time between tremor episodes increased.
9/25/92	4:00 p.m.	00:00*	Update	Yellow	Continued low-level seismicity.
9/28/92	4:00 p.m.	00:00*	Update	Yellow	Continued low-level seismicity.
9/29/92	4:00 p.m.	00:00*	Update	Yellow	Continued low-level seismicity.
9/30/92	4:00 p.m.	00:00*	Update	Yellow	Continued low-level seismicity.
10/1/92	4:00 p.m.	00:00*	Update	Yellow	Continued low-level seismicity.
	7:00 p.m.	03:00*	Seismicity	Yellow	Continuous volcanic tremor began.
	11:00 p.m.	07:00*	Update	Yellow	Weak, continuous volcanic tremor resumed under Crater Peak; increased SO ₂ emission; heightened possibility of eruption during the next several days.
10/2/92	6:17 p.m.	02:17*	Calldown	Red	Announcement of color code Red based on intense seismicity.
	6:30 p.m.	02:30*	Update	Red	Announcement of color code Red: intense seismic tremor at Crater Peak; tremor amplitude comparable to levels preceding the June 27 and September 17 eruptions; large eruption seemed likely within 24 to 48 hrs.
	6:30 p.m.	02:30*	Seismicity	Red	Tremor abruptly ended.
10/3/92	4:00 p.m.	00:00*	Update	Red	Irregular tremor signified unrest of the volcano. Eruption could occur with little or no warning; color code remained at Red.
10/4/92	9:00 a.m.	17:00	Update	Red	Continued volcanic unrest indicated by banded tremor with tremor bursts that lasted 1 to 2 hours.
	5:00 p.m.	01:00*	Update	Yellow	Announcement of downgrade to color code Yellow. More quiet time between tremor periods; banded tremor, which often precedes eruption, continued; volcano still considered restless.
10/5/92	4:00 p.m.	00:00*	Update	Yellow	Continued banded tremor and an 18-hour period of sustained tremor.
10/6/92	4:00 p.m.	00:00*	Update	Yellow	Intermittent tremor persisted.
10/7-11/6/92	- -	--	Updates	Yellow	Daily weekday reports of continued irregular levels of seismicity and occasional gas emission, as well as other observations of the volcano. Caution issued that an eruption could occur with little warning.
11/2/92	2:00 a.m.	11:00	Seismicity	Yellow	Rate of shallow volcano-tectonic earthquakes increased.
11/9/92	12:00 m.	21:00	Update	Orange	Announcement of upgrade to color code Orange based on shallow seismic swarm; eruption could occur without additional warning.
	4:00 p.m.	01:00*	Update	Orange	Increased shallow seismicity.
	9:00 p.m.	06:00*	Update	Orange	Number of seismic events was 1.5/min for 3.5 hrs; this was the most energetic seismic swarm of 1992 eruptive period.

Table 1. Chronology of major events and Alaska Volcano Observatory notifications pertaining to the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska—Continued.

DATE	TIME		EVENT	COLOR CODE	EVENT DESCRIPTION OR NOTIFICATION SUMMARY ISSUED
	Local	Universal			
	10:07 p.m.	07:07*	Update	Red	Announcement of upgrade to color code Red. Significant increase in seismicity on all stations; large eruption seemed likely within 24 to 48 hrs.
11/10/92	3:00 p.m.	00:00*	Update	Orange	Color code downgraded to Orange. Seismicity decreased significantly since the buildup of 11/9/92.
11/11/92	3:00 p.m.	00:00*	Update	Orange	Decreased seismicity from the swarm of November 9; continued unrest could result in eruption without additional warning.
11/12/92	3:00 p.m.	00:00*	Update	Yellow	Color code downgraded to Yellow based on 2 days of relative seismic quiescence; however, volcanic unrest continued.
11/13– 12/18/92	--	--	Updates	Yellow	Daily weekday reports of continued irregular levels of seismicity and occasional gas emission, as well as other observations of the volcano, still indicated that an eruption could occur with little warning.
12/21/92	3:00 p.m.	00:00*	Update	Yellow	Color code would remain at Yellow until early January because: (1) continued high level of seismicity, and (2) eruptions and seismic crises since June eruption occurred at 55-, 30-, and 53-day intervals; and by early January, 60 days had passed since the seismic swarms on November 9–10.
12/24/92	12:00 m.	21:00	Update	Yellow	Swarm of small earthquakes occurred at depth of 8 to 12 km beneath Crater Peak during past 5 days; continued unrest indicated that the volcano was still unstable and eruption could occur with little warning.
12/28/92	3:00 p.m.	00:00*	Update	Yellow	Swarm of small earthquakes of December 20 continued through December 26, followed by increased shallow 2-hour periods of tremor-like events beneath Crater Peak on December 27; seismicity then decreased; volcano still considered unstable.
12/31/92	12:00 m.	21:00	Update	Yellow	Continued low-level seismicity including two earthquakes beneath Crater Peak and two low-amplitude tremor episodes during the previous 24 hrs. Decision made for color code to stay at Yellow.
1/4–2/26/93	--	--	Updates	Yellow	Return to weekly Friday reporting of low-level seismicity, steam plumes, and other observations.
3/5/93	2:00 p.m.	23:00	Update	Green	Color code downgraded to Green. Seismicity continued at a low level for a time period longer than that between the 27 June and 18 August 1992 eruptions (see 12/21/92); possibility of steam and ash explosions still existed.
3/12/93– 1/14/94	--	--	Updates	Green	Mention of Mount Spurr seismicity and other observations in weekly updates.
1/21/94	--	--	Update	Green	Announcement that level of seismicity at Mount Spurr had returned to background levels and that specific mention of Mount Spurr in weekly updates would be discontinued.

of 1992, however, updates stated that the increased seismicity was not considered precursory to eruption. The heightened seismicity continued, punctuated with stronger pulses, including a prominent pulse beneath Crater Peak in May. In an internal memorandum dated June 3, 1992, AVO scientists interpreted the seismicity as the result of increasing magmatic pressure at depth, which caused intrusion into conduits below Crater Peak and the summit dome complex and minor movement of caldera faults. The memo urged increased monitoring, as well as advising government agencies

and the aviation industry to review plans for coping with an eruption. Both of these recommendations were implemented.

The number of daily locatable VT earthquakes reached 28 on June 5 and then declined to about 3 to 6 events per day. Volcanic tremor began on June 6, and an overflight on June 8 revealed upwelling in the crater lake, which had turned in color from green to gray. AVO released a special update describing the situation and stating that there would be an increase in overflights and in sensitivity of the 24-hour alarm

system at the AVO Seismology Laboratory in Fairbanks, along with an added evening shift for AVO personnel. As many as 24 shallow tremor bursts per day, each burst lasting 1 to 10 minutes, were recorded at stations within 10 km of Crater Peak from June 6 to 26 (McNutt and others, this volume). A field party spent 6 hours in the crater on June 11, measured a lake temperature of 50°C and pH of 2.5, and collected water samples for chemical analysis (Keith and others, this volume). The party observed small geysers—never before seen in the crater—at the base of the north crater wall. Geysering did not coincide with seismic tremor events.

On June 17, AVO informed ADES of the evolving situation at Mount Spurr. On June 19, AVO sent briefing materials to the Governor, noting the “well above normal” activity and continuing episodes of tremor. The Anchorage Daily News published a front-page article on June 21 that described how Mount Spurr was being monitored and what the public might expect from an eruption.

JUNE 27 ERUPTION

Seismic behavior changed ominously on June 24 when a tremor episode lasted 154 minutes, followed 12 hours later by a similar episode that lasted 142 minutes. Eight additional tremor bursts occurred within the next 8 hours. The weekly update of 10:30 a.m. Alaska daylight time (ADT) on June 26 reported “well above normal” seismic activity but still cautioned that an eruption might not be imminent in view of an absence of long-period (LP) earthquakes. At about the same time, aerial observations of the crater revealed that the lake had almost completely drained and that several large rocks had impacted the resulting mud flat. At 12:04 p.m. ADT, tremor that was continuous and stronger than earlier bursts began. AVO formally issued a warning of level of concern color code Yellow at 4:30 p.m. ADT and went on 24-hour duty. At 3:00 a.m. ADT on June 27, a swarm of VT earthquakes struck at 0- to 2-km depths beneath Crater Peak; their rate soon increased to about one every 2 minutes. Three LP events accompanied this swarm. Tremor amplitude abruptly doubled at 7:04 a.m. ADT. This increase in amplitude was later interpreted to represent the onset of eruption, although weather clouds prevented visual verification. At 7:16 a.m. ADT, AVO began an emergency calldown announcing level of concern color code Orange. About the same time, telemetry was lost from the seismic station 400 m from the vent and an Alaska Airlines pilot reported that an eruption plume had risen 5,000 m above the cloud cover. AVO announced color code Red at 9:10 a.m. ADT. Tremor amplitude gradually increased, peaking

between 9:35 a.m. and 10:25 a.m. ADT and registering on stations more than 100 km away. Pilots estimated the plume at mid-morning as high as 9,000 m and the NWS measured a maximum plume height of 14,500 m with C-band radar (Alaska Volcano Observatory, 1993; Rose and others, this volume). The tephra cloud moved northward and ash began falling on Denali National Park at 10:30 a.m. ADT. Debris flows swept southward down narrow drainages and entered the Chakachatna River in three places. Most debris followed the course of the 1953 lahar. At about 11:30 a.m. ADT, seismicity decreased abruptly and the eruption was over. Weather and steam obscured Crater Peak and the eruption plume track from aerial observation on the afternoon of June 27, but paths of the debris flows were visible on the lower flanks of the volcano (Meyer and Trabant, this volume).

At 9:00 a.m. ADT on June 28, the level of concern was downgraded to Yellow. Aerial observation revealed a black northward-broadening swath of tephra on snow fields and glaciers. Ash thickness was about 1 to 2 mm at Denali National Park and Manley Hot Springs, 260 and 420 km north of Crater Peak, respectively (Neal and others, this volume). This region is sparsely populated and there were no other reports of tephra there. The ash cloud continued northward to the Beaufort Sea, then it turned southeast into Canada and the coterminous United States, where it became indistinguishable from weather clouds about July 2.

The level of concern color code was downgraded to Green on July 8. This was done partly on the basis of greatly reduced seismicity and SO₂ emission, but also on the basis of analogy to the single-eruption pattern of 1953. The downgrading to color code Green proved to be premature.

AUGUST 18 ERUPTION

Following the June eruption, seismicity remained low through the first half of August. Only one shallow and two deep earthquakes were recorded between August 12 and 17. Because the closest operational seismic station at that time was 5 km from the vent, several attempts were made to reestablish a seismic station on the crater rim. These were unsuccessful because of poor weather conditions. At 3:37 p.m. ADT on August 18, a 16-minute episode of weak tremor and several LP events began, but these rather obscure events were not identified until post-eruption analysis of the data. At 3:48 p.m. ADT, a pilot reported an ash-rich plume. With confirmation of this plume at 4:25 p.m., AVO began a calldown announcing level of concern color code Yellow. The main eruption be-

gan at 4:42 p.m. ADT, when strong tremor was recorded by all Mount Spurr seismic stations. AVO began a calldown announcing color code Orange at 4:47 p.m. ADT, but repeated the calldown process 11 minutes later to raise the color code to Red. By 4:58 p.m. ADT, a subplinian column had risen through low clouds to a height of 11,000 m, and it ultimately reached nearly 14,000 m. From an aircraft only 2.5 km away, AVO staff observed and videotaped a dark roiling cloud that was periodically surrounded by lenticular shock waves. Large bombs were thrown 800 m above the vent. Small-volume pyroclastic flows of breadcrusted blocks descended the east and southeast flanks of Crater Peak; these flows formed coarse, clast-supported lobate deposits with steep-fronted margins. Other flows mixed with snow and ice high on the cone and became lahars. A late fusillade of mostly lithic ballistic projectiles, some as large as 1 m, were hurled as far as 10 km southeastward (Waitt and others, this volume). More than 170 lightning strikes were detected by the AVO lightning detection system (LDS) during the second half of the eruption (Paskievitch and others, this volume). The eruption ended at 8:10 p.m. ADT.

Upper-level winds moved the eruption plume east-southeast directly over Anchorage, where it deposited as much as 3 mm of sand-sized ash (Neal and others, this volume). A satellite image 44 minutes after the onset of eruption shows the plume extending 80 km east from the volcano over an area of 2,000 km². Three hours after onset of eruption, the leading edge of the plume was 300 km southeast of Mount Spurr, and its area had grown to 21,000 km².

Ashfall forced the closing of Anchorage International Airport for 20 hours (N.W. Gibson, Anchorage International Airport, written commun., 1993). Air quality alerts were issued in Anchorage during the fallout period and also on the following day, as vehicle traffic stirred up ash again (R.B. Stewart, Office of Emergency Management, Municipality of Anchorage, written commun., 1993). Reworked windblown ash continued to reduce air quality until the first snow of autumn, and then it reappeared during the summer of 1993.

SEPTEMBER 16-17 ERUPTION

Following the August eruption, deep seismicity gradually increased, and by mid-September it had returned to levels comparable to mid-June. AVO teams visiting the crater on September 7 and 16 noted nothing unusual. At about 7:30 p.m. ADT on September 16, however, discrete seismic events and weak tremor were detected by the newly reinstalled crater rim sta-

tion. Tremor amplitude increased at 10:25 p.m. ADT, and at 10:33 p.m. ADT, AVO declared concern color code Red and began the emergency calldown. An eruption began at 10:36 p.m. ADT that lasted 11 minutes. Incandescence was recorded on the video camera at AVO-Anchorage and on the telemetered slow-scan television camera at Kasilof, 120 km southeast of Crater Peak just south of Kenai on the Kenai Peninsula. Weak tremor through the next hour foreshadowed the main phase of eruption, which began at 12:03 a.m. ADT on September 17. Intermittent bright incandescence could be seen from Anchorage.

The September 17 eruption lasted 3.6 hours. Pyroclastic flows swept down the south-southeast and southeast flanks of Crater Peak and mixed with snow and ice to become lahars. These flows were similar in appearance to pyroclastic flows, but they were cool and water saturated hours after emplacement. Tephra fallout on the Kidazgeni Glacier generated a debris flow that temporarily dammed the Chakachatna River. Once again, a narrow ballistic field extended at least 10 km from the vent along the south margin of the tephra plume. The eruption closed with a strong swarm of about 50 VT shocks between 5 and 10 km in depth, which may reflect readjustment of the conduit after magma withdrawal.

The plume moved eastward, dusted the north edge of Anchorage and deposited about 1.5 mm of ash in Palmer, Wasilla, and nearby communities in the Matanuska-Susitna Valley north of Anchorage. Very light ashfall was reported in the town of Glenallen, 350 km east of Crater Peak.

SEISMIC SWARMS: OCTOBER 2-6, NOVEMBER 9-10, AND DECEMBER 21-27

The level of seismicity remained elevated, and several seismic crises caused concern at AVO. Tremor between 7:00 p.m. ADT on October 1 and 6:30 p.m. ADT on October 2, followed by quasiperiodic tremor over the next 72 hours, prompted AVO to announce concern color code Red twice during that period, but no eruption ensued.

At 10:07 p.m. Alaska standard time (AST) on November 9, AVO again announced concern color code Red an hour into an intense seismic swarm that lasted 3.5 hours and contained 170 detected earthquakes. These earthquakes had mixed frequency contents, occurred within a very restricted volume 1.2 km beneath the vent, and had magnitudes as large as 1.7. Again, seismicity receded without eruption. Because the duration of this swarm was similar to each of the three eruptions and its seismic energy release was the largest of the 1992 eruptive period, the swarm may be

viewed as a "failed" eruption, in which magma was rapidly emplaced at shallow depth (Power and others, this volume).

A final flurry of earthquakes beneath Crater Peak on December 21 through 26 was followed by a few hours of shallow tremor-like activity on December 27. This time, AVO's advisories remained at concern color code Yellow.

At the beginning of 1993, Mount Spurr returned to a low level of seismic activity: 1 to 10 small earthquakes per day at depths of 0 to 40 km and rare periods of possible weak tremor. This background seismic "noise" fluctuated daily and seasonally (G. Tytgat and others, unpub. data). Some shallow seismic events were correlated with observed landslides, rockfalls, glacial melt runouts, and even airplanes flying over the volcano. Some seismicity may be attributed to post-eruptive readjustments of the magmatic and hydrothermal systems, and some remains unexplained. On March 5, 1993, after a time period greater than the seismic quiescence between the June 27 and August 18, 1992, eruptions had passed uneventfully, AVO lowered the level of concern color code to Green.

Vigorous white steam plumes caused false reports of eruption several times during 1993. However, volcanic gas (SO_2 and CO_2) emissions remained at or below the limit of detectability (less than 1,000 tonnes per day CO_2 , less than about 4 tonnes/day SO_2 ; Doukas and Gerlach, this volume). New concern arose on June 29, 1993, when tremor-like seismicity appeared on records from several of the Crater Peak seismic stations. A field party was fortuitously present near Crater Peak and observed the source to be an outburst flood from Kidazgeni Glacier on the southeast slope of Mount Spurr (Nye and others, this volume).

During and following the 1992 eruptions, field investigations were carried out to map and sample the eruption products. Work in the vent area of the crater was precluded through the summer of 1993 for safety reasons. This was a lesson drawn from the sudden explosions at Galeras Volcano, Columbia, and Guagua Pichincha Volcano, Ecuador, in early 1993, that resulted in the deaths of eight volcanologists, and the September 1992 eruption of Crater Peak, which followed a period of seismic quiescence and occurred just hours after an AVO team had left the vent.

GENERAL CHARACTERISTICS OF THE ERUPTIONS

The June, August, and September 1992 eruptions produced 44, 52, and 56 million m^3 of tephra (12, 14, and 15 million m^3 dense rock equivalent or DRE), respectively. In each case, juvenile tephra clasts ranged

from proximal breadcrusted bombs to distal ash in a narrowly defined band of fallout extending downwind (Gardner and others, 1993; Neal and others, this volume). The vent migrated about 50 m westward from June to August to September, and so in September it was nestled against the west-northwest crater wall. Near-field ejecta from each eruption were, therefore, directed progressively more strongly to the east-southeast. Otherwise, vent geometry changed remarkably little. Distribution of pyroclastic flow deposits, ballistic showers, and lahars shows that all flanks of a volcanic cone are hazardous during subplinian activity (Miller and others, this volume; Waitt, this volume).

The eruptions were strongly magmatic in character. The dominant juvenile component of 1992 ejecta is dense to scoriaceous dark gray andesite (57 weight percent SiO_2 bulk composition) with a more silicic (60–62 weight percent SiO_2) microlite-rich matrix of brown glass (Harbin and others, this volume; Nye and others, this volume; Swanson and others, this volume). A subordinate volume of bulk-chemically identical andesite is more crystal-rich and contains clear matrix glass of rhyolitic composition (73–75 weight percent SiO_2). Trace amounts of rhyolite pumice with sparse plagioclase and quartz phenocrysts were also erupted. Pyroclastic flows of August and September contained large partially melted and greatly inflated felsic metamorphic xenoliths (Miller and others, this volume). The estimated total volume of magma for the three Crater Peak eruptions of 41 million m^3 DRE is markedly smaller than the estimated 200 million m^3 DRE (Gardner and others, 1994) released in some 20 significant eruptive events during the 1989–90 Redoubt activity.

ANALYSIS OF RESPONSE

AVO's response plan went into effect as soon as unrest was recognized at Mount Spurr in August of 1991, and it continued through post-eruptive readjustment to inactive status. For alerting a community to an imminent volcanic danger, Mount Spurr provided a favorable situation. Its Holocene deposits had been mapped and some 40 tephra layers from its 5,000-year-old Crater Peak vent had been identified (Riehle, 1985; Nye and Turner, 1990); these deposits place it among the more prolific volcanoes in the eastern Aleutian Arc. Moreover, it had erupted during historical time (Juehle and Coulter, 1955; Wilcox, 1959). In retrospect, the singularity of the 1953 eruption proved misleading for what was to come, but it did provide a useful example of Mount Spurr's behavior. Because of the relatively well documented record of activity,

Mount Spurr had already been instrumented and monitored for over a decade when it reawakened. This lengthy period of monitored quiescence allowed the August 1991 seismicity to be accurately interpreted as anomalous. Mount Spurr's reawakening was gradual, and thus it permitted plans to be made and warnings to be given in a timely fashion. It was not so fast that AVO was caught off guard, nor so slow that the warnings of a future eruption could be dismissed. Care must be taken when turning observations of increasing seismicity into an appropriate warning, because not all cases of increasing seismicity, even if caused by rising magma, lead to eruption. Often, magma can solidify in the conduit without venting (Newhall and Dzurisin, 1988), as may have occurred during the seismic swarms of May, October, November, and December of 1992 (table 1). In retrospect, informing a community about the possibility of an eruption a month in advance, and issuing a serious warning a day in advance, is a useful time frame. Eruption warnings of October and November that were not fulfilled were considered by the public to be proper caution rather than mistaken judgment, coming as they did late in the episode.

Although seismic monitoring was the primary tool in AVO's response, geologic insight was necessary both for informed warnings of what was to come and for the original decision to monitor the seismic activity of Mount Spurr. A broad host of observations and techniques was used to evaluate the volcano's condition.

INTERPRETATION OF SEISMICITY

Following the establishment of AVO, seismic stations were added to the existing Mount Spurr network, making a total of 10 stations within 15 km of the summit (McNutt and others, this volume; Power and others, this volume). Seismicity that began in August 1991 eventually defined an active conduit without, unlike Redoubt Volcano, a hint of a volumetrically large storage zone. That elevated seismicity occurred throughout the volcanic massif indicates that stresses induced by magma intrusion and perhaps even magmatic vapor affected the entire Mount Spurr system. Helicorders, continuous digital data, event-triggered data, real-time seismic amplitude measurements (RSAM; Endo and Murray, 1991) and seismic spectral amplitude measurements (SSAM) were all used for analysis of the seismic information (McNutt and others, this volume).

Three types of shallow seismicity indicative of magma movement were identified: (1) tremor bursts that lasted 1 to 10 minutes, (2) continuous tremor lasting 2 hours to several days, and (3) eruption tremor (McNutt and others, this volume). Fourteen hours of continu-

ous tremor and 4 hours of shallow VT earthquakes preceded the June eruption. However, this seismic pattern that made forecast of the June eruption possible was not repeated in subsequent events. Only 3 hours of increased seismicity preceded the September eruption; accordingly, the warning from AVO was issued just minutes ahead of the onset of eruptive activity and an hour and a half before the main pulse. Unfortunately, no detectable seismic activity foreshadowed the August eruption. Had it been possible to replace the near-vent seismic station destroyed during the June eruption, shallow precursory tremor bursts might have been recorded. However, the absence of the near-vent station does not explain the lack of locatable VT earthquakes before that eruption. Presumably, the strongest precursory signals were generated at the outset of the eruptive episode because rising magma and gases were reopening a cold, plugged conduit.

EVIDENCE FROM HYDROTHERMAL ACTIVITY AND GAS EMISSION

The sense of urgency engendered by May and June seismicity was heightened by observations of the small crater lake that had existed since the 1952 eruption. Days prior to the first eruption, lake water changed from blue-green to gray, became turbid with areas of upwelling, and finally evaporated or drained (Keith and others, this volume). A single pair of measurements, one during dormancy and the other in June, revealed a two-orders-of-magnitude increase in SO_4/Cl in the lake.

The increase in SO_4/Cl may relate to the failure of airborne SO_2 emission measurements to detect the rise of magma (Doukas and Gerlach, this volume). Airborne measurements were begun when anomalous seismicity was first detected. Not until the June eruption did the flux of SO_2 from the vent increase significantly, and SO_2 concentrations returned to low values 2 days later. SO_2 flux was also low prior to and following the August and September eruptions. Yet, the National Aeronautical and Space Administration's (NASA) Total Ozone Mapping Spectrometer (TOMS) indicated 200 ± 60 , 230 ± 70 , and 400 ± 120 kt of SO_2 emitted during the June, August, and September eruptions, respectively. It is likely that SO_2 released during pre-eruptive intrusion was being scrubbed and the resulting H_2S and sulfate dissolved in the hydrothermal system. Doukas and Gerlach (this volume) and Bluth and others (this volume) conclude from independent techniques that a significant amount of H_2S was emitted during the 1992 eruptions. Airborne CO_2 measurements made following the September eruption did detect substantial magmatic outgassing, which suggests that monitoring of this less water-soluble gas might have provided a more diagnostic tool than monitoring of SO_2 (Doukas and Gerlach, this volume).

VISUAL OBSERVATION OF ERUPTIONS

Visual observations, when they can be obtained, are the most direct means of assessing eruptive activity. AVO's remote, slow-scan, continuously recording television camera at Kasilof on the Kenai Peninsula was turned to view Crater Peak, 120 km away, on May 6, 1992; it provided real-time views for AVO staff in both Anchorage and Fairbanks. In addition, a video camera was set up on the roof of AVO in Anchorage and wired into the operations room. Both cameras recorded the incandescence at the crater rim during the September 16–17 eruption (Neal and others, this volume). When weather permitted, the information from these cameras was critical in responding to false eruption alarms.

Especially helpful to AVO were the pilot reports (PIREPS) of visual observations passed on by FAA and NWS. For example, although seismicity indicated that the June eruption was imminent, the ash cloud was reported by a PIREP before seismic verification could be made and while clouds prevented ground-based visual observations of the volcano. A PIREP revealed the August eruption, because there was no precursory seismicity. PIREPs also provided a valuable record of eruption column height as a function of time.

DETECTION OF VOLCANIC LIGHTNING

Once volcanic activity has been detected seismically or visually, the most important task is to establish whether ash is being emitted. This can be difficult, given weather and light conditions in Alaska. AVO has established a lightning detection system (LDS) for the Cook Inlet region; this system is based on observations that lightning is commonly associated with vigorous, ash-laden eruptive columns at volcanoes throughout the world (Hoblitt, 1994; Paskievitch and others, this volume). Detected lightning might therefore indicate significant ash emission. During the Crater Peak eruptions, this technique proved only partially successful. A ring of lightning 10 km in diameter, centered 5 km east of the vent, was detected during the August eruption, but LDS did not detect lightning during the June and September eruptions. That lightning accompanied the June and September eruptions is indicated by interruptions in signals from seismic stations (Davis and McNutt, 1993).

RADAR OBSERVATION OF ERUPTIVE COLUMNS

One technique for detection of eruptive columns is C-band radar. An experiment conducted in cooperation with NWS by Rose and others (this volume)

at Kenai provided a first test with excellent results. Radar imaged the Crater Peak eruption columns regardless of weather, and it also allowed estimates of column height and mass flux to be made. A limitation of radar observation is that only coarser particles are reflective, and so neither the top of the eruptive column, nor the cloud tens of minutes after cessation of eruption, can be seen.

PROJECTION OF PLUME PATHS

During the 1980 eruptions of Mount St. Helens, graphic plots of wind direction and velocity at several altitudes, used with permission of NWS, proved useful in determining the path of eruption plumes (Miller and others, 1981). Their usefulness was demonstrated again at Mount Spurr, particularly in calculating the time for the eruption cloud of August 18 to reach Anchorage. In addition, it was immediately obvious that the June 27 eruption cloud would drift north toward sparsely populated areas and would not affect Anchorage. Winds for the September 18 eruption were flowing northeastward from the vent. A warning of possible ashfall was given to Anchorage, although the model correctly showed the ash-laden winds passing just north of the city. The model correctly predicted that ash would fall mainly on the Matanuska-Susitna Valley. Also important in tracking ash were calls from citizens informing AVO of ashfall locations and thicknesses.

TRACKING OF PLUMES BY SATELLITE

The hazards of volcanic ash emitted during eruptions extend well beyond local communities. For the first several days, a cloud from an explosive eruption is dangerous to high-flying jets (Casadevall, 1992; Casadevall and Krohn, this volume). During the Mount Spurr eruptions, advanced very high resolution radiometry (AVHRR) data were processed using a new high-resolution picture transmission (HRPT) information processing system (HIPS) at NWS for real-time tracking of volcanic clouds (Schneider and others, this volume). The August and September eruption clouds were followed for more than 80 hours and thousands of kilometers (Schneider and others, this volume). Further development of this technique may make it possible to determine particle-size distribution and particle density in volcanic clouds, so the severity of hazard they pose can be more quantitatively evaluated.

TOMS satellite images were likewise used to track the eruption clouds for as many as 7 days after formation (Bluth and others, this volume). Sulfur diox-

ide content of clouds from the August and September eruptions increased during their first 2 days, probably by atmospheric oxidation of H_2S . At least half the 400 ± 120 kt of SO_2 emitted by the three eruptions reached the lower stratosphere.

DISSEMINATION OF INFORMATION

AVO's use of updates, level of concern color code, calldown lists, and direct personal communications worked effectively to notify civil authorities and the general public of expected eruptions and resulting hazards from Mount Spurr. Intensive media coverage, based on interviews with AVO personnel by both local and national radio and television and the local newspaper, added to AVO's outreach.

AVO's color code is central to rapid communication of the level of hazard to concerned constituencies, particularly the aviation industry. As indicated in figure 4, increasing levels of concern convey both increasing likelihood of eruption and increasing potential or actual size of eruption. No simple one-dimensional scale can adequately describe an impending or ongoing event. Yet, simplicity of communication is essential during an emergency. In practice, use of the code is tempered by awareness of the consequences that communication of the concern will cause. During the 1992 Crater Peak eruptions, color was used primarily to indicate likelihood of eruption. Red meant that an eruption was imminent or ongoing; Red was used because all events were major ones with eruptive columns that reached flight levels and produced regional tephra falls.

Two occasions show the kinds of decisions that must be made in applying the code and that include factors beyond the simple criteria in the code. The volcano's behavior arguably fell within the condition "restless" beginning in August of 1991. However, uncertainty as to the significance of the early seismicity, and knowledge that premature use of warning colors would entail unnecessary fueling and routing costs for the airlines, led to deferral of use of color code Yellow. A more literally based interpretation of the code was used on August 18, 1992, when AVO downgraded the concern level from Red to Orange at the end of the eruption, but 3 hours before ash stopped falling on Anchorage. The color code was not meant to address conditions away from the volcano, so the downgrading probably did not seem an adequate portrayal of the situation to citizens of Anchorage who were just beginning to be affected by the eruption. Never-

theless, the AVO decision was appropriate to signal to the airlines and others that voluminous ash production at the vent had ceased and the end of ashfall in Anchorage could be predicted.

ECONOMIC EFFECTS OF THE 1992 ERUPTIONS

Economic losses from the Crater Peak eruptions were mainly effects of ashfall at Anchorage International Airport and associated disruption of air traffic from the August 18 eruption. Significant losses also included cleanup costs for the Municipality of Anchorage, adjacent military bases, and towns in the Matanuska-Susitna Valley. The cost of cleanup from the August 18, 1992, ashfall for Anchorage International Airport, Merrill Field, and Elmendorf Air Force Base, including cleaning buildings, runways, and aircraft on the ground, was at least \$683,000 (Anchorage International Airport, unpub. report, 1993). Revenue lost from the closure of Anchorage International Airport for 20 hours is estimated at \$276,000 (Casadevall and Krohn, this volume). The Municipality of Anchorage reported nearly \$2 million in damage, office closures, and cleanup costs from the August eruption (Municipality of Anchorage, unpub. report, 1993). In addition, there were numerous, small-scale losses incurred by businesses and citizens from the initial ashfall and reworked volcanic ash that affected equipment such as vehicle engines, building air-circulation systems, and computers in banks, businesses, and local military installations. Economic impacts of ashfall on the upper Kenai Peninsula, Cordova, and the Matanuska-Susitna Valley have not been evaluated. AVO received calls from these areas following all three eruptions reporting school closures, poor road conditions, and health concerns. No loss of life resulted directly from the eruptions. Two heart attacks, one fatal, from shoveling ashfall were reported in Anchorage.

In contrast to the estimate of \$160 million in losses from the 1989–90 Redoubt eruptions (Tuck and Huskey, 1992), total losses from Crater Peak eruptions are informally estimated to be from \$5 to 8 million. This large difference arises from at least two causes. First, there were more eruptive pulses and greater eruptive volume in the Redoubt episode. Second, Redoubt losses were dominated by damage sustained by five jets (Casadevall, 1994). Damage to one of these aircraft nearly resulted in a catastrophe. The Crater Peak eruptions carried the same potential for disaster. However, increased awareness of the ash-cloud

hazard within the aviation community and emplacement of an effective warning system by AVO, in concert with other agencies, resulted in minimal economic loss to air carriers.

VOLCANOLOGICAL IMPLICATIONS OF THE 1992 ERUPTIONS

The 1992 eruptions of the Crater Peak vent of Mount Spurr provide interesting insights into eastern Aleutian volcanism as well as andesitic volcanism in general. Some salient points are described below:

1. Less silicic magma composition does not necessarily mean less explosive eruptions. Silicic volcanism is often equated with explosive eruptions, and it is tempting to think of more mafic eruptions as more likely to produce lava flows. The Crater Peak eruptions were distinctly more mafic than those of either Redoubt or Augustine. Yet, these eruptions were entirely explosive, whereas a significant proportion of recent Redoubt and Augustine activity has been effusive. A rationale often applied to the case of effusive mafic eruptions is that in more fluid magma gas bubbles escape more readily and thus preclude significant buildup of pressure. However, a fluid magma can also ascend more rapidly than a viscous one. Rapid ascension of magma would prevent release of vapor from magma in the subsurface prior to eruption (Jaupart and Allegre, 1991) and could explain the explosiveness of the Crater Peak eruptions (Eichelberger, 1995).

2. Each of the three eruptions was remarkably similar in style and volume (Gardner and others, 1993; Neal and others, this volume). Quiet periods between events were of similar duration. The simplest interpretation is that the erupted magma was stored in the same way prior to each eruption. Storage at the same depth would lead to the same initial volatile conditions. One might envisage a conduit that gradually filled to the same volume prior to each event, emptied in each event, and was not physically altered during each event. In such a view, cessation of the complete eruptive episode represents decline in source region overpressure (Tait and others, 1989), and cessation of an individual eruption represents exhaustion of conduit-stored magma. Storage in a conduit extending to the base of the seismogenic zone at 40 km (Power and others, this volume) would require a conduit radius of 10 m.

3. Like Redoubt, there is evident heterogeneity in melt composition in the system (Nye and others, this volume; Swanson and others, 1994), but unlike

Redoubt, there is little heterogeneity in bulk composition (Nye and others, this volume; Nye and others, 1994). This phase heterogeneity and bulk homogeneity probably represents the existence of magmatic domains that had similar origins but different pressure/temperature histories, and that have not interacted chemically with the mid- to upper-crust (despite the presence of abundant melted felsic xenoliths), nor lingered and fractionated there. Again unlike Redoubt, melt heterogeneity increases with time through the episode (Harbin and others, this volume). Perhaps this difference relates to the clearly different styles of magma storage beneath the two volcanoes, but that relation is not clear.

4. Despite the similarities of eruptive events within the 1992 episode, the character of associated seismicity varied greatly (Power and others, this volume; McNutt and others, this volume). There was seismicity preceding and following eruption, seismic quiescence preceding and following eruption, and seismicity without eruption. Because much of the seismicity defines a near-vertical line from the active vent to 40 km in depth, there seems little doubt that the seismicity is associated with flow of magma. It must be concluded, however, that Mount Spurr's magma can also move aseismically, and that significant shallow seismically detected movement of magma can occur without eruption. There is a mirror symmetry to the pattern, that is seismicity occurred without eruption at the beginning and end of the episode and aseismic eruption occurred in the middle. This pattern is perhaps consistent with thermal waxing and waning of the shallow system. When the system was hottest, less seismogenic brittle fracturing or interaction between magma and hydrothermal fluids occurred. This symmetry, however, is not complete in terms of distribution and type of seismicity, and much early seismicity occurred far from plausible sites of intrusion. The pattern is complex. Apparently, multiple processes associated with the rise of magma generated the seismicity.

5. Although the most recent prior eruption was a good guide to the style of the 1992 events, its singularity was not. Past activity should be interpreted cautiously, especially when the data set is limited.

Interpretation of available data for the Mount Spurr volcanic system suggests the rapid migration of relatively fluid, volatile-rich magma through a narrow, deeply rooted conduit. Useful lessons have been drawn in terms of eruption prediction. For example, the importance of even weak tremor in presaging eruption and the variability of the relation between seismicity and impending eruption. However, much re-

mains to be learned about underlying processes. The origin of seismicity accompanying magma flow is not clearly understood, nor are the mechanisms by which new magma reopens an existing conduit. Whereas petrologic inferences can be drawn about equilibrium conditions of magma storage, the chemical and physical changes that accompany magma rise and produce the large variations in eruptive behavior are poorly known. A hope for the future is that geophysical and geologic insights from active systems such as those gained from the Crater Peak eruptions will provide the basis for a better understanding of magma's ascent to the surface. Such understanding could lead to better predictive capability—that is, to predict whether magma will merely intrude or will erupt, and to predict whether eruptions are likely to be explosive or effusive.

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Real-Time C-Band Radar Observations of 1992 Eruption Clouds from Crater Peak, Mount Spurr Volcano, Alaska

By William I. Rose, Alexander B. Kostinski, and Lee Kelley

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ABSTRACT

Repeated aircraft hazards in Alaska related to volcanic clouds have resulted in the use of a mobile C-band radar devoted to volcanic-cloud monitoring. The radar is located at Kenai, in range of several volcanoes in the Cook Inlet area. Three significant eruptions from the Crater Peak vent of Mount Spurr volcano (about 80 km from Kenai) in 1992 provided the first tests of the radar. The system constructs maps of the eruption columns and the drifting ash clouds for short periods (as long as 30 minutes) after eruption. The radar gives direct information about active eruptions in any weather conditions and allows estimates of the altitude of the column, which are useful for three-dimensional trajectory models of ash-cloud transport. It also allows an estimate of the eruption rate based on the ash-column height. However, such estimates may be lower than the true values because the very top of the eruption column, which may not contain coarse ash, may not be detected by the radar.

We conclude that the radar detects mainly ash particles, sized from about 1 millimeter to a few centimeters on the basis of three sources—the brief duration of the reflected radar signal, data from independent ground observations on the mass and size of particles which fell out of the reflected cloud, and the intensity of the reflected signal. The most intense reflections come from ash clouds with particles that range from 2 to 20 mm in diameter and with a total particle mass concentration of less than .01 to 1 g/m³.

The data are useful for constructing models of ash columns and deposition of coarse tephra. Radars are not useful for long-term volcanic cloud tracking because the large ash particles, that provide for strong radar signals fall out soon after an eruption. The radar does not detect smaller (<1–100 μ m in diameter) ash particles in drifting volcanic clouds that can persist in the atmosphere for several days or more.

INTRODUCTION

Since 1986, three volcanoes (Augustine, Redoubt, and Mount Spurr) have erupted in the Cook Inlet area of Alaska (fig. 1). Each of the eruptions has had a considerable impact on commercial aviation in south-central Alaska, particularly in Anchorage. Consequently, the National Weather Service (NWS) has significantly improved its capability of measuring and tracking ash clouds, in order to advise the aviation community about how to avoid ash clouds, which can be hazardous to jet aircraft. This report summarizes and interprets radar observations of ash clouds, made in real time during the 1992 Crater Peak eruptions, by a leased WR100-2 EEC meteorological C-band radar operated at Kenai, Alaska, opposite the volcanoes (fig. 1).

The radar was installed just prior to the 1992 activity. Three eruptions of the Crater Peak of Mount Spurr volcano in 1992 provided a first, robust test of the radar. The results provide data that are useful for real-time ash-cloud tracking and for forecasting cloud movement. Communication between the NWS office, the U.S. Geological Survey (USGS), and the Federal Aviation Administration (FAA) at Anchorage provides for the exchange of data and for timely and effective dissemination of eruption information to the public.

The data collected are also of volcanological interest. The usefulness of weather radar systems at active volcanoes was demonstrated at Mount St. Helens (Harris and others 1981; Harris and Rose, 1983). The heights of ash columns, which were measured during the eruptions, were used to calculate the eruption rate of ash. The radar system also mapped the extent of the most reflective parts of the ash cloud as it moved across Washington State.

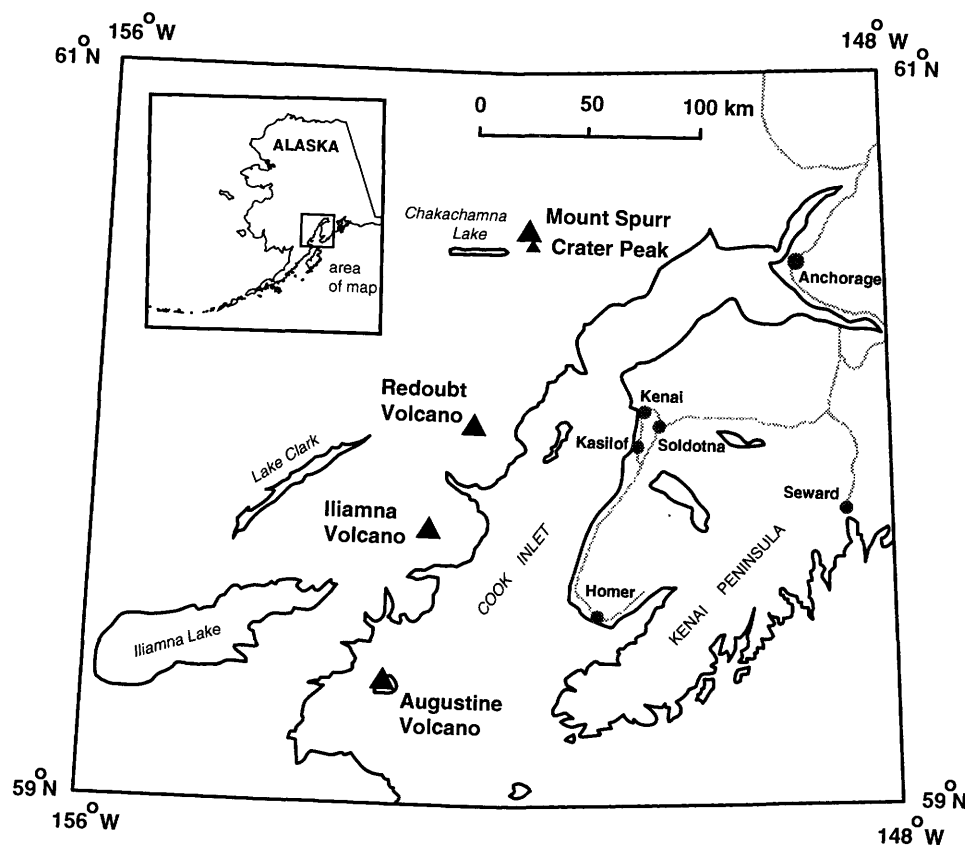


Figure 1. Index map of Cook Inlet area. The solid triangles represent volcanoes. Modified from Brantley (1990).

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THE 1992 CRATER PEAK ERUPTIONS

A summary of data on the three Crater Peak eruptions of 1992 is given in table 1. The events were similar in many respects. They each lasted 3 to 4 hours and were marked by andesitic ash eruptions, which

produced roughly similar volumes of new ash (Neal and others, this volume). Each eruption resulted in narrow fallout blankets, which stretched for many kilometers from the volcano, and clouds that were much larger than those produced by the 1986 eruptions of Augustine Volcano (Holasek and Rose, 1991) and the 1989–90 eruptions at Redoubt Volcano (Schneider and Rose, 1994).

DESCRIPTION OF THE RADAR SYSTEM

The Kenai radar is remotely operated by meteorologists stationed in the NWS forecast office in Anchorage, who communicate with the radar via modem. The main data displays are the range-height indicator (RHI), and the plan-position indicator (PPI). The RHI display is a vertical cross section of a reflecting cloud with a range and height scale (fig. 2). The PPI display is a horizontal cross section of the volcanic ash cloud on a scaled map background of Cook Inlet (fig. 2). The characteristics of the radar used are given in table 2. The radar's operation and performance were described in more detail by Serafin (1990).

Table 1. Data on 1992 Crater Peak eruptions, Mount Spurr volcano, Alaska.

[ADT, Alaska daylight time; AVO, Alaska Volcano Observatory; NWS, National Weather Service; *, data unavailable; Sources of AVO data were Steve McNutt, Game McGimsey, Christina Neal, and Thomas Miller (oral commun., 1993). Anchorage International Airport closed at 6:10 p.m. ADT on 8/18/92]

	6/27/92	8/18/92	9/17/92
AVO Observations			
Eruption start, ADT -----	07:04 a.m.	4:42 p.m.	12:03 a.m.
Eruption end, ADT -----	11:07 a.m.	8:10 p.m.	3:39 a.m.
Volume of ash ($\times 10^6 \text{ m}^3$) -----	12	14	15
NWS Radar Observations			
Eruption start, ADT -----	*	*	12:10 a.m.
Eruption end, ADT -----	*	6:13 p.m.	3:08 a.m.
Minimum volume of ash ($\times 10^6 \text{ m}^3$) ----	*	*	16
Maximum column height, km -----	14.5	13.7	13.9
Time of maximum height, ADT -----	10:23 a.m.	4:55 p.m.	2:21 a.m.
Anchorage International Airport Information			
Service from airport -----	open	closed	open
Aircraft rerouted -----	yes	yes	yes

Table 2. Specifications of WR 100-22 EEC meteorological C-band radar used at Kenai, Alaska.

Parameter	Specification for the Kenai Radar
Primary power -----	115 volts, 60 cycle, single phase
Wavelength -----	5 cm
Peak power -----	250 KW
Pulse length -----	2 microseconds
Pulse repetition frequency -----	256 Hertz
Magnetron type -----	7156 A
Radio frequency -----	5500 to 5600 Megahertz
Antenna beam width -----	1.6 degrees
Receiver minimum detectable signal ----	-106 dBm (minimum)
Receiver IF -----	30 Megahertz
IF bandwidth -----	log 2 Megahertz
Receiver dynamic range -----	78 dB minimum

The reflectivity of the volcanic clouds is measured by comparing them to a standard equivalent reflectivity expressed in dBZ, which stands for decibels relative to a rain target with a reflectivity factor of $1 \text{ mm}^6/\text{m}^3$, for example, $\text{dBZ} = 10 \log_{10} (Z/1 \text{ mm}^6/\text{m}^3)$; this allows inclusion of the natural range of meteorological precipitation clouds, within values of about 1 to more than 50 (Rinehart, 1991; Sauvageot, 1992). By choosing a range of dBZ values, contours of equal reflectivity are displayed on both the RHI and PPI images. Particle size and concentration may then be estimated from the measured reflectivities.

USE OF RADAR FOR MEASUREMENTS OF ERUPTING COLUMN HEIGHT

When operated in the RHI (range-height indicator) mode and pointed directly toward the volcano (azi-

muth of 328 in the case of Spurr), the radar can detect the top of the erupting column (fig. 3). Repeated measurements during the 3-hour eruption of September 16–17, 1992, give a record of the dynamics of the eruption of Crater Peak. The height measurements can be converted to eruption rates by the inversion of the Morton Equation (Morton and others, 1956; Wilson and others, 1978). Ignoring the effects of entrained atmospheric water vapor, we can then calculate the final height of a buoyant plume, given a constant eruption rate, to give an estimate of intensity of eruption with time (fig. 3). In simplified form, the Morton equation is $H = 236.6 Q^{1/4}$. H is the height of the erupting column in meters, and Q is the eruption rate in kilograms/sec. The constant, 236.6, includes an empirical constant in the correct dimensional form.

This measure of intensity compares well with the raw seismic signal measured during eruption, based

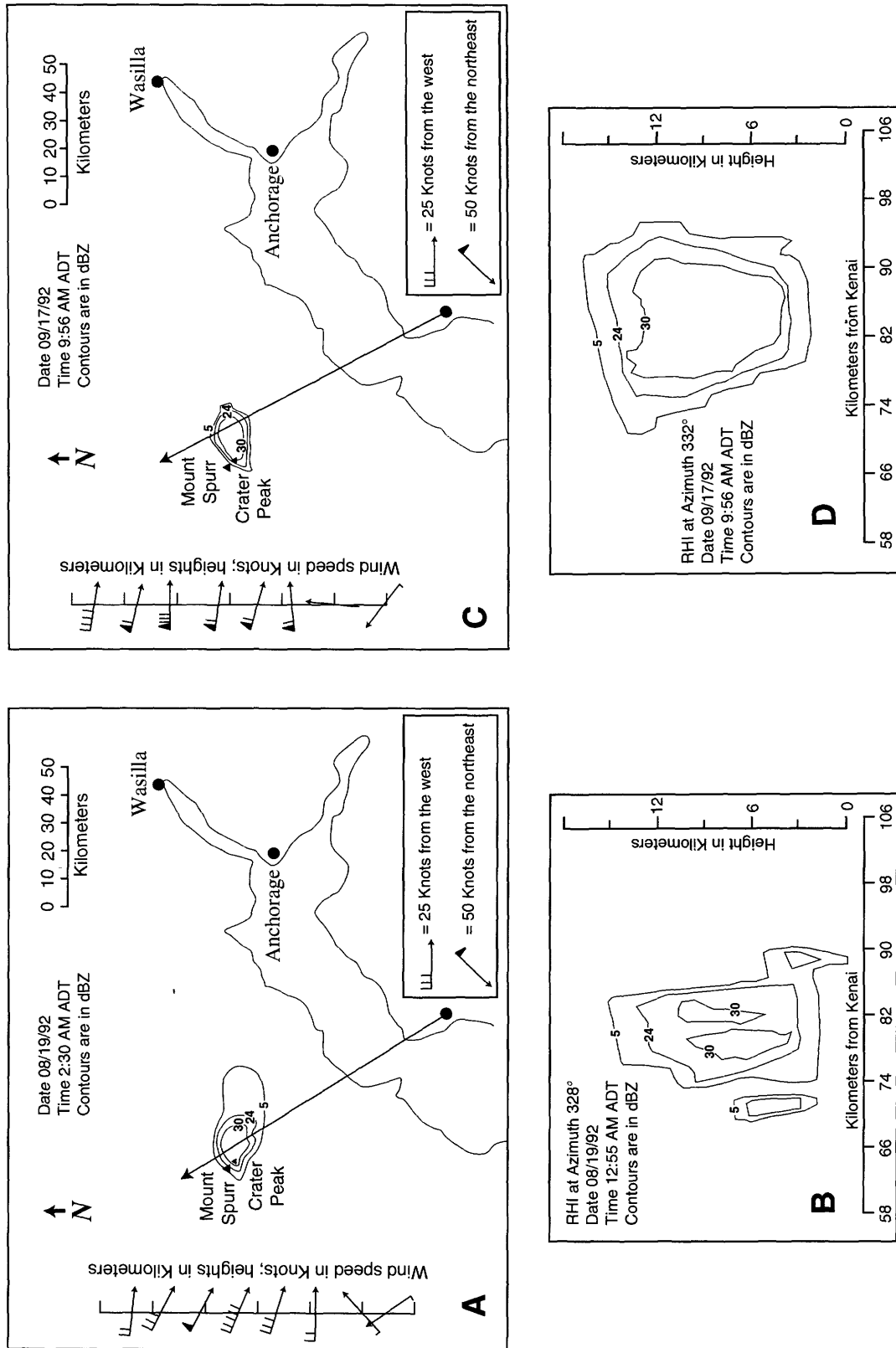


Figure 2. Data from C-band radar collected during the August (A and B) and September (C and D) Spurr eruptions. Times are listed in GMT. Both PPI and RHI formats are shown. The azimuth of the depicted RHI image is plotted on the PPI image. NWS upper-level wind data are plotted to the left of the PPI maps (altitudes shown in kilometers). The reflectivity of the volcanic clouds is measured by comparing them to a standard equivalent reflectivity expressed in dBZ, which stands for decibels relative to a rain target with a reflectivity factor of $1 \text{ mm}^6/\text{m}^3$.

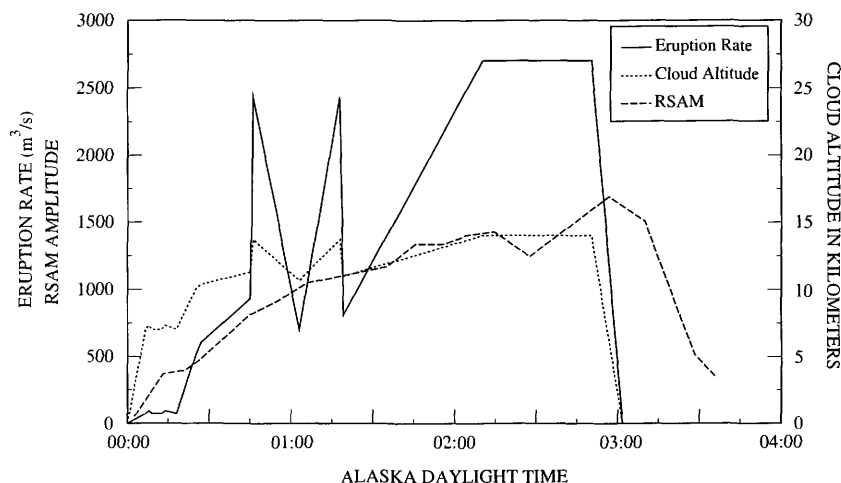


Figure 3. Plots of eruption cloud maximum height versus time for the September 17, 1992, Spurr eruption as measured by C-band radar from Kenai. RHI recordings, with azimuths of 328° – 332° (toward Crater Peak from Kenai), were used. A beam width correction was applied. Eruption rate versus time for the eruption is also plotted, derived from the radar height data and application of the Morton equation. RSAM seismic results from AVO (S. McNutt, oral commun., 1993), which are also measures of eruption intensity, are plotted for comparison. RSAM is a 1-minute average of seismic amplitudes from station BGL, 10 km from Crater Peak vent of Mount Spurr.

on real-time seismic amplitude measurement (RSAM) (Murray and Endo, 1992; fig. 3). The apparent eruption rate, estimated by the Morton equation and radar, can be integrated over the duration to get an estimate of the volume of ash emitted. The result is about $16 \times 10^6 \text{ m}^3$, which is almost the same as the amount estimated from ash fallout ($15 \times 10^6 \text{ m}^3$). However, the radar could underestimate the column height because larger particles do not rise as high as smaller, less reflective ones (Carey and Sparks, 1986).

In summary, the radar offers a method of measuring the height of the eruption column, which is important for trajectory models of ash cloud movements. The systematic measurement of heights of the column is a way to monitor the eruption intensity and also gives a minimum estimate of mass of ash erupted.

SIZE OF ASH PARTICLES DETECTED BY RADAR

The intensity of radar reflections is described by the Probert-Jones equation (Probert-Jones, 1962), which can be expressed in simplified form as $P_r = R_c K^2 Z/R^2$. P_r is the radar echo power received, and R_c is a radar constant containing the wavelength relations specific to the radar (peak power, antenna gain, beam width, and pulse length). R is the range of the

target. K^2 is a refractive index factor, and Z is the target reflectivity factor, which is proportional to the product of the number density of reflecting particles and their radii, raised to the sixth power. Because the radar echoes are so strongly dependent on the size of the particles (r^6 in the small dipole or Rayleigh approximation), the return signal detected by the radar is dominated by larger ash particles suspended in the cloud.

We observed that the radar system detected ash clouds only near the volcano and within about one half hour or less of actual eruption. The drifting cloud remains highly reflective to radar for a distance of only about 15 to 25 km from the volcano (fig. 2), which at the rate of winds measured, represents a transport time of about 15 to 30 minutes. We used data from figure 4 (Lapple, 1974) to estimate the size of ash particles that would fall to the ground in less than one half hour from heights of 14.2 km (47,000 feet) or less—time and height constrained by the Crater Peak radar observations. On the basis of these estimates, spherical particles with diameters of 1,000 microns (1 mm) or more would have terminal velocities (600 cm/sec) that would cause them to fall out within a half hour from 14.2 km. Because of the particles' extremely size-sensitive reflectivities and the rapid disappearance of the clouds from the radar, we think that the radar reflections we measured were caused by particles with diameters of 1,000 microns (1 mm) or larger.

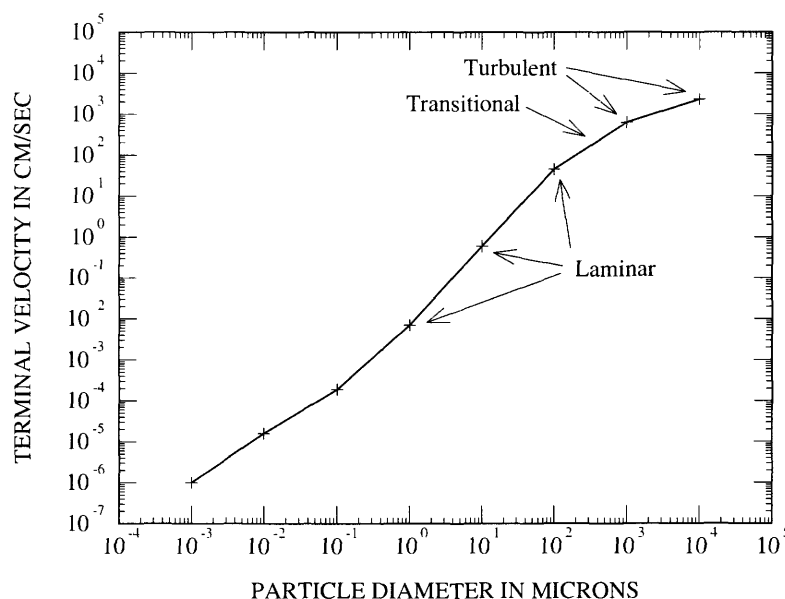


Figure 4. Terminal gravitational settling velocities of various size spheres with a density of 2.0 g/cm^3 in dry atmosphere at 25°C and 1 atm (Laple, 1974).

A cross section of the cloud that was about 7 km downwind from the crater is shown in figure 3, and this cloud section shows very strong reflections (30 dBZ). A radar reflection is caused by the summation of all of the particles multiplied by their individual reflectivities. Consequently, a large number of small particles can have a reflectivity equal to a small number of large ones. Therefore, other data are necessary to determine the characteristics of the cloud.

In this case, we have excellent additional data: the time it takes for the radar image to disappear and the size and mass of the particles that fell out of the section imaged. These radar reflections cannot be produced from very large ash particles (greater than 3 cm), because their higher terminal velocities would cause them to fall out before drifting 7 km. On the basis of prevailing winds at appropriate altitudes (about 50–60 knots or 80–96 km/hr) at the time, we estimate that the cloud 7 km downwind is about 5 minutes old. Spheres with a density of 2 g/cm^3 , which were larger than about 3 cm, would be completely removed from a 14-km-high cloud within 5 minutes. We should expect that ash particles, which are not spherical, will have greater atmospheric drag and, therefore, slower terminal velocities, but particles with diameters much larger than a few centimeters are not likely to be present in the cloud section shown in figure 2.

We compared the data on radar reflectivity of volcanic ash (Harris and Rose, 1983; Rose and Kostinski, 1994) to see what combinations of particle mass and size could explain the observed intensity of

radar reflection. The theoretical reflectivity results, calculated for identical, spherical particles with a volcanic ash refractive index factor (K^2) of 0.3, which is about one-third that of liquid water (0.93), are shown in figure 5. The maximum reflection, seen in the drifting cloud 7 km downwind (30 dBZ), corresponds to more than 100 g/m^3 of ash particles with a radius of 0.25 mm, 20 g/m^3 of ash particles with a radius of 0.5 mm, 0.7 g/m^3 of ash with a radius of 1 mm, or less than 0.01 g/m^3 with a radius of 4 mm. Because the radar reflection shown in figure 2 is thousands of meters high, a particle-mass concentration of 20 g/m^3 implies a fallout thickness of much more than 1.5 m. Consequently, the mass concentrations of 20 g/m^3 or higher are impossible because the observed fallout data are inconsistent with them.

The isopach maps of fallout on September 17, 1992, show that the thickness 7 km down the dispersal axis is about 10 cm, and the grain size is mainly between 1 and 30 mm (R.G. McGimsey, oral commun., 1993). This data, along with our observation that the radar signal disappeared after 30 minutes, implies that most of the particles which reflected the radar signal were 2 to 20 mm in diameter and the particle concentrations of the reflecting cloud ranged from less than 0.01 to about 1 g/m^3 (shown by the shaded area of fig. 5). In comparison, Harris and Rose (1983) estimated that the particle-mass concentrations for the Mount St. Helens eruption clouds of May 18, 1980, and March 19, 1982, were 3 g/m^3 and 0.2 g/m^3 respectively, similar to the Crater Peak estimates.

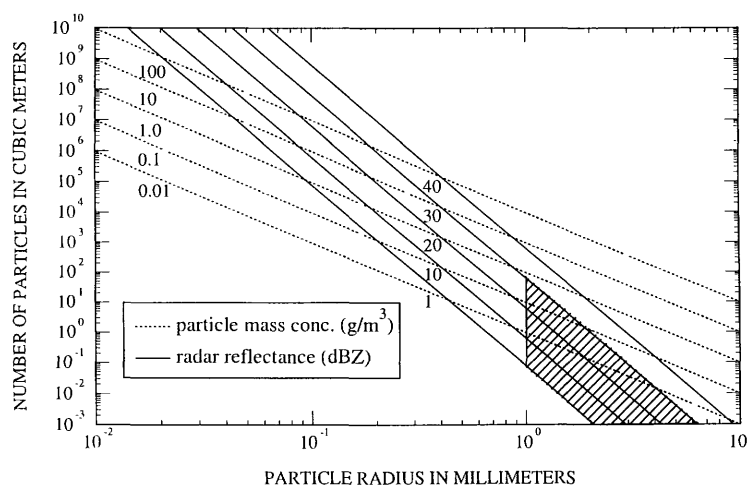


Figure 5. Relation between radar reflection returned (measured in decibels relative to a rain target or dBZ), the total mass of ash particles per unit volume (measured in grams per cubic meter), the number of particles per cubic meter, and the radius of the particles in mm. This figure is based on idealized assumptions: single size for particles, perfect spherical shape of all particles, and a refractive index factor which is about one-third of that for water droplets. Cross-hatched portion probably most closely approximates conditions in the 1992 eruption clouds from Crater Peak of Mount Spurr, which were detected by the C-band radar. This approximation is based on the duration of radar reflection and observed size of fallout.

OPTIMAL USE OF RADAR SYSTEMS DURING AN ERUPTION

Aircraft traveling at high speeds must get timely warning of ash clouds. During eruptions, radar measurements of ash can be directly communicated from the NWS to the Anchorage Air Route Traffic Control Center (ARTCC) so that aircraft in immediate danger can be diverted and decisions on possible airspace closures can be made.

Because the returns of the radar signal come mainly from large (1–30 mm) ash particles, the radar measurements can only be made during and immediately following actual eruptions. At the beginning of the eruption, the radar in the RHI mode makes initial detection and height measurements.

During the September 16–17, 1992, eruption of Crater Peak vent, the radar was able to verify the Alaska Volcano Observatory's (AVO) seismic indications of activity at Crater Peak within a few minutes. The eruption began at 10:36 a.m. ADT on September 16, and the initial activity lasted 16 minutes. At 12:03 a.m. ADT on September 17, AVO called NWS alerting them that an eruption was likely in progress. At 12:10 a.m. ADT, the Kenai radar verified a volcanic ash cloud above Mount Spurr. Later that morning, at just after 3:00 a.m. ADT, AVO called NWS with a notification

that the eruption had ended. During the call from AVO, the volcano was scanned, and the ash cloud observed by the radar was detached from the volcano. Within 20 minutes, the radar could detect no ash cloud.

Radar detection is valuable in many ways. Radar-based detection is especially important during conditions of poor visibility, when other types of observations are difficult. Radar height measurements of the ash cloud are inputs to the production of a three-dimensional ash dispersion tracking model, which the NWS is implementing to forecast volcanic cloud positions. In cooperation with AVO, the NWS works with the Alaska Division of Emergency Services to identify those areas of the state that will most likely be affected by ash fallout.

A new generation WSR-88D radar system was installed in 1994 at Kenai, Alaska. The WSR-88D is a modern S band ($\lambda = 10\text{--}11\text{ cm}$), doppler radar, which could provide some advantages for volcanic cloud mapping, including information on radial velocities in and around the ash cloud.

CONCLUSIONS

The 5-cm radar operated from Kenai was successfully used to map ash clouds from Crater Peak about 80 km away. The radar signal is dominated by

ash particles larger than about 1,000 microns in diameter. We were able to use ground measurements of the sizes and masses of fallout particles that fell from the ash cloud as well as the dynamics of fallout constrained by the duration of the radar reflection to tell us that the largest radar reflections were mainly caused by ash particles that were between 2 and 20 mm in diameter.

Radar can also be used to measure the height of the ash cloud, but it may underestimate the height because it does not respond to smaller ash particles, which rise higher. C-band radar does not detect distal parts of the ash cloud, which have an atmospheric residence time longer than 30 minutes, because the larger more reflective ash particles drop out.

The Kenai radar is the first type of remote sensor to actually view ash clouds. Its all weather capability fills the data gaps of satellite data as well as other observational methods and the real-time use of a radar provides an excellent way to corroborate the seismic indications of actual eruptions. Radar provides vital, immediate warning and information to the aviation community.

The radar system offers, for the first time, the opportunity of mapping internal structure of dense proximal ash plumes during eruptions. This capability can supply data to volcanologists who are interested in modeling eruption columns (Wilson and others 1978; Carey and Sparks, 1986; Bursik and others, 1992). Because of the radar's sensitivity to larger particles, such a system would be particularly valuable in testing models for the fallout of larger particles, such as the one developed by Bursik and others (1992).

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Tracking of 1992 Eruption Clouds from Crater Peak Vent of Mount Spurr Volcano, Alaska, Using AVHRR

By David J. Schneider, William I. Rose, and Lee Kelley

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ABSTRACT

The 1992 Crater Peak vent eruptions of Mount Spurr provided a robust test of the new High Resolution Picture Transmission (HRPT) Information Processing System (HIPS), which was installed in Anchorage for real-time tracking of volcanic clouds. One principle objective of the tracking system is to avoid aviation hazards. The system receives Advanced Very High Resolution Radiometer (AVHRR) images from polar-orbiting satellites and processes them using a two-channel technique. The resulting processed imagery reflects the extensive real-time capability now available. The frequency of data collection (about once every few hours) allows for an understanding of the patterns of change in spectral response that occur during the evolution of a volcanic cloud. Volcanic clouds imaged during and shortly after eruption are optically thick and can contain abundant water droplets and (or) ice, and these characteristics cause their spectral signal to closely resemble a meteorological cloud. As the volcanic cloud disperses, the spectral properties of the cloud change, first at the edge and then throughout. These changes produce a volcanic cloud signal that can be distinguished by using a brightness temperature difference determined from thermal bands 4 and 5 of the AVHRR. The usefulness of this technique of tracking clouds for long distances was investigated using archived data. Two clouds were tracked for more than 80 hours and thousands of kilometers, during which the band 4 minus 5 brightness tempera-

ture signal decayed. These clouds traveled southward over the continental United States and posed a potential hazard to aircraft. To alleviate this potential hazard, the measurable brightness temperature difference signal needs to be correlated with ash concentration and particle size to determine when drifting volcanic clouds are a threat to aircraft.

INTRODUCTION

Since 1986, three volcanoes (Augustine Volcano, Redoubt Volcano, and Mount Spurr) have erupted in the Cook Inlet area of Alaska. Each of these eruptions has had a significant impact on commercial aviation at Anchorage. The National Weather Service has greatly improved its capability to measure and track ash clouds in order to better advise the aviation community about the location of hazardous ash clouds. Two significant improvements, which were made following the Redoubt activity in 1989-90, were the installation of a High Resolution Picture Transmission (HRPT) Information Processing System (HIPS) workstation for rapid digital-image processing of Advanced Very High Resolution Radiometer (AVHRR) data from polar-orbiting weather satellites, and the leasing of a meteorological C-band radar. Weather satellite data have been used by many investigators to observe ash clouds (Hanstrum and Watson, 1983; Matson, 1984; Sawada, 1987; Prata, 1989; Holasek and Rose, 1991; Schneider and Rose, 1994) but until now, real-time analysis of processed images has been difficult.

The AVHRR data complement the radar observations because the radar can map clouds near the volcano, within minutes of eruption (Rose and others, this volume), while the AVHRR can track the clouds for up to several days. Three eruptions of the Crater Peak vent of Mount Spurr provided a real-time test of the equipment's ability to detect and map significant ash clouds. The results obtained with the new HIPS workstation equipment, and the results obtained on the long-range tracking of the cloud over North America using archived AVHRR data are described herein.

ACKNOWLEDGMENTS

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TWO-CHANNEL AVHRR DETECTION OF VOLCANIC CLOUDS

AVHRR satellite data are received from the NOAA 11 and 12 polar-orbiting satellites at the National Weather Service Forecast Offices in Anchorage, Alaska, through the HIPS system and is linked to meteorologist workstations in forecast offices of the Alaska Region Operations Network. Full passes (4.4-km resolution) over active volcanoes are available as soon as 10 minutes after pass completion, with subsectors (1.1-km resolution) available within 40 minutes. All five spectral bands, from visible to thermal infrared, are received (band 1: 0.58–0.68 μm ; band 2: 0.73–1.10 μm ; band 3: 3.55–3.93 μm ; band 4: 10.3–11.3 μm ; and band 5: 11.5–12.5 μm). Satellite coverage extends from the western edge of the Kamchatka Peninsula to east of British Columbia; this coverage encompasses four major air routes near the volcanoes of Alaska and Kamchatka. Because of orbital irregularities and the movement of volcanic clouds, coverage frequency is variable and gaps as long as 8 hours can occur.

We used band 4 minus band 5 brightness temperature difference images to detect volcanic clouds, and to distinguish them from meteorological clouds. Thermal data from bands 4 and 5 were converted from raw sensor counts to radiance, using the technique of Lauritson and others (1988). A slight nonlinearity in bands 4 and 5 was corrected using a quadratic function of radiance. The radiance values were converted into brightness temperature values, using the inverse Planck function, and band 4 minus 5 brightness temperature images were created. Volcanic clouds are known to have negative band 4 minus 5 brightness temperature differences (Prata, 1989; Schneider and Rose, 1994), whereas meteorological clouds generally

have positive brightness temperature differences (Yamanouchi and others, 1987). For volcanic clouds, the magnitude of the negative brightness temperature difference depends on many characteristics. These characteristics include the optical thickness of the cloud; amounts of water, volcanic ash, and sulfuric acid in the cloud; mean size and size distribution of particles in the cloud; and temperature contrast between the cloud and the underlying surface (meteorological clouds, land, or water) (Prata, 1989; Wen and Rose, 1994).

The volcanic cloud signal detected by the brightness temperature difference method has well-defined edges, which allow us to define the area of the cloud. Cross sections through volcanic clouds show a clearly defined edge at a brightness temperature difference value of -0.5°C (fig. 1). The ability to detect cloud edges is important in reducing aircraft hazards, because it is a snapshot of cloud position that can be quickly reported. It also serves as a periodic check on the trajectory models for the ash cloud.

Table 1. List of AVHRR images used for this study showing date and time of satellite coverage, hours since start of eruption, and satellite used.

Time (UT)	Hours since start of eruption	Satellite
August 19, 1992		
01:26	1.5	NOAA A-11
03:38	3.6	NOAA A-11
05:10	5.2	NOAA A-12
13:38	13.6	NOAA A-11
17:20	17.3	NOAA A-12
18:57	18.9	NOAA A-12
23:30	23.1	NOAA A-11
August 20, 1992		
13:23	37.4	NOAA A-11
23:15	47.3	NOAA A-11
August 21, 1992		
23:00	71.0	NOAA A-11
August 22, 1992		
11:21	83.4	NOAA A-11
September 17, 1992		
12:40	3.7	NOAA A-11
17:00	8.0	NOAA A-11
22:44	13.7	NOAA A-11
September 18, 1992		
11:00	26.0	NOAA A-11
20:45	35.7	NOAA A-11
September 19, 1992		
09:00	49.0	NOAA A-11
18:53	57.9	NOAA A-11
September 20, 1992		
07:00	70.0	NOAA A-11
17:04	80.1	NOAA A-11

OBSERVATIONS OF THE CRATER PEAK VOLCANIC CLOUDS

Eleven images of the volcanic cloud resulting from the August 18, 1992, eruption of Crater Peak and nine images of the volcanic cloud from the September 16–17, 1992, eruption were analyzed for this study (table 1). The evolution of the volcanic cloud signal can be seen in figure 2, a series of images col-

lected by the HIPS system of the August 18 volcanic cloud. The earliest image, collected about 90 minutes after the start of the eruption (fig. 2A), shows a cold, circular volcanic cloud in band 4 and high thermal contrast between the volcanic cloud and the lower, warmer meteorological clouds. The band 4 minus band 5 brightness temperature difference algorithm does not work well on this image (fig. 2B). The second image in this series was collected 2 hours later and shows a

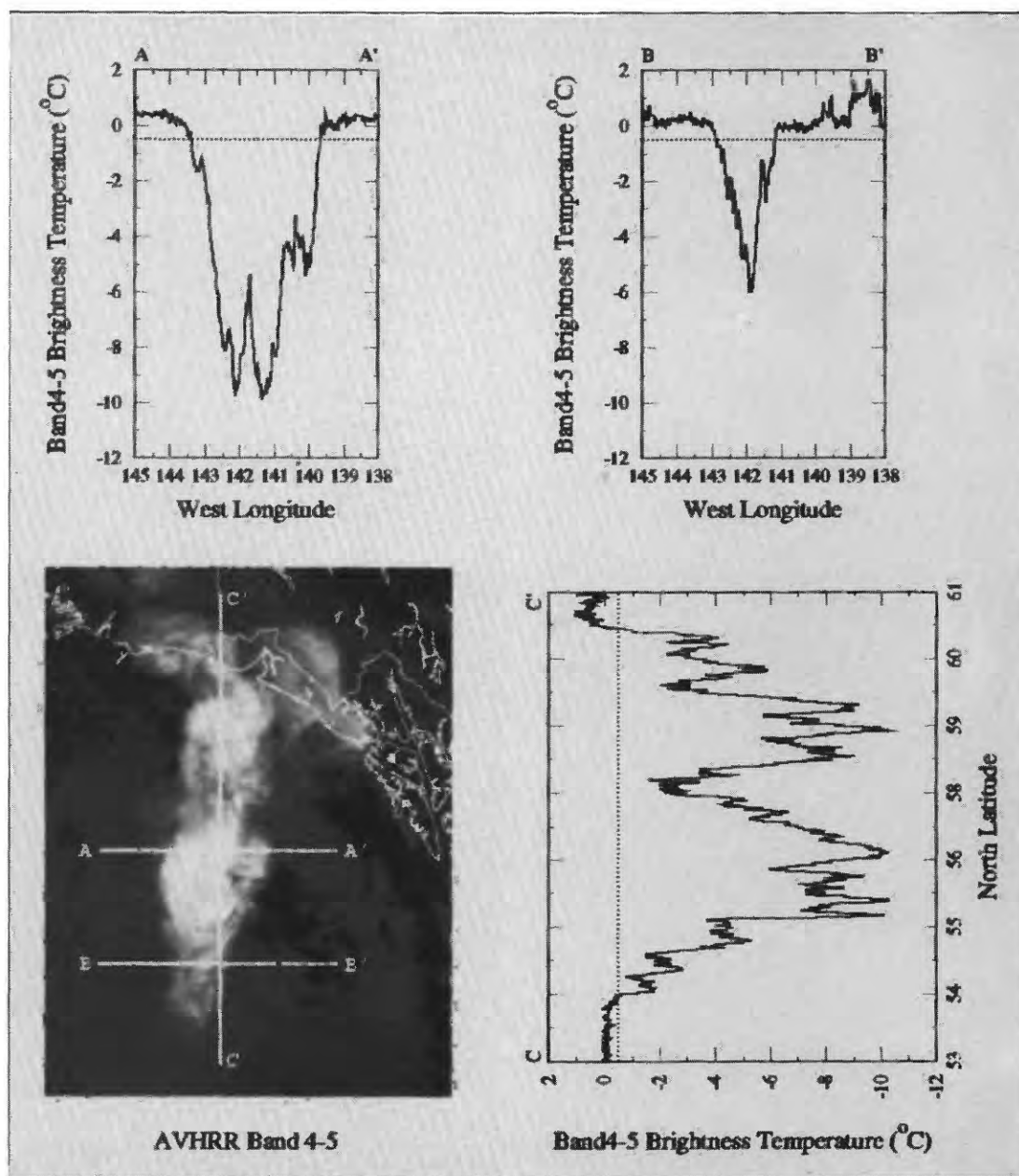


Figure 1. Cross sections through a band 4 minus band 5 image that compare brightness temperature difference of a volcanic cloud to those of meteorological clouds. Interior volcanic cloud shown has negative band 4 minus band 5 brightness temperature differences as large as -14°C , whereas meteorological clouds, land, and open water generally have brightness temperature differences greater than zero. Dotted line at -0.5°C shown for reference.

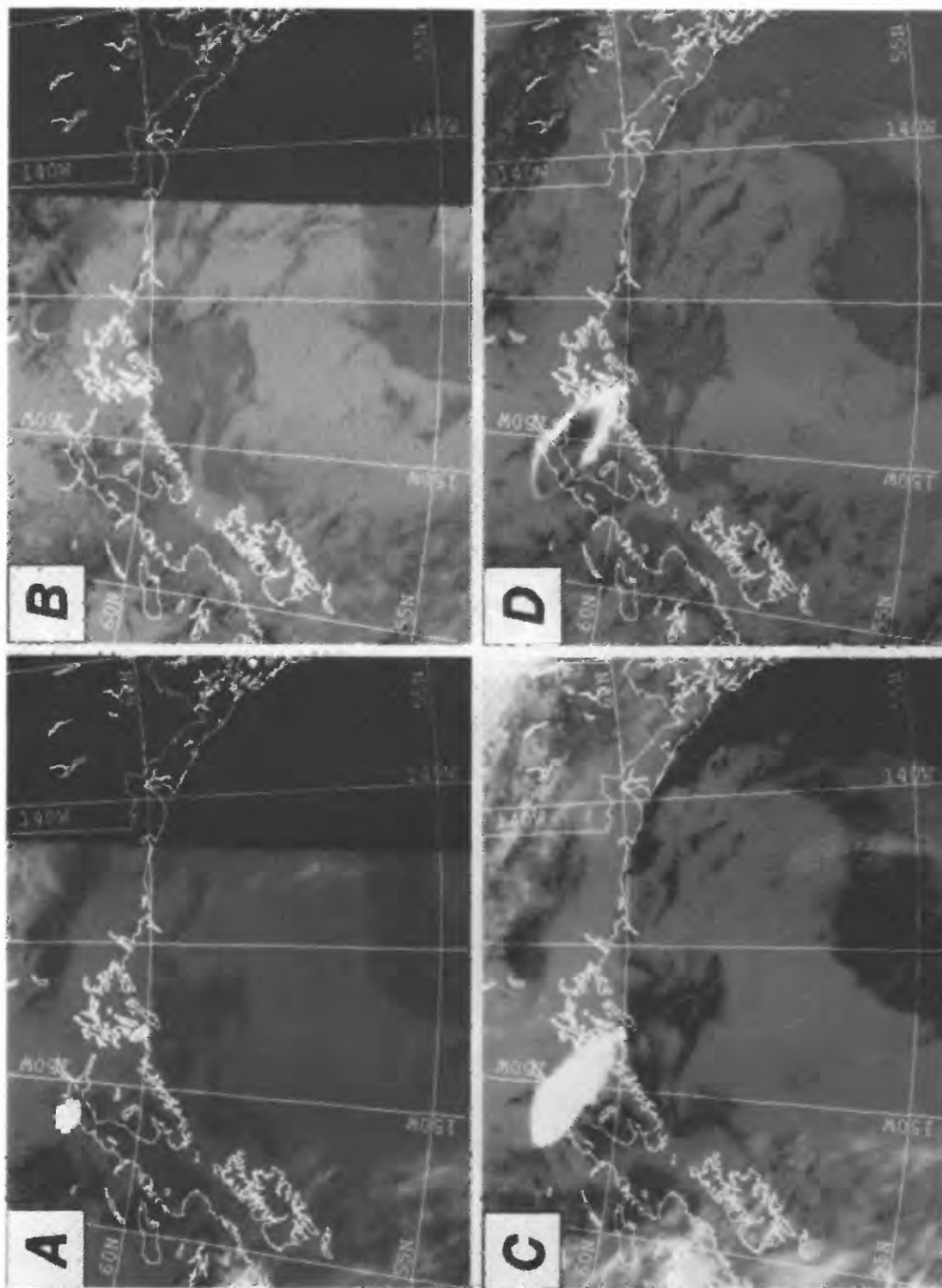


Figure 2. Sequence of four images received from the HIPS system of volcanic cloud from August 18, 1992, eruption of Crater Peak vent, Mount Spurr, Alaska. Sequence shows evolution of the volcanic cloud. See text for definition of HIPS and AVHRR. In the band 4 images, cold features are bright and warm features are dark. In the band 4 minus band 5 images, negative brightness temperatures are bright and positive brightness temperatures are dark. A, AVHRR band 4 image at 5:26 p.m. ADT (August 19, 01:26 UT). B, Corresponding band 4 minus band 5 brightness temperature image at 5:26 p.m. ADT (August 19, 01:26 UT). C, AVHRR band 4 image at 7:38 p.m. ADT (August 19, 03:38 UT). D, Corresponding band 4 minus band 5 brightness temperature image at 7:38 p.m. ADT (August 19, 03:38 UT).

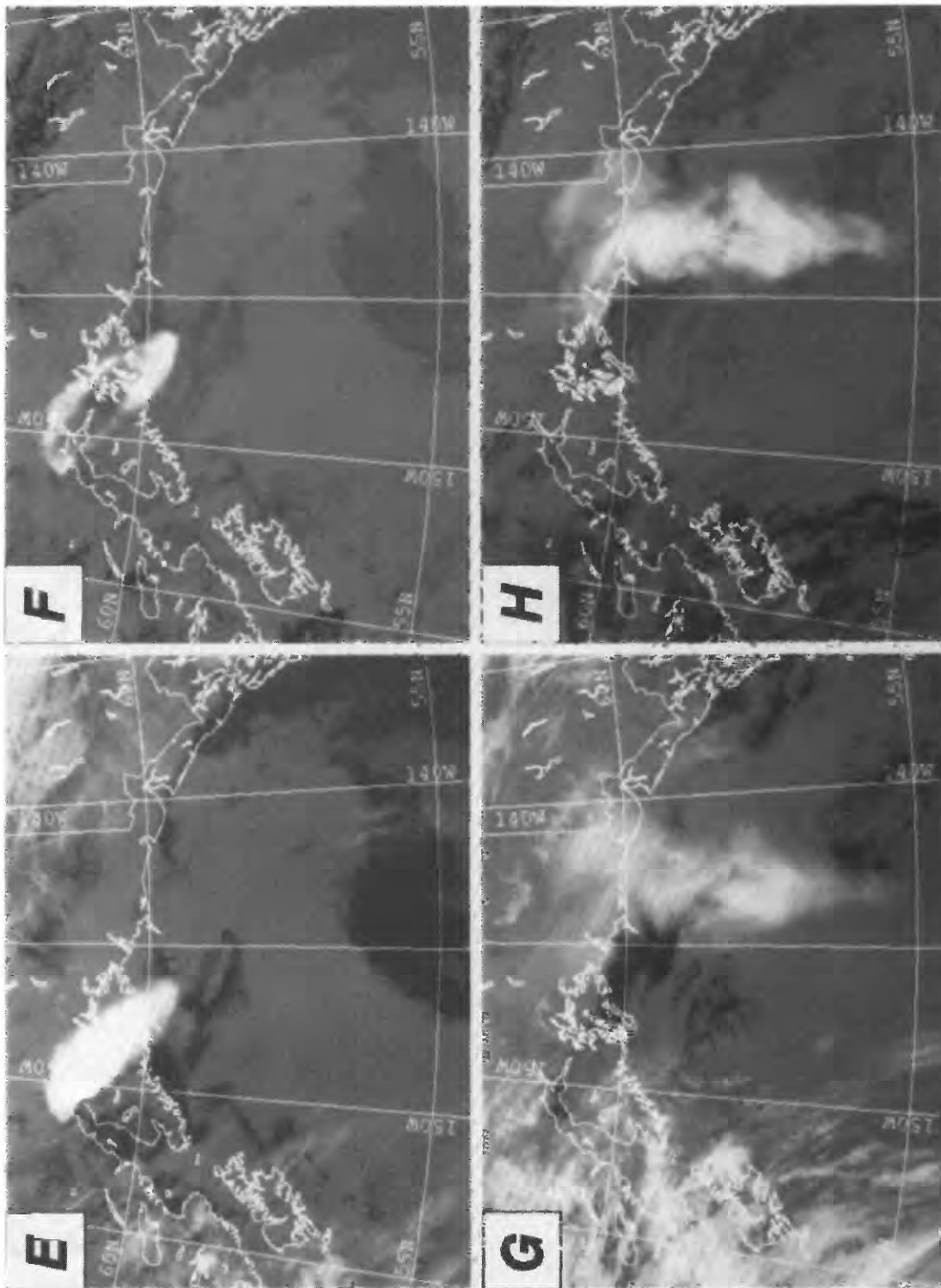


Figure 2. Continued. *E*, AVHRR band 4 image at 9:10 p.m. ADT (August 19, 05:10 UT). *F*, Corresponding band 4 minus band 5 brightness temperature image at 9:10 p.m. ADT (August 19, 05:10 UT). *G*, AVHRR band 4 image at 5:38 a.m. ADT on August 19 (13:38 UT). *H*, Corresponding band 4 minus band 5 brightness temperature image at 5:38 a.m. ADT on August 19 (13:38 UT).

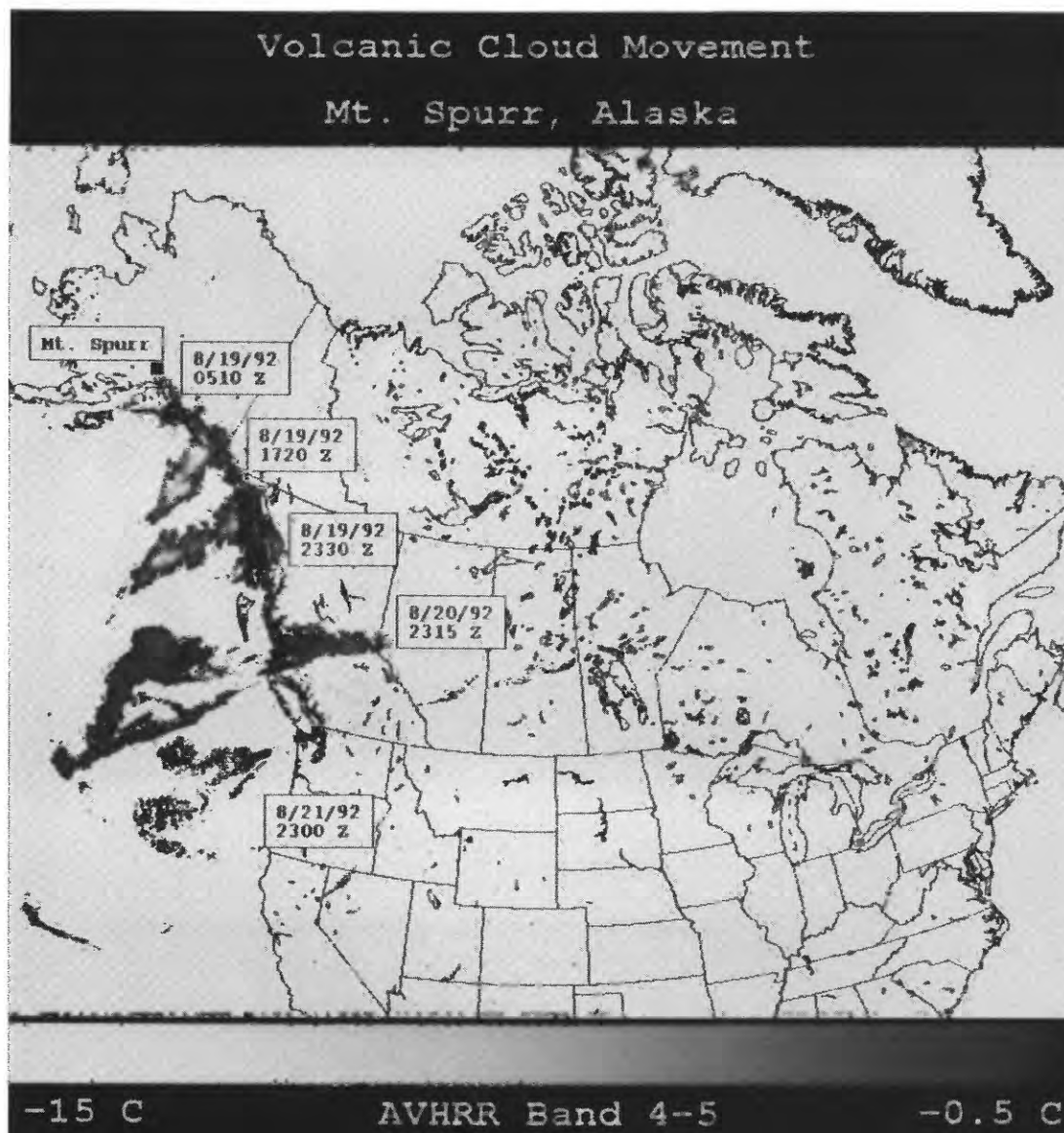


Figure 3. Composite of five band 4 minus band 5 images, showing the long-distance transport of the August 18, 1992, volcanic cloud from the Crater Peak vent of Mount Spurr, Alaska. This cloud was detected in 11 images, but for clarity, only five images are shown. Alaska daylight time equals Z (same as UT) - 8 hours.

larger cloud in the band 4 image (fig. 2C) and a fringing response from the band 4 minus band 5 process (fig. 2D). The third image (fig. 2E) was collected 90 minutes later. By this time the eruption had ended, and the fringe of negative band 4 minus band 5 values had grown and encircled the entire volcanic cloud (fig. 2F). The fourth image was collected 8 hours later. The volcanic cloud is difficult to distinguish from the meteorological clouds in the band 4 image (fig. 2G), but it is easily distinguished on the band 4 minus 5 image (fig. 2H) in which the entire cloud shows negative brightness temperature difference values.

The utility of the band 4 minus band 5 operation to track clouds for days following an eruption was tested using two archived AVHRR data sets. A composite of five images of the August volcanic cloud is shown in figure 3, and a composite of nine images of the September volcanic cloud is shown in figure 4. In these images, only the pixels with negative band 4 minus band 5 brightness temperature values are shown. Because many factors affect the magnitude of the negative brightness temperature difference, the regions of the cloud with the greatest temperature difference do not necessarily have the highest concentration of particles.

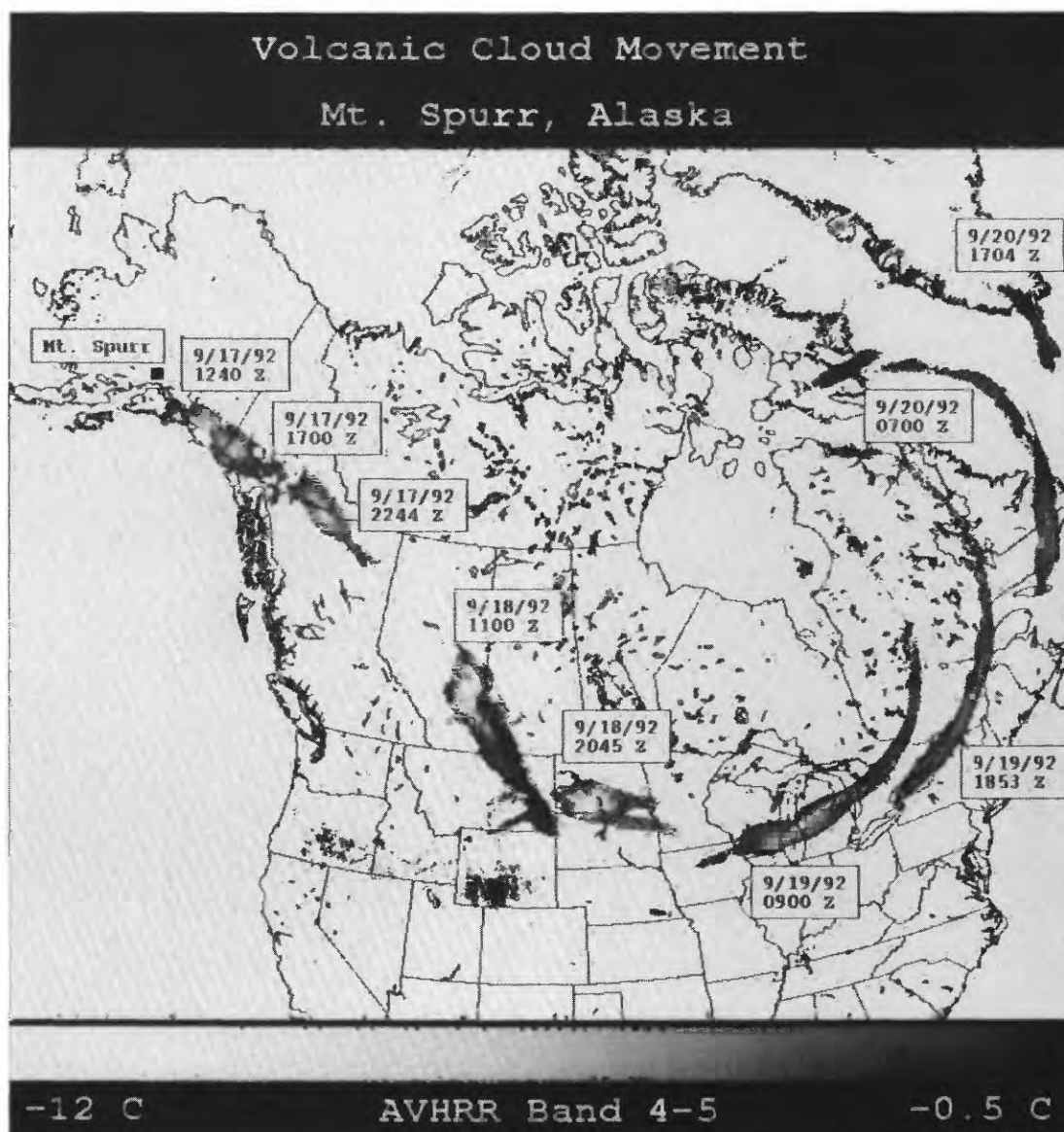


Figure 4. Composite of nine band 4 minus band 5 images, showing the long-distance transport of the September 16-17, 1992, volcanic cloud from the Crater Peak vent of Mount Spurr, Alaska. Alaska daylight time equals Z (same as UT) - 8 hours.

The detected area of the August cloud increased during the first 47 hours at a rate of 13,000 km²/hr and reached a maximum area of about 600,000 km². The detected area of the September cloud increased at an average rate of 12,000 km²/hr during the first 26 hours and reached a maximum area of 310,000 km² (fig. 5A). The leading edge of the August cloud moved about 3,000 km in 83 hours for an average velocity of 36 km/hr, and the leading edge of the September cloud moved about 4,400 km in 36 hours for an average velocity of 120 km/hr (fig. 5B). For both clouds, the magnitude of the negative band 4 minus band 5 brightness temperature difference first increased and then decreased during the cloud evolution (fig. 5C).

DISCUSSION

The sequence of images of the Crater Peak volcanic clouds defined the evolution of the spectral properties of volcanic clouds better than any previous observations. Volcanic cloud optical thickness is a critical parameter in distinguishing volcanic clouds. Opaque volcanic clouds (optical thickness = ∞) have positive band 4 minus band 5 brightness temperature values, and transparent volcanic clouds have negative band 4 minus band 5 brightness temperature values (Prata, 1989). The transition from the opaque stage to the transparent stage is seen in images of the August 18, 1992, cloud (fig. 2).

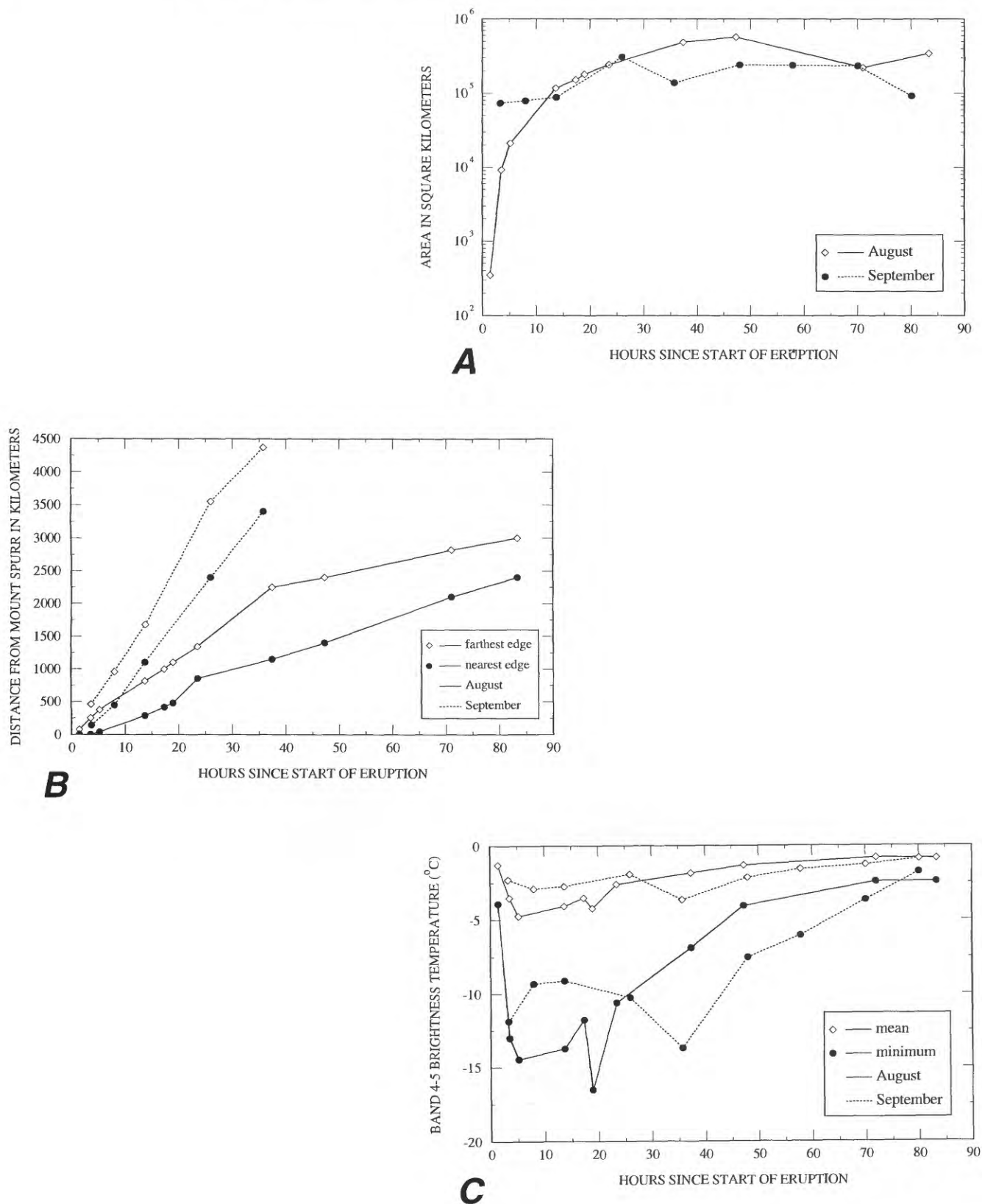


Figure 5. Changes, which occurred during transport, in the August 18 and September 16–17 volcanic clouds from the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska. *A*, Volcanic cloud area. *B*, Distance cloud was transported from the volcano. *C*, Brightness temperature differences calculated from bands 4 and 5 of AVHRR. Mean temperature of entire cloud and minimum value of each cloud are shown. See text for definition of AVHRR.

Initially, the volcanic cloud is infused with water vapor from entrainment of moist lower tropospheric air. The formation of water droplets and ice crystals produce an opaque cloud (fig. 2A), which is spectrally similar to meteorological clouds in the band 4 minus 5 image (fig. 2B). As the cloud ages, it disperses along its perimeter to produce a transparent fringe. The transparent portion of the cloud has negative band 4 minus band 5 brightness temperature values, which are depicted as bright white regions, while the opaque core has positive values, which are depicted as dark gray regions (fig. 2D). The fringe of negative values grows and the core of positive values shrinks as dispersion proceeds (fig. 2F). This continues until the entire cloud is transparent and has negative band 4 minus band 5 values (fig. 2H). As the opaque area of the cloud decreases, the magnitude of the brightness temperature difference increases and then decreases as the cloud becomes more transparent over time (fig. 5C).

Because the band 4 minus band 5 algorithm only discriminates transparent volcanic clouds, the magnitude of the brightness temperature difference and the area of the detected cloud are influenced by the temperature of the surface underlying the volcanic cloud. The detected area of the volcanic cloud becomes discontinuous, and the magnitude of the brightness temperature difference decreases when it overlies cold meteorological clouds (see fig. 3, 8/21/92 at 23:00 UT). However, the magnitude of the temperature difference increases when a cloud overlies a warm land surface (see fig. 4, 9/18/92 at 20:45 UT). The brightness temperature difference continues to decrease until the particle concentrations are so low that the cloud is undetectable by this technique. Because several factors affect the discrimination of volcanic clouds, the minimum detectable particle concentration cannot be directly related to brightness temperature difference values.

The signal detected with the band 4 minus band 5 brightness temperature difference needs to be linked to aircraft hazard, possibly through a radiative transfer model, which can quantify the brightness temperature difference values and estimate particle size and burden (Wen and Rose, 1994). Validation of radiative transfer models could be accomplished by direct sampling of a cloud throughout its evolution.

CONCLUSIONS

AVHRR data were successfully analyzed to detect and track the position of volcanic clouds during the Crater Peak eruptions of 1992, using a new capability of the National Weather Service in Anchorage. By detecting the volcanic clouds for longer periods of time, AVHRR complements the C-band radar sys-

tem employed by the NWS to image the volcanic clouds near the volcano in their first 30 minutes. Analyses of archived AVHRR data show that the volcanic clouds can be detected and tracked for several days following their eruption and throughout several thousands of kilometers of transport.

An algorithm based on a brightness temperature difference from the thermal infrared bands 4 and 5 of the AVHRR was used. A detailed sequence of images from the August eruption defined the evolution of the spectral properties of volcanic clouds better than any previous observations. The brightness temperature difference signal changes as the cloud evolves. It is more enhanced in the first few hours after the cloud forms and becomes transparent, and then it slowly decays to background values over the next few days. Because the magnitude of the brightness temperature difference is also affected by the temperature of the surface beneath the volcanic cloud, regions with the greatest temperature difference do not necessarily have the highest particle concentration.

Two of the Crater Peak clouds dispersed in the direction of the continental United States, raising the issue of whether Alaskan eruptions could pose hazards to principal domestic air routes in North America. Because the volcanic clouds were readily detectable when they crossed the continental United States, real-time processing of AVHRR data could have provided information to air traffic control outside Alaska. We need to calibrate the temperature difference signal to understand when volcanic clouds may be hazardous.

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Explosive Emissions of Sulfur Dioxide from the 1992 Crater Peak Eruptions, Mount Spurr Volcano, Alaska

By Gregg J.S. Bluth, Courtney J. Scott, Ian E. Sprod, Charles C. Schnetzler, Arlin J. Krueger, and Louis S. Walter

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ABSTRACT

Sulfur dioxide clouds from three explosive eruptions of Crater Peak vent of Mount Spurr during the Summer of 1992 were detected, tracked, and measured by NASA's Total Ozone Mapping Spectrometer (TOMS). From the TOMS data we calculated that the August 18 eruption produced the largest sulfur dioxide emission of the three events at 400 ± 120 kilotons (kt) SO_2 , followed by the September 16–17 and June 27 eruptions (230 ± 70 kt and 200 ± 60 , respectively). The June SO_2 plume remained primarily over eastern Alaska and Canada's Yukon province, and it was observed by TOMS for 7 days. The August cloud drifted southeast along Canada's Pacific coast before being carried eastward, and after 8 days it was finally viewed passing over England. The September cloud traveled the most rapidly of the three: southeast across Canada and the northern United States, then northeast over Toronto, and finally over Greenland 5 days later. TOMS data for the August and September eruptions revealed that the two gas plumes increased in SO_2 amounts over the first 2 days—well after the eruptions had ceased. The TOMS data and other available evidence suggest that significant quantities of hydrogen sulfide, in addition to sulfur dioxide, may have been explosively outgassed by these two eruptions. The maximum injection altitudes of the ash clouds, matching of SO_2 cloud positions to known wind con-

ditions, and the dispersion rates of the gas clouds all suggest that at least half of the sulfur dioxide emitted by the Crater Peak eruptions reached the lower stratosphere.

INTRODUCTION

Mount Spurr, Alaska ($61^\circ 30' \text{N}$, $152^\circ 25' \text{W}$) had been inactive since 1953 but during the Summer of 1992 erupted violently three times and sent large amounts of ash and gas into the atmosphere. Because of prevailing meteorological conditions, each eruption cloud followed unique dispersal paths. The SO_2 produced by these three eruptions was observed by the Total Ozone Mapping Spectrometer (TOMS), on board NASA's Nimbus-7 satellite. We used the TOMS data to analyze the eruption clouds for their SO_2 contents, areal extents, and dispersion patterns.

The Nimbus-7 satellite is in a polar sun-synchronous orbit and crosses the equator every 26 degrees ($2,800$ km) of longitude at local noon, observing the whole Earth once a day (13.7 orbits per day). Scans consisting of 35 data scenes cover the $2,800$ -km-wide orbit swaths; each scan takes about 8 seconds to complete. The ground resolution of the TOMS ranges from 50 km at nadir to approximately 200 km at the orbit edges (average resolution is 66 km). The TOMS instrument has provided global SO_2 and ozone data since late 1978 by measuring the ultraviolet albedo, which is the ratio of backscattered Earth radiance to incoming solar irradiance (Krueger, 1983). The amount of SO_2 is calculated on the basis of characteristic attenuation of the ultraviolet signal. The TOMS instrument gives no direct indications of SO_2 cloud altitude; altitude can be derived by comparison of cloud movement to known or modeled wind velocity and directions (see Schoeberl and others, 1993).

Sulfur dioxide concentrations are calculated in milli-atmosphere centimeters (matm-cm), which express the column thickness of the gas at standard temperature and pressure (STP). The tonnage of SO_2 over

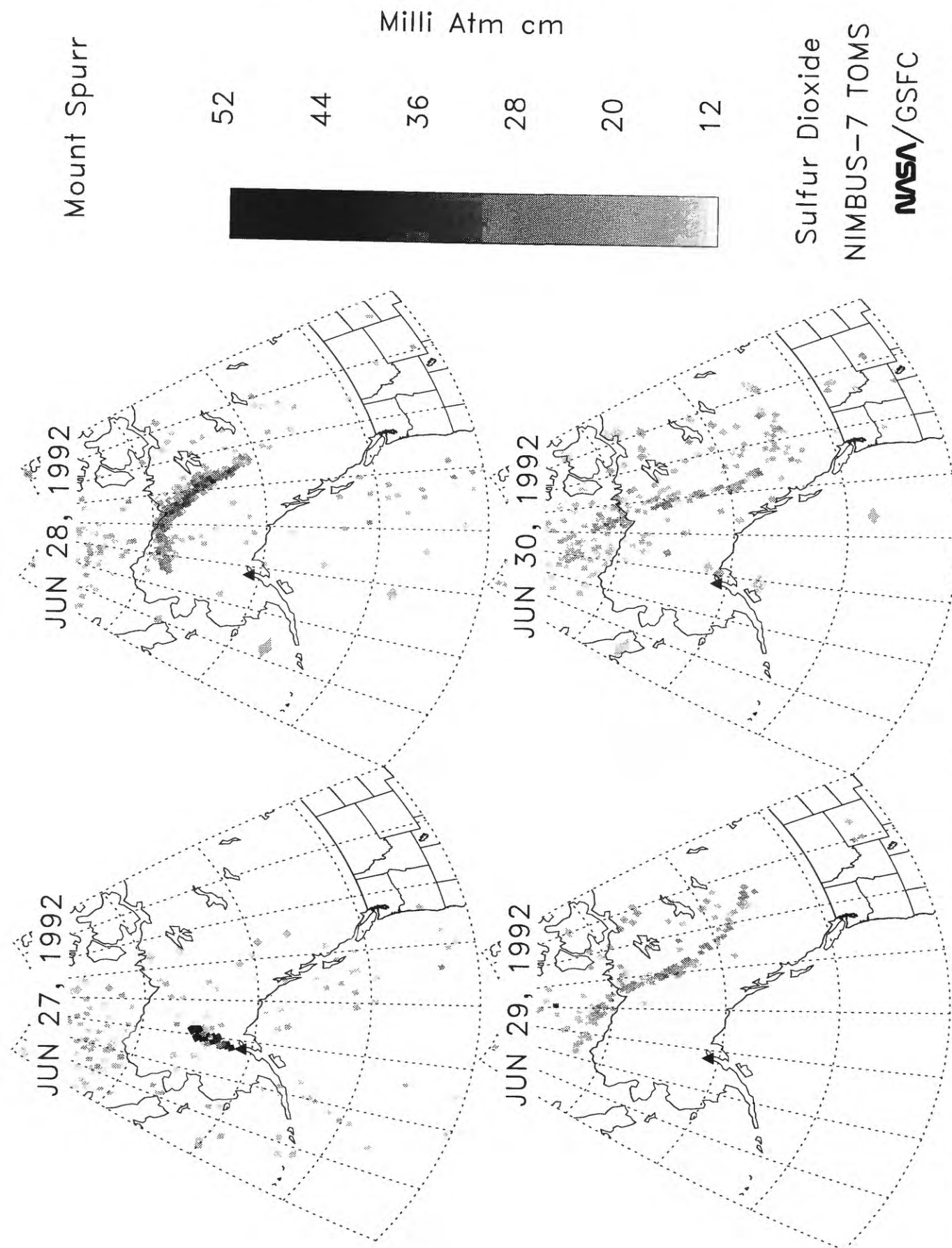


Figure 1. Total Ozone Mapping Spectrometer (TOMS) images of the SO_2 cloud for 4 days after the June 27, 1992, eruption of Crater Peak vent of Mount Spurr in Alaska. The SO_2 scale is in milli-atmosphere centimeters, a measure of the total SO_2 gas present in a column between the satellite and the reflecting surface (Earth). These measurements are then converted to a mass of SO_2 . Approximate overpass times, in ADT, are: June 27: 10:49 a.m.; June 28: 9:25 a.m.; June 29: 9:41 a.m.; June 30: 9:58 a.m. (ADT = UT - 8 hours). Mount Spurr location is shown by a filled triangle.

a given region is then obtained by multiplying the column amounts by their ground surface areas. Empirical corrections are used to offset the linear effects that variations in ozone and surface reflectance can impart on the SO_2 calculations. Eruption cloud tonnages are estimated from measurements of the SO_2 cloud compared to background measurements. Background data are taken adjacent to the SO_2 cloud to minimize latitudinal (such as, solar zenith angle and climatological variations) differences, and they are also taken in identical scan positions to avoid scan bias (such as, changes in instrument resolution from nadir to scan-edge positions).

The typical TOMS detection limit of a fresh, tightly constrained eruption cloud is roughly 5 to 10 kilotons (kt) SO_2 . Older, more dispersed clouds are difficult to discern because of low signal-to-noise ratios. Uncertainties arising from background variations and theoretical limitations of the SO_2 algorithm combine to give a typical total error of approximately 30 percent in cloud tonnage, as calculated by Krueger and others (1990). The location of Mount Spurr in the northern latitudes makes TOMS data collected there more difficult to analyze owing to high solar zenith angles (the longer path lengths increase background noise), but it is also advantageous in that more observations are possible because of orbit overlap.

ACKNOWLEDGMENTS

This work was supported by funding from NASA Headquarters Solid Earth Program. Scott Doiron and Elizabeth Duncan provided technical and scientific support to the TOMS SO_2 research. This paper was improved by the helpful reviews of Terry Gerlach and Bill Rose.

JUNE 27 ERUPTION

This first violent eruption of Mount Spurr began at 7:04 a.m. Alaskan daylight time (ADT = UT - 8 hours), and it lasted over 4 hours (Power and others, 1992). A composite image of TOMS observations is shown in figure 1. The first TOMS overpass occurred 10:49 a.m. ADT, and it caught the SO_2 emission in progress. A second view of the erupting cloud was obtained in the subsequent, overlapping orbit, at approximately 12:23 p.m. ADT. The gas cloud was observed initially drifting northward and its movement was consistent with a 10-km altitude, as determined from wind data from the National Meteorological Center (NMC). By the next day the cloud was elongated and had begun to drift to the east and southeast. By

June 29 the cloud became further elongated; it stretched from Point Barrow, Alaska, to Edmonton, Alberta, in Canada. However, on June 30 the cloud unexpectedly ceased its southeastward drift, and it rotated about 10° clockwise from its position on the previous day. The SO_2 cloud became more difficult to discern from background, but observations over the next 3 days indicated that the cloud continued its slow clockwise rotation. From NMC data, the shearing action on the 29 and 30 and subsequent cloud motions of the cloud suggested that the upper portion of the SO_2 cloud had reached altitudes of around 12 km. The last recognizable view of the cloud was obtained July 3 when it was oriented along 140°W longitude and stretched from Dawson in the Yukon province to the Gulf of Alaska.

AUGUST 18 ERUPTION

The Crater Peak vent of Mount Spurr erupted again at 4:42 p.m. ADT on August 18 for approximately 3.5 hours (McGimsey and Dorava, 1992). The SO_2 cloud emitted from this eruption drifted southeast, and it was observed by the TOMS instrument at 10:51 a.m. ADT on August 19 as a discrete mass over the Gulf of Alaska (fig. 2). By the next day the cloud had begun to spread and show evidence of shearing on both southern and northeastern fronts. On the August 21 TOMS image, the SO_2 cloud consisted of a main portion located off the coast of Oregon, and an elongated section that extended over 3,000 km to the east over Lake Winnipeg, Canada. The western edge of the cloud could not be delineated because of TOMS data losses during that particular overpass (for comparison with the extent of the ash component of the cloud on this day, see Schneider and others, this volume). By August 22, the elongated eastern section had separated from the main cloud mass, and it was barely detectable near Hudson Bay. NMC data suggest that this part of the cloud had been sheared away by winds at an approximately 14-km altitude. The main cloud was located over Oregon and showed an eastward drift that indicated an altitude of about 12 km. On subsequent days this main cloud mass elongated and followed a tortuous route along the jet stream northeast over Hudson Bay, south of Greenland, then across the Atlantic. The last remnant of the cloud visible to TOMS was observed over England on August 26.

SEPTEMBER 16-17 ERUPTION

At 12:03 a.m. ADT (September 17), the third major eruption of Mount Spurr began, and it also lasted for approximately 3.5 hours (Alaska Volcano Obser-

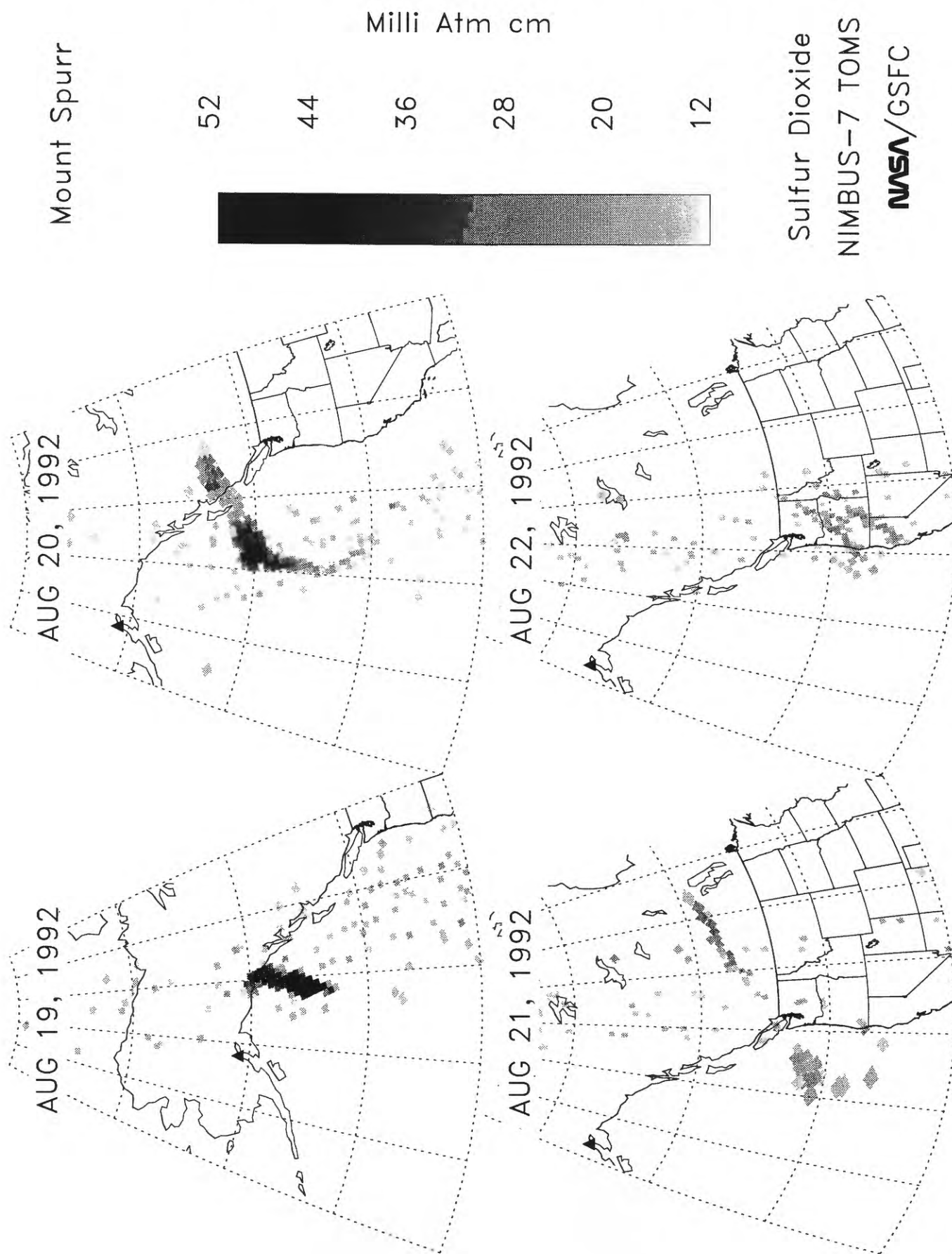


Figure 2. SO₂ cloud progression from the August 18, 1992, eruption of Crater Peak vent of Mount Spurr in Alaska. Approximate overpass times, in ADT: August 19: 10:51 a.m.; August 20: 11:07 a.m.; August 21: 9:41 a.m.; August 22: 9:57 a.m. This cloud was tracked for several more days as it crossed the Atlantic and reached England on August 26. Mount Spurr location is shown by a filled triangle.

vatory, 1993). The SO_2 cloud drifted eastward into Canada, then southward over the United States at a remarkable rate; it reached North Dakota (nearly 3,700 km away) approximately 34 hours after the eruption began (fig. 3). The cloud remained fairly compact during this time, and its behavior was consistent with NMC wind patterns over a range from 12 to 16 km. NMC data located the SO_2 cloud on a high-pressure ridge during the first 2 days after the eruption. However, by the following day the cloud was drawn into a region of high vertical shear, the steep wind gradient between high- and low- pressure systems over the Great Lakes area on the 19th. Accordingly, in the September 19 image the cloud changed drastically; it formed a 2,000-km-long arc from Detroit to northern Quebec. The extent of shearing observed was consistent with the SO_2 cloud extending from 12 to 16 km in altitude. The following day's image showed the shearing had abated, but the cloud continued its northeasterly drift over the Labrador Sea. The last day the SO_2 cloud could be discerned was on September 21 as a small mass off the east coast of Greenland.

DISCUSSION

Relations of cloud tonnage and area as a function of time can be useful in monitoring the physical and chemical processes that affect an erupted SO_2 cloud. The amount of TOMS-observed SO_2 generally decreases because of both chemical conversion of SO_2 to H_2SO_4 and physical dissipation of SO_2 at cloud margins to below the TOMS detection limits. The (chemical) decay rate of SO_2 gas can ideally be described by an exponential decay function, with an $e/2$ or e -folding time. For large eruption clouds emplaced in the stratosphere, in which chemical processes dominate physical effects, the TOMS measurements can be used to derive an e -folding time for the cloud (for example, approximately 35 days for the 1991 Mount Pinatubo SO_2 cloud; Bluth and others, 1992). However, the dispersion rate of SO_2 is altitude dependent; it is typically more persistent if emplaced in the relatively drier, less turbulent, and less reactive stratosphere rather than in the troposphere. Sulfur dioxide dispersion is also dependent on cloud size and shape; the larger the surface area to mass ratio, the greater the effects of dissipation and chemical reaction at cloud boundaries.

Because TOMS observes volcanic SO_2 clouds for only a short period of their total lifetimes in the atmosphere, the clouds' dispersion rates within this time can be approximated by a linear daily decrease. Empirically, SO_2 clouds emplaced in the stratosphere generally decrease in mass by about 10 percent per day

(Doiron and others, 1991; Bluth and others, 1992), whereas tropospheric clouds can decrease by up to 50 percent per day (Bluth and others, 1994). In contrast, cloud area typically increases for the first few days, reflecting expansion of the gas cloud. As dispersion processes become more prevalent, the areal extent of SO_2 observable by TOMS then decreases owing to chemical conversion and physical dissipation. Rapid changes in area and (or) SO_2 amount generally indicate that the clouds have undergone physical changes as a result of wind shear and other turbulence. The daily tonnages and areal extents of the SO_2 clouds for the three Crater Peak eruptions are shown as a function of time in figure 4.

The eruption of June 26 was in progress when initially observed by TOMS, and the first two SO_2 and area values in figure 4A show that the eruption cloud was still developing. Two more observations were made on June 27, but the second orbit missed approximately 5 to 10 areal percent of the cloud's leading edge, as reflected by the small drop in SO_2 amounts determined from one orbit to the next. Both the SO_2 mass and area of the June cloud decreased at fairly constant rates; these decreases, along with the NMC wind data and relatively minor amount of observed drift, suggest that the local meteorological conditions were fairly stable at that time.

Both the August and September eruption clouds showed marked variations in area as they underwent fairly drastic changes as a result of wind shear (figs. 4B and 4C). But the most remarkable pattern observed for these two eruptions is in the daily SO_2 tonnages. After both eruptions, measured SO_2 increased from the first to second day, followed by gradual cloud dispersion. In contrast to the June eruption, both eruption clouds were observed after complete separation from the volcano—accordingly, there was no possibility that the increase could have occurred from additional outgassing after the TOMS overpass. Nor was there evidence of severe physical disturbance during this time.

We can offer several reasons why measured SO_2 may have increased in the clouds. First, the data variation falls within our stated accuracy of cloud tonnage estimations (± 30 percent; however, we do not believe the pattern is an artifact of the uncertainty in TOMS data or analysis. When analyzing cloud tonnages on successive days, the relative precision of the TOMS on these daily measurements of an SO_2 cloud is approximately 10 percent. Thus, the SO_2 increases observed for these eruption clouds are probably well above uncertainty levels.

Second, the SO_2 retrieval on the first day may have been incomplete because of interference from co-erupted ash. Data from the Advanced Very High

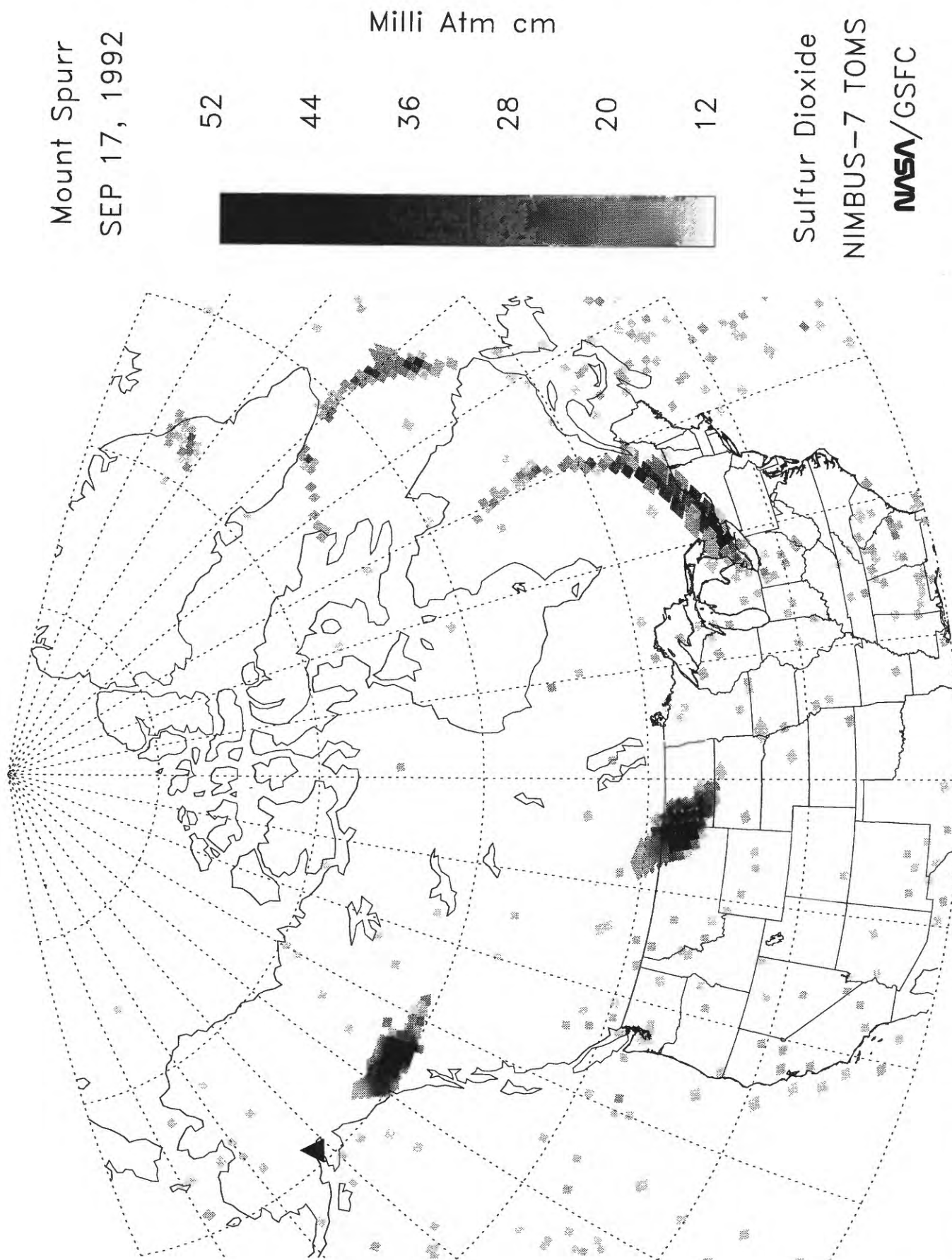


Figure 3. Composite image of the September 16-17, 1992, SO₂ cloud, observed for 5 days after the eruption of Crater Peak vent of Mount Spurr in Alaska. Approximate overpass times for each cloud observation, in ADT: September 17: 10:47 a.m.; September 18: 9:16 a.m.; September 19: 6:07 a.m.; September 20: 4:45 a.m.; and September 21: 3:20 a.m. Mount Spurr location is shown by a filled triangle.

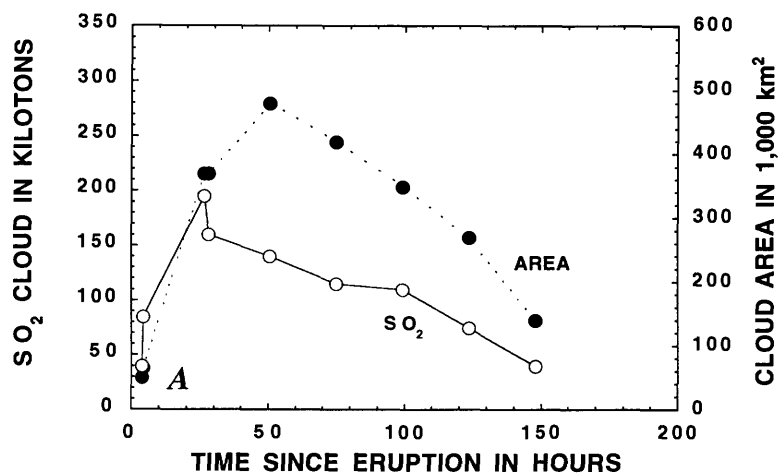


Figure 4. Comparisons of cloud sulfur dioxide mass (open circles) and area (filled circles) as a function of time after 1992 eruptions of Crater Peak vent of Mount Spurr in Alaska. A, June 27 eruption.

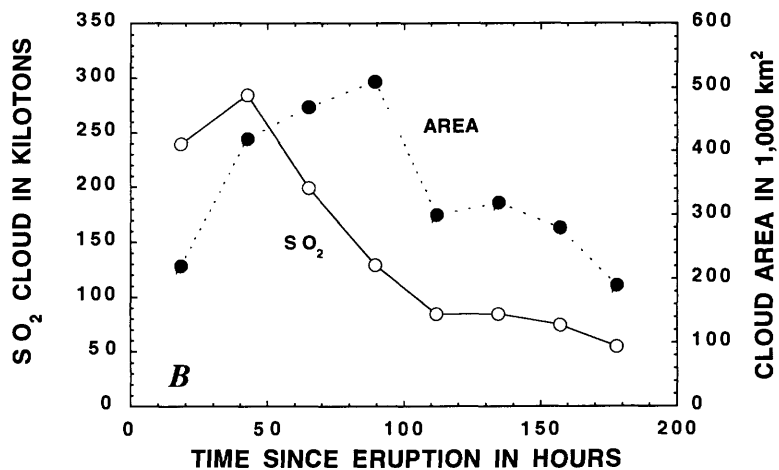


Figure 4. Continued. B, August 18 eruption.

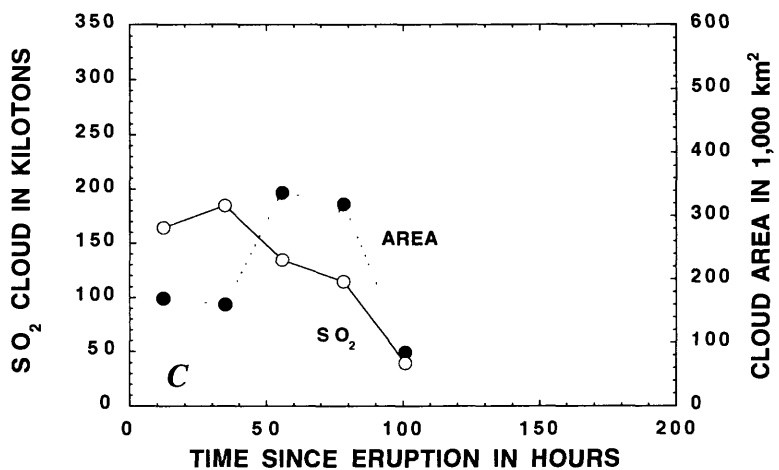


Figure 4. Continued. C, September 16-17 eruption.

Resolution Radiometer (AVHRR; Schneider and others, this volume) show that the geographic positions and dimensions of the ash clouds were remarkably similar to the gas clouds. Because of the coincidence of the two, ultraviolet (UV) light scattering by the ash may have partially masked detection of SO_2 . However, TOMS reflectivity data (taken at two wavelengths where ozone and SO_2 are nonabsorbing) showed no evidence of UV attenuation from ash particles. The fact that only the first TOMS observations were affected, 18 hours (August) and 12 hours (September) after eruption, may also provide some constraints based on the fallout rates of the tephra. For instance, the slow fallout rate of micron-sized particles in the cloud (Rose, 1993) suggests that their effect on TOMS measurements should have lasted for weeks, rather than only a single day. One of the keys to understanding the effect of ash on SO_2 detection is determining the UV absorption by the variously sized particles in the cloud.

A third possibility is that significant amounts of H_2S gas were emitted, along with sulfur dioxide, from the Crater Peak eruptions. Doukas and Gerlach (this volume) report on evidence and a possible mechanism for H_2S emissions from the volcano, and Graedel (1977) calculated that H_2S in the atmosphere, derived from oxidation reactions, has a lifetime of approximately 1 day. If significant quantities of H_2S were emitted, the oxidation of this species to SO_2 subsequent to the first day could explain the tonnage increases. The amount of H_2S required to produce the SO_2 pattern observed in the TOMS data is roughly 75 kt H_2S for the August eruption (a 3:1 SO_2 - H_2S ratio), and 25 kt H_2S for the September eruption (a 6:1 SO_2 - H_2S ratio).

This pattern of SO_2 signals from an eruption cloud was observed previously in the TOMS data after the eruption of Mount St. Helens in 1980. The TOMS SO_2

data for this eruption showed an increase of about 10 percent from the first to the second day of observation, then decreasing tonnage measurements on the following days (unpub. data). On May 18, only hours after the major eruption, Hobbs and others (1981) sampled the eruption cloud and reported high concentrations of H_2S . However, we do not have enough data to make definitive conclusions regarding TOMS detection of H_2S gas emissions from either the Mount St. Helens or the Crater Peak eruptions.

The characteristics of the three eruptions of Mount Spurr are summarized in table 1. The emitted SO_2 is calculated from the TOMS data, assuming a linear daily loss of SO_2 . Data from each day were used until cloud losses through shearing were evident. A best-fit line was determined using 7 days after the June eruption, days 2 through 4 for each of the August and September eruptions, and projected back to the approximate end point of the eruption.

The August eruption was calculated from the TOMS data to have produced the greatest amount of SO_2 , at 400 ± 120 kt. The September eruption produced 230 ± 70 kt, and the June eruption emitted 200 ± 60 kt SO_2 . There was no apparent correlation of gas production with tephra emission that the Alaskan Volcano Observatory estimated for the June, August, and September eruptions. These estimates were 12, 14, and 15 million m^3 , respectively (W. Rose, oral commun., 1993).

The Mount Spurr SO_2 tonnage data corresponded to a 20 to 30 percent daily loss of SO_2 for each of the three eruptions; this loss suggests the clouds were injected at middle altitudes between purely stratospheric and purely tropospheric layers. At the latitudes of Alaska and southern Canada, the summertime tropopause height fluctuates but is generally around 12 km altitude. The maximum injection altitudes of the June,

Table 1. Observations of SO_2 clouds produced by the 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska.

Eruption	TOMS data			SO_2 cloud altitude ¹	Max ash altitude ²	Emitted tephra ³
	Days observed	Maximum areal extent	Estimated SO_2 emitted			
June 27	7	480,000 km^2	200 ± 60 kt	10-12 km	14.5 km	$12 \times 10^6 \text{ m}^3$
August 18	8	510,000	400 ± 120	12-14	13.7	14
September 16-17	5	340,000	230 ± 70	12-16	13.9	15

¹Calculated from NMC wind data and actual cloud positions from TOMS data

²Rose and others (this volume)

³Alaska Volcano Observatory (1993)

August, and September ash clouds ranged from 13.7 to 14.5 km (table 1, modified from Rose and others, this volume). Each of these eruptions should therefore have reached the lower stratosphere; however, comparisons of daily wind data to actual cloud positions indicate the SO₂ clouds are most accurately considered to extend several kilometers vertically. The June SO₂ cloud corresponded to the lowest altitude (10–12 km) on the basis of its drift speed and directions, and most of this cloud probably remained in the troposphere. In contrast, most of the September cloud (ranging from 12–16 km by our calculations) was in the lower stratosphere. A large part of the August cloud (12–14 km) also appears to have reached the stratosphere. From the available data we therefore estimate that at least half of the SO₂ emitted by the three Crater Peak eruptions was emplaced in the stratosphere.

SUMMARY

During the summer of 1992, Mount Spurr ended a 39-year period of quiescence and emitted large clouds of gas and ash during three separate eruptions within a 3-month period. The TOMS data suggest that the August 18 eruption was the largest of the three with 400 ± 120 kt SO₂ emitted, followed by the September 16–17 (230 ± 70 kt) and June 27 eruptions (200 ± 60 kt). Data for the August and September eruptions suggest that either co-emitted ash affected the TOMS measurements, or that significant quantities of H₂S (25–75 kt), in addition to SO₂, may have been explosively outgassed by these two eruptions. However, TOMS does not measure H₂S, and thus we can only, at this time, infer this possibility from SO₂ data. Based on the maximum injection altitudes of the ash clouds, matching of ash and gas cloud positions to known wind conditions, and the dispersion rates of the gas clouds, the explosive eruptions of Crater Peak vent of Mount Spurr injected at least 50 percent of their SO₂ into the lower stratosphere.

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Sulfur Dioxide Scrubbing During the 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

By Michael P. Doukas and Terrence M. Gerlach

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ABSTRACT

Sulfur dioxide scrubbing by liquid water masked SO₂ emissions from shallow magma during the 1992 eruptions of Crater Peak and effectively prevented observation of SO₂ emissions from shallow magma both before and after explosive eruptions and seismic crises. Airborne ultraviolet correlation spectrometer (COSPEC) measurements from July 22, 1991, to September 24, 1992, indicate only background to minor (<100 t/d) noneruptive SO₂ emissions from Crater Peak, even though this period included the onset of precursory seismicity beneath Crater Peak (August 1991), the peaking in frequency of volcano-tectonic earthquakes and the initiation of volcanic tremor bursts (June 5, 1992), and three explosive eruptions that produced SO₂ emissions of 200 to 400 kilotons (kt). The low, nonexplosive SO₂ emissions are explained by hydrolysis of SO₂ to aqueous H₂S and sulfate from interactions with liquid water: $4\text{H}_2\text{O}(\text{l}) + 4\text{SO}_2(\text{aq}) \rightarrow \text{H}_2\text{S}(\text{aq}) + 3\text{H}^+(\text{aq}) + 3\text{HSO}_4^-(\text{aq})$. Sulfur dioxide hydrolysis also explains the increase in the sulfate content of Crater Peak lake water prior to the first eruption, the strong H₂S odor during periods of background to low SO₂ emission, the TOMS evidence for significant H₂S emissions during the explosive eruptions, and the observed decline of SO₂ during periods of volcanic tremor. Abundant, local sources of melt water and a high permeability for the Mount Spurr volcanic edifice are probably the

chief factors responsible for masking SO₂ emissions by scrubbing, and possibly for quenching shallow intrusions that were ascending. Large SO₂ emissions unencumbered by scrubbing were only possible during the three explosive eruptions when magma penetrated through liquid water zones under Crater Peak and reached the surface. Nonexplosive SO₂ emissions of as much as 750 t/d were possible, however, for a brief period when dry pathways to the surface existed from September 25 until about October 10, 1992. Airborne infrared spectrometer (MIRAN) measurements of CO₂ emissions indicate that in addition to the degassing of magma through dry pathways, degassing through boiling water with the loss of SO₂ by scrubbing was also important during that time. The CO₂ emission data indicate that magma degassing was taking place, and CO₂/SO₂ values calculated from MIRAN and COSPEC data are in the range 10 to 100, which supports the hypothesis of SO₂ loss by scrubbing. Because of its strong preference for the vapor phase during boiling, CO₂ emissions from degassing magma are less likely to be masked by the presence of water, whereas SO₂ emissions may be lost totally from interactions with water; thus misleading COSPEC results are obtained. We recommend prompt and early monitoring of CO₂ when Cook Inlet volcanoes become restless.

INTRODUCTION

The USGS Volcano Hazards Program supports airborne ultraviolet correlation spectrometer (COSPEC) and infrared spectrometer (MIRAN) flights to monitor sulfur dioxide and carbon dioxide gas emissions from convergent-plate volcanoes. These methods have been successfully applied to explosive eruptions at Mount St. Helens in the Cascade volcanic system (Casadevall and others, 1981, 1983; Harris and others, 1981; McGee, 1992) and at Redoubt Volcano in the Cook Inlet part of the Aleutian volcano chain (Casadevall and others, 1994). These investigations show that the emplacement and degassing of shallow magma in the crust beneath active volcanoes can be

monitored by measuring the emission rates of SO_2 and CO_2 from volcanoes to the atmosphere. A common feature observed to date in the fluxes of these gas species is their tendency to decline exponentially through time after major explosive episodes. During the 1992 eruptions at Crater Peak vent on Mount Spurr volcano, however, we discovered SO_2 gas fluxes that were markedly lower than those observed in past investigations. We suggest that the anomalously low SO_2 degassing patterns observed at Crater Peak are related to the hydrolysis of magmatic SO_2 to aqueous H_2S and sulfate by liquid water within the volcano.

Herein, we report the results and interpretation of SO_2 and CO_2 emission measurements obtained at Crater Peak from late 1991 through early 1993, a period that included three subplinian eruptions. (For the eruption chronology, see Eichelberger and others, this volume). Sulfur dioxide measurements were taken at Mount Spurr volcano in July of 1991 as part of the Alaska Volcano Observatory (AVO) gas-monitoring program for the volcanoes of the Cook Inlet. However, few SO_2 flux measurements were made until after the third eruption on September 16–17, 1992, because of budget constraints, the surprisingly low SO_2 flux observed at Crater Peak following eruptions on June 27 and August 18, 1992, and the perception that the activity would cease soon after the initial erup-

tions, as it did during the 1953 eruption at Crater Peak. Shortly after the September 16–17 eruption, AVO intensified SO_2 monitoring and added CO_2 flux determinations in the gas-monitoring program at Mount Spurr volcano. The CO_2 measurements were motivated by the working hypotheses that (1) both SO_2 and CO_2 were degassed noneruptively from magma and that (2) subsequent dissolution in liquid water within the volcano resulted in scrubbing of SO_2 by hydrolysis reactions while CO_2 remained available for emission when the liquid boiled.

ACKNOWLEDGMENTS

We are grateful for the aircraft support provided through the Office of Aircraft Services (OAS), Department of Interior, Anchorage. We are most appreciative of the pilots who skillfully and safely flew under and through the Crater Peak gas plume, often under less than ideal flying conditions. Our special thanks to Mark Iven of the Cascades Volcano Observatory who promptly checked the condition of the MIRAN instrument and shipped it to AVO after the September 16–17 eruption. We thank Terry Keith, Ken McGee, Steve McNutt, and Jeff Sutton for their reviews of a draft of this report and for suggesting several improvements.

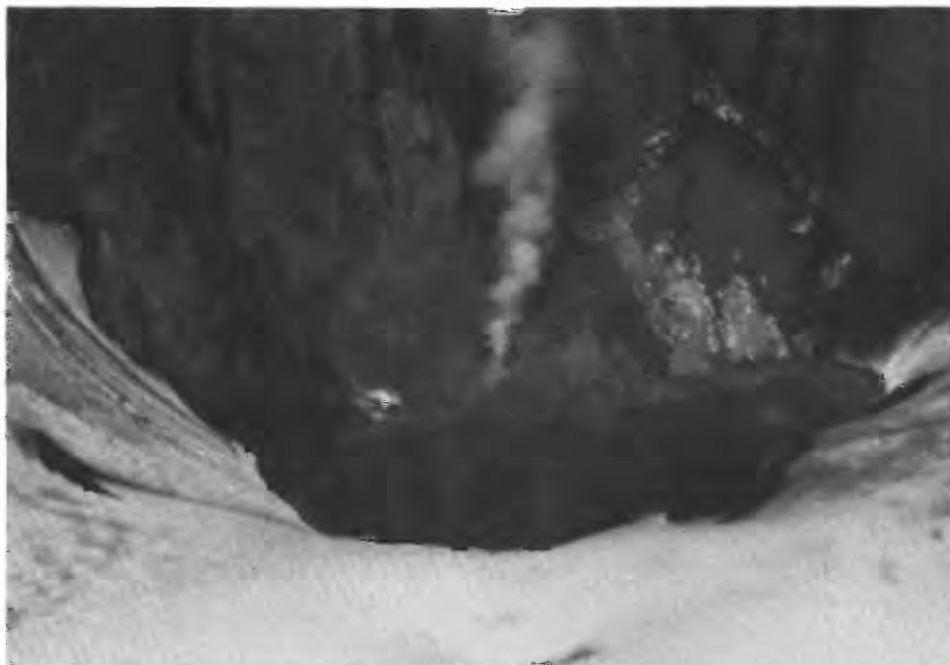


Figure 1. View of pre-eruption crater lake in summit vent of Crater Peak, Mount Spurr volcano before the June 27, 1992, eruption. Main fumarole (center) located in talus pile in northwestern quadrant of vent. Whitish upwelling located right center on lake is vapor off the lake surface, which indicates an elevated lake temperature.

PRE-ERUPTION DEGASSING

Pre-eruption features at the Crater Peak vent on Mount Spurr included a summit crater lake occupying the 1953 explosion crater (fig. 1), a vigorous fumarole emitting through talus debris at the northwest side of the lake, and wispy fumaroles emitting steam from pyroclastic deposits on the eastern interior of the crater. The crater lake normally was blue-green in color with no apparent bubbling or surface roiling. A few weeks prior to the June 27, 1992, eruption, the lake water changed to a gray color, its temperature was measured at 49°C, and active boiling and upwelling became apparent. Observers reported H₂S odor at the crater rim several tens of meters above the lake's surface. Analyses of crater lake samples collected on June 11, 1992, showed about a fourfold increase in sulfate content compared to lake samples collected on August 9, 1970 (Keith and others, this volume). By the day before the June 27 eruption, the lake level had dropped significantly and exposed a muddy beach and shoreline. The remaining lake and western talus cone were probably ejected during the June 27 eruption. The main vent present after the June 27 eruption included the site of the principal pre-eruption fumarole (fig. 2).

PROCEDURES AND METHODS

SULFUR DIOXIDE MEASUREMENTS

Sulfur dioxide measurements were made from a fixed-wing airplane. The COSPEC method uses solar ultraviolet light scattered by the Earth's atmosphere as an illumination source. With the instrument directed toward the sky, the COSPEC measures the difference between the amount of ultraviolet light absorbed by SO₂ molecules present along the optical path through the plume and the ambient total ultraviolet light available. The COSPEC measures the amount of absorption in units of concentration pathlength (ppm-m) or burden. Calibration is performed by using an internal standard with a known SO₂ burden of 135 ppm-m. A profile of SO₂ burden across the plume is measured at right angles to the plume trajectory. The product of the burden (ppm-m), the width of the plume (m), and the plume speed (m/s) gives the emission rate of SO₂ (ppm-m³/s), which is reported in this study in metric tons per day (t/d; 1 metric ton (t) = 10³ kg). Reported values typically consist of four to six separate traverses averaged together and rounded to two significant figures (Casadevall and others, 1981, 1983, and 1994; McGee, 1992).



Figure 2. View of crater in Crater Peak vent, Mount Spurr volcano, Alaska, after the June 27, 1992, eruption. Dark deposit in middle ground is snow-free pyroclastic debris within crater. Dark area in background is vertical south-facing crater wall.

Measured SO₂ emission rates can vary owing to natural causes such as variation in emission rate, changes in sun angle, and amount of ultraviolet light and plume opacity due to the presence of suspended ash. Uncertainty of COSPEC measurements arises from several operational factors including operator variance (pilot, instrument manipulation, data reduction) and measurement of plume velocity (Casadevall and others, 1994).

At Crater Peak, the accurate measurement of plume velocity was usually the largest source of uncertainty. We assumed that the plume was carried along at the same speed as the ambient winds. Wind direction and speed near the vent are controlled by adjacent mountainous topography. Winds passing over the location of Crater Peak in this part of the Alaska Range can be turned, shifted, and back-eddied. Local wind directions were occasionally 180 degrees from the regional trend reported by the National Weather Service (NWS) in Anchorage, Alaska. Wind speed and direction were determined in several ways. One method relied on wind and speed determinations made twice daily at Anchorage International Airport by the National Oceanic and Atmospheric Administration (NOAA) radiosonde. One of these measurements is made at 2:00 p.m. local time, which coincides with local solar noon (highest ultraviolet levels) and the normal times of the SO₂ flights. When forecast winds were light and variable, we assigned 5 knots as the operative value. Another method was to compare onboard navigational instruments or global positioning satellite (GPS) receivers with NWS forecasted winds aloft.

Aircraft speed was determined in several ways. During close orbits of the vent, geographical landmarks were identified and noted on the SO₂ chart recorder paper. The chart recorder was then used as a clock to time one orbit. Using the distance from the vent as a radius, the aircraft ground speed was calculated. At other times, radar tracking from Anchorage International Airport was used to calibrate onboard instruments for reduction of aircraft airspeed to ground speed. Ideally, a moderately high emission rate is desirable (greater than 500 t/d of SO₂) along with wind speeds of nearly 10 nautical miles per hour. During times of low SO₂ flux or low wind speeds, orbits within a kilometer of the vent were required. Plume positions during these times, or when wind directions drove the plume into mountainous terrain, would not allow the aircraft to fly safely beneath the entire plume; these alternate routes resulted in minimum values being measured for SO₂.

CARBON DIOXIDE MEASUREMENTS

Carbon dioxide measurements were obtained from a fixed-wing aircraft by flying directly through the plume and using the infrared spectrometer technique described by Harris and others (1981) for measuring the amount of excess CO₂ in the plume relative to that in the ambient atmosphere. This method is capable of detecting excess CO₂ down to 1 or 2 ppm (by volume) and CO₂ emission rates down to about 200 t/d. The excess CO₂ concentration profiles along a number of passes through the plume at different altitudes are combined to define the plume area and CO₂ concentration in a cross section of the plume. The CO₂ emission rate is calculated from the plume area, the CO₂ concentration in the cross section, the density of CO₂ from the altitude and temperature of the plume, and from the wind speed (Harris and others, 1981). During times of low emission and low wind speeds, orbits within a kilometer of the vent were required, which gave low CO₂ emission rates because of slow instrumental response times. Thus, some CO₂ emission rates reported here are minimum values.

RESULTS

The SO₂ and CO₂ emission rate data obtained for 1991 through 1993 are presented in table 1. The data as a time series that includes the three 1992 eruptions of June 27, August 18, and September 16–17 are shown in figure 3; periods of low UV radiation are also shown in this figure. Data obtained after the September 16–17 eruption when COSPEC flights were more frequent and the MIRAN instrument was used for CO₂ measurements are shown in figure 4. All emission rate data are reported in metric tons/day (t/d) above the background level upwind of the volcano.

COSPEC measurements did not detect significant increases above background in SO₂ emissions prior to the June 27, 1992, eruption. Some COSPEC flights made during times of increased seismicity recorded SO₂ emissions slightly elevated above background levels. For example, two measurements in May and June were 88 and 21 t/d, respectively, during a period of increased seismic activity. COSPEC measurements on June 29, just 2 days after the first eruption, showed SO₂ emission rates of only 5 t/d. Measurements on September 10, a week prior to the September 16–17 eruption, were at background levels, and measurements 4 and 6 days after the eruption were only slightly above background (23 and 24 t/d). Strong H₂S odors

Table 1. SO₂ and CO₂ fluxes measured at Crater Peak, Mount Spurr volcano, Alaska, 1991–1993.[d-m-y, day-month-year; t/d, metric tons per day; sd, standard deviation; no., number of SO₂ determinations; BG, background; NM, not measured; >, minimum value]

Date (d-m-y)	SO ₂ avg (t/d)	SO ₂ max (t/d)	SO ₂ min (t/d)	SO ₂ sd (±t/d)	SO ₂ no.	CO ₂ (t/d)	CO ₂ /SO ₂ (wt)
22-07-91	BG	BG	BG	-	-	NM	-
• <i>Precursory seismic activity begins beneath Crater Peak, August 1991</i>							
21-08-91	BG	BG	BG	-	-	NM	-
29-08-91	BG	BG	BG	-	-	NM	-
03-09-91	85	97	72	18	2	NM	-
22-11-91	BG	BG	BG	-	-	NM	-
14-05-92	88	96	80	10	2	NM	-
• <i>Volcano-tectonic earthquakes peak, June 5, 1992</i>							
• <i>First volcanic tremor bursts, June 6, 1992</i>							
08-06-92	21	24	18	4	2	NM	-
• <i>Eruption, June 27, 1992</i>							
29-06-92	5	10	2	3	5	NM	-
• <i>Eruption, August 18, 1992</i>							
10-09-92	BG	BG	BG	-	-	NM	-
• <i>Volcanic tremor episodes, September 12–25, 1992</i>							
• <i>Eruption, September 16–17, 1992</i>							
21-09-92	23	38	10	12	5	NM	-
23-09-92	24	33	22	7	4	NM	-
24-09-92	79	92	72	11	3	NM	-
25-09-92	300	370	240	54	4	11,000	53
28-09-92	190	270	180	25	4	12,000	90
29-09-92	750	830	690	75	3	8,700	17
• <i>Volcanic tremor episode, October 2–5, 1992</i>							
02-10-92	450	450	450	0	2	NM	-
03-10-92	220	320	110	110	4	4,800	32
04-10-92	210	320	170	72	4	3,500	24
05-10-92	360	560	230	170	4	2,900	12
10-10-92	240	360	160	100	3	NM	-
• <i>Volcanic tremor episode, October 12, 1992</i>							
14-10-92	47	50	44	4	3	2,400	74
15-10-92	64	67	62	4	2	NM	-
16-10-92	>200	-	-	-	-	>2,000	-
19-10-92	BG	BG	BG	-	-	BG	-
23-10-92	24	26	22	3	2	2,100	127
29-10-92	BG	BG	BG	-	-	BG	-
• <i>Earthquake swarms, November 9, 1992</i>							
10-11-92	BG	BG	BG	-	-	BG	-
11-11-92	BG	BG	BG	-	-	BG	-
12-11-92	BG	BG	BG	-	-	1,100	•
13-11-92	BG	BG	BG	-	-	1,000	•
15-11-92	BG	BG	BG	-	-	1,000	•
08-12-92	BG	BG	BG	-	-	>1,000	•
11-12-92	BG	BG	BG	-	-	>500	•
15-12-92	BG	BG	BG	-	-	3,200	•
18-12-92	BG	BG	BG	-	-	>1,000	•
• <i>Swarm of small VT earthquakes, December 21–26, 1992</i>							
23-12-92	BG	BG	BG	-	-	>1,000	•
21-01-93	BG	BG	BG	-	-	>1,000	•
18-02-93	BG	BG	BG	-	-	>400	•
• <i>Seismicity at pre-August 1991 levels, April 1993</i>							
24-04-93	BG	BG	BG	-	-	>930	•

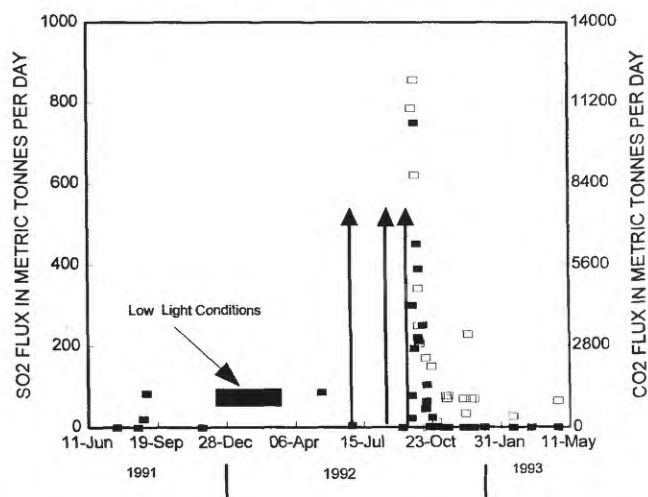


Figure 3. Sulfur dioxide and CO₂ measurements at Crater Peak of Mount Spurr volcano, Alaska, from 1991 to 1993. Left vertical axis and solid squares give SO₂ emission rate (in metric tons per day, or t/d); right vertical axis and open squares give CO₂ flux (t/d). Arrows are eruptions dates (June 27, August 18, and September 16–17, 1992). Bar indicates low ultraviolet light period during winter when COSPEC measurements are not possible.

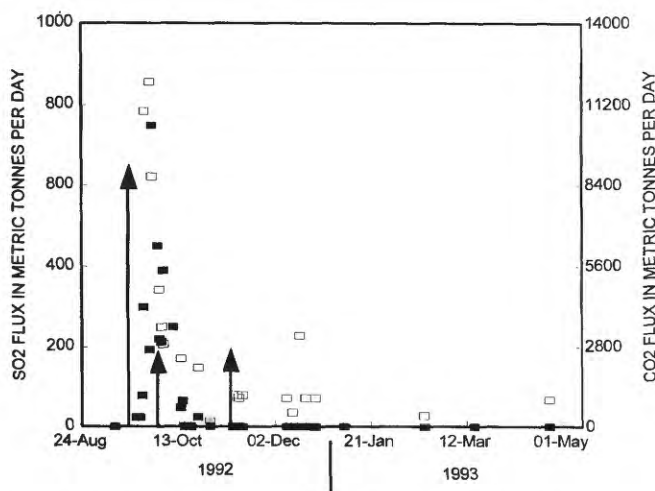


Figure 4. Sulfur dioxide (solid squares) and CO₂ (open squares) emission rates after September 16–17, 1992, eruption of Crater Peak vent of Mount Spurr volcano, Alaska. Long arrow denotes September eruption; short arrows are periods of elevated seismic energy release on October 1–2 and November 9, 1992.

were frequently noted during airborne monitoring flights throughout the course of the 1992 eruptions. The periods of strongest H₂S odor correlated with times of minor to background SO₂ emission rates.

On September 24, SO₂ emission rates reached 79 t/d. By September 25, they rose to 300 t/d, just as

the first CO₂ measurements showed 11,000 t/d. Measurements on September 28 showed CO₂ at 12,000 t/d. Sulfur dioxide values peaked on September 29 at 750 t/d, coincident with a CO₂ flux of 8,700 t/d.

In the week following the October 2–5 tremor episode, SO₂ remained above background levels, although it declined somewhat from 450 to 240 t/d, as CO₂ declined from 4,800 to 2,900 t/d. During this period, aerosol plumes of sulfate fume derived from volcanic SO₂ became visible for the first time downwind from the crater. These plumes extended beyond the water vapor condensation cloud near the vent and spread out along an inversion layer in the atmosphere (fig. 5). On October 3, 4, and 5, SO₂ fluxes were measured before and after the onset of volcanic tremor. After about 2 hours into each tremor episode, SO₂ emission rates decreased to approximately half their pre-tremor values (fig. 6). Sulfur dioxide emission rates, but not CO₂ rates, were also down sharply after the tremor episode on October 12. Measurable SO₂ levels gradually declined over the next 2 weeks and returned to near background levels by the end of October 1992. Three of five CO₂ measurements over the same period gave emission rates over 2,000 t/d; the other two measurements showed background levels.

Both MIRAN and COSPEC surveys were made of the atmosphere around the volcano immediately after the November 9 earthquake swarms. Sulfur dioxide emission rates never rose above background levels following the crisis, but CO₂ emissions rose above background to a level of about 1,000 t/d by the third day after the crisis and stayed at about this level through April 1993. Occasional later measurements showed lower CO₂ levels of only a few hundred metric tons per day, which may represent the normal baseline CO₂ fluxes from Crater Peak at times of inactivity.

Integrating the noneruptive CO₂ emission rate data for the period September 25, 1992, to April 24, 1993 (fig. 4) gives a total CO₂ emission of approximately 110 kt. A similar treatment of the COSPEC SO₂ flux data results in a total SO₂ emission of only about 6 kt.

DISCUSSION

COSPEC measurements (table 1) from July 22, 1991, to September 24, 1992, indicate only background to minor (<100 t/d) noneruptive SO₂ emissions from Crater Peak, although this period included (1) the onset of precursory seismicity beneath Crater Peak (August 1991), (2) the peaking in frequency of volcano-tectonic earthquakes and the initiation of volcanic tremor bursts (June 5, 1992), and (3) the three explosive eruptions (June 27, August 18, and September 16–17, 1992).



Figure 5. View of Crater Peak of Mount Spurr volcano from the south on October 2, 1992, shows condensation plume above Crater Peak vent (left center) and volcanic aerosol plume extending to the east.

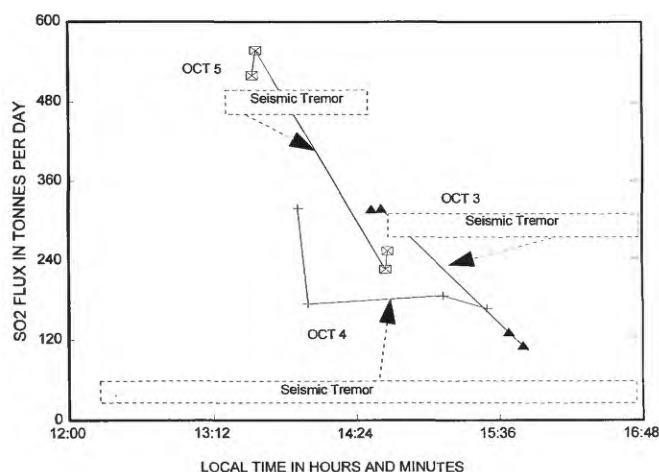


Figure 6. Relation between periods of seismic tremor and decreasing SO_2 emission rates (symbols connected by solid lines) on October 3, 4, and 5, 1992. Dotted lines with arrows point to SO_2 flux measured during tremor on particular days.

Total Ozone Mapping Spectrometer (TOMS) satellite measurements (Bluth and others, this volume) show large eruptive SO_2 emissions by comparison for the explosive events: 200 kt (June 27), 400 kt (August 18), and 230 kt (September 16–17). The integrated noneruptive SO_2 emission of about 6 kt is small compared to the total TOMS-based explosive SO_2 emission of 830 kt. The 1989–90 eruptions of nearby Redoubt Volcano involved similar amounts of andesitic magma and produced approximately 175 kt of SO_2 in explosive eruptive emissions, as determined by TOMS (Schnetzler and others, 1994), and approximately 800

kt of SO_2 in noneruptive emissions, as determined by COSPEC (Casadevall and others, 1994). Total SO_2 emissions are similar for both eruptions, but their eruptive and non-eruptive components are very different.

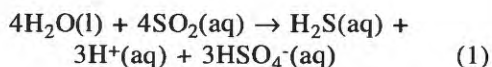
The great difference between noneruptive and eruptive SO_2 degassing may reflect the role of liquid water at Crater Peak in hydrolyzing most of the noneruptively emitted SO_2 gas; an abundance of water may also have played a role in effectively quenching shallow intrusions during their ascent to the surface. After obtaining an average SO_2 emission rate of only 5 t/d (table 1) at Crater Peak just 2 days after the June 27 eruption, we speculated that the extremely low, noneruptive SO_2 emissions resulted from the hydrolysis of SO_2 to aqueous H_2S and sulfate by liquid water present in the volcano. One likely source of water is a liquid-dominated hydrothermal system postulated to exist between 1.5 and 2 km beneath Crater Peak (Motyka and Nye, 1993). The presence of a hydrothermal system does not, however, guarantee that SO_2 emissions from shallow magma will be effectively absorbed. For example, Mount Pinatubo contained a hydrothermal system sufficiently impressive to attract exploration drilling for geothermal energy in 1988–89 (Delfin and others, 1992). Nevertheless, SO_2 emission rates at Mount Pinatubo reached values as high as 13 kt/d in the weeks before the climactic June 15, 1991, eruption (Daag and others, 1994). The Mount Pinatubo hydrothermal system had a low permeability—too low for geothermal-energy production (Delfin and others, 1992)—which may have inhibited its capacity to rapidly recharge water lost from boiling and to prevent (or greatly diminish) pre-eruption SO_2 emissions. We suggest that, in addition to the presence of

a hydrothermal system, important factors contributing to SO_2 scrubbing at Crater Peak may have included a highly permeable volcanic edifice and large sources of liquid water from snow melt and melt from glaciers directly above and surrounding Crater Peak, thus permitting rapid and abundant recharge of water to the volcano. A 1-km-long zone of warm springs (40°C) discharges from the base of the volcano about 4 km directly downslope from Crater Peak. Total warm-water flow from this system is estimated to be approximately 1,000 l/min, and stable isotope data indicate that the meteoric recharge for the ground-water system supplying the springs is derived from local snow melt and precipitation falling at middle to upper elevations (Motyka and Nye, 1993).

The main gases, in addition to water vapor, released from shallow degassing magma of convergent-plate volcanic systems are CO_2 , SO_2 , H_2S , and HCl . All these gases are soluble in liquid water that may be present in volcanic edifices. When the water boils, the dissolved gases will partition preferentially into the vapor phase. The dissolution of gases in water and their partitioning between vapor and liquid during boiling are topics beyond the scope of this report. We refer readers to a discussion of these and related topics in Henley and others (1984); much of the interpretation that follows is based on data and concepts they present. We focus entirely on liquid-vapor interactions. In most cases, water-rock interactions are likely to proceed at a slow rate relative to the time scales of volcanic degassing and liquid-vapor reactions during boiling. We assume, therefore, that water-rock interactions introduce only secondary effects that can be ignored here.

Because aqueous HCl is a strong acid, it dissociates in water to hydrogen and chloride ions ($\text{HCl}(\text{aq}) \rightarrow \text{H}^+(\text{aq}) + \text{Cl}^-(\text{aq})$). The concentration of $\text{HCl}(\text{aq})$ will therefore be negligible, and its exsolution as HCl gas in the vapor phase during boiling will be negligible, except for unusually acidic solutions at temperatures much greater than 250°C . Thus, water in a volcano will dissolve HCl released from degassing magma, but because of its strong dissociation, HCl will not normally exsolve to a vapor phase when the resulting solution is boiled.

Aqueous SO_2 is thermodynamically unstable in water relative to sulfate species and aqueous H_2S , except at temperatures greater than 300°C and pH values less than or equal to 2. The hydrolysis of SO_2 by the disproportionation reaction:



proceeds strongly in the direction indicated. Consequently, the concentration of $\text{SO}_2(\text{aq})$ is very low and the contribution of SO_2 to the vapor phase during boiling is negligible compared to the bulk amount of SO_2 gas that may have been injected initially into the water during magma degassing.

The gases CO_2 and H_2S dissolve in water as $\text{CO}_2(\text{aq})$ and $\text{H}_2\text{S}(\text{aq})$. Their dissociation ($\text{H}_2\text{O}(\text{l}) + \text{CO}_2(\text{aq}) = \text{H}^+(\text{aq}) + \text{HCO}_3^-(\text{aq})$; $\text{H}_2\text{S}(\text{aq}) = \text{H}^+(\text{aq}) + \text{HS}^-(\text{aq})$) is not appreciable in ground water and hydrothermal waters provided pH is below 9. Unlike SO_2 , they are relatively stable with respect to water, although some $\text{H}_2\text{S}(\text{aq})$ may react with dissolved Fe^{++} to form pyrite (FeS_2). The concentrations of $\text{CO}_2(\text{aq})$ and $\text{H}_2\text{S}(\text{aq})$ in waters that have absorbed magmatic gases are, therefore, significant compared to those for aqueous HCl and SO_2 . Carbon dioxide and H_2S are thus distributed between vapor and liquid phases during boiling. The distribution of either gas between vapor and liquid can be represented by a distribution coefficient, B_i , defined as follows:

$$B_i = (n_i/n_{\text{H}_2\text{O}})_{\text{vap}} / (n_i/n_{\text{H}_2\text{O}})_{\text{liq}} \quad (2)$$

where $(n_i/n_{\text{H}_2\text{O}})_{\text{vap}}$ is the molar ratio of the gas i to H_2O in the vapor, and $(n_i/n_{\text{H}_2\text{O}})_{\text{liq}}$ is the ratio of the gas to H_2O in the liquid. Giggenbach (1980) derived the following regression equations, valid for 100°C to 340°C , from experimental gas-solubility data for CO_2 and H_2S in water:

$$\log B_{\text{CO}_2} = 4.7593 - 0.01092T \quad (3)$$

$$\log B_{\text{H}_2\text{S}} = 4.0547 - 0.00981T \quad (4)$$

where T is in $^\circ\text{C}$. Over temperatures from 100°C to 340°C , the respective values for B_{CO_2} range from 4,650 to 11; those for $B_{\text{H}_2\text{S}}$ range from 1,185 to 5. Thus, both CO_2 and H_2S will move preferentially into the vapor phase during boiling, and this preference is stronger the lower the temperature of boiling.

With equations (3) and (4) and other equations provided in Henley and others (1984), it is possible to calculate the distribution of CO_2 and H_2S between vapor and liquid as boiling progresses. The results depend on the temperature of boiling, the extent of boiling (that is, the fraction of the total water in the system transferred to the vapor phase as steam), the isothermal or isoenthalpic nature of the boiling process, and the open or closed character of the system. The percent of CO_2 and H_2S remaining dissolved in the liq-

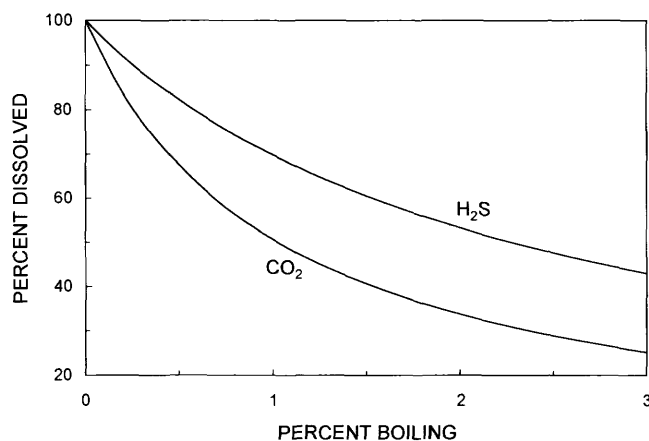


Figure 7. Percent of CO₂ and H₂S dissolved in liquid water, relative to the initial amounts in solution, as a function of the percent isenthalpic boiling for a closed system initially at 260°C. Only 25 percent of the initial CO₂ and 43 percent of the initial H₂S remain dissolved after just 3 percent boiling. The temperature after 3 percent boiling is 250°C. The curves are based on calculations at intervals of 1 percent boiling from equations (3) and (4) in the text and other equations provided in Henley and others (1984).

uid, relative to the initial amounts in solution, as a function of the percent isenthalpic boiling for a closed system initially at a temperature of 260°C are shown in figure 7. For these conditions, only 25 percent of the initial CO₂ and 43 percent of the initial H₂S remain dissolved in the liquid after just 3 percent boiling. The temperature after 3 percent isenthalpic boiling is 250°C. The curves produced by closed-system, isothermal boiling at 260°C are similar but somewhat above those shown in figure 7 for the same range of percent boiling. For example, 29 percent of the initial CO₂ (instead of 25 percent) remains in the liquid after 3 percent boiling. Because the values for B_{CO_2} and $B_{\text{H}_2\text{S}}$ increase as the system cools, CO₂ and H₂S will partition into the vapor at a faster rate from isenthalpic boiling than from isothermal boiling. In open-system boiling, gases are removed from solution at exponential rates and cause marked changes in gas content of the liquid after only a few percent boiling. If open-system rather than closed-system conditions are assumed, less than 5 and 30 percent of the initial CO₂ and H₂S, respectively, remain dissolved in the liquid after just 3 percent boiling. At lower temperatures (higher B_i -values), less boiling is required to produce the same changes in the gas content of the liquid for both open and closed systems. In general, boiling is more effective in stripping CO₂ and H₂S from aqueous solutions the lower the temperature of boiling and the more open and isenthalpic the behavior of the system during boiling.

Our observations on SO₂, CO₂, and H₂S emissions during the 1992 eruptions of Crater Peak are consistent with the foregoing discussion of SO₂ hydrolysis and the partitioning of CO₂ and H₂S during boiling. We suggest that scrubbing of SO₂ by liquid water in the volcano effectively masked noneruptive magmatic emissions of SO₂ from the time of precursory seismic activity in August of 1991 to September 24, 1992 (table 1). We particularly stress this interpretation for the COSPEC measurements from June 8 to September 25, 1992, which show SO₂ emissions of background to only 88 t/d at times before, after, and between the three eruptions. At these times, SO₂ emissions of at least several hundred t/d would normally be expected on the basis of past experience (Casadevall and others, 1981, 1983, 1994; McGee, 1992). In addition, the hydrolysis of SO₂ accounts for the increase in the sulfate content of the Crater Peak lake water prior to the first eruption. The persistently strong H₂S odor during this period, and during later periods when SO₂ emissions were at or near background levels, is also consistent with this interpretation. In addition to direct dissolution of H₂S degassed from shallow magma, the hydrolysis of aqueous sulfur dioxide by reaction (1) converts 25 percent of the degassed SO₂ into aqueous H₂S. Because magmas generally emit much more SO₂ than H₂S, the production of H₂S by hydrolysis probably was significantly greater than the H₂S from magmatic degassing. Aqueous H₂S derived both from direct magma degassing and from hydrolysis of degassed SO₂ would be preferentially partitioned into the vapor phase and released in emissions during boiling of water. Hydrogen sulfide has a lifetime in the atmosphere of only about 1 day, because of oxidation reactions (Graedel, 1977). Thus, while it is possible that the minor (5–88 t/d) SO₂ fluxes during the June 8 through September 25 period represent magmatic degassing of SO₂ through rare dry pathways to the surface at this time, we cannot rule out the atmospheric oxidation of H₂S emissions from boiling water as the source of the detected SO₂.

Large SO₂ emissions were possible during the three explosive eruptions because magma ascended rapidly to the surface and could degas directly to the atmosphere. This effectively prevented contact with liquid water and loss by scrubbing of most of the SO₂ released during the eruptions. The TOMS data indicate that 15 to 25 percent of the sulfur released in the three eruptions was emitted as H₂S (Bluth and others, this volume). This suggests the possibility that H₂S emissions also occurred from the boiling of water during the explosive events. Hence some of the sulfur released in the explosive eruptions may have been degassed noneruptively as SO₂ at an earlier time, converted to aqueous H₂S by hydrolysis, and then re-

leased during boiling from heating associated with the ascent of the erupting magma.

On September 25, about a week after the September 16–17 eruption, the system finally dried out enough to permit SO_2 fluxes in excess of 100 t/d (table 1). The week of time may have been necessary to transfer the heat required to dry out a sufficient volume of the rock surrounding the shallow magma so that significant amounts of SO_2 could degas and escape to the atmosphere without encountering liquid water. It is possible that some additional shallow magma was supplied during this week; this magma augmented the supply of magmatic gas and heat needed to create and maintain a dry escape route to the surface. We suggest that during this period the liquid-vapor boiling surface migrated outward from the shallow magma, because the rate at which heat was supplied for boiling outpaced the rate at which recharging water removed heat.

The CO_2 emission rate data (table 1) indicate that magma was degassing, and these data support the hypothesis of SO_2 loss by scrubbing. Together, the CO_2 and SO_2 fluxes from September 25 to about October 10 indicate that magma was degassing through boiling water as well as through a zone of dry rock. The values of CO_2/SO_2 weight ratios for magmatic gases from convergent-plate volcanic systems are less than 15, except for gases collected from degassed domes several years after emplacement (Symonds and others, 1994). Most of the CO_2/SO_2 values at Crater Peak are in the 10 to 100 range (table 1). A higher CO_2/SO_2 ratio is expected if a significant fraction of the degassing takes place through boiling water. Magma degassing through boiling water would cause relatively minor fractionation of CO_2 , because of its tendency to partition preferentially into a vapor phase, compared to its effect on SO_2 , which is susceptible to hydrolysis. Consequently, the CO_2/SO_2 values of the gases emitted to the atmosphere will be higher than the magmatic gas values, and they will increase as the fraction of degassing through boiling water increases relative to degassing through a zone of dry rock. The high CO_2 emission rates of 11 and 12 kt/d at times of relatively modest (190 to 300 t/d) SO_2 emissions and the high CO_2/SO_2 values of 53 to 90 on September 25 and 28 suggest that significant degassing of magma through water was also occurring in the week immediately after the September 16–17 eruption, even though SO_2 fluxes were then very low (<100 t/d), presumably because of SO_2 scrubbing by boiling water.

Sulfur dioxide emission rates peaked near the end of September, only about a week after the dry-out zone was established. By early October, removal of heat by recharge of meteoric water apparently be-

gan to outpace the supply of magmatic heat again, the dry escape route through the wet edifice began to close, and SO_2 emissions started to drop. The exhaustion of gas, as well as heat, from the shallow magma probably contributed to the drop in SO_2 emission rate. The coincident decline in CO_2 also suggests the exhaustion of volatiles from the shallow magma.

We attribute the drop in SO_2 emission rates during tremor episodes in early October (fig. 6) to the recharge of liquid water back into the dry-out zone around and above the shallow magma. As water invaded fractures within this zone, vigorous boiling ensued, presumably giving rise to tremor, and the adjacent hot rock and melt cooled sufficiently to permit the presence of an increasing amount of liquid water. As the process proceeded, SO_2 was increasingly lost by scrubbing, SO_2 emissions dropped, and CO_2/SO_2 and $\text{H}_2\text{S}/\text{SO}_2$ values rose. After the tremor episode of October 12, the system remained effectively waterlogged without a dry zone for SO_2 escape. Sulfur dioxide emissions declined to background levels where they have remained.

Carbon dioxide emissions of greater than 2 kt/d suggest that episodes of boiling continued to release CO_2 and H_2S after the October 12 tremor episode. Carbon dioxide emissions and strong H_2S odors were evident again 3 days after the November 9 earthquake swarms. We speculate that these swarms were related to one or more intrusions too small to penetrate through the water-saturated zone but sufficiently large to produce thermal perturbations that led to boiling and release of CO_2 and H_2S . Carbon dioxide fluxes of 1 to 3 kt/d continued for over 2 months. A measurement on April 24, 1993, still showed a flux of about 1 kt/d (table 1). We suspect these emission levels signify continuing restlessness because measurements on July 7, 1994, did not detect CO_2 above the background (Doukas, 1995).

CONCLUSIONS

Sulfur dioxide scrubbing by liquid water masked SO_2 emissions from shallow magma during the 1992 eruptions of Crater Peak and created the unusual SO_2 degassing patterns marked by the absence of exponentially declining fluxes after the three explosive eruptions. Hydrolysis reactions effectively prevented observation of SO_2 emissions from shallow magma both before and after explosive eruptions and seismic crises, except for a brief period when dry degassing pathways to the surface existed from September 25 to about October 10, 1992. Measurements of CO_2 emissions initiated on September 25, 1992, indicate that in addition to the degassing of magma through dry path-

ways, degassing through boiling water with the loss of SO_2 by scrubbing was also important at this time. The hydrolysis of SO_2 also explains the increase in the sulfate content of Crater Peak lake water prior to the first eruption, the strong H_2S odor during periods of background to minor SO_2 emissions, the TOMS evidence for significant H_2S emissions during the explosive eruptions, and the observed decline of SO_2 emissions during periods of volcanic tremor. Large SO_2 emissions unencumbered by scrubbing only occurred during the three explosive eruptions when magma penetrated through liquid water zones under Crater Peak and reached the surface. Abundant, local sources of melt water and a high permeability for the Mount Spurr volcanic edifice are probably the chief factors responsible for the masking of SO_2 emissions by scrubbing, and possibly for quenching shallow intrusions in the process of ascending to the surface.

We strongly recommend early monitoring of CO_2 when Cook Inlet volcanoes become restless. Because of its strong preference for the vapor phase during boiling, CO_2 emissions from degassing magma are less likely to be masked by the presence of water, whereas SO_2 emissions may be lost totally from interactions with water and thus render misleading COSPEC results. We also recommend that CO_2 emission rate data sets be obtained for inactive Cook Inlet volcanoes so that baseline fluxes will be available for comparison when these volcanoes become restless.

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Chemistry of Crater Lake Waters Prior to the 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

By Terry E.C. Keith, J. Michael Thompson, and Robert G. McGimsey

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ABSTRACT

Waters sampled from the shallow lake in the Crater Peak vent of Mount Spurr volcano during June of 1992 were much higher in SO_4 and lower in Cl than waters sampled at the same locality in August of 1970. The major change in water composition is an effect of the degassing of shallow magma into the volcano-hydrothermal system prior to the June 27, 1992, eruption. Though the timing of the chemical change is unknown, it may have been concurrent with a change in the lake observed in early June of 1992. At that time the lake changed from clear, blue-green to turbid, battleship gray. Volcanic unrest was also accompanied by upwellings and bubbling in the lake prior to its disappearance owing to evaporation and (or) draining.

INTRODUCTION

A nearly circular lake, approximately 100 m long, developed in the bottom of the satellitic Crater Peak vent on the south flank of Mount Spurr at an unknown time following the eruption of July 9, 1953. Air photos taken in September 1952 show ice filling the crater (Motyka and Nye, 1993) 10 months prior to its first known historical eruption. Careful study of tephra from Mount Spurr volcano show that the Crater Peak vent has erupted 35 times in the past 6,000 years, including 15 times within the past 500 years (Riehle, 1985). The lake in Crater Peak was visited

and its water sampled by Bruce Reed and Marvin Lanphere during U.S. Geological Survey (USGS) mapping in the region during the Summer of 1970. Fumaroles at Crater Peak were sampled in 1982 and also visited in 1985, but in 1985 they were too diffuse to sample (Motyka and Nye, 1993). No further sampling of the lake water or fumaroles was done until 2 weeks prior to the June 27, 1992, eruption.

Thermal springs low on the west flanks of Mount Spurr volcano below Crater Peak were discovered and sampled in 1985 by Motyka and Nye (1993). Unfortunately, other priorities precluded sampling these springs during the 1991–92 period of unrest and eruption of the volcano. They were resampled on June 2–3, 1993 (R. Motyka, unpub. data, 1993).

OBSERVATIONS AND SAMPLING OF CRATER LAKE WATERS

1970 SAMPLING AND OBSERVATIONS

Reed and Lanphere visited the Crater Peak lake on August 9, 1970, and collected two water samples for analysis (table 1). They were not equipped to measure temperature or pH. Reed's field notes (unpub. data) state that "both samples from approximately 6 inches depth; 2.5 to 3.5 ft from shore; water is milky gray, too hot to hold hand in; had difficulty collecting the samples because of the temperature—could hold hand in for 1–2 seconds, very hot, water is being diluted by snow meltwater . . ." These observations suggest temperatures in the range 55 to 65°C.

1992 SAMPLING AND OBSERVATIONS

Observations and photographs from fixed-wing aircraft overflights in May and the first week of June 1992 recorded the lake color as a clear blue-green, as had been observed during previous years. On the observation flight of June 8, however, the lake was reported as being a gray color with several small areas of upwelling or bubbling, the most prominent of which

occurred on the northeast side of the lake. During this flight, the usual scattered, wispy fumarolic activity was noted, and the most vigorous fumarole was emitting a strong steam plume from a talus pile adjacent to the lake on the west side of the crater floor.

On June 11, a descent was made into the Crater Peak vent for general observations. The lake was a murky, battleship gray color. Continued small upwellings were observed; the measured temperature of the lake water (49°C) indicated that the upwellings were caused by rising noncondensable gases, rather than

boiling. Geyser activity, previously unreported, was observed in the talus pile adjacent to the north edge of the lake where the most vigorous fumarole in the Crater Peak vent had been active for at least several months. Motyka and Nye (1993) reported the most vigorous fumarolic activity as being on the east side of the lake in 1982 and on the north side in 1985. The fairly rapid migration of active fumarolic vents suggests redirection of fumarolic channels by shallow seismic disturbances; local self-sealing of fumarolic channels by hydrothermal alteration must have occurred to some extent but probably was not the controlling factor in this case.

On the June 11 trip, two whole-water samples were taken from the lake. One sample (CP-1) was taken adjacent to the shore and the other (CP-2) was taken about 0.5 m out from shore. The samples were brought back to Anchorage approximately 4 hours following collection. Half of each sample was filtered and acidified. Chemical and isotopic analyses (table 1) were made following the procedures of Thompson (1990).

The pale gray, very fine-grained filtrate was analyzed for mineral phases using X-ray diffraction. Phases identified were poorly crystalline sulfur, alunite, pyrite, and kaolinite. Following the mineral analysis, the same filtrate was chemically analyzed using semi-quantitative emission spectrographic methods to determine minor-element concentrations (table 2).

Two weeks later, on the afternoon of June 26, geologists from the Alaska Division of Geological and Geophysical Surveys who were returning from fieldwork by helicopter passed by Crater Peak and observed that the shallow crater lake had "drained." Gray mud covered the crater floor, and a substantial steam cloud rose from the crater. Lake waters evaporated when heated by the rising magma in addition to lake water draining into the subsurface along fractures caused by seismic disturbances in the volcanic edifice.

ISOTOPES

Oxygen and deuterium isotopic analyses were obtained only on the 1992 lake water samples (table 1). Comparison of these data to the World Meteoric Water Line (Craig, 1961) and to data presented by Motyka and Nye (1993) for meteoric waters from Spurr and Augustine volcanoes show that the Crater Peak lake waters of June 8, 1992, fall well off the meteoric water line (fig. 1). In a hydrothermal system, only oxygen shifts to heavier values during water-rock interaction but deuterium remains stable (Craig, 1966). Evaporation shifts both δD and $\delta^{18}O$ to heavier values. Thus, isotopic values for the Crater Peak lake waters indicate that evaporation dominated any water-rock interaction. The isotope values for the Crater Peak lake waters could also be interpreted to be a

Table 1. Chemical and isotopic analyses of Crater Peak (Mount Spurr) lake waters, Alaska.

	*CP92-1	*CP92-2
Conditions		
T°C -----	49	49
field pH -----	2.5	2.5
lab pH -----	2.48	2.49
conductance -----	3130	3120
Chemical species, in milligrams per liter		
SiO ₂ -----	263	312
Al -----	41.8	50
Fe -----	33.2	31.6
Mn -----	2.95	2.94
As -----	0.16	0.26
H -----	7.21	8.18
Ca -----	184	187
Mg -----	40.5	41.3
Na -----	56.0	55.1
K -----	3.51	3.69
Li -----	0.034	0.035
HCO ₃ -----	0	0
SO ₄ -----	1450	1490
Cl -----	53.2	54.5
F -----	0.42	0.42
B -----	3.36	3.18
SO₄:Cl ratio		
SO ₄ :Cl -----	27.3	27.3
Isotopic hydrogen and oxygen composition		
**dD -----	-116	-113.5
**d18O -----	-7.2	-7.2
***70AR-201 (south edge) ***70AR-202 (north edge)		
Conditions		
T°C -----	nd	nd
field pH -----	nd	nd
lab pH -----	2.17	2.30
Chemical species, in milligrams per liter		
Ca -----	192	182
Mg -----	71	70
Na -----	75	75
K -----	7.5	7.3
HCO ₃ -----	0	0
SO ₄ -----	442	315
Cl -----	1180	1100
B -----	24	24
SO₄:Cl ratio		
SO ₄ :Cl -----	0.4	0.3

*analyst: J. Michael Thompson, USGS, Menlo Park, CA

**analyst: L.D. White, USGS, Menlo Park, CA; values relative to SMOW

***unpublished analyses of water samples collected by Marvin Lanphere and Bruce Reed (USGS) on 8/9/70 were provided by Bruce Reed; analyst: Roberta Barnes, USGS, Menlo Park, CA; nd, not determined

Table 2. Semiquantitative emission spectrographic chemical analysis of minor elements in filtrate from Crater Peak (Mount Spurr, Alaska) lake waters of June 11, 1992.

[Analyst, E.A. Bailey; ppm, parts per million]

element, in weight percent	
Ca	1.
Fe	5.
Mg	1.
Na	1.5
P	0.2
Ti	0.7
element, in ppm	
Ag	1.5
As	<200
Au	<10
B	50
Ba	>5,000
Be	<1
Bi	<10
Cd	~20
Co	20
Cr	70
Cu	150
Ga	30
Ge	<10
La	50
Mn	300
Mo	<5
Nb	<20
Ni	50
Pb	20
Sb	<100
Sc	50
Sn	20
Sr	1000
Th	<100
V	200
W	<20
Y	10
Zn	<200
Zr	70

result of meteoric (hydrothermal) waters mixing with magmatic waters, using the field of Sheppard and others (1969).

SOLID PHASES

Crater Peak vent andesitic rocks, in contact for years with warm, acidic lake waters of the 1970 composition (table 1) and vent fumaroles (Motyka and Nye, 1993), must have altered somewhat to common acid-sulfate products such as alunite, kaolinite, pyrite, Fe oxides and hydroxides, as well as residual noncrystalline to poorly crystalline silica phases. The fine-grained gray mud of the filtrate is similar in appearance and mineralogy to that in areas of characteristic acid-sulfate alteration, including mudpots of Yellowstone National Park (White and others, 1988), acid altered andesite at White Island, New Zealand (Giggenbach, 1987), and the gray muds in the Ruapehu volcano crater lake (Christenson and Wood, 1993). Upwelling disturbance of the lake waters in Crater

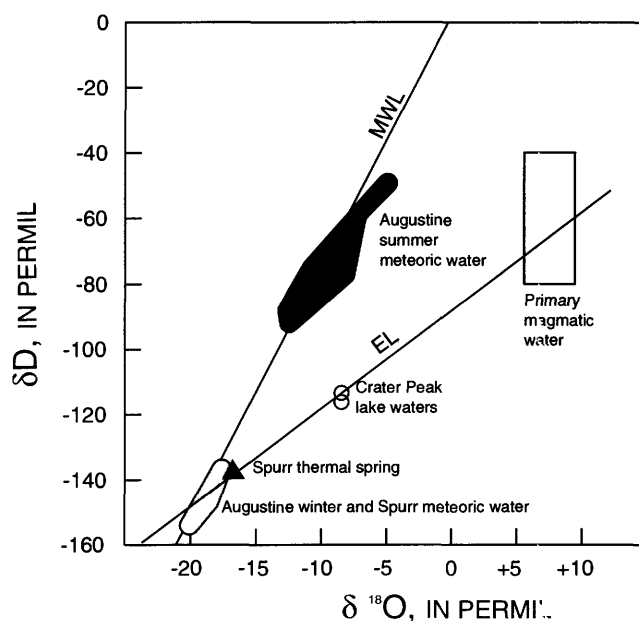


Figure 1. Plot of $\delta^{18}\text{O}$ versus δD of waters from Mount Spurr and Augustine volcanoes compared to the World Meteoric Water Line (MWL) of Craig (1961) and an evaporation line (EL) for the Crater Peak (Mount Spurr) lake waters. Data for Augustine summer meteoric waters, Mount Spurr meteoric waters (including a cold stream, snowmelt), and Augustine winter meteoric waters, and the triangle representing the thermal spring on the lower south flank of Mount Spurr below Crater Peak from Motyka and Nye (1993). Data for primary magmatic waters from Sheppard and others (1969). Open circles, Crater Peak thermal lake waters of June 11, 1992.

Peak suspended the fine-grained alteration products in the water. Thus the poorly crystalline filtrate minerals (alunite, pyrite, kaolinite) identified by X-ray diffraction are residual rock alteration products, and the sulfur is from the oxidation of increased magmatic H_2S in the lake water. Although silica phases were not detected by X-ray diffraction and the amount of filtrate was too small to analyze for silica, some form of silica was probably present because the lake waters contained substantial dissolved silica.

DISCUSSION

The striking difference in chemical composition between the two sets of water samples is the high $\text{SO}_4:\text{Cl}$ ratio in samples taken in 1992 relative to those taken in 1970. The increased SO_4 in water sampled in June 1992 is most likely a result of increased sulfur species degassing from rising magma in the Crater Peak vent and disproportionation of magmatic SO_2 in the hydrothermal system and lake waters to form

aqueous H_2S and sulfate, which are much more stable in water than SO_2 (Giggenbach, 1987; Doukas and Gerlach, this volume). The Cl concentration in the crater lake water decreased in 1992 relative to 1970. The 1992 decrease in Cl was apparently simultaneous with increased sulfur degassing from the magma. Generally, the concentration of dissolved Cl in the volcano-hydrothermal system and (or) the overlying shallow lake water would be expected to increase as a result of magma degassing into water.

The relevant question may be why Cl was so high in 1970 rather than why Cl was so low in 1992. One possibility is that the subsurface hydrothermal system was not well developed in 1970 following the 1953 eruption, and so Cl from degassing magma could reach the surface waters, although the large amount of precipitation in the area should have reestablished the hydrothermal system very quickly. The relatively high concentration of B in the 1970 lake waters supports the hypothesis of a well-developed hydrothermal system at that time. High Mg concentrations of the lake waters in both years suggest substantial water-rock interaction between the acidic lake water and the andesitic host rock of the volcanic edifice. Perhaps Cl was more concentrated in the August 1970 lake waters because of evaporation of thermal water (where Cl remained dissolved) accompanied by the precipitation of sulfate as mineral phases, thus being removed from the lake water.

The volcano-hydrothermal system within the south flank of Mount Spurr volcano and beneath the Crater Peak vent was undoubtedly disturbed by the intrusion of magma leading to the June 27, 1992, eruption. As the magma began its ascent, the water of the hydrothermal system was further heated and boiling increased; these processes released H_2O , O_2 , and H_2S , which condensed in the lake and surface waters. Accordingly, the lake waters sampled 16 days prior to eruption in 1992, are probably a mixture of pre-existing lake, magmatic, hydrothermal, and meteoric waters as well as some component of dissolved magmatic gases including sulfur species, Cl, and F. The volcano-hydrothermal system has since been severely disrupted by the three 1992 eruptions and the non-eruptive intrusions in late 1992 (Eichelberger and others, this volume).

The changes in character of the lake in the Crater Peak vent prior to the June 27, 1992, eruption are probably a result of major disturbances within the shallow part of the volcano. Similar changes were reported in 1985 at Ruapehu volcano, New Zealand (Scott, 1987; Christenson and Wood, 1993) during heating episodes, some of which resulted in eruption. Crater lake disturbances such as color change and upwellings are frequently associated with increased heating by rising magma.

The fact that ice occupied the Crater Peak vent during September 1952 prior to the July 1953 eruption has interesting implications for monitoring Mount Spurr volcano. The Crater Peak vent has a frequent eruptive pattern. Riehle (1985) has documented 16 eruptions during the past 500 years, and this is a minimum of events because small eruptions that left no tephra records could also have occurred. Observations on the status of the ice in the crater or the development of a lake prior to the 1953 eruption are not known to exist. The development and maintenance of a warm crater lake in the vent between the 1953 and 1992 eruptions suggests that the volcano did not return to the level of inactivity that existed prior to the 1953 eruption, which was cool enough to permit ice to form in the crater. The presence or absence of ice in the crater and the characteristics and chemistry of any crater lake are important indicators of the future status of unrest. Sampling of flank thermal springs during periods of unrest and eruptions might indicate how well they are connected to the hydrothermal system near the vent.

CONCLUSIONS

The presence of a warm lake in the Crater Peak vent between its 1953 and 1992 eruptions indicates that the volcano had not returned to its pre-1953 state of inactivity when conditions were cold enough that ice occupied the crater only 10 months prior to eruption. Variations in crater lake appearance (color and turbidity) and chemical composition, especially the SO_4/Cl ratio, can be an indicator of volcanic unrest. Such variations may signal the intrusion of new magma. At Crater Peak, the lake disturbances could indicate imminent potential for magmatic or phreatic eruption. Although sampling of the lake waters in the crater may be useful, safety factors often preclude it.

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Tephra-Fall Deposits from the 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska: A Preliminary Report on Distribution, Stratigraphy, and Composition

By Christina A. Neal, Robert G. McGimsey, Cynthia A. Gardner, Michelle L. Harbin, and Christopher J. Nye

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ABSTRACT

Most ejecta from the 1992 eruptions of Crater Peak was deposited as tephra fall from eruption columns that ascended rapidly to altitudes of greater than 14 km above sea level. Coarse bomb- to fine ash-sized tephra fell along relatively narrow trajectories extending north, northeast, and east from the vent; these patterns reflect prevailing wind directions in upper Cook Inlet. Each brief (3.5- to 4.0-hour), small-volume (12 to 15 x 10⁶ m³ dense rock equivalent or DRE) eruption involved two distinct juvenile andesitic tephra components that differ in color and density only; no systematic differences in glass chemistry, bulk composition, or mineralogy have been identified. The stratigraphic relations of these components in the August 18 and September 16-17 deposits suggest that the upper part of an inferred pre-eruption magma body was enriched in volatiles relative to the lower part. Clast-

size grading may reflect a progressive increase in magma discharge during the first 2 to 3 hours of tephra emission.

INTRODUCTION

After 39 years of quiescence, the Crater Peak vent on the south flank of Mount Spurr volcano reawakened with a series of three short lived but violent eruptions of andesitic tephra on June 27, August 18, and September 16-17 of 1992 (Eichelberger and others, this volume; Alaska Volcano Observatory, 1993). Each eruption lasted 3.5 to 4.0 hours and generated vulcanian to subplinian eruption columns that reached altitudes of more than 14 km above sea level (ASL). Carried downwind, these tephra plumes resulted in narrowly distributed fall deposits composed principally of poorly vesiculated, crystal-rich juvenile andesite that ranged in size from coarse bombs to fine ash. Plumes from the three eruptions extended north, northeast, and east away from the volcano. Two of the three 1992 events deposited significant amounts of fallout tephra onto the most heavily populated areas of south-central Alaska (fig. 1A).

Tephra-fall deposits make up the bulk of the eruptive volume for each event (12, 14, 15 x 10⁶ m³ DRE for the three eruptions) and hence contain critical information regarding temporal changes in chemistry, eruption dynamics, and plume behavior. Each eruption also produced small pyroclastic flows and lahars, however, together these deposits make up less than 1 percent of the total bulk volume for the entire eruption series (Miller and others, this volume; D. Meyer, oral commun., 1994). This paper summarizes field characteristics and preliminary compositional analyses of proximal tephra-fall deposits, and it also proposes a model to explain observed physical and sedimentologic variations.

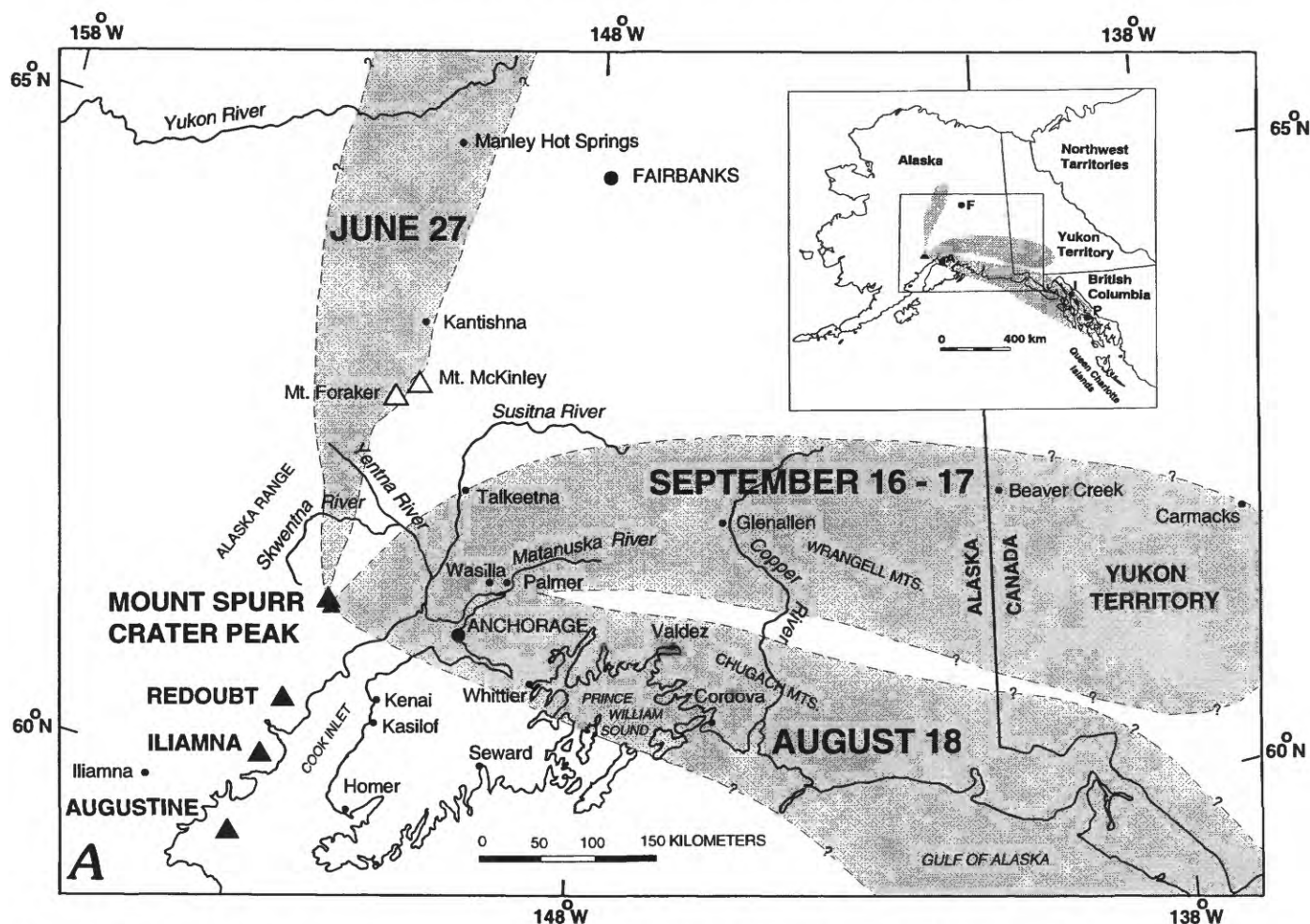


Figure 1. A, Generalized map showing extent of the three tephra-fall deposits from the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska. Shaded areas denote approximate limits of tephra-fall inferred from ground and aerial surveys within 250 km of the vent as well as phone contact with distant communities within several days of the eruptions. Due to very sparse population and poor road access, the actual downwind extent of each deposit is unknown.

ACKNOWLEDGMENTS

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DISTRIBUTION OF TEPHRA-FALL DEPOSITS

Prevailing winds in upper Cook Inlet (fig. 1B) blow from the southeast, southwest, and northwest; the dominant wind direction is from the south to southwest. Consequently, tephra erupted from any Cook Inlet volcano is likely to be carried over populated areas and into the busy air traffic corridor of south-central Alaska; this phenomenon underscores the need for timely eruption notification and tracking of tephra plumes. As part of eruption-response procedures at the Alaska Volcano Observatory (AVO), hypothetical plume trajectory plots provided by the NWS were generated at AVO (Brantley, 1990, fig. 32). Based on twice-daily atmospheric soundings of wind velocity at various altitudes, these maps are used to forecast probable dispersal patterns of tephra plumes and resultant fallout

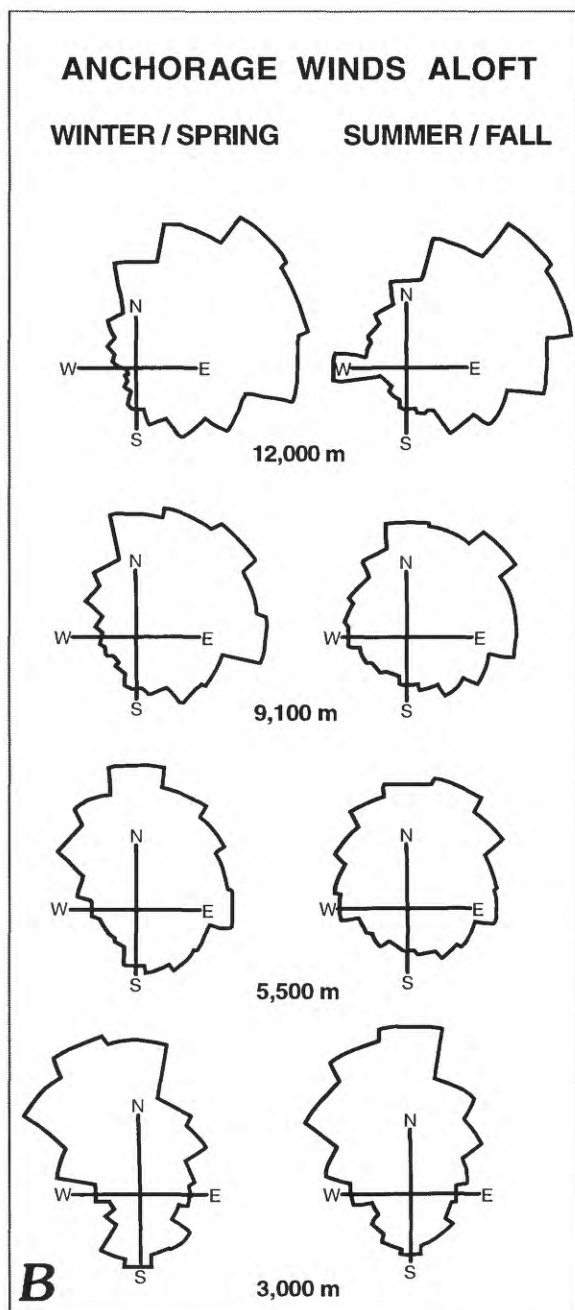


Figure 1. Continued. *B*, Compass rose diagrams indicate percent of time that winds blow toward a given direction over Anchorage (adapted from Till and others, 1993, and based on data from the National Climatic Data Center, Asheville, North Carolina for the period 1948 to 1972).

(Murray and others, 1994). All three tephra plumes produced during the 1992 eruptions at Crater Peak followed trajectories in good agreement with these plots.

The June 27 eruption occurred during a period of strong southerly winds and heavy overcast. As determined by C-band (5.4-cm-wavelength) radar operated by NWS, the vertical eruption column reached a maximum altitude of 14.5 km ASL by 10:23 a.m. Alaska Daylight Time (ADT), approximately 3.3 hours into the eruption (Rose and others, this volume). The tephra plume traveled north from Crater Peak (fig. 1A) over the relatively unpopulated Yentna and Skwentna River basins, the Alaska Range near Mount Foraker, and the settlement of Kantishna. The most distal report of ash fall was received from Manley Hot Springs, 140 km west of Fairbanks and 420 km north of Crater Peak. To the north of this community, there are few settlements.

During the August 18 eruption, winds were westerly and the axis of the main tephra fall extended east from Crater Peak, across Cook Inlet and over the Anchorage basin where 1 to 3 mm of primarily coarse ash was deposited. The plume continued southeastward over Prince William Sound, the eastern Chugach Mountains, and the Gulf of Alaska (fig. 1A). Petersburg, Alaska, 1,200 km southeast of Crater Peak (fig. 1A, inset) reported a light dusting of fine ash on August 19. Farther south, in the Queen Charlotte Islands and along coastal British Columbia, Canadian Flight Service Station personnel noted only the passage of a hazy layer overhead. In distribution, this fallout pattern is similar to that of the 1953 Crater Peak eruption; tephra accumulation in Anchorage, however, was greater in 1953 (3–6 mm; Wilcox, 1959). The maximum radar-discernible plume height during the August 18, 1992, eruption was 13.7 km ASL at 4:55 p.m. ADT (Rose and others, this volume), only 14 minutes after the onset of the main phase of the eruption. Pilot estimates of the top of the visible plume suggested actual heights may have been in excess of 18 km.

The September 16–17 eruption occurred during southwesterly winds and the plume traveled northeast over the communities of Wasilla and Palmer in the Matanuska-Susitna valley, parts of the Talkeetna Mountains, the Copper River Basin, and east into the Yukon Territory of Canada (fig. 1A). The farthest reported ash fall was in Carmacks, Yukon Territory, 860 km from Crater Peak. Light and variable low-level winds on September 16 and 17 accounted for numerous reports of very light dustings of ash as far south as Anchorage. Maximum radar-discernible plume height was 13.9 km ASL by 2:21 a.m. ADT, September 17 (Rose and others, this volume), more than 2 hours after the eruption began. Because the eruption occurred at night, there were few pilot reports of column height to corroborate this estimate.

VOLUME OF TEPHRA-FALL DEPOSITS

Helicopter-supported mapping and road access to large areas affected by the August 18 and September 16–17 eruptions allowed for rapid measurement of tephra-fall deposit thicknesses within south-central Alaska. Most data and samples were gathered within several days to 1 week of the eruption. With the exception of some wetting and compaction owing to rain-fall or incorporation of water from melting snow and ice substrate, no significant reworking of the deposits was evident during collection of most samples. For each of the three tephra-fall deposits, we collected between 5 and 50 measured area samples (typically 20 x 20 cm) along and across the deposit axis as far as 190 km from Crater Peak. Based on mass per unit area data (Scott and McGimsey, 1994; Fierstein and Nathenson, 1992), the DRE volumes (using $2,600 \text{ kg/m}^3$) calculated for each event are $12 \times 10^6 \text{ m}^3$, $14 \times 10^6 \text{ m}^3$, and $15 \times 10^6 \text{ m}^3$ for the June 27, August 18, and September 16–17 eruptions, respectively. Bulk volumes calculated using the same method and a density of 700 kg/m^3 are $44 \times 10^6 \text{ m}^3$, $52 \times 10^6 \text{ m}^3$, and $56 \times 10^6 \text{ m}^3$.

THICKNESS OF TEPHRA-FALL DEPOSITS

Stable wind conditions during each brief eruption contributed to fairly uniform changes in thickness and grain size with increasing distance from the vent. Proximal tephra-fall deposits for each eruption range similarly in thickness from about 1 m at the base of Crater Peak (approximately 1 km from vent) to 10 cm at a 10-km distance (fig. 2). At a distance of 100 km, the total ash fall is only a few millimeters thick. Over the same distance, modal grain size varies roughly from cobble gravel near the vent to coarse sand at 100 km. Additional grain-size analyses are in progress to further characterize the deposits.

COMPOSITION OF TEPHRA-FALL DEPOSITS

Most tephra from each eruption consists of juvenile andesite fragments of two distinct colors, tan and gray. For the August 18 and September 16–17 tephra sequences, lower and upper stratigraphic units (layers A and B, respectively) reflect a distinct change in the proportion of the two colors (table 1). Overall, the deposits are volumetrically dominated by tan to light-brown, slightly pumiceous, angular, equant to pyramidal fragments of crystal-rich andesite. Many clasts were fractured during ejection or deposition and have one or more flat or slightly concave faces. Lapilli

(1–3 cm) commonly show gradational color variations from tan exteriors to dark-brown interiors. Inclusions of vesiculated basement clasts; dense gray, aphyric volcanic lithics; and rounded, sugar-textured, plagioclase-pyroxene crystal clots are common and are similar to inclusions found in the breadcrusted bombs of the pyroclastic-flow deposits (Miller and others, this volume; Nye and others, this volume; Harbin and others, this volume). Second in abundance are light- to medium-gray, microvesicular, crystal-rich andesite clasts that predominate in the upper part of the August 18 and September 16–17 tephra sections. These clasts are smaller in grain size and denser (table 2) than accompanying tan clasts, but they contain similar lithic, crystal-clot, and exotic pumiceous inclusions. They are angular to subrounded and generally equant in shape.

Accidental clasts make up a small volume of the tephra-fall deposits. These clasts consist of pre-1992 Crater Peak volcanic rocks recognized on the basis of alteration, fragments of coarse-grained metamorphic and possibly plutonic basement, crystal clots, and vesiculated clasts of basement rock (table 1; Harbin and others, this volume; Miller and others, this volume). A representative proximal sample of each tephra-fall unit was selected, and between 900 and 3,400 clasts were sorted and counted for each sample; percentages listed below and in table 1 represent the proportion of counted clasts. Fragments were classified on the basis of gross visual characteristics. Although these counts represent only a limited view of each deposit, we noted no field evidence of gross spatial changes in deposit character, and hence consider these percentages to be good first-order representations of each unit.

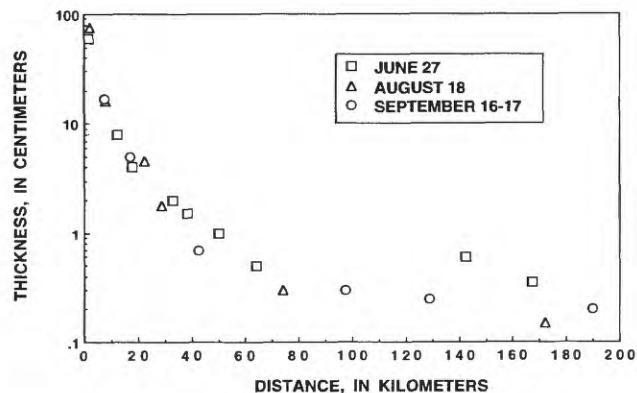


Figure 2. Thickness of tephra-fall deposits from the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska, measured downwind along an approximate axis of each plume. The two most distant points for June 27 tephra were measured in the late summer of 1993 and are considered anomalously thick owing to percolation of the fine sand-sized tephra across a recrystallized snow layer.

Table 1. Percentage of clast types for tephra-fall deposits from the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska.

[Populations determined by hand picking from size fraction larger than -1.75 phi (approximately 3 mm). Similar calculations based on weight or volume of each clast type produce nearly identical results. See text for description of each clast type.]

	Tan andesite	Gray andesite	Pre-1992 volcanics	Metamorphic and plutonic	Total clasts
June 27 -----	78	14	7	1	1761
August 18 layer A-----	70	23	6	1	3355
August 18 layer B-----	6	91	3	0	1067
September 16-17 layer A-----	79	15	5	1	2630
September 16-17 layer B-----	26	64	10	0	918

Table 2. Densities, (in g/cm³) for selected juvenile tephra clasts from the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska.

[Note the similarity in component densities for August and September. Overall, June clasts of both types are 10 to 15 percent denser than clasts from the other two eruptions. Densities were measured by M. Couchman following techniques described in Hoblitt and Harmon (1993).]

June tan	June gray	August layer A tan	August layer A gray	August layer B tan	August layer B gray	September layer A tan	September layer A gray	September layer B tan	September layer B gray
1.98	2.21	1.59	1.98	1.65	1.98	1.61	1.85	1.54	1.83
1.56	2.08	1.52	2.07	1.57	1.98	1.60	1.75	1.65	1.93
1.76	1.91	1.93	1.86	1.32	1.85	1.37	1.89	1.53	1.92
1.86	1.98	1.47	2.47	1.69	2.14	1.69	2.02	1.80	2.70
1.78	2.14	1.67	2.07	1.55	1.80	1.86	1.87	1.73	1.99
1.92	2.38	1.69	1.87	1.48	2.04	1.76	1.59	1.14	2.09
1.66	2.03	1.34	1.89	1.48	2.46	1.16	1.73	1.64	2.05
1.44	1.93	1.40	1.88	1.55	2.02	1.53	1.47		
1.93	2.72	1.87	2.07	1.53	1.93	1.71	2.41		
1.82	2.04	1.43	1.96	1.29	1.81	1.62	2.38		
1.86	2.69	1.41	2.36	1.62	1.95	1.88	2.61		
1.58	2.76	1.48	1.97	1.58	2.56	0.98	2.71		
1.58	2.45	1.57	2.48	1.47	1.81	1.59	2.25		
1.91	2.67	1.51	2.09	1.43	2.03	1.76	1.85		
1.87	2.10	1.70	2.02	1.43	1.96	1.33	1.79		
2.01		1.67	1.97	1.41	1.89	1.62			
1.91		1.62	2.33	1.31	2.10	1.41			
1.79		1.38	1.88	1.61	1.90	1.33			
1.43		1.45	1.94	1.55	2.01	1.25			
1.49		1.38	1.86	1.55	2.01	1.51			
Average densities									
1.76	2.27	1.55	2.05	1.50	2.01	1.53	2.01	1.58	2.07

STRATIGRAPHY OF TEPHRA-FALL DEPOSITS

JUNE 27

The June 27 tephra-fall deposit is light gray to tan in color, moderately well sorted, and shows weak reverse grading. Angular clasts of tan to light-brown, poorly vesicular, porphyritic juvenile andesite are the dominant component (78 percent; table 1), followed by a smaller population (14 percent) of light-gray, denser, and more finely crystalline andesite fragments. The color and textural variation between the two components is subtle and quite difficult to discern in clasts smaller than about 5 mm across. Pre-1992 volcanic clasts, some altered to a brick-red color, represent 7 percent of the deposit. Nonvolcanic fragments are a

minor component (1 percent). Dry-bulk densities of five samples with median grain sizes of 0.5 to 4.0 mm range from 1.07 to 1.32 g/cm³. The tan and gray juvenile clasts from the most proximal sample have average densities of 1.76 and 2.27 g/cm³, respectively (table 2).

Less than 1 percent of the June 27 tephra consists of subangular to subrounded clasts of light-tan, light-gray or white, sugar-textured, and pumiceous (density less than 1.0 g/cm³) mixtures of feldspar, quartz, scattered mafic minerals and silica-rich glass, here referred to as "silicic pumice" (Nye and others, this volume). These exotics represent entrained basement rock (Miller and others, this volume; Harbin and others, this volume). They are found predominantly on the surface of the June 27 (and subsequent) tephra-fall sequences, which probably reflects their low density and longer residence time in the tephra plume.

AUGUST 18

The main August 18 tephra-fall sequence (fig. 3) consists of a weakly developed, fine-grained basal unit (Layer A'), a moderately well sorted, light-gray to tan, reversely graded unit (layer A) that represents from 60 to 75 percent of the total thickness of any given section, and an upper, well sorted, dark gray, ungraded unit (layer B). Layers A and B are discernible in proximal sections and in thin (less than 2 mm thick) distal tephra deposits as much as 170 km east of Crater Peak.

At three localities 6 to 14 km downwind from Crater Peak, a discontinuous, thin (less than 1–2 mm) dark-gray, medium-grained sand horizon is present beneath layer A. This layer (A') was too thin to sample and was not observed with confidence at all August 18 tephra sections.

Approximately 70 percent of fragments within layer A consists of juvenile, tan, slightly pumiceous, phenocryst-rich clasts. Twenty-three percent of layer A consists of smaller juvenile fragments of andesite that are light to medium gray, microvesicular to dense, and contain sparse lithic inclusions. Clasts that contain both tan and gray components are rare, and in these clasts the boundaries between the tan and gray components are sharp and easily recognized on the basis of color and vesicularity. In addition to the tan and gray clasts, there is a small (approximately 6 percent) population of altered volcanic rock fragments

and a small number of white, dense metamorphic clasts. No silicic pumice has been recognized in layer A samples. Dry-bulk densities of five samples range from 0.96 to 1.09 g/cm³ for layer A. The tan and gray juvenile clasts from layer A have average densities of 1.55 and 2.05 g/cm³, respectively (table 2).

Most fragments (91 percent) within layer B consist of juvenile, microvesicular to dense, light to medium gray andesite that are similar to the gray fragments of layer A. Tan clasts similar to the dominant component of layer A make up only 6 percent of layer B. Three percent of layer B consists of pre-1992 volcanic clasts and there are only rare fragments of dense, white metamorphic rock. Scattered on the surface of layer B are subangular to subrounded clasts of light-tan to white silicic pumice fragments. Layer B is finer grained at a given distance from the vent than layer A, and it has a dry-bulk density of 1.13 to 1.34 g/cm³, which is approximately 30 percent higher than that of layer A. Average tan and gray clast densities are 1.50 and 2.01 g/cm³, respectively (table 2).

In addition to the continuous tephra sequence described above, a prominent ballistic impact zone occurs along the southern margin of the proximal August 18 tephra fall. Extending as much as 10 km from the vent (Waitt and others, this volume), this shower of juvenile and accidental blocks as large as 1 m across occurred primarily late in the eruption, following the main phase of tephra emission (Miller and others, this volume; Waitt and others, this volume).



Figure 3. A 16-cm-thick tephra-fall deposit from the August 18, 1992, eruption, located approximately 6.4 km east of Crater Peak, Mount Spurr volcano, Alaska. Layers A and B are labeled. Note reverse grading of layer A. Section rests on sandy tephra from a 1990 eruption of Redoubt Volcano (light horizon; see arrow). Spatula is 23 cm long.

SEPTEMBER 16–17

The September 16–17 fallout deposit contains three layers (layers A', A, and B) similar to those for the August 18 tephra, however the color change that marks the layer A-B transition is not as visually distinct (fig. 4).

At one locality 6.5 km from the vent, a coarse sand unit (layer A') was noted beneath the main sequence. A sample of this unit consists of coarse sand-sized, crystal-rich, dense to slightly pumiceous fragments. Both tan and gray colors are represented, estimated visually to be about 60 and 30 percent, respectively. A significant (approximately 10 percent) number of particles are strongly oxidized brick red to purple in color, phenocryst-rich volcanic rock fragments or white to gray crystalline, possibly plutonic or metamorphic fragments.

Layer A makes up the lower 70 to 80 percent of the tephrafall, is moderately well sorted, reversely graded, and composed of 79 percent very light tan, slightly pumiceous, phenocryst-rich clasts similar to those in the August 18 tephra. Some of these clasts show a slight yellow tint and are among the most pumiceous of Crater Peak juvenile andesite tephra; a few are highly pumiceous and probably represent the "light-green-gray andesite" (a slight differentiate of the juvenile andesite) of Harbin and others, (this volume).

Fifteen percent of layer A consists of smaller fragments of juvenile microvesicular to dense, light- to medium-gray andesite. Pre-1992 volcanic fragments including distinctive rounded clasts of glassy black vesicular rock make up 5 percent of layer A. About 1 percent consists of light-colored plutonic or metamorphic clasts. Dry-bulk densities of two layer A tephra samples are 1.09 and 1.16 g/cm³. The tan and gray juvenile clasts from layer A have average densities of 1.53 and 2.01 g/cm³, respectively (table 2).

The September 16–17 layer B is ungraded and finer grained at a given distance from the vent than layer A. Sixty-four percent of the clasts in layer B are light- to medium-gray, microvesicular to dense andesite. Twenty-six percent of layer B clasts are light tan, slightly pumiceous and phenocryst-rich. Dry-bulk densities of two samples are 1.24 and 1.32 g/cm³, 15 percent greater than layer A. The lower bulk density of September 16–17 layer B compared to August 18 layer B is reflected in the higher percentage of tan clasts in the September 16–17 deposit (table 1). Average clast densities for tan and gray components of layer B are 1.58 and 2.07 g/cm³, respectively. Pre-1992 volcanic fragments make up nearly 10 percent of layer B. At more distal locations (20 to 30 km from the vent) scattered silicic pumice fragments litter the surface of layer B.

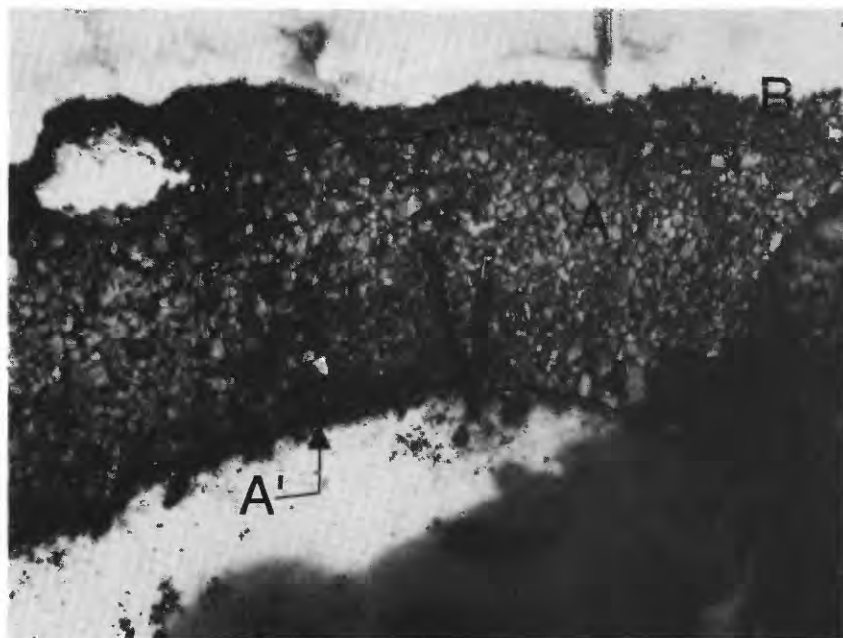


Figure 4. A 17-cm-thick tephra-fall deposit from the September 16–17, 1992, eruption, located approximately 7.2 km northeast of Crater Peak, Mount Spurr volcano, Alaska. Layers A', A, and B are labeled. Note reverse grading of layer A. Section rests on snow. Pencil in center of photograph is approximately 10 cm long.

With increasing distance along the axis of the deposit, the color change in the September 16–17 tephra fall becomes more distinct. About 250 km away, at the northern edge of the deposit, the ash is very light in color, presumably reflecting dominance of the tan juvenile component (layer A?). At the southern margin, the ash is noticeably dark and probably represents the gray juvenile component (layer B?). Given measurable density differences in both dominant clast and bulk densities between layers A and B of 15 to 30 percent (table 2), such asymmetry could reflect within-plume sorting due to slight variations in wind direction with altitude or over the course of the eruption, similar to that documented at Mount St. Helens in 1980 (Sarna-Wojcicki and others, 1981; Waitt and others, 1981).

Ballistic emplacement of coarse breadcrusted bombs consisting principally of tan and gray juvenile components and a smaller lithic component occurred during the September 16–17 eruption. Within approximately 10 km of the vent, along the southern margin of the tephra-fall field, lapilli and bombs were deposited along a narrow corridor. In the snowpack, the ballistic horizon consisted of isolated bombs and lapilli in a matrix of snow and ice that extended down as much as a meter (fig. 5). This disseminated deposit was produced when hot ballistic ejecta impacted snow, differentially melting downward depending on the size

of each clast. Melted entry trails were subsequently refrozen to form ice stringers. In part because of the variable settling, there was no time sequence preserved in this ballistic deposit, hence we could not evaluate any stratigraphic information. The collapsed thickness of the deposit after drying ranged from 3 to 4 cm between 6 and 12 km from the vent.

RELATIONS OF TEPHRA-FALL DEPOSITS TO OTHER OBSERVATIONS

To help interpret the proximal tephra-fall record in relation to eruptive processes, the following section summarizes aspects of the timing of tephra fall to pyroclastic flow formation, ballistic ejection of large blocks, and other observations.

JUNE 27

Cloud cover precluded direct observation of the development of the June 27 eruption column or tephra plume. Moreover, strong southerly winds carried the plume northward, 180° away from the path of hot-pyroclast and snow mixtures that swept down the south and southeast flanks of Crater Peak (Meyer and Trabant, this volume). Hence, there is no stratigraphic infor-

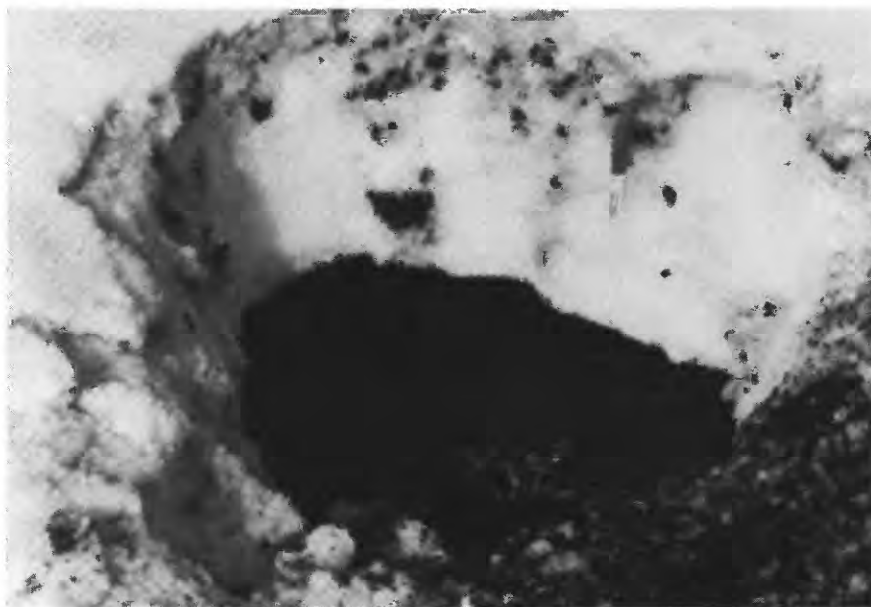


Figure 5. Ballistic tephra-fall deposit from September 16–17, 1992, eruption, located 7 km northeast of Crater Peak, Mount Spurr volcano, Alaska. Projectiles shown here penetrated snow upon impact and left a disseminated horizon of lapilli, bombs, snow, and ice. Dark horizon below is the August 18 tephra section. At the top is a fine ash layer that probably reflects aeolian reworking of material subsequent to the September 16–17 event.

mation to help determine the timing of pyroclastic flow emplacement. Observations inside Crater Peak following the eruption revealed meter-sized bombs littering the surface of a greater than 10-meter-thick deposit of finer grained pyroclastic material interpreted by Miller and others (this volume) as surge deposits. Similarly, numerous coarse juvenile and lithic bombs and associated impact craters dotted the landscape north and east of Crater Peak. Thus, ejection of large bombs, perhaps caused by intermittent phreatomagmatic explosions in the upper part of the conduit, occurred at the end of the June 27 eruption.

Radar images of the June 27 eruption plume are of insufficient quality and pilot reports are too few to establish a good record of the development of the column. However, the maximum radar-discernible column height was recorded 3.3 hours into the eruption (Rose and others, this volume). This is remarkably consistent with RSAM (real-time seismic amplitude measurements; a crude measure of eruption intensity; Murray and Endo, 1989) values, which increased gradually, peaked nearly 3.5 hours into the eruption, and then decreased abruptly during the final hour (fig. 6; McNutt and others, this volume).

AUGUST 18

The opening phase of the August 18 eruption coincided with 12 minutes of tremor recorded on one Spurr seismic station between 3:41 p.m. and 3:52 p.m. ADT (station CRP; McNutt and others, this volume). This event produced a minor tephra plume that, according to pilot reports, had reached approximately 2.5 km above the vent (4.5 km ASL) by 4:20 p.m. ADT. A thin (less than 1 mm), medium sand deposit from this event (layer A') was identified at three localities 4 to 16 km east of the vent. The deposit was

not sampled, and the amount of juvenile material vented during this opening phase of eruption is not known.

The main phase of the August 18 eruption occurred from 4:42 p.m. to 8:10 p.m. ADT. Parts of this event were photographed by Paul Palmer on a commercial flight from Anchorage to Iliamna and also by

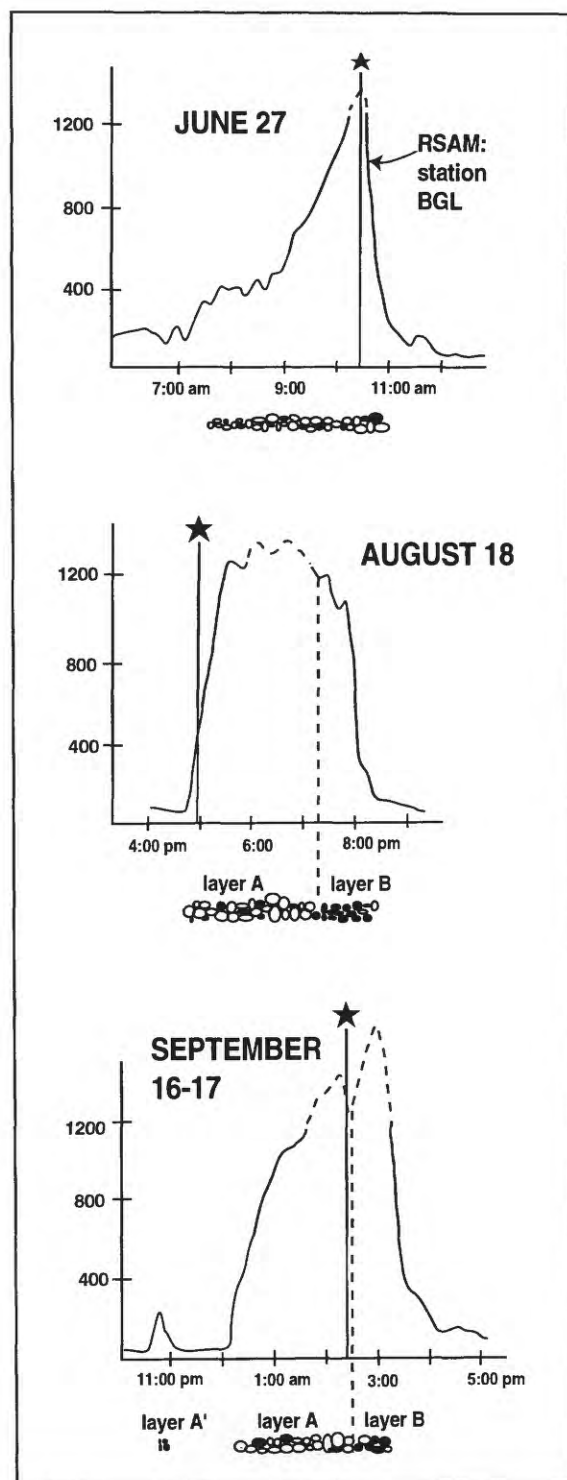


Figure 6. Schematic relation in time between RSAM, maximum eruption column height, tephra production and tephra-fall stratigraphy for the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska. Curves are RSAM changes at station BGL, located 7.5 km west of Crater Peak; dashed where values represent instrument saturation and RSAM is unreliable. Vertical axis is digital counts. Stars denote time of maximum column height determined from C-Band radar (Rose and others, this volume). Below each curve is a schematic section of tephra-fall deposit indicating approximate proportion of tan and gray components, changes in grain size and, for August 18 and September 16–17 deposits, position of the layer A–layer B break (long dashes). For September 16–17, sandy layer A' from the opening phase of the eruptions is also shown.

an airborne AVO crew who reached the mountain at about 5:50 p.m. ADT (McGimsey and Dorava, 1994). According to the Palmer photographs, pyroclastic flows, indicated by a light-tan elutriation cloud adjacent to the main eruption column, had already descended the southeast flank by approximately 4:55 p.m. ADT (fig. 7; P. Palmer, oral commun., 1993; Miller and others, this volume). Scaling of photographs indicates an altitude of between 10.5 and 11.5 km ASL for the top of the main eruption plume, consistent with pilot estimates of at least 10 km ASL.

Seismicity indicates that vigorous emission of tephra continued until approximately 8:11 p.m. ADT (McNutt and others, this volume). Intermittent production of pyroclastic flows may have continued throughout the eruption, however the final pyroclastic flows did not extend as far down the flank of Crater Peak. In photographs and from a brief field inspection, these late flows were also darker in color and contained a higher percentage of gray juvenile bombs. Ballistic emplacement of dense bombs, indicated by impact craters on the surface of primary pyroclastic-flow and tephra deposits east of Crater Peak, occurred during the final phase of the eruption (Miller and others, this volume; Waitt and others, this volume).

No complete exposures of the tephra sequence inside Crater Peak were accessible following the August 18 eruption. However, we were able to briefly examine the upper part of the section on the southwest rim of a basin bounded by a tephra rampart (see fig. 1 of Miller and others, this volume). The uppermost unit was a fine, stratified, brown mud approximately 2 to 6 cm thick. This mud overlay an approximately 15-cm-thick sequence of stratified, poorly sorted,

brown to black, sandy debris that included a large percentage of oxidized fragments. The lowermost unit was a gravel-boulder horizon, more than several meters thick. It showed strong upward coarsening and many of its angular clasts were highly oxidized. Among the coarsest blocks at the top of this horizon were dense, prismatic jointed, dark-brown juvenile andesite, as much as 50 cm across, and abundant dense, light-colored, aphyric to porphyritic, volcanic bombs as much as 2 m across. This coarse ejecta within the crater is probably correlative with the ballistic field on the southeast flank of Crater Peak (Waitt and others, this volume) and represents a series of closely spaced phreatomagmatic explosions that increased in intensity with time at the very end of the eruption. The sand to silt-sized material capping the entire sequence suggests that low-level phreatic (?) emission of fine material continued after emplacement of ballistics. This may explain the intermittent, low-frequency seismicity in the 24 hours following the eruption, as well as a pilot report of a low (a few hundred meters above Crater Peak) ash and steam column visible above Crater Peak on the morning of August 19.

The radar record of the August 18 eruption is only partially complete, but the maximum recorded height of a radar-discernible plume occurred 13 minutes into the main phase of the eruption (Rose and others, this volume). RSAM amplitude at station BGL reached maximum levels rapidly, and it stayed elevated throughout most of the eruption (fig. 6). Although the record is hampered by saturation problems, RSAM values apparently began to decrease by about 7:50 p.m. ADT, approximately 3 hours into the 3.5-hour eruption (McNutt and others, this volume).



Figure 7. Photograph of the August 18 eruption column and spreading tephra cloud through the window of a commercial flight en route to Iliamna from Anchorage. View direction is approximately north-northwest. The summit of Mount Spurr is just visible to the left of the base of the eruption column. Crater Peak is hidden below a cloud layer, the top of which is about 8,500 feet (2,590 m) ASL. Small, lighter-colored cloud (arrow) is elutriated ash from pyroclastic flows. Time of photograph is approximately 4:55 p.m. ADT, 13 minutes after the onset of the main phase of the eruption. The top of the eruption cloud has reached an altitude of about 11 km ASL. Photograph by Paul Palmer. Used with permission.

SEPTEMBER 16-17

The September 16-17 eruption occurred at night under clear skies. The main phase of the eruption was preceded by an 11-minute burst of seismicity and explosive tephra emission that started at 8:36 p.m. ADT. This eruption was vigorous enough to produce bright flashes on the AVO slow-scan TV camera located at Kasilof, 120 km southeast of Mount Spurr volcano (fig. 1A). In contrast to August 18, this opening phase was detected on C-band radar, hence the tephra column must have extended at least 2 km above Crater Peak (Lee Kelley, NWS, oral commun., 1993). Tephra fall from the first phase was difficult to identify, but at one locality the deposit consisted of a few millimeters of coarse ash and lapilli of both juvenile and accidental material.

The main phase of eruption began just after midnight. Intermittent bursts of incandescence near the base of the eruption column were visible from Anchorage and recorded by the slow-scan camera. Of note is that many of the elongated light flashes captured on film deviate from vertical by several tens of degrees. Miller and others (this volume) relate this directed ejection of incandescent material to the formation of coarse pyroclastic flows on the east flank of Crater Peak. The ballistic field on the southern edge of the September 16-17 tephra deposit may have also formed during episodes of this directed eruption column. Hunters camped 18 km southeast of Crater Peak described incandescent bursts related to projectile impacts into debris on the glaciers east of the cone during the first hours of the eruption; their observations suggest that ballistic ejection occurred over an extended period. Bomb craters on the surfaces of primary pyroclastic-flow and lahar deposits on the glacier east of Crater Peak indicate that ballistic blocks and bombs of appreciable size also fell *after* generation of pyroclastic flows on September 16-17. There were no bomb craters noted in the lahar deposits that descended the south flank of Crater Peak during the September 16-17 eruption. This observation is consistent with projectiles directed predominantly eastward from the vent, as indicated on the slow-scan images.

Intracrater deposits from the September 16-17 eruption are exposed in a series of scalloped headwalls cut into terraces that ring the inner vent. On the south margin of the crater, approximately the upper meter consists of oxidized, bedded sandy surge deposits capped by fine silt, which was water saturated upon examination a few weeks after the eruption. These sandy surges thicken toward the east sector of the crater where they overlie a massive, poorly sorted very coarse sand to boulder-gravel deposit, the base of which was not exposed.

The C-band radar record of the September 16-17 event is the best of the series and documents a fluctuating tephra column, which finally reached a sustained maximum height of 13.9 km about 2.3 hours into the eruption (Rose and others, this volume). RSAM values at station BGL increased gradually, remained at high levels between about 1:30 a.m. and 3:00 a.m. ADT and decreased abruptly by 3:10 a.m. ADT, about 3 hours into the eruption (fig. 6.; McNutt and others, this volume).

DISCUSSION

The presence of tan and gray andesite components in all three eruptions is a primary characteristic of the 1992 Crater Peak tephra. Although petrographically quite similar, the two components differ in density: the tan component is consistently 22 to 24 percent less dense than the gray component (table 2). And, whereas the entire tephra sequence is mixed to some degree for all three eruptions, for August 18 and September 16-17, the tan (less dense) component dominates the lower 60 to 80 percent of the deposits (layer A) and the gray component (more dense) dominates the capping unit (layer B). One explanation of this density stratification is pre-eruption zonation in volatile content. Tan-colored clasts of layer A represent the first material erupted from the upper, volatile enriched part of the magma body; gray-colored clasts represent fragments of melt stored slightly deeper and depleted, relatively, in volatiles. The imperfect separation of components (mixing of tan and gray within layers A and B) may reflect nonuniform withdrawal of magma from the conduit (Blake and Ivey, 1986).

If the layer A-B divisions within the August 18 and September 16-17 deposits reflect a change in magmatic volatile content, a coincident and measurable change in the behavior of the eruption column might be expected. To test this, we superimposed a schematic representation of the fallout record on plots of RSAM, the only independent instrumental measure of eruption intensity available for the entire eruption series (fig. 6). To construct figure 6, seismicity, pilot reports, radar data, the few eyewitness observations, and slow-scan video were used to roughly constrain the period of tephra production. Note that we have assumed by the superposition of the RSAM record over the schematic tephra-fall section that there is an instantaneous relation between seismicity and tephra accumulation; in reality, there must be some delay.

For the June 27 eruption, there is no break in tephra-fall stratigraphy where the proportions of tan to gray components change. Hence, based solely on field observations and the samples in hand, we can-

not simply relate the RSAM curve to the proportion of clast type. For the August 18 eruption, radar data (Rose and others, this volume) and pilot reports allow fairly accurate limits to be placed on the timing of tephra production. In figure 6, the layer A-B transition was placed approximately 65 to 70 percent through the eruption on the basis of field measurements of relative thicknesses. Projecting this break onto the RSAM curve suggests that the later stage of tephra production (dominated by the gray component of layer B) was associated in time with a period of declining RSAM. This association could be explained by tapping of a deeper, more degassed part of the magma body, which resulted in lower eruption column vigor. Unfortunately, radar data are lacking for the later stages of the eruption and we cannot independently evaluate column height. For the September 16–17 eruption, the layer A-B transition occurs during a period of instrument saturation, hence it is difficult to assess the relation of stratigraphy to RSAM data for station BGL. To summarize, only data for the August 18 eruption suggest a possible correlation between eruption intensity and clast type erupted.

Another prominent characteristic of the August 18 and September 16–17 layer A deposits, and to a lesser extent the June 27 deposits, is reverse grading. Reverse grading in tephra-fall deposits can reflect a decrease in the angle of pyroclast ejection, changes in wind velocity, or increasing magma discharge over the course of an eruption. For both the August 18 and September 16–17 eruptions, photography and slow-scan TV images do not support any progressive changes in the inclination of the eruption column. Furthermore, the narrow fallout patterns indicate that local winds were steady in speed and direction over the course of all three events. Thus, we propose that the reverse grading in the June 27 and lower August 18 and September 16–17 deposits is a result of an increase in magma discharge at the vent.

To test this, we compared the stratigraphic position of maximum clast size within the resulting deposits and the times of maximum plume heights and RSAM peaks (fig. 6). For June 27, there is a good correlation among all three parameters. For August 18 layer A, maximum grain size is approximately coincident with the RSAM peak (although saturation makes this uncertain). More importantly, the amplitude of volcanic tremor recorded at more distant stations reached a maximum about halfway through the August 18 eruption (S. McNutt, written commun., 1993). Similarly, the maximum clast size is present very near the middle of the deposit. Maximum plume height on August 18, however, occurred only minutes after the onset of the eruption. For September 16–17 layer A, maximum clast size and maximum plume

height are approximately congruent. While the RSAM data for September 16–17 are difficult to interpret, the strongest tremor occurred during the middle of the eruption (S. McNutt, written commun., 1993). Thus, grain size maximums for all three eruptions occur at stratigraphic positions that approximately mirror the times of maximum seismicity (as measured by a combination of RSAM and tremor amplitude). The large offset in time of maximum plume height for the August 18 eruption and the maximum clast size in the resulting tephra-fall deposit remains an enigma and may be an artifact of poor radar sampling (Rose and others, this volume). A more careful analysis of changes in grain size and clast characteristics with stratigraphic height for all three deposits is needed to determine the exact mechanism behind the development of reverse grading.

In other stratigraphic studies of tephra-fall deposits, color and clast-size changes are associated with small but measurable chemical or mineralogic variations, interpreted to reflect pre-eruption chemical zonation in the magma chamber (Bursik and others, 1992; Criswell, 1987; Carey and Sigurdsson, 1987). Whole-rock major- and trace-element analyses of individual tan and gray clasts for Crater Peak, however, show no *significant* differences between the two clast types either within or through the eruption series.

Major- and trace-element chemical data for tephra normalized to average 1992 andesite blocks retrieved from pyroclastic-flow deposits are shown in figure 8. Tan tephra are nearly identical to the andesite with the possible exception of low Cs and Rb concentrations in two samples. But even these are less than 25 percent different from the andesite values (0.1 ppm Cs and 3 ppm Rb), and they are nearly within two standard deviations of the andesite mean. Over half of the gray tephra samples analyzed to date have concentrations of all elements similar to the andesite blocks. However, some gray tephra samples are more heterogeneous. These samples are highly enriched in Cs (as much as 3.7 times the andesite mean) and also enriched in Sc, Cr, and Ni. The paired enrichment in alkalis and compatible transition metals suggests contamination by metamorphic country-rock represented by partially melted xenoliths found in the 1992 pyroclastic-flow deposits (see Nye and others, this volume). However, simple xenolith contamination is not controlling the high Cs in these samples (fig. 9). Those samples with very high Cs do not have the low CaO or high Th that xenolith-contaminated samples would have.

Processes controlling the chemistry of the gray tephra samples are not clearly known but may include accumulation of mafic phenocrysts, vapor-phase transport of some metals, or incorporation of highly fractionated material from the margin of a sub-Crater Peak

magma chamber. This final possibility is supported to some extent by the recovery of rare, light gray, moderately siliceous pumice lapilli from the June 27 eruption, which are similar to the divergent tephra samples in Cs, Ca, Th, U and, to a lesser extent, Cr and Ni (Nye and others, this volume).

Finally, the gray tephra samples that are chemically unlike the andesite have compositions governed by chemical heterogeneities on a very small scale (about 10 g of material). Only when small sample amounts

were analyzed were heterogeneities found, and the magnitude of those heterogeneities varied between splits. When larger samples (greater than 50 g) were processed, the tephra analyses were generally closer to that of the andesite. In summary, both tan and gray tephra are chemically similar or identical to the 1992 andesite and to each other. Differences that can be found are slight and related to contamination by poorly understood components on a very localized scale.

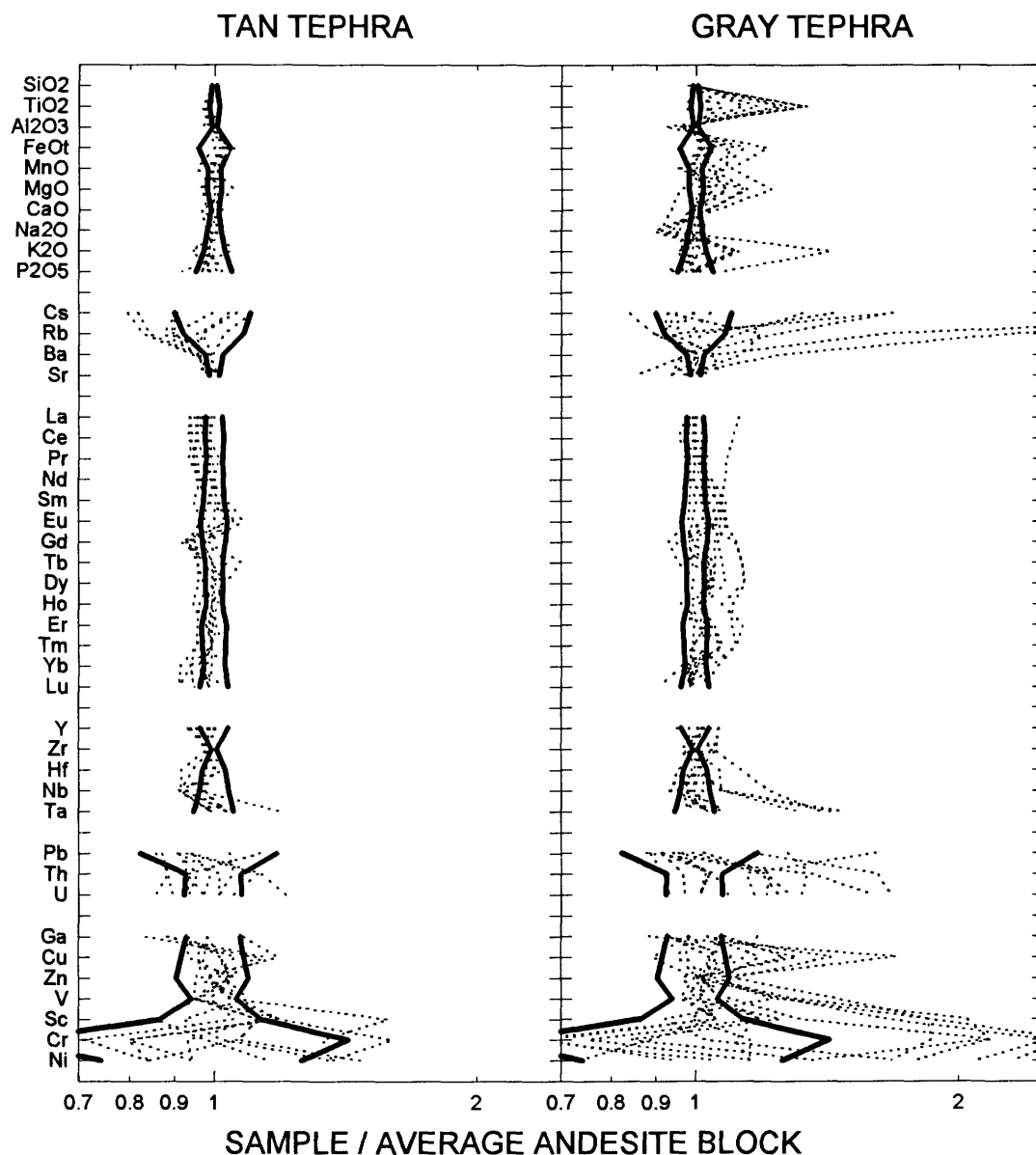


Figure 8. Whole-rock major- and trace-element compositions of selected tan and gray tephra lapilli from the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska. FeO_t = FeO total. Dark lines represent one standard variation about the mean of 18 juvenile andesite blocks from all phases of each eruption (see Nye and others, this volume).

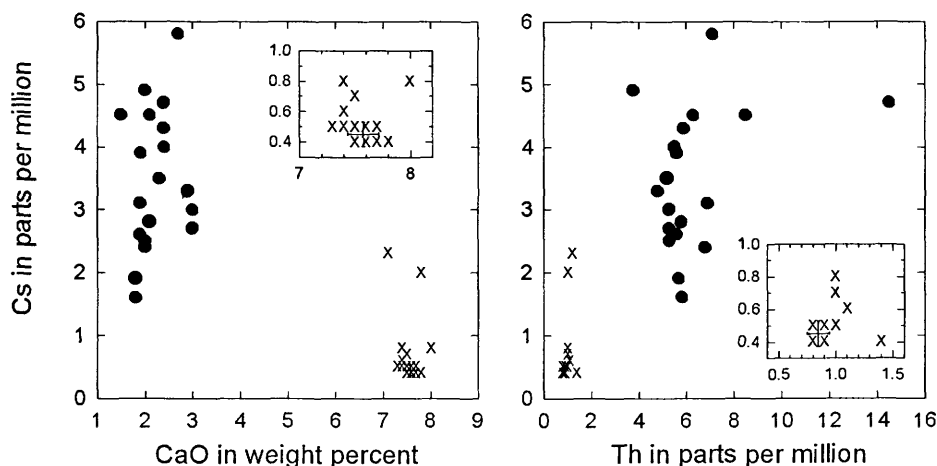


Figure 9. Variations of CaO and Th with Cs. Filled circles, analyses of partially melted aluminous metamorphic xenoliths ejected during the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska (see Nye and others, this volume); X, tan and gray tephra analyses of fig. 8; and +, average analysis of the 1992 juvenile andesite blocks (+ spans one standard deviation about the mean of 18 samples). Note that the majority of tephra are similar to the andesite and that those dissimilar from the andesite do not form mixing arrays with the xenoliths.

SUMMARY

Three vulcanian to subplinian eruptions of Crater Peak at Mount Spurr volcano during 1992 produced tephra-fall deposits that extend north, northeast, and east of the volcano. The eruptions were similar in magnitude, style, and duration and produced, respectively, 44, 52, and 56 $\times 10^6$ m³ of bulk tephra (12, 14, and 15 $\times 10^6$ m³ DRE) in 3.5 to 4.0 hours of vigorous eruption. Measured thicknesses ranged from about 1 m at the base of the vent to less than 1 mm at a distance of 200 km from Crater Peak.

Fallout tephra from the 1992 eruptions of Crater Peak consists predominantly of two texturally distinct but chemically similar juvenile andesitic components. Density data suggest that prior to each eruption, a physically zoned magma body existed beneath Crater Peak. The zonation is best explained by upward migration of volatiles and formation of a gas-charged upper part of the chamber. The presence of both components in all three eruptions suggests that such zonation can develop rapidly (that is, within 4 to 6 weeks, the time intervals between eruptions). Alternatively, each eruption may have tapped only a part of an initially zoned chamber emplaced prior to the first eruption on June 27. Minor settling of mafic phenocrysts, contamination of juvenile andesite by remelted country rock, incorporation of a highly fractionated derivative of the juvenile andesite, and vapor-phase transport of some elements may account for observed small geochemical variations of individual tephra clasts.

Future analytical work on the Crater Peak tephra-fall deposits will focus on clarifying the extent and nature of small-scale geochemical heterogeneities; quantifying vesicularity, phenocryst, and groundmass textures to address the root cause of the systematic density differences (Gardner and others, 1993); and examining clast grading and density characteristics in more detail to reconstruct changes in eruption-column energetics through each event.

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Pyroclastic Flows of the 1992 Crater Peak Eruptions: Distribution and Origin

By Thomas P. Miller, Christina A. Neal, and Richard B. Waitt

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ABSTRACT

Two of the three 1992 eruption events at Crater Peak generated pyroclastic flows on the flanks of the cone. The pyroclastic-flow deposits are very coarse, generally fines poor, poorly sorted, and clast supported; 60 to 80 percent of the clasts are dense to slightly vesiculated cobble to boulder-sized blocks of juvenile andesite. Accessory and accidental clasts are chiefly fragments of lava but include some spectacularly vesiculated metasedimentary rocks (buchites). The pyroclastic flows traveled a maximum distance of about 3.2 km and had a combined bulk volume of about 400,000 m³, or less than 0.5 percent of the bulk volume of the entire eruption. They were avalanches of hot debris that produced sluggish, poorly inflated dense-clast flows. The flows may have been generated at least in part by directed ballistic ejection of debris rather than by "boilover" through a low point on the crater rim. The lack of correlation between the distribution of pyroclastic flows and the low part of the crater rim has important implications for hazard analysis of vulcanian and subplinian eruptions elsewhere because it reminds us that all flanks of a volcano can be at risk from pyroclastic flows.

INTRODUCTION

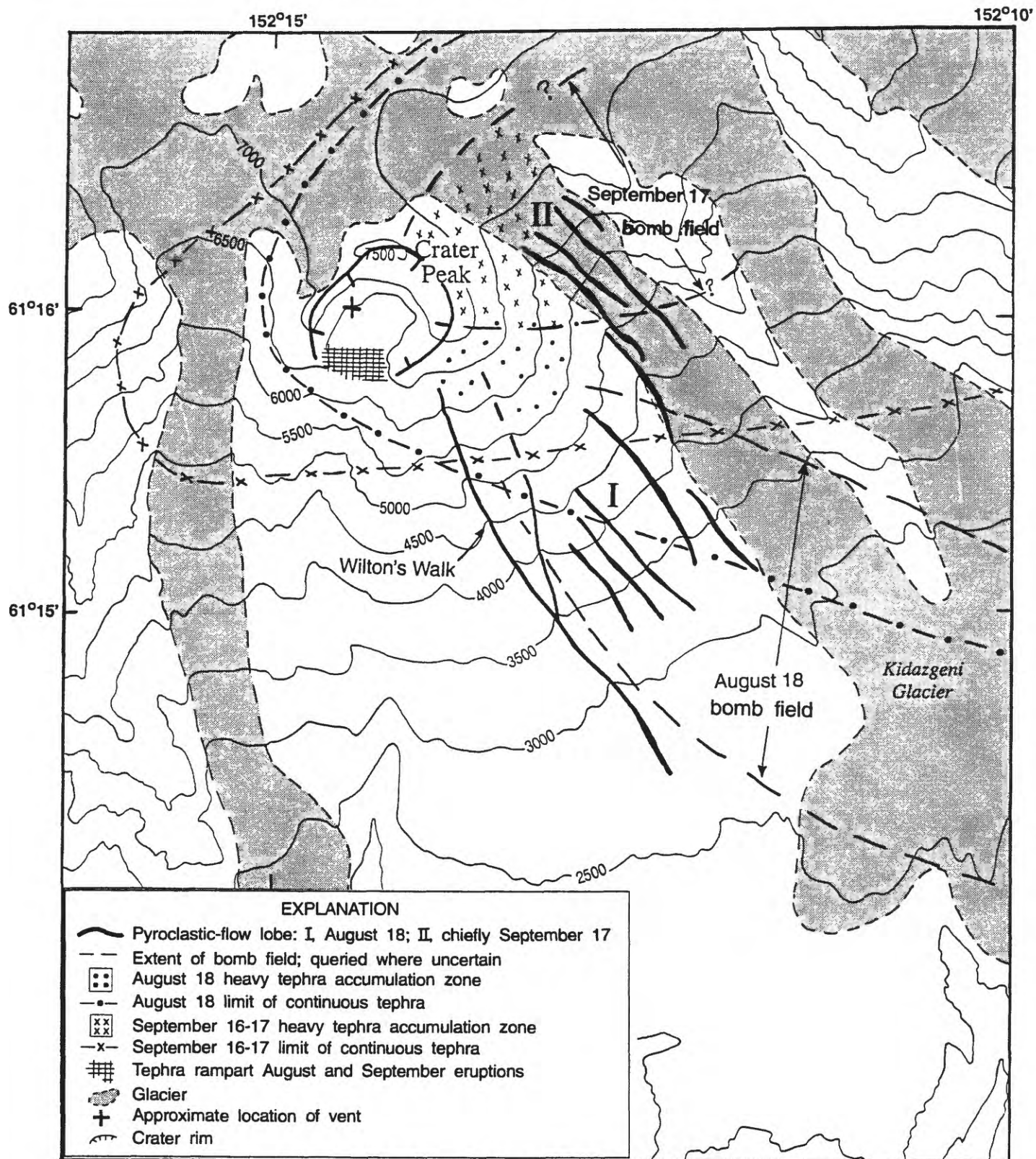
Crater Peak, a 2,309-m-high satellite cone on the south flank of the Mount Spurr volcanic center 125 km west of Anchorage erupted on June 27, August 18, and the night of September 16–17, 1992 (see Eichelberger and others, this volume) after 39 years of quiescence. The three eruptions were similar in that

they each produced vulcanian to subplinian eruption columns as high as 18 km above sea level, formed extensive bomb fields, lasted 3 to 4 hours, erupted pyroclastic material only, had ejecta bulk volumes of 44 to 56 x 10⁶ m³ each (see Eichelberger and others, this volume), and were of andesitic composition. The July 9, 1953, eruption, the only other historic Crater Peak eruption, had similar characteristics, according to the fragmentary accounts available (Juhle and Coulter, 1955).

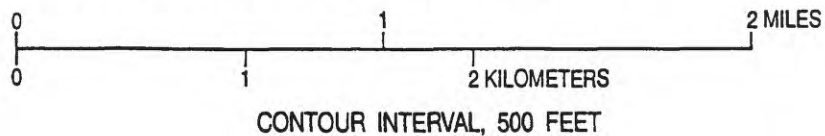
Crater Peak is capped by a circular summit crater 300 m deep and 800 m in diameter (fig. 1); rim elevation ranges from 1,980 m (6,500 ft) at the low point on the south-southwest to 2,309 m (7,575 ft) on the northeast. Before the initial June 27 event, the crater contained a considerable thickness of ice, snow, and an 800-square-meter lake. Most of the ice and snow was melted or removed during the 1992 eruptions, particularly the one on June 27: the lake disappeared a few days before the June 27 eruption. A terrace of 15 m or more of massive, locally well bedded surge deposits capped by 1 to 2 m of fine ash formed in the southern third of the crater during the June 27 eruption; 1- to 2-m lithic blocks were scattered across the surface of this terrace. Later eruptions modified this feature into two tephra-covered terraces that now occupy the southern two-thirds of the crater.

The crater walls changed little during the 1992 eruptions other than steepening of the north wall. The eruptive vent, estimated by aerial observations to be less than 30 m across at the surface, is in the northwest and deepest part of the crater about 70 m below the surface of the lower terrace.

Small-volume, dense-clast pyroclastic flows moved down the outer south and east flanks of Crater Peak cone during the August 18 and September 16–17 eruptions (fig. 1). During the June 27 event, either pyroclastic flows were not generated or were turned quickly into "snowflows" or "mixed" flows (see Waitt, this volume). The total volume of the pyroclastic-flow deposits is estimated at 0.4 x 10⁶ m³, or less than 0.05 percent of the approximately 150 x 10⁶ m³ bulk volume of the 1992 eruption products (see Eichelberger and others, this volume).



Base from U.S. Geological Survey, Tyonek, A-6, A-7, B-6 and B-7, Alaska, 1:63,360, 1958



Similar pyroclastic flows, often called scoria flows, have been reported from historic eruptions at Ngauruhoe volcano in New Zealand (Nairn and Self, 1978) and at Manam volcano in the western Bismarck arc near Papua, New Guinea (McKee, 1981). Such flows are commonly thought to have formed by column collapse or by boilover of a low eruption column (Cas and Wright, 1987) and flowage through low points on crater rims. The limited distribution of the Crater Peak pyroclastic flows and the lack of any association with low points on the crater rim suggest that other mechanisms were responsible for their formation. This possibility is important for assessing hazards from similar eruptions elsewhere. Because pyroclastic flows were not observed during daylight hours, the processes that formed them must be inferred.

ACKNOWLEDGMENTS

We are indebted to Game McGimsey and Cynthia Gardner for assistance in the field during and after the eruptions. Reviews by Ric Hoblitt and Wes Hildreth greatly improved the manuscript.

DISTRIBUTION

Crater Peak is bounded on the north, west, and east by glaciers (fig. 1). The August 18 pyroclastic flows moved down the steep (31° to 7°) south and southeast flanks of Crater Peak to elevations of 820 to 1,000 m (2,700 to 3,300 ft) at distances of as much as 3.2 km from the crater rim. The flows were channeled into gullies, where they formed elongate lobes of debris (fig. 1) with steep-sided margins and flow fronts 0.5 to 1.5 m high. The westernmost flows of the August 18 eruption were channeled into a well-developed, intermittent stream drainage (informally known as Wilton's Walk; figs. 1 and 2) on top of thin snowflows of the June 27 eruption. The easternmost pyroclastic flows followed swales and low gullies lined with brush and other low vegetation. None of the pyroclastic flows descended to tree line.

Most of this report is based on study of the August 18 pyroclastic-flow deposits, which are the most accessible—although lahars from the September 16–17 eruption overlapped many of the August pyroclastic-flow deposits in Wilton's Walk gully. The September 16–17 pyroclastic flows moved down the east-north-

east flank onto a high ice field (fig. 1), causing much melting of the ice and were soon transformed into lahars; access to these pyroclastic-flow deposits is difficult and hazardous.

The pyroclastic-flow deposits of the August 18 and September 16–17 eruptions have a close spatial association with large bomb fields (figs. 1 and 3) that covered areas of 7 and 2 km², respectively. Field relations show that for the August eruption, most of the bombs were ejected after the pyroclastic flows were emplaced (see Waitt and others, this volume).



Figure 2. View of south flank of Crater Peak cone (summit obscured by clouds) with pyroclastic-flow deposit in foreground. Prominent drainage, locally known as Wilton's Walk, was the site of numerous pyroclastic flows during August 18 eruption and lahars during the September 16–17, 1992, eruption.

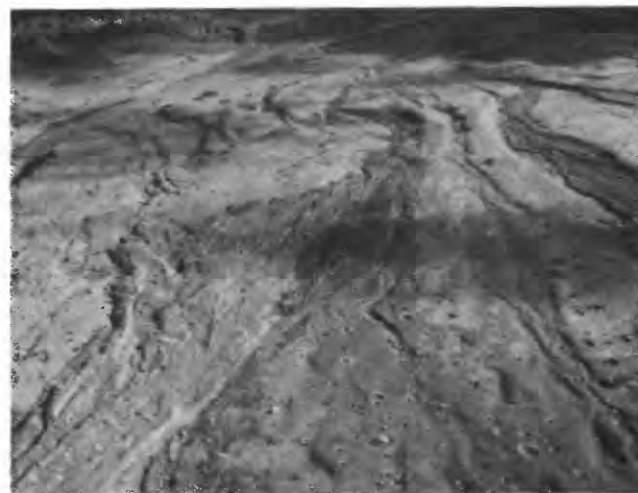


Figure 3. View toward southeast of bomb field developed on top of the August 18 pyroclastic-flow deposits, southeast flank of Crater Peak, Mount Spurr volcano, Alaska. Craters in foreground are 3 to 5 m in diameter.

◀ **Figure 1.** Crater Peak area of Mount Spurr volcano, Alaska, showing distribution of pyroclastic-flow deposits and tephra accumulation zones from the August 18 and September 16–17, 1992, eruptions and the August 18 bomb field. Bomb field from September 16–17 eruption not completely mapped because of inaccessibility.

COMPOSITION

The pyroclastic-flow deposits are very coarse, clast supported, and generally fines poor (fig. 4). Most clasts (60 to 80 percent) consist of dense to slightly vesiculated (15 to 20 percent vesicles), dark gray to brownish-black compositionally homogeneous, porphyritic juvenile andesite (56 to 57 percent SiO_2 , table 1; see Nye and others, this volume); this material forms blocks as large as 1 m in diameter. The andesite contains about 20 percent phenocrysts consisting of plagioclase, clinopyroxene, orthopyroxene, iron oxides, and sparse hornblende in a microlite-rich groundmass of brown andesitic glass. A second and comparatively minor type of juvenile clast is a light gray-green andesite scoria that is compositionally similar to the dark andesite in spite of containing a clear rhyolite glass in the groundmass (see Nye and others, this volume).

Lithic clasts, mostly accessory but including a few accidental ones, are the second most abundant clast type (20 to 40 percent) and are typically quite large (as much as 2 m in diameter). The accessory clasts are chiefly unaltered andesite and basalt that were probably picked up from the conduit and the crater walls; accidental clasts were obtained from lava flows exposed on the flanks of Crater Peak. Some clasts of indurated pre-1992 pyroclastic material are also present. Accessory and accidental lithic contents of several pyroclastic-flow deposits from August 18 increase from west to east across the southeast flank of the cone.



Figure 4. Dense clast, fines-poor August 18 pyroclastic-flow deposits in Wilton's Walk. Photo taken August 20, 1992, 2 days after eruption of Crater Peak vent, Mount Spurr volcano, Alaska.

Table 1. Chemical analyses of juvenile andesite clasts in August 18 and September 16–17 pyroclastic flows (see Nye and others, this volume).

Table 1. Chemical analyses of juvenile andesite clasts in August 18 and September 16–17 pyroclastic flows (see Nye and others, this volume).

	August 18	September 17
SiO_2	56.5	55.75
Al_2O_3	19.2	18.7
FeO	6.9	7.0
MgO	3.6	3.7
CaO	7.6	7.5
Na_2O	4.1	4.0
K_2O	.9	.9
TiO_2	.67	.68
P_2O_5	.29	.28
MnO	.15	.15
Total	99.9	98.6
Number of analyses	3	4

Particularly exotic metamorphic xenoliths make up less than 0.05 percent of the clasts in the flows, but they are important for determining the type (and depth) of wallrock enclosing the crustal magma chamber in which the andesitic magma may have resided immediately before eruption. These xenoliths consist of highly expanded pumice-like material that forms smooth rounded clasts as much as 60 cm in diameter. This partly glassy material is composed of plagioclase+orthopyroxene±sillimanite±biotite±cordierite±spinel±garnet±quartz and has SiO_2 contents of about 65 percent (see Harbin and others, this volume). The clasts have a well-developed compositional banding and range from slightly to highly vesiculated. Similar material at Ngauruhoe, New Zealand, was described as “quartzofeldspathic gneisses” or “buchites” that have been “frothed” (Steiner, 1958; Nairn and Self, 1978; Graham and others, 1988).

A less abundant but equally exotic component is a white, fine-grained, sugary-textured rock composed of plagioclase+quartz+glass. This material is friable but occurs in blocks as large as 60 cm and it makes up less than 5 percent of the exotic lithics. Rounded pieces of this material as large as 15 cm across have been found in juvenile andesite blocks. Contacts between the two materials are sharp suggesting only minor interaction between the “white-rock” and the andesitic magma. This relation was observed in tephra fragments as well (see Neal and others, this volume; Harbin and others, this volume).

A few clasts of dense white garnetiferous calc-silicate skarn as much as 20 cm in diameter were found in the eastern pyroclastic-flow deposits of August 18. The garnets form large (1.5 cm) euhedral amber-colored porphyroblasts in a dense mosaic of wollastonite and plagioclase.

Metasedimentary rocks of these compositions have not been recognized within many tens of kilometers of the volcano. Although the basement surrounding the Mount Spurr volcanic center consists chiefly of Mesozoic plutonic rocks (Magoon and others, 1976) of the Alaska-Aleutian Range batholith, few plutonic rocks have been found in the 1992 ejecta. The almost total domination of vesiculated metamorphic rocks compared to other basement rocks suggests that they form the wallrock for a crustal magma reservoir that fed the 1992 Crater Peak eruptions. The lack of plutonic xenoliths indicates that the magma reservoir could be deeper than any upper-crustal granitic pluton.

DEPOSIT CHARACTERISTICS

At least some of the August 18 pyroclastic flows formed early in the eruption on the basis of observations made during an hour-long aerial overflight that began about 20 minutes into the eruption on August 18 (McGimsey and Dorava, 1992). A thin but persistent cloud layer surrounded the volcano at about the 2,400 m (7,800 foot) level. Brief glimpses through breaks in the cloud indicated that pyroclastic flows had come down the south and east flanks of the cone to an estimated elevation of 1,000 m. In addition to a vigorously convecting gray-black eruption column that rose to an altitude of 18 km, a lower (3.5 to 4.5 km) lighter colored cloud was visible along the south and southeast edge of the column and probably represented fine material elutriated from the pyroclastic flows. Blocks larger than 2 m were observed falling out at the perimeter of the column 300 m above the vent. These blocks probably contributed to the large post-pyroclastic-flow bomb field southeast of the cone.

Emplacement of the September 16–17 pyroclastic flows may have taken longer. Inspection of videotape from slow-scan TV records of the night-time eruption of September 16–17, taken from a camera located 120 km south of Crater Peak, shows that inclined incandescent fountaining (and probably pyroclastic flow generation) took place sporadically throughout the 4-hour eruption.

Field characteristics indicate that the dense-clast pyroclastic flows of August 18 (and probably those of September 16–17) began as hot avalanches that were poorly inflated and slow moving. Minor sand-blasting of brush took place adjacent to the flows, but brush within 10 m of the termini of the flows showed no damage. The andesite blocks that constitute most of the pyroclastic flow deposits have a characteristic orbicular shape with irregular cauliflower-like exteriors and slightly breadcrusted surfaces; they fracture along poorly developed polygonal joints. These features suggest that extensively degassed juvenile andesitic magma

was ejected explosively as hot plastic fragments, which were incandescent upon ejection (according to eyewitness reports and slow-scan video views of the September 16–17 eruption) but were too viscous to attain appreciable degrees of rounding during flight.

The pyroclastic flows of both August 18 and September 16–17 were emplaced as hot juvenile clasts and sparse silt to sand-size matrix material were still hot to the touch when first examined about 40 hours after emplacement. Pyroclastic flows were hot enough (about 300°C) to char vegetation in contact with the andesite clasts. Vegetation adjacent to the flows was not charred, indicating low temperatures in the air around the descending flows. For several weeks after the August 18 eruption, a linear axial zone of weak fumarolic activity was present in the 2-m-thick center channel of the pyroclastic-flow deposits in Wilton's Walk near their distal end.

The flows were not fluidized, and gas played little if any role in their transport. Most flows probably moved so slowly that little or no air ingestion took place. Absolute flow rates are unknown; however, at a point 2.4 km from the crater rim and 0.8 km from the terminus of the flow in Wilton's Walk, a pyroclastic flow overrode a 4-m-high vertical bank and left a thin overbank lag deposit of juvenile andesite clasts. The pyroclastic flow at this point is estimated to have been traveling at about 22 km/hr on the basis of the relation $v^2 = gh$ where h is the height of bank, v is the velocity, and g is the acceleration owing to gravity (see Cas and Wright, 1987, p. 179 for a discussion).

Well-developed levees and surface ridges along the margins of the pyroclastic flow indicate that the denser main body of the flow had sufficient yield strength and momentum to push earlier deposited material aside. As the flows came to rest, clasts were dumped and formed steep lobate flow fronts 1 to 1.5 m high (figs. 4 and 5).

Fallback of pyroclastic material on interfluvial areas high on the cone deposited a veneer of chiefly juvenile debris lying at the angle of repose. The steep dip favored formation of numerous small secondary pyroclastic flows for several days after the August 18 eruption, and so most of these were soon stripped away.

Pyroclastic flows channeled into valleys and other topographic depressions attained a maximum thickness of 1 to 3 m. The thickest flows formed in Wilton's Walk, the widest gully on the southeast flank of Crater Peak. Most of these pyroclastic flows were later covered by September 16–17 lahars that descended to an elevation of about 975 m (3,200 ft). Considerable erosion in this drainage by lahars and associated flooding during the June 27 event cut deep channels from 975 m (3,200 ft) down to the 790 m (2,600 ft) level prior to deposition of the August 18 pyroclastic flows.

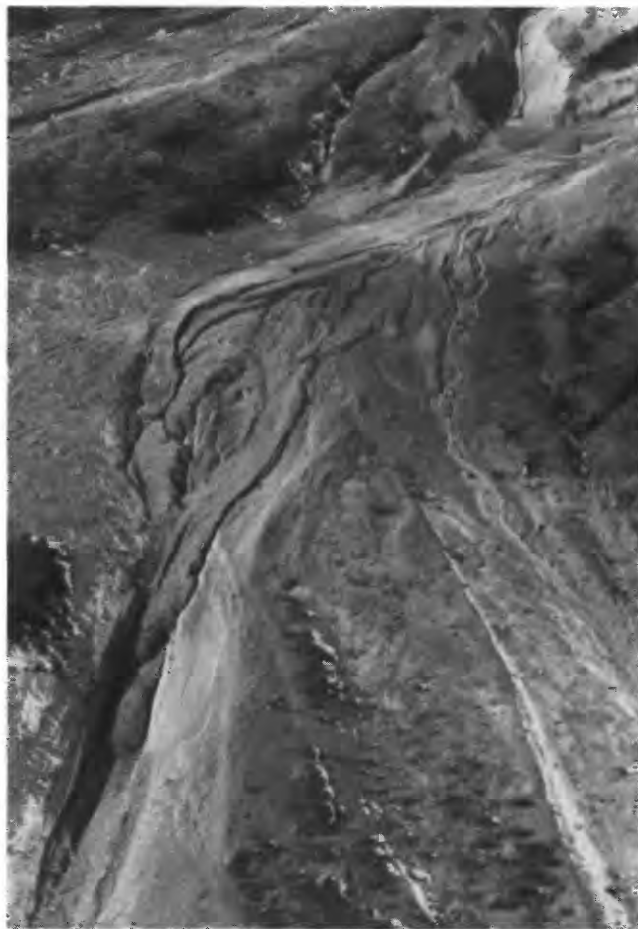


Figure 5. View to northwest of terminus of overlapping August 18 pyroclastic-flow deposits in Wilton's Walk. These deposits were produced from eruption of Crater Peak vent, Mount Spurr volcano, Alaska. Width of pyroclastic-flow deposit at terminus is about 15 m.

DISCUSSION

The pyroclastic flows at Crater Peak were probably emplaced episodically during each eruptive episode. The distribution of the small-volume pyroclastic-flow deposits indicates that they did not result from boilover through low points on the crater rim or through column collapse because apparently no pyroclastic flows traveled through the low point of the Crater Peak rim during any of the three eruptions. Indeed, the pyroclastic-flow deposits of August 18 and September 16–17 are downslope from some of the highest parts of the crater rim (2,200 to 2,290 m; fig. 1) on its eastern sector.

Some of the pyroclastic flows, particularly from the September 16–17 eruption, could have originated from asymmetrical collapse of an eruption column, because they were entirely contained within the te-

phra-fall area of that eruption (fig. 1). The short (about 3 km) runout distances of the pyroclastic flows suggests that the debris forming the pyroclastic flows did not rise high in the eruption column. Prevailing winds would seem unlikely to have a strong influence on the collapse pattern of such coarse material.

The zones of accumulation for the August and September pyroclastic flows are each spatially associated with large ballistic bomb fields (fig. 1). This association suggests that ballistic ejection from a confined and inclined vent may have been a factor in their formation. Numerous modest explosions from an inclined or shielded vent may have repeatedly reamed the conduit of high-viscosity andesitic magma that was already extensively degassed and that occupied the uppermost part of the conduit. Some of these explosions may have resulted from clasts too large to be expelled accumulating in the vent until it became temporarily choked. The resulting increase in pressure caused the expulsion of a mass of coarse debris. These slugs of stiff, brittle, and fragmenting magma were expelled to altitudes generally high enough (probably <1 km) to clear the northeast and east crater walls (about 310 m) before falling back to form heavy, near-vent tephra deposited near the angle of repose on steep slopes. This material was quickly mobilized into coarse avalanche-like pyroclastic flows that moved downslope and into gullies that lead out of the fall-out zone.

The difference in distribution of the August and September bomb fields may reflect a change in the location and inclination of the active vent. Aerial observations indicate that the active vent shifted between eruptions from the north-central part of the crater in June to a location against and slightly beneath the overhanging northwest crater rim by September. The steep north and west walls of the vent area and a resistant east-west buttress immediately south of the vent would block ballistic material going in those directions and favor ejection to the east, as was apparently noticed on the remote slow-scan TV on September 16–17.

The lack of pyroclastic flows associated with the June 27 eruption may reflect ballistic material from this eruption being deposited in the proximal area around the cone on a high, relatively flat, undulating ice field rather than on a steep-sided cone surface. The thick pre-eruption crater fill may also have inhibited the early generation of pyroclastic flows by ballistic fountaining, whereas during the later eruptions an "open vent" situation prevailed.

The lack of correspondence between the distribution of pyroclastic flows and the low part of the crater rim and the possibility of ballistic emplacement of pyroclastic flows has important implications for hazard analyses of volcanoes prone to vulcanian or

subplinian eruptions. Certainly drainages downslope of low points on crater rims would always be areas susceptible to be filled with pyroclastic flows that utilize the breach during a boilover. But the distribution of pyroclastic flows at Crater Peak is a reminder that such flows can be generated by directed ejection and inclined column collapse. Thus hazards may exist on any flank of the volcano, even without major vertical column collapse. This possibility needs to be considered in hazard analysis.

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Ballistic Showers During Crater Peak Eruptions of Mount Spurr Volcano, Summer 1992

By Richard B. Waitt, Larry G. Mastin, and Thomas P. Miller

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ABSTRACT

Crater Peak vent at Mount Spurr volcano in Alaska ejected large ballistic blocks and bombs southeast from the vent near the end of the August 18, 1992, subplinian eruption. This eruption produced the most extensive bombfield of the 1992 eruptions. Lithic and juvenile andesite ballistic projectiles range in diameter from less than 10 cm to 2 m, and they range in density from 1.20 to 2.75 g/cm³. Lithic ballistic clasts are angular and blocky while cauliflower bombs are common among juvenile clasts, both typical of hydromagmatic explosions. As the eruption waned, ground water from the cone may have poured into the vent along with ejecta fallback and talus to fuel hydromagmatic explosions.

We estimate initial velocity needed to eject blocks to the observed 2 to 4 km and more from the vent by iteratively calculating the points of hypothetical trajectories. Drag influences the range of the blocks but is poorly constrained for irregularly shaped volcanic ballistics because it varies enormously with various combinations of shape, orientation, roughness, and velocity. For ballistics of 25 cm and larger that landed within 3.5 km of the vent, drag coefficients vary between 0.06 and 1.5 for spheres and cubes, and they

yield calculated vent velocities of 155 to 840 m/s; smaller and more distal clasts yield values improbably as much as several times higher. Ballistics larger than 10 cm apparently traveled much of their trajectories in a range of very high Reynolds numbers, between 4×10^5 and 1×10^7 , where drag on smooth spheres is as low as 0.06. In contrast to previous studies on volcanic ballistics, we infer that with low drag coefficients relatively small ballistic clasts can achieve fairly long range even with moderate ejection velocities.

Many impact craters along the southeast azimuth from the vent are aligned not directly southeastward but more nearly south. This southward deviation is typically 20° to 40°, the maximum deviation 67°—despite westerly wind that tended to shift trajectories east. A pitched baseball curves in the direction of applied spin (Magnus effect). Thus many projectiles apparently had strong angular velocity about near-vertical axes. The deviant projectiles at Crater Peak apparently acquired clockwise spin (when viewed from above) while being ejected from the vent.

INTRODUCTION

Crater Peak vent at Mount Spurr volcano ejected large ballistic blocks and bombs during all three eruptions of 1992. Ballistics flew north during the June 27 eruption, east-southeast during the August 18 eruption, and east during the September 16–17 eruption. The bombfield of the August eruption, which is scattered over an area of at least 12 square kilometers and extends more than 8 km east-southeast of the vent (figs. 1, 2, 3), is the most extensive and accessible bombfield of the 1992 eruptions.

Between the June 27 and August 18 eruptions, the vent shifted to the northwesternmost part of the crater, tight against and even slightly beneath the steep northwest (southeast-facing) crater wall (fig. 2). This position promoted ejection of debris preferentially southeastward. Thus pyroclastic flows of the August eruption were launched over the southeast crater rim rather than flowing south through a deep, wide breach in the crater rim (fig. 2) (Miller and others, this volume; Waitt, this volume).

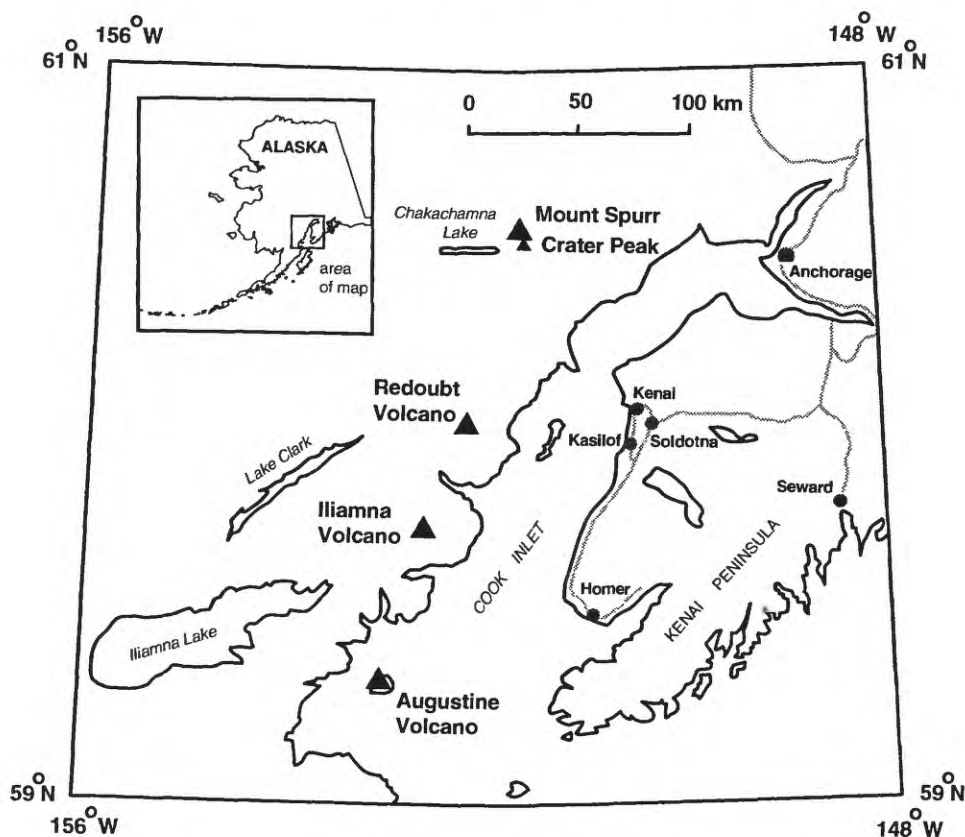


Figure 1. Index map of Cook Inlet area, Alaska, showing location of Mount Spurr volcano.

On the southeast flank of Crater Peak cone the August 18 bombfield comprises thousands of craters, bombs, and bomb fragments. Visible on 1:30,000-scale aerial photographs, the largest (>2 m diameter) and most concentrated craters extend up to altitude 1,375 m (4,500 ft) at range 1.8 km southeast of the vent. Above that level the slope consists of rock and is too steep for bombs to make impact craters. The abundantly cratered part of the field extends down to altitude 835 m (2,740 ft) at range 3.4 km from the vent. The density of large (2–5 m diameter) bomb craters is five per 100 square meters, estimated from aerial photographs. The actual number of bombs and craters including those below the resolution of the photographs (less than 2 m diameter) is at least an order of magnitude larger.

From field and aerial-photograph observations at several adjacent sites at range about 3.3 km, nearly as many large impact craters (>1 m diameter) per unit area are apparent on pyroclastic-flow deposits of the August eruption as just beyond flow margins. At range 8 km where the ballistics overlapped the field of continuous subplinian fall, the ballistics formed small craterlets in the much finer grained fall deposit. Therefore most of the ballistics came near the end of the

eruption after pyroclastic-flow and fall deposits were emplaced. Yet near range 3.3 km a pyroclastic-flow margin overlapped two impact craters, and so some of the ballistics clearly fell before the most extensive pyroclastic flows had occurred.

CHARACTER OF PROJECTILES

PROPERTIES

Ballistic projectiles range in diameter from about 2 m to 8 cm or less. Those larger than 30 cm are lithic. About half the projectiles smaller than 15 cm are fresh juvenile andesite, like that composing pyroclastic-flow deposits of the eruption (Miller and others, this volume); half are lithic cognate andesite and accidental basaltic andesite. The measured densities of 56 projectiles collected at sites 2.9 to 8 km from the vent show a wide range of 1.20 to 2.75 g/cm³ (fig. 4). If

Figure 2. Distribution of August 18, 1992, eruption products, including field of bomb craters. ➤

a dense-rock equivalent density (ρ_d) of 2.6 g/cm³ (Hall, 1987, table 9) is used for andesite, these clast densities (ρ_c) can be converted into vesicularity percent ($V(\%)$) by the relation $V(\%) = 100(\rho_d - \rho_c)/\rho_d$ (Houghton and Wilson, 1989). The median vesicularity for eight density classes ranges from 0 to 50 percent (fig. 4).

The shapes and densities of gravel-sized clasts reveal something of their history (Houghton and Smith,

1993). Most Crater Peak ballistic clasts are angular and blocky. All clasts denser than 1.9 g/cm³ are but slightly vesiculated, have angular edges and blocky shapes (fig. 5), and are very hard under the hammer—clearly lithic. Most juvenile clasts are less dense than 1.8 g/cm³, have microcauliflower surface texture on unbroken faces, are conspicuously vesiculated, and are completely unaltered. Many juvenile clasts have bro-

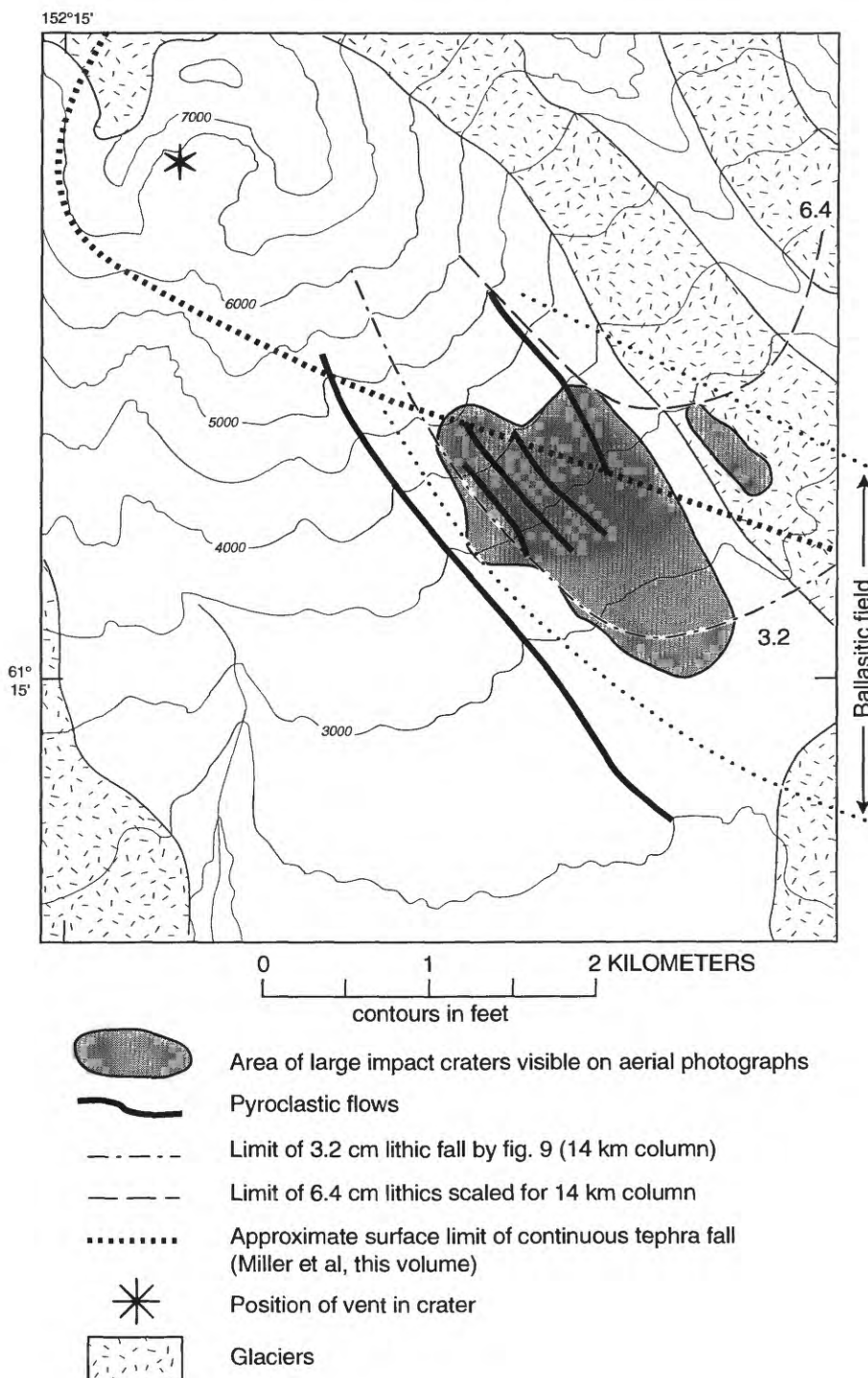




Figure 3. View north of Crater Peak on south flank of Mount Spurr volcano, Alaska. Right (south) margin of field of ballistics thrown by August 18, 1992, eruption approximately shown by dashed line. Field probably extends north beneath area covered by fall and flows of September 1992 eruption. Photograph taken July 27, 1993.

Figure 4. Density distribution of ballistic clasts of August eruption from sites 2.9 to 8.0 km from vent. Calculated vesicularity is also shown (at tops of columns), calculated from an assumed dense-rock-equivalent density of 2.6 g/cm^3 . Negative vesicularities result from some basaltic-andesite accidental lithic clasts being denser than the assumed 2.6 g/cm^3 ; the area above the line drawn through the two right-hand columns hypothetically distinguishes them.

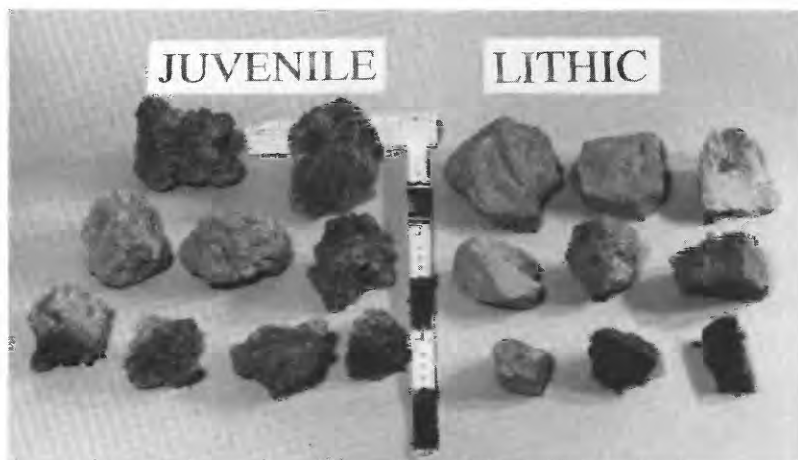
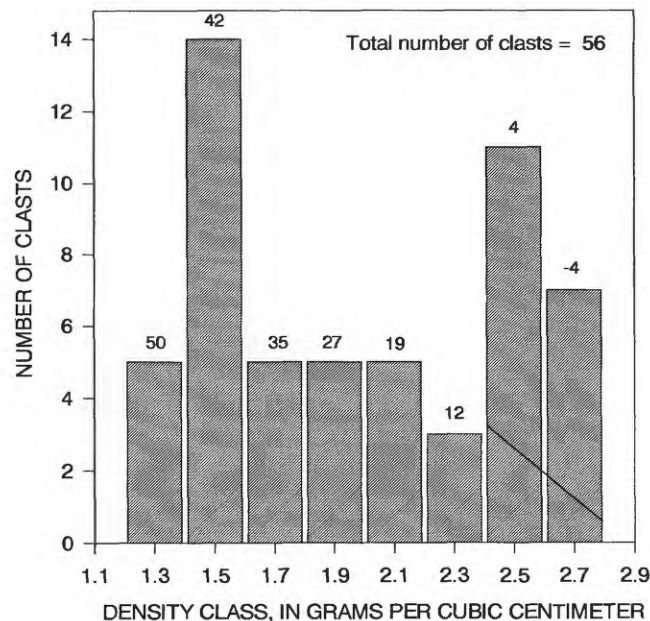


Figure 5. Shapes of small ballistic clasts from the August 18, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska. At center of juvenile group is a small cauliflower bomb, a scarcely vesicular juvenile lithic. The three in lower left are fragments of cauliflower bombs. Bar scale divisions are 5 cm.

ken faces (fig. 5) and show a more vesicular interior and a poorly vesicular exterior. Some juvenile lithic clasts have a more vesicular interior and a scarcely vesicular "lithic" surface. The wide range in density does not vary with throw distance (fig. 6): lithic and juvenile andesite apparently were ejected and traveled together.

Cauliflower-shaped bombs are common among the juvenile clasts (fig. 7). Where best formed, the distinctive texture is nearly identical to bombs photographed from clearly hydromagmatic deposits at Surtsey in Iceland, West Eifel maars in Germany, and Ukinrek maars in Alaska (Lorenz, 1974; Lorenz and Büchel, 1980; Self and others, 1980; Lorenz and Zimanowski, 1984). At Crater Peak, bombs 10 to 15 cm in diameter that were embedded in turf 3 to 4 km downrange

are cauliflowered on top and bottom alike. Because their bottoms were not notably flattened nor their fine texture destroyed, they must have cooled enough to be no longer plastic and deformable upon impact—though they remained hot enough to scorch vegetation beneath them. Many juvenile clasts 2 to 5 cm in diameter are *fragments* of cauliflower bombs. Typically a third to a half of such a clast is gray to light brownish gray (10YR 6/1 to 6/2) and has a micro-cauliflower texture, while the clearly broken faces bounded by sharp angles on other parts of the clast are dark gray to very dark gray (5YR 4/1 to 3/1). In most such clasts, vesicularity notably decreases from 50 to 30 percent at broken (formerly interior) faces to 15 to 5 percent within 5 mm of the microcauliflowered surficial face.

The size of largest projectiles decreases sharply with throw distances (range) southeast from the vent (fig. 8). Three axes—the maximum, intermediate, and minimum diameters—were measured for many clasts, but because most of the projectiles are subequant, only the intermediate axis is recorded for most of them. The outer limit of large projectiles is fairly sharp: those 40 to 200 cm in intermediate diameter stop at horizontal range 3.2 km, those 20 to 40 cm in diameter stop at 3.6 km, and those 10 to cm in diameter stop at 3.8 km. Projectiles 1 to 10 cm in diameter lie throughout the range 2.5 to 8 km. The plot can be interpreted in several ways. First, the steep decrease in large clasts and the long range of small ones may reflect a steeply logarithmic decay of a single population. But second, an outer limit of one population of large (>15 cm) clasts may exist whereas a second population of 1- to 10-cm clasts extends out to at least 8 km.

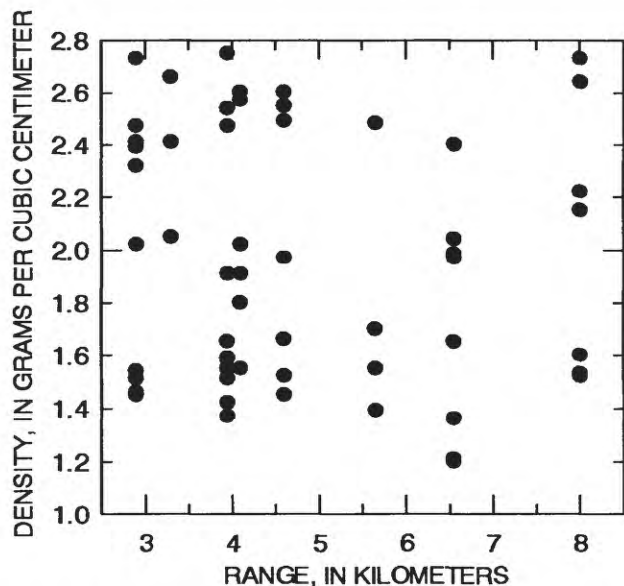


Figure 6. Density of ballistic clasts of August 18, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska, plotted against range of throw.

Figure 7. Cauliflower bombs thrown by August 18, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska. One on left is 12 cm across, that on right 9 cm.



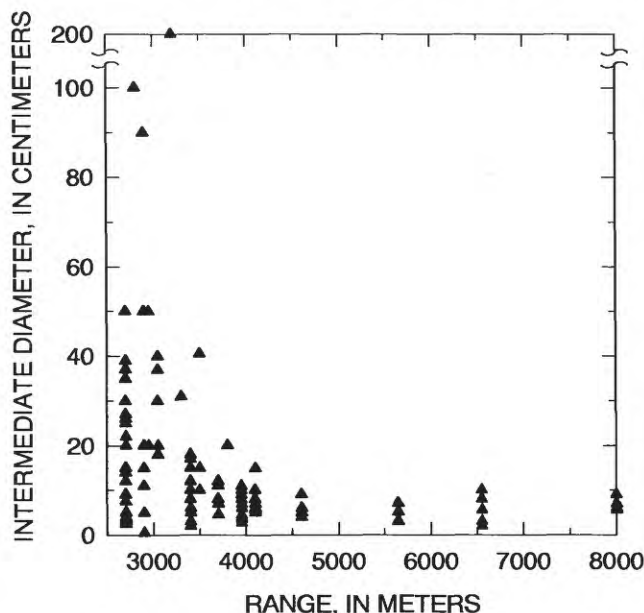


Figure 8. Intermediate projectile diameter plotted as a function of range of throw. Projectiles from August 18, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska.

UNEVEN TEMPERATURES

Many projectiles, both lithic and juvenile, landed in a mat of ground-hugging tundra plants such as caribou moss and scrub blueberry during growth season. Many of the vesicular, clearly juvenile clasts larger than about 10 cm that landed within 4 km of the vent notably scorched the vegetation on which they came to rest—as did juvenile clasts that tumbled off pyroclastic flows and descended into vegetated areas (Miller and others, this volume; Waitt, this volume). Yet beneath lithic ballistic clasts 5 to 50 cm in diameter, identical vegetation was pulverized but scarcely if at all scorched. Among the clasts of the ballistic shower(s), the juvenile ones clearly were much hotter than the lithic ones.

This pattern indicates that the juvenile and lithic clasts did not reach thermal equilibrium before ejection or during flight. The lithic clasts must not have resided at the vent long enough to acquire as high a temperature as the juvenile clasts before explosion(s) ejected them both.

Figure 9. Impact craters of blocks thrown ballistically by 1992 eruptions of Crater Peak vent, Mount Spurr volcano, Alaska. A, shattered largish block of August 18, 1992, bomb field; a third of the original clast is the angular block outside crater. B, Shattered block of September 17, 1992, that impacted pyroclastic-flow deposit emplaced earlier in same eruption.

IMPACT CRATERS

The largest impact craters range from 3.5 to 5.5 m in diameter and are as deep as 1.5 m. The “target” material is fairly loose, coarse, lithic pyroclastic debris of the 1953 and prehistoric eruptions of Crater Peak, mantled by a thin loam soil and tundra turf (fig. 9). Whereas the largest craters contained bombs as large as 50 cm, most projectile blocks larger than about 35 cm split or shattered upon impact, and parts of them lie outside the crater. When reconstructed, the largest clasts, which lie in and about the largest craters, measure 0.5 to 2 m in intermediate diameter. Craters larger than 3 m in diameter have 2 to 4 rays of ejected material radiating as much as 5 m downrange from the crater rim through 90° to 160° arcs. Clasts as small as 8 cm made impact craters as much as three times their diameter.



RELATION TO WIND, DISTINCTION FROM PROXIMAL FALL

The south limit of fall from the convective vulcanian column (Neal and others, this volume) is uncertain in the area within 10 km of the vent. The south limit of continuous fall is near the apparent north limit of the ballistic field. But within a few kilometers of vent, clasts that fell from the column can be difficult to distinguish from small clasts that moved on ballistic trajectories. During the Mount St. Helens eruptions of 1980, fall from eruption columns occurred through low-level winds oblique to the high-level winds driving the plume. Along the upwind *lateral* margin in proximal areas, the fall deposit was typically *discontinuous*—pumiceous clasts of sizes between medium pebbles and small cobbles and denser lithic clasts of small- to medium-pebble size (Waitt and others, 1981, figs. 269–273). At Crater Peak (Mount Spurr) on August 18, 1992, low-level winds were east-northeastward, oblique to the upper-level winds carrying the plume east and east-southeast (table 1).

Within the ballistic field 2 to 3.5 km from the vent lie clasts of all sizes from 2 m to 1 cm (fig. 8). The large (>40 cm in diameter) clasts, none of them low-density pumice, are unarguably ballistics, as probably are lithic clasts down to 15 cm or smaller. Yet there is no distinguishable minimum-size limit of bal-

listics detectable in the field. Did the 1- to 10-cm clasts depicted on figure 8 arrive as ballistics or as fall? The discontinuous clasts smaller than 8 cm appear similar to the proximal falls at Mount St. Helens along the upwind lateral margins. Yet at Crater Peak, lithic clasts as small as 8 cm formed impact craters as far as 8 km from the vent in the much finer fall deposits there. Because the larger clasts have higher terminal velocities, they could not have traveled with the finer fall materials without landing first, not last. Therefore the larger clasts that formed small impact craters clearly were ejected later, and thus appear to have been ballistic projectiles. Without additional highly specific field data, distinction between the coarser ballistics and the finer fall is somewhat arbitrary.

Sparks (1986) modeled eruption columns whose diameters typically increase upward as a funnel shape, and Carey and Sparks (1986) analyzed fall from such columns (fig. 10). The highest reach of the Crater Peak eruption column of August 18, 1992, was 14 km (Rose and others, this volume). For a column 14 km high and a downrange wind velocity of 10 m/s, Carey and Sparks's (1986, fig. 14c) model predicts the outer limit of fall of 3.2-cm lithic clasts near the outer edge of the large August 18 impact craters; an extrapolated limit for 6.4-cm clasts lies farther ventward of the main ballistic field (fig. 2). Thus by the plume model of Carey and Sparks (1986), lithic clasts only as large

Table 1. Wind data for August 18, 1992, eruption in vicinity of Mount Spurr volcano, Alaska.

[A. Projected winds beginning 7:00 p.m. August 18, 1992]

Altitude (ft)	Azimuth toward (degree)	Wind velocity (m/s)	Wind toward azimuth 135 (m/s)
5,000	17	4.8	-2.3
10,000	71	5.2	2.7
18,000	72	11.1	5.0
23,000	96	11.1	8.7
30,000	104	14.7	12.7

Source: National Weather Service

[B. Balloon sounding above Anchorage at 4:00 p.m. August 18, 1992]

Altitude (ft)	Azimuth toward (degree)	Wind velocity (m/s)	Wind toward azimuth 135 (m/s)
6,000	350	6.2	-3.5
8,000	30	5.1	-1.3
12,000	50	7.2	0.6
16,000	70	9.3	3.9
20,000	85	10.8	6.9
25,000	115	15.9	15.0
30,000	130	22.1	22.0

Source: National Weather Service

as 3 to 6 cm should have fallen as tephra within the area of the August 18, 1992, Crater Peak ballistic field. Were column height somewhat higher or winds higher than reported, somewhat larger clasts could have fallen in the shaded area shown in figure 2. By stretching the envelope of uncertainty, one can imagine lithic clasts as large as 8 cm being possibly related to airfall near the southeast part of the shaded area, but it doesn't help to explain larger clasts, which therefore must be ballistics.

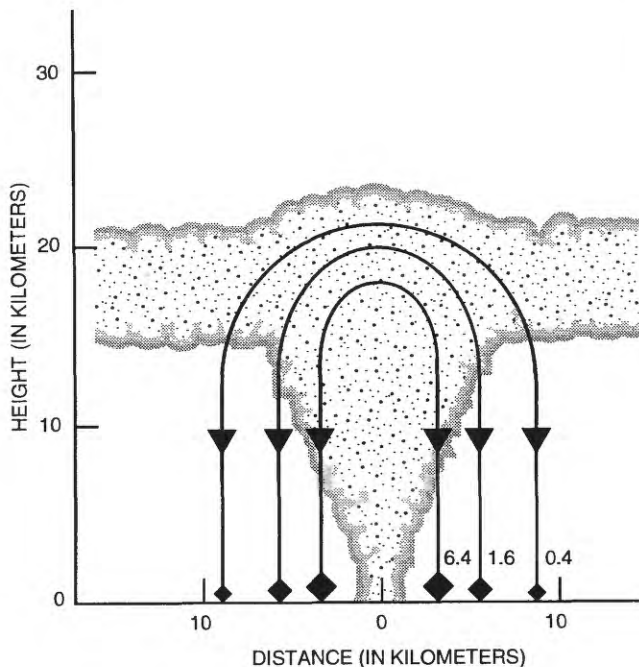


Figure 10. Fall of lithic (density = 2.5 g/cm^3) clasts from modeled column roughly 1.5 times the height of the maximum reported for the August 18, 1992, eruption of Crater Peak, Mount Spurr volcano, Alaska eruption (illustration from Carey and Sparks, 1986, fig. 6). Scale is true (no vertical exaggeration). Upward flare of model, about 23° from the vertical, is about that shown by Crater Peak eruption photographs of August 18, 1992. Symbols show theoretical outer limit of lithic clasts 6.4, 1.6, and 0.4 cm in diameter. Outer limit of 6.4-cm lithic clasts is 3.3 km from plume axis, the limit of 1.6-cm clasts 5.5 km. The modeled column is much higher than the Crater Peak column, thus the distance of lateral fall from plume axis is maximum-limiting case.

RECONSTRUCTED VELOCITIES OF BALLISTIC PROJECTILES

Some elements of the trajectory of a ballistic projectile can be inferred from its impact point, including minimum ejection velocity. From minimum velocity of an ejected block, vent pressure required to expel the block can be calculated within very broad limits and with considerable uncertainties.

METHOD

Wilson (1972) reconstructed volcanic-ballistic trajectories utilizing data on drag forces on cylinder-shaped projectiles. To reach a given distance from the vent, a small projectile (say 10 cm) requires a far higher launch velocity than does a large (1 m) one, because for a small particle the ratio of surface area to mass is much higher, thus drag per unit mass much higher.

The velocity required to eject blocks to the distances observed at Crater Peak is estimated by computing ballistic trajectories for blocks of known properties using an algorithm outlined in Wilson (1972). Components of velocity (v) in the horizontal (x) and vertical (z) directions are given by:

$$v_x = v \cos \theta \quad (1a)$$

$$v_z = v \sin \theta \quad (1b)$$

where θ is trajectory angle measured from the horizontal. At any given point along a particle's trajectory, the rates of change of v_x and v_z are given by:

$$\frac{dv_x}{dt} = \frac{-v_x \rho_a v A C_d}{2m} \quad (2a)$$

$$\frac{dv_z}{dt} = \frac{-v_z \rho_a v A C_d}{2m} - g \quad (2b)$$

where ρ_a is air density, A is cross-sectional area of the block, C_d is drag coefficient on the block, m is the block's mass, g is gravitational acceleration, and t is time. These equations are identical to equations 3 and 4 in Wilson (1972), though variables have been rearranged and different symbols are used to designate them. The variables A and m depend on the shape and size of the block, assumed in this case to be spherical (that is, $A = \pi r^2$ and $m = (4/3)\pi r^3 \rho_r$), with block radius (r) equal to the intermediate measured block semidiameter, and rock density (ρ_r) equal to 2.5 g/cm^3 . Air density (ρ_a) varies with height and was recalculated at each point in the trajectory using the following equation, which is a best-fit curve through published altitude-density data (Ahrens, 1991, p. 546):

$$\rho_a = .0012232 - 1.1389 \times 10^{-7}(z_0 + z) + 3.3036 \times 10^{-12}(z_0 + z)^2 \quad (3)$$

where z in this case is the vertical position above the vent, and z_0 is the altitude of the vent above sea level (air densities in g/cm^3 , altitude in meters).

Equations 2 above were integrated throughout the trajectory using a fourth-order Runge-Kutta method, as were those of Wilson (1972). This algorithm was made more versatile by incorporating corrections for tailwind velocity (w) downrange from the crater, and an altitude difference (ξ) between the launch and landing sites (ξ is <0 for sites that are lower than launch site, like all calculated herein). The correction for altitude difference was included by simply calculating the trajectory until the vertical position (z) of the block reached its landing elevation (z_f), which generally differs from launch altitude ($z=0$). The wind correction was made by substituting $(v_x - w)$ for v_x in equation

2a, and $\sqrt{v_z^2 + (v_x - w)^2}$ for v in equations 2a and 2b.

DRAW

Drag coefficient C_d influences the final range of the blocks enormously but is poorly constrained for use with volcanic ballistics, for the experimental data show that drag on a projectile varies greatly with its shape, orientation, and roughness. Experimental results (for example, Hoerner, 1965) indicate that C_d for variously shaped objects traveling through fluids varies with two dimensionless numbers that indicate flow regime: Reynolds number (Re) and Mach number (M). Reynolds number relates the relative importance of viscous versus inertial forces in the fluid by the relation (for external flow) $Re = v d \rho / \eta$, where v is particle velocity, d is particle diameter, ρ is fluid

density, and η is fluid viscosity ("fluid" here being air). The range of particle sizes and calculated velocities give Reynolds numbers (Re) greater than or equal to 4×10^5 . Drag coefficient C_d for smooth spheres is 0.06 at Re of 4×10^5 and rises gradually to 0.2 at Re greater than about 1×10^7 (fig. 11) (Hoerner, 1965, figs. 3-10, 3-11; Achenbach, 1972, fig. 4).

Most authors analyzing volcanic ballistics have used high drag coefficients. In experiments on falling pyroclasts, Walker and others (1971) showed that for terminal velocities of small particles falling at low velocities (hence low Reynolds numbers), the relatively high drag coefficients for cylinders were more appropriate than the lower coefficients for spheres. Thus Wilson (1972) and Fagents and Wilson (1993) assumed that C_d measured for cylinders applied to irregularly shaped blocks. For such objects several other authors have also assumed rather high values of C_d , usually around 1.0 and no lower than 0.7 (Minakami, 1942; Fudali and Melson, 1972; Steinberg and Lorenz, 1983; Self and others, 1980; Wilson, 1980; Mastin, 1991).

Mach number ($M \equiv v/c$) gives the speed of the object (v) relative to the fluid's sonic velocity (c). For air under ambient conditions, the sonic velocity is about 350 m/s. As long as objects are traveling at speeds less than about 70 percent of the sonic velocity (that is, less than about 250 m/s), C_d varies with Reynolds number only. But at Mach numbers above about 0.7, the production of sound waves (and shock waves when $M \geq 1$) dissipates energy. Drag coefficients rise dramatically for Mach numbers in the range 0.7 to 1.5 and stabilize for Mach numbers above 1.5 (fig. 12).

To calculate ballistic trajectories for the irregularly shaped blocks at Crater Peak, we assumed drag coefficients for spheres (which have low drag) and

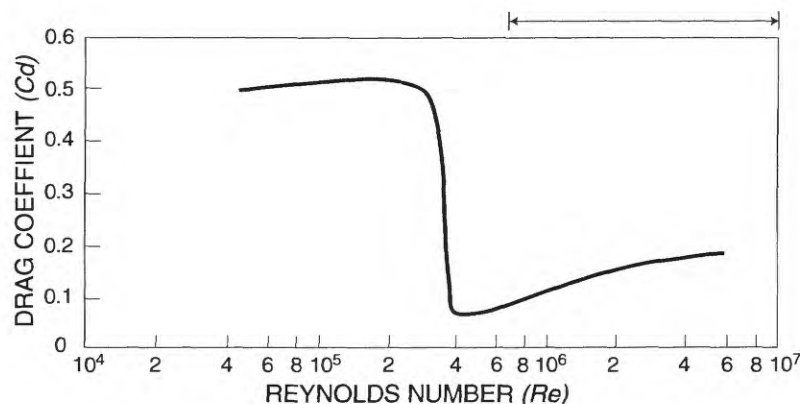


Figure 11. Low drag coefficients for smooth spheres at very high Reynolds numbers (from Achenbach, 1972, fig. 4). Bar with arrowheads shows range of Reynolds numbers calculated from initial velocities computed for ballistic clasts from the 1992 eruptions of Crater Peak, Mount Spurr volcano, Alaska.

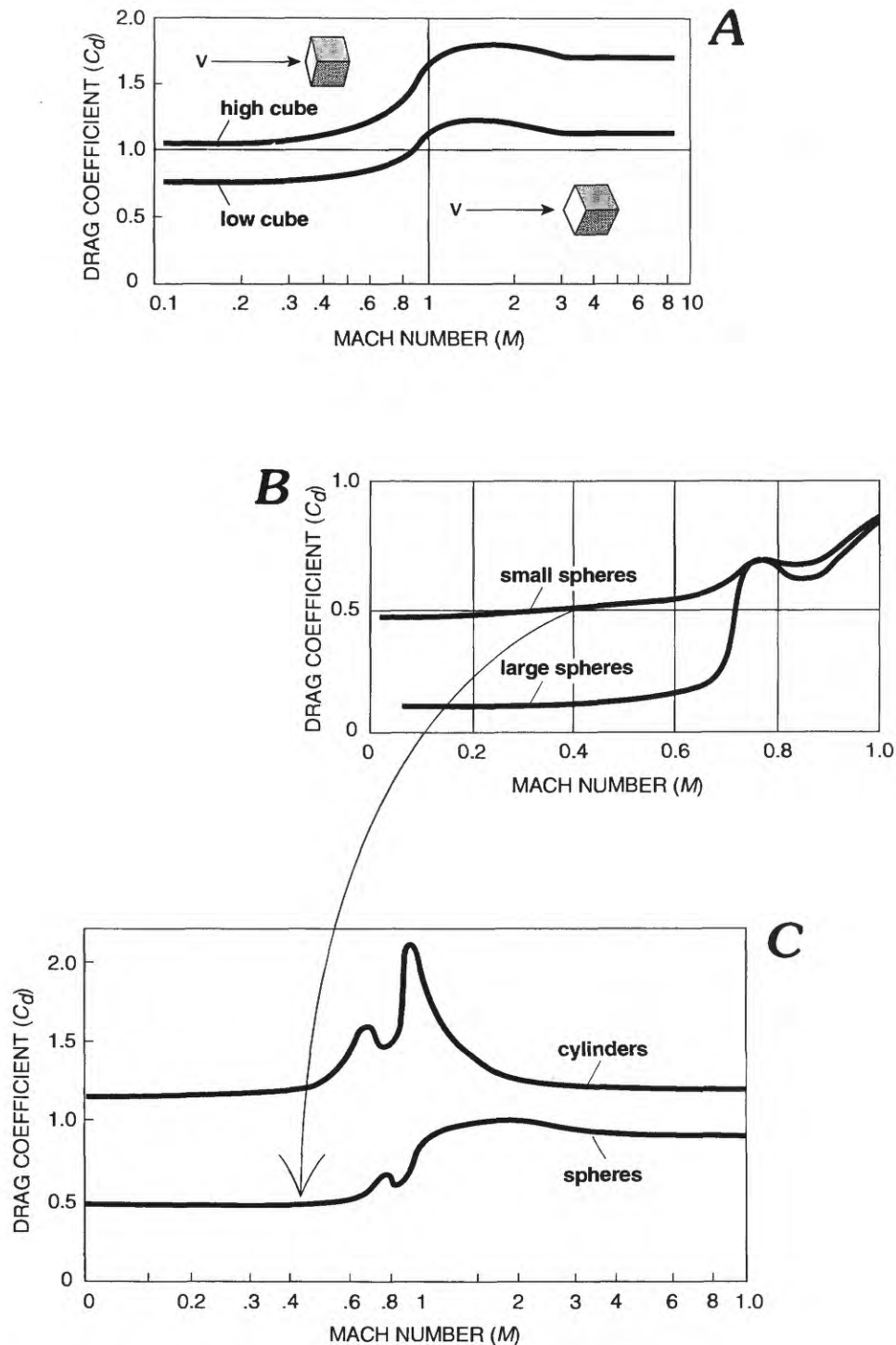


Figure 12. Drag coefficients for (A) cubes at transonic velocities, (B) spheres at velocities approaching sonic, and (C) cylinders (top curve) and spheres (lower curve) at transonic velocities. Top curve in (A) represents drag coefficients for cubes oriented with one face perpendicular to the direction of motion. Bottom curve represents drag for cubes oriented with the vertex between three faces at the leading edge of the cube. Top curve in (B) represents drag coefficients for small spheres at Reynolds numbers below the critical transition ($Re \approx 2 \times 10^5$), whereas the lower curve represents drag coefficients for large spheres at Reynolds numbers above the critical transition. The ballistic model calculates Reynolds number at each point in the trajectory and uses the drag appropriate at the calculated Reynolds and Mach numbers (for all blocks analyzed in this study, the appropriate drag coefficient is represented by the large-sphere curve). Calculations for supersonic velocities used drag coefficients for spheres illustrated in (C). Modified from Hoerner (1965).

for cubes (with high drag) (table 2). At each point in the trajectory the Mach number (taken from fig. 12) was used for the next calculation. Results from this program using some typical input parameters are shown in figure 13 for a launch angle of 45° .

We back-calculate initial velocities for the largest one or two projectiles from each small sampled area. Altitude of impact site (Z_{final} or Z_f) was measured in the field by altimeter and plotted on a topographic map, from which range (x_f) was read. We assume a lithic particle density of 2.5 g/cm^3 and launch angles between 30° and 35° —the most efficient launch angles calculated by the program under the conditions of range less than 4 km and with impact altitude far below launch altitude. We assume downrange windspeed of 0 m/s; calculations with the highest permissible windspeed (10 m/s) do not substantially affect results.

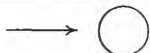
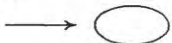



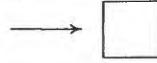
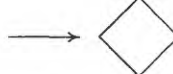
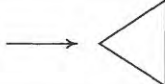
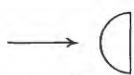
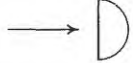
Our results for ballistic clasts thrown by Crater Peak are shown by table 3, calculated using values of C_d measured for spheres and cubes. At the very high Reynolds numbers ($\text{Re} > 4 \times 10^5$) typical for Mount Spurr

ballistics, C_d for smooth spheres can be as low as 0.06 (fig. 11) but for very rough objects may be much higher.

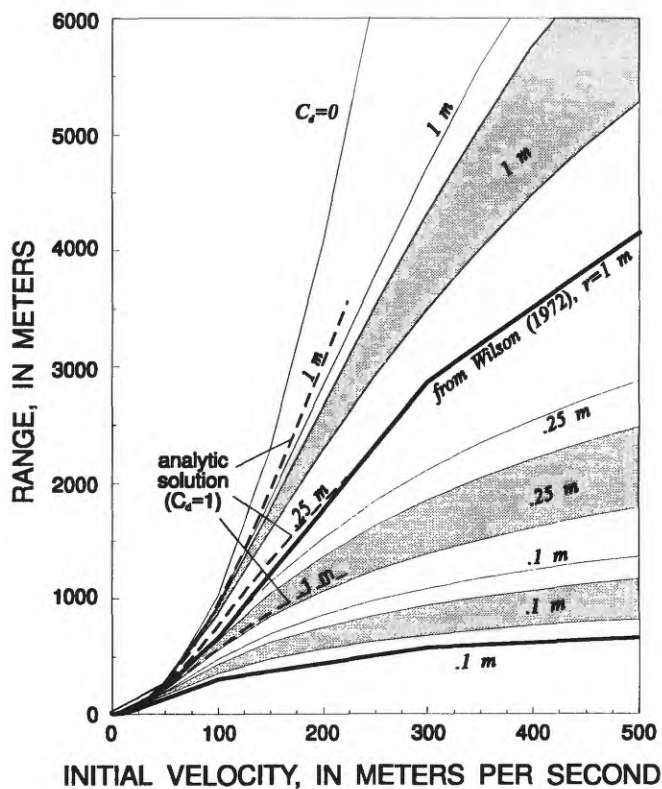
For comparison to the Mount Spurr data is a synopsis of ballistics thrown by other eruptions, calculated with a constant high drag coefficient ($C_d=1$) used by investigators of those eruptions. At drag coefficient $C_d=1$ (table 3, col. 8), our calculated initial velocities at Ukinrek maars and Inyo Craters—where ballistic clasts are fairly large and maximum range only about 700 m—are near the values 85 to 100 m/s reported by Self and others (1980) and Mastin (1991) respectively. The calculated initial velocities for the large ballistics at moderate to long distance at Asama in 1938, Arenal in 1968, and Ngauruhoe in 1975, (195–322 m/s) are also well below sonic velocities, more so than originally reported (Minakami, 1942; Fudali and Melson, 1972; Nairn, 1976).

If we use a constant $C_d=1$ at Crater Peak (Mount Spurr), subsonic velocities are given only by the largest (2 m) clast thrown to moderate range, whose calculated v_i is 183 m/s (table 3, col. 8). The only other

Table 2. Drag coefficients for three-dimensional shapes.

Shape		Laminar flow	Turbulent flow
Sphere		0.47	0.10
Ellipsoidal body with circular cross section			
2 : 1		0.27	0.06
4 : 1		0.20	0.06
8 : 1		0.25	0.13
Disk			1.17
Cube			1.05
			0.80
60° cone			0.49
Solid hemisphere:			
Convex surface to flow			0.38
Flat face to flow			1.17

Adapted from Mironer (1979, table 8.2)



clasts yielding subsonic (barely) initial velocities are 50 cm or more in diameter. Ballistics 37 to 39 cm in diameter even in the *nearest* part of the sampled range (2.7 km) yield much higher supersonic velocities (>400 m/s). Calculated velocities above 700 m/s (Mach number ≈ 2) for projectiles of smaller size and (or) those thrown to greater range seem unreasonably high.

If now these clasts are calculated using a variable drag coefficient as low as $C_d=0.06$ —the values for smooth spheres when Re exceeds 4×10^5 (fig. 11)—ballistic clasts with diameters as small as 15 cm have

Figure 13. Initial velocity of blocks ejected plotted as a function of range assuming wind velocity $w=0$ m/s, impact altitude below vent altitude $\xi=0$ m, density of projectiles $\rho=2.5$ g/cm³, and the initial trajectory angle is 45° from horizontal. Shaded regions represent results using drag coefficients for cubes, thin solid lines represent results using drag coefficients for spheres. Heavy solid lines are values from Wilson (1972) based on drag coefficients for cylinders. Dashed lines show values from the approximate analytical solution of Minakami (1942) (see also Self and others, 1980; Mastin, 1991) that assumes a constant drag coefficient $C_d = 1$.

Table 3. Ballistics of August 1992 eruption of Crater Peak, Alaska.

[Vent altitude, 1,952 m. C_d , drag coefficient; Re , Reynolds number; M , Mach number]

Station No. 1993	Range (m)	Altitude (m)	Altitude difference (m)	Azimuth (degree)	Diameter (cm)	INITIAL VELOCITY			
						Vacuum $C_d=0$ (m/s)	Fixed C_d $C_d=1$ (m/s)	C_d variable with Re and M	
								Sphere (m/s)	Cube (m/s)
.01	2700	976	-976	140	37	140	432	156	342
.02	2700	976	-976	140	39	140	398	156	319
.02	2700	976	-976	140	25	140	1025	165	761
.03	2700	968	-984	140	9	140	Infinite	194	Infinite
.04	2900	961	-991	133	50	145	331	160	277
.05	2950	930	-1022	133	50	146	340	162	283
.06	2950	930	-1022	133	20	146	3400	184	2411
.08	3200	891	-1061	132	200	154	183	157	175
.09	3400	857	-1095	133	17	159	Infinite	219	Infinite
.10	3400	851	-1101	134	18	159	Infinite	216	Infinite
.11	3300	866	-1086	134	31	156	914	188	839
.12	3500	836	-1116	137	40.5	162	672	189	534
.12	3500	836	-1116	137	15	162	Infinite	245	Infinite
.13	3700	817	-1135	138	11	167	Infinite	604	Infinite
.14	3800	787	-1165	139	20	169	Infinite	248	Infinite
.15	4100	790	-1162	140	15	177	Infinite	582	Infinite
.19	3950	747	-1205	137	11	172	Infinite	936	Infinite
.20	6550	628	-1324	134	10	231	Infinite	Infinite	Infinite
.24	8000	686	-1266	121	9	262	Infinite	Infinite	Infinite
Asama 1938	3500		-1100		100		266	(Minakami, 1942)	
Ngauruhoe	2800		-980		80		237	(Nairn, 1976)	
Arenal	5000		-500		200		291	(Fudali and Melson, 1972)	
Ukinrek	700		0		200		87	(Self and others, 1980)	
Inyo	700		40		40		108	(Mastin, 1991)	

subsonic velocities (table 3, col. 9). The larger projectiles (intermediate diameters between 200 and 17 cm) thrown downrange 2.7 to 3.5 km give reconstructed initial velocities between 156 and 248 m/s. For smooth spheres the variable C_d of table 3, column 9 is low throughout most or all of a clast's range, except where required initial velocity exceeds Mach number 0.7, which triggers the use of higher drag coefficients taken from figure 12. Under the drag assumptions for smooth spheres, implied vent velocities for the particles of 18 cm and larger thrown from Crater Peak during its August eruption were between 160 and 250 m/s.

Figure 14 illustrates the highly variable drag for a sphere initially shot at 400 m/s. During the initial part of its flight, drag coefficient is about 0.85 and falls to 0.65 in the first one-sixth of its flight. Then as Mach number drops below 0.7, drag falls off rapidly to 0.2. From there it drops even lower as veloc-

ity reaches a minimum at the top of the trajectory, and then rises back to about 0.2 as velocity increases on the downward leg of the trajectory.

Volcanic particles are not spheres and many have angular faces. Wind-tunnel drag data are readily available (Hoerner, 1965) for cubes. If we test the Crater Peak ballistics with these drag coefficients for cubes, all but the largest block would have required initial velocities above 400 m/s. Many of those exceed even maximum theoretical velocities for eruptions under nearly all reasonable initial pressure conditions (about 2000 m/s, Mastin, 1995). Of the many apparently ballistic projectiles smaller than 18 cm in diameter that landed downrange 2.7 to 8.0 km from the vent (fig. 8), implied initial velocities are 500, 1000, 2000 m/s and higher. The more the calculated values for initial velocity exceed 400 m/s, the less believable they are.

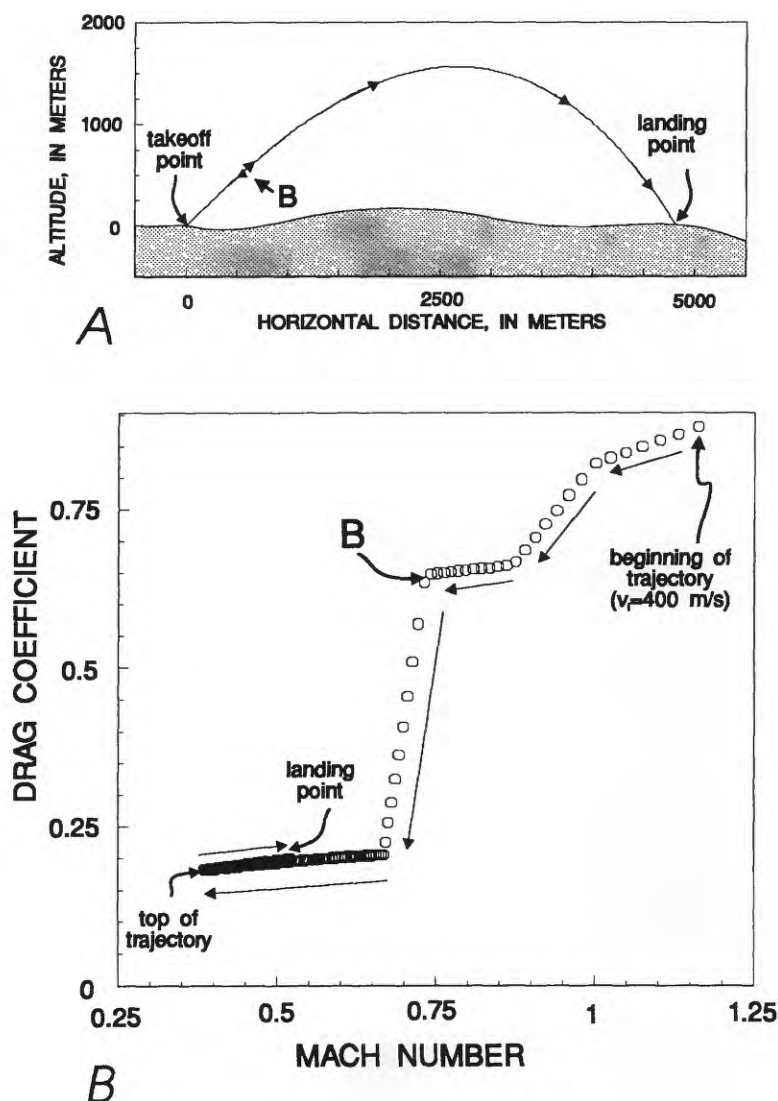


Figure 14. Successively, sharply lower drag coefficients of typical ballistic block thrown from Crater Peak whose velocity is initially supersonic. A, Block's vertical trajectory during flight. B, Change in drag coefficient during flight.

DISCUSSION

Most or all students of volcanic ballistics have followed Walker and others (1971) in assuming high drag coefficients (as for cylinders) in calculating trajectories of volcanic ballistics, and no other experiments have appeared to test that analysis. But calculations using such high drag give impossibly high initial velocities for the smaller of the Crater Peak clasts that are clearly ballistics.

Drag data on cubes compiled by Hoerner (1965) also entails fairly high drag coefficients (fig. 12A). High drag in the computations yields the high initial velocities for a clast even of moderate size (say, 20–40 cm) to be thrown to moderate range (say, 2.5 to 3.5 km). These results are then compounded where Mach numbers exceed 0.6, above which drag further increases significantly (fig. 12B,C).

Velocities of 100 to 250 m/s of clasts 0.1 to 1.0 m in diameter generate Reynolds numbers between about 7×10^6 and 1×10^7 . This is the range of very high Reynolds numbers explored by Achenbach (1972) for smooth spheres, whose drag coefficients he measured to be very low—below 0.2, even as low as 0.06 (fig. 11). Terminal velocities exceeding 40 m/s of ballistics larger than 5 cm in diameter keep them from “mov-

ing leftward” on figure 11 below Reynolds numbers smaller than 5×10^6 and into a range of much higher drag coefficients, even toward the end of their flight. If the flight of a particle occurs mostly or entirely within this envelope of low drag, 10- to 40-cm particles thrown even to range 4 to 8 km would require initial velocities far below sonic velocity (table 3, col. 9). We sense that the much lower drag coefficients permitted by wind-tunnel data on spheres where Reynolds numbers are greater than 4×10^6 is the key to unlock the enigma of far-traveled small ballistics.

While drag coefficients for very rough volcanic particles generally exceed those for smooth spheres, there may be compensating effects. For instance, slight roughening of a smooth sphere increases turbulence in the boundary layer and causes the wake to be much smaller (fig. 15); thus drag is reduced. Photographs illustrating this principle are shown in many textbooks on fluid mechanics (Rouse, 1946, plate 16; Mironer, 1979, fig. 8.19; White, 1979, fig. 7.14). A golf ball “roughened” by inverse topography (“dimples”) travels farther than a smooth one (Jorgensen, 1993, p. 62–64).

The surfaces of volcanic ballistics may be irregular in shape, but few resemble the least aerodynamic geometric solid (cube) tested in wind tunnels

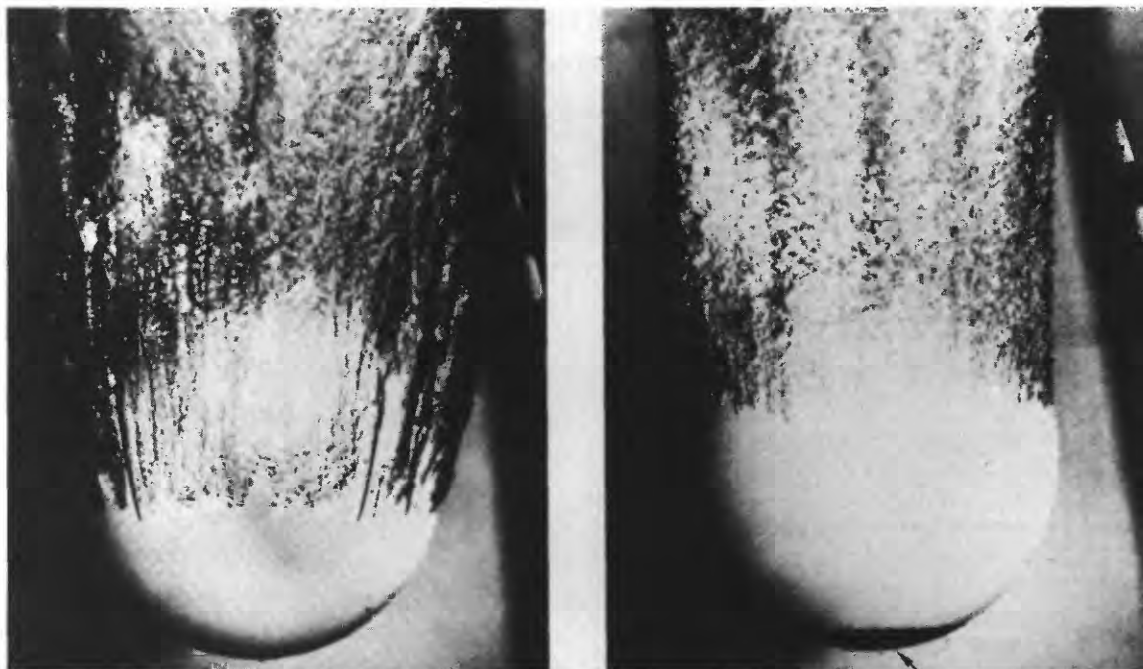


Figure 15. Bowling ball in two experiments generating similar Reynolds numbers. The roughened surface of ball on right (sandpaper patch at base [arrow] causes “premature” boundary-layer turbulence and consequent decrease in area of wake in lee of sphere, compared with that of smooth ball. The smaller wake decreases the relative vacuum behind the object that exerts a small backward force on it, and hence decreases total drag. From Waugh and Stubstab (1973, fig. 15.6).

(table 2). Until actual volcanic ballistics are tested in wind tunnels, one can only infer that drag coefficient will be somewhere between about 0.06 and 1.5 at substantially subsonic velocities. If the lower end of this enormous range proves at all close to reality, the evidently "impossible" great range achieved by some smallish ballistic particles will be partly explained.

FACTORS BESIDES DRAG

Other factors besides a cannonlike ejection into a continuous air-drag environment must influence the great distance traveled by many of the smaller blocks. These other factors could include: (1) much higher tailwinds than assumed in the model, (2) convective lift during flight, (3) large blocks traveling mostly intact but then disaggregating before landing, (4) smaller blocks drawn behind larger blocks or ejected as aggregates with other blocks and thus collectively subject to less drag than imposed on blocks separately, and (5) shock waves and other forward air movement attending ejection of blocks, which reduce air drag during the first part of the trajectory. These five factors are examined below.

Using a range of hypothetical tailwind velocities, we calculate initial ejection velocities for the 40.5-cm block thrown 3.5 km from the vent (table 4). Even with a low drag coefficient appropriate for smooth spheres and the lowest appropriate for cubes, tailwind velocities exceeding 35 m/s are required throughout the trajectory to bring ejection velocities within the less than 400 m/s range of velocities previously observed. With the higher drag coefficients for cubes, required tailwind velocities would be about 60 m/s. Such hypothetical wind speeds are 5 to 30 times those shown by the NOAA data for August 18, 1992, (table 1).

Table 4. Hypothetical effect of tailwind in carrying blocks to greater distance.

[Example of a 40.5-cm-diameter block thrown to range 3,500 m and landing 1,116 m below vent altitude]

Wind Velocity (m/s)	Initial Velocity (m/s)
0	1200
10	770
20	575
40	330
60	210
80	160

Convective lift cannot explain the distance traveled by the larger blocks. The terminal velocity of blocks 40 cm in diameter is 100 to 300 m/s in air (depending on drag coefficient chosen) as calculated from a standard equation (Mironer, 1979, p. 277). Only updrafts a large fraction of terminal velocity would greatly lengthen a block's trajectory. Convective updrafts of tens of meters per second or more occur in sustained plinian eruptive columns (Woods, 1988), but it is unlikely that transient bursts such as at Crater Peak could sustain such drafts. Yet clasts smaller than 5 cm diameter, whose terminal velocities are less than 50 m/s, could be somewhat affected by small updrafts. Perhaps this factor, operating during the subplinian phase of the first 3 hours of the eruption (I'ual and others, this volume) especially in an upward-funneling column as analyzed by Sparks (1986) and by Carey and Sparks (1986), could explain the many lithic and juvenile clasts smaller than 5 cm scattered throughout the ballistic range as elements of fall. Yet because most of the ballistics apparently were ejected near the end of the eruption, some traveling as far as 8 km from the vent even after the subplinian fall had ended, sustained updrafts seem unimportant for many of the small ballistic clasts.

If larger blocks somehow disaggregated in flight just before landing, the ballistic blocks should lie in clusters of broken fragments of like lithology. Such a clustering was nowhere decipherable, except purely hypothetically if applied to juvenile clasts scattered throughout the impact area. The many fragments of cauliflower bombs seemed randomly and widely dispersed, as though they were shattered upon explosion *at the vent*, not near the end of their flightpath. Some juvenile clasts are completely rimmed with quenched margins (figs. 5, 6) and clearly did not fragment in the air or on impact. In-flight fragmentation therefore does not help explain the large distances traveled by many of the blocks.

Numerous images of explosive "cocktail plumes" captured on movie film, video, and photographs (for examples, Williams and McBirney, 1979, p. 251; G. Rosenquist in Voight, 1981, p. 75-76) show each fingerlike projection to be led by a large block in whose wake trails a cloud of smaller debris. At Crater Peak this could help explain many small blocks within the 3.5-km-range limit of large blocks. Yet the smaller blocks extend much farther (fig. 8), and they are about as plentiful just beyond the limit of large blocks as just inside that limit. Thus smaller fragments drawn behind large ones cannot be important in the outer half of the August ballistic field.

Forward air movement attending the whole complex process of ejection has been suggested to aid the flight of ballistics (Wilson, 1980), a concept supported by visual observation of shock waves accom-

panying explosive eruptions (Nairn and Self, 1978; Ishihara, 1985). Fagents and Wilson (1993) included this effect in their ballistic analysis by considering that the blocks were ejected within an envelope of air whose velocity (initially the same as that of the blocks) decayed exponentially with time.

DISCUSSION

Convective uplift helps explain many small *apparent* ballistics lying generally outside the continuous-airfall field as clasts that dropped from a funnel-shaped eruption column as shown by Carey and Sparks (1986, figs. 3, 13). Several researchers think that gas streaming from the vent could account for small particles thrown to the same or even greater range as large particles (Lorenz, 1970; Self and others, 1980; Mastin, 1991). Fagents and Wilson (1993) offer a numerical modification of Wilson (1972) that takes into account a gas-expansion phase 25 to 150 m outward from the vent, which reduces but does not eliminate the unreasonably high initial velocities that calculations for smaller particles yield. Further analysis of a gas-expansion phase during the Crater Peak ballistic explosions and of drag coefficients is in progress (L.G. Mastin and R.B. Waitt, unpub. data).

HYDROVOLCANIC NATURE OF EXPLOSIONS

Violent explosions caused by rapid intermixing of molten material and water have been analyzed in industrial and volcanic settings (Colgate and Sigurdsson, 1973; Peckover and others, 1973; Wohletz, 1983; Wohletz and McQueen, 1984). Swift interaction of magma with external water could help explain the transiently high pressure generated at apparently shallow depths near the end of the August eruption of Crater Peak. The August eruption lasted about 4 hours. As magmatic vent pressure waned, ground water from the cone may have poured into the vent along with ejecta fallback and talus from the overhanging northwest crater wall. Mixing of hot rock and water in the vent, contained beneath collapsing crater walls and slumping talus, could fuel hydroexplosions. Most of the ballistic blocks from Crater Peak are angular, blocky lithic clasts and cauliflower bombs (mostly fragmented cauliflower bombs). Vesicularity among juvenile clasts is only low to moderate (generally <50 percent), and the range in vesicularity is high. These characteristics resemble hydromagmatic deposits elsewhere and contrast with strombolian deposits having higher vesicularity and fluidal external shapes (Lorenz and Büchel, 1980; Self and others, 1980; Lorenz and Zimanowski, 1984; Houghton and Hackett, 1984;

Houghton and Schmincke, 1989; Houghton and Nairn, 1991). Thus the explosions late in the August eruption appear to have been propelled by hydromagmatic processes rather than by vesiculation of magma.

Eruptions of some stratovolcanoes have evolved from magmatic to hydromagmatic as an eruption wanes. As the Vesuvian A.D. 79 magmatic eruption dwindled, interaction of magma with external water produced the final surge beds (Sheridan and others, 1981; Sigurdsson and others, 1985). Increased interaction with water also occurred near the end of many other Vesuvian eruptions from Pleistocene to recent (Sheridan and others, 1981; Rosi and Santacroce, 1983; Rosi and others, 1993; Bertagini and others, 1991).

HORIZONTALLY CURVING TRAJECTORIES

Many ballistic projectiles at Crater Peak are embedded in the downrange side of the impact crater, or they shattered and the fragments spread downrange from the crater, which indicate the projectiles arrived at distinct angles to the vertical. Most impact craters are elongate and have an asymmetric ejecta rim. The largest craters have ejecta rays radiating on one side. Both phenomena register the azimuth of the block's impact.

In the accessible right-marginal part of the August 1992 ballistic field investigated, some impacts point directly away from the Crater Peak vent on an azimuth of 135°, yet most of them along this general azimuth point more southward than in a straight line from the vent (fig. 16). A typical southward deviation is 20° to 40° south, the maximum 67°. In this downrange right side of the ballistic field, the deviation is systematically southward; no trajectory has the opposite displacement from a straight line. The southward tendency at the end of the ballistic trajectories was *despite* the ambient wind toward the east-southeast (toward about azimuth 115°), which tended to straighten out (that is, shift eastward) any south-curving trajectories. The angular deviation from a straight-line trajectory seems to be largely independent of a projectile's diameter.

A pitched baseball or softball will always curve in the direction of a sharply applied spin. This phenomenon, known as the "Magnus effect," has been variously analyzed by Briggs (1959) and Adair (1990), among others. The systematic map-view pattern at Crater Peak—of which we have found no other example in the literature on volcanic ballistics—suggests that many of these clasts had a notable angular velocity about strongly vertical axes. The deviant clasts that curved systematically rightward (viewed from above), if caused by spinning, would have spun clockwise.

In several movies and videos showing typical low-energy strombolian-type explosions, the clasts do not spin much, certainly less than 1 rev/s. If spinning is the cause of deviations in the horizontal component of trajectory of the Crater Peak ballistics, it must be a phenomenon induced by high-energy ejection.

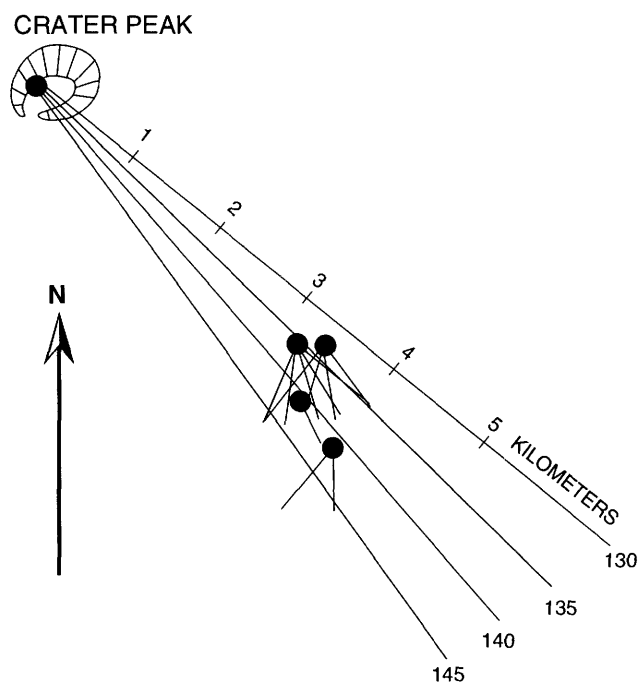


Figure 16. Azimuths of projectiles arriving at right margin bomb field of August 18, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska.

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Hybrid Wet Flows Formed by Hot Pyroclasts Interacting with Snow During the 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

By Richard B. Waitt

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ABSTRACT

During each of the three 1992 eruptions of Crater Peak at Mount Spurr volcano in Alaska, hot pyroclasts interacted with snowpack to form mixed snow-pyroclast flows. Some flows were mostly of snow that melted slightly to be mobile but still cold; others were mostly of pyroclasts whose snow component melted entirely to make the flows more mobile and cooler than ordinary pyroclastic flows. Some pyroclastic flows were large enough to substantially melt and erode glacier ice to generate a lahar that greatly enlarged a moraine breach. The largest wet flows coalesced, swept down and off the volcano flanks, and transformed into lahars. On the volcano flanks were a wide range of hybrid flows between conventional pyroclastic flows and conventional lahars. The hybrid flows of mixed debris from Crater Peak resemble some formed by swift pyroclast-snow interactions at Mount St. Helens, Washington, during explosive events between 1980 and 1991; at Nevado del Ruiz, Colombia, in 1985; at Augustine Volcano, Alaska, in 1976 and 1986; and at Redoubt Volcano, Alaska, in 1989–90. Wet pyroclastic flows are hazardous insofar as their reduced internal friction projects destructive flows down valleys beyond the reach of dry pyroclastic currents. The trajectory of pyroclasts launched over the crater rim by a partly obstructed, inclined vent produced flows on the south-east and east flanks in addition to those along a southward route through the low point of the crater rim.

INTRODUCTION

Crater Peak at Mount Spurr volcano, Alaska (fig. 1), erupted three times in 1992: June 27, August 18, and September 16–17. These eruptions, each lasting 3 to 4 hours, were the first activity since a single eruption on 9 July 1953 (Juhle and Coulter, 1955). During each of the 1992 eruptions, hot pyroclasts turbulently interacted with snowpack to form mixed snow-pyroclast flows.

Pyroclastic flows and surges on the one hand, and lahars on the other, are commonly treated as distinct phenomena. Hybrid flows lie somewhere in between. This report describes the deposits and the erosion of glaciers by hybrid pyroclast-snow flows on the flanks of Crater Peak, gives photographic documentation, qualitatively discusses pyroclast-snow interaction processes, and analyzes flow-initiation mechanisms. Of the 1992 Crater Peak eruptions, chronology and overview are addressed by Eichelberger and others (this volume), pyroclastic flows by Miller and others (this volume), tephra deposits by Neal and others (this volume), and sizable lahars on the lower volcano flanks by Meyer and Trabant (this volume).

The behavior of a hybrid mixed flow depends partly on the proportion of components. Two end members of mixed flows are: (1) mostly snow that only partly melts but including a minor pyroclastic component; and (2) mostly pyroclasts but including a minor proportion of snow that entirely melts. The pyroclast component can vary between lithic and pumiceous end members. A pyroclastic flow that incorporates only a small proportion of snow will behave more or less as a conventional hot and dry pyroclastic flow but be somewhat cooler. A small pyroclastic flow that triggers a huge snow avalanche will behave more or less as a snow avalanche—albeit a wet one because of some melting. Intermediate pyroclast-snow mixtures form a wide variety of fluid flows ranging from wet pyroclastic flows, to slushflows, to debris flows, to floods.

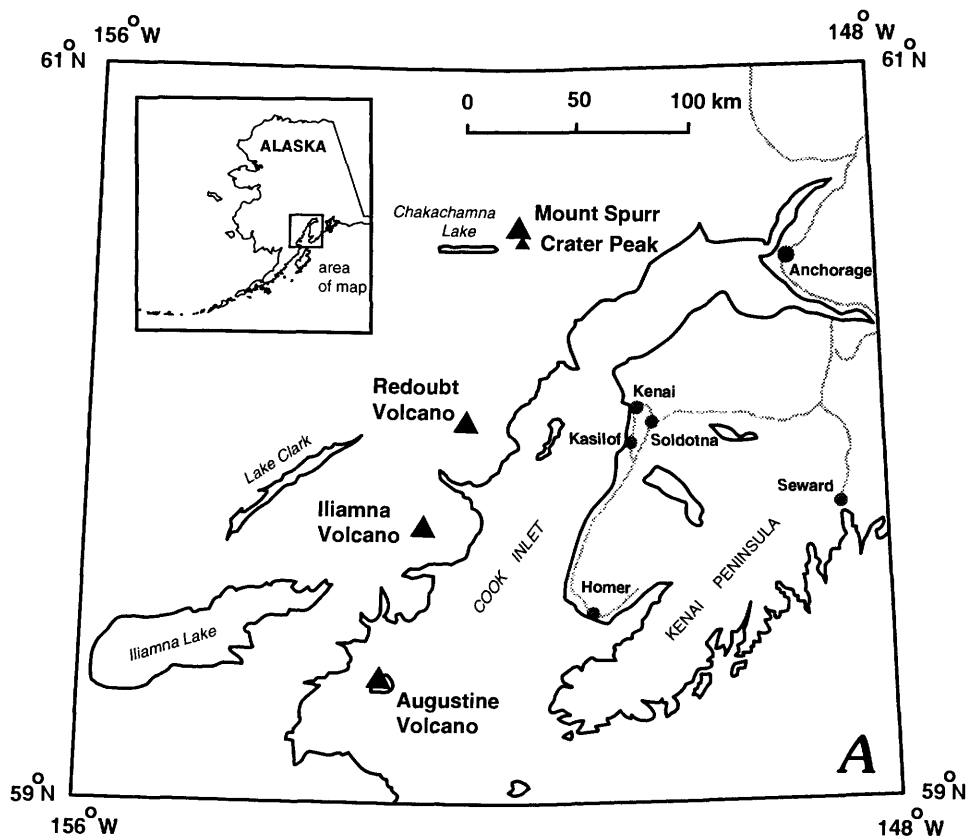
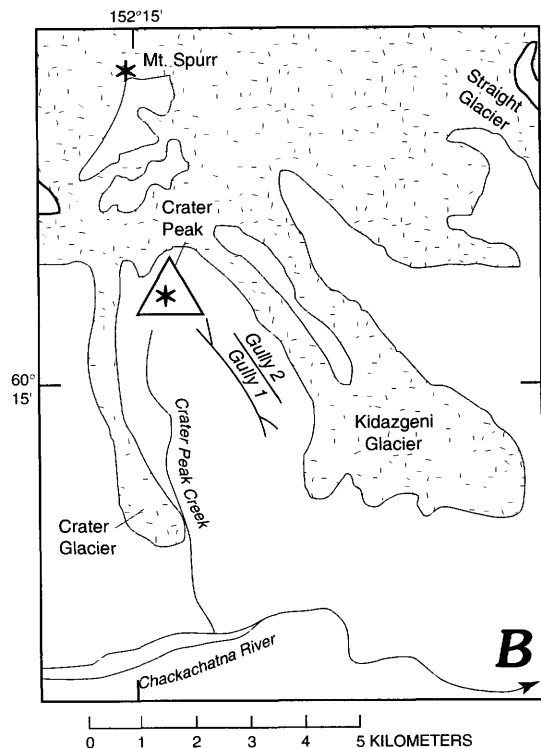


Figure 1. Index maps of (A, above) Cook Inlet area and (B, below left) Mount Spurr and Crater Peak area, Alaska (stipple pattern, glacier).



JUNE ERUPTION

Before the June 27 eruption, snowpack discontinuously clad Crater Peak from the crater rim (2,135–1,950 m) down to 1,370 m, thickest in gullies. Hot pyroclastic debris mixed with this snow during the June eruption to form mixed snow, ice, and rock avalanches that flowed south down creeks between Kidazgeni and Crater Peak Glaciers. These flows entered the Chakachamna River 6 km from the crater in two places. Most of the debris flowed down Crater Peak Creek drainage, the course followed by a lahar that briefly dammed the river during the 1953 eruption.

The largest flow exited the south-facing crater breach and followed a sinuous deep gully to Crater Glacier and Crater Peak Creek (fig. 2A,B) (Meyer and Trabant, this volume). Gully 1, which descends the south-southeast flank of Crater Peak, does not head at the south-facing breach in the crater rim (1,950 m). A snowy debris flow of volume of $1 \times 10^5 \text{ m}^3$ or less descended gully 1, broadly overflowed its sides (fig. 3), and fingered out; the flow in the main channel continued down below altitude 740 m. At a sharp bend

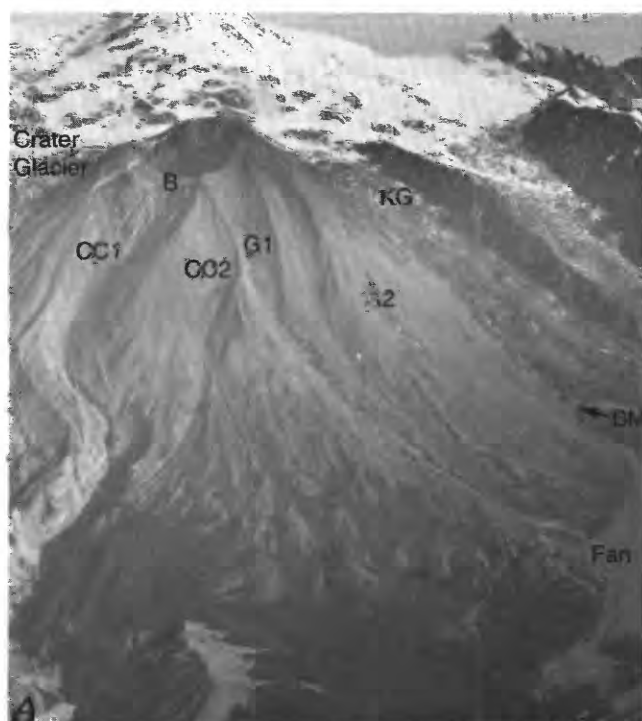


Figure 3. Oblique view northwestward of gully 1 showing distribution of mixed-flow and pyroclastic-flow deposits of 1992 Crater Peak eruptions, Mount Spurr volcano, Alaska. J, broadly distributed June pyroclast-snowflow deposit; A, August pyroclastic flows; S, confined September wet pyroclastic-lahar flows; sr, site of spillover of June and runoff of August flows where velocity was measured. Photograph taken September 26, 1992.

Figure 2. Oblique aerial views north showing Crater Peak, Alaska, on south flank of Mount Spurr volcano, shown at top of both photographs. Between Mount Spurr and Crater Peak is the common head of Crater Glacier and Kidazgeni Glacier. Crater Peak Creek heads at the breach (B) in the crater rim; the southeast crater rim forms a barrier between the crater and the two heads of gully 1 (G1) and especially gully 2 (G2). Arrows show paths of main pyroclast-snow flows of June 27, 1992. CC1 and CC2 show paths of the main and a smaller flow that combined to form main lahar down Crater Peak Creek (see Meyer and Trabant, this volume). BM (photograph A) denotes breached moraine where flood of September eruption on Kidazgeni Glacier entrained much debris, deposited on fan below. A, Photograph taken July 27, 1993. B, Photograph taken September 26, 1992; steam plume shows position of vent against west crater wall.

in the creek at about altitude 790 m, the flow ran as high as 8 m up a steep wall nearly transverse to flow. The equation $v = (2gh)^{1/2}$ (v , velocity; g , gravitational acceleration; h , runup height) relating kinetic energy converted to potential energy at a runup gives a velocity of 13 m/s or 47 km/hr.

At summer's end, the flow deposit was a loose, unsorted diamict with a sandy matrix that carried lithic clasts (no pumice) as large as 0.5 m. The deposit had an intricately rough surface texture in places including small (<20 cm deep) kettles (fig. 4) and an outer margin only 10 to 30 cm high. Scrub willow abraded by the flow was unscorched and uncharred, and by late summer perennials such as fireweed had sprouted up through the deposit where it was thinner than 30 cm—both of which suggest that this flow had not been

hot. A high proportion of internal voids formed during summer by melting of a substantial snow component caused the deposit to suddenly deflate when jarred underfoot. The deflatable character and the rough, kettled surface of this deposit was like that of mixed lithic pyroclast-snow flows emplaced during several eruptions at Mount St. Helens between 1981 and 1986 (Waitt and MacLeod, 1987). After a winter season of compaction beneath snow, a spring season of wetting, and a summer season of episodic high winds, the deposit had become fully deflated and its surface smoother. The deposit shows no evidence of having been channeled by free surface water during emplacement, except along the gully axis where a wetter phase carved a broad channel (fig. 5) and transformed into a small debris flow downchannel (Meyer and Trabant, this volume).



Figure 4. View in gully 1 at altitude 770 m of unsorted mixed-flow deposit from June 27, 1992, eruption of Crater Peak, Mount Spurr volcano, Alaska. Mixed-flow deposit in foreground is deflated by melting of its former snow constituent. It is overlapped by August pyroclastic-flow deposit (just behind person). Photograph taken September 26, 1992.



Figure 5. View northwest in gully 1 on Crater Peak, Mount Spurr volcano, Alaska, showing channel (C) cut through snow-flow deposits (J) of June eruption by trailing watery phase of same flow, then partly filled by pumiceous pyroclastic flow of August eruption (A). Photograph taken July 27, 1993.

The trajectory of the Crater Peak eruption debris not only southward through the low point on the crater rim but also broadly over the higher south-south-east rim was influenced by the position of the vent against the northwest crater wall (fig. 2A,B). This vent position directed explosions southeastward; thus voluminous hot fragmental material lofted over the crater rim and onto the outer flank of the cone where it swiftly and turbulently mixed with snowpack. The ejected debris was mainly lithic, apparently derived from the vent walls and crater wall. Southeastward ejection occurred despite wind that carried the eruption's tephra plume north (Neal and others, this volume).



Figure 6. Steep, lobate deposits of hot pyroclastic flows that descended gully 2 during August 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska. Concentrated pumice blocks are at flow margin, beneath which vegetation is charred. Width of lobe in foreground is about 3 m.

AUGUST ERUPTION

The August 18 eruption sent small-volume pyroclastic flows of breadcrusted dark andesite pumice down the east and southeast flanks of Crater Peak (Miller and others, this volume). The deposits at the surface are coarse, clast supported, and fines poor. In gully 1, levees of angular pumice mark the extents of successive flow lobes. They rode up 7 m on a steep nearly transverse side ($v = (2gh)^{1/2} = 12 \text{ m/s}$ or 43 km/hr), and they overflowed the gully sides in a few places but far less than did the June mixed flow (fig. 5). Steep-fronted deposits of the flow terminate in the main channel at about altitude 745 m (fig. 3), well upvalley of the June mixed-flow terminus.

Several pumiceous pyroclastic flows swept the east-southeast flank of Crater Peak, down to 885-m altitude in gully 2, where they left typical steep, fines-poor, lobate margins 1 to 2 m high (fig. 6). The deposits remained hot for days, and blocks at the edges of these flows charred vegetation beneath them.

One of these pyroclastic flows in gully 2 mixed with snow high on the cone to form a flow much more mobile but compositionally identical (pumiceous) to one that descended the west side of the gully down to altitude 780 m (fig. 7). In contrast to typical pyroclastic-flow deposits upslope, this deposit has gently sloping margins only 20 to 50 cm high. The flow did not char vegetation that it had abraded; pumiceous blocks at the flow edge did not scorch underlying tundra plants. Small boulders jammed up behind large ones (fig. 8) much more than in the steep-margined flows—evidence of a crude sorting process that trapped coarse clasts and evidence that flow continued after depositing jams. A water phase must have caused the higher mobility. Traced upgully to altitude 885 m, this mobile-flow deposit merges with a steep-fronted typical pyroclastic-flow lobe on the east side of gully 2. Thus the part of the flow in the deeper part of the gully detached and flowed much farther. This relation is evidence that a single pyroclastic flow divided because part of it mixed with snow in the gully axis, became wet and lost some internal friction, and thus was able to flow farther even on gentle slopes.

These flows descended gully 2 despite that the gully heads on the high southeast segment of the crater rim far from the south-facing crater breach (fig. 2). Downslope from the even higher segment of the crater rim, pyroclastic flows also mixed with snow to form a small watery lahar that flowed down the west side of Kidazgeni Glacier, through a moraine breach, and to Chakachatna valley (Meyer and Trabant, this volume). The pyroclastic debris did not spill through the southward gap in the crater rim but was explosively launched east and southeast with enough vol-



Figure 7. View upvalley of mobile-flow deposit of the August 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska, that greatly outdistanced normal pyroclastic flows in gully 2. Margins are low, clasts are sorted (water winnowed), and vegetation is abraded but not charred. Person (to right) is standing in ballistic impact crater. Boulder jam shown in figure 8 marked by arrow. Compare with figure 6.



Figure 8. View downflow of jam of cobbles and small sorted boulders on upstream side of 1.5-m boulder in August 1992 mixed-flow deposit in gully 2. Handle of shovel (center of photograph) is about 50 cm long. Crater Peak, Mount Spurr volcano, Alaska.

ume to erode and melt snow that lingered even in late summer. Ballistic projectiles—lithic clasts as large as 1.5 m and smaller juvenile projectiles—were hurled at least 4 km southeast of Crater Peak (fig. 7) (Waitt and others, this volume). This ballistic shower, which mostly occurred after the pyroclastic flows were emplaced on the same flank, is further evidence that the vent explosively directed debris southeast over the crater rim.

SEPTEMBER ERUPTION

The eruption during the night of September 16–17 sent at least two wet pyroclastic (laharic) flows down the south-southeast and southeast flanks of Crater Peak. In gully 1 a fluid flow with gentle termini fingered out 300 m upgully of the August pyroclastic-flow termini. A second deposit in gully 1 has a sharp steep margin 0.5 to 1.5 m high (fig. 9) of several

lobes whose crosscutting patterns delineate slight differences in arrival times. That second flow apparently was far more viscous than the slightly earlier flow, and it terminated abruptly about 50 m upvalley of the first wet-flow terminus. The September deposits consist largely of breadcrusted to angular clasts of juvenile andesite pumice superficially resembling pyroclastic-flow deposits. Yet 12 to 15 hours after the eruption both were cool and water saturated (R.M. McGimsey, oral commun., 1992). The younger of the two has arcuate pressure ridges indicating a flow with internal strength (fig. 9). Rather than remaining loose as in typical pyroclastic flows, the matrix hardened within days as the wet clay and silt particles adhered to each other and bound the coarse clasts. Unlike the largely pumiceous August flows, only two-thirds of gravel clasts of the September wet flows are juvenile; the rest are diverse porphyritic andesite clasts of older Crater Peak rocks incorporated upgully.

Pumice-lithic pyroclastic flows also descended the east-northeast side of Crater Peak cone and turned southeast down Kidazgeni Glacier. An eyewitness situated 18 km to the southeast reported the heaviest incandescent flows of the night eruption on the east flank. At altitude 1,000 m their collective deposit is a poorly sorted pebble gravel that consists mostly of dark brownish gray andesite and lithic clasts. Many ice blocks as large as 1 m were entrained as the flow descended an ice fall. This deposit broadly overlies gently sloping parts of the glacier between altitudes 1,050 and 800 m (fig. 10B). Along the west side of the glacier from altitude 2,000 down to 1,000 m (gradients 500 to 150 m/km), meltwater stripped off this deposit along

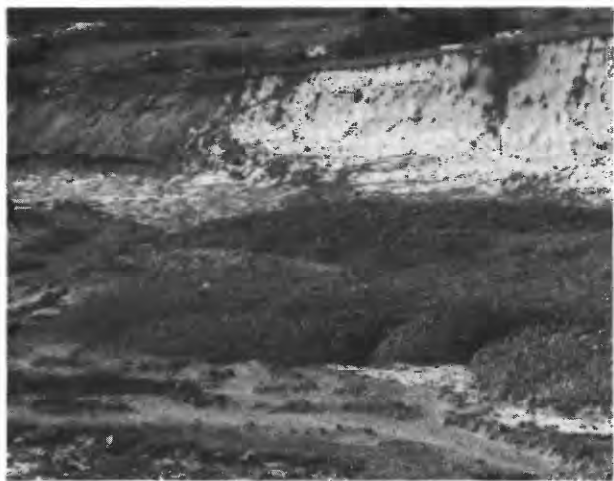


Figure 9. Terminus of lobate laharcic flow of the September 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska in gully 1. Note delicate pressure ridges on flow surface. Person for scale. Photograph taken September 26, 1992.

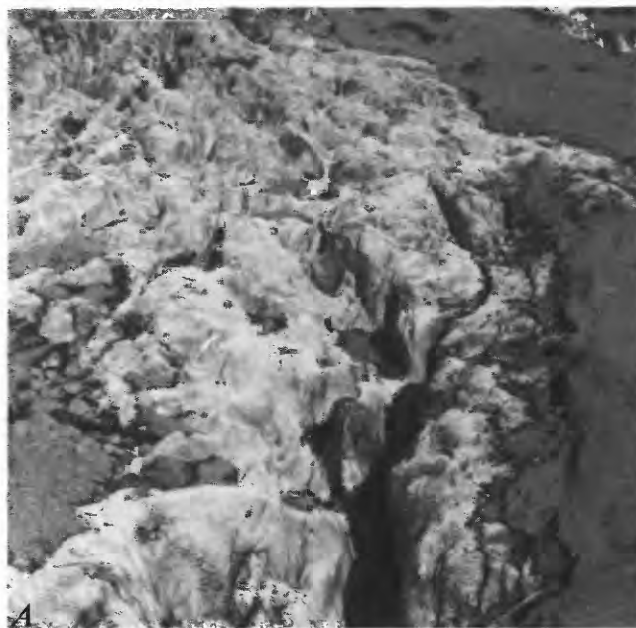


Figure 10. Low-level aerial views (photographs taken September 29, 1992) of ice scabland produced by flood during September 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska, along west side of Kidazgeni Glacier. A, September pyroclastic-flow deposit veneer stripped along September ice-scabland channels; width between scarps in deposit in lower part of photograph is about 5 m.

a tract 5 to 15 m wide while eroding a stepped series of cataracts and plunge pools several meters wide and deep (figs. 10, 11). This essentially subfluvial topography is a small-scale scabland with 0.5- to 1-m scarps along its margins, analogous to the bizarre topography of Washington's channeled scabland, two to four orders of magnitude larger cut in basalt by gigantic floods during the Pleistocene (Bretz, 1928, 1959; Waitt, 1994). Within the ice scabland on Kidazgeni Glacier are pockets of poorly sorted diamict containing lithic boulders as large as 1.5 m indicative of debris flow, and moderately sorted gravel in the form of small expansion bars (also a channeled-scabland analogy) at the lower ends of some eroded ice-scabland channels, indicative of sediment-laden streamflow. The debris flow formed as meltwater mixed with the new pyroclastic-flow deposit; the streamflow ensued as the proportion of meltwater outstripped the available supply of pyroclasts. The watery phase of flow down the glacier closely followed the debris-flow phase.

From the ice-scabland tract the flows emerged onto a flatter (gradient 100 to 50 m/km) part of the glacier and spread out as a broad fan 50 to 75 m broad. Later incision of this deposit exposed more than 1.2 m of compact, nearly massive pebble gravel (diamict) capped by 20 cm of loose openwork pebble



Figure 10. Continued. *B*, View directly up main ice-scabland tract; width of roughest part of tract about 8 m.



Figure 10. Continued. *C*, Erosional ice-scabland and patches of pyroclastic-flow deposit, view upglacier; height of largest scabland elements about 2.5 m.



Figure 11. Close view (photograph taken September 29, 1992) upglacier of ice scabland on Kidazgeni glacier produced by flood during September 17, 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska. Depth of scour shown here is about 2.5 m.

gravel (fig. 12). This two-layer deposit bespeaks a voluminous debris flow phase followed by a voluminous watery phase.

Detailed grain-size analyses are beyond the scope of this study, but differences obvious in the field are quantified by histograms and cumulative-frequency grain-size distributions obtained by sieving representative samples (fig. 13A,B). Mass-flow deposits are poorly sorted, the water-flow deposit moderately sorted (tables 1,2). In Folk's (1980) classification the pyroclastic-flow deposit is medium-pebble gravel, the debris-flow deposit sandy small-pebble gravel (but both are diamicts), and the watery flow deposit a large-pebble gravel (fig. 13C). The large deviation in clast sizes within the pyroclastic-flow deposit bespeaks its polymodal source and lack of winnowing during flow. The even poorer sorting of the debris-flow deposit probably reflects its incorporation of supraglacial de-

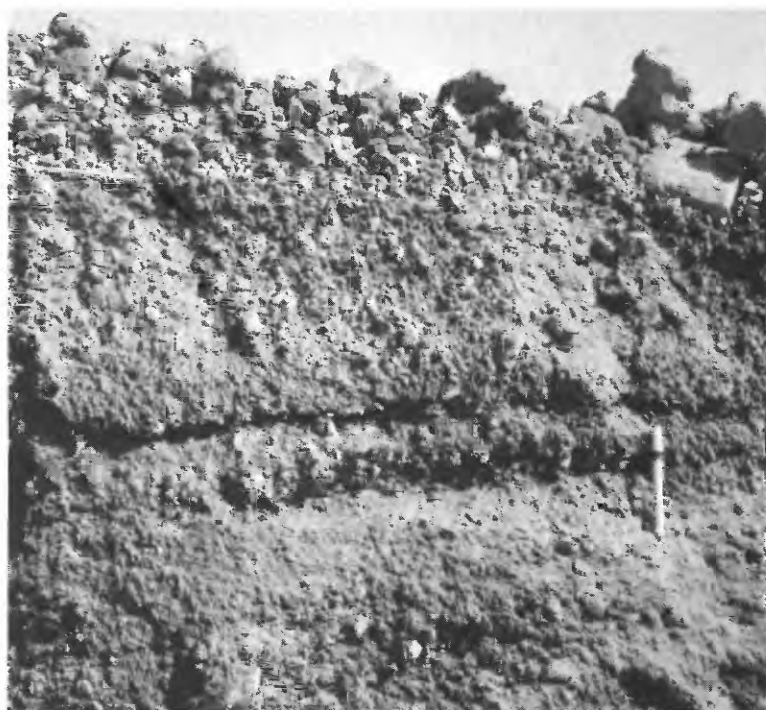


Figure 12. Section exposing September 13, 1992, flowage deposits overlying Kidazgeni Glacier, Crater Peak, Mount Spurr volcano, Alaska. Matrix-bearing debris-flow deposit with poorly developed plane bedding is overlain at pencil (arrow) by openwork pebble-cobble gravel watery-flow deposit. Pen to right is 13 cm long.

bris; the far lower deviation of the waterlaid deposit demonstrates a winnowing out of the fines and perhaps also the coarsest clasts.

The composite debris and watery flow descended to the glacier side and through an existing breach in the right-lateral moraine about at altitude 1,068 m. The flow apparently greatly enlarged the erosional gash, where it incorporated much coarse sediment before spreading gravel across a fan west of the glacier at altitude 825 to 670 (fig. 14). From there a small lahar tongue descended to the Chakachatna River (Meyer and Trabant, this volume).

The September pyroclastic flows and derivative water-containing flows were concentrated on the east-northeast through south-southeast segment of the cone rather than downslope from the south-facing crater breach. This distribution indicates that the pyroclasts were launched by directed explosions, as they were during the June and August eruptions. Over the course of the three eruptions the vent was observed to migrate progressively west, and so after the September eruption it was tight against the west-northwest crater wall (T.P. Miller, oral commun., 1992). With each eruption the migrating vent directed energetic explo-

sions successively farther counterclockwise, from south-southeastward to eastward, obliquely away from the south-facing crater-rim breach.

Table 1. Sieve analyses for three samples of flows of the September 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska.

[Values are weight percent of entire sieved sample.]

Size class in phi (φ)	Pyroclastic flow	Debris flow	Watery flow
-6	11.5		
-5	5.8	5.3	34.5
-4	11.8	10.9	51.0
-3	22.1	25.3	12.7
-2	17.5	20.7	0.1
-1	16.7	12.9	0.0
0	10.9	8.7	0.0
1	2.7	5.1	0.0
2	0.7	7.3	0.1
3	0.3	3.3	0.2
4	0.1	0.1	0.3
>4	0.0	0.4	1.2
Total -----	100.0	100.0	100.0

Table 2. Statistical data from sieved samples of flows of the September 1992 eruption of Crater Peak, Mount Spurr volcano, Alaska.

[Parameters of Folk, 1980]

Parameter	Pyroclastic flow	Debris flow	Watery flow
Weight percent gravel -----	85.4	75.1	98.3
Weight percent sand -----	14.6	24.5	0.6
Weight percent mud -----	0	0.4	1.1
Mean (ϕ) -----	-3.1	-2.2	-4.8
	<i>medium pebble</i>	<i>small pebble</i>	<i>large pebble</i>
Sorting (ϕ) -----	1.8	2.1	0.8
	<i>poor</i>	<i>very poor</i>	<i>moderate</i>
Skewness -----	0.02	0.29	0.0
	<i>symmetrical</i>	<i>fine skewed</i>	<i>symmetrical</i>

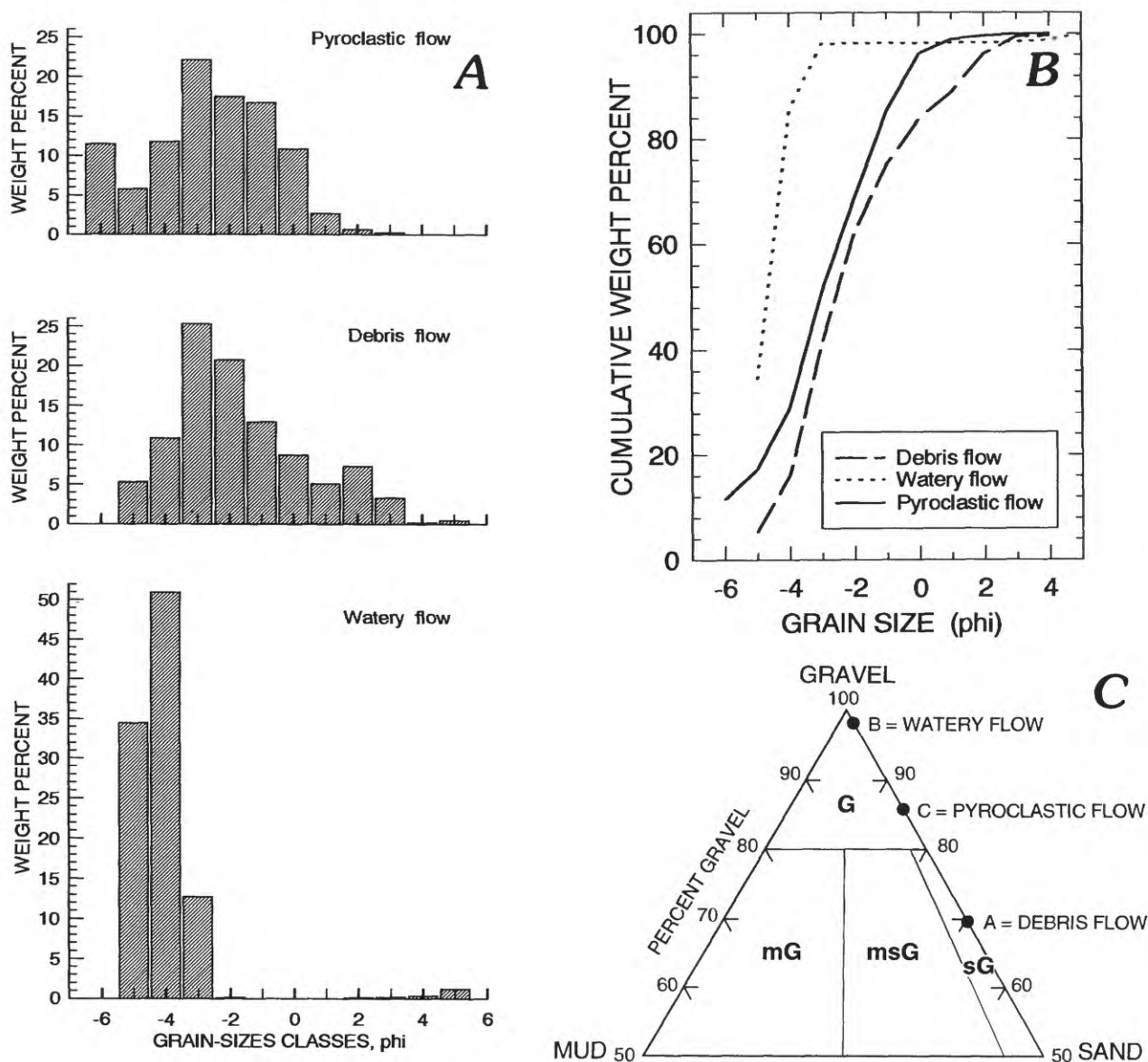
**Figure 13.** Grain-size character of pyroclastic-flow deposit of August 18, 1992, eruption on Kidargeni glacier and deposits derived from it by water-flow processes. Crater Peak, Mount Spurr volcano, Alaska. A, Histograms showing grain-size distributions. B, Cumulative grain-size-frequency distributions. C, Names according to Folk's (1980) classification. Only top (gravel) portion of Folk's classification triangle shown. G, gravel; sG, sandy gravel; msG, muddy sandy gravel; mG, muddy gravel.



Figure 14. View north showing fan of flood debris below the right-lateral moraine of Kidazgeni Glacier and other effects of September 16–17, 1992, eruption of Crater Peak, Mount Spurr volcano, Alaska. Mount Spurr is high, snow-capped peak to upper left; Crater Peak is peak on far left. G1, gully 1; G2, gully 2; P, pyroclastic-flow deposit on glacier; S, water-scoured west side of glacier. B, breach enlarged by flood through lateral moraine. Debris fan resulted from distribution of material from moraine breach. At the mouth of gully 1, the September fan cuts across small fan of June flow that descended gully 1.

COMPARISON WITH ERUPTIONS AT OTHER SNOWCLAD VOLCANOES

The deposits of hybrid flows from Crater Peak resemble those of mixed pyroclast-snow flows during several other recent eruptions. Swift pyroclast-snow mixing occurred at the base of the huge pyroclastic surge off Mount St. Helens during the first minutes of the cataclysmic eruption of May 18, 1980. Stratigraphy shows that large laharcic flows down the east and west volcano flanks trailed the pyroclastic surge and were clearly caused by it (Waite, 1989). Between 1981 and 1991 many small to moderate explosive eruptions at Mount St. Helens caused turbulent, rapid mixing of hot pyroclasts with snowpack on the crater walls and floor to form various mixed flows (Waite and others, 1983; Waite and MacLeod, 1987; Mellors and others, 1988). Each mixed-flow deposit included a relatively dry marginal facies consisting largely of snow

that did not melt and which froze soon after emplacement. Many pyroclast-snow mixing events also produced far more mobile slushflows along deeper and more axial parts of flowpath. Further melting transformed some of them into watery flows that deeply channeled snowpack along the lowest axial part of the flowpath.

During the eruption of Nevado del Ruiz, Colombia, in November 1985, initial pyroclastic surges dynamically mixed with snowpack, firn, and ice to form unusual, mixed ice-debris deposits (Pierson and others 1990, figs. 8, 9). Ensuing pyroclastic flows did far more turbulent mixing and melting to form lahars that flowed down surrounding valleys, where they enlarged with debris and streamwater into huge lethal flows.

At Augustine Volcano, Alaska, pyroclastic flows of two recent eruptions turbulently mixed pyroclasts with snowpack to yield watery flows. During the 1976 winter eruption, pumiceous pyroclastic flows incorporated winter snowpack to form pumiceous deposits more sorted and with more rounded clasts than adjacent pyroclastic-flow deposits (Kamata and others, 1991, fig. 9). The 1986 spring eruption produced typical pumiceous pyroclastic flows that merge downslope into small scabland erosional tracts overlain by patchy, matrix-free water-washed coarse gravel (unpub. data, 1988), evidence that pyroclastic flows swiftly melted snowpack they overrode.

During the 1989–90 eruptions of Redoubt Volcano, Alaska, repeated collapses of fragmented hot dome rock onto glacier ice formed debris flows and watery floods. The initial eruption of December 15, 1989, formed an expansive “dry” ice-diamict deposit on both the north and south volcano flanks (Waite and others, 1994). The degree of melting was small because pyroclasts were swiftly cooled by a huge volume of readily incorporated snow and firn. But during many later dome collapses, fragmented dome rock swiftly interacted with and greatly eroded the northside Drift glacier (Trabant and others, 1993) to yield large watery floods.

At Crater Peak the wet, mixed flows were fairly small. The dry-appearing hybrid-flow deposit of June 27 closely resembles the surface texture of the “dry” phases of mixed-flow deposits at Mount St. Helens between 1982 and 1986 and also the melted ice-diamict surface at Redoubt (Waite and others, 1994, figs. 11, 12). The small scabland eroded into Kidazgeni Glacier in September was similar to the ice-erosional forms carved in Drift glacier by the 1989–90 Redoubt eruptions (Trabant and others, 1994).

Hybrid flows are probably far more common than past reports suggest. Those from Crater Peak and the other cited eruptions were documented because of ac-

cess by helicopter soon after eruption. Because the flows are cool or cold, snowfall quickly obscures them in winter. Because some of the flows contain much snow, much of the original volume disappears and the resulting deposit can resemble a typical pyroclastic-flow or lahar deposit. The subtle evidence of wetness can be mostly lost in the geologic record, indeed easily overlooked even in recent surface deposits.

HAZARDS

Had the wet or snowy flows at Crater Peak been larger they might have dammed Chakachatna River and induced large floods, but they would not have endangered life or property in that remote setting. Yet these small hybrid flows illustrate a hazard common to snowclad volcanoes. When hot pyroclasts turbulently mix with snow or flow over glacier ice, they swiftly generate wet flows that can soon travel beyond the limits of dry, hot pyroclastic flows. Thus their damaging effects extend beyond the mountain flanks. At volcanoes in Indonesia and the Philippines that are affected by monsoons, most lahars initiate secondarily when torrential rains act on loose, hot, unstable pyroclastic debris (Rodolfo and Arguden, 1991; Pierson and others, 1992). The type of snowflows and floods produced at Mount Spurr volcano, Mount St. Helens, Redoubt Volcano, and other snowclad volcanoes, are *primary*, produced as they are swiftly *during* a pyroclastic eruption by turbulent hot flows moving over snow or ice. Large, lethal examples of such floods originated at Öræfajökull in Iceland in 1362 (Thorarinsson, 1958), at Cotopaxi in Ecuador in 1877 (Wolf, 1878), and at Nevado del Ruiz in Colombia in 1985 (Pierson and others, 1990). In a literature review, Major and Newhall (1989) compiled data on historical eruptions at snowclad volcanoes, concluding that pyroclastic flows and surges acting on ice and snow are the most important processes generating *primary* large lahars and floods.

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Whole-rock Major- and Trace-element Chemistry of 1992 Ejecta from Crater Peak, Mount Spurr Volcano, Alaska

By Christopher J. Nye, Michelle L. Harbin, Thomas P. Miller, Samuel E. Swanson, and Christina A. Neal

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ABSTRACT

Juvenile ejecta from the three 1992 eruptions of the Crater Peak vent of Mount Spurr are calcalkaline andesite and contain about 57 percent SiO_2 by weight. Samples are chemically uniform within analytical error for all but a few of the most mobile of the 43 elements whose concentrations were determined. The andesite is unlike prehistoric Crater Peak andesites of similar silica content in having high concentrations of Al_2O_3 , Na_2O , and Sr and low concentrations of compatible transition metals. These differences suggest that the immediate parent to the 1992 andesite crystallized relatively more ferromagnesian silicate (probably clinopyroxene) and less plagioclase than previous Mount Spurr andesites, which requires either higher water pressure, greater residence depth, or both. The 1992 andesite is also unlike previous Crater Peak andesites, including that erupted in 1953, in having lower concentrations of some incompatible elements yet higher amounts of silica. These observations suggest that the magma that fed the 1992 eruptions was not derived from the same magma body that fed the 1953 eruption, but it instead is new magma that came from the source region in the deepest crust or mantle. The

1992 eruptions also ejected partially melted aluminous metamorphic xenoliths characterized by high concentrations of both incompatible and compatible trace elements despite moderate silica contents and by mineral assemblages unlike exposed country rock. Small quantities of highly siliceous xenoliths characterized by dramatic depletion of all but a few trace elements (Rb, Ta, U) were also erupted.

INTRODUCTION

Mount Spurr volcano, 125 km west of Anchorage, is the next-to-last Holocene volcano at the northeastern end of the Aleutian arc. It is about 100 km above the Benioff zone and 500 km from the trench in a direction parallel to the movement between the Pacific and North American plates. The region surrounding Mount Spurr is underlain by a Jurassic to mid-Tertiary granitic batholith, which intrudes Paleozoic to Mesozoic limestone, basalt, and flysch of generally low metamorphic grade (Wilson and others, 1985).

Spurr volcano is composed of an ancestral Mount Spurr (fig. 1) whose construction began at least 255 ka (Nye and Turner, 1990). It is predominantly built of 58 to 60 percent SiO_2 two-pyroxene andesite flows that are interbedded with a few more mafic flows. Lava flows dominate the upper part of the cone, and pyroclastic rocks dominate the lower part. The growth of the ancestral Mount Spurr was terminated by Bezymianny-type avalanche caldera formation and the production of a debris avalanche with a minimum runout of 25 km and a volume of a few cubic kilometers (Nye and Turner, 1990). The resultant amphitheater is 5 km by 6 km, predominantly ice filled, and breached to the south. The debris avalanche is immediately overlain by partially welded ashflows more silicic (61–63 percent SiO_2) than ancestral Mount Spurr andesites. Intrusion of this silicic magma high into the cone was presumably the cause of stratocone failure. The age of caldera formation is not precisely known. It post-

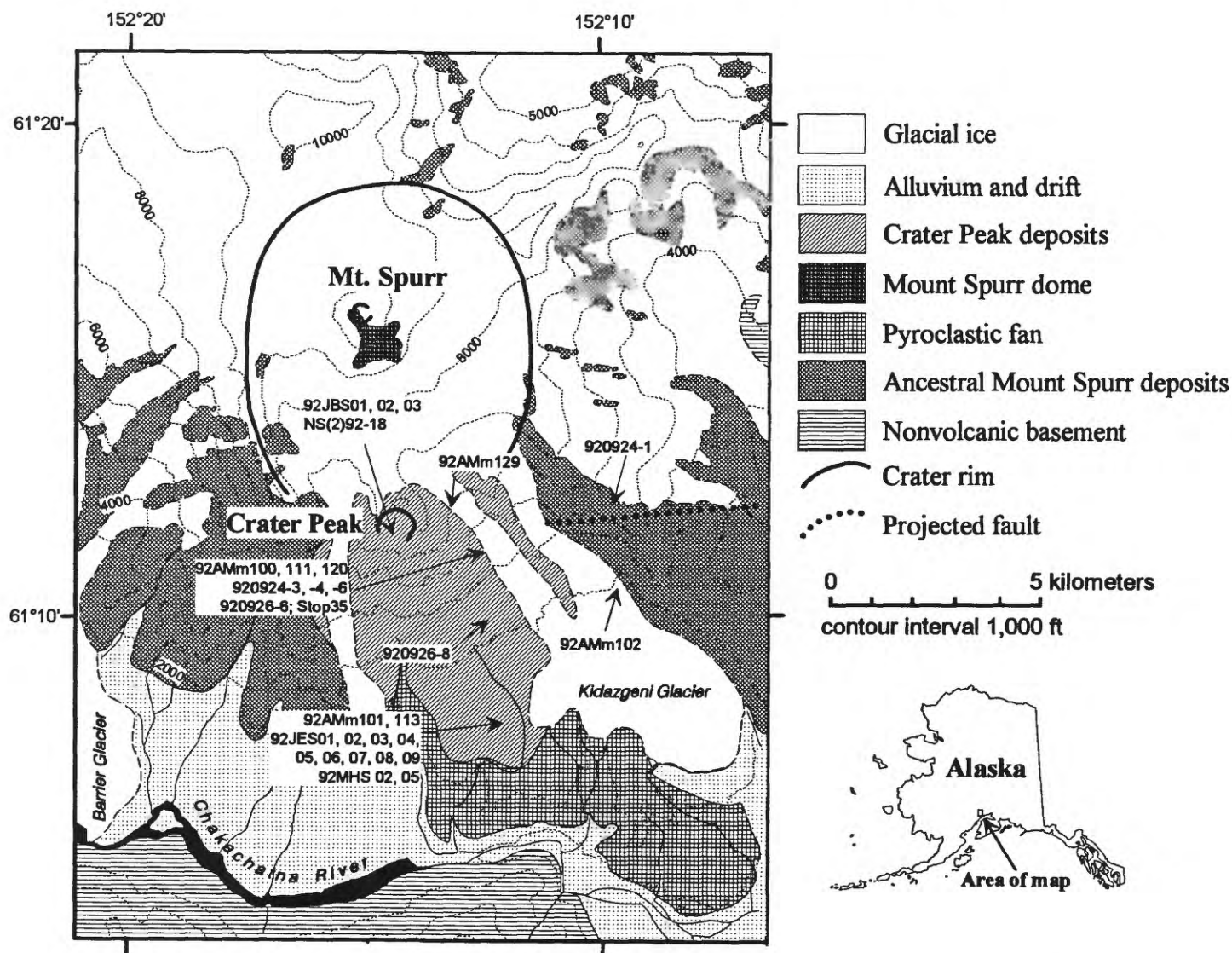


Figure 1. Geology (modified Nye and Turner, 1990) and geography of the area immediately surrounding Mount Spurr, Alaska. Locations of samples analyzed for this study are also shown.

dates the youngest cone-building andesite (youngest measured age is 59 ka; Nye and Turner, 1990) and predates the oldest post-caldera tephra (>7 ka; Riehle, 1985). The lack of a tephra blanket suggests that caldera formation occurred during a time of widespread ice cover, ergo a Pleistocene age. The lack of erosion of the debris avalanche, despite its location in major glacial valley, suggests a late Pleistocene age of caldera formation.

After caldera formation, a large dome formed in the center of the caldera (fig. 1). It forms the current summit of Mount Spurr and is ice covered except for periodically exposed patches of dome lava with scattered, diffuse boiling-point fumaroles. The few samples recovered from the dome have 61 to 63 weight percent SiO_2 , plagioclase, two pyroxenes, and occasional hornblende xenocrysts and have composi-

tions similar to the ashflows overlying the debris avalanche (Nye and Turner, 1990). Tephra from this vent do not appear to be younger than about 5 ka (Riehle, 1985). Crater Peak, 6 km south of the Mount Spurr summit (fig. 1), sits in the caldera breach and is the historically active vent. It consists of lava and pyroclastic flows with 54 to 57 percent SiO_2 —more mafic than cone-building andesites, which contain plagioclase, two pyroxenes, and sparse hornblende phenocrysts. Tephra from Crater Peak range in age from 5 ka to the present (Riehle, 1985). However, the presence of block-and-ash flows produced by the eruption of magmas that are apparent mixtures of Crater Peak and Spurr summit dome magmas as well as hornblende xenocrysts derived from Crater Peak magmas in Spurr summit dome magmas, suggest that the two magmatic systems overlapped in time (Nye and Turner,

1990). Remnants of an older cone in the position of the current Crater Peak lie just west and east of Crater Peak. These are referred to as ancestral Crater Peak here, but they were called proto-Crater Peak by Nye and Turner (1990).

The Crater Peak vent of Mount Spurr erupted on June 27, August 18, and September 16–17, 1992 (Eichelberger and others, this volume). Eruptions were 3.5 to 4 hours in duration, and they produced subplinian columns which deposited extensive tephra blankets and minor proximal pyroclastic flows and other debris flows. These eruptions produced six types of ejecta. In order of decreasing volume, they are: (1) widespread blankets of mostly juvenile andesitic tephra with a combined dense-rock-equivalent volume of about $35 \times 10^6 \text{ m}^3$ (Neal and others, this volume); (2) brown to dark-gray, breadcrusted andesite blocks deposited in pyroclastic flows, lahars, and ballistic bomb fields with a total volume of about one percent of that of the tephra (Miller and others, this volume; Waitt and others, this volume); (3) aluminous xenoliths composed of highly inflated, partially melted blocks of metamorphic country rock, with a total volume of about one percent of the volume of the andesite blocks (Harbin and others, this volume); (4) light gray-green andesite, which is distinguished most readily from the dark andesite by clear, rhyolitic groundmass glass rather than the brown, andesitic glass of the dark andesite; (5) volumetrically minor metamorphic xenoliths including a white glass-quartz-plagioclase rock and wolastonite-bearing skarn; and (6) accidental blocks of older Crater Peak andesite. Representative samples of most of these types of ejecta were analyzed in bulk for 10 major and 33 trace elements.

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SAMPLE SUITE AND ANALYTICAL QUALITY

The samples discussed herein include 19 juvenile andesite blocks (three from the June 27 eruption, nine from the August 18 eruption, and seven from the September 16–17 eruption). These blocks span the

known emplacement modes and stages of each eruption. We also analyzed five light gray-green andesites (four from August 18 and one from September 16–17); 20 aluminous xenoliths (all from August 18); and 3 white glass-quartz-plagioclase xenoliths from August 18 deposits. Sample locations for andesite blocks and xenoliths are shown in figure 1. In addition we analyzed a total of 20 proximal tephra samples consisting of two or three samples from light and dark and upper and lower tephra from each eruption. These are described more fully in Neal and others (this volume). Separate splits of each sample were powdered in a tungsten-carbide shatterbox for analysis by X-ray fluorescence (XRF) and in a steel shatterbox for analysis by inductively coupled plasma mass spectrometry (ICP/MS). Crushing in tungsten-carbide avoids compatible-transition-metal contamination and crushing in steel prevents high-field-strength element contamination. Milling and analysis were done at the GeoAnalytical laboratory at Washington State University. One duplicate was chosen at random and analyzed for each batch of samples submitted. These duplicates form the basis for our reported analytical precision (table 1). Overlap of analyzed elements between techniques (high-precision Rb, Ba, Y, and Nb data were obtained by both XRF and ICP-MS) provides another method of monitoring analytical quality.

ANDESITE

HOMOGENEITY OF 1992 ANDESITE

Two types of andesite erupted from Crater Peak in 1992: dense to scoriaceous brown to dark-gray andesite and much less abundant light gray-green andesite. The dark andesite is fairly homogeneous throughout the deposits from all three 1992 eruptions (tables 1, 2). For most elements the percent deviation of all 1992 Spurr andesites is close to the analytical precision of the laboratory. Some of the highly incompatible elements (Cs, Rb, U, Th, and Pb) have variations in 1992 ejecta outside analytical precision, but even for these elements the total variation among samples is small. For example, Cs variability between samples is about 9 percent, but this only represents a 0.04-ppm standard deviation of the analyses. Pb variability is apparently quite high (18 percent), but this reflects the inclusion in the mean of a single sample from the September 17 eruption with 9.54 ppm Pb. Omission of this sample from the calculation results in a relative deviation of only 4 percent. We did not analyze this sample in duplicate to verify its Pb content. Some of the compatible transition metals (Fe, Sc, Cr, and Ni) also show variability outside analytical precision.

Table 1. Whole-rock composition of selected historic and prehistoric andesites erupted from the Crater Peak vent of Mount Spurr, Alaska.

[Sample locations, modes and descriptions for samples not described herein given by in Nye and Turner (1990). Mean for dark 1992 andesite is for 19 juvenile andesite blocks from 1992 eruptions of Crater Peak; std dev., one standard deviation of the analyses in preceding column; % dev. is the standard deviation expressed as a percentage of the mean for each element; precision refers to analytical precision, in percent, of these analyses as indicated by duplicate studies. Major elements normalized to 100 percent anhydrous. FeOt is total iron as FeO. Number of reported decimal places may exceed analytical precision.]

Sample	Prehistoric Crater Peak andesite				1953 andesite		Dark 1992 andesite			Light grey-green 1992 andesite					
	PCP-B01	PCP-A02	CP-02	CP-12	PUM-16	PUM-17	mean	std. dev.	%dev.	precision (2σ)	92AMm 102GS	92AMm 113C	92JES -01	92MHS 10	92AMm 120A
Major-element oxides, in weight percent															
SiO ₂	57.44	56.61	56.61	56.30	54.89	54.35	56.73	0.34	0.6	0.2	56.59	58.84	56.47	56.16	56.52
TiO ₂	0.85	0.79	0.87	0.79	0.86	0.91	0.69	0.01	1.2	1.2	0.67	0.64	0.70	0.76	0.71
Al ₂ O ₃	17.23	17.14	18.00	18.55	18.80	18.61	19.05	0.11	0.6	0.6	19.33	18.23	18.95	18.77	19.02
FeOt	6.91	7.00	7.05	6.99	7.68	7.68	6.95	0.29	4.2	1.6	6.82	6.53	7.10	7.39	7.07
MnO	0.14	0.14	0.14	0.14	0.15	0.14	0.15	0.00	1.7	1.2	0.15	0.14	0.16	0.16	0.15
MgO	5.04	6.33	4.61	4.32	4.58	4.85	3.65	0.07	1.9	1.2	3.52	3.46	3.75	3.93	3.74
CaO	7.54	7.59	7.38	7.38	8.02	8.14	7.57	0.08	1.0	0.8	7.65	6.69	7.74	7.91	7.61
Na ₂ O	3.41	3.25	3.85	4.21	3.81	4.07	4.02	0.07	1.8	1.0	4.08	4.10	3.92	3.77	3.98
K ₂ O	1.24	0.95	1.25	1.05	0.95	0.99	0.91	0.03	2.8	1.4	0.90	1.14	0.93	0.88	0.93
P ₂ O ₅	0.21	0.19	0.25	0.26	0.26	0.26	0.28	0.01	4.8	0.6	0.29	0.23	0.28	0.28	0.28
Trace elements, in parts per million															
Cs	0.17	0.14	0.28	0.22	0.60	0.19	0.45	0.05	10.0	2	0.57	0.41	0.83	0.53	0.51
Rb	26.98	19.79	29.82	21.85	16.70	18.53	15.12	1.20	7.9	2	15.56	18.04	15.44	14.57	18.39
Ba	459.53	453.07	448.69	358.99	385.17	357.34	385.24	8.81	2.3	1	397.00	411.78	387.39	352.53	381.20
Sr	558.00	554.00	612.00	681.00	620.00	699.00	658.84	8.76	1.3	1	671.00	564.00	657.00	629.00	650.00
La	13.82	11.32	14.79	10.01	11.12	12.10	11.05	0.25	2.2	4	11.46	11.28	11.02	10.30	10.76
Ce	28.04	22.56	29.67	21.60	22.89	24.86	22.63	0.55	2.4	4	23.87	22.04	22.59	21.47	21.65
Pr	3.56	2.84	3.74	2.82	3.03	3.26	2.96	0.06	2.1	4	3.15	2.82	2.84	2.88	2.93
Nd	15.53	12.60	16.39	12.71	13.64	14.68	13.24	0.32	2.5	4	13.93	12.23	12.87	12.83	12.94
Sm	3.72	3.26	4.13	3.22	3.67	3.71	3.33	0.09	2.7	4	3.51	3.08	3.21	3.43	3.23
Eu	1.18	1.14	1.27	1.06	1.26	1.26	1.14	0.04	3.5	4	1.23	0.98	1.08	1.17	1.12
Gd	3.51	3.12	3.62	2.80	3.38	3.30	3.08	0.09	3.1	4	3.08	2.82	2.98	3.30	2.91
Tb	0.58	0.50	0.61	0.48	0.57	0.55	0.51	0.01	2.3	4	0.52	0.48	0.51	0.55	0.51
Dy	3.53	3.08	3.58	2.87	3.36	3.30	3.09	0.07	2.3	4	3.18	2.86	2.98	3.26	3.01
Ho	0.72	0.63	0.71	0.56	0.68	0.66	0.63	0.01	2.1	4	0.65	0.58	0.61	0.68	0.63
Er	2.05	1.77	1.99	1.54	1.95	1.87	1.81	0.06	3.3	3	1.88	1.68	1.81	1.90	1.76
Tm	0.28	0.24	0.27	0.21	0.27	0.26	0.25	0.01	3.0	4	0.25	0.23	0.25	0.26	0.25
Yb	1.78	1.49	1.76	1.27	1.68	1.61	1.64	0.05	2.8	4	1.69	1.53	1.59	1.71	1.58
Lu	0.28	0.24	0.27	0.20	0.27	0.25	0.26	0.01	3.7	4	0.28	0.25	0.25	0.28	0.26
Y	18.97	16.58	20.65	9.40	18.43	17.90	17.24	0.63	3.6	6	17.60	14.88	17.54	18.65	17.14
Zr	110.00	103.00	115.00	106.00	100.00	106.00	104.79	0.77	0.7	1	103.00	106.00	103.00	102.00	104.00
Hf	2.63	2.25	2.48	2.21	2.10	2.16	2.15	0.07	3.1	6	2.12	2.15	2.10	2.15	2.06
Nb	9.60	4.08	6.05	4.93	5.46	4.54	4.55	0.18	3.9	6	4.73	4.16	4.70	5.15	4.41
Ta	1.28	0.27	0.34	0.27	0.22	0.25	0.25	0.01	5.3	6	0.28	0.26	0.26	0.31	0.25
Pb	4.00	4.33	5.09	5.28	4.83	5.02	5.51	0.98	17.7	6	5.17	5.02	5.25	4.64	5.31
Th	2.24	1.31	2.17	0.29	1.01	1.04	0.85	0.06	7.2	6	0.89	1.35	0.90	0.80	0.89
U	0.55	0.32	0.59	0.36	0.35	0.37	0.26	0.02	7.5	6	0.26	0.39	0.28	0.25	0.28
Ga	17	18	17	20	19	19	20.37	1.46	7.2	12	21	18	18	21	19
Cu	28	55	33	49	53	67	48.16	3.98	8.3	14	44	59	50	58	44
Zn	76	80	88	85	83	86	90.26	8.56	9.5	6	91	87	95	82	92
V	205	190	202	173	213	222	153.00	9.08	5.9	6	156	150	169	168	154
Sc	29	28	27	27	22	30	15.16	2.03	13.4	10	16	17	19	16	17
Cr	88	198	69	26	32	43	22.53	9.58	42.5	10	89	29	28	33	23
Ni	21	83	25	22	15	23	9.53	2.46	25.8	16	16	15	11	16	11

Table 2. Whole-rock compositions of dark andesite from 1992 eruptions of Crater Peak vent of Mount Spurr, Alaska.

[Major elements normalized to 100 percent anhydrous. FeOt is total iron as FeO. Number of reported decimal places may exceed analytical precision.]

Sample	June 27 eruption										August 18 eruption										September 16-17 eruption									
	Major-element oxides, in weight percent										Major-element oxides, in weight percent										Major-element oxides, in weight percent									
	92JBS01	92JBS02	92JBS03	9209268	92amm	92amm	92amm	92amm	92amm	92amm	92JES-03	92JES-06	92JES-07	92JES-08	NS(2)92-18	920924-1	920924-3	920926-6	92AMm-120B	92AMm-120C	92AMm-129	Stop33								
	100	100	101	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B	111B								
SiO ₂	56.68	56.64	56.61	56.74	56.51	56.45	56.50	56.39	56.42	56.87	56.44	56.96	56.86	56.62	57.05	56.54	56.55	57.19	57.85											
TiO ₂	0.68	0.68	0.68	0.69	0.68	0.68	0.69	0.69	0.70	0.68	0.68	0.67	0.70	0.68	0.70	0.70	0.69	0.68	0.69											
Al ₂ O ₃	19.14	19.08	19.13	18.98	19.27	19.13	19.05	19.05	18.79	18.95	19.19	18.97	18.99	19.17	19.12	18.87	18.98	18.99	19.13											
FeOt	6.97	7.03	7.01	7.07	6.91	7.07	7.12	7.10	7.40	6.93	6.93	6.93	6.93	7.02	6.59	7.23	7.13	6.78	5.93											
MnO	0.15	0.15	0.15	0.15	0.15	0.15	0.15	0.15	0.16	0.15	0.15	0.15	0.15	0.15	0.15	0.16	0.15	0.15	0.15											
MgO	3.56	3.58	3.54	3.67	3.58	3.62	3.65	3.70	3.59	3.64	3.69	3.60	3.77	3.66	3.72	3.81	3.69	3.59	3.65											
CaO	7.64	7.59	7.61	7.58	7.58	7.58	7.60	7.64	7.79	7.57	7.63	7.49	7.51	7.58	7.55	7.52	7.56	7.43	7.44											
Na ₂ O	4.03	4.06	4.09	3.94	4.12	4.14	4.05	4.09	3.95	4.04	4.08	4.03	3.89	3.94	3.92	3.95	4.05	3.97	3.95											
K ₂ O	0.88	0.87	0.88	0.88	0.90	0.90	0.91	0.90	0.91	0.90	0.91	0.93	0.92	0.89	0.94	0.94	0.92	0.94	0.97											
P ₂ O ₅	0.29	0.29	0.29	0.29	0.30	0.29	0.28	0.29	0.29	0.28	0.29	0.28	0.28	0.30	0.26	0.28	0.28	0.28	0.24											
Trace elements, in parts per million																														
Cs	0.4	0.4	0.4	0.42	0.43	0.42	0.43	0.44	0.46	0.45	0.44	0.48	0.48	0.42	0.57	0.45	0.46	0.51	0.53											
Rb	16.09	15.99	16.06	13.86	15.13	15.55	14.32	14.48	13.37	13.51	13.28	14.71	15.28	13.71	15.30	17.03	16.46	15.86	17.27											
Ba	384	386	382	371	402	398	380	379	385	379	382	398	383	381	387	378	371	398	397											
Sr	662	670	666	671	667	663	664	661	653	660	665	654	647	671	651	649	655	649	640											
La	11.4	11.17	10.87	11.09	11.31	11.54	10.86	11.14	10.96	10.97	11.10	11.20	10.96	11.13	10.80	10.53	10.57	11.09	11.15											
Ce	23.11	23.08	22.98	22.95	23.39	23.71	22.08	22.45	22.40	22.47	22.64	22.81	22.41	22.96	22.07	21.60	21.44	22.67	22.82											
Pr	3.03	3.04	3.01	3.04	3.01	3.07	2.91	2.89	2.90	2.88	2.94	2.98	2.89	2.96	2.95	2.89	2.85	2.97	2.99											
Nd	13.7	13.68	13.72	13.57	13.65	13.7	13.01	12.95	13.29	13.08	13.03	13.09	12.99	13.24	13.05	12.73	12.70	13.24	13.18											
Sm	3.48	3.46	3.39	3.40	3.36	3.42	3.31	3.27	3.26	3.26	3.34	3.40	3.16	3.36	3.26	3.23	3.15	3.39	3.36											
Eu	1.22	1.2	1.18	1.15	1.2	1.19	1.13	1.09	1.13	1.12	1.11	1.13	1.10	1.13	1.10	1.14	1.07	1.15	1.14											
Gd	3.23	3.19	3.19	3.10	3.24	3.15	2.92	3.01	3.10	3.06	3.00	3.17	2.93	3.13	3.03	3.00	3.02	3.11	2.99											
Tb	0.52	0.52	0.53	0.51	0.54	0.53	0.51	0.50	0.50	0.50	0.51	0.51	0.50	0.52	0.50	0.51	0.51	0.50	0.50											
Dy	3.15	3.2	3.17	3.16	3.16	3.14	2.98	3.04	3.00	3.05	3.07	3.13	2.95	3.09	3.01	3.07	3.06	3.16	3.07											
Ho	0.65	0.66	0.65	0.64	0.66	0.64	0.62	0.63	0.62	0.64	0.61	0.64	0.62	0.64	0.62	0.62	0.62	0.63	0.64											
Er	1.87	1.96	1.86	1.81	1.86	1.88	1.80	1.81	1.75	1.72	1.78	1.82	1.79	1.79	1.72	1.72	1.78	1.82	1.80											
Tm	0.26	0.27	0.26	0.26	0.25	0.26	0.24	0.25	0.25	0.25	0.24	0.26	0.25	0.26	0.24	0.24	0.25	0.25	0.25											
Yb	1.71	1.75	1.65	1.64	1.69	1.71	1.60	1.65	1.61	1.64	1.64	1.64	1.58	1.65	1.60	1.58	1.59	1.63	1.65											
Lu	0.27	0.28	0.27	0.26	0.28	0.28	0.27	0.27	0.26	0.27	0.26	0.26	0.25	0.26	0.26	0.25	0.25	0.26	0.26											
Y	17.89	17.83	18.21	17.36	18	17.75	15.34	17.43	17.32	17.12	16.75	16.55	16.86	17.05	16.69	17.02	17.26	17.38	17.58											
Zr	104	105	104	105	105	105	106	105	105	103	105	106	104	105	105	104	104	105	106											
Hf	2.21	2.24	2.16	2.15	2.22	2.24	2.08	2.11	2.13	2.05	2.07	2.15	2.11	2.21	2.18	2.13	2.00	2.19	2.21											
Nb	4.82	4.73	4.86	4.63	4.74	4.71	4.06	4.56	4.55	4.53	4.45	4.46	4.46	4.53	4.41	4.46	4.40	4.62	4.57											
Ta	0.27	0.28	0.26	0.24	0.27	0.27	0.24	0.25	0.25	0.25	0.24	0.25	0.24	0.25	0.26	0.24	0.23	0.25	0.26											
Pb	5.83	5.35	5.14	5.00	5.44	4.96	5.58	5.29	5.12	4.96	5.09	5.44	5.31	9.54	5.21	5.25	5.25	5.22	5.69											
Th	0.83	0.81	0.8	0.82	0.91	0.83	0.86	0.90	0.79	0.77	0.74	0.82	0.92	0.79	0.88	0.87	0.85	0.95	0.98											
U	0.26	0.25	0.25	0.25	0.27	0.26	0.26	0.25	0.23	0.23	0.21	0.25	0.28	0.24	0.26	0.27	0.27	0.28	0.29											
Ga	20	18	21	19	20	19	20	20	22	22	19	22	23	21	20	22	19	22	18											
Cu	45	54	57	49	49	50	51	50	51	46	44	41	49	50	48	46	46	49	40											
Zn	84	91	92	81	91	88	95	90	91	93	96	94	101	84	68	96	85	111	84											
V	135	150	138	165	161	141	160	154	150	147	147	142	159	163	166	161	156	158	154											
Sc	17	18	16	12	16	16	18	15	16	14	14	16	15	17	15	14	11	11	17											
Cr	17	19	19	14	16	16	21	23	17	22	22	59	30	16	30	23	19	23	22											
Ni	8	9	6	9	6	10	8	11	8	9	8	17	13	8	10	10	9	12	10											

Four of the five light gray-green andesites that were analyzed are compositionally identical to the dark andesite, except for slightly higher and more variable Cs (table 1). This similarity is remarkable in view of the drastically different composition of the groundmass glass in the light gray-green andesites (rhyolite glass) versus dark andesites (andesite to dacite glass) (Harbin and others, this volume). The similar whole-rock chemistry in spite of different groundmass chemistry presumably reflects, on the scale of individual blocks, more extensive crystallization of the dark andesite to produce the light gray-green andesite, perhaps at the chamber margin. The fifth sample (92AMm113c) is more evolved, and it contains 58.8 percent SiO_2 . This sample has no distinguishing petrographic characteristics, and it may contain a higher ratio of glass to crystals.

COMPARISON TO PREHISTORIC CRATER PEAK MAGMAS

Prehistoric Crater Peak lavas are hornblende-bearing two-pyroxene calcalkaline basaltic andesite and andesite (Nye and Turner, 1990). Besides being slightly more mafic than all but a very few ancestral cone-building andesites, their most important distinguishing feature compared to other Mount Spurr lavas is the presence of centimeter-sized, holocrystalline, cumulus pyroxenite clots (Nye and Turner, 1990). The 1992 magma is also calcalkaline andesite, but it falls outside chemical trends of prehistoric magmas in many ways. The 1953 and prehistoric lava compositions normalized to average 1992 andesite are shown in figure 2. Data for selected samples are presented in table 1. The prehistoric andesites that are plotted have SiO_2 within 1 weight percent of the 1992 andesite. Important differences between the 1992 andesite and previous andesites, at nearly constant SiO_2 , are higher concentrations of Al, Na, and Sr, and lower concentrations of Ti, V, Mg, Sc, Cr, and Ni. These two observations suggest more protracted crystallization of the parent to the 1992 andesite at mid- to lower-crustal depths, where the increased pressure favors ferromagnesian silicate crystallization over plagioclase crystallization. MgO and Al_2O_3 versus SiO_2 are shown in figure 3 for the entire Mount Spurr suite. These plots suggest that increased plagioclase (high Al_2O_3) and decreased pyroxene (low MgO) components in the 1992 andesites are shared by the 1953 andesites, but not by most older magmas. Low MgO by itself does not uniquely demonstrate pyroxene crystallization, but pyroxene is likely to be the most abundant ferromagnesian phase. Greater-than-normal reservoir depths may be characteristic of latest Holocene Crater Peak magmatism.

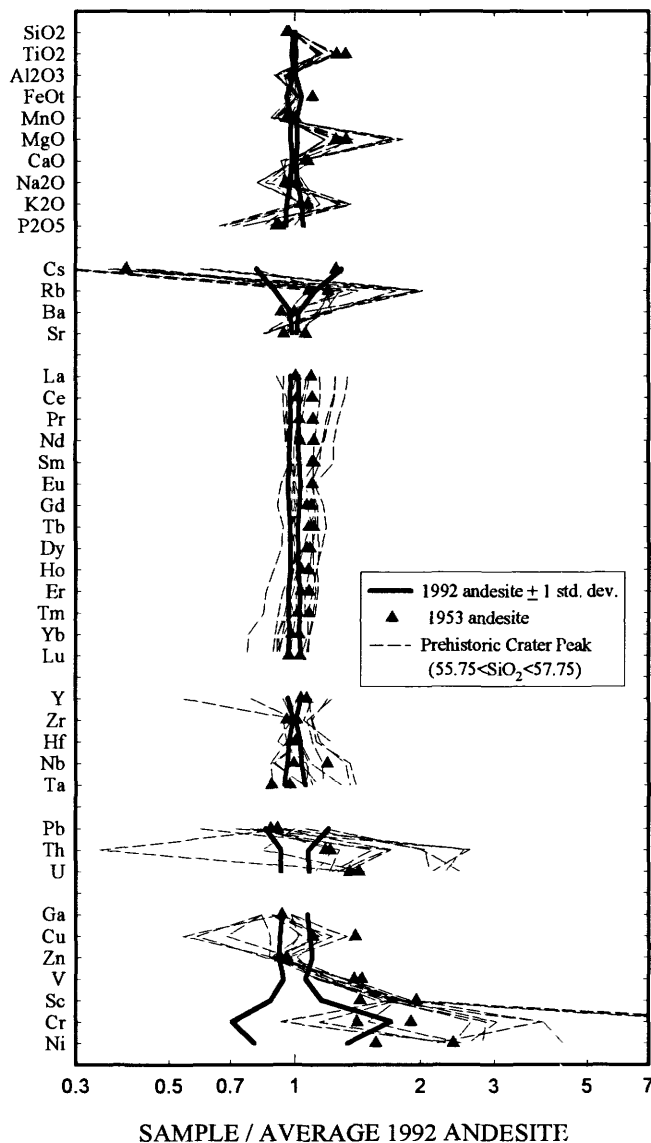


Figure 2. Mean composition of andesite erupted during 1992 from Crater Peak volcano, Mount Spurr, Alaska, compared to 1953 andesite and prehistoric Crater Peak andesite. Each line is a single sample's composition divided by the mean composition of 1992 andesite blocks. Gaps separate groups of geochemically similar elements. FeO_t is total iron as FeO .

There are also major differences in the abundances and ratios of highly incompatible trace elements between 1992 andesite and previous andesite of similar SiO_2 . Specifically, 1992 andesite has Cs concentrations about two times higher than older andesite, yet slightly lower concentrations of the other alkali elements K and Rb. These differences are about 20 times analytical uncertainty. Compared to most previous andesite, 1992 andesite has low concentrations of Th and U, yet it has about the same Pb content. These differences in the relative concentrations of

highly incompatible elements cannot be generated by fractional crystallization and must reflect some process in the source that varies for individual batches of magma. Andesite from the 1992 eruptions has a slightly lower La/Yb ratio, and thus it is slightly less light-rare-earth-element enriched than previous andesite. Concentrations of the high-field-strength elements are comparable for some, but not all, samples of previous Crater Peak andesite.

Rb-SiO₂ variations in samples from the 1992 and 1953 eruptions of Crater Peak are shown in figure 4; also shown are Rb-SiO₂ variations from prehistoric flows from Crater Peak and the eastern and western remnants of the ancestral Crater Peak. The relative stratigraphic positions of samples are known, with the exception that the relative ages of the east and west ancestral Crater Peak sections are unknown. Three groups of samples (Crater Peak and west and east ancestral Crater Peak) are connected by lines in order of stratigraphic succession. In each case the oldest samples are the most silicic. In many cases Rb variations in near-neighbor mafic flows are so large that they cannot be related by fractional crystallization or mixing. Rb variations at 54 to 55 percent SiO₂ are large enough that simple crystal sorting also cannot

explain the variations. These relations suggest that relatively small volumes of chemically unrelated magma can migrate from deep within the system to feed just a few eruptions.

COMPARISON TO 1953 MAGMA

An important volcanological question is whether the 1992 andesite erupted from new magma that recently migrated to the upper crust, or whether it erupted from the remnant of the upper-crustal magma body that fed the 1953 eruption. Analyses of two samples from proximal 1953 pumice deposits are shown in table 1 and plotted in figures 2, 3, and 4. Andesite from 1992 is unlike 1953 andesite in that, despite its higher SiO₂, it has lower concentrations of some incompatible elements such as Rb, U, and the middle rare-earth elements. These differences cannot be produced by fractional crystallization, and they suggest that the 1992 andesite indeed represents a new magma. Calculated fractional crystallization paths for 1953 magmas are shown in figure 4. The crystallization calculation is from Nye and Turner (1990) and is based on detailed least-squares calculations to relate SiO₂ in-

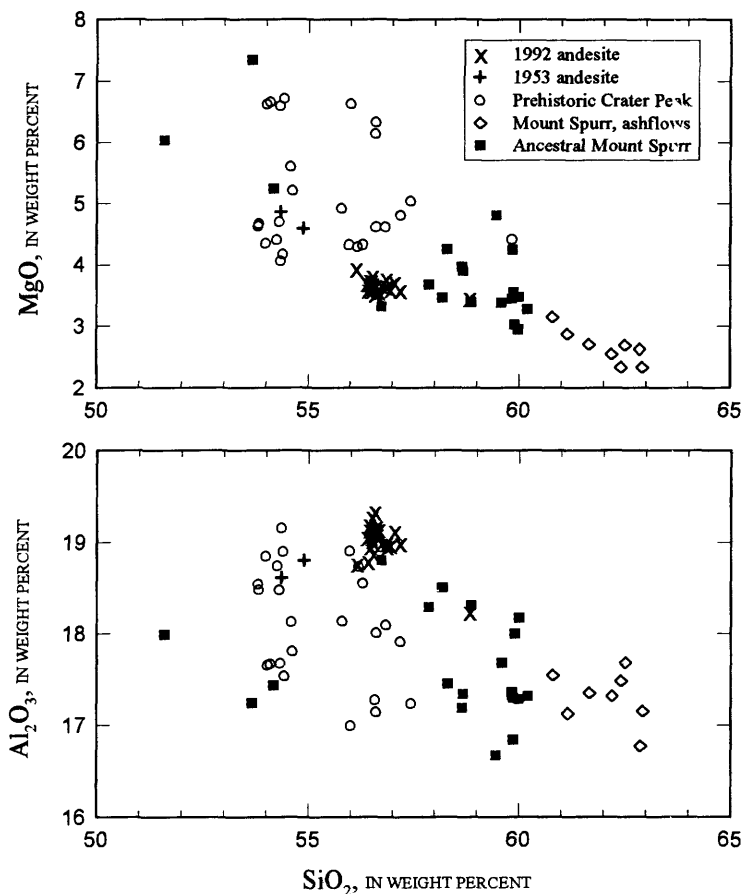


Figure 3. MgO and Al₂O₃ versus SiO₂ (in weight percent) of Pleistocene, Holocene, and historic samples from Mount Spurr volcano, Alaska. Samples not discussed in this paper are discussed in Nye and Turner (1990). Note the high Al₂O₃ and low MgO of 1992 andesite compared to previous Crater Peak andesite and other Mount Spurr lavas.

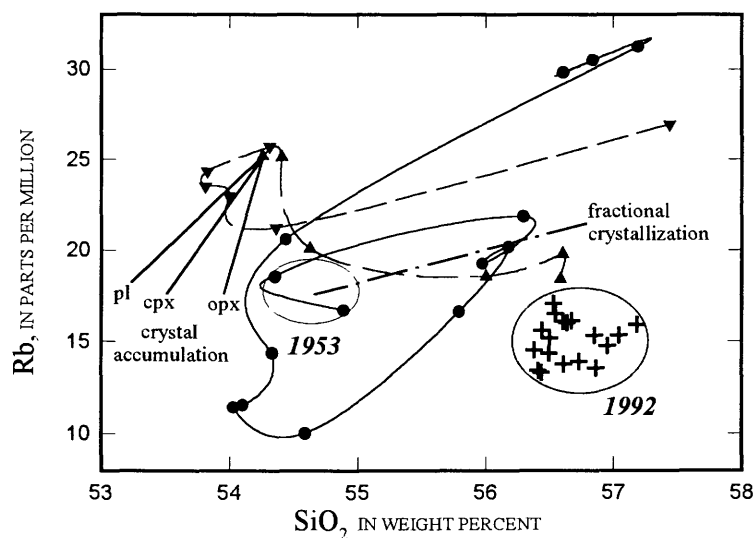


Figure 4. Rb (in ppm) versus SiO_2 (in weight percent) for samples from the Crater Peak vent of Mount Spurr volcano, Alaska. Pre-1992 samples are discussed in Nye and Turner (1990). The plot also shows data from east (inverted triangles) and west (upright triangles) ancestral Crater Peak and the current Crater Peak (filled circles). Lines connecting symbols represent stratigraphic order, with the most siliceous end of each line being the oldest.

crease to the amount of fractionation. If the 1992 andesite was derived by simple fractional crystallization of the 1953 magma, then about 16 to 19 percent (by weight) of crystals would have to have separated and the concentrations of Rb should have increased to about 21 ppm, compared to the 15 ppm Rb in the 1992 andesite. The differences in Rb between 1953 and 1992 andesite are also unlikely to have been produced by open-system processes such as fractionation combined with crustal assimilation, which also causes increasing Rb with increasing SiO_2 .

Andesite from 1953 is also unlike 1992 andesite in its high Ti, Mg, Sc, Cr, and Ni. Some or all of this is to be expected because the 1953 andesite has lower SiO_2 , and it is therefore less evolved.

TEPHRA

Tephra samples were analyzed in a search for compositional diversity in 1992 ejecta and to test for compositional differences between the light-colored (less dense) and dark-colored (more dense) tephra components (Neal and others, this volume). A more extended discussion is presented by Neal and others (this volume). The main conclusions are that all the light tephra and half of the dark tephra samples are identical to 1992 andesite in their concentrations of 10 major and 33 trace elements. Those dark tephra samples that are different have anomalously high Cs, Rb, and com-

patible transition metals including Cr and Ni. The anomalous compositions are the products of chemical heterogeneities on the scale of a few grams of material. When large masses (>50 grams) of carefully chosen lapilli were crushed together, the anomalous compositions were not found. The contaminant has not yet been identified, but it is not small fragments of the xenoliths discussed in the next section.

XENOLITHS

ALUMINOUS CORDIERITE- AND SILLIMANITE-BEARING XENOLITHS

1992 ejecta also include a minor proportion of xenoliths. These are partially fused metamorphic rocks with 60 to 70 percent SiO_2 , which contain glass, plagioclase, and some combination of cordierite, sillimanite, garnet, biotite, pyroxene, and spinel. Their bulk compositions fall off the trend of Mount Spurr lavas toward high Ti, Fe, K, Cs, Rb, Ba, REE, Nb, Ta (but not Zr and Hf), Y, U, Th, Pb, V, Cr, and Ni and low Na, Ca, and Sr (table 3, figure 5). The bulk-xenolith composition may not reflect the composition of the protolith because the liquid (now glass) may have moved with respect to the crystal residue. Additional work on glass separates is in progress. The high transition metal concentrations and low Na make it unlikely that xenoliths such as these have been routinely

Table 3. Whole-rock compositions of selected metamorphic xenoliths from 1992 eruptions of Crater Peak vent, Mount Spurr, Alaska.

[Major elements normalized to 100 percent anhydrous. FeOt is total iron as FeO. Number of reported decimal places may exceed analytical precision.]

Sample #	Aluminous xenoliths										Siliceous xenoliths									
	92AMm	92AMm	92AMm	92AMm	92JES	92AMm	92AMm	92AMm	92AMm	92JES	92AMm	92MHS	92MHS	92MHS	92AMm	92AMm	92JES	92JES	92JES	92JES
	102D	111A-1	102F	111A-2	-05	102A	113B	04a	113A	-09	02a	05b	05b	05b	113D	113D	-02	-02	-04	-04
Major-elements, in weight percent																				
SiO ₂	50.34	60.15	61.03	62.34	63.17	64.66	65.02	65.39	65.83	67.55	68.38	69.19	76.46	76.17	75.74	75.74	76.17	76.17	75.74	75.74
TiO ₂	1.33	1.01	0.87	0.92	0.88	0.87	1.02	0.84	0.84	0.77	0.79	0.75	0.00	0.02	0.02	0.02	0.02	0.02	0.02	0.02
Al ₂ O ₃	24.69	18.83	18.06	19.07	17.32	16.18	17.79	16.40	17.24	15.86	15.39	15.81	14.05	13.13	14.04	14.04	13.13	13.13	14.04	14.04
FeOt	6.83	8.57	7.65	7.20	7.81	6.77	6.73	5.96	6.63	5.82	6.07	4.40	0.57	0.71	0.48	0.48	0.71	0.71	0.48	0.48
MnO	0.06	0.19	0.24	0.28	0.16	0.15	0.11	0.12	0.14	0.13	0.13	0.09	0.02	0.07	0.03	0.03	0.07	0.07	0.03	0.03
MgO	3.68	3.81	3.34	3.33	3.71	2.88	2.51	2.96	2.44	2.85	2.21	2.12	0.00	0.01	0.00	0.00	0.01	0.01	0.00	0.00
CaO	7.95	2.08	2.66	1.75	1.94	2.96	1.49	3.04	1.99	2.05	1.91	2.34	1.10	3.20	1.91	1.91	3.20	3.20	1.91	1.91
Na ₂ O	3.75	2.45	2.90	2.73	1.99	2.92	2.13	2.24	2.64	2.16	2.28	2.21	7.11	5.85	6.47	6.47	5.85	5.85	6.47	6.47
K ₂ O	1.27	2.62	2.98	2.18	2.69	2.36	3.00	2.85	2.05	2.58	2.63	2.92	0.52	0.79	1.15	1.15	0.79	0.79	1.15	1.15
P ₂ O ₅	0.10	0.29	0.25	0.20	0.33	0.27	0.19	0.20	0.21	0.23	0.23	0.16	0.18	0.06	0.16	0.16	0.06	0.06	0.16	0.16
Trace-elements, in parts per million																				
Cs	0.47	4.48	5.76	1.94	3.10	3.02	4.46	2.74	2.40	2.79	2.57	3.45	0.20	0.17	0.26	0.26	0.17	0.17	0.26	0.26
Rb	30.84	81.65	115.4	55.96	83.47	71.48	85.48	69.42	66.81	76.72	75.99	68.82	9.75	13.26	22.79	22.79	13.26	13.26	22.79	22.79
Ba	1045	1231	1258	2174	791	944	883	837	760	773	931	679	58	154	174	174	154	154	174	174
Sr	793	221	236	153	228	304	231	296	229	216	243	282	45	123	116	116	123	123	116	116
La	15.35	29.95	27.64	27.82	26.79	24.57	27.45	23.35	26.88	23.82	24.46	22.76	1.75	2.54	3.48	3.48	2.54	2.54	3.48	3.48
Ce	25.31	56.93	54.39	46.72	51.80	46.29	52.96	44.43	51.04	45.27	46.79	43.42	3.33	4.60	6.40	6.40	4.60	4.60	6.40	6.40
Pr	2.78	6.38	6.34	6.21	5.72	5.44	6.21	5.25	6.00	5.26	5.54	5.11	0.35	0.51	0.69	0.69	0.51	0.51	0.69	0.69
Nd	11.31	25.66	26.04	25.60	23.14	22.64	26.15	21.22	24.52	21.75	22.80	20.84	1.04	1.85	2.51	2.51	1.85	1.85	2.51	2.51
Sm	2.06	5.83	5.92	5.49	5.22	4.98	6.70	4.79	5.70	4.76	5.36	4.72	0.39	0.50	0.67	0.67	0.50	0.50	0.67	0.67
Eu	2.2	1.26	1.1	1.04	1.39	1.44	1.29	1.17	1.42	1.22	1.39	1.23	0.02	0.17	0.17	0.17	0.17	0.17	0.17	0.17
Gd	2.53	5.04	5.74	5.27	4.58	4.71	6.49	4.38	5.00	4.46	5.08	4.02	0.29	0.50	0.53	0.53	0.50	0.50	0.53	0.53
Tb	0.26	0.82	0.9	0.85	0.80	0.77	1.16	0.72	0.90	0.76	0.88	0.68	0.08	0.10	0.08	0.08	0.10	0.10	0.08	0.08
Dy	1.47	4.74	5.33	5.01	5.19	4.62	6.95	4.41	5.59	4.60	5.37	3.96	0.49	0.52	0.45	0.45	0.52	0.52	0.45	0.45
Ho	0.3	0.94	1.09	1.04	1.13	0.98	1.43	0.91	1.16	0.93	1.14	0.81	0.09	0.10	0.08	0.08	0.10	0.10	0.08	0.08
Er	0.8	2.62	3.28	2.96	3.39	2.86	3.99	2.58	3.32	2.69	3.30	2.28	0.24	0.26	0.20	0.20	0.26	0.26	0.20	0.20
Tm	0.11	0.36	0.47	0.40	0.49	0.4	0.55	0.35	0.47	0.37	0.46	0.33	0.05	0.04	0.03	0.03	0.05	0.05	0.04	0.04
Yb	0.75	2.41	3.07	2.66	3.22	2.58	3.39	2.40	3.12	2.40	2.95	2.07	0.41	0.29	0.20	0.20	0.29	0.29	0.20	0.20
Lu	0.14	0.39	0.5	0.42	0.53	0.43	0.53	0.38	0.50	0.38	0.47	0.33	0.06	0.04	0.03	0.03	0.06	0.06	0.04	0.04
Y	7.87	23.08	29.5	22.77	31.32	26.29	35.80	25.09	28.01	25.41	30.94	21.77	2.71	3.45	2.78	2.78	3.45	3.45	2.78	2.78
Zr	190	170	145	171	149	149	209	152	153	136	155	160	44	31	33	33	44	44	31	31
Hf	3.75	4.68	3.37	3.87	4.05	4.02	5.01	4.17	4.17	3.59	4.15	4.28	2.28	1.13	1.14	1.14	2.28	2.28	1.13	1.13
Nb	8.97	12.23	12.02	8.39	12.22	11.4	12.04	10.57	9.12	10.30	10.21	9.50	2.26	3.44	3.49	3.49	3.44	3.44	3.49	3.49
Ta	0.64	0.81	0.84	0.59	0.77	0.69	0.87	0.71	0.61	0.76	0.61	0.65	0.40	0.70	0.90	0.90	0.70	0.70	0.90	0.90
Pb	4.23	12.69	14.12	16.13	13.53	14.2	11.67	8.87	9.74	12.43	11.89	11.77	10.18	9.20	10.37	10.37	9.20	9.20	10.37	10.37
Th	0.8	8.48	7.05	5.67	6.93	5.31	6.31	5.28	6.81	5.84	5.55	5.18	0.40	0.64	0.75	0.75	0.64	0.64	0.75	0.75
U	0.38	2.15	2.11	2.17	1.66	1.5	2.30	1.65	1.07	2.02	1.21	1.58	4.47	2.23	2.18	2.18	2.23	2.23	2.18	2.18
Ga	28	24	20	29	20	20	22	21	18	17	20	16	17	17	16	16	17	17	16	16
Cu	266	110	72	87	36	59	59	70	44	9	47	19	21	18	18	18	21	21	18	18
Zn	120	144	146	139	164	109	132	80	111	106	87	74	7	11	2	2	7	7	11	11
V	299	242	206	226	201	171	172	167	171	163	173	130	10	14	9	9	10	10	14	14
Sc	13	25	22	21	21	23	20	16	19	16	19	15	6	1	0	0	6	6	1	1
Cr	148	133	118	140	206	119	125	116	98	117	92	83	6	6	4	4	6	6	4	4
Ni	48	57	56	63	97	51	46	48	52	49	47	30	6	9	8	8	6	6	9	9

assimilated by Mount Spurr magmas. Nye and Turner (1990) appealed to a large-ion lithophile element (LILE)-rich, high-field-strength element (HFSE)- and heavy rare-earth element (HREE)-poor crustal contaminant in Spurr magmas. This component could logically be derived from small degrees of partial melting of country rock leaving such phases as garnet, spinel, zircon, and rutile in the residue. Whereas these xenoliths are not that component, their glass may be.

SILICEOUS XENOLITHS

A few siliceous, bright white, plagioclase-quartz-glass xenoliths of uncertain parentage also occur. They make up a very small percentage of erupted material. They have rhyolitic SiO_2 (table 3), but they are strongly depleted in all elements except Na, Rb, Ta, U, and Ni (fig. 5). Both mobile and immobile trace elements are depleted. Concentrations of K, Rb, Ba, Zr, and Y are all very low; they average around 0.8 percent, 17 ppm, 130 ppm, 36 ppm, and 2 ppm, respectively. For comparison, concentrations of these elements in Novarupta rhyolites, are 3.2 percent, 63 ppm, 907 ppm, 150 ppm, and 48 ppm (Hildreth, 1983; Nye, unpublished data). The siliceous xenoliths are unlikely to represent crustal minimum melts or any kind of a high-temperature crustal distillate, particularly because of their low LILE. The parents to these xenoliths may be some highly depleted protolith, such as hydrothermal veins, or they may be strongly leached rock of some other lithology. Xenoliths that are macroscopically similar to these are also found in prehistoric Crater Peak andesites.

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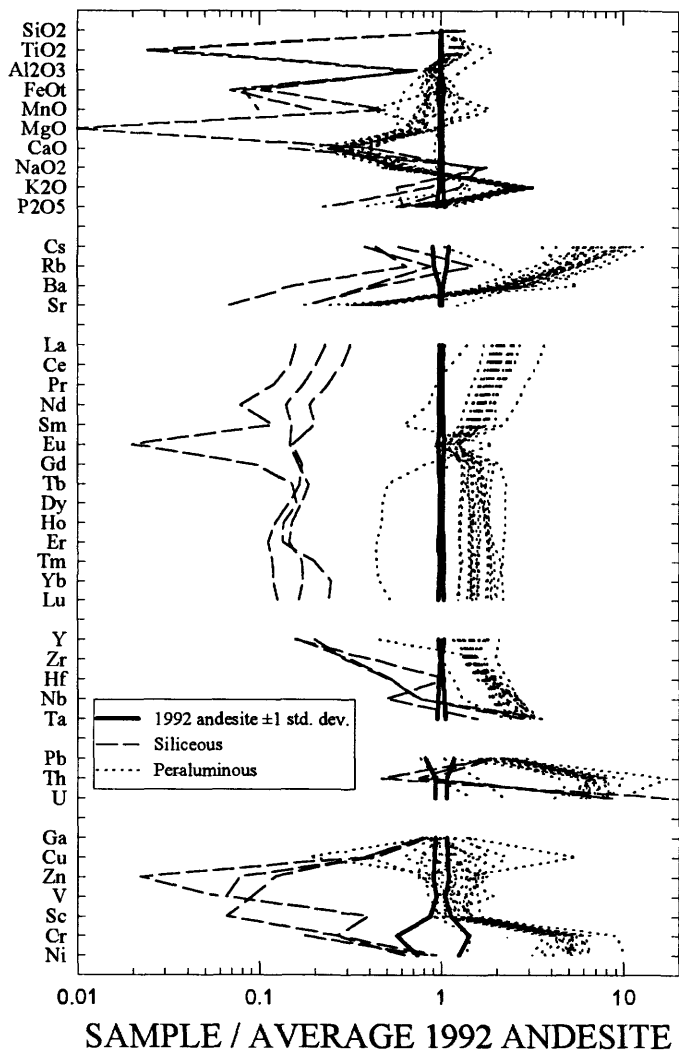


Figure 5. Compositions of aluminous and siliceous xenoliths compared to the mean composition of 1992 andesite. Each xenolith sample was normalized to the mean of all 1992 juvenile andesite blocks. FeO_t is total iron as FeO .

Use of Volcanic Glass from Ash as a Monitoring Tool: An Example from the 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

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ABSTRACT

Compositions of volcanic glass from ash ejected during the 1992 eruptions of the Crater Peak vent of Mount Spurr are dominantly andesitic (SiO_2 of 61–62 weight percent), but they include much smaller amounts of dacitic (SiO_2 of 63–69 weight percent) and rhyolitic (SiO_2 of 74–77 weight percent) glass. In each eruption (June, August, and September), glass in the ash and matrix glass in rocks from proximal deposits show similar chemical trends.

The rhyolitic glass component was first recognized in the ash from the June 27, 1992, eruption of Crater Peak. Subsequent work on proximal samples resulted in the identification of rhyolitic glasses in the matrices of juvenile volcanic blocks and in partially melted metamorphic xenoliths (buchites). Re-examination of prehistoric tephra materials from Crater Peak reveals a similar rhyolitic component; this

discovery suggests that generation of minor amounts of silica-rich melt is common at Crater Peak. Rapid analysis of volcanic glass in ash from the 1992 Crater Peak eruptions proved a useful monitor of the chemistry of the erupting magma.

INTRODUCTION

Volcanic ash is a common component of eruptions of andesitic magmas. The ash can be hazardous, especially to aircraft and other air-filtering cultural activities (Swanson and Kienle, 1988; Casadevall, 1992). Volcanic ash is fine-grained material that enters the air during an eruption. Ash may be injected into the atmosphere directly as part of the explosive phase of the eruption or as an elutriating cloud from advancing pyroclastic flows or avalanches. Materials in the ash are typically glass (quenched melt) as well as mineral and rock fragments. The mineral and rock fragments may be juvenile magmatic material or part of the vent fill, crater, or conduit walls that were entrained during eruption.

Petrologists have been reluctant to monitor the petrologic evolution of volcanic systems using ash because of the potential for mechanical sorting during ash transportation and deposition (Steen-McIntyre, 1977). Another problem is the heterogeneity of ash; it may contain various proportions of rock fragments, mineral grains, and volcanic glass, any of which can be accidental or comagmatic. Even so, ash can be easily sampled and can provide samples of erupted material *before* proximal materials can be sampled. Volcanic ash can also reveal something about the nature of the ongoing eruption (that is, whether or not new magma is involved, Swanson and others, 1991). Bulk compositions of volcanic ash are subject to sorting or contamination, however, compositions of individual phases within the ash may provide important clues about an erupting magmatic system, especially if one can demonstrate that the phase is clearly a juvenile magmatic product.

The glass phase (quenched melt) in the ash is a sensitive monitor of the magmatic conditions during an eruption (Swanson and others, 1991). Most, if not all, the glass in andesitic ash represents the fractionated melt after crystallization of some mineral phases. Accordingly, the composition of the melt is controlled by the mineralogy and proportion of the crystallizing phases, which, in turn, are controlled by magmatic conditions (for example, temperature, pressure, water content, and oxygen fugacity). Using the morphology of glass particles (delicate, vesicular, often shard-like), it is relatively easy to identify juvenile material in the ash. Small differences in glass compositions produced by different fractionating mineral assemblages can quickly and routinely be measured using an electron microprobe. Glass compositions from different eruptions will be different depending on the fractionation process, and there may even be changes in glass compositions during an eruption. This lesson was learned from numerous tephra studies of individual eruptions (Mount St. Helens, Scheidegger and Federman, 1982; Pinatubo, Pallister and others, 1992; 1912 Novarupta, Federman and Scheidegger, 1984, Avery, 1992; Redoubt Volcano 1989–90 eruption, Swanson and others, 1994).

Volcanic glass is generally enriched in incompatible components with respect to the bulk-rock composition (see Swanson and others, 1994). Thus it is common for andesite to have a matrix glass with a rhyolitic composition produced by crystallization of plagioclase, pyroxenes, FeTi oxides, and perhaps hornblende from the bulk magma.

If glass in volcanic ash is used to monitor magmatic systems during volcanic eruptions, then it must be established that changes in the magmatic system are clearly reflected in the glass chemistry in a timely fashion. During the 1989–90 eruptions of Redoubt Volcano, there was a change from bimodal dacite-rhyolite matrix glass early in the eruption to solely rhyolite matrix glass (Swanson and others, 1994). This change in matrix glass compositions (determined from blocks and pumice bombs) was perfectly mimicked by the glass in the volcanic ash (Swanson and others, 1991).

Questions about the nature of the erupting material and what that means in terms of eruption phenomena are some of the first questions asked during the early stages of a volcanic eruption (Sarna-Wojcicki and others, 1981). Recognition of juvenile magmatic material (typically represented by glass) in an eruption can result in an upgrading of the hazard assessment for that event (Fiske, 1984). Some workers (Sparks and others, 1977; Gourgaud and others, 1989; Pallister and others, 1992) think that magma mixing is an important triggering process in some eruptions.

Thus, the timely recognition of mixed magma during an eruption may indicate a higher hazard level than for single-magma eruptions.

Volcanic ash provides an ideal material for monitoring magmatic changes during an eruption. Glass compositions in the Redoubt ashes (measured after the eruption) reflected changes in the magmatic system. This observation prompted us to consider using volcanic ash glass compositions as a way of monitoring magmatic changes during a protracted volcanic eruption (Swanson and others, 1991). We developed techniques for rapid sample preparation and analysis of glass from volcanic ash. The eruption of Crater Peak in 1992 provided the opportunity to test our technique of monitoring an ongoing volcanic eruption using glass from volcanic ash.

1992 ERUPTIONS OF CRATER PEAK

Crater Peak, a vent on the flank of Mount Spurr volcano, erupted three times during the summer and fall of 1992 (Eichelberger and others, this volume). The eruptions were from the same vent that last erupted in 1953 (Juhle and Coulter, 1955). Each of the eruptions was similar to the 1953 event: a moderate subplinian eruption that lasted for several hours and then stopped. Each of the eruptions was accompanied by a large ash cloud that was carried upward from 6 to 15 km and was then dispersed to the north and east by prevailing winds (Neal and others, this volume). Pyroclastic flows accompanied the eruptions in August and September, and ash was produced from these flows (Miller and others, this volume).

ERUPTION OF JUNE 27, 1992

Ash from the June 27 eruption came from an erupting column with a maximum height estimated by radar to be 14.5 km (Rose and others, this volume). Prevailing winds carried the ash north over Mount McKinley, just west of Fairbanks, and over the Yukon River just downstream from Fort Yukon on June 27 (Neal and others, this volume).

The ash provided the first material from the June 27 eruption for analysis. Samples of the ash were obtained from near the volcano and from distal sites (200 to 450 km north of Mount Spurr) in the Mount McKinley area, Lake Minchumina, and Manley Hot Springs. All of the ash samples are similar. Maximum grain size of the ash is about 0.4 mm within 50 km of the volcano and 0.11 to 0.13 mm in the distal samples from beyond Mount McKinley. Most of the ash is composed of dark-brown to black porphyritic volcanic rock

fragments. The rock fragments contain micrphenocrysts set in a groundmass of brown glass with abundant microlites. Isolated phenocryst fragments are also found in the ash. Plagioclase is the most common phenocryst phase, but small amounts of augite, hypersthene, hornblende, and FeTi oxides are also present in the ash and in the volcanic rock fragments. The hornblende is pleochroic from green to brown. Colorless, microlite-free blocky or microvesicular fragments of glass are rare (much less than 1 percent).

At first, shortly after the eruption, it was not clear whether juvenile magmatic material had been erupted. However, the presence of delicate shards of colorless glass (fig. 1) that could not have survived "recycling" in this eruption confirmed the presence of at least some juvenile material in the eruptive products. Subsequent geochemical work (Nye and others, this volume) shows the eruptive products are compositionally distinct from other recent Mount Spurr lavas.

The June ash is a fine-grained version of the blocks deposited near the volcano both in terms of mineralogic composition and relative abundance. Most of the proximal lapilli and blocks deposited on June 27 consist of black, vesicular andesite, but rare (much less than 1 percent) white to light-gray compositionally banded pumice lapilli are also present (Harbin and others, this volume). The dark andesite contains phenocrysts set in a matrix of brown glass with abundant microlites, whereas the light-colored pumice lapilli contain phenocrysts in a matrix of colorless, microlite-free glass. Details of the petrography are given in Harbin and others (this volume). Nye and others (this volume) discuss the bulk geochemistry of the pumice and andesite.



ERUPTION OF AUGUST 18, 1992

Strong explosive eruptions carried ash to altitudes of 15 km during the eruption of Crater Peak on August 18, 1992. Winds carried the ash east, directly over Anchorage (Neal and others, this volume) and then to the south and east over southeastern Alaska.

With the fall of ash on Anchorage, collection of ash samples was easy. One suite of samples (AVO-PAT) was collected over timed intervals (about 30 to 60 minutes) as the ash fell on the patio of the Alaska Volcano Observatory at Gould Hall, 124 km downwind from the volcano. Most of the ash fell during a period of about 3 hours.

Initial maximum grain size of the ash at Anchorage was about 0.5 mm in diameter, and it decreased to about 0.3 mm by the end of the ash fall. The color of the ash also changed from a light yellowish-brown to a darker gray-brown by the end of the eruption. The lighter color appears to be partly the result of light-gray, fine-grained ash that coats the darker, larger fragments and partly the result of a difference in vesicularity.

The ash consists of dense, dark volcanic rock fragments with phenocrysts of plagioclase, pyroxene, hornblende, and FeTi oxides in a matrix of brown glass rich in microlites. Fragments of these same phenocrysts are also found in the ash. The amount of plagioclase crystal fragments is greatest followed by pyroxene, hornblende, and FeTi oxides. These observations are based on visual estimates. About 1 percent of the ash is composed of microvesicular or blocky fragments of colorless microlite-free glass. There is no apparent variation in the proportions of these phases within the suite of ash samples collected during the 3-hour ashfall.

The August 18 ash is virtually identical to the ash from the June eruption and the variation in the ash also mimics the variation in the bombs and blocks deposited near the volcano. Black, breadcrusted bombs of andesite and lighter vesicular greenish-gray andesite blocks are the most common component in the lahar and pyroclastic-flow deposits (Miller and others, this volume) from the August eruption. This black andesite corresponds to the dark volcanic rock fragments in the ash, but the vesicular greenish-gray andesite does not have an obvious counterpart in the ash. Details of andesite petrography are given by Harbin and others (this volume), and Nye and others (this volume) discuss the bulk composition of the andesite.

Figure 1. Secondary electron image of a rhyolitic glass shard from ash of the June 27, 1992, eruption of Crater Peak taken with an electron microprobe at the University of Alaska, Fairbanks by Michelle Harbin. White scale bar is 20 μ m.

Table 1. Composition of glass from volcanic ash and pumice lapilli erupted on June 27, 1992, from Crater Peak vent, Mount Spurr volcano, Alaska.

[* , total Fe as FeO; number of points analyzed in parentheses below sample number]

Sample no.	Volcanic ash		Pumice lapilli					
	SS-2 (13)	SS-2 (1)	JB-4A (9)	JB4B1 (6)	JB-4B2 (21)	JB-4B2 (15)	JB-4B3 (14)	JB-4B3 (19)
glass color	brown	colorless	colorless	colorless	colorless	light brown	colorless	brown
SiO ₂	60.51±0.93	76.60	74.28±1.26	75.66±1.37	71.59±1.23	67.41±1.45	75.87±1.37	61.00±0.80
TiO ₂	.60±0.25	0.37	0.40±0.18	0.19±0.13	0.33±0.13	0.44±0.12	0.22±0.14	0.62±0.14
Al ₂ O ₃	17.22±1.00	12.60	13.40±0.95	12.16±0.87	14.41±0.48	15.06±0.55	12.45±1.47	16.97±0.69
FeO*	6.46±0.89	1.11	2.31±0.30	1.24±0.11	2.60±0.42	4.28±0.52	0.85±0.25	6.23±0.71
MgO	2.34±0.85	0.24	0.50±0.08	0.15±0.03	0.55±0.08	1.04±0.16	0.18±0.15	2.56±0.48
CaO	5.45±0.61	0.37	1.01±0.25	0.66±0.12	1.28±0.38	3.02±0.38	0.79±0.15	5.85±0.31
Na ₂ O	4.35±0.34	3.54	2.56±0.14	0.80±0.36	3.27±0.31	3.72±0.52	0.42±0.13	4.36±0.44
K ₂ O	1.36±0.10	4.99	4.40±0.14	6.83±0.40	3.99±0.30	2.67±0.16	7.06±0.33	1.43±0.45
Cl	0.21±0.04	0.17	0.02±0.01	0.04±0.02	0.03±0.03	0.15±0.04	0.11±0.03	0.21±0.03
total	98.50	99.98	98.88	98.74	98.05	97.79	97.97	98.74

Buchites, partially melted blocks of metamorphic rocks, form pumiceous blocks in the proximal August deposits (Harbin and others, this volume). Highly vesicular, colorless, microlite-free glass is present in all of the buchites, and it ranges in abundance from trace to 30 percent. The glass in the buchites is identical in appearance to the small amounts of colorless glass found in the ash.

ERUPTION OF SEPTEMBER 16–17, 1992

The major phase of the September eruption of Crater Peak began just after midnight local time on September 17, 1992. A series of violent explosions again carried ash high into the air. Winds carried ash to the east, to the north of Anchorage, and eventually into the Yukon Territory of Canada.

Ash was sampled at 45 and 110 km downwind from Crater Peak. The ash that fell 45 km downwind from the volcano had a maximum grain size of 0.7 mm, whereas the ash that fell 110 km downwind had a maximum grain size of 0.5 mm. Petrographically, the ash from both sites is very similar. Blocky fragments of porphyritic andesite (containing fragments of phenocrysts of plagioclase, hornblende, augite, hypersthene, and FeTi oxides in a matrix of brown glass with abundant microlites) were the most common constituent of the ash. Isolated grains of the phenocryst minerals were also common as angular fragments in the ash. A few grains of quartz and cordierite were also found in this ash during microprobe analysis. Fragments of colorless, microlite-free, highly vesiculated glass with a few plagioclase microphenocrysts were noted in the finer grained (about 0.5 mm) fraction of the ash.

Blocks from the September eruption show the same pattern of petrographic variation as observed in the ash. Dense, black glassy blocks of andesite contain phenocrysts in a matrix of light-brown glass with abundant microlites. Some blocks of light-gray andesite contain phenocrysts in a matrix of colorless glass rich in microlites of plagioclase, pyroxene, and FeTi oxides. Buchites, similar to those from the August deposits, were also erupted in September. Harbin and others, and Nye and others (both this volume) provide more information on the petrology and geochemistry of the andesites.

COMPOSITION OF THE 1992 VOLCANIC GLASSES

Compositions of volcanic glass from each of the Mount Spurr eruptions were determined by electron microprobe at the University of Alaska Fairbanks. Glasses from the 1992 eruptions were analyzed on a four-channel electron microprobe using an 8 µm beam at 15 Kev with a 10-sec counting time and a 0.01 microamp sample current. Natural glasses were used as standards. The analytical routine was designed to minimize alkali loss in these relatively anhydrous glasses. Experience with this same analytical routine on the Redoubt glasses (Swanson and others, 1994) showed no significant alkali loss occurs during the analytical routine when sodium is measured (first 10 seconds of the analysis). All of the analyses were made on polished grain mounts of the ash. Multiple points were selected for analysis (one point per grain) and the microprobe was run in the automated mode. Analy-

ses that obviously included mineral grains and glass were eliminated from the data set, as were analyses with totals outside of the range 98 to 101 weight percent oxide.

Compositions of the prehistoric tephra glasses were determined on a nine-channel microprobe at a U.S. Geological Survey laboratory. The glasses were analyzed at 15 keV, with a 10-second counting time, 0.01-microamp sample current, and a beam size of 10 μ m. Natural glasses were used as standards and rhyolitic glass was analyzed repeatedly as an internal standard. Somewhat lower analytical totals (96–99 weight percent) for the prehistoric glasses are probably related to the hydration of these glasses.

COMPOSITION OF THE JUNE 27, 1992, VOLCANIC GLASSES

Both proximal and distal tephra samples were analyzed from the eruption of June 27, 1992. The distal sample (SS-2) was collected from Manley Hot Springs, approximately 450 km north of the Crater Peak vent on Mount Spurr. Proximal samples (JB-4A, JB-4B1, JB-4B2, and JB-4B3) consisted of small pumice lapilli and were collected a few kilometers north of Crater Peak.

The ash and pumice lapilli both contain a heterogeneous suite of glass compositions (table 1). The common brown glass with abundant microlites is typically andesitic with an SiO₂ content of about 61 weight percent (columns 1 and 8, table 1). However, brown glass in one of the pumice lapilli (JB-4B2) is dacitic (SiO₂ of 67.4 weight percent, column 6, table 1). The colorless microlite-free glass in the ash (SS-2) is rhyolitic (SiO₂ of 76.6 weight percent, column 2, table 1). Similar microlite-free colorless rhyolitic glass is found in the pumice lapilli (columns 3, 4, 5, 7; table 1).

COMPOSITION OF THE AUGUST 18, 1992, VOLCANIC GLASSES

Glass in the ash of August 18, 1992, is heterogeneous. Three distinct populations of glass compositions are found in the ash (table 2). The most common (more than 90 percent) component is a brown microlite-rich andesitic glass (SiO₂ of about 62 weight percent, column 1, table 2). A few percent of very light brown dacitic glass (SiO₂ of 65 weight percent, column 2, table 2) with abundant microlites are also present. A rare (less than 1 percent) colorless, microlite-free rhyolitic glass (SiO₂ of about 76 weight percent, column 3, table 2) is also found in the ash erupted in August.

COMPOSITIONS OF THE SEPTEMBER 16–17, 1992, VOLCANIC GLASSES

Ash analyzed from the eruption of September 16–17, 1992 (sample 19B, table 3) was collected 110 km downwind from Crater Peak. The sample is a bulk sample for the entire eruption.

Four compositionally distinct glass populations are present in the September ash (table 3). Three brown dacitic glasses (SiO₂ of about 63–69 weight percent, columns 1, 2, 3, table 3) with abundant microlites are the most common, but colorless, microlite-free rhyolitic glass is also present in the suite (column 4, table 3).

COMPARISON OF GLASS COMPOSITIONS FROM THE ASH AND BLOCKS IN THE 1992 ERUPTIONS

The pattern of glass compositional variation found in the ash is generally mimicked by the variation of matrix glass compositions measured in the blocks (Harbin and others, this volume). Similar results were

Table 2. Composition of glass from volcanic ash erupted on August 18, 1992, from Crater Peak vent, Mount Spurr volcano, Alaska, collected on a patio at Alaska Volcano Observatory in Anchorage.

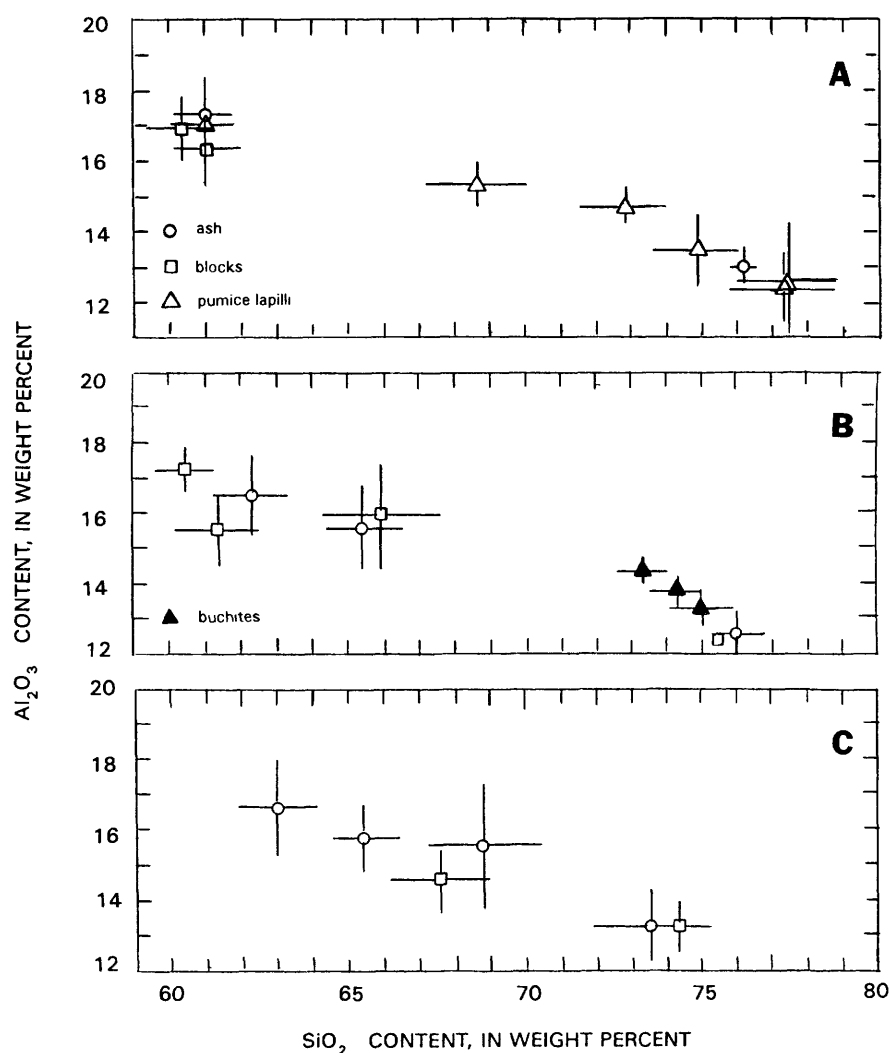
[*, total Fe as FeO; number of points analyzed in parentheses below sample number; AVO-PAT, patio at Alaska Volcano Observatory]

Sample no.	AVO-PAT (107)	AVO-PAT (18)	AVO-PAT (12)
glass color	brown	light brown	colorless
SiO ₂ -----	62.25±1.32	65.49±0.92	75.93±1.36
TiO ₂ -----	0.59±0.17	0.59±0.11	0.35±0.17
Al ₂ O ₃ -----	16.45±1.11	15.62±1.16	12.55±0.45
FeO* -----	5.62±0.96	4.82±1.11	2.29±0.28
MgO -----	2.01±0.61	1.36±0.48	0.46±0.05
CaO -----	5.37±0.58	4.42±0.51	0.91±0.10
Na ₂ O -----	5.26±0.33	5.32±0.40	3.06±0.16
K ₂ O -----	1.39±0.19	1.80±0.25	4.07±0.27
Cl -----	0.20±0.05	0.23±0.07	0.02±0.01
Total -----	99.14	99.62	99.64

Table 3. Composition of glass from volcanic ash erupted on September 17, 1992, from Crater Peak vent, Mount Spurr volcano, Alaska.

[* , total Fe as FeO; number of points analyzed in parentheses below sample number]

Analysis no.	19B (159)	19B (75)	19B (14)	19B (14)
glass color	brown	light brown	light brown	colorless
SiO ₂ -----	63.02±1.14	65.61±1.00	69.16±1.41	73.37±1.68
TiO ₂ -----	0.53±0.16	0.52±0.16	0.48±0.16	0.51±0.24
Al ₂ O ₃ -----	16.64±1.32	15.87±0.92	15.71±1.68	13.28±1.02
FeO* -----	5.14±1.12	4.65±1.04	3.06±1.40	2.39±1.01
MgO -----	1.87±0.71	1.39±0.51	0.88±0.80	0.46±0.46
CaO -----	5.20±0.62	4.30±0.58	3.20±0.96	1.54±0.66
Na ₂ O -----	5.34±0.36	5.38±0.36	5.17±0.92	3.80±0.89
K ₂ O -----	1.43±0.25	1.80±0.51	2.45±0.40	4.01±1.09
Cl -----	0.20±0.06	0.21±0.06	0.14±0.08	0.15±0.16
Total -----	99.37	99.73	100.25	99.51

**Figure 2.** Glass compositions from the June 27, 1992 (A), 18 August 1992 (B), and September 17, 1992 (C) eruptions of Crater Peak. Bars on the points represent the one sigma standard deviations about the mean of the data. Data on blocks and buchites are from Harbin and others (this volume). Data on ash and pumice lapilli are from tables 1-3.

found in the 1989–90 Redoubt eruption (Swanson and others, 1991) where matrix glass and glass from volcanic ash showed similar patterns of variation on the same time scale.

The June ash contains an andesitic and rhyolitic component (table 1). Comparison of these glasses to the matrix glass in blocks from the June eruption (Harbin and others, this volume) shows that the andesitic glass is derived from the dark gray to black “cauliflower” bombs noted in the proximal June deposits (Miller and others; Harbin and others, this volume). After recognition of the rhyolitic component in the June ash, a careful search was made of the June proximal deposits for possible block equivalents. A few small lapilli of light-gray pumice were found in the proximal deposits. Some of these lapilli contained bands with brown andesitic matrix glass. However, the dominant glass in the lapilli was a colorless, microlite-free rhyolite ranging from 68.6 to 76.3 weight percent SiO_2 (fig. 2). The rhyolitic glass in the gray pumice lapilli is very similar to the rhyolitic glass in the ash (fig. 2).

Compositions of glass from the ash and blocks were also similar in the August eruption (fig. 2). The brown andesitic glass in the ash is represented in the blocks by the dark “cauliflower” bombs of andesite. A dacitic glass is found in the ash (table 2, column 2) and in a block of greenish andesite scoria (Harbin and others, this volume). The matrix glass in the andesitic scoria is very light brown and the glass contains abundant microlites. Rhyolitic glass is found in the ash (table 2, column 3), in vesicular, gray-green andesite, and in buchites in the August deposits. The gray-green andesite has colorless matrix glass without microlites and is very similar to the rhyolitic glass found in the ash (fig. 2). Glass in the buchites is also colorless and microlite-free, but is slightly lower in SiO_2 than either the ash or the greenish-gray andesitic glasses (fig. 2).

Rhyolitic and dacitic glasses are found in both the ash and in the matrix of blocks produced in the

September eruption. The rhyolitic matrix glass is colorless and contains abundant microlites. Rhyolitic glass from the ash (column 4, table 3) is identical in composition to the rhyolitic matrix glass found in a light-gray prismatically-jointed andesite block (fig. 2). High-silica dacitic glass (about 69 weight percent SiO_2) is found in the September ash (table 3, column 3) and in the matrix of a black prismatic andesite block. The two high-silica dacitic glasses are very similar in composition (fig. 2). Brown, microlite-rich low-silica dacitic and high-silica andesitic glasses are found in the matrix of blocks from the September eruption (Harbin and others, this volume), and the low-silica dacitic matrix glass is similar to glass in the ash (fig. 2).

GLASS COMPOSITIONS OF TEPHRA FROM PREHISTORIC ERUPTIONS OF CRATER PEAK VENT, MOUNT SPURR

Brown andesitic glass, similar to that erupted in 1992, is the common glass in the prehistoric tephra from Crater Peak vent of Mount Spurr (Riehle, 1985) and is easily recognized in the Quaternary deposits of south-central Alaska (Riehle, 1985; Beget and others, 1994). These are the only Cook Inlet volcanoes that have produced andesitic glass—the other Cook Inlet volcanoes produce dacitic or rhyolitic glasses (Riehle, 1985).

A volumetrically small but persistent rhyolitic glass phase is also found in the prehistoric tephra from Crater Peak (Riehle, 1985). The rhyolitic glass is colorless and microlite-free. Commonly the rhyolitic glass is vesicular. Reexamination of microprobe analyses reveals that rhyolitic glass is present in most prehistoric Crater Peak tephra samples.

New analyses of this rhyolitic component (table 4) have been done on some prehistoric tephras to compare these rhyolitic glasses with those of the 1992 eruption. The rhyolitic glasses in the prehistoric te-

Table 4. Rhyolitic glass compositions of prehistoric basal tephras and pyroclastic-flow samples from Mount Spurr volcano, Alaska.

[*, total Fe as FeO ; number of points analyzed in parentheses below sample number]

	Pyroclastic flow		Basal tephra	
SiO_2 -----	74.58±0.75	74.06±2.35	74.61±1.75	73.46±1.95
TiO_2 -----	0.39±0.03	0.45±0.16	0.45±0.09	0.46±0.07
Al_2O_3 -----	11.89±0.72	12.57±0.44	13.31±0.67	12.53±1.52
FeO^* -----	1.45±0.14	2.11±0.80	1.86±0.47	2.10±0.33
MnO -----	0.01±0.02	0.08±0.03	0.05±0.04	0.06±0.03
MgO -----	0.13±0.06	0.04±0.45	0.27±0.12	0.43±0.23
CaO -----	0.72±0.37	1.42±0.43	1.40±0.30	1.48±0.66
Na_2O -----	3.00±0.21	2.97±0.36	2.96±0.26	2.84±0.21
K_2O -----	4.45±0.16	3.96±0.78	4.01±0.17	3.85±0.87
Total -----	96.62	98.02	98.92	96.94

phras are very similar to those from the 1992 eruptions. All of the Crater Peak vent of Mount Spurr rhyolitic glasses are peraluminous and they show similar ranges of alkali content (fig. 3).

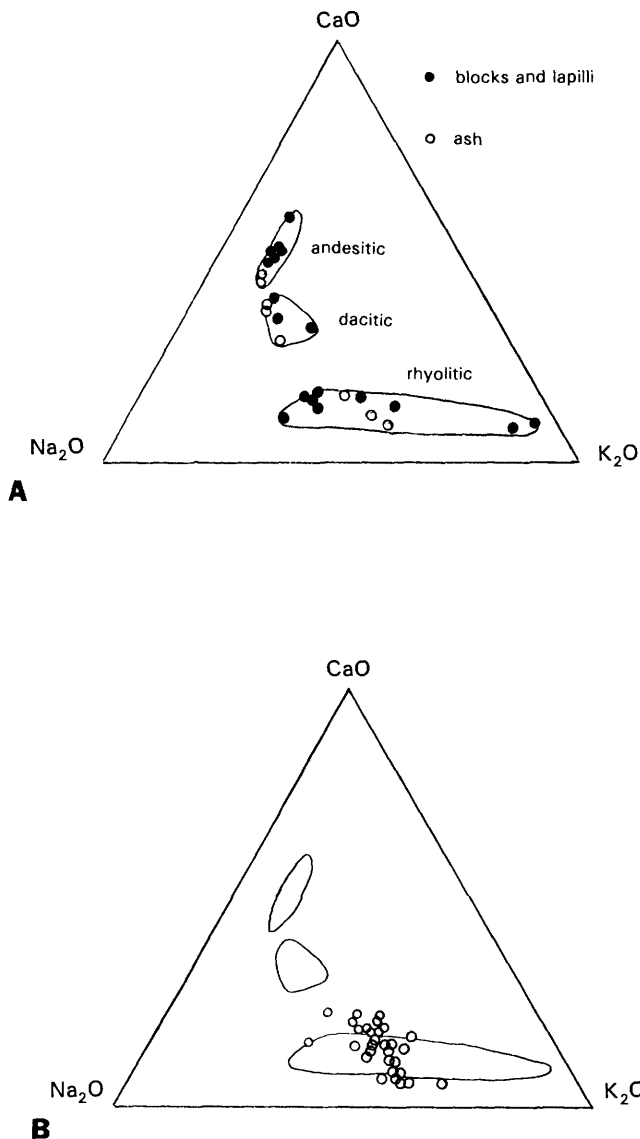


Figure 3. A, Compositions of the 1992 Crater Peak glasses in terms of CaO-Na₂O-K₂O. Compositional fields outline andesitic, dacitic, and rhyolitic glasses. Data sources as in figure 2. B, Individual analyses (averages given in table 4) of rhyolitic glasses in prehistoric Crater Peak tephra in terms of CaO-Na₂O-K₂O. Composition fields for the 1992 glasses are reproduced from figure 3A.

DISCUSSION AND FUTURE WORK

Glass compositions in the ash from the 1992 eruptions of Crater Peak provided a valuable means of monitoring changes in the magmatic system. The glass compositions in the ash and matrix glass from blocks all show the same patterns for each of the eruptions.

Rhyolitic glasses similar to those in the 1992 eruption occur in prehistoric tephra from Mount Spurr and Crater Peak, indicating that the 1992 eruption was typical of other prehistoric eruptions. Rhyolitic glass occurs in ash, the matrix of blocks, and in buchites from the August 1992 eruption (fig. 2). All of these rhyolitic glasses are very similar (fig. 3), supporting an origin of the rhyolitic melt by partial melting of metamorphic basement rocks. The variable alkali contents of the rhyolitic glasses (for example, compare sodium and potassium of columns 3–8, table 1) are typical of the disequilibrium melts produced in the initial stages of melting pelitic buchites (Grapes, 1986). The occurrence of these rhyolitic glasses in prehistoric deposits suggests that similar buchites should be found in prehistoric Crater Peak deposits.

During an eruption of multiple magmas it is commonly difficult to estimate the relative proportions of each of the magmas. Glass compositions in volcanic ash provide a relatively easy technique for estimating relative proportions of magmas in an eruption. For example, table 3 shows averages of glass analyses that represent over 250 individual analyses of glass in a bulk sample of the September ash. All of the points were selected randomly in a polished grain mount. Most of the glass in the September sample is represented by low-silica dacitic compositions (234 out of 262 total analyses); high-silica dacitic and rhyolitic glasses each amounted to only 14 out of 262 analyses. This provides one quantitative way of determining the relative proportions of the magmatic components that erupted in September. Some caution is needed in the application of this technique because of the potential for differential sorting of comagmatic phases in the ash clouds. The fragments of crystal-rich andesite glass might be expected to fall closer to the volcano than the relatively crystal free shards of rhyolite glass, but this has not been apparent in the petrographic examination of proximal and distal samples. When ash can be sampled sequentially, either as it falls or from well-preserved stratigraphy within the ash layer, detailed analysis of the subsets provides useful clues to magmatic variation during an eruption.

CONCLUSIONS

Volcanic ash provides important clues to the nature of an erupting magmatic system. The composition of glasses in the 1992 Crater Peak ashes proved that new (different from prehistoric) magma was being erupted and that a small rhyolitic component was involved in the eruption. Because the ash was dispersed great distances from the volcano, it was possible to sample the magma at safe, distal locations.

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Preliminary Petrology and Chemistry of Proximal Eruptive Products: 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

By Michelle L. Harbin, Samuel E. Swanson, Christopher J. Nye, and Thomas P. Miller

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ABSTRACT

The 1992 eruptions of Crater Peak, Mount Spurr, Alaska, produced two types of andesite distinguished macroscopically by color and texture. The two andesites also have slight petrographic differences. The mineral assemblage of plagioclase, orthopyroxene, clinopyroxene, amphibole, and FeTi oxides is the same for both andesites, and compositions of the mineral phases are nearly identical. However, the color and composition of the groundmass glass differ; one type is brown and andesitic (60–62 weight percent SiO₂) and the other type is colorless and rhyolitic (73–75 weight percent SiO₂). Intermediate dacitic glass compositions are also present. The stability of hornblende phenocrysts also differs between the andesites. In the andesite containing brown andesitic glass, hornblende typically shows little or no reaction corona, whereas in andesite with rhyolitic groundmass glass hornblende phenocrysts have well-developed coronas. Xenoliths of partially melted metamorphic rocks were also ejected during the eruptions.

INTRODUCTION

Mount Spurr, Alaska, located approximately 125 km west of Anchorage, is the northernmost historically active volcanic center of the eastern Aleutian arc. Prehistoric Mount Spurr was an andesitic stratovolcano that underwent avalanche caldera formation approximately 10,000 years ago (Nye and Turner, 1990). Following subsequent ash flows, a dacite dome was emplaced in the caldera and now forms the summit of Mount Spurr. Crater Peak, a flank vent, was built in the breach of the southern caldera rim and has been the site of historic volcanism, including the eruptions in 1953 and 1992.

The 1992 eruptions were subplinian (after Fisher and Schmincke, 1984) and occurred on June 27, August 18, and September 16–17. Each eruption lasted 3.5 to 4 hours (AVO, 1993; Eichelberger and others, this volume). Most of the eruptive material consisted of tephra fall (Neal and others, this volume; Swanson and others, this volume); pyroclastic flows (Miller and others, this volume), debris flows (Meyer and others, this volume), and ballistic showers (Waitt and others, this volume) were less significant. This study concentrates on juvenile blocks emplaced either as bombs or as clasts in the pyroclastic flows during all three eruptions (table 1). Originally, there was some debate as to whether juvenile material was ejected during the June eruption. However, presence of fresh vesicular glass fragments in the distal tephra fall (Swanson and others, this volume), small pumice lapilli, and geochemistry (Nye and others, this volume) indicate that juvenile magma was indeed erupted. Recycling of 1953 olivine-bearing material may be reflected in two olivine-bearing samples—one collected from the June deposits, and one collected from a September pyroclastic flow. The ejecta have been broadly classified into three groups, two of which are juvenile andesite on the basis of hand-sample characteristics. The first group consists of brown-black, breadcrusted, cauliflower-shaped andesite (approximately 95 percent by

Table 1. Eruption date, location and description, and bulk and groundmass glass weight percent SiO₂ of individual andesite samples from 1992 eruptions of Mount Spurr, Alaska.

[Bulk-rock data from Nye and others, this volume. Average values for both bulk-rock and groundmass glass normalized to 100 percent. NA, not analyzed]

<u>Sample number</u>	<u>Eruption date</u>	<u>Location and description</u>	<u>SiO₂ content bulk rock</u>	<u>SiO₂ content groundmass glass</u>
92JBS-01	June 27	Dense, brown, cauliflower-shaped andesite bomb (23 cm in diameter)	56.68	61.74±0.84
92JBS-03A	June 27	Black rounded breadcrusted andesite bomb (8 cm in diameter)	56.61	61.36±0.85
92JBS-04A	June 27	White pumice lapilli (1 cm in diameter)	NA	74.85±0.54
92JBS-04B1	June 27	Pumice lapilli (1 cm in diameter)	NA	77.04±1.49
92JBS-04B2	June 27	Pumice lapilli (1 cm in diameter)	NA	68.05±2.27; 72.31±1.63
92JBS-04B3	June 27	Pumice lapilli (1 cm in diameter)	NA	61.02±0.77; 77.18±1.65
92CNS-03B	June 27	Rounded breadcrusted (cauliflower-shaped) andesite bomb (8 cm in diameter) from north side of Crater Peak (2,149 m elevation)	NA	60.59±0.98
92MHS-12	Aug. 18	Brown cauliflower-shaped andesite (9 cm in diameter) from pyroclastic flow (762 m elevation)	NA	62.18±0.78
92AMm-101	Aug. 18	Brown andesite from pyroclastic flow	56.45	60.44±0.64
92AMm-102GS	Aug. 18	Greenish colored scoria from fallout	56.59	65.87±1.60
92AMm-111A-1	Aug. 18	Pumiceous metamorphic clast	60.15	73.42±0.67
92AMm-111A-3	Aug. 18	Pumiceous metamorphic clast	66.68	74.44±0.75
92AMm-111B	Aug. 18	Brown andesite from east side of Crater Peak	56.50	61.82±1.00
92AMm-113A	Aug. 18	Pumiceous metamorphic material from pyroclastic flow on the south side of Crater Peak	65.02	74.97±0.90
92AMm-113C	Aug. 18	Sugary, grayish-green andesite from pyroclastic flow on the south side of Crater Peak	58.84	75.56±0.38
92AMm-120A	Sept. 17	Grayish-green prismatic andesite from 1,402 m elevation on the west tongue of Kidazgeni Glacier	56.52	74.03±1.24
92AMm-120B	Sept. 17	Dark-gray to black prismatic andesite from 1,402 m elevation on the west tongue of Kidazgeni Glacier	56.54	67.47±1.68
92AMm-129a	Sept. 17	Dark-gray to black andesite from 1,829 m elevation of Kidazgeni Glacier	57.19	64.36±1.18
Stop 35	Sept. 17	Brown vesicular andesite from 1,402 m elevation of Kidazgeni Glacier	57.85	62.34±1.04; 64.60±0.83
920924-3	Sept. 17	Dense, dark-gray to black andesite from 1,402 m elevation of Kidazgeni Glacier	56.62	65.44±0.70
920926-6	Sept. 17	Dark-gray to black andesite from 1,402 m elevation of Kidazgeni Glacier	57.05	64.72±1.06
				64.35±1.37

volume of pyroclastic flows). The second group consists of greenish-gray andesite (approximately 1 percent by volume of pyroclastic flows) erupted in August and September. The third group consists of xenoliths and old volcanic rocks (approximately 4–5 percent by volume of pyroclastic flows) (Miller and others, this volume).

PETROLOGY AND CHEMISTRY

Petrographically, the andesite of the first two groups are nearly identical. Both contain phenocrysts of plagioclase, orthopyroxene, clinopyroxene, hornblende, and FeTi oxides. One difference between the andesites is the groundmass glass. The brown andesite has a groundmass of very microlite-rich (plagioclase, pyroxene, FeTi oxides) brown glass. The greenish-gray andesite has a colorless groundmass glass with or without microlites. The other striking difference between the two andesite types is the apparent stability of hornblende phenocrysts. Hornblende in the brown andesite does not have reaction coronas, but hornblende phenocrysts in the greenish-gray andesite have well-developed reaction coronas.

Metamorphic xenoliths are present in small pumice lapilli from the June eruption and as large blocks in the pyroclastic flows and bomb fields of August and September. These xenolithic blocks are as much as 50 cm in maximum diameter and typically are highly inflated (40–60 percent vesicles). Phases represented in the xenoliths include vesicular colorless rhyolitic glass, quartz, plagioclase, pyrrhotite±orthopyroxene±cordierite±spinel±magnetite±ilmenite±sillimanite±biotite±garnet±apatite±graphite±corundum.

ANALYTICAL TECHNIQUES

Analyses of polished thin sections were performed on the four-channel electron microprobe at the University of Alaska, Fairbanks. Mineral analyses were done using an accelerating voltage of 15 Kev, a beam current of 10 nA, a beam diameter of 5 μ m, and a counting time of 10 sec per element. Well-characterized natural minerals were used as standards. Glass analyses were done at 15 Kev, 10 nA, 8 μ m, and a 10-sec counting time per element. Natural glasses were used as standards for all major elements. The scapolite standard was used as a standard for chlorine. This analytical routine was designed to minimize Na loss during analysis of these glasses. Analysis of Redoubt Volcano samples using this same analytical routine (Swanson and others, 1994) indicates no significant loss of Na during an analysis when Na is measured

within the first 10 seconds of the analysis. Analyses that clearly included mineral grains were eliminated from the data set. Statistical tests were then performed on the data set to develop a good total range of acceptable analyses. The microlite-rich glasses were the most contaminated, and some of the heterogeneity of these glasses can be attributed to the presence of the microlites.

GROUNDMASS GLASS

The brown-black andesite of the June and August eruptions contains a microlite-rich brown glass, but most samples from the September eruption and some from the August eruption contain a microlite-rich, lighter-brown glass. However, it is nearly impossible to macroscopically identify samples with the light-brown glass because they are similar to samples with a brown groundmass glass in August deposits, but the clasts from September are lighter in color. August and September ejecta also contain greenish-gray andesite with a colorless groundmass glass.

Color differences among the groundmass glass types identified to date correlate with differences in the composition of the glass. Compositionally, the brown microlite-rich glass from the June and August eruptions is andesitic (60–62 weight percent SiO₂) (fig. 1, table 2). Small pumice lapilli on the surface of the June deposits contain brown andesitic (60–62 weight percent SiO₂) (the same as that in the large brown-black blocks) and dacitic (65–71 weight percent SiO₂) glass and colorless glass that is dacitic (68–73 weight percent SiO₂) and rhyolitic (75–79 weight percent SiO₂) in composition. The boundaries between the two glasses are sharp. One sample from the August eruption has a light-brown microlite-rich glass that is dacitic (approximately 65 weight percent SiO₂).

Brown-black andesite from the September eruption is lighter in color and contains lighter, more silicic groundmass glass relative to that of the June and August brown-black andesite. September glasses are mostly dacitic (63–68 weight percent SiO₂) (fig. 1, table 2) but show a wide compositional range among and within samples. This change in dominant glass composition was also observed in analyses of distal tephra fall (Swanson and others, this volume).

One sample from September (Stop 35, table 1) contains all three colors of glass (brown, light brown, and colorless) with diffuse contacts between the three glasses; some of the brown glass has an andesitic composition similar to that of June and August samples (table 2), and some is more similar to the light-brown glass in the sample. Both the light-brown glass and colorless glass are dacitic, but the light-brown glass

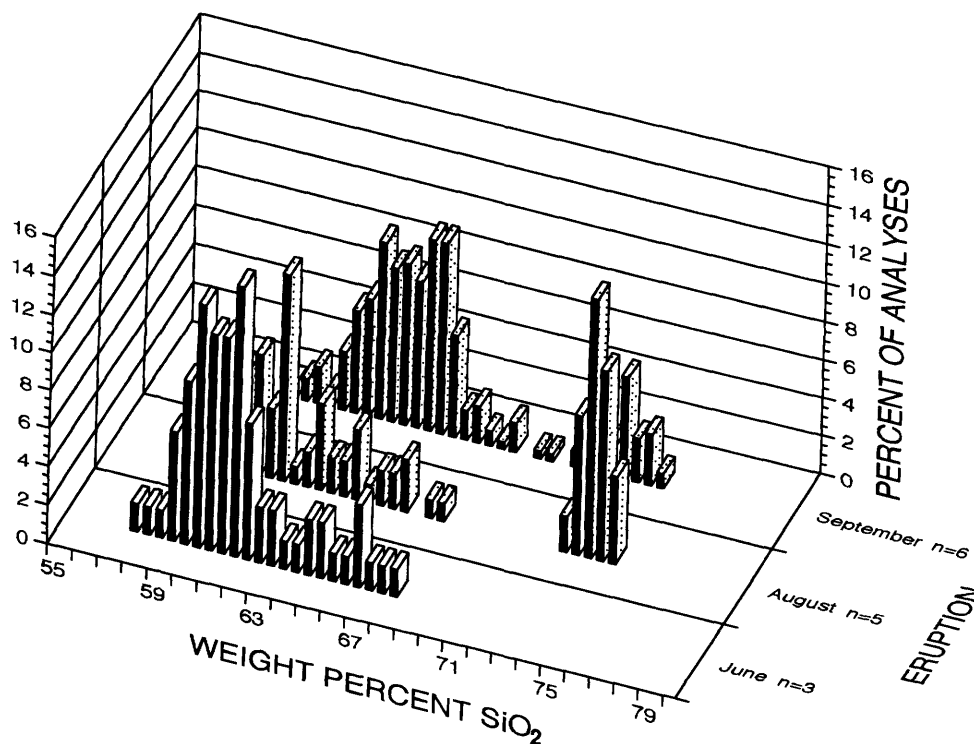


Figure 1. Histogram of groundmass glass compositions analyzed for each of three eruptions of Mount Spurr volcano, Alaska. June and August brown glasses are nearly identical, but the peak is shifted toward more silicic compositions in September. The August and September colorless glasses are also nearly identical. "Percent of Analyses" is the percentage of total analyses of a given weight percent SiO_2 . The total number of analyses varies for each eruption, 71 for June samples, 111 for August, and 261 for September. n, number of samples analyzed from each eruption.

Table 2. Average groundmass glass compositions for andesite from 1992 eruptions of Mount Spurr volcano, Alaska.

[*, total Fe as Fe_2O_3 ; n, number of analyses; †, sample from June eruption; ‡, sample from August eruption; §, sample from September eruption]

	Andesite glass				Dacite glass			Rhyolite glass	
	92JBS01†	92JBS03A†	92AMm111B‡	Stop 35§	92AMm102GS‡	92AMm129a§	92AMm120B§	92AMm113C‡	92AMm120A§
SiO_2	60.39±1.63	61.34±0.81	60.98±0.89	61.65±0.80	65.24±2.04	63.27±1.36	67.08±1.61	74.28±0.54	73.46±1.27
TiO_2	0.67±0.12	0.66±0.18	0.65±0.18	0.50±0.07	0.48±0.15	0.56±0.21	0.60±0.18	0.39±0.16	0.26±0.12
Al_2O_3	16.24±0.60	15.95±0.71	15.44±1.11	16.2±0.41	15.76±1.53	15.86±0.93	14.61±1.07	12.21±0.14	13.41±1.01
Fe_2O_3^*	6.92±0.69	7.64±0.63	7.34±0.90	7.03±1.20	5.24±1.03	6.09±0.90	5.34±0.82	2.68±0.21	2.49±0.40
MgO	2.14±0.43	2.33±0.67	2.80±0.94	2.26±0.32	1.21±0.70	1.75±0.46	0.94±0.45	0.39±0.03	0.26±0.04
CaO	5.61±0.31	5.73±0.33	5.34±0.48	5.09±0.35	4.35±0.92	4.50±0.55	3.67±0.64	1.36±0.19	1.57±0.67
Na_2O	4.38±0.17	4.91±0.25	4.54±0.29	4.45±0.19	4.75±0.37	4.50±0.26	4.85±0.29	3.72±0.15	4.41±0.26
K_2O	0.67±0.12	0.66±0.18	0.65±0.18	0.50±0.07	1.80±0.23	1.59±0.20	2.14±0.31	2.96±0.08	3.02±0.28
Cl	0.12±0.04	0.21±0.04	0.16±0.04	0.17±0.04	0.21±0.04	0.18±0.04	0.20±0.05	0.33±0.04	0.35±0.07
Total	97.81±1.83	99.97±0.53	98.65±0.69	98.90±1.36	99.04±0.95	98.31±1.30	99.42±0.51	98.31±0.56	99.23±0.62
n	14	44	24	7	22	38	32	46	44

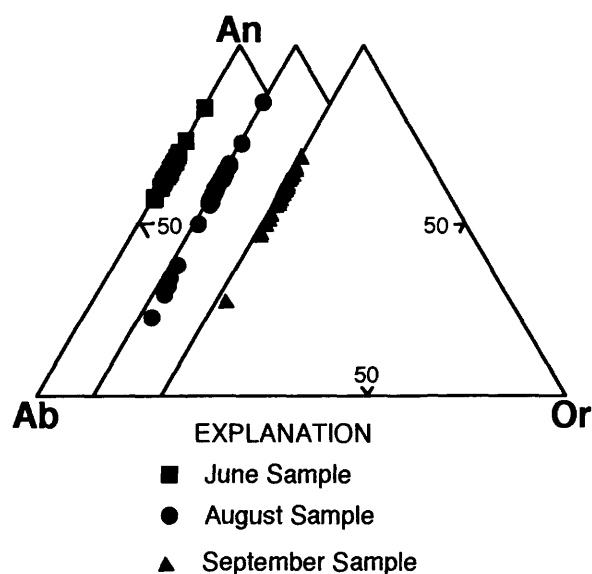
is slightly less Si rich. The three glasses also contain different microlite abundances. The brown glass is very rich in microlites, the light brown glass has abundant microlites, and the colorless glass is relatively free of microlites.

Blocks of greenish-gray andesite in August and September ejecta have a colorless groundmass glass that is rhyolitic (73–75 weight percent SiO_2) and peraluminous (table 2). The sample from the August eruption contains no microlites in the groundmass glass, but the sample from the September eruption has abundant microlites.

MINERALOGY

PLAGIOCLASE

The modal percentage of plagioclase ranges from approximately 17 to 31 percent (dense rock). It makes up 70 to 90 percent of the phenocrysts, and it typically occurs as euhedral crystals as much as 1 mm long with normal, reverse, oscillatory, and patchy zoning. Sieved (abundant glass inclusions) cores or zones parallel to grain boundaries are very common in phenocrysts.



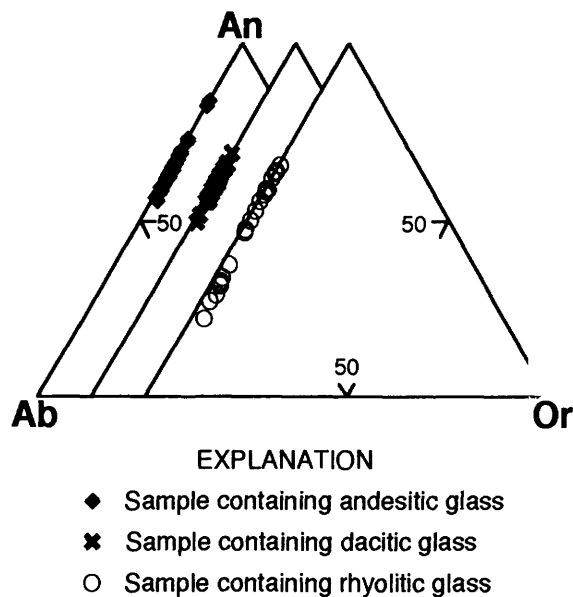
A

Figure 2. A, Range of plagioclase rim compositions of andesites for each of the 1992 Mount Spurr eruptions.

Rim compositions vary over the eruption sequence and correlate with the composition of the groundmass glass; rhyolitic glass samples contain the more sodic compositions. In June samples, rims range from An_{82} to An_{56} , in August samples, they range from An_{74} to An_{22} , and in September samples, they range from An_{64} to An_{27} (fig. 2A, table 3). Samples from throughout the eruption series that contain similar groundmass glasses have similar compositional ranges. A range of An_{84} to An_{55} characterizes andesitic glass, An_{55} to An_{50} characterizes dacitic glass, and An_{75} to An_{22} characterizes rhyolitic glass (fig. 2B).

CLINOPYROXENE

Light-green, weakly pleochroic clinopyroxene phenocrysts make up less than 1 percent of the mode (dense rock) in both andesite types. Crystals are subhedral to euhedral and less than 0.5 mm in length. The restricted compositional range of En_{38-48} Fs_{14-19} Wo_{36-44} (rim compositions) (fig. 3, table 3) is nearly identical to that of prehistoric Mount Spurr clinopyroxene (Nye and Turner, 1990). There is no clear change in composition with changing groundmass glass composition nor any compositional change through the eruption series.



B

Figure 2. Continued. B, Range of plagioclase rim compositions for different groundmass glass compositions.

Table 3. Representative analyses of plagioclase and clinopyroxene of andesites from 1992 eruptions of Mount Spurr volcano, Alaska.

[*, total Fe as Fe₂O₃; pyroxene recalculation scheme from Lindsley (1983) and represents rim compositions; FeO_t, total measured iron as FeO; Mg#, Mg/(Mg+Fe)•100; †, sample from June eruption; ‡, sample from August eruption; §, sample from September eruption]

Plagioclase

	<u>92JBS01†</u>	<u>92AMm111B‡</u>	<u>92AMm102GS‡</u>	<u>92AMm113C‡</u>	<u>92AMm129a§</u>	<u>92AMm120A§</u>
SiO ₂	49.39	50.34	46.89	58.90	46.84	53.88
Al ₂ O ₃	32.27	31.29	33.98	26.04	33.69	28.52
Fe ₂ O ₃ *	0.52	0.66	0.50	0.19	0.29	0.56
CaO	14.92	13.44	16.45	7.19	16.54	10.97
Na ₂ O	2.86	3.44	2.15	6.93	2.15	5.20
K ₂ O	0.03	0.01	0.02	0.34	0.04	0.08
Total	99.99	99.18	99.99	99.60	99.57	99.21
An	74.11	68.25	80.76	35.68	80.75	53.61
Ab	25.69	31.64	19.14	62.25	19.05	45.99
Or	0.20	0.10	0.10	2.07	0.20	0.40

Clinopyroxene

	<u>92JBS03A†</u>	<u>92CNS03B†</u>	<u>92AMm111B‡</u>	<u>92AMm102GS‡</u>	<u>92AMm113C‡</u>	<u>92AMm120A§</u>
SiO ₂	50.33	51.89	52.41	52.17	51.73	52.11
TiO ₂	0.54	0.38	0.37	0.44	0.44	0.35
Al ₂ O ₃	3.28	2.00	1.95	2.40	2.42	1.70
Fe ₂ O ₃	2.79	1.63	1.04	0.55	1.37	1.51
FeO	6.74	7.40	8.70	8.47	7.56	8.23
MnO	0.41	0.43	0.28	0.32	0.30	0.42
MgO	15.72	16.30	16.97	15.72	15.36	16.32
CaO	20.03	19.78	18.24	19.79	20.66	19.27
Na ₂ O	0.33	0.26	0.27	0.26	0.35	0.26
Total	100.17	100.07	100.23	100.12	100.19	100.17
FeO _t	9.25	8.87	9.64	8.96	8.79	9.59
En	44.23	45.61	47.59	44.71	43.48	45.58
Fs	15.25	14.60	15.62	14.82	14.45	15.70
Wo	40.52	39.79	36.79	40.47	42.07	38.71
Mg#	74.36	75.75	75.29	75.11	75.06	74.38

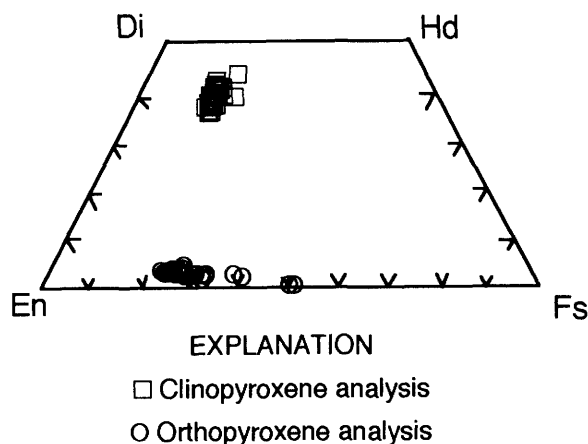


Figure 3. Representative range of clinopyroxene and orthopyroxene rim compositions of andesite samples from three eruptions of Mount Spurr, Alaska, in 1992. Di, diopside; Hd, hedenbergite; En, enstatite; Fs, forsterite.

Orthopyroxene phenocrysts are the most abundant mafic phase in all andesite samples, typically 3 to 7 percent of the mode (dense rock). Crystals are subhedral to euhedral and as much as 1.2 mm long, pale greenish-brown, and weakly to moderately pleochroic. Some crystals have zones (typically cores) that are more strongly colored. The abundance of such "cored" crystals appears to increase in samples from each succeeding eruption. The wide compositional range of En₄₉₋₇₄ Fs₂₂₋₅₁ Wo_{0.3-5} (rim compositions) (fig. 3, table 4) is similar to that for prehistoric orthopyroxene (Nye and Turner, 1990). Unlike clinopyroxene, the orthopyroxene shows a compositional change with groundmass glass composition. Orthopyroxene phenocrysts in andesitic and dacitic glass samples have a similar range of En₆₁₋₇₄ Fs₂₂₋₃₇ Wo_{0.9-5}, but rhyolitic

glass samples contain orthopyroxene with more iron-rich compositions ($\text{En}_{49-73} \text{Fs}_{24-51} \text{Wo}_{0.3-5}$).

AMPHIBOLE

Pargasitic amphibole phenocrysts (Hawthorne, 1981) form less than 2 percent of the mode (dense rock) in all samples. Crystals are euhedral to anhedral, less than 2 mm in largest dimension, and pleochroic in shades of green and brown. Coronas of plagioclase, orthopyroxene, and FeTi oxides surround amphibole in some samples and presumably record disequilibrium. Well-developed coronas are dominant in samples with

a colorless rhyolitic groundmass glass, whereas samples containing brown groundmass glass typically lack evidence of disequilibrium. However, amphibole both with and without coronas are found in some samples. Compositionally, the amphibole is high in Al ($\text{Al(IV)} = 1.35-2$, based on 23 oxygens) and alkalis ($\text{Na} + \text{K}$ in A site = $0.6-0.85$) (table 4), similar to the range for prehistoric Mount Spurr amphibole (Nye and Turner, 1990). Pargasitic Mount Spurr amphibole falls at the end of the Aleutian amphibole trend (Kay and Kay, 1985) and is similar to other eastern Aleutian pargasitic amphibole (from Augustine Volcano (Daley, 1986) and Redoubt Volcano (Swanson and others, 1994)) (fig. 4).

Table 4. Representative analyses of orthopyroxene and hornblende of andesite from 1992 eruptions of Mount Spurr volcano, Alaska.

[Pyroxene recalculation scheme from Lindsley (1983) and represents rim compositions; FeO_t , total measured iron as FeO; Mg#, $\text{Mg}/(\text{Mg}+\text{Fe})\cdot 100$; †, sample from June eruption; ‡, sample from August eruption; §, sample from September eruption]

Orthopyroxene

	<u>92JBS03A[†]</u>	<u>92CNS03B[†]</u>	<u>92AMm111B[‡]</u>	<u>92AMm102GS[‡]</u>	<u>92AMm113C[‡]</u>	<u>92AMm120A[§]</u>
SiO ₂	53.21	54.21	53.17	53.04	51.31	52.01
TiO ₂	0.19	0.12	0.16	0.26	0.19	0.16
Al ₂ O ₃	1.66	0.96	1.09	1.90	1.25	2.53
Fe ₂ O ₃	1.87	1.33	1.47	1.32	0.18	2.00
FeO	13.82	13.86	16.62	14.79	25.39	15.69
MnO	0.63	0.47	0.69	0.52	0.77	0.51
MgO	26.58	27.54	25.54	26.29	19.24	25.32
CaO	2.11	1.49	1.20	1.62	0.87	1.49
Na ₂ O	0.03	0.06	0.00	0.00	0.01	0.00
Total	100.10	100.04	99.94	99.74	99.21	99.71
FeO _t	15.50	15.06	17.94	15.98	25.55	17.49
En	71.54	73.78	69.29	71.61	55.54	69.37
Fs	24.38	23.35	28.37	25.23	42.65	27.69
Wo	4.08	2.87	2.34	3.16	1.81	2.94
Mg#	74.58	75.96	70.95	73.95	56.57	71.47

Hornblende

	<u>92CNS03B[†]</u>	<u>92AMm111B[‡]</u>	<u>92AMm102GS[‡]</u>	<u>92AMm113C[‡]</u>	<u>92AMm120B[§]</u>
SiO ₂	42.84	42.94	42.87	42.64	43.36
TiO ₂	2.14	2.39	2.70	2.61	2.38
Al ₂ O ₃	13.51	12.74	11.23	11.92	12.45
FeO _t	10.64	12.69	12.60	12.85	12.09
MnO	0.07	0.11	0.27	0.15	0.30
MgO	15.01	13.34	14.26	13.53	13.91
CaO	11.10	10.75	10.70	11.01	10.77
Na ₂ O	2.36	2.30	2.45	2.42	2.28
K ₂ O	0.26	0.38	0.35	0.35	0.31
Cl	0.04	0.08	0.05	0.06	0.03
F	0.14	0.23	0.37	0.37	0.05
H ₂ O	1.99	1.90	1.83	1.83	2.02
F=O	-0.06	-0.10	-0.16	-0.16	-0.02
Cl=O	-0.01	-0.02	-0.01	-0.01	-0.01
Total	100.02	99.73	99.52	99.56	99.92

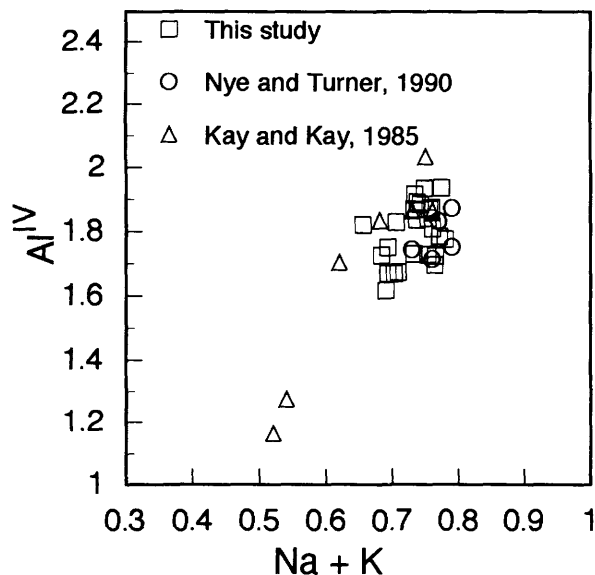


Figure 4. Al(IV) versus Na + K for hornblende of this study and hornblende from Mount Spurr and other Aleutian volcanoes from Nye and Turner (1990) and Kay and Kay (1985).

OPAQUE MINERALS

Opaque minerals are dominantly magnetite, but include a trace amount of pyrrhotite, and they may make up as much as 5 percent of the mode (dense rock). Magnetite occurs as anhedral to subhedral crystals less than 0.2 mm in maximum dimension. Pyrrhotite occurs as anhedral masses as much as 0.5 mm in maximum dimension.

CRYSTAL CLOTS

Small (≤ 2 mm) crystal clots of different minerals are found in all samples (1–2 percent of the mode). The most common clots are (1) plagioclase and magnetite; (2) orthopyroxene and magnetite; (3) plagioclase, orthopyroxene, and magnetite; and (4) poikilitic hornblende and plagioclase. Clinopyroxene is also present in some clots. The pyroxene crystals are subhedral to anhedral, whereas the plagioclase is subhedral to euhedral and typically somewhat sieved.

XENOLITHS

Metamorphic inclusions are present in many samples. The inclusions are of three types: (1) plagioclase and spinel \pm FeTi oxides, (2) polycrystalline quartz, and (3) plagioclase, quartz, sillimanite, spinel, FeTi oxides, pyrrhotite \pm biotite \pm cordierite \pm corundum. Grains of cordierite are present in some andesites and

are probably derived from xenoliths. Nye and Turner (1990) also described plagioclase-spinel clots in Mount Spurr lavas, which they inferred to be altered anorthosite. The ubiquitous nature of spinel in the large metamorphic blocks suggests that they are of metamorphic origin. Xenoliths of older volcanic material are also present in some samples.

DISCUSSION

The brown-black and greenish-gray andesites are petrographically and geochemically similar (Nye and others, this volume); they differ only in their groundmass glass compositions. Groundmass glass compositions in the brown-black andesite from the June and August eruptions are nearly identical. However, in the samples from the September eruption the dominant glass is more silicic and the compositions shift relative to those of samples from the June and August eruptions (fig. 1). This shift in the dominant melt composition from August to September could possibly be the result of progressive crystallization. Presumably this crystallization occurred in the magma body throughout the eruption sequence. Thus, the shift in melt composition suggests the presence of a single magma body undergoing progressive crystallization and tapped for three eruptions. The presence of a single magma body is supported by the geochemistry of the andesite (Nye and others, this volume).

The rhyolitic groundmass glasses from the August and September eruptions also show identical compositional ranges (fig. 1). This observation suggests that the same process generated andesite with rhyolitic groundmass glass in August and September.

Differences in the groundmass glass compositions also result in differences in the range of plagioclase and orthopyroxene compositions. The brown-black andesite with both andesitic and dacitic glasses has plagioclase rim compositions from An_{84} to An_{56} , whereas in the greenish-gray andesite with rhyolitic glass rim compositions are very sodic and extend to An_{22} . The orthopyroxene in samples with andesitic and dacitic glass has rim compositions of En_{61-74} and in the rhyolitic glass samples, En_{49-73} .

There are three possible mechanisms to explain the origin of the various glasses and the resulting sympathetic plagioclase and orthopyroxene compositional differences. These mechanisms are fractional crystallization, assimilation, and magma mixing. Incorporation of more felsic material either through melting of continental crust (assimilation) or magma mixing could cause changes in the melt and crystallizing phases. However, the similarity of andesite geochemistry argues for a closed system (Nye and others, this volume) during the eruption sequence. Therefore, the more

likely explanation is fractional crystallization. Groundmass glass major-element variation diagrams (fig. 5) show predominantly linear trends between the andesitic and rhyolitic glasses, with the exception of Na_2O and MgO . Na_2O vector modeling of plagioclase crystallization indicates it would be possible to drive glass compositions from the dacitic glass to rhyolitic glass through crystallization of plagioclase with compositions of approximately An_{35-45} .

Hornblende phenocrysts, although identical in composition, are stable in the andesitic glasses and to a degree in the dacitic glasses, but they are clearly unstable in the rhyolitic glasses. There are three possible explanations for this: differences in water content of the different glasses (melts); the peraluminous nature of the rhyolitic melts; and reaction kinetics (Rutherford and Hill, 1993). It is unclear which of these three mechanisms is the most likely explanation.

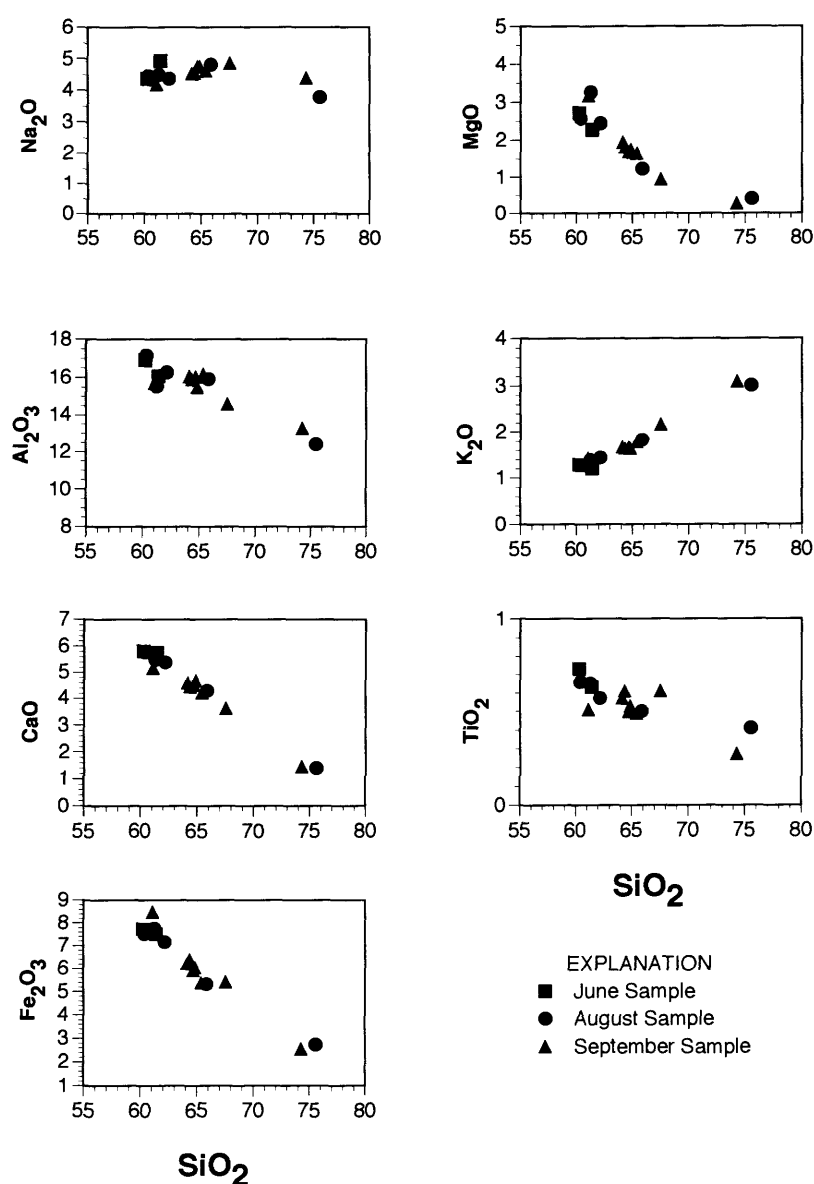


Figure 5. Average groundmass glass compositions (in weight percent) of major oxides plotted against weight percent SiO_2 . Glass from andesite samples collected following three eruptions of Mount Spurr, Alaska, in 1992.

SUMMARY

The 1992 eruptions of Mount Spurr provided several types of products for examination, including two types of andesite. Derivation of the greenish-gray andesite with rhyolitic glass appears to have been accomplished through fractional crystallization of brown andesite with andesitic glass. Fractional crystallization is supported by the compositional changes observed in plagioclase and orthopyroxene. Examination of the andesite groundmass glasses suggests that a single magma body undergoing progressive crystallization was present. It seems likely that progressive crystallization took place as microphenocrysts and microlites. Work will be done on groundmass crystallinity to determine any differences.

The role of the metamorphic xenoliths in the Crater Peak magmatic system is not understood. The system was apparently closed during the eruption sequence so any affect on the magma must have occurred prior to the June eruption. Nye and Turner (1990) suggested that crustal sweats have influenced the Mount Spurr magmatic system. The presence of the xenoliths suggests that they are the origin of these crustal sweats. Future work will involve characterizing the xenoliths to determine their role.

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Seismicity and Forecasting of the 1992 Eruptions of Crater Peak Vent, Mount Spurr Volcano, Alaska: An Overview

By John A. Power, Arthur D. Jolly, Robert A. Page, and Stephen R. McNutt

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ABSTRACT

Seismicity associated with the 1992 eruptions of the Crater Peak vent of Mount Spurr was monitored by a network of 6 to 10 seismometers within 15 km of the mountain's summit. Precursory seismicity began in August of 1991. It was marked by a conspicuous swarm of shallow volcano-tectonic (VT) earthquakes beneath Crater Peak, the first such swarm in a decade of monitoring, and by the gradual onset of seismicity between depths of 5 and 40 km, in a volume of rock that had been essentially aseismic for nearly a decade. The volcano erupted three times during 1992: on June 27, August 18, and September 16–17. The June eruption was preceded by about 14 hours of continuous tremor and a vigorous 4-hour swarm of shallow VT earthquakes. Following the June eruption, shallow seismicity decreased to normal background levels, and the August eruption occurred without any detectable precursory seismicity. A few hours of shallow seismicity preceded the September eruption, after which seismicity in the depth range of 5 to 40 km peaked. From an initial analysis of the seismic data, we suggest the 1992 eruption sequence resulted from the in-

trusion of magma into the mid-crust that began in mid-1991. The magma had migrated to shallower depths by early June of 1992, when the distribution of hypocenters changed and volcanic tremor began. The character of seismograms suggests the deep seismicity included a variety of source processes ranging from brittle failure of the country rock in response to injection and withdrawal of magma to the vibration of fluid-filled cracks associated with magma transport.

Close monitoring of Mount Spurr seismicity allowed the Alaska Volcano Observatory to issue public statements describing the increased activity beginning in August of 1991. Forecasts were issued prior to the June 27 and September 16–17 eruptions, and notifications were issued shortly after all three events. Two additional forecasts were issued on the basis of observed increases in seismicity that did not culminate in eruptions.

INTRODUCTION

During the summer of 1992, Crater Peak erupted three times—on June 27, August 18, and September 16–17. Each of the eruptions lasted 3.5 to 4 hours and produced large tephra plumes, which drifted downwind across large parts of Alaska (Neal and others, this volume).

Rudimentary seismic monitoring of Mount Spurr volcano began in 1971, and since 1989 the Alaska Volcano Observatory (AVO) has operated a 6- to 10-station seismic array (herein referred to as the Spurr network) centered on the volcano. The reawakening of Crater Peak was preceded by roughly 10 months of elevated earthquake activity that spanned the caldera and extended to depths of 40 km. The seismicity associated with the eruptions constitutes a complex sequence of events including shallow volcano-tectonic (VT) earthquakes, long-period (LP) events, volcanic tremor, and periods of seismic quiescence (see "Instrumentation and Data Analysis" section for definitions).

In this paper we describe the seismicity at Mount Spurr preceding and accompanying the 1992 eruption sequence. First we review the seismic instrumentation, data acquisition, and data analysis. We then briefly summarize the seismic data from the Mount Spurr area beginning in 1981 as well as develop a more detailed seismic chronology for the period between August of 1991 and December of 1992. This chronology is based primarily on earthquake hypocenter plots and helicorder records, and to a lesser degree on Real-time Seismic Amplitude Measurements (RSAM) (Endo and Murray, 1991) and Seismic Spectral Amplitude Measurements (SSAM) (Stephens and others, 1994). We offer an initial volcanological interpretation of the patterns of seismicity, and we conclude with a review of the role that seismic observations played in forecasting the 1992 eruptions.

ACKNOWLEDGMENTS

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eruptions. Gail Davidson provided valuable assistance with data acquisition and computer management. The success in forecasting the Crater Peak eruptions is shared with our colleagues, most notably John Lahr, Bernard Chouet, Scott Stihler, Christopher Stephens, Guy Tytgat, and John Davies. We thank Elliot Endo and Randy White for their reviews of the manuscript and figures.

INSTRUMENTATION AND DATA ANALYSIS

Seismic monitoring on Mount Spurr volcano has a relatively long history. Station SPU was installed roughly 14 km from the summit of Mount Spurr in August of 1971 (fig. 1). The addition of stations CRP and CGL in the fall of 1981 allowed small earthquakes near Mount Spurr to be located. Station CGL was not maintained between late 1985 and early 1989, so only two stations were operating near the volcano. Following the establishment of AVO, three new stations, CKL, BGL, and NCG, were added to the seismic network in late 1989, and stations CKN and CPK were in-

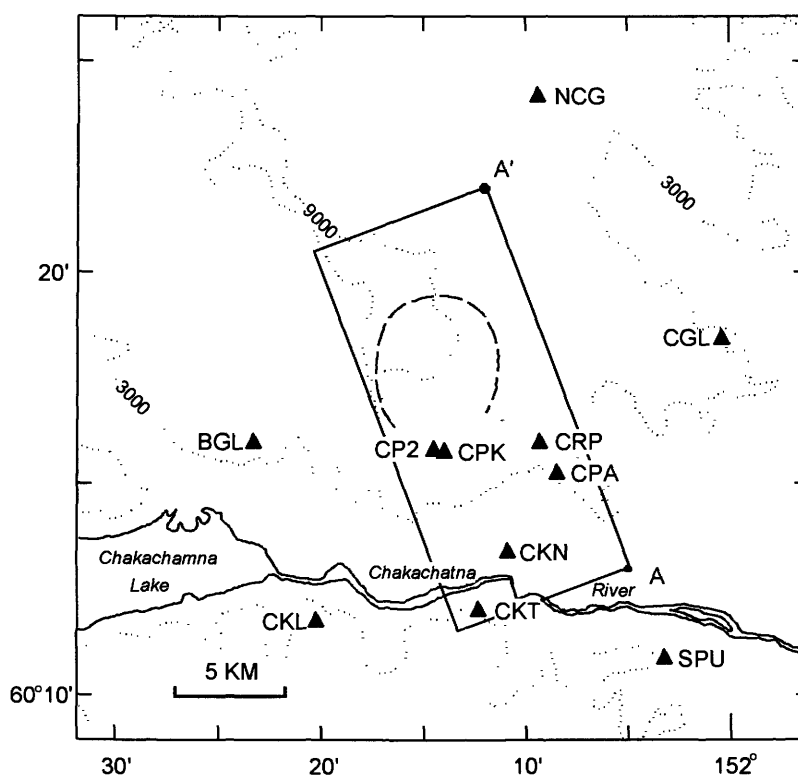


Figure 1. Seismic stations near Mount Spurr volcano, Alaska. Dotted lines represent 3,000- and 9,000- foot contours, dashed line represents approximate location of caldera rim. Box and line A-A' denote spatial extent of hypocenters shown in figure 4, and the location of projected plane for cross sections in figure 3.

stalled in the fall of 1991. Station CPK was destroyed during the June 27, 1992, eruption. A new instrument was installed at the same location on September 10, and it was subsequently destroyed during the September 16–17 eruption. Station CPK was replaced by station CP2 on October 23. Stations CKT and CPA were added to the network on September 20 and October 29, respectively, which brought the total number of stations in the seismic network to ten. All stations were short-period single-component vertical seismometers except CPA and the original CPK, each of which had horizontal components. The original CPK also had a low-gain vertical component.

Between 1981 and 1989, data from the Spurr seismic network was recorded on 16-mm photographic film. Starting in July of 1988 these data were digitally recorded on a computer system at the AVO office in Fairbanks (AVO staff, 1990). In October of 1989, we also began recording the Spurr network on a personal computer using the program MDETECT, which digitizes and records 16 channels of event-detected data (Lee and others, 1989). This acquisition system provided the primary data set used in this study. Not all the stations from the Spurr network were recorded on the personal computer system during 1990 (fig. 2) because stations monitoring the ongoing eruptions of Redoubt Volcano were given higher priority for available channels on the system (Power and others, 1994). Beginning January 1, 1991, a complete suite of Spurr stations was established on this system and the triggering parameters were held constant through February 1993. Stations in the Spurr network were also monitored by an RSAM (Endo and Murray, 1991) and an SSAM (Stephens and others, 1994) system. March and Power (1990) described the various computer systems and their networking.

Event-detected waveforms were transferred to a microcomputer and phase arrivals were determined using the program XPICK (Robinson, 1992). Hypocenters and local magnitudes were determined using the program HYPOELLIPSE (Lahr, 1989) with a flat-layered velocity model and station corrections to account for local variations in seismic-wave velocities. The model we used was developed specifically for the Mount Spurr area by Jolly and others (1994). In this model, earthquake depths are permitted to occur as much as 3.2 km above sea level; the elevation of the summit of Mount Spurr. Consequently, depths are referenced to sea level; negative depths refer to height above sea level.

In describing the various waveforms and event types at Mount Spurr volcano, we use terminology and event classifications similar to that developed by Chouet and others (1994), Lahr and others, (1994), and Power and others, (1994) during the 1989–90 eruptions of Redoubt Volcano. This classification system

is based on our present understanding of the physical processes associated with various seismic sources. Volcano-tectonic earthquakes refer to sources representing purely elastic processes. Long-period events are thought to represent a more complex process in which fluids (gas and liquid phases) play an active role in the generation of seismic waves (Chouet, 1992). Volcano-tectonic and LP events represent end members of a spectrum; events combining these two characteristics are referred to as hybrid events. We do not identify hybrid events in this paper because identification of their spectral characteristics requires additional analysis, which has not yet been attempted (see Lahr and others, 1994). The seismic events associated with the forcible ejection of gas and ash from the volcano are referred to here as explosive eruptions.

SEISMIC CHRONOLOGY

OCTOBER 1981 THROUGH JULY 1991

Mount Spurr and the surrounding area has been seismically active since local monitoring began in 1981. During periods between October 1981 and August 1991 when three or more stations were operating in the Spurr network, an average of 8.5 earthquakes per month greater than magnitude (M_L) 0.0 were located within 20 km of the volcano (fig. 2). Most of this activity concentrates beneath the summit at depths shallower than 3 km. Very few shocks were located deeper than 5 km (fig. 3A). The largest located event in the Spurr area is a magnitude 3.3, which occurred on November 16, 1989, at a depth of 16.5 km about 12 km northwest of the summit. Most located events range between magnitudes -0.5 and 1.5.

A second concentration of epicenters lies on the north rim of the Mount Spurr caldera. Activity in this area occurred in two clusters; one is about 2 km north of the main summit cluster, and the other is roughly 3 km northeast of the main summit cluster (fig. 3A). The one northeast of the summit is the site of two small swarms of shallow earthquakes that occurred in early 1982 and in early 1989 (fig. 4A). Hypocenters could not be calculated for the 1989 swarm, because only two stations were operating at the time. On the basis of similarities in waveforms between the two swarms, however, we surmise that the 1989 swarm occurred in the same location as the 1981 swarm. A more detailed description and analysis of the seismicity during this period is given by Jolly and others (1994).

AUGUST 1991 THROUGH JUNE 27, 1992

The first sign of reawakening at Mount Spurr volcano was a small swarm of VT earthquakes be-

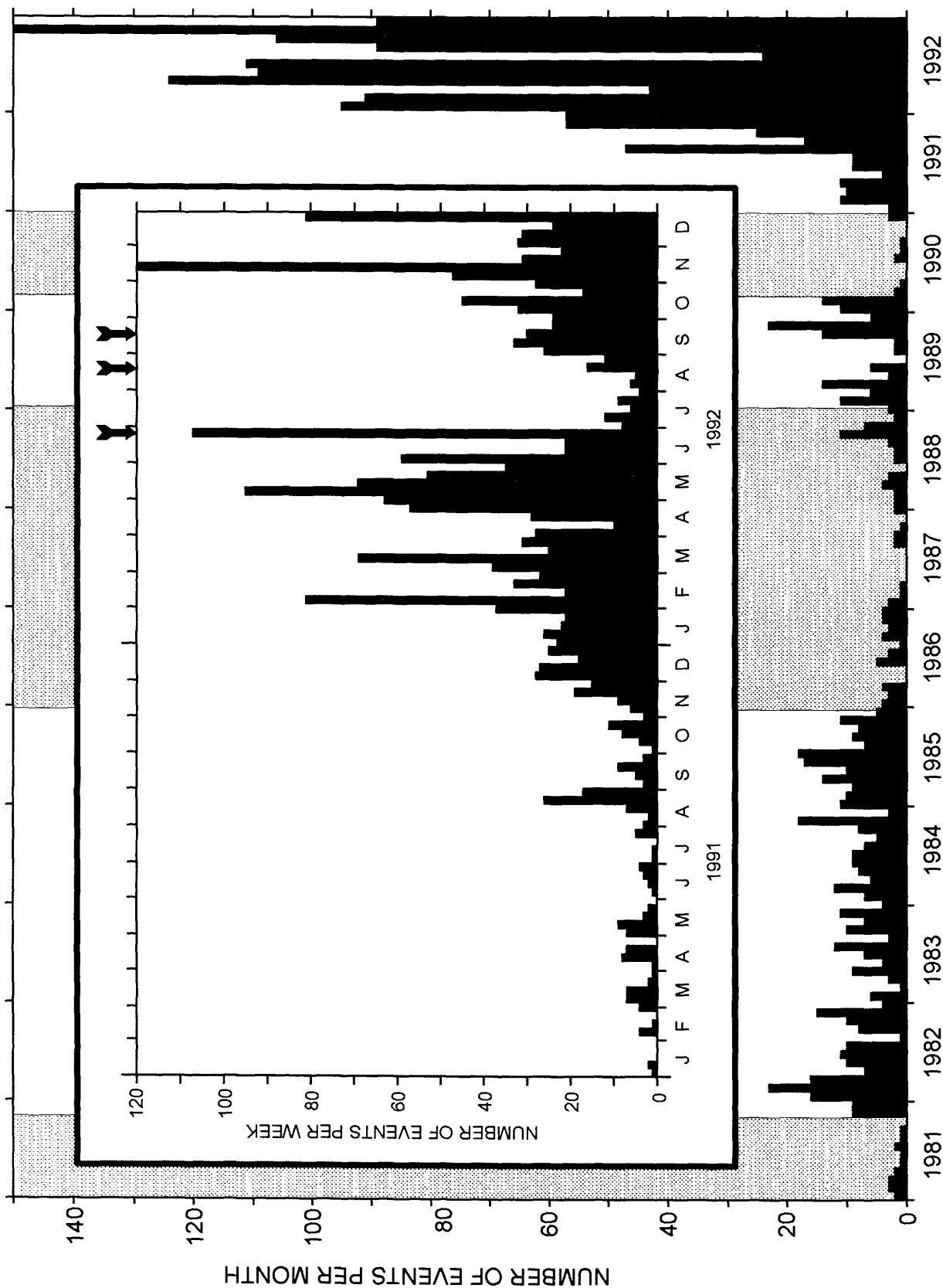


Figure 2. Histogram of monthly number of earthquakes greater than M_L 0.0 located within 20 km of the summit of Mount Spurr volcano between 1981 and 1992. Shaded pattern in 1981 and between late 1985 and early 1990 reflects periods when fewer than three stations were operating in the Spurr network. Shaded pattern in 1990 denotes a period when not all stations in the Spurr network were recorded on the computerized data-acquisition system (see text for details). Inset shows weekly number of events located regardless of magnitude during 1991 and 1992. Arrows denote approximate times of eruptions. Note peaks in activity in February, March, and May, as well as relative quiescence following the June 27 eruption. Strong peaks in June and November correspond to the swarm immediately preceding the June 27 eruption and the November 9–10 swarm.

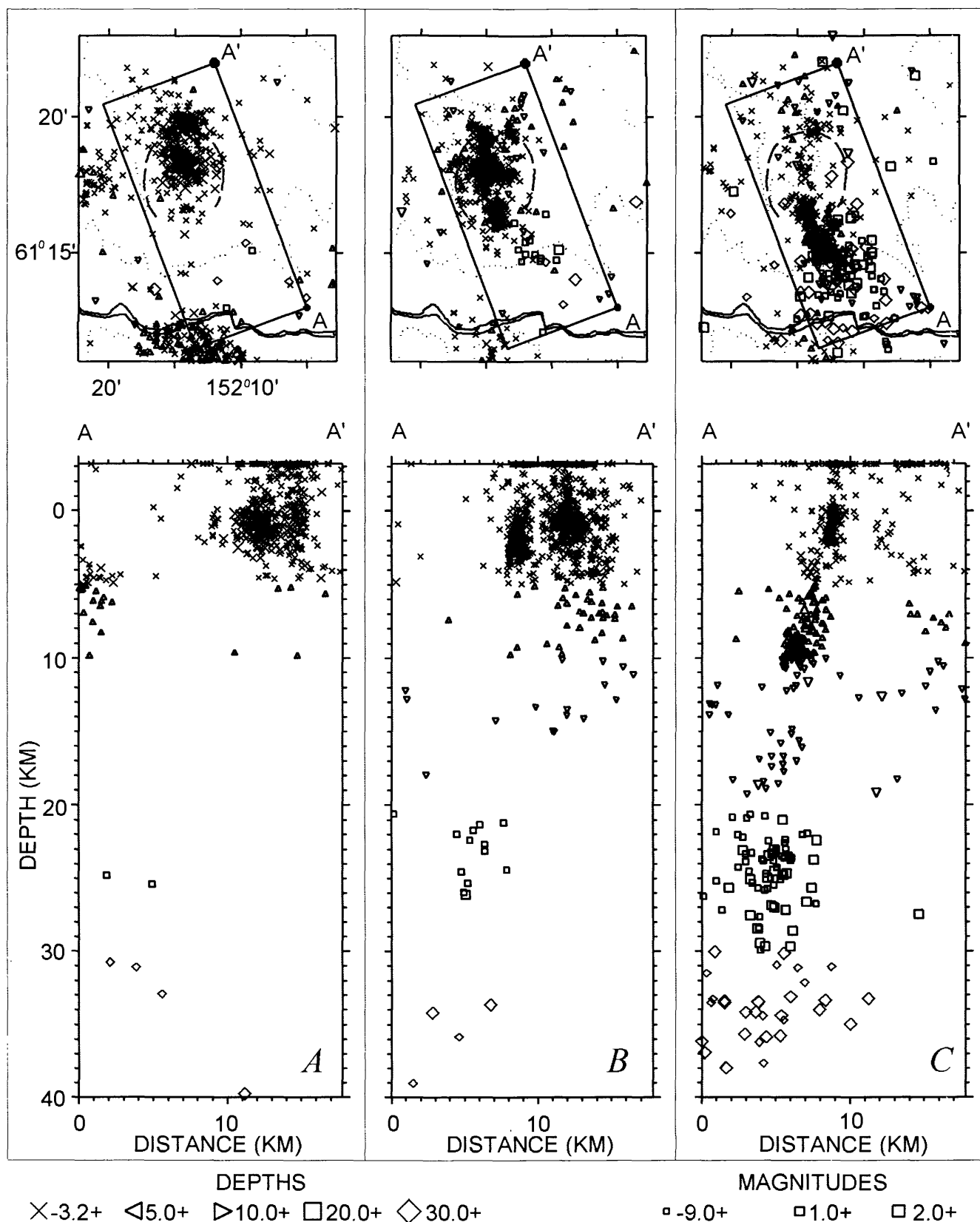


Figure 3. Epicenter maps and cross sections (A to A', figure 1) for three successive periods: A, background, October 1981 through July 1991; B, precursory, August 1991 through June 1992; and C, eruptive, July through December 31, 1992.

neath Crater Peak in August of 1991 (fig. 4B). Most of these shocks were less than M_L 1.0 and ranged in depth from 1 to 4 km. This swarm marked the first sustained seismic activity observed beneath Crater Peak since earthquake hypocenters could be reliably determined in 1981 (fig. 4A).

The August activity beneath Crater Peak was followed by two months of quiescence, and then by continuous elevated activity punctuated by more intense periods of seismicity in February and May (fig. 4B).

The August 1991 Crater Peak swarm initiated an increase in shallow seismic activity, which spanned the Mount Spurr caldera. Except for the Crater Peak area, this increase was mostly confined to areas that had been historically active—the summit and the north caldera rim. The number of VT earthquakes beneath the summit of Mount Spurr began to increase in October of 1991 (fig. 4B). Activity beneath the summit then increased nearly continuously until early June of 1992. Hypocenters were generally concentrated between -1.0 and 5.0 km in depth, although several shocks were recorded as deep as 15 km (fig. 3A). The largest event during this period was a M_L 1.7 on February 18, 1992, at a depth of 0.3 km.

Activity beneath the north caldera rim was more discontinuous (fig. 4B). The first perceptible increase was a small group of earthquakes between 0 and 2 km depth in mid-December. More vigorous swarms occurred in March, April, and May of 1992. These events were concentrated near the two clusters of events from the previous decade (compare figs. 3A and 3B). The distribution of depths in the north cluster is more diffuse, generally ranging from -2.0 to 6.0 km. The largest event during this period beneath the north rim is a M_L 1.6, which occurred on May 3, 1992, at a depth of 0.3 km.

On June 5, 1992, the number of located earthquakes at Mount Spurr volcano reached 28, the highest number recorded in any 24-hour period. These shocks ranged in depth between -1.0 and 1.0 km and were clustered beneath Crater Peak. Coincident with this swarm, short bursts of volcanic tremor were recorded on the stations closest to Crater Peak for the first time (McNutt and others, this volume). Following the June 5 swarm, the number of locatable events declined throughout the entire caldera, most notably beneath the summit and Crater Peak (figs. 2 and 4B).

The other noticeable change in seismicity, which occurred between August of 1991 and June 27, was the onset of activity in the depth range of 10 to 40 km beneath and southeast of Crater Peak (figs. 3B, 4B). The first of these events was located in August of 1991, and more than 40 occurred before the June 1992 eruption. These events occurred in a volume that had been practically aseismic during the previous decade (compare fig. 3A and 3B). Waveforms from events

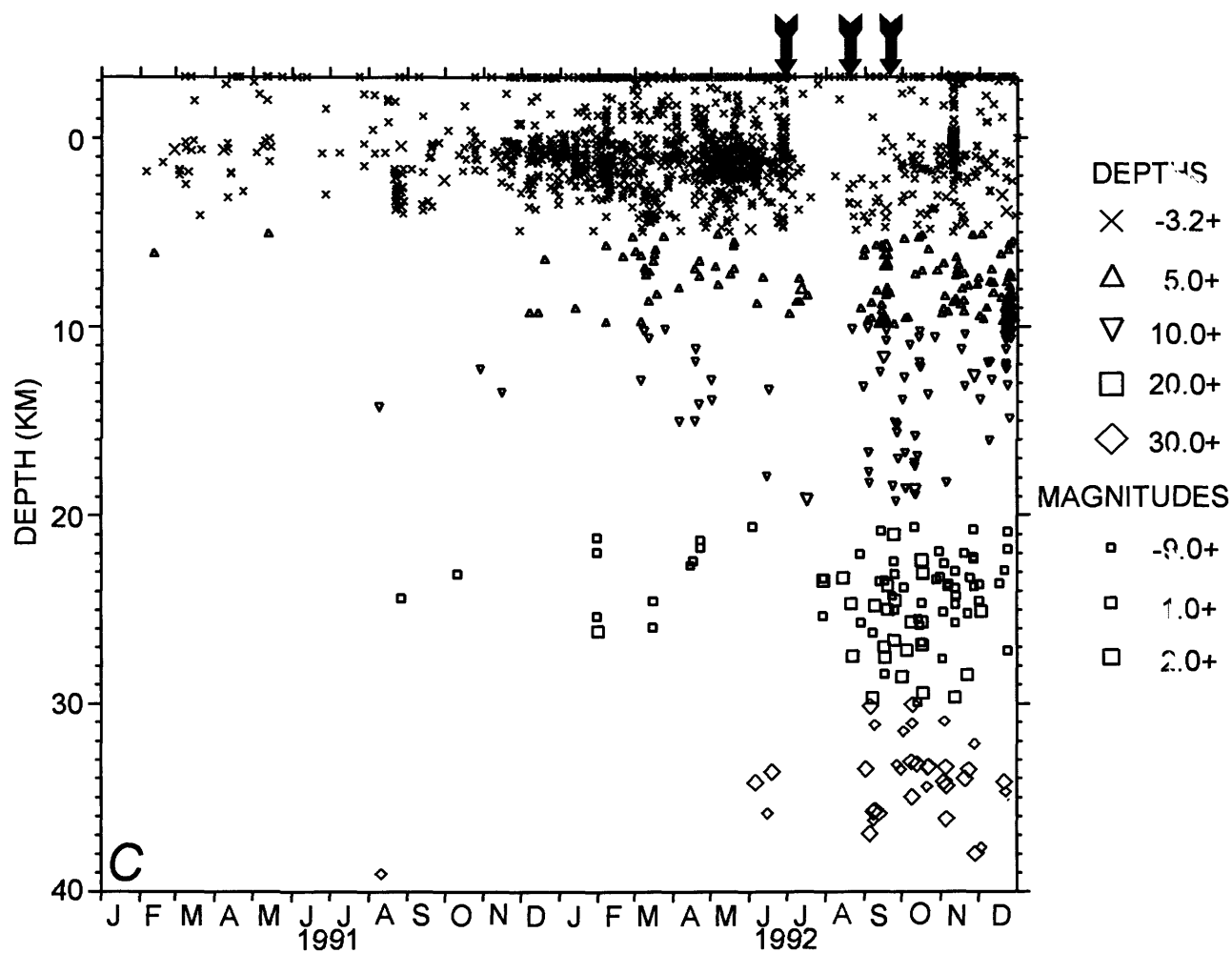
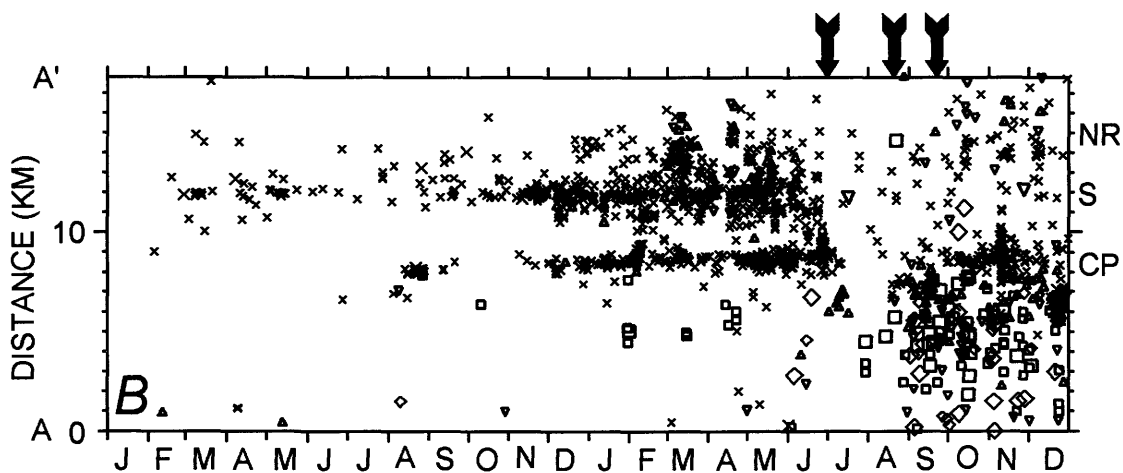
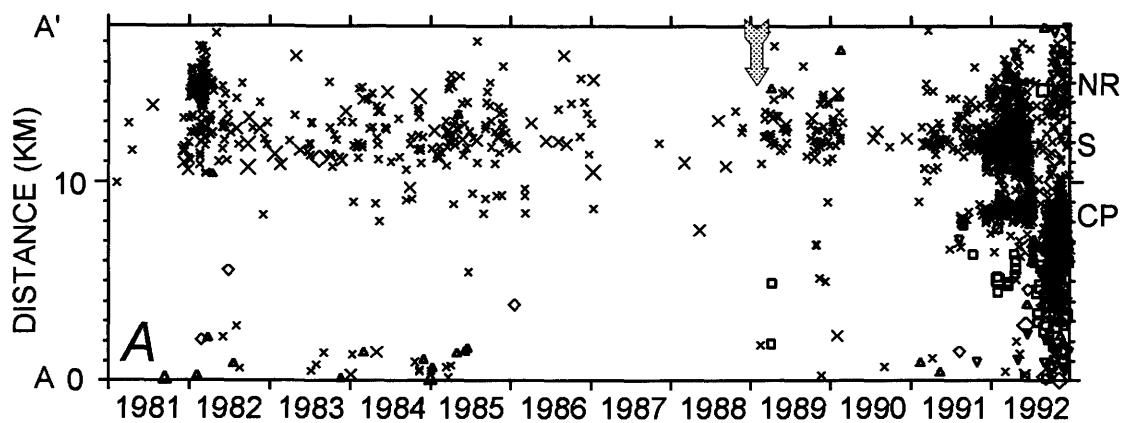
at depth display a broader range of spectral characteristics that range from high-frequency earthquakes, which appear to be volcano-tectonic to low-frequency events, which have a strong long-period character. Attenuation of the seismic waves from these deep events makes their classification difficult. Those events that we feel are volcano-tectonic generally have well-defined phases and a dominant frequency between 3 and 4 Hz, whereas low-frequency events lack well-defined phases, have peak frequencies of about 2 Hz, and have extended coda. The variability in observed waveforms for two events from roughly 33 km depth is shown in figure 5. A detailed spectral analysis of all of the events below 10 km has not yet been completed.

At 12:04 p.m. ADT on June 26 continuous volcanic tremor began, which was easily visible on all stations within 14 km of the volcano (McNutt and others, this volume). At roughly 3:00 a.m. ADT on June 27, just 4 hours before eruptive activity began, a vigorous swarm of shallow VT earthquakes began. These shocks were located between -1.0 and 1.0 km in depth and ranged in magnitude between -0.6 and 1.2. The magnitudes of events increased through the duration of the swarm. To date, only three LF events have been identified in this swarm. The eruption on June 27 lasted 4 hours and 3 minutes (McNutt and others, this volume). The June 27 eruption was followed by a short-lived shallow swarm of LP events, which lasted for about 24 hours.

JUNE 27 THROUGH SEPTEMBER 17, 1992

Following the June 27 eruption, the volcano entered a period of relative quiescence in which the shallow seismicity was comparable to or lower than the pre-August 1991 level (fig. 4B). In contrast, the deeper seismicity continued at a low, steady rate comparable to its August 1991 through June 1992 level. The area beneath the summit remained nearly aseismic throughout the remainder of 1992, and only low levels of activity occurred beneath the north caldera rim (figs. 3C, 4B).

Figure 4. Time-distance (A, B) and time-depth (C) plots for hypocenters within the box shown in figure 1. A and B are time distance plots along the line A-A' for periods 1981 through 1992 and 1991 through 1992, respectively. Shaded arrow in A shows approximate location of 1989 swarm. C, data from 1991–92. Solid arrows denote the three eruptions on June 27, August 18, and September 16–17. ➤



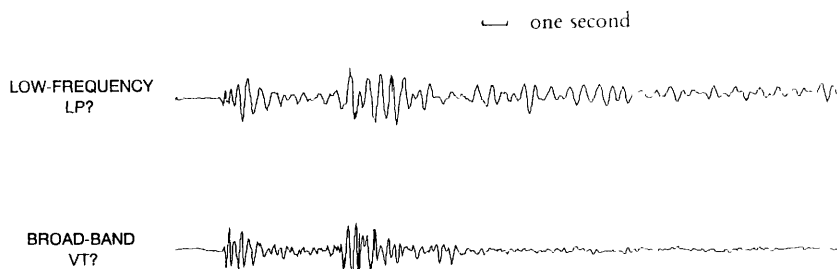


Figure 5. Comparative waveforms recorded at seismic station CGL on Mount Spurr volcano, Alaska, from two events located at a depth of approximately 33 km. These waveforms illustrate the diversity of signals observed in events at depths greater than 10 km. Note the monochromatic character and long duration of the low frequency event (upper trace). Such features are typical of long-period (LP) events.

The August 18 eruption occurred with no obvious seismic precursor at CRP, the closest station to the crater that was operating at the time. Eruptive activity on August 18 began with a small ash-laden plume at 3:37 p.m. ADT, which was accompanied by 16 minutes of weak tremor and several LP events. The main eruption began at 4:42 p.m. ADT and lasted about 3 hours and 28 minutes (McNutt and others, this volume).

Following the August 18 eruption, seismicity below depths of 2 km began to slowly increase, while that at shallower depths remained low (fig. 4C). Activity in the depth range of 2 to 10 km and below 20 km appears to have stabilized at a new, higher background level at about the time of the September eruption (fig. 4C). On September 16, shallow tremor, which was interspersed with small discrete events, began about 7:30 p.m. ADT. A small eruption, which lasted about 11 minutes, began at 10:36 p.m. ADT. Seismic activity remained elevated, and at 12:04 a.m. ADT on September 17 the main eruption began. The intense seismicity associated with this eruption lasted 3 hours and 36 minutes (McNutt and others, this volume). The waning phase of the eruption included a series of VT earthquakes between 2 and 11 km depth, the largest of which was a M_L 1.9.

SEPTEMBER 17 THROUGH DECEMBER 31, 1992

The seismicity that followed the September 16–17 eruption differed markedly from seismicity that followed the two earlier eruptions. Strong volcanic tremor occurred for roughly 9 days on stations close to the volcano. The rate of VT earthquakes between depths of 0 and 5 km increased and remained elevated over the next 6 weeks (figs. 4B, 4C). The occurrence rate of shocks deeper than 10 km also increased following the September eruption.

Shallow volcanic tremor returned several times between October 2 and 7. Early on November 9 the rate of shallow VT earthquakes began to increase on

the stations closest to the volcano. By 9:00 p.m. ADT the events were occurring at a rate of approximately 1.5 per minute. This activity continued for roughly 3.5 hours and comprised more than 170 shocks. The magnitude of the largest earthquake was 1.7. This was the most energetic swarm recorded at Mount Spurr associated with the 1992 eruption sequence.

Following the swarm on November 9 and 10, activity remained elevated at all depths. In early December the rate of events between depths of 20 and 40 km began to decline slowly. A swarm of small ($M_L < 1.5$) VT earthquakes occurred between December 21 and 26. These events originated between depths of 5 and 12 km, and most of the events were clustered between 8 and 10 km. Activity slowly declined through early 1993 and by late April the number of locatable earthquakes in the Mount Spurr area approached a level comparable to that observed prior to August of 1991.

DISCUSSION

VOLCANOLOGICAL INTERPRETATION

From our observations of the spatial and temporal patterns of hypocenters, waveforms of the various seismic events, and amplitude and spectral measurements, we offer some interpretations. This discussion should be viewed as preliminary, as it is based on primary observations rather than a detailed analysis of all available seismic data.

Between October 1981 and July 1991, Mount Spurr seismicity was characterized by steady-state activity beneath the summit, punctuated by two swarms under the north caldera rim. There is a marked absence of seismicity at depths greater than 5 km during this period. Jolly and others (1994) model this aseismic zone as a volume of hot buoyant rock, heated by earlier episodes of intrusion, which deformed plastically rather than elastically.

The first unusual activity identified prior to the 1992 eruptions was a shallow short-lived swarm of VT earthquakes beneath Crater Peak in August of 1991 (fig. 4B). This swarm was followed within several months by an increase in shallow activity throughout the caldera. A more subtle increase was also observed at depths greater than 10 km. Its onset may have preceded slightly the August 1991 Crater Peak swarm, but there are too few deep events to draw a definitive conclusion. We propose that most of the seismicity in the 10-month period prior to the June 27 eruption resulted from an intrusion of magma between depths of 5 and 15 km. Thus we suggest a deep origin for the initiation of the eruption sequence. The caldera-wide seismicity likely resulted from a mechanical response to deformation and stresses associated with this intrusion. We suggest that the intrusion produced strain rates that were sufficient to drive rock in the aseismic volume identified by Jolly and others (1994) from plastic to brittle behavior.

That seismicity beneath the summit declined following the June 5 swarm lends support to the mechanical model for the precursory seismicity. The June 5 swarm is coincident with the onset of volcanic tremor at Crater Peak. We suggest that magma had migrated to shallower depths by this time and had activated the shallow hydrothermal system (McNutt and others, this volume). This migration concentrated stress at Crater Peak, which resulted in the decline in activity in other areas of the caldera.

Following the June eruption, seismicity entered a 7-week period of near-quiescence comparable to or less than the level of activity observed prior to August of 1991. This quiescence suggests that stress and related deformation, as well as fluid pressures within the shallow hydrothermal system, were greatly reduced following the eruption. Activity below a depth of 10 km did not lapse into quiescence, although the overall rate of deeper events remained relatively low and did not increase until after the August 18 eruption.

That no increase in shallow seismicity preceded the August 18 eruption and that the September 16–17 eruption was preceded by only several hours of shallow activity visible on CPK suggest that the vent and shallow conduit system were left relatively unobstructed following the June eruption.

Deeper seismicity (>10 km) increased markedly following the August 18 eruption and peaked shortly after the September eruption. The waveforms of these events cover a broad spectrum ranging from VT events to rather monochromatic, low-frequency signals suggestive of LP events (fig. 5). These events might reflect the recharge and migration of magma at these depths. Alternatively, these events might result from readjustments of stress in response to the removal of magma from these depths. Both these explanations are

quite plausible in view of the variety of observed waveforms. However, additional analyses are required to demonstrate that the low-frequency waveforms are related to the characteristic of the source process rather than the result of attenuation of higher frequency energy along the wave path.

The strong swarm of events on November 5 and 10 is likely related to a shallow intrusion of magma. Preliminary analysis of these events suggests they occurred at three locations a few hundred meters apart between 0.0 and -1.0 km depth (Chris Stephens, written commun.). The cluster of events in late December is perhaps related to an additional deeper intrusion at 8 to 10 km depth.

ERUPTION FORECASTING

In this section we review the role that seismological observations played in formulating eruption forecasts and the factors that influenced our interpretations at the time. The forecasting strategy used by AVO relied on the synthesis of data from a number of monitoring techniques, which include a variety of seismic methods (hypocenters, seismicity rate, RSAM, SSAM, and waveform characteristics), visual observations, measurements of gas emissions, and geochemical and petrologic analyses. In reviewing AVO's public statements concerning volcanic activity at Mount Spurr volcano, we use the terms forecast and eruption notification. A forecast describes statements issued prior to expected eruptive activity. Forecasts were relatively imprecise statements concerning the timing and nature of expected eruptive activity. Our use of the term forecast differs somewhat from that of Swanson and others (1985) because at Mount Spurr forecasts were formulated on the basis of observed changes in seismic measurements and observations of the volcano. We avoid the term prediction because statements issued for Mount Spurr contained little precise information on the timing of expected eruptive activity. Eruption notifications were purely factual statements that contained information about the onset time and size of eruptions. Eruption notifications described here belong in the category of factual statements in the terminology of Swanson and others (1985). The primary vehicle by which AVO disseminated forecasts and eruption notifications to other government agencies, industry, and the public was based on a color code system similar to that developed during the 1989–90 eruptions at Redoubt Volcano (Brantley, 1990). The color code definitions used at Redoubt were modified slightly on the afternoon of June 26 to more appropriately describe the situation and conditions at Mount Spurr.

Both the June 27 and September 16–17 eruptions were successfully forecast. Whereas no forecast

was issued for the August 18 eruption, seismic confirmation allowed AVO to issue an eruption notification within minutes of the onset of both the premonitory and the main eruption. Forecasts were also issued during a period of heightened tremor on October 2 and during the November 9–10 earthquake swarm, but no subsequent eruption occurred.

The 10-month period of precursory VT earthquakes makes the June 27 eruption one of the more anticipated eruptions on record. The unusual character of the August 1991 swarm was recognized as the events were occurring. This was possible because of the long-term operation of the seismic network and the availability of the associated earthquake catalog. Between August of 1991 and June of 1992, the increasing caldera-wide seismicity was discussed in weekly AVO public information releases. On June 8 in response to the escalating seismicity, AVO issued a special public advisory, which stated that the present seismic unrest could be an early forerunner of an eruption. In response to the onset of continuous tremor on June 26, AVO issued a forecast that cited potential hazards associated with an eruption of Crater Peak. The basis for this forecast was the onset of strong continuous volcanic tremor (McNutt and others, this volume) as well as observed changes in the hydrothermal regime at Crater Peak (Eichelberger and others, this volume). While this warning did discuss the increased probability of an eruption, it did not state that an eruption was likely within the next few days. Our hesitation regarding the imminence of an eruption on the evening of June 26 largely stemmed from two factors: two strong episodes of volcanic tremor had occurred earlier on June 24 and 25 (McNutt and others, this volume), and the absence of swarms of LP events. We were perhaps overly focused on the occurrence of swarms of LP events, because of our recent experiences at Redoubt Volcano (Chouet and others, 1994, Stephens and others, 1994) and Mount Pinatubo (Harlow and others, 1991), at which strong LP swarms preceded many of the eruptions. Once the swarm of VT earthquakes began at about 3:00 a.m. ADT on June 27, we had little doubt that an eruption was close at hand. In hindsight, the strength of the tremor on June 26 differed enough from earlier episodes to justify a stronger forecast.

In contrast to late June, no unusual shallow precursory activity was observed prior to the August 18 eruption, although station CPK had not yet been repaired. The eruption ended the 7-week period of shallow seismic quiescence, which followed the June event. The eruption was not expected as the level of seismicity had declined sharply after the June eruption and remained low. Additionally, the 1953 eruption at Crater Peak consisted of a single event (Juhle and Coulter, 1955). The caldera-wide volume activated

seismically prior to the June eruption raised the question of the volume of magma involved and the possibility of additional eruptive events. However, in view of the long duration of seismic quiescence following the June eruption, the timing of the August eruption was a surprise. Seismic monitoring did allow confirmation of a small premonitory eruption at 3:35 p.m. ADT, which was reported by a pilot. The occurrence of this small event prompted AVO to issue a public advisory of increased activity. When the main phase of the eruption began at 4:42 p.m. ADT, AVO issued an eruption notification within 2 minutes of its onset.

The September 16–17 eruption was preceded by 3 hours of identifiable shallow precursory activity (McNutt and others, this volume). The overall level of seismicity below a depth of 3 km had been increasing steadily since the August 18 eruption. Following the August 18 eruption, we were particularly mindful that future eruptive events might occur with little or no precursory activity. Close monitoring of the seismicity allowed identification of the precursory tremor shortly after it began and the recognition of the small premonitory eruption at 10:36 p.m. ADT. A forecast was issued at 10:45 p.m. ADT, more than an hour before the main phase of the eruption began at 12:04 a.m. ADT on September 17, and an eruption notification was issued shortly after its onset.

The two unsuccessful forecasts issued during the 1992 eruption sequence were based on an increase in tremor amplitude on October 2, and on an increase in the rate of earthquake occurrence during the November 9–10 swarm. The tremor on October 2 was similar in character to that which immediately preceded the June 27 eruption. Average tremor amplitudes increased dramatically between October 1 and 5 (McNutt and others, this volume). The swarm on November 9–10 was the most energetic observed during the eruption sequence and is likely related to the emplacement of a shallow intrusion. At the height of this swarm, earthquakes were occurring at a rate of roughly 1.5 per minute. While neither of these seismic episodes culminated in an eruption, the forecasts were justified on the basis of the continuing increases in seismicity that were larger than those observed prior to the August and September eruptions. These events serve as a reminder that not all strong increases in seismicity result in eruptions. At this time, there is no seismological basis for differentiating these sequences of events from those that resulted in eruptions.

CONCLUSIONS

Seismological observations suggest that precursory seismicity resulted from an intrusion at mid-crustal depths, perhaps 5 to 15 km. The shallow precursory seismicity that occurred throughout the Mount Spurr

caldera was likely a mechanical response to the intrusion. We also observed an increase in seismicity in a previously identified aseismic volume of rock, between depths of 10 and 40 km depth (Jolly and others, 1994). The onset of this seismicity suggests the intrusion generated sufficient strain rates to induce brittle failure in a volume that had previously deformed plastically.

The June 27 eruption was preceded by 10 months of shallow VT earthquakes, several weeks of shallow volcanic tremor, and immediately by 14 hours of tremor and a 4-hour swarm of VT earthquakes between depths of -1.0 and 1.0 km. In contrast the August 18 eruption had no precursory activity visible on the stations operating at the time, and the September 16–17 eruption had only about 3.5 hours of shallow precursory activity. We conclude that the vent and the upper portions of the conduit were left largely unobstructed following the June eruption.

Seismicity between depths of 10 and 40 km exhibits a variety of waveforms ranging from VT to LP. The diversity of waveforms suggests a variety of source processes ranging from stress adjustments in the country rock caused by withdrawal or intrusion of magma, or the transport of magma. Planned analysis of individual deep events should allow us to narrow our interpretations.

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Preliminary Analyses of Volcanic Tremor Associated with 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

By Stephen R. McNutt, Guy C. Tytgat, and John A. Power

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ABSTRACT

The occurrence of volcanic tremor played an important role in monitoring and eruption forecasting at Mount Spurr by signifying changes in rates of precursory seismicity and indicating the beginnings and ends of eruptions. For example, the onset of continuous tremor provided the principal basis for changing the level of concern color code 15 hours before the first eruption on June 27, 1992. The onset of strong tremor marked the beginning of each of the three main eruptions, and it was the basis for issuing updates and notifications. The declines in tremor amplitude signaled the ends of the eruptions. Tremor also provided physical insight concerning eruptive processes based on spatial and temporal variations in amplitudes and frequencies versus time. The volcano is monitored by a network of 10 seismic stations within 22 km of the vent, in addition to regional seismic stations at greater distances. Three types of tremor were observed at Mount Spurr in association with the 1992 eruptions—tremor bursts lasting 1 to 10 minutes, continuous tremor lasting 2 hours to several days, and tremor during eruptions. Tremor bursts occurred from June 6 to 26 in higher numbers at first which then declined. The overall cumulative energy curve is concave upward, but individual segments are all concave downward; this phe-

nomenon suggests that sources of energy such as heat or water formed quickly then gradually lost energy. Tremor energy increased by one order of magnitude in the 3 days before the first eruption on June 27, 1992. Comparison of spectra from June 26 and June 27 showed that spectral peaks were shifted to lower frequencies by 18 percent; this shift suggests an increase in source length or a decrease in velocity caused perhaps by formation of bubbles. RMS amplitude ratios computed for near (400 m) and far (5 km) stations showed an increase by a factor of 2.7 at the near station from June 11 to June 25, which we model as a shoaling of the source depth. Eruption tremor was substantially stronger than any of the pre-eruption tremor, and was recorded on stations at distances of as much as 138 km. Calculated reduced displacements ranged from 16 to 30 cm² for the strongest eruption tremor. The duration of the strongest eruption tremor was crudely proportional to the tephra volume for the three eruptions. Future efforts will focus on modeling the source mechanisms in greater detail.

INTRODUCTION

Active volcanoes produce a great variety of seismic signals. In addition to earthquakes and various low-frequency events, volcanic tremor is often observed. The typical appearance of tremor is that of an irregular sinusoid, and a key distinguishing feature is its long duration compared with earthquakes of the same amplitude. Continuous signal durations of days, weeks, and longer are common. Tremor has been recorded at 129 volcanoes worldwide (McNutt, 1992).

Volcanic tremor has received considerable attention in the literature over the past few years. Research efforts have been concentrated in four main areas: (1) sources, including shock-waves, (2) propagation effects, (3) magma flow and bubble distribution, and (4) nonlinear models. Theoretical work on tremor by Chouet (1992) focused chiefly on source effects of fluid pressurization in resonant cavities excited by a variety of mechanisms. Chouet numerically treated a rectangular crack that was excited by pressure pulses

acting at various locations on the surface of the crack. He considered his source model to apply to both long-period (LP) events and to volcanic tremor, by assuming that superposition of LP events produced tremor. More recently, Chouet and others (1994) have extended this model to the LP swarms accompanying the 1989–90 eruptions of Redoubt Volcano, Alaska. They suggest, in the new model, flow-induced shock waves at a depth of 1.4 km below the vent as the source of the pressure pulses. Dawson and others (1992) modeled a small, low-velocity zone at the same depth, and they suggested the presence of a small magma body.

Gordeev (1992) studied propagation effects, using near-surface pressure fluctuations such as explosions as sources, and he modeled waveforms propagating through low-velocity layers overlying a half space. His conclusion was that the frequencies and waveforms of tremor are chiefly caused by propagation effects. Hurst (1992) performed stochastic simulations of tremor using both white noise and Poisson forcing functions to excite resonance in a simple harmonic oscillator. He also allowed the oscillator to have a random damping factor. His simulations produced synthetic signals that closely resemble observed volcanic tremor at Ruapehu in New Zealand.

Montalto and others (1992) qualitatively modeled the effects of bubble distribution on tremor and eruption energy for the 1990 eruptions of Mount Etna in Sicily. They suggest that low-amplitude tremor is caused by bubbly flow of magma, whereas high-amplitude tremor is caused by slug flow.

Ukawa (1993) made an important observation for the 1989 Ito-oki submarine eruption. He found that low-frequency (1 Hz) tremor within 50 km of the source consisted primarily of surface waves trapped in sedimentary layers. Middle-frequency tremor (2–7 Hz) observed at the same time at more distant stations was composed primarily of body waves, thus highlighting the importance of propagation effects. Julian (1994) studied a nonlinear model for tremor generation, using a lumped parameter model of flow through a channel that is represented by a mass-spring system, which includes damping. He argued that self-oscillations of this system can produce most of the main features observed in volcanic tremor. In summary, all this recent work suggests that a variety of source and propagation effects are present in volcanic tremor, and there is as yet no consensus about the tremor source. In fact, there are such a variety of observations and source models that there are probably several sources for tremor, not just one.

The occurrence of volcanic tremor played an important role in eruption monitoring and forecasting at Mount Spurr by signifying changes in rates of precursory seismicity and indicating the beginnings and

ends of eruptions. For example, the onset of continuous tremor provided the principal basis for changing the level of concern color code 15 hours before the first eruption on June 27, 1992. The onset of strong tremor marked the beginnings of each of the three main eruptions, and it was the basis for issuing updates and notifications. The declines in tremor amplitude signaled the ends of the eruptions. Analyses of tremor helped provide insight into physical processes that occurred prior to, during, and after the three eruptions of June 27, August 18, and September 16–17, 1992. Three types of volcanic tremor were recorded at Mount Spurr: (1) tremor bursts of 1 to 10 minutes duration, with irregular amplitudes, occurring from June 6 to June 26; (2) continuous tremor lasting 2 hours to several days, with uniform amplitudes, occurring June 24–27, September 17–25, October 2–20, and November 9–10; and (3) tremor during eruptions on June 27, August 18, and September 16–17. The purpose of this paper is to show examples of all three types of tremor and to present results from the preliminary analyses that have been performed as of April 1993. We also discuss possible models for tremor at Mount Spurr, compare its tremor to that at other volcanoes, and discuss the role played by tremor in eruption forecasting.

MOUNT SPURR TREMOR DATA

Mount Spurr is monitored by a network of 10 seismic stations within 22 km of the summit (fig. 1). The active vent, Crater Peak, is located 400 m west of seismic station CPK. Eight of the stations have a 1-Hz vertical geophone, (CPK and CPA are three-component) and data are telemetered in analog form via radio and microwave to the Alaska Volcano Observatory offices at the Geophysical Institute of the University of Alaska Fairbanks. Power and others (this volume) describe the instrumentation and various data-acquisition systems in detail. One data-acquisition system used in Fairbanks digitizes the signals at a sampling rate of 120 samples per second using a Masscomp computer system. In addition to triggered event recording, this system records data continuously from about 140 stations of the Alaska regional seismic network 24 hours per day on DAT tapes (Sonnafrank and others, 1991). This computer provided the primary data set analyzed in this paper. Seismic data are also recorded on a Real-time Seismic Amplitude Measurement (RSAM) system (Endo and Murray, 1991), a Seismic Spectral Amplitude Measurement (SSAM) system (Stephens and others, 1994), and on a PC/AT computer, which runs continuously in event-detected mode (Lee, 1989).

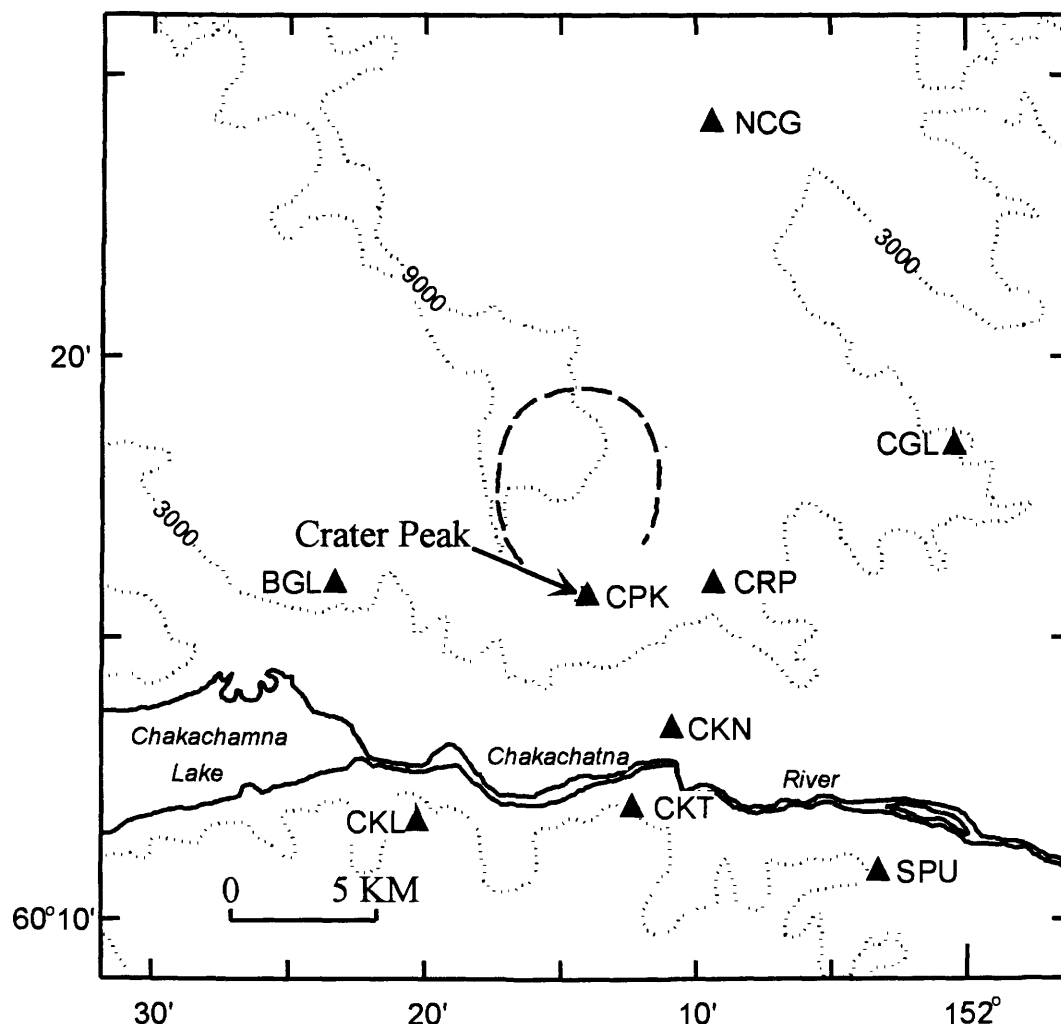


Figure 1. Seismic stations within 20 km of Mount Spurr volcano, southwestern Alaska. Station CPK was destroyed 46 minutes after the start of the June 27, 1992, eruption and was reinstalled in early September. Station CPK was destroyed again on September 17 and was replaced by station CP2. Station CRP became inoperative about 2 hours after the start of the August 18 eruption, and was repaired several days later. Dashed line is caldera boundary. Contours in feet.

Both RSAM and SSAM are rapid-analysis tools used routinely by the Alaska Volcano Observatory. RSAM provides plots of 1-minute averages of seismic amplitude or energy, and SSAM provides plots of spectral amplitude in predetermined narrow frequency bands. Both of these monitoring tools are useful for quantitatively determining the gross features of tremor automatically during crises when detailed manual analyses are not possible. RSAM and SSAM are especially useful digital recording techniques when signal amplitudes on helicorder drums begin to overlap or exceed the physical dynamic range of analog recording, and so reliable amplitudes cannot be easily obtained. At Mount Spurr, RSAM and SSAM were most useful during tremor episodes on June 26–27, October 1–6, and November 9–10.

Volcanic tremor first appeared at Mount Spurr on June 6, 1992, following a 9.5-month swarm of volcano-tectonic earthquakes (Power and others, this volume). The tremor appeared as irregular events 1 to 10 minutes in duration, and had relatively high amplitudes at station CPK (400 m from the vent; CPK is a co-located low-gain station) and lower amplitudes at other stations (fig. 2). Dominant frequencies were between 3 and 6 Hz at station CPK, and about 2 Hz at other stations. We called the events tremor bursts on the basis of their similarity to events recorded at other volcanoes, such as Mammoth Mountain (Hill and others, 1990). We use the term “bursts” for tremor lasting 1 to 10 minutes and the term “episodes” for tremor lasting hours or longer. The only tremor with durations between 10 minutes and 1 hour was associated with eruptions on August 18 and September 16–17.

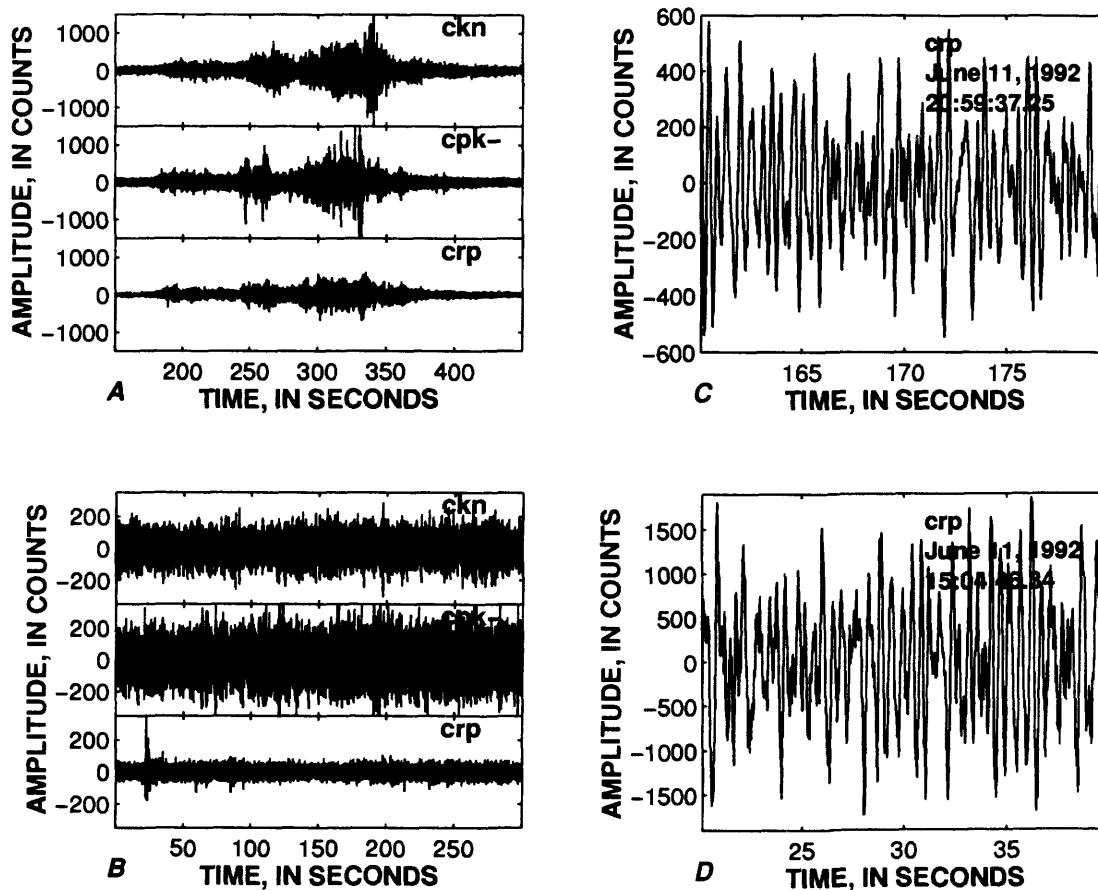


Figure 2. Seismograms of volcanic tremor from Mount Spurr volcano, Alaska. *A*, tremor burst from June 11, 1992, on stations CKN, CPK, and CRP, 300 sec of data shown; *B*, tremor episode June 24–25, 1992, same 3 stations, 300 sec of data shown; *C*, continuous tremor from June 26, 1992, station CRP, 20 sec of data shown; *D*, eruption tremor, June 27, 1992, station CRP, 20 sec of data shown.

A total of 174 tremor bursts occurred between June 6 and June 27 (Appendix 1). A peak rate of 24 tremor bursts occurred on June 11, and one burst each occurred on June 17, 18, and 21. We measured durations in seconds for each event on seismograms from station CPK, and peak-to-peak amplitudes in mm on seismograms from station CRP because CPK was frequently saturated (see Appendix 1). Squaring the amplitude and multiplying by the corresponding duration of each burst gave a measure of relative energy. We then summed the durations and energies for each 6-hour period. The resulting plots of cumulative durations and energy rates were approximately constant from June 6 to June 13, then declined (became less steep) until June 21, when they increased slightly (fig. 3). On June 24, a 154-minute long tremor episode occurred, followed 12 hours later by another similar episode lasting 142 minutes. Continuous tremor began at 12:03 p.m. ADT (20:03 UT) on June 26 and lasted 19 hours until the onset of the eruption on June 27 at 07:04 a.m. ADT (15:04 UT). The amplitude fluctuated

during this time (fig. 4) and the signal resembled banded tremor (Kieffer, 1984; McKee and others, 1981). The three tremor episodes from June 24–27 represented an order of magnitude increase in energy over all the tremor bursts recorded until June 24 (fig. 3). Tremor amplitude approximately doubled on June 27, 1992, at 07:04 a.m. ADT (15:04 UT), signaling the start of the eruption. This was recognized within 2 minutes, and the level of concern color code was formally changed several minutes later.

Being able to consistently determine eruption start times, end times, and durations (table 1) precisely was problematic because of station failures. Start times were determined as the time when the amplitude on the station CRP helicorder in Fairbanks first exceeded 3 mm peak-to-peak (1.5 mm peak to-peak for September owing to a gain change). On June 27, this occurred about 12 seconds after the apparent start of the eruption on station CPK, which was the closest station, located on the crater rim 400 m from the vent. The 12-sec delay was partly due to propagation time

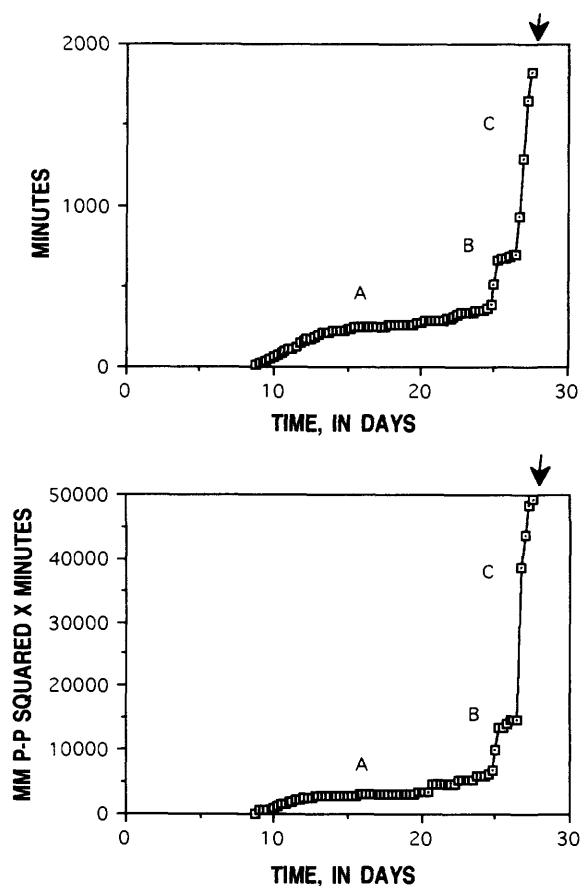


Figure 3. Mount Spurr tremor cumulative duration and energy versus time for June 1992. Data computed for 6-hour intervals. Eruption occurred on June 27 (arrows). The energy units are mm squared times minutes, and differ from standard units such as ergs by a multiplicative constant. Parts labeled A, B, and C are discussed in text.

and partly due to attenuation at the 5-km distance of station CRP. Unfortunately, station CPK was destroyed, so it could not be used to determine the start times

for all the eruptions, so station CRP was used instead. The end times were determined when the peak-to-peak amplitude on the station BGL helicorder in Fairbanks dropped below 1 mm (2 mm for September). This is a fairly high threshold because of the greater distance of station BGL (7.7 km), so the indicated times correspond to the strong phases of activity. Station BGL was used for end times because station CRP became inoperative after a lightning strike during the eruption on August 18. All times and durations are shown in table 1.

This first eruption lasted 4 hr and 3 min (table 1). No tremor was recorded over the next 52 days until August 18, when a 12-minute burst of weak tremor including 3 entrained LP events coincided with pilot reports of a small ash plume. About 1 hour later, the main phase of the second eruption began, accompanied by strong tremor (as much as 30 cm² reduced displacement) lasting 3 hr and 28 min. No tremor was recorded over the next 29 days until September 16, about 3 hours before the September 16–17 eruption. Medium-strength tremor (maximum about 5 cm² reduced displacement) occurred for 11 min during the first phase of the eruption at 10:36 p.m. ADT, and strong tremor (as much as 25 cm² reduced displacement) occurred for 3 hr and 36 min during the main phase of the eruption which began at 12:03 a.m. ADT September 17 (table 1). Unlike the aftermath of the June 27 and August 18 eruptions, continuous tremor occurred for 1 week following the September 17 eruption.

Tremor reappeared on October 2, and it occurred almost continuously with variable amplitude until October 20. The last tremor occurred during a swarm of earthquakes on November 9–10, 1992, that may have been associated with an intrusion (Stephens and others, unpub. data).

Altogether, during the 1992 activity of Mount Spurr, more than 460 hours of volcanic tremor were recorded.

Table 1. Mount Spurr volcano, Alaska, eruption start times, end times, and durations, as determined from tremor amplitudes, for 1992 eruptions.

[All times expressed in Alaskan daylight time.]

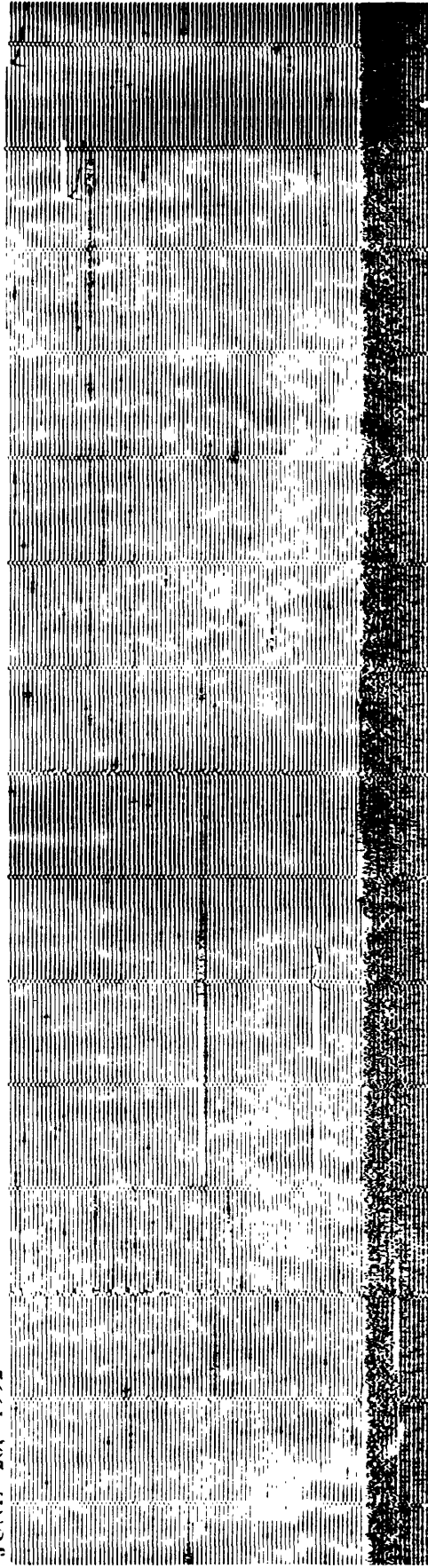
Date	Start time CRP>3mm	End time BGL<1mm	Duration hr:min	Tephra volume [^] x10 ³ m ³
6/27	07:04:49 a.m.	11:07:51 a.m.	4:03	44
8/18	03:41:06 p.m.	03:52:55 p.m. #	0:12	(visible ash plume)
8/18	04:42:25 p.m.	08:10:45 p.m.	3:28	52
9/16	10:36:18 p.m.	10:47:40 p.m.	0:11	(2-km ash plume)
9/17	12:03:48 a.m.	03:39:52 a.m. *	3:36	56

[^] R. McGimsey, written commun., 1993

measured at station CRP

* BGL<2mm

JUNE 26, 1992



STATION CRP

JUNE 27, 1992



Figure 4. Helicorder records from station CRP, 5 km west of Crater Peak vent on Mount Spurr volcano, southwestern Alaska, for June 26 and 27, 1992. Note (A) continuous tremor starting at 12:03 p.m. ADT (20:03 UT) June 26, and (B) earthquake swarm starting 03:00 a.m. ADT (11:00 UT) June 27. Eruption tremor (arrow) started at 07:04 a.m. ADT (15:04 UT) June 27. Tick marks are at one-minute intervals.

ANALYSES AND RESULTS

We studied (1) duration and energy versus time for all the tremor bursts and the first three episodes of continuous tremor during June 6–27; (2) amplitude ratios for selected tremor bursts (June 8–22) and one episode of continuous tremor on June 24–25; (3) amplitude variations for October 1–6 continuous tremor; (4) spectra for a tremor burst, the June 24–25 tremor episode, a sample of June 26 tremor, and a sample of June 27 tremor during the first eruption; (5) spectral variations for October 1–6 continuous tremor; and (6) a comparison of tremor during eruptions at Mount Spurr with tremor recorded during eruptions at other volcanoes.

DURATION AND ENERGY VERSUS TIME

Tremor bursts began on June 6, 1992, and we tabulated the events and plotted the cumulative duration and cumulative energy as a function of time (fig. 3 and Appendix). We observed that the overall pattern of the plots was concave upward, but individual segments were all concave downward. For example, the cumulative-duration plot was concave downward from June 6 to June 21 (A in fig. 3) indicating that the rate of occurrence of the tremor bursts was decreasing with time. On June 11, a field party observed that the crater lake had changed color from light blue to dark gray, and that adjacent fumaroles had increased in activity (T. Miller and R. McGimsey, oral commun., 1992). These observations and shallow earthquake locations suggested that a change had occurred in the shallow geothermal system, specifically an increase in heat caused by magma intruding to shallow levels in the crust. The concave downward part of figure 3 suggests that the source of heat moved up and stopped, and it gradually lost energy as a function of time. Alternatively, the geothermal system could have used up its supply of water or some other geochemical change could have occurred, such as partial sealing of cracks and pores. The low amplitudes of the tremor bursts (about 1 cm² reduced displacement) were in the same range as tremor observed elsewhere in association with hydrothermal activity, lending further evidence to the hypothesis that the tremor bursts probably represented some type of boiling phenomena.

The two 2-hr long tremor episodes on June 24–25 occurred during poor weather, so no direct field observations were made. However, on June 26 a field party observed that the crater lake was half empty, and lithic block impact marks were photographed on the mud of the former lake bottom (C. Nye, oral commun., 1992). The cumulative duration and cumu-

lative energy plots from June 24 to June 26 (B in fig. 3) are also concave downward, suggesting that a heat source (magma) moved upward and stopped, losing its heat or using up the available water. The impact marks suggest that mild phreatic activity may have occurred. The last part of the tremor (C in fig. 3) on June 26 and 27 is also concave downward in energy, but straight in cumulative duration because the same duration of tremor occurred in each 6-hour time step once it became continuous. Inspection of the seismograms in figure 4 shows that the continuous tremor occurred in four bands. In each case the onset was relative abrupt and strong, and the tremor then gradually declined in amplitude over several hours. Each of these four bands follows the same pattern as parts A, B, and C in figure 3, again suggesting that the magma moved upward to a shallower level and stopped, losing heat or using up the available water. Following the 4-hr VT earthquake swarm on June 27 (fig. 4) the onset of strong tremor marked the start of the eruption, confirming that magma had indeed moved to shallow levels in the volcano.

AMPLITUDE RATIOS

Except in rare cases when the onset is abrupt (Aki and Koyanagi, 1981), volcanic tremor cannot be routinely located because no clear phases such as P-wave or S-wave arrival times are present in the waveforms. Therefore, indirect techniques are commonly used to infer changing tremor source locations. At Mount Spurr we examined amplitude ratios between near and far stations to determine whether any significant depth changes had occurred. We used data from stations CPK, 400 m from the vent, and CRP, 4.8 km from the vent (fig. 1). The requirement of on-scale seismograms for CPK and good signal-to-noise ratio for CRP restricted our search to only a few samples, yet sufficient data have been analyzed to provide reliable results.

We computed rms (root-mean-square) amplitudes for CPK and CRP data for six samples (table 2) using the program SAC (Tapley and Tull, 1991). Amplitudes for a tremor burst on June 11 and the 2.5-hr episode on June 24–25 at stations CPK and CRP are shown in figure 5. The plots are normalized so that the amplitude on station CRP appears the same (left side, fig. 5), whereas the CPK amplitude varies. We observe that the amplitude ratio CPK/CRP for June 11 is 1.5, and the ratio for June 24–25 is 4. The amplitude on the close station CPK has increased significantly compared with the amplitude at the more distant station CRP. We interpret this as evidence for a shoaling of the depth of the source. Recall that the

Table 2. Root mean square amplitudes of Mount Spurr volcanic tremor.

Date	Time (ADT)	CKN rms amplitude	CPK rms amplitude	CRP rms amplitude	CPK/CRP ratio
June 8	08:20 a.m.	92	63	41	1.5
June 11	07:00 a.m.	199	204	138	1.5
June 17	12:16 a.m.	74	63	50	1.3
June 22	08:44 a.m.	119	141	58	2.4
June 24	04:24 p.m.	44	68	17	4.0

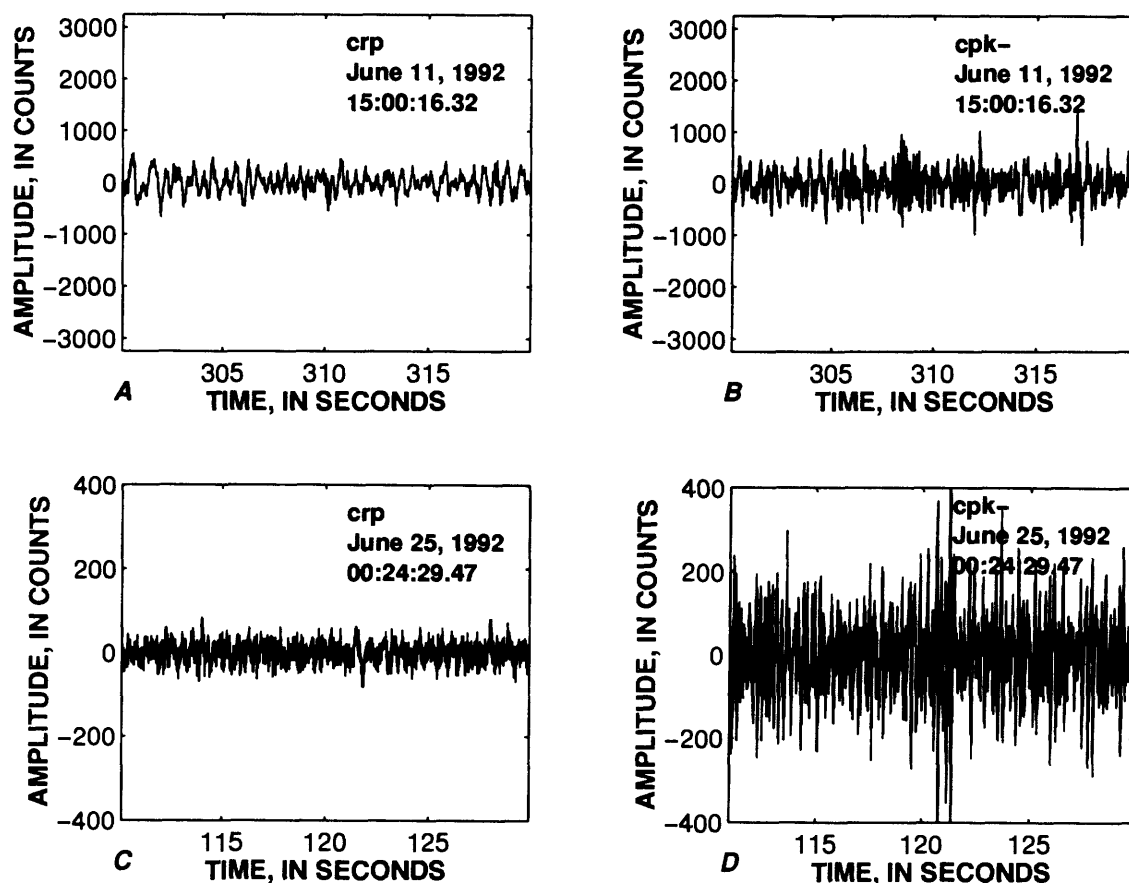


Figure 5. Amplitude ratios for two tremor samples from seismic stations on Mount Spurr volcano, Alaska. For each sample, data from stations CRP (A and C) and CPK (B and D) are shown. The plots are normalized to the CRP rms amplitude. The ratio CPK/CRP on June 11 is about 1.5 (compare A and B), whereas on June 24–25 the ratio is about 4 (compare C and D). This suggests that the June 24–25 event occurred closer to station CPK, or at shallower depth since station CPK is located only 400 m from the vent whereas station CRP is located 5 km from the vent. Universal Time.

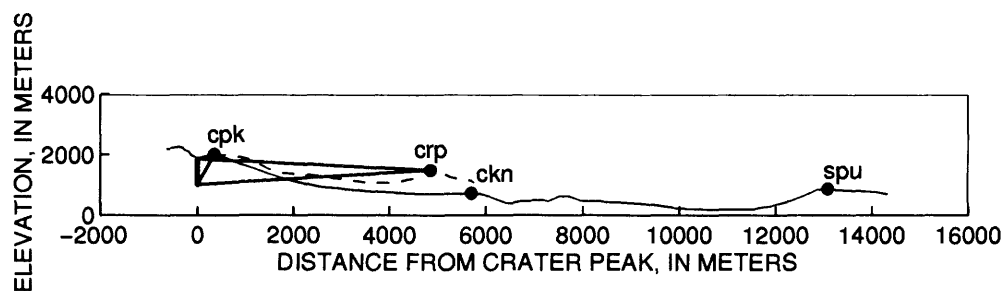


Figure 6. Schematic diagram of postulated depth distribution that explains amplitude ratios shown in figure 5. Two true-scale topographic profiles are shown. Straight line ray paths are drawn to stations CPK and CRP. The rays to station CRP are nearly the same length, whereas those to station CPK are scaled to the changing amplitude ratio. The straight rays to station CRP are unrealistic because of the topography, nevertheless, they illustrate the concept.

crater lake was half empty, with lithic block impacts visible, when it was viewed June 26 prior to the onset of the continuous tremor (Alaska Volcano Observatory, 1993; C. Nye, oral commun., 1992). This observation suggests that mild phreatic activity may have occurred, probably during the June 24–25 tremor episodes.

If we assume a point source at the surface for the June 24–25 tremor, then we can perform a simple inversion for the depth of the tremor burst on June 11. The geometry is shown in figure 6. The distance of the ray path to the distant station CRP (4.8 km) changes very little with a change in source depth, whereas the distance to the close station CPK (400 m) changes significantly. We use the amplitude ratios 1.5 and 4 given above, and assume $1/r$ geometric spreading for body waves (energy decays as $1/r^2$, so amplitude decays as $1/r$). This means the distance must increase by a factor of $4/1.5=2.67$. We assume straight-line ray paths, and use the Pythagorean theorem to solve for the depth change. This yields a depth of 920 m for the June 11 tremor burst. If the June 24–25 source is located at a depth of 100 m rather than at the surface, then the depth of the tremor burst would be 934 m. This calculation is a first-order approximation only (probably reliable to only one significant figure), and it ignores effects such as attenuation, reflections, and refractions. If the source was not deeper but was farther away in some other direction, the distances would be different than those determined above. Also, if the source is a line source or some other geometry, the calculation is more complicated, and it has not yet been attempted.

OCTOBER 1–6 AMPLITUDE VARIATIONS

Volcanic tremor continued for 1 week following the September 17 eruption and ceased on September 25. After a 1-week hiatus, tremor resumed on October 1. The tremor sequence in October showed significant temporal variations, including patterns similar to those that have preceded eruptions at Mount Spurr and elsewhere.

The RSAM plot (fig. 7) shows tremor starting on October 1 and then increasing in amplitude at a nearly exponential rate on October 2, after which it stopped rather abruptly. This episode was similar to the RSAM plot for the days preceding the January 2, 1990, eruption of Redoubt Volcano (Brantley, 1990). On October 3 a series of 5 tremor episodes occurred, each about 1.5 hr long. This signal resembled banded tremor (Kieffer, 1984), and it was also similar to the tremor preceding the June 27 eruption of Mount Spurr. On October 4, tremor returned and the amplitude again

increased nearly exponentially. When this tremor declined in amplitude on October 5, another series of tremor episodes occurred, each about 2 hr long, similar to those of October 3, but followed by a gradual decline.

From October 1 through October 6 a steam plume was being emitted from Crater Peak to heights of 2,000 to 12,000 ft above the volcano, and approximately steady amounts of SO_2 and CO_2 gas were being emitted (400 and 2,900 tons/day, respectively). Doukas and Gerlach (this volume) noted a drop in SO_2 flux shortly after the onset of several tremor episodes. There was no clear correlation between plume height and tremor amplitude, and no eruption occurred.

The RSAM plots were valuable for near-real-time monitoring to quickly compare the current tremor level to that of the previous few days and to monitor the temporal patterns. Although the amplitude varied considerably, the tremor was quite stationary in its frequency characteristics and in its relative amplitudes at different stations. We speculate that this tremor was caused by magma degassing at shallow depths.

The tremor observed in October demonstrated that patterns of occurrence alone, such as exponential increases in amplitude or periodic banding, are not sufficient basis for forecasting eruptions.

SPECTRA

We computed Fourier spectra for several representative portions of the tremor data to search for significant spectral peaks that might help identify source properties. We selected a typical tremor burst on June 11, and samples of tremor episodes on June 25, June 26, and during the eruption on June 27 (fig. 2). For each sample, we chose four adjacent windows, each 10 sec long, computed Fast Fourier Transform spectra, then stacked the spectra. This adds stability because constructive interference enhances the real peaks whereas destructive interference diminishes the noise. Results are shown in figure 8. For all the samples, significant spectral amplitudes occur between 0.5 and 5 Hz, and the highest amplitude peaks are near 2 Hz. In detail, however, two of the spectra show something quite remarkable.

If the June 26 spectrum is horizontally compressed and superimposed on the June 27 spectrum, as shown in figure 9, we see that nearly all of the main spectral peaks coincide. The amount of compression is 18 percent. Mori and others (1989) showed a shifting in frequency of a single prominent peak at Langila Volcano in Papua New Guinea, and Kamo and others (1977) showed a series of spectral peaks from Sakurajima Volcano in Japan whose positions changed sys-

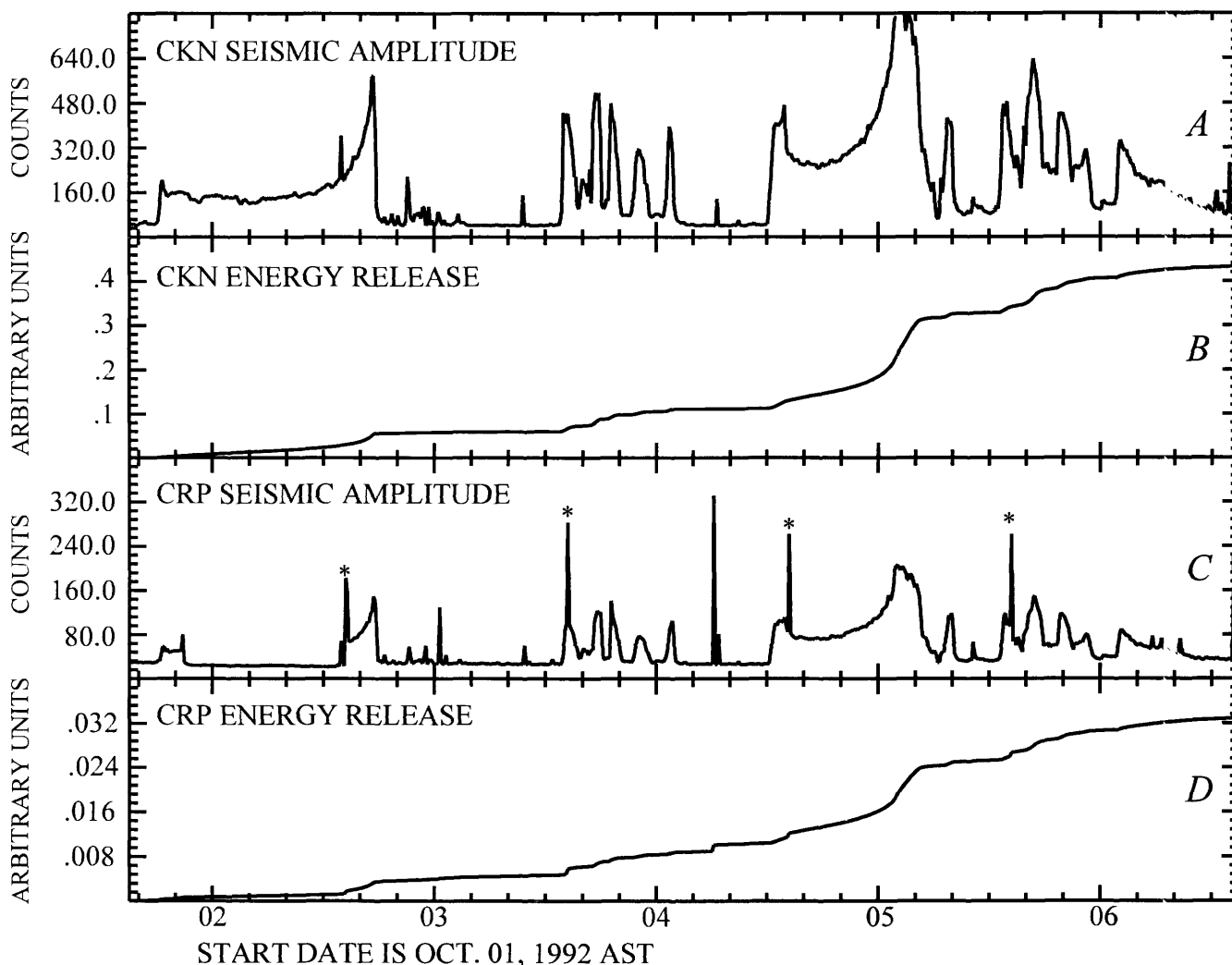


Figure 7. RSAM (Real-time Seismic Amplitude Measurement) plot for tremor data from Mount Spurr volcano, Alaska, for October 1–6, 1992. Horizontal axis is time in days; vertical axis is digital counts for *A* and *C*, arbitrary units for *B* and *D*. *A*, amplitude, station CKN; *B*, cumulative energy release, station CKN; *C*, amplitude, station CRP; *D*, cumulative energy release, station CRP. In *D* on October 2 the signal gain ranged for part of the time. Equally spaced spikes (asterisks) are daily calibration sequences.

tematically with time. In all three cases the shifting of peaks implies either that the length of a resonator changed by 18 to 40 percent, or that the acoustic wave speed changed, as could happen, for example, with the introduction of bubbles. The presence of even a small amount of bubbles dramatically lowers acoustic speeds of fluids (Kieffer, 1977).

The Mount Spurr data differ from these other data in several significant respects. First, the Mount Spurr eruptive conditions changed from no eruption to vigorous subplinian eruption. The Langila and Sakurajima data, on the other hand, were recorded during times of relatively constant conditions. Second, the Langila case was recorded over 2 minutes, whereas the Spurr data were recorded over 19 hours. Third, the Langila and Sakurajima seismograms appear to

be very even and regular, whereas the Spurr data are far more irregular (fig. 2). On closer inspection of figure 9, the Spurr spectra consist of a broadband triangular shaped part (apex upward) from 0.5 to 4 Hz, with evenly-spaced sharp individual peaks superimposed. This has important implications for the source model. The broadband part suggests that a series of random pulses was acting at the source (see Hurst, 1992). These could be, for example, bubble cavitations, flow perturbations, or impacts of blocks on the side of a conduit. The evenly-spaced sharp peaks, on the other hand, suggest resonance modes of an oscillator. The single set of peaks suggests a one-dimensional oscillator, with each of the peaks representing an eigenfrequency of vibration (for example, Schick and others, 1982; Kamo and others, 1977; McNutt, 1986).

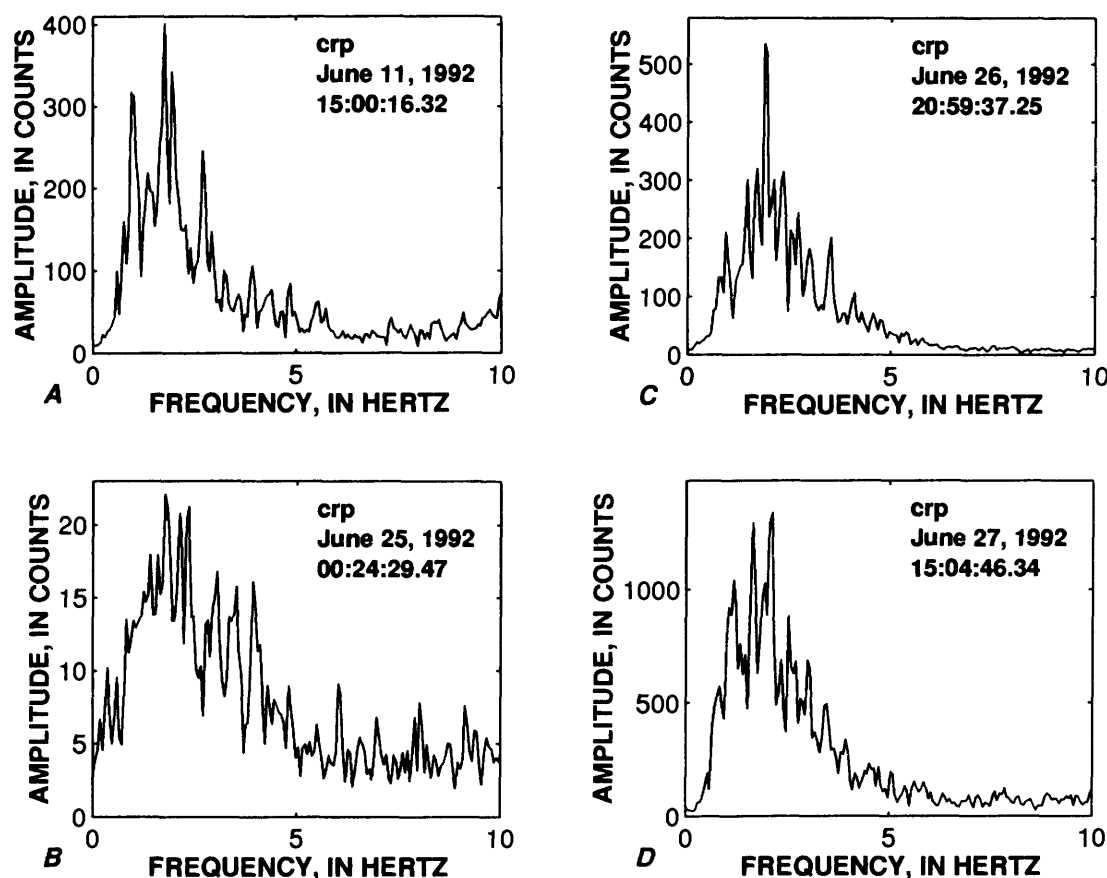


Figure 8. Fourier spectra for the time series shown in figure 2. Each spectrum is a stack of four individual spectra, each for a 10-sec sample of data. A, tremor burst, June 11, 1992. B, tremor episode 154 minutes long, June 24–25, 1992. C, continuous tremor prior to the eruption, June 26, 1992. D, eruption tremor, June 27, 1992. Universal Time.

Models of this type involve a one-dimensional oscillator such as the familiar organ pipe, whose permissible frequencies of resonance are determined by the length, velocity, and boundary conditions. A pipe with two open or two closed ends has resonance modes $n\lambda/2$ (where λ = wavelength and $n=1,2,3,\dots$). A pipe with one open end and one closed end has resonance modes $n\lambda/4$ (where λ = wavelength and $n=1,3,5,\dots$). In each case the spectral peaks are evenly spaced, however, the spacing from the origin to the first peak is different. The interpeak and origin-to-peak spacing is the same for $n\lambda/2$ modes, whereas the origin-to-peak distance is $\lambda/2$ that of the interpeak distance for $n\lambda/4$ modes. This feature provides a simple way to identify modes on a spectrum.

The data suggest that both sets of modes are present in spectra from June 26 and June 27, before and during the eruption. The peaks on figure 9 have been numbered for identification, and we observe that alternate peaks change amplitudes systematically from before to during the eruption. Peaks 1, 3, 5, 7, and 9

had the same or only slightly different amplitudes, whereas peaks 2, 4, 6, 8 and 10 became significantly larger during the eruption. The pattern is most clear for peaks 1, 2, 3, and 4. The most straightforward interpretation for this pattern is that there was a change in the boundary conditions. We suggest that the spectral peaks represent mixed modes of organ pipe vibration, that is, both open-ended and close-ended modes are present. If the boundary conditions are changed, for example by breaking an impermeable cap, changing the shape, or transmitting lava, then one set of modes would be enhanced relative to the other. It is best to think of the boundary conditions in terms of reflection and transmission, rather than open-ended or close-ended. However, the open-ended and close-ended modes give a clear conceptual picture. We observe that the close-ended modes (that is, one open end and one closed end) became stronger during the eruption, which is what would be expected as the vent is reamed out. Although preliminary, to the best of our knowledge, this is the first observation of this type that has

been made on volcanic tremor spectra. This is also a clear source effect. The fundamental mode for the even-numbered peaks (closed-end modes) does not appear on figure 9 because its frequency is too low to be recorded on the 1-Hz seismometer we used.

While we have presented our preferred preliminary model, other possibilities exist. The shape of a resonator could have changed, for example from a one-dimensional pipe-like structure to a two-dimensional rectangular crack (Chouet, 1992), and the systematic variation of spectral amplitudes could be caused by preferential excitation of modes in one direction over the other. It is also possible that no resonator was involved, and that the observed amplitude variation is caused by some other process (Julian, 1994). It is conceivable that the alternate peak spectral amplitude variations are due to chance alone, however, we think that it is very unlikely that random fluctuations would be so systematic.

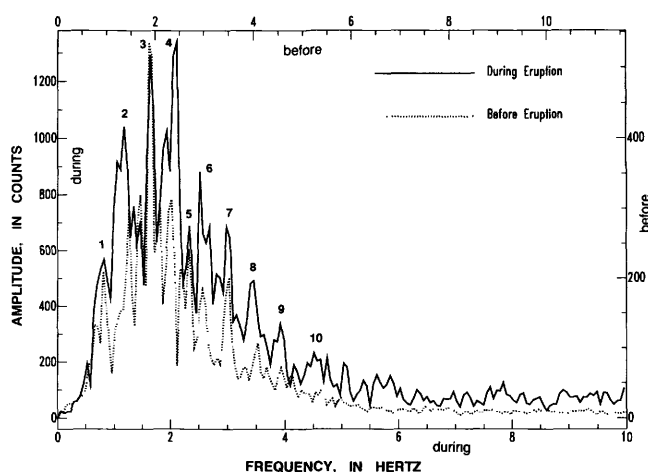


Figure 9. Spectra from figure 8C and 8D superimposed with the horizontal axis compressed by 18 percent for the June 26 spectrum. Left and bottom scales refer to tremor during the eruption on June 27. Top and right scales refer to pre-eruption tremor on June 26. Peaks numbered for reference.

OCTOBER 1-6 SPECTRAL MEASUREMENTS

In addition to computation of spectra for selected tremor samples as discussed above, we also studied spectra using SSAM (Stephens and others, 1994). SSAM runs continuously in the background, and plots can be made at any time for stations and times of interest. SSAM plots show spectral amplitudes in pre-determined bands that are several tenths of a Hertz wide. SSAM is thus a good tool to use when the spectra do not vary with time, but it lacks the resolution to identify small changes in spectral position.

The SSAM plot (fig. 10) shows data from October 2-6, the same period as figure 7. Here we see persistent high amplitudes in the 1.9 to 2.5 Hz bands, the same frequencies that are high amplitude in the spectra shown in figure 8. We also observe that as the signal becomes stronger (fig. 7), the spectral content becomes somewhat more broadband. This may be partially caused, however, by saturation of the spectral bands for strong signals. In general, the spectral content is remarkably uniform over the period October 2-6. Based on the results of both RSAM and SSAM, we conclude that volcanic tremor with a single source occurred during October 1992. Our preliminary conclusion is that shallow-level degassing of magma was the source. It is not known whether the degassing was juvenile, hydrothermal, or a mixture of the two, although we suggest the latter was the more likely gas source (Doukas and Gerlach, this volume).

Tremor occurred again for several hours on November 9-10 in association with a swarm of shallow earthquakes. This tremor was observed mainly on the closest station CP2 (a replacement for CPK, fig. 1) and consisted of a high signal level recorded between discrete events of the earthquake swarm. The frequency content of this tremor was somewhat higher than that of previous tremor (>5 Hz). No detailed analyses have yet been carried out on this tremor.

TREMOR DURING ERUPTIONS

Time histories for the three eruptions using RSAM data for station BGL, which operated continuously during all three eruptions, are shown in figure 11. Station BGL has a 1-sec vertical geophone, and it is located 7.7 km west of the Crater Peak vent (fig. 1). Because the station is fairly close to the vent, its data are partially saturated (clipped), especially for higher amplitudes above 1,200 counts. The total durations and volumes of each eruption are shown in table 1. However, the duration of the strongest eruption tremor was roughly proportional to the tephra volume of each eruption. For example, the duration of tremor greater than 800 counts on June 27 was 1 hour, on August 18 about 2.8 hours, and on September 16-17 about 2.5 hours (recall that the end of the September eruption was "contaminated" by an earthquake swarm), whereas the tephra volumes were 44, 52, and 56 million cubic meters, respectively.

The amplitude of Mount Spurr eruption tremor is compared to that from eruptions at other volcanoes in figure 12 (modified from McNutt, 1994). Representative values from the June 27 and August 18 eruptions are shown; the maximum amplitude during the August eruption was approximately double the June

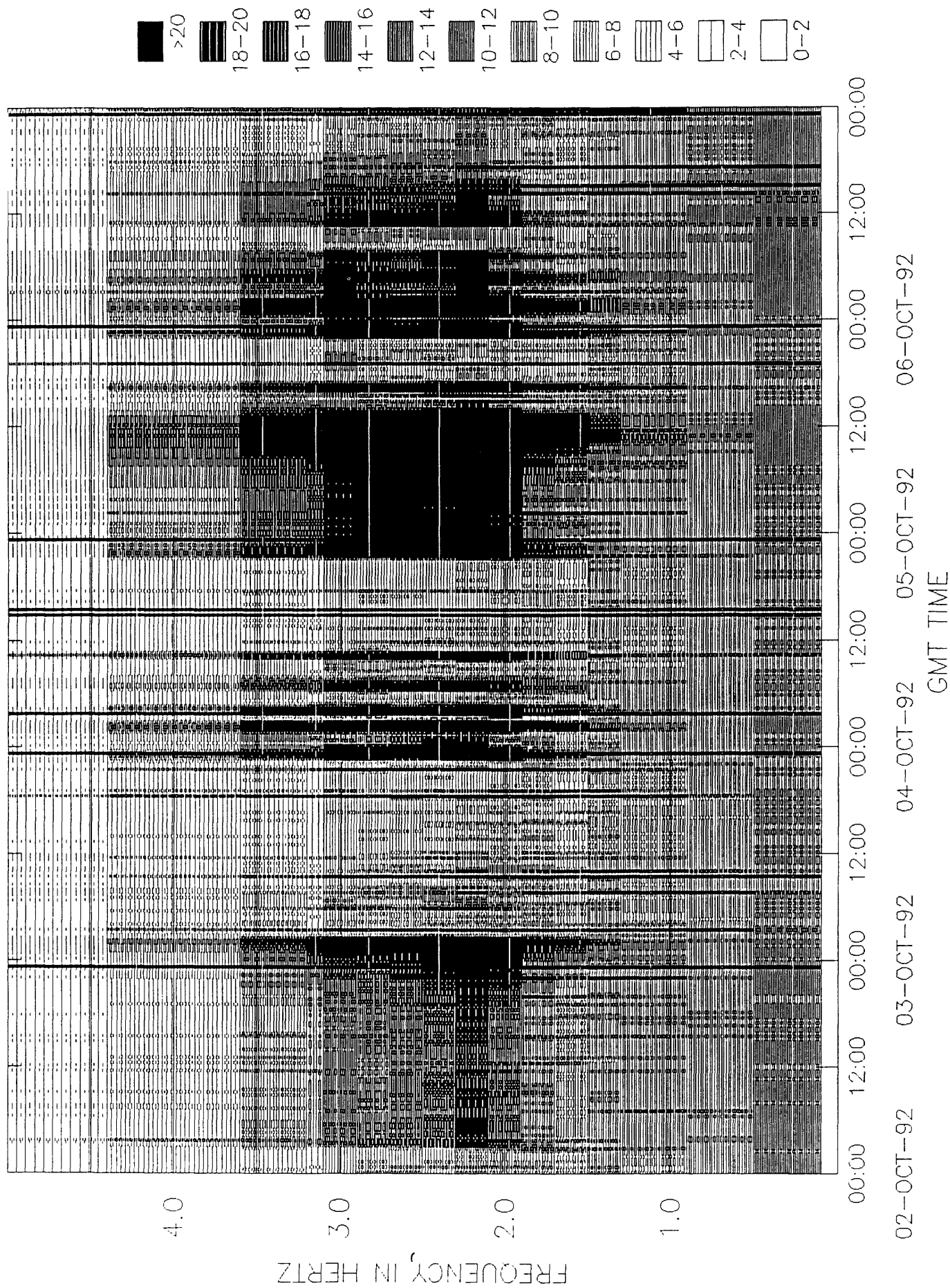


Figure 10. SSAM (Seismic Spectral Amplitude Measurement) record for October 2-6, 1992, for station CRP on Mount Spurr volcano, Alaska, plotted with high values in 18-minute cells. Spectral amplitudes are shown as various shades and patterns, with darker colors representing higher amplitudes. Frequency bands are 0.2 Hz wide. Note persistent high amplitudes in the 1.9 to 3.2 Hz bands.

maximum. Each datum represents the tremor amplitude normalized to reduced displacement, which corrects for distance, geometric spreading, and instrument magnification (Aki and Koyanagi, 1981; Fehler, 1983), and is plotted versus the Volcanic Explosivity Index (VEI) of Newhall and Self (1982). Each of the three main eruptions of Mount Spurr were VEI=3, as determined from tephra volumes (table 1) and ash-column heights of 9 to 15 km. The Mount Spurr tremor amplitudes are similar to values obtained elsewhere for eruptions of similar size, although this does not imply that the exact eruption mechanisms were iden-

tical. The mechanisms of the Mount Spurr eruptions are discussed by Neal and others (this volume), and they are characterized as subplinian.

The amplitude of a representative tremor burst from June 11 is also shown in figure 12 for comparison. The VEI was set at 0 because of the less than 1 km steam plume and the impact block craters in the mud flat within a few tens of meters of the vent. This value also falls in the range observed at other volcanoes for eruptions of similar size and characteristics (Gil Cruz and others, 1987; Leet, 1988; McNutt, 1994).

DISCUSSION AND CONCLUSIONS

We used a variety of methods to study tremor at Mount Spurr, choosing specific techniques for different time periods based on station operation, signal strength, and overall data sufficiency. In this section, we synthesize the key results into a preliminary qualitative model of tremor generation.

The tremor bursts first appeared on June 6, the day after the highest daily number of volcano-tectonic (VT) earthquakes, 28 events on June 5 (Power and others, this volume), of the 9.5-month swarm up to that time. During the three weeks over which tremor bursts were observed, the seismicity remained at about 5 to 10 locatable VT earthquakes per day. This suggests that magma or volatiles had reached shallow levels in the crust, and instead of building up stresses and triggering VT earthquakes, the magma or volatiles interacted with ground water to produce tremor bursts. Recall that geochemical changes were observed in the crater lake shortly after the tremor bursts first appeared. For 18 days the cumulative duration and cumulative energy plots were concave downward (fig. 3), which means that the rates of tremor occurrence and energy were declining with time. Amplitudes were small (about 1 cm² reduced displacement) and source depth was about 1 km. We infer that magmatic heat was interacting with the shallow geothermal system.

A significant change occurred on June 24–25 when a 2.5-hour tremor episode occurred, followed by another similar episode 12 hours later (fig. 3 and Appendix). The durations and energies of these two episodes alone exceeded those of all the earlier bursts together. The source was becoming shallower, as indicated by changing amplitude ratios (figs. 5 and 6) and the crater lake was partially emptied, with lithic block impact marks visible. The next significant change was the onset of continuous tremor 19 hours before the first eruption. The cumulative duration increased again by a factor of about 3, the cumulative energy also increased by a factor of 3 (fig. 3), and the source remained shallow. Evenly-spaced, narrow spectral peaks suggest that a one-dimensional resonator (pipe or con-

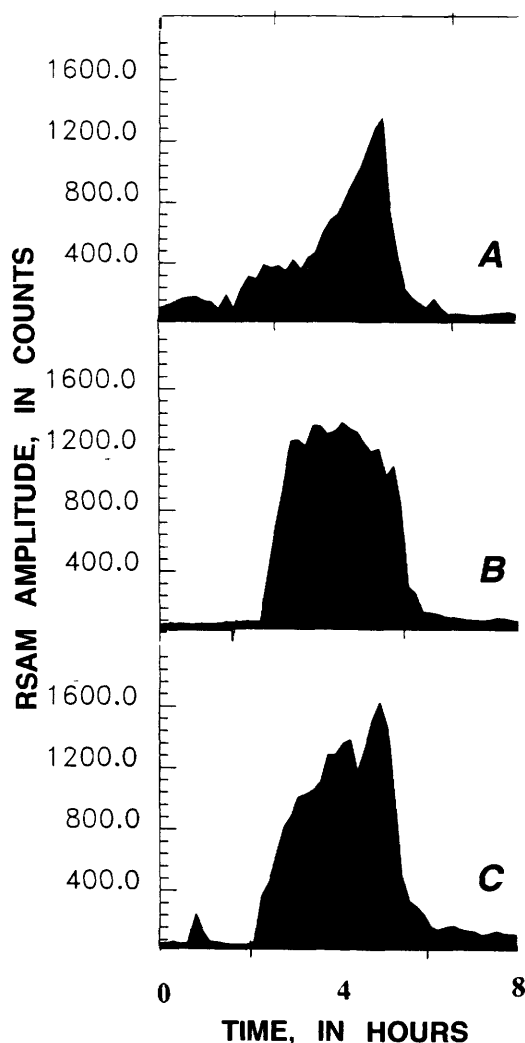


Figure 11. RSAM (Real-time Seismic Amplitude Measurement) plot for volcanic tremor amplitudes during the three 1992 eruptions of Mount Spurr in southwestern Alaska. Station BGL is located 7.7 km west of Crater Peak, the active vent. Vertical axis is digital counts. Partial saturation of highest amplitudes (above 1200 counts) occurred during all three eruptions. A, June 27; B, August 18; C, September 16–17.

duit) was active at this time. We believe the conduit was filled with a mixture of water, lithic materials, magma, and magmatic fluids. The concave downward parts of the energy plot and the similar amplitude patterns on seismograms (figs. 3 and 4) suggest that heating and upward movement occurred intermittently and abruptly, rather than smoothly and steadily.

The onset of a shallow (1 to 2 km beneath the vent) VT earthquake swarm preceded the first eruption by 4 hours (Power and others, this volume); the onset suggests that stresses were again building up as magma made its final ascent to the surface. Tremor was constant and of low amplitude during this swarm. We speculate this means that all the fluids had been boiled off by this time. The eruption began quite abruptly and the tremor looked much the same as tremor recorded previously except that it was stronger by a factor of two. Spectra computed for tremor during the eruption showed the same peaks shifted in frequency by 18 percent. The presence of the same peaks implies that the same structure was active; namely, the conduit previously occupied mainly by water or a water-rich mixture was now occupied mainly by magma. The shifting of the spectral peaks to lower frequencies (fig. 9) suggests that either the resonator (conduit) length increased or the acoustic velocity decreased within the conduit. Either hypothesis is reasonable and we cannot yet distinguish between these possibilities. The systematic change in amplitudes of alternate spectral peaks (fig. 9) suggests that boundary conditions changed from closed (reflection) to open (transmission) at one end of the conduit from before to during the eruption. This observation is consistent with disruption of an impermeable cap at the top of the conduit during the initial stages of the eruption. Tremor during the eruptions was stronger than any previous tremor and was similar in strength (16 to 30 cm² reduced displacement) to that recorded at other volcanoes (fig. 12).

All three eruptions were similar to each other in their durations, volumes, ash-column heights, chemistry, and other parameters (table 1, fig. 11; Alaska Volcano Observatory, 1993). Tremor during the eruptions was also similar. However, patterns of tremor occurrence before and after the eruptions varied considerably. In June, most tremor preceded the eruption; we suggest that this tremor was chiefly hydrothermal in origin. No tremor followed the June eruption. In August, only a single short tremor episode occurred about 1 hour before the eruption. We suggest that the hydrothermal system had not yet been fully re-established, or had been re-established with different geometry or boundary conditions. In September, weak tremor occurred for 3 hours before the eruption, but in contrast to June and August, tremor occurred after the eruption continuously for 1 week. Because the hydro-

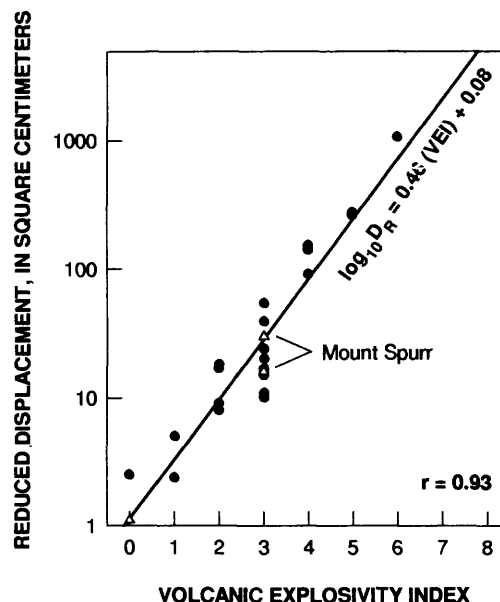


Figure 12. Volcanic tremor reduced displacement versus the Volcanic Explosivity Index (VEI) of Newhall and Self (1982). Data represent 24 eruptions at 15 different volcanoes. The plotted line is a linear regression fit to the data excluding Mount Spurr volcano in southwestern Alaska and has a correlation coefficient r of 0.93. Representative values from the June 27 and August 18 eruptions of Mount Spurr, and from a tremor burst, shown as triangles. Modified from McNutt (1994).

thermal system had been seriously perturbed by the eruption, we suggest that continued degassing of unerupted shallow magma was the source of this tremor. Similar conditions and tremor occurred from October 1 to 20. The last tremor on November 9–10 may be different than most other tremor at Mount Spurr; its frequency content was higher and its amplitude distribution between stations was different. However, thorough analyses of this last tremor are not yet complete.

Several of the observations made during this preliminary work have implications for our monitoring efforts. For example, spectra are routinely computed for tremor and there are many published examples in the literature. Nevertheless, our observation of the systematic change in the amplitude of alternate peaks appears to be unique and we are not aware of any published examples. This is clearly a source effect, and it may be possible to exploit this feature in the future to improve eruption forecasting. The shifting of the spectral peaks has been observed elsewhere, but the Mount Spurr observations occurred over a longer time scale and when eruption conditions had clearly changed. Future analyses will elucidate the timing more precisely. SSAM was not useful for determining the

shifting of the spectral peaks and the amplitude variations of alternate peaks because the frequency bins were fixed and did not coincide with the spectral peaks.

Several of our most important observations depended on data from individual seismic stations. Station CPK, which was located only 400 meters from the vent, proved to be especially valuable. This station enabled us to identify the occurrence of tremor bursts in early June when they were small enough that they were not clearly recognized on more distant stations. This station also enabled us to determine dramatic changes in the rms amplitude ratios (fig. 5), which we believe were most likely caused by a shoaling of the source. The station CPK was destroyed twice during eruptions, and its replacement was worth the cost because of the insight into physical characteristics of tremor we obtained from this key station. We suggest that every monitored volcano would benefit from having at least one station within a few hundred meters of the active vent.

Our analyses and monitoring efforts greatly benefited from having a variety of tools at our disposal, including helicorders, continuous digital data, event-triggered data, RSAM, and SSAM. Our experience in forecasting eruptions at Mount Spurr illustrates the importance of closely monitoring volcanic tremor using a variety of techniques and methods. Future work at Mount Spurr will focus not only on more detailed analyses of the data presented here, but also on improvements in the methods for recording and analyzing tremor in general.

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Appendix 1. Tremor bursts measured at seismic stations CPK and CRP between June 6 and June 26, 1992, Mount Spurr volcano, Alaska.

[All dates expressed in Universal Time; duration expressed in seconds; amplitude expressed in millimeters]

Time (UT)	CPK duration (sec)	CRP amplitude (mm)	Time (UT)	CPK duration (sec)	CRP amplitude (mm)	Time (UT)	CPK duration (sec)	CRP amplitude (mm)
June 6, 1992			June 10, 1992—Continued			June 14, 1992		
15:00	190	5.5	07:19	90	2.0	04:19	100	3.0
16:57	130	4.0	08:28	160	3.0	06:24	200	3.0
18:16	80	5.0	08:41	160	2.0	09:00	120	2.0
20:43	90	3.5	10:54	120	6.0	09:39	110	1.0
22:04	160	2.5	11:39	110	5.0	16:32	80	1.5
June 7, 1992			13:04	170	7.0	17:39	90	3.5
01:26	190	1.5	14:09	150	6.0	June 15, 1992		
02:28	130	5.0	15:25	70	1.5	00:08	260	1.5
04:58	220	3.0	15:31	120	5.0	04:05	90	5.0
05:47	80	3.5	15:44	110	3.0	12:48	160	<1.0
07:26	180	2.5	16:55	70	3.0	13:16	120	1.0
08:23	200	2.0	17:09	150	3.0	13:50	150	1.5
09:40	170	2.5	17:20	160	2.0	14:04	70	1.0
10:58	180	1.5	19:38	110	1.5	19:24	100	6.0
11:56	180	1.0	20:43	190	5.0	June 16, 1992		
13:38	190	4.0	21:52	230	6.0	08:44	250	1.0
14:04	100	1.0	June 11, 1992			17:39	230	3.0
14:47	200	3.0	00:16	100	4.0	20:50	90	0.0
15:29	110	3.5	00:31	70	3.0	June 17, 1992		
15:48	140	3.0	00:35	140	2.5	20:16	110	3.5
16:06	110	2.5	00:45	180	1.5	June 18, 1992		
18:24	140	3.5	02:53	180	1.5	14:07	120	2.0
19:01	170	2.5	03:48	100	2.5	June 19, 1992		
19:35	140	1.0	03:50	110	2.5	04:57	370	5.0
June 8, 1992			04:39	100	10.0	18:35	260	2.0
06:30	320	3.0	07:05	70	2.0	19:08	110	4.0
07:58	310	1.0	07:29	150	3.0	19:36	180	2.0
6/8/92			14:26	120	2.0	20:02	360	3.5
09:38	250	5.5	14:45	150	2.0	June 20, 1992		
09:46	80	1.5	15:03	230	5.0	10:54	220	<1.0
10:24	240	4.0	16:07	100	2.0	20:59	550	12.0
11:00	120	1.0	17:15	>90	1.5	June 21, 1992		
11:36	130	3.5	17:18	100	4.0	23:05	130	2.0
12:02	100	2.0	18:09	120	4.0	June 22, 1992		
12:35	90	3.0	18:33	240	2.0	03:25	320	3.5
12:52	80	2.0	20:04	90	3.5	03:50	80	<1.0
14:17	130	2.0	21:07	220	3.5	05:49	110	3.0
15:04	170	1.5	21:34	90	2.5	07:19	270	2.0
15:53	160	1.5	21:36	110	3.0	08:35	130	2.0
17:13	170	3.0	22:08	110	3.5	09:29	210	2.0
17:40	170	2.0	23:47	110	3.5	09:47	100	2.0
18:28	130	1.5	June 12, 1992			16:45	600	6.0
19:52	180	5.0	01:01	230	1.0	22:16	370	1.5
21:38	>150	5.0	03:17	250	6.0	22:47	190	3.5
June 9, 1992			05:59	180	2.0	June 23, 1992		
00:58	270	5.0	06:25	220	1.5	14:28	220	6.5
01:16	220	6.0	10:49	240	2.5	22:06	80	10.5
01:26	~150	2.0	11:04	200	2.5	22:09	210	5.0
02:02	80	4.0	11:12	80	2.5	22:33	170	3.5
02:37	100	3.0	15:10	130	3.0	June 24, 1992		
03:22	90	3.0	15:23	180	1.0	06:01	260	5.0
04:55	190	4.0	20:47	360	4.0	15:09	350	6.5
15:27	>120	2.5	21:00	220	4.0	23:34	9,240	5.0
15:49	160	1.5	21:08	80	4.0	June 25, 1992		
16:02	180	2.5	June 13, 1992			06:52	350	5.0
16:33	80	2.0	00:15	200	3.0	11:13	310	5.5
18:07	150	2.0	00:45	210	22.	11:32	8,520	4.5
20:29	170	3.5	09:27	190	2.5	15:58	240	7.5
22:15	140	3.5	09:30	110	2.0	20:11	130	9.0
June 10, 1992			09:43	150	1.5	23:55	360	7.0
01:01	160	4.0	10:09	120	1.0	June 26, 1992		
03:04	190	2.0	10:29	70	1.5	04:51	450	10.
04:56	250	6.0	11:29	150	3.0	05:36	150	2.0
05:03	140	2.0	15:03	230	1.5	11:16	390	3.5
06:12	130	8.0	18:04	120	3.5	20:03	68,460	~10.0
			23:10	100	2.0			
			23:39	120	1.0			

Lightning Associated with the August 18, 1992, Eruption of Crater Peak Vent, Mount Spurr Volcano, Alaska

By John F. Paskievitch, Thomas L. Murray, Richard P. Hoblitt, and Christina A. Neal

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ABSTRACT

The Alaska Volcano Observatory's lightning detection system detected and located 171 lightning strokes during the August 18, 1992, eruption of the Crater Peak vent of Mount Spurr volcano in Alaska. The strokes, predominantly intracloud, were detected during a 70-minute interval that began more than an hour into the eruption. All detected strokes were of positive polarity. The spatial distribution of the strokes in the horizontal plane defines a ring-like pattern approximately 10 km in diameter and displaced roughly 5 km to the east of Crater Peak. Although lightning was observed during the September 16–17, 1992, eruption of Crater Peak, no lightning was detected by the lightning detection system.

INTRODUCTION

Although the occurrence of lightning in volcanic eruption clouds is well documented, few attempts have been made to use lightning detection and location to monitor eruptions. Such a monitoring approach could potentially allow for the detection of an ash cloud even when meteorological conditions might prevent observations from satellites and ground-based radar. In 1990, the Alaska Volcano Observatory (AVO) experimented with a lightning detection system (LDS) used by the Bureau of Land Management (BLM) in their forest fire program. BLM's network is configured to locate typical meteorologic lightning strikes

that could potentially cause forest fires. A subset of BLM's network was modified and configured to enhance the potential to detect and locate volcanogenic lightning caused by the eruptions of Redoubt Volcano, Alaska. The system successfully detected and located volcanogenic lightning during several eruptions (Hoblitt, 1994). In 1991, AVO installed its own LDS and configured it to detect lightning in the Cook Inlet area.

The typical lightning discharge, or flash, is composed of a number of strokes. The series of strokes associated with a lightning flash occur within a few tenths of a second. A lightning discharge is termed a "strike" only if it makes contact with some conductive surface such as the Earth, trees, or aircraft. Each discrete stroke of lightning generates a broadband radio signal that radiates at the speed of light in an omnidirectional pattern.

The LDS at AVO incorporates lightning stroke detectors in the Alaska cities of Palmer, Iliamna, and Homer (fig. 1). These detectors monitor a portion of the radio frequency spectrum for a lightning-induced instantaneous rise in signal amplitude. A stroke that generates a radio signal with an amplitude exceeding the detector's threshold will be recorded at each detector at a different instant in time. The time of arrival of a signal is measured and recorded to sub-microsecond resolution. Using these precise times of arrival and the known locations of the detector sites, a lightning stroke's location can be determined.

A central analyzer in Anchorage calculates the locations of strokes. Communication between the analyzer and the remote detectors is via modem and phoneline (fig. 2). At regular intervals, the central analyzer automatically dials and queries each of the three remote detectors. Data is downloaded by the central analyzer and quickly correlated to determine if lightning strokes were detected. Various parameters are determined for each stroke including amplitude, polarity, and location.

There are two primary modes of recording LDS data: in a log file, and in the video information system (VIS). Processed and raw data (including system-status reports) are automatically transferred to and

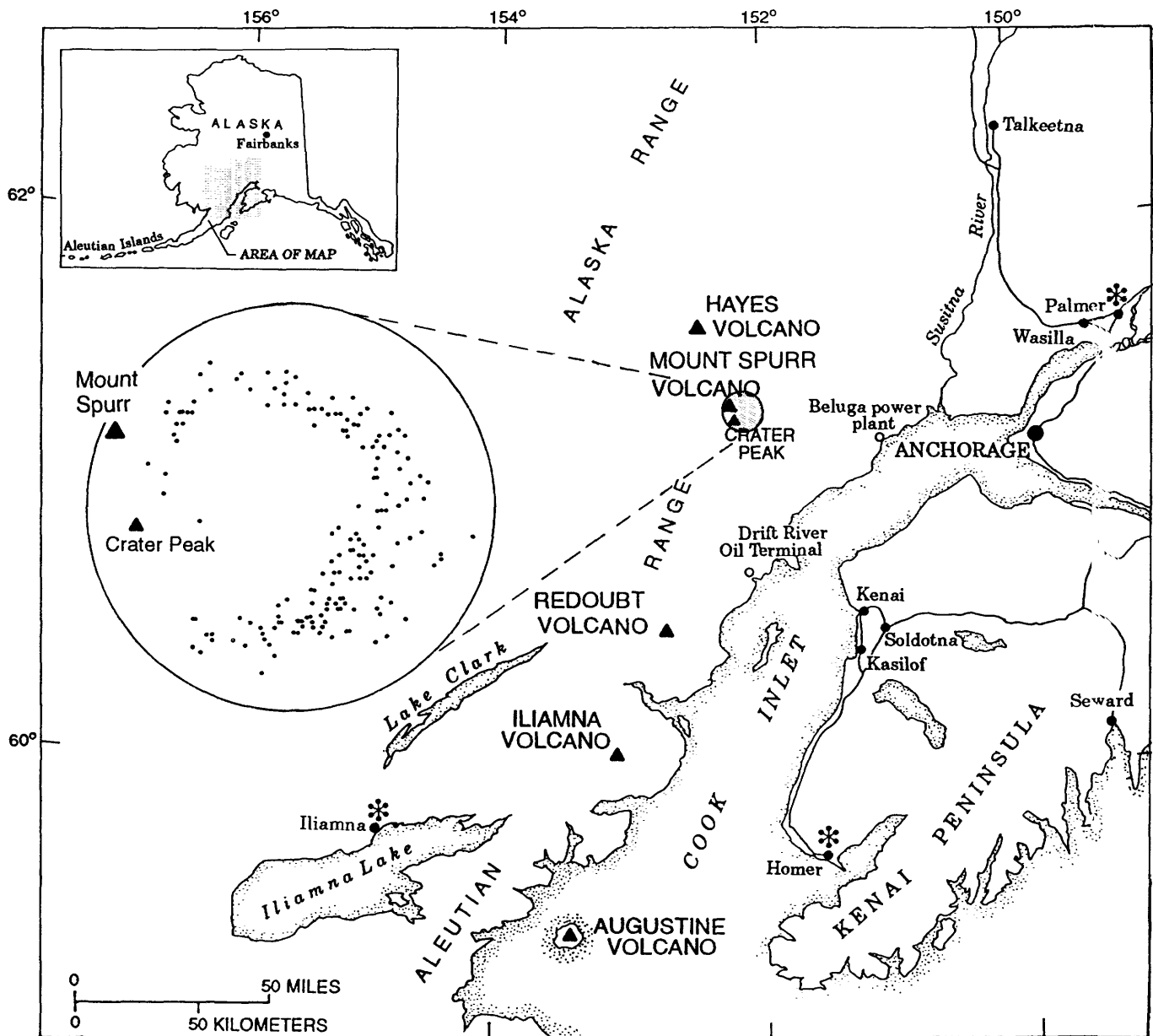


Figure 1. Cook Inlet area and locations of the active volcanoes. Asterisks show locations of LDS receivers. Enlarged view (inset) of area surrounding Mount Spurr shows 171 lightning strokes.

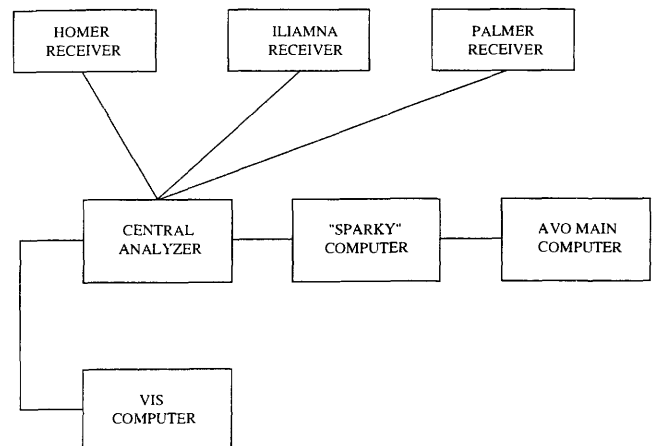


Figure 2. Data storage and communication links for Alaska Volcano Observatory's (AVO) lightning detection system. The central analyzer queries (via modem and phone line) each remote receiver at user-defined intervals. Incoming data are stored in the central analyzer and the computer named "Sparky." Data can be manipulated and viewed graphically through the video information system (VIS) and the AVO main computer.

stored in a log file on an associated computer. This file essentially contains the entire central analyzer output. The VIS data are initially stored in a buffer in the central analyzer. Before lightning data are sent to the VIS, a stroke must meet the criteria of being regionally comparable to, and occurring within 5 minutes of a preceding stroke. These criteria are designed to ensure the reliability of data. The VIS provides a graphic display of stroke locations and allows access to various quantitative stroke parameters.

THE 1992 ERUPTIONS OF CRATER PEAK

Mount Spurr volcano is located 125 km west of Anchorage at the northeastern end of the Aleutian arc (fig. 1). Following 10 months of elevated seismicity, the volcano erupted explosively on June 27, August 18, and September 16–17, 1992. The site of the three eruptions was Crater Peak, a satellite cone at 2,300 m elevation on the southern margin of the caldera and 3.2 km south of Mount Spurr's summit cone (Nye and Turner, 1990).

The three eruptions of Crater Peak were similar in character. The eruptions were explosive with sudden and impulsive onsets to each main phase. Duration of the main phase of each eruption were 3.5 to 4 hours, material ejected largely consisted of andesitic tephra, dense-rock equivalent erupted volumes for each of the eruptions were similar (Neal and others, this volume), and maximum radar-discernible column heights were approximately 13 to 15 km above sea level (Rose and others, this volume).

LIGHTNING RECORDED DURING THE AUGUST 18 ERUPTION

Although lightning was visually observed during the August and September eruptions, the LDS recorded lightning only during the August 18 eruption. The main phase of the August 18 eruption began at 4:42 p.m. Alaska daylight time (ADT). The LDS was unable to establish communication with the detector at Iliamna and therefore was not fully operational prior to 6:30 p.m. ADT. This prevented direct comparison of the onset of the eruption with the onset of lightning. Although the LDS first detected and located lightning at 6:30 p.m. ADT, it is known that lightning did occur before this time. First, an independent LDS operated by the BLM and designed to detect only ground strokes, or strikes, recorded 14 strikes associated with the eruption between 5:43 p.m. ADT and 7:51 p.m. ADT (Keith Pollock, BLM, oral commun., 1992). Sec-

ondly, the two operational detectors at Homer and Palmer recognized common events believed to be lightning strokes associated with the eruption as early as 5:30 p.m. ADT. Third, airborne observers at the volcano saw lightning at 5:45 p.m. ADT (G. McGirsey, USGS, oral commun., 1992).

Once communication to all three detectors was established, the LDS recorded 171 strokes during a 70-minute interval. The horizontal spatial pattern of recorded stroke locations forms a crude circle roughly 10 km in diameter (fig. 1). There is a concentration of strokes on the eastern side of the circle. The center of the circle of strokes is displaced approximately 5 km in a direction 110° (true north) from the eruptive vent. This displacement is consistent with the migration of the eruption cloud resulting from the westerly winds during the eruption (Neal and others, this volume).

At the time of lightning detection, the eruption column had extended to its maximum altitude of about 14 km on the basis of C-band radar estimates (Rose and others, this volume), or 18 km (unpublished pilot reports). Satellite imagery near this time depicts an elongate plume extending to the east. The National Weather Service's C-band radar provides a cross-sectional view of the plume (not the eruption column) displaced to the east of the vent. The image shows zones that are defined by regions of like particle size (Rose and others, this volume).

The spatial pattern made by the stroke locations is very uniform, which suggests that it is not a random function. Stroke occurrence and location are most likely controlled by concentrations of particle-size fractions or particle densities within the plume that in turn control charge densities and polarities.

Analyses of stroke signals suggest that more than 70 percent of the strokes were intracloud, and the remaining 30 percent were ground strokes, or strikes. However, caution must be exercised when considering the distinction between these intracloud and cloud-to-ground strokes. The LDS determines stroke type with a model that is based on typical meteorologic lightning. The time between peak stroke signal and the first polarity change is measured. If this time interval is less than 10 microseconds, the stroke is classified as intracloud. If the time interval exceeds 10 microseconds, a ground stroke is indicated. However, volcanogenic lightning might behave quite differently than meteorologic lightning in this regard. For instance, the conductive ionized channel that precedes volcanogenic lightning could perhaps be a slower path than that for meteorologic lightning (Rodney Bent, Atmospheric Research Systems, Inc., oral commun., 1992). If this were the case, some volcanogenic intracloud strokes could be misclassified as ground strokes.

Stroke amplitudes, calculated by AVO's central analyzer, for the 50 detected ground strokes ranged from 15 kA to 40 kA with an average amplitude of 22 kA. These values fall within the range of ground strokes observed in local area thunderstorms but approach neither the low nor high values that describe this observed range (3.5 kA to 350 kA). The LDS is unable to calculate amplitudes for intracloud strokes.

UNDETECTED LIGHTNING

Lightning was neither detected nor observed during the June eruption. Although the occurrence of lightning was observed during the September eruption, none was detected by the LDS even though the system appeared to be operational. Indeed, the system probably did not detect all of the lightning that occurred during the August 18 eruption, even when the system was fully operational. There are several possibilities for this. The simplest explanation is that the signal from intracloud lightning has a lower amplitude than cloud-to-ground lightning. Data from the August 18 eruption indicate that intracloud lightning predominates. Possibly the typical signal levels from these strokes fall below the detection threshold of the LDS. The subset of strokes recorded for the August 18 eruption may have produced greater signal strengths and were therefore the only ones detected.

Considering the similarities between the three eruptions, it is puzzling that strokes were detected for only the August eruption. Perhaps subtle factors are involved in an eruption that would raise the stroke amplitudes or affect other stroke characteristics that would allow for detection. Also, atmospheric and meteorologic conditions at the time of an eruption may influence volcanogenic lightning characteristics.

SUMMARY

An LDS was installed at AVO to further explore the potential to determine the occurrence of volcanic ash clouds through the detection of associated lightning. Such a monitoring approach would allow for the detection of an ash cloud even when meteorologic conditions might prevent observation from satellites and ground-based radar. During one of the three 1992 Crater Peak eruptions, it was demonstrated that the system could function as anticipated. The LDS was able to detect and locate at least a subset of the lightning associated with the August 18 eruption. The circular pattern of the stroke locations give an indication of both size and movement of the ash cloud.

Although lightning was observed during the September eruption, none was detected by the LDS. The LDS did not record all of the lightning associated with the August eruption. A plausible explanation for this is that the relatively low-amplitude intracloud lightning that was predominant during these eruptions did not, with the exception of the 171 lightning strokes recorded during the August eruption, generate strong enough signals to be detected at all three detector locations. Installing a fourth detector closer to the monitored volcanoes may enhance the ability to detect volcanogenic lightning during future eruptions of Cook Inlet volcanoes.

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Lahars from the 1992 Eruptions of Crater Peak, Mount Spurr Volcano, Alaska

By David F. Meyer and Dennis C. Trabant

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ABSTRACT

Lahars, which are flows of volcanic sediment and water, are commonly the most immediate hydrologic hazard resulting from volcanic eruptions. Lahars often change character as they flow downstream, and a single flow can behave as a debris flow at one time and place and as streamflow at another time and place. Crater Peak, a vent on the flank of Mount Spurr volcano, erupted on June 27, August 18, and September 16-17, 1992. Lahars generated during these eruptions illustrate the effects of clay content on lahar behavior, processes of flow transition, and thresholds of glacier slope and surface roughness, below which lahars are not generated.

On June 27, snow, ice, and rock avalanches were triggered by the eruption and flowed down several steep gullies on the south flank of Crater Peak. These flows became debris flows and entered the Chakachatna River in three places. During the August 18 eruption, at least four pulses of hot pyroclastic flow covered a small area on the middle part of Kidazgeni Glacier and flowed down three gullies on the south flank of Crater Peak, but no lahars resulted and no flows entered the Chakachatna River. On September 16-17, pyroclastic flows covered and traversed a larger area on Kidazgeni Glacier and eroded snow and ice along steep, crevassed icefalls then coalesced near the toe of the glacier to form a debris flow that deposited most of its sedi-

ment along a more gently sloping reach. The debris flow then eroded more material as it descended a narrow gully, and the resulting deposit at the mouth of that gully dammed the Chakachatna River.

Comparison of the behavior of the 1992 Crater Peak lahars with lahars that resulted from the 1980 eruption of Mount St. Helens, Washington, indicates that if less than 1 percent clay-size material is present in the flow, coarse-grained material is deposited on relatively steep slopes, and debris flows do not extend as far downstream as do debris flows with larger proportions of clay-size material.

INTRODUCTION

Lahars often present the most serious of all volcanic risks in populated areas. The hazard these flows of sediment and water represent depends on their character. Lahars can range from Newtonian, sediment-laden streamflow to non-Newtonian debris flow and can change as they move downstream. The character of the flows depends on initiating mechanisms, nature of the material incorporated in the flow, water content, and changing channel and valley geometries. Although numerous lahars have been documented, especially since the devastating eruptions of Mount St. Helens, Washington, and Nevado del Ruiz, Columbia (see Janda and others, 1981; Pierson and others, 1990; Scott, 1988; 1989), documentation of the range of factors that produce lahars remains incomplete.

Lahars were produced during two of the three 1992 eruptions of Crater Peak, a vent on Mount Spurr volcano. During the June 27 eruption, avalanches of snow and volcanoclastic debris flowed down Crater Peak creek (*informal name*) and Wilton's Walk gully (*informal name*; fig. 1) and evolved into debris flows. During the September 16-17 eruption, pyroclastic flows eroded and incorporated snow and glacier ice on Kidazgeni Glacier, and they transformed into debris flows that eventually dammed the Chakachatna River at the mouth of Bench gully (*informal name*; fig. 1). During the August 18 eruption, pyroclastic flows extended down the gullies on the upper flanks of Crater

Peak and onto Kidazgeni Glacier and produced small meltwater flows that did not evolve into debris flows.

In both eruptions in which debris flows were generated (June 27 and September 16–17), the volume and character of the flows changed as older volcanoclastic material was entrained and as sediment was deposited along the flow path. Most of the sediment was deposited near the base of the volcano, and the lahars did not extend far downstream.

Lahars generated during the July 9, 1953, eruption of Crater Peak (Juhle and Coulter, 1955) affected the same drainages that were affected during the 1992 eruptions and were about the same size. They flowed down Crater Peak creek, across an area known informally as "the Bench," and down Bench gully and an unnamed stream draining from Kidazgeni Glacier, 2 km downstream from Bench gully (figs. 1 and 2). The Crater Peak creek lahar deposited debris that formed

a fan in the channel of the Chakachatna River that dammed the river and formed a lake. The lake still existed in 1995, although it was not as large. Smaller fans were deposited at the mouths of Bench gully and the unnamed stream 2 km farther downstream. Apparently, considerable tephra was deposited on Kidazgeni and Straight Glaciers, but little, if any, tephra was deposited on Crater or Barrier Glaciers. The summit crater of Crater Peak was filled with snow and ice prior to the 1953 eruption. The snow and ice were completely removed by the 1953 eruption.

Currently no structures or people are at risk from flows down the south flank of Mount Spurr. However, hydroelectric and hydrothermal development and extensive coal mining have been or are being considered in the region. The significant hazard presented by lahars from Mount Spurr is integral to land-use planning for areas surrounding the volcano.

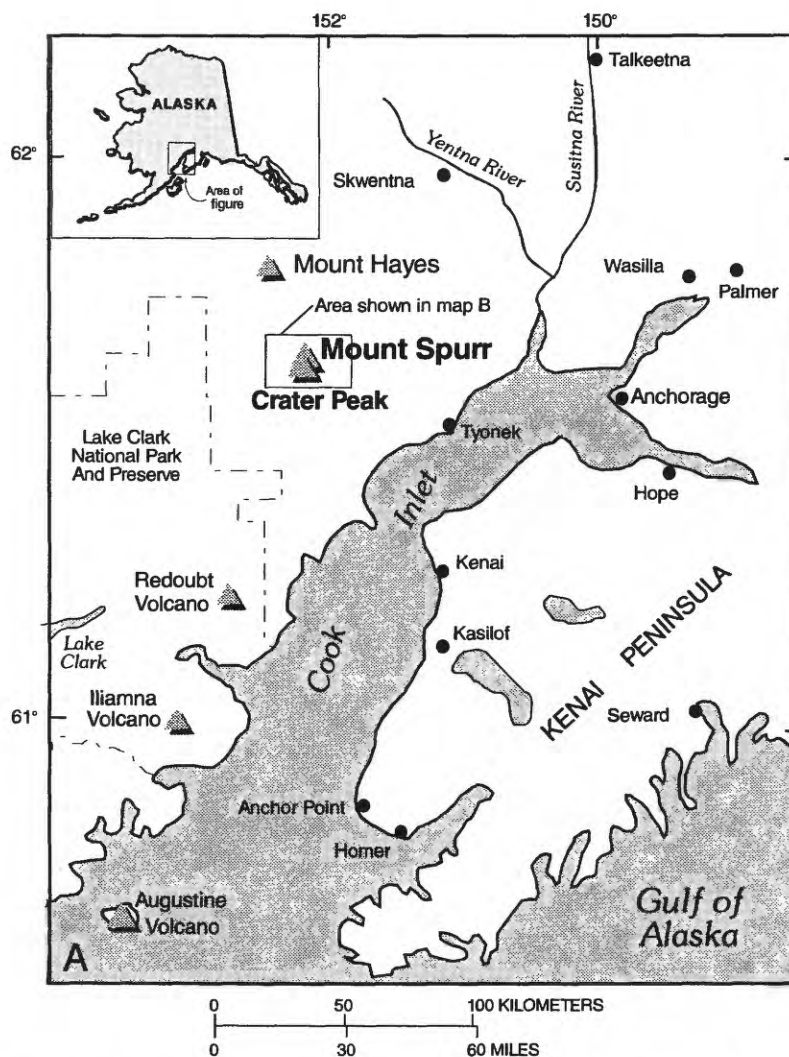


Figure 1. A, Location of Mount Spurr volcano and its active vent, Crater Peak, in southwestern Alaska.

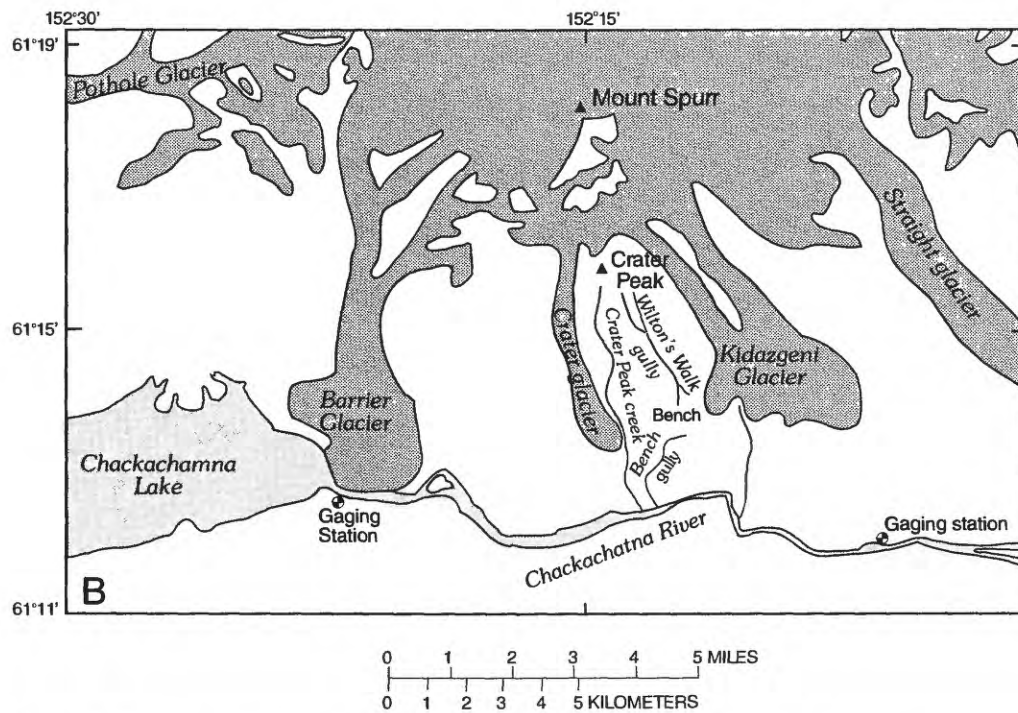


Figure 1. Continued. B, Hydrologic features affected by 1992 eruptions of Crater Peak.

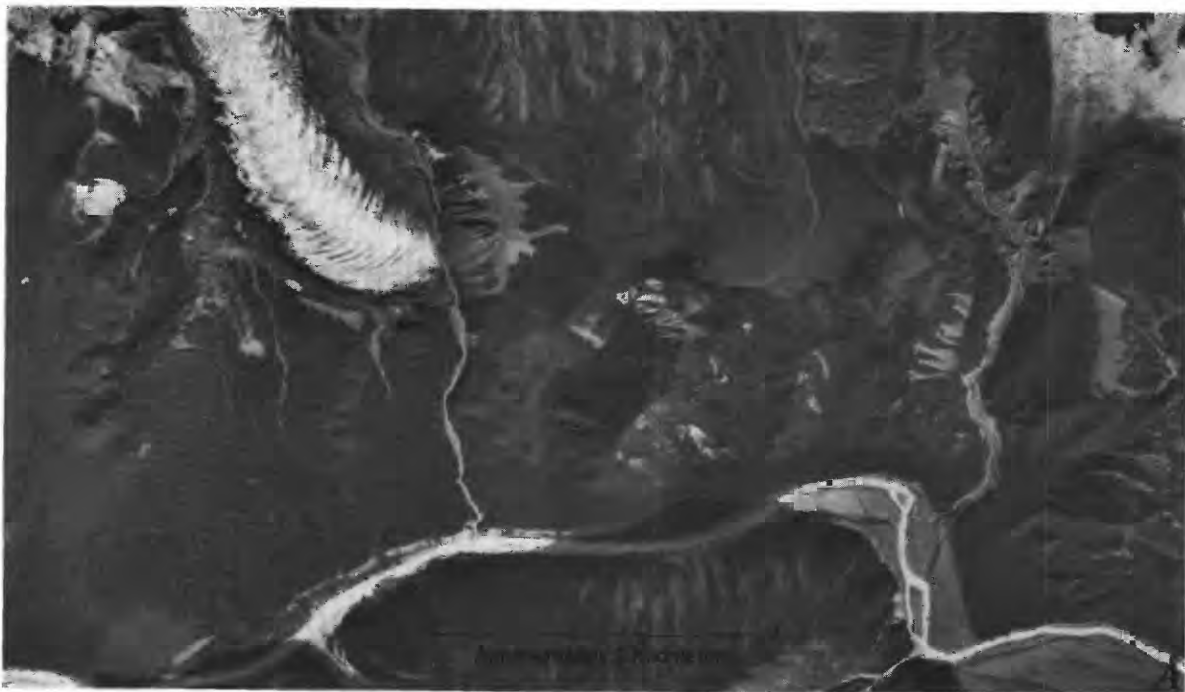


Figure 2. A, South flank of Mount Spurr volcano, southwestern Alaska, on September 2, 1952.



Figure 2. Continued. *B*, South flank of Mount Spurr volcano, southwestern Alaska, on August 29, 1954. Lahar deposits from July 9, 1953, eruption of Crater Peak visible in center of photograph.

PHYSIOGRAPHY AND HYDROLOGY

About 67 km³ of snow and ice, covering 360 km², are stored in the glaciers on Mount Spurr. The main caldera of Mount Spurr is filled by snow and glacier ice and is the source of Kidazgeni Glacier, which flows around the east side of Crater Peak, and of Crater glacier (*informal name*), which flows around the west side of Crater Peak (fig. 1). Prior to the 1953 eruption of Crater Peak, the crater at its summit was filled with snow and ice that extended a short distance down the flanks (Juhle and Coulter, 1955). Barrier Glacier is the largest glacier in the Chakachatna River drainage and partly obstructs the river southwest of Mount Spurr; this obstruction resulted in the formation of Chakachamna Lake. Barrier Glacier heads on the west flank of the main caldera of Mount Spurr, about 5 km northwest of Crater Peak.

The Chakachatna River drains the west, south, and east flanks of Mount Spurr, including the areas most affected by the 1992 eruptions of Crater Peak. The largest of numerous canyons and small gullies on the flanks of Crater Peak is Crater Peak creek, a deep, V-shaped gorge cut into pyroclastic and laharic deposits from the Crater Peak vent (Nye and Turner, 1990). Several hot springs are present along the creek. Stream slopes of all of these drainages are steep, rang-

ing from 0.3 to 1.1 meters per meter (m/m) in Crater Peak creek within the first 3 km from the rim of the crater (fig. 3). At an altitude of about 520 m, Crater Peak creek is bounded by the toe of Crater glacier on the west and by a late Pleistocene debris-avalanche deposit on the east. At this point, channel slopes decrease and range from 0.04 to 0.09 m/m. Wilton's Walk gully, which drains the south-southeast flank, ranges in slope from 0.3 to 0.8 m/m upstream from the point where it drains onto the Bench, a gently sloping plain underlain by late Pleistocene debris-avalanche deposits and late Holocene outwash from Kidazgeni Glacier. Slopes across the Bench range from 0.04 to 0.2 m/m. The east flank of Crater Peak is drained by a stream that flows on and along the margin of Kidazgeni Glacier, and then through a breach in the Kidazgeni lateral moraine onto the Bench. The Bench is drained by Bench gully, a steep (0.1 to 0.6 m/m) gully that flows onto an alluvial fan in the Chakachatna River valley.

Mean annual precipitation near Mount Spurr ranges from less than 1,000 mm near the coast to more than 2,000 mm along the crest of the north end of the Aleutian Range (National Weather Service, 1972); on Mount Spurr, most of the precipitation falls as snow. The snowpack reaches its maximum depth in late May or early June.

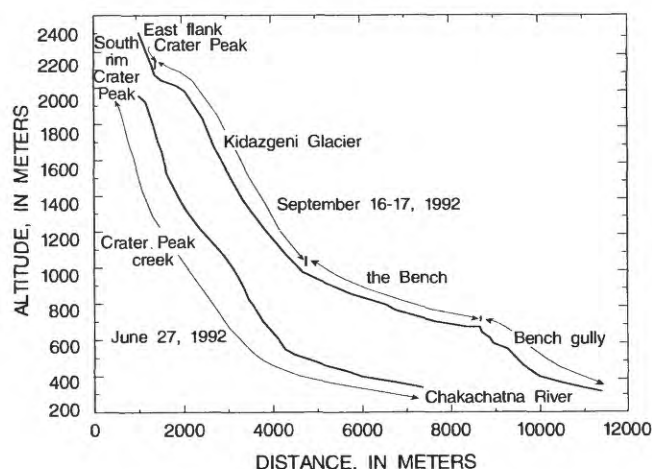


Figure 3. Longitudinal profiles for paths of June 27, 1992, lahar and September 16-17, 1992, lahars produced during eruptions of Crater Peak vent, Mount Spurr volcano, southwestern Alaska.

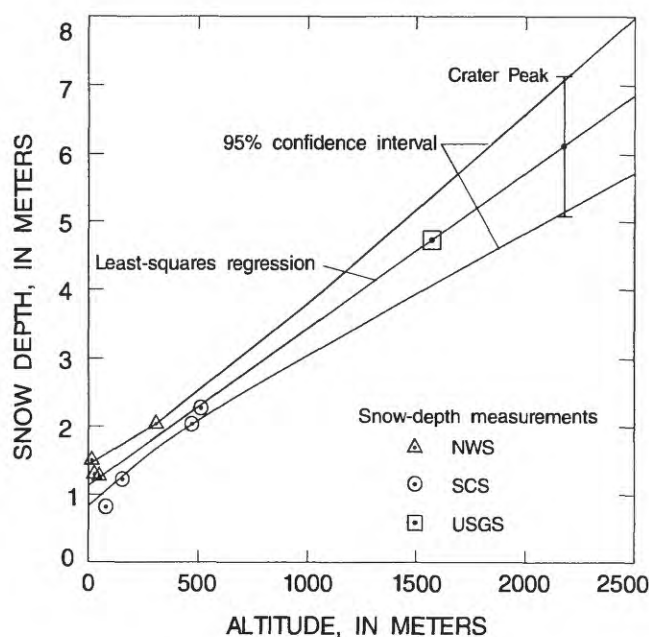


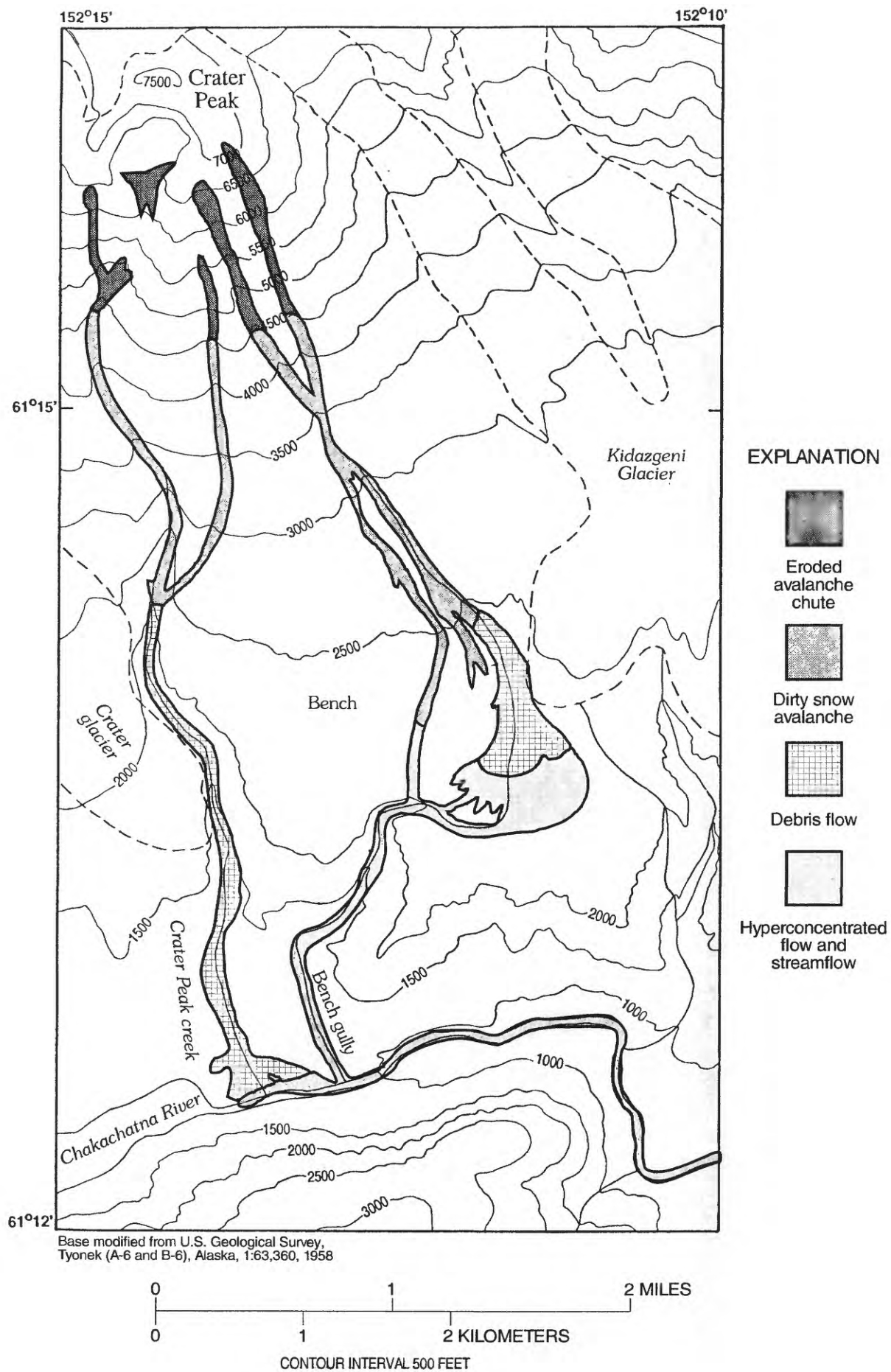
Figure 4. Maximum snow depth versus altitude for four climatological National Weather Service (NWS) stations and four Soil Conservation Service (SCS) snow courses closest to Mount Spurr volcano, southwestern Alaska, and one snowpack measurement from a site on Double Glacier, about 75 km south of Crater Peak that was taken by U.S. Geological Survey (USGS) scientists. Maximum snow depth was measured during March of 1992 at the low-altitude (NWS and SCS) sites and on April 10, 1991, at Double Glacier site. Extrapolation of regression based on these data indicates that snowpack on south rim of Crater Peak was 6 ± 1 m thick.

The mean annual discharge from the 2,900 km² basin upstream from the mouth of Chakachamna Lake was 107 m³/s from 1959 to 1972, when a streamflow gaging station was operated there. Annual high flow is generated from glacier melt, and streamflow is generally greatest during August; the August mean monthly discharge is 340 m³/s (Scully and others, 1978). A flow with a probability of being equaled or exceeded once every 10 years has a magnitude of 595 m³/s (Lamke, 1979). The largest flood measured on the Chakachamna River had a peak discharge of 13,300 m³/s, and it occurred on August 11, 1971, when part of the Barrier Glacier ice dam was eroded. Barrier Glacier extends most of the way across the Chakachamna River valley. The Chakachamna River occupies a narrow canyon between the toe of Barrier Glacier to the north and the bedrock valley wall to the south. During the period from 1959 to 1971, Barrier Glacier had advanced faster than ice was removed from its toe, and the outlet to the lake became progressively smaller. During high flow generated by a regional rainfall, the toe of the glacier was eroded laterally, which enlarged the outlet of the lake and resulted in the release of about 2.9×10^8 m³ water (Lamke, 1972; Scully and others, 1978).

JUNE 27, 1992, LAHARS

The June 27 eruption occurred near the time of the maximum snow accumulation on Crater Peak. The snow depth on relatively flat surfaces, such as glaciers near 2,000 m altitude, is estimated to have been about 6 m of snow, on the basis of regional extrapolation (fig. 4). The water equivalent of that snowpack is estimated to have been 2.2 m on the basis of an average density (0.36 g/cm³) for the five snowpack measurements (made during April 1991 and March 1992) used to construct figure 4. However, snow was not evenly distributed prior to the eruption, and snow in the avalanche-source areas was probably deeper than 6 m; gullies were filled with snow that had avalanched or blown from the ridges. A small perennial snow and ice layer was present on the east inside wall of the crater prior to the eruption.

The eruption triggered large avalanches of snow, ice, and rock that flowed down the four major drainages on the south and southeast flanks of Crater Peak (fig. 5). Massive tephra deposits blanketed an area of about 0.08 km² on the crater rim, upstream from the source areas for the avalanches, but no specific evidence of pyroclastic flows was observed (see Miller and others, this volume, fig. 1). The avalanches apparently occurred during the later part of the erup-



tion, as the avalanche deposits were not extensively covered by fallout tephra, whereas adjacent areas were blanketed by fallout tephra deposits. The avalanches eroded snow, ice, and rock from a reach beginning near the south crater rim at 2,100 m to 2,000 m altitude, down to 1,400 m altitude (fig. 6), and they deposited material along the channel margins from 1,400 m to 600 m in altitude in Crater Peak creek, and to at least 800 m in altitude on the Bench. When fresh, the avalanche deposits had a characteristic dark color, caused by wet rock, relative to adjacent surficial material, and they contained abundant snow and ice particles. After the snow and ice melted and drained away, the avalanche deposits were light-colored, fluffy, and marked by rare small kettle-like pits. These deposits were scoured from the center of the channels by later flows.

The snow, ice, and rock avalanches and the later flows that scoured the center of the channels apparently evolved into debris flows downstream. Debris flows left extensive deposits below an altitude of 430 m along Crater Peak creek, and across the Bench from 800 m to about 700 m in altitude (fig. 5). The massive debris-flow deposits were overlain by well-sorted, waterlaid gravel that had been reworked either by streamflow that followed the debris flows or by flows generated locally as the debris-flow deposit dewatered.

Canyon walls and valley-floor alluvium were eroded along the steep, V-shaped Crater Peak creek canyon. Spring snowmelt probably had saturated the valley-floor alluvium prior to the eruption, and the eroded material combined with the dirty-snow avalanche to produce the debris flow that left deposits from an altitude of 430 m down to the Chakachatna River. A 0- to 2.5-m-thick, 100- to 200-m-wide diamict (a nonsorted deposit having a wide range of grain sizes) covered the lower 1 km of the channel and the alluvial fan at the mouth of Crater Peak creek (fig. 7). Most of the diamict was deposited below the break in slope between 520 m and 550 m in altitude. This lower reach, within 2 km of the mouth of Crater Peak creek, has a channel slope between 0.04 and 0.09 m/m (fig. 3).

At a location 0.5 km upstream from the mouth of Crater Peak creek, the debris flow had an estimated peak discharge of 2,000 m³/s. This estimate was made using the superelevation of the flow around a bend to estimate the velocity of the flow (Chow, 1959, p. 448). Mid-channel bars in the Chakachatna River at the mouth of Crater Peak creek were veneered by deposits of sand and gravel, but the river was not blocked as it had been during the 1953 Crater Peak eruption. The debris flow did not extend beyond the alluvial fan at the mouth of Crater Peak creek.

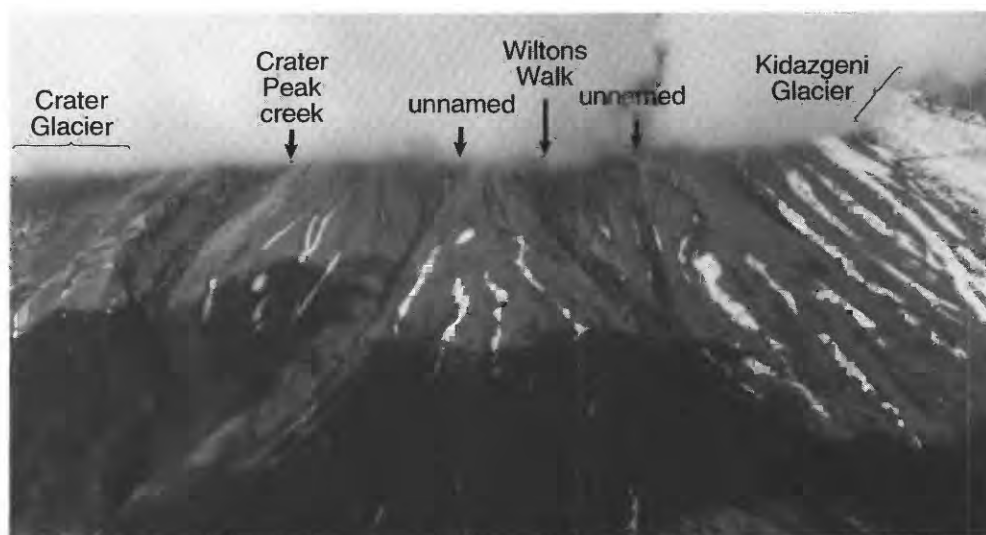


Figure 6. Four large drainage channels on south-southeast flank of Crater Peak vent of Mount Spurr volcano, southwestern Alaska, on June 29, 1992, 2 days after an eruption of Crater Peak. Dark deposits in channels are snow, ice, and rock avalanche deposits. Dark areas covering lower slopes are cloud shadows.

◀ **Figure 5.** Distribution of deposits from June 27, 1992, Crater Peak (Mount Spurr volcano, Alaska) avalanches and lahars associated with eruption. Dashed lines show approximate boundaries of glaciers.

In Wilton's Walk gully and two unnamed channels, neither the dirty snow avalanches nor the subsequent flows eroded significant amounts of material from the poorly defined channels. As the flows spread across the Bench, coarse sediment was deposited as a diamict sheet in proximal areas, indicative of a debris flow, but the deposit was thinner, bedded, and better sorted along its margins, indicative of sediment-laden streamflow. Individual units that could be correlated to the dirty snow avalanche or to the later erosive flows upstream could not be distinguished within the diamict on the Bench. The two types of flow evident upstream may have combined to form a single deposit on the Bench or the later erosive flows may not have left deposits on the Bench. Nearly all of the sediment in the flow was deposited on the Bench, and only relatively clear water flowed down Bench gully and into the Chakachatna River. At the mouth of Bench gully, the surface of a fan was covered by thin, waterlaid deposits.



Figure 7. June 27, 1992, debris-flow deposit resulting from eruption of Crater Peak vent (Mount Spurr volcano, Alaska) near mouth of Crater Peak creek, 100 m upstream from Chakachatna River. Notebook in center of photograph is 20 by 11 cm and rests on pre-eruption surface. Deposit is about 2.5 m thick.

Figure 8. Small meltwater channel along west margin of Kidazgeni Glacier, which is covered with wet, dark tephra deposits from August 18, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska. Northeast rim of Crater Peak is in left background. Photograph taken August 20, 1992.

AUGUST 18, 1992, WATERY FLOWS

Several small pyroclastic flows descended the Wilton's Walk gully on the southeast side of Crater Peak on August 18, but they stopped short of the debris-flow deposits emplaced during the June 27 eruption. Four or more overlapping pyroclastic flows descended the east flank of Crater Peak and covered 0.1 to 0.2 km² of the western half of the west arm of Kidazgeni Glacier (see Miller and others, this volume, fig. 1). Eight to twelve cm of fallout tephra was deposited beyond the margins of the pyroclastic-flow deposits, but none was deposited on the pyroclastic-flow deposits (see Miller and others, this volume; Neal and others, this volume).

The pyroclastic flows ran onto a relatively smooth glacier surface, except at the highest area of contact with Kidazgeni Glacier, near the extreme northeastern sector of Crater Peak, where some glacier ice was



mechanically entrained by the flows. Glacier slopes ranged from 0.21 to 0.26 m/m. Channels scoured by small watery flows were observed downstream from the multipulsed pyroclastic-flow deposits along both the east and west margins of the lower part of the west arm of Kidazgeni Glacier (fig. 8), but no debris-flow deposits were observed in direct association with either the pyroclastic flows or the watery flows. The water flows were probably generated by melting of small amounts of snow and ice in and under the pyroclastic-flow deposits. The small amount of snow and ice on Crater Peak prior to the August 18 eruption was confined to small patches in gullies. Water flowing down the east margin of the west arm of Kidazgeni Glacier ponded in depressions near the glacier terminus; water flowing down the west margin of the west arm of Kidazgeni Glacier flowed through a small gap in the lateral moraine onto the Bench. Cobble-size bed material was mobilized by these flows, and minor bank erosion occurred within a kilometer of the glacier, but the overall effects were comparable to those from annual glacier-melt generated flows.

SEPTEMBER 16–17, 1992, LAHARS

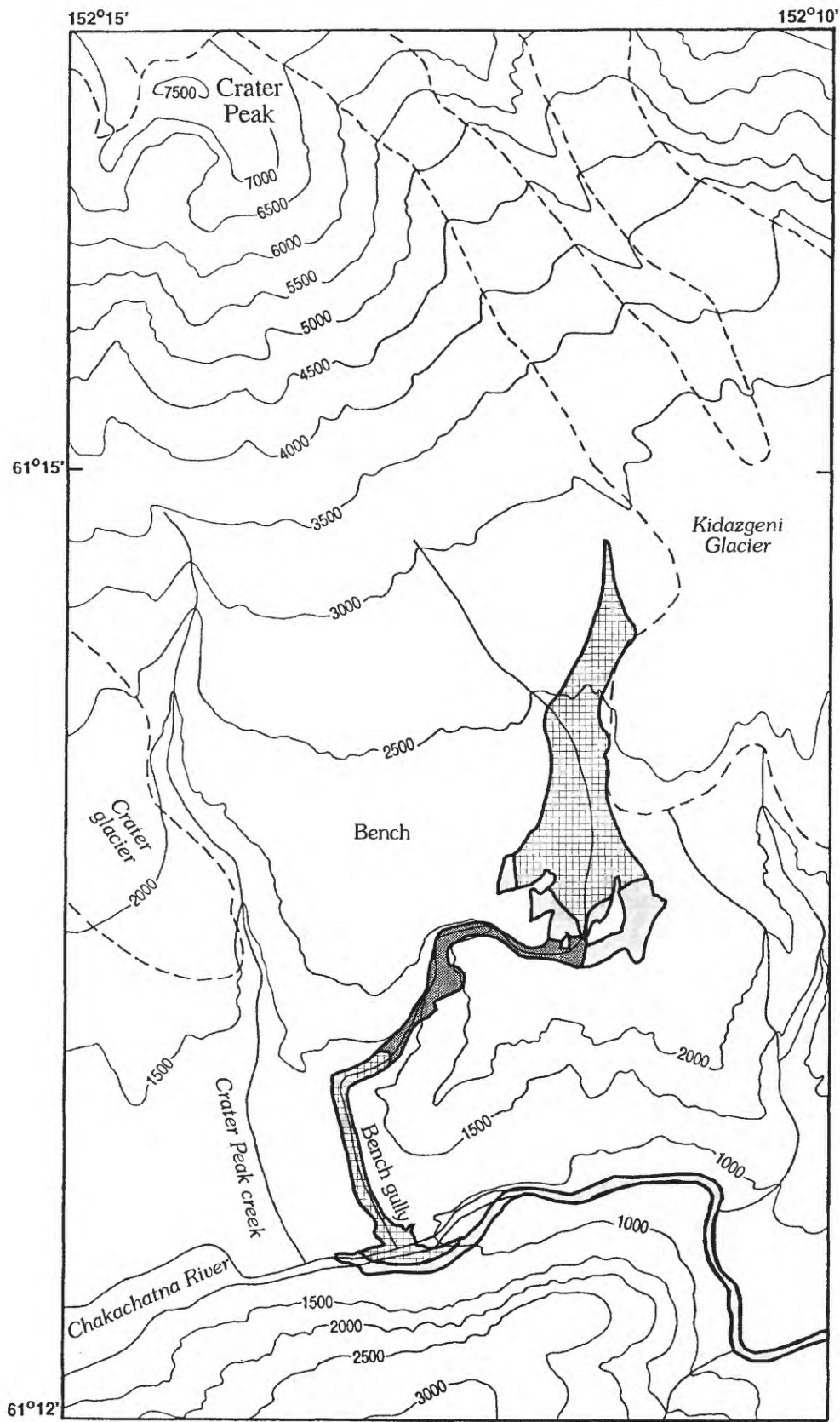
Pyroclastic flows generated during the September 16–17 eruption flowed down the northeast flank of Crater Peak and onto the west arm of Kidazgeni Glacier, near the upstream end of a steep icefall (see Miller and others, this volume). These pyroclastic flows were thicker, and they traversed a more extensive and steeper area of the glacier, covering about 0.3 km² of Kidazgeni Glacier having slopes ranging from 0.25 to 0.38 m/m. Much of this area was broken by crevasses. Because of the steeper slope, thicker flow, and greater surface roughness, the September 16–17 pyroclastic flows probably showed greater shear at their base than the flows that occurred during the August 18 eruption, and they eroded considerable snow and glacier ice. As the pyroclastic flows incorporated water melted from eroded snow and ice, they were transformed into a lahar that deposited both massive diamicts and bedded streamflow deposits along channels scoured into the glacier surface (see Waitt, this volume).

The lahar fanned across the Bench surface, widening from about 100 m on the proximal Bench to about a kilometer on the distal Bench (fig. 9). Most of the debris-flow phase of the lahar was deposited as a continuous, 1- to more than 2.5-m-thick diamict on the Bench surface. Steep bouldery flow-front lobes formed a few hundred meters upstream from the headcut of Bench gully, which drains the Bench sur-




face. The flow-front lobes are depleted in sand- and gravel-size sediment. Deposits of planar- or cross-bedded pebbles and sand cap the surface of the diamict (fig. 10), and they extend downfan from the bouldery flow-front lobes as splay deposits (fig. 11). In some places, the splay deposits extend only a few tens of meters into brushy vegetation; in other places, they extend hundreds of meters into gullies and channels that are tributary to Bench gully (fig. 9).

One lobe of the debris flow extended far enough downstream to flow into Bench gully in a channel 10 m wide, where it combined with the more watery flows. An estimated 10,000 to 100,000 m³ of material was eroded from the banks of Bench gully by the debris flow and subsequent incision during and following the night of September 16–17 when arcuate failures of 10- to 20-m-high banks as thick as 10 m occurred along a 1,500-m-long reach between the Bench and the Chakachatna River. Large bank failures were still occurring during the afternoon of September 17. An estimated 50,000 to 100,000 m³ of coarse debris was deposited as a 0- to 3-m-thick diamict over the 50,000 m² alluvial fan where Bench gully enters the Chakachatna River. This sand and gravel diamict contains boulders as large as 1 m in diameter, similar on a gross scale to the diamict on the Bench. Subsequent watery flows, probably during the waning stages of the lahar, reworked the surface of the fan and incised a channel 3 to 4 m deep and less than 100 m long along the centerline of the fan.

The diamict formed a temporary dam across the Chakachatna River, and so the river rose to depths of 0.5 to 1 m directly upstream from the fan (fig. 12). A 20-m-wide channel was eroded along the distal margin of the fan by the Chakachatna River; through this channel pulsed a flood wave with an estimated peak discharge of 800 m³/s just downstream from the debris fan. Peak discharge estimates (800 m³/s and 400 m³/s) were made using the slope-conveyance technique, which applies the Manning's equation to a single cross section to estimate the flow velocity (Chow, 1959, p. 98). The estimated discharge is larger than the largest measured, meteorologically generated flow on the Chakachatna River, but it was less than one-tenth of the peak discharge that resulted from erosion of the Barrier Glacier ice dam in 1971 (Lamke, 1972; Scully and others, 1978). Deposits downstream from the fan that mark the passage of the flood wave consist of medium- to coarse-grained, planar-bedded sand. The peak attenuated quickly from 800 m³/s at the downstream margin of the fan to about 400 m³/s at a location 6 km downstream, where flood deposits were fine- to medium-grained, planar-bedded sand.



EXPLANATION

-  Eroded gully
-  Debris flow
-  Hyperconcentrated flow and streamflow

Base modified from U.S. Geological Survey,
Tyonek (A-6 and B-6), Alaska, 1:63,360, 1958

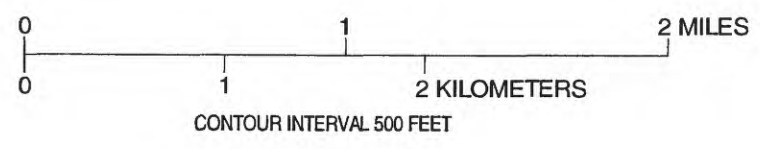




Figure 10. September 16–17, 1992, lahar deposits resulting from Crater Peak vent eruptions, Mount Spurr volcano, Alaska. Deposits are on proximal part of the Bench; they consist of a massive diamict (below dashed line) overlain by planar- and cross-bedded sand and gravel (above dashed line). Shovel in center of photograph is 0.6 m long.



Figure 11. September 16–17, 1992, splay deposits resulting from Crater Peak vent (peak issuing steam in the right background) eruptions, Mount Spurr volcano, Alaska. Deposits crop out along distal part of the Bench just upstream from head of Bench gully and consist of boulders depleted of fine material in the center and right foreground and well-sorted sand and gravel that form flat surfaces in left foreground.

◀ **Figure 9.** Distribution of deposits from lahars resulting from September 16–17, 1992, eruption of Crater Peak vent, Mount Spurr volcano, Alaska. Dashed lines show approximate boundaries of glaciers.



Figure 12. Alluvial fan at mouth of Bench gully, including eroded debris-flow dam (upper left outlined area) across Chakachatna River (light colored river in center). Extent of debris-flow and sediment-laden streamflow deposits indicated by solid lines, dashed where inferred. Channel at toe of alluvial fan (shown at bottom of photograph) is about 20 m wide. View is upstream.

GRAIN SIZE, LAHAR TRANSFORMATION, AND LAHAR GENERATION

The behavior of debris flows depends, among other factors, on the grain-size distribution of the material contained in the flow, especially the proportion of clay (Scott, 1988; Scott and others, 1992). Cohesive lahars from Mount St. Helens, Washington, both those that occurred in 1980 and at earlier times, and lahars from Mount Rainier, Washington, maintained fairly constant rheologic and sedimentologic properties when they contained more than 3 percent clay particles (those smaller than 0.004 mm in diameter). The May 18, 1980, North Fork Toutle River lahar from Mount St. Helens contained more than 3 percent clay, and it flowed as a coherent debris flow more than 100 km along slopes ranging from 0.007 to 0.002 m/m. The May 18, 1980, South Fork Toutle River lahar and the March 19, 1982, North Fork Toutle River lahar, both from Mount St. Helens, contained less than 3 percent clay and were noncohesive. They changed from debris flow to hyperconcentrated streamflow along 63 and 58 km channel reaches, respectively, that ranged from 0.01 to 0.002 m/m in slope. However, these transformations occurred as the lahars were diluted when they overrode streamflow. Eventually, the flows were diluted to such an extent that they became sediment-laden streamflow.

The size of the March 19, 1982, North Fork Toutle River lahar on the flanks of Mount St. Helens was similar to the June and September 1992 lahars from Crater Peak. However, as the 1982 lahar spread onto the fan at the base of the north flank of Mount St. Helens, it traversed highly erodible 1980 pyroclastic-flow and debris-avalanche deposits, and it not only bulked in volume but also incorporated sediment with a higher proportion of clay-size material (Pierson and Scott, 1985, fig. 9).

The 1992 lahars at Crater Peak contained less than 1 percent clay (fig. 13) and transformed from debris flow to sediment-laden streamflow as they moved downvalley. The flow transformations that were observed within the May 18, 1980, South Fork Toutle River lahar and the March 19, 1982, North Fork Toutle River lahar were quite different than those of the Crater Peak lahars. First, the Crater Peak lahars did not leave extensive deposits of massive, moderately sorted, sand deposits, which are indicative of hyperconcentrated streamflow (Scott, 1988, table 4). Secondly, the Crater Peak lahars moved through dry channels and overbank areas and the transformations could not have resulted from the lahar overriding streamflow. Thirdly, the lahars did leave matrix-poor levees at the distal margins of the flow. Fourth, transformation of the Crater Peak flows resulted in the debris flow coming to rest on slopes as steep as 0.04 to 0.09 m/m. The extremely low proportions of clay in the 1992 Crater Peak flows may have caused these differences.

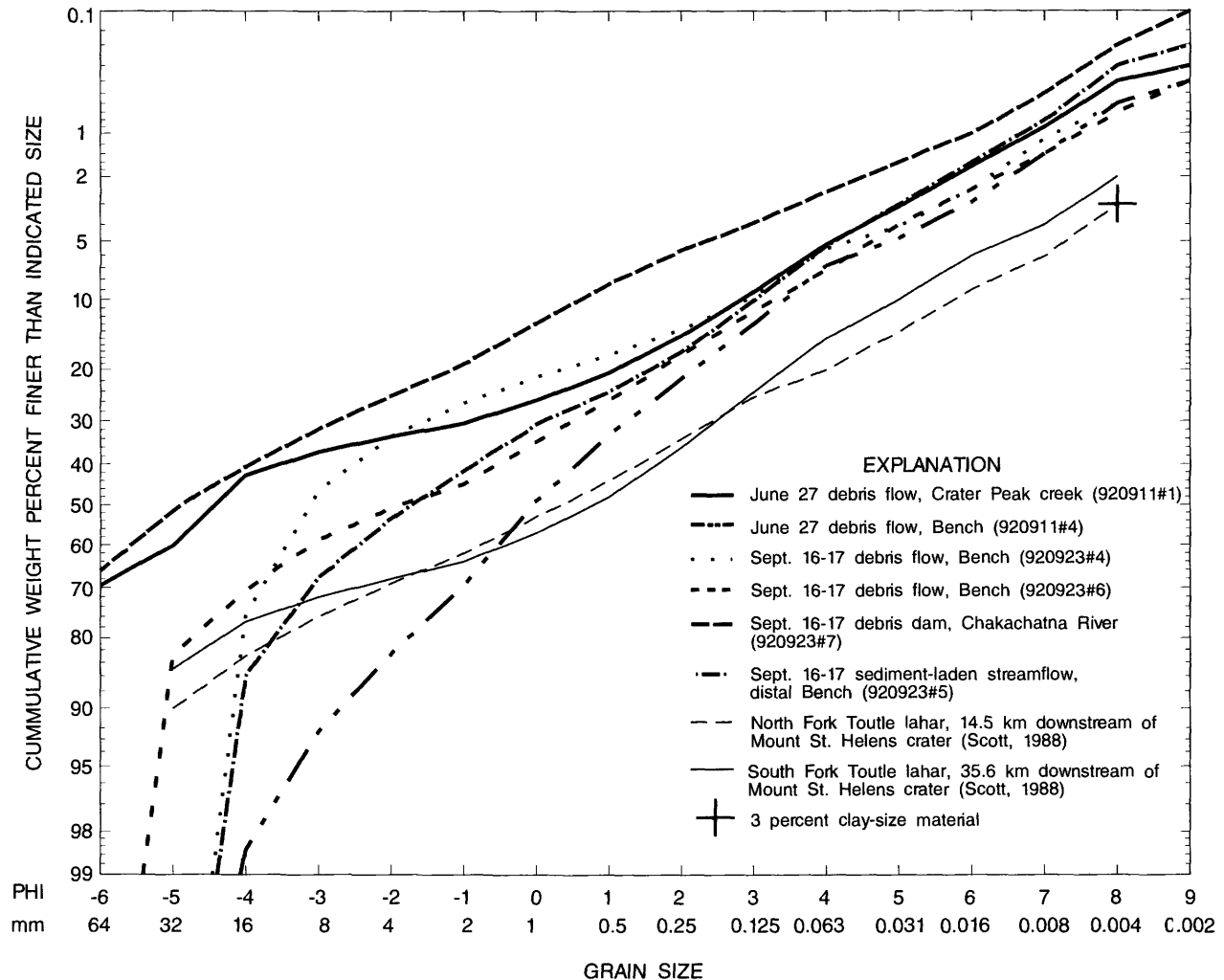


Figure 13. Grain-size distribution of matrix-only samples of June 27, 1992, debris flows and of September 16–17, 1992, debris flows, debris dam, and sediment-laden streamflow deposits from eruptions of Crater Peak vent, Mount Spurr volcano, Alaska. None of the samples include material coarser than 64 mm in diameter. Material coarser than 64 mm in diameter composed as much as 50 percent of the Mount Spurr debris-flow deposits but was not present in sediment-laden streamflow deposits. Also shown are grain-size distributions of two lahars from May 18, 1980, eruption of Mount St. Helens (Scott, 1988). Large plus sign is at a point representing 3 percent clay-size material, the percentage of clay-size material thought to distinguish cohesive and noncohesive lahars.

The distal lobes of the 1992 Crater Peak lahar deposits are composed of concentrations of coarse boulders devoid of fine-grained matrix. These lobes are unlike the contemporaneous Crater Peak lahar deposits in the proximal Bench or the 1980 South Fork Toutle River lahar deposits, in which coarse boulders are supported by a finer grained, poorly sorted, but clay-poor matrix. The distal lobes appear to have lost the fine-grained sandy matrix during or after deposition of the large boulders. Slurry flows, that is flows that show plastic flow behavior in the field yet become partially liquefied when remolded, commonly consolidate at a rate at which pore fluid can drain out, but most sediment remains in the deposit (Pierson and Costa, 1987,

p. 7). This phenomenon is due, in part, to viscous forces, which are dominant over inertial forces in finer grained or slower moving slurry flows. Such was true of the Toutle River lahars, which all contained more than 1 percent clay. The Crater Peak lahars, being generally coarser grained, would have been more dependent (than the Toutle lahars) on inertial forces to sustain plastic fluid flow that characterizes debris flow (Pierson and Costa, 1987, p. 7). When the Crater Peak deposits came to rest, viscous forces were insufficient to hold the matrix in place between the boulders, and the matrix was probably eroded by pore water as it drained rapidly from the deposits.

Distal lobes were not formed by most of the Mount St. Helens lahars or by the June 27 Crater Peak creek lahar; these lahars were diluted when they overrode streamflow as they flowed downstream and where they entered larger rivers.

Several alternative hypotheses could also explain the boulder levees and lobes that are devoid of matrix:

1. The fine-grained matrix could have been washed from the deposits by waning, watery stages of the lahar. This process probably did account for some of the general reworking of the diamict and the levees. Channels were eroded into the surface of the diamict upstream from the levees and on the glacier. However, these channels were commonly devoid of boulders larger than about 1 m, and they were observed only as discrete channels on the diamict, whereas the matrix-poor boulder levees were common near the distal margins of the diamict.

2. The fine-grained matrix could have been washed out by later rain. The June 27 deposits were not sampled and photographed until September 11, so they would have been reworked to an unknown degree by intervening rainfall and runoff, even though the deposits appeared unaltered where they were sampled. The September 16–17 deposits were sampled and photographed on September 23. Minor amounts of snow and little or no rain fell during the intervening period, and no reworking by surface runoff was observed on fresh, vertical cut banks within the diamict.

3. The lahar could have transported large concentrations of boulders that were deposited together as the flow traversed the Bench. This process apparently did occur, but it was responsible for deposits unlike the matrix-poor levees. Large (larger than 2 m) boulders were observed in the diamict whose flat surfaces were covered by sand and gravel similar to the reworked matrix deposits. The bases of these boulders were not observed, but similar deposits present in a debris-flow deposit on an alluvial fan at Olancho Creek in southeastern California rest directly on rooted vegetation. As the flow at Olancho Creek spread out over the alluvial fan, increasing in width, it probably became progressively shallower. As its depth became roughly equal to the diameter of the largest boulders, those boulders became grounded, while finer material, although still coarse, continued to flow downfan. At Olancho Creek, very large boulders (larger than 2 m) were grounded near the head of the fan, and grounded boulders are progressively smaller downfan.

At Crater Peak, small flows generated by water draining from the diamict reworked the surface of the deposit (fig. 10). Larger flows, generated from water draining from the distal margins of the diamict, co-

alesced to form flows large enough to erode unconsolidated deposits farther downstream, for example, along Bench gully. Similar flow transition was observed in debris flows having extremely low (less than 1 percent) proportions of clay-size material in O'lancho Creek, California (D.F. Meyer, unpub. data) and in Ophir Creek, western Nevada, by P.A. Glancy (oral commun., 1993) of the U.S. Geological Survey.

At Mount St. Helens and at Mount Rainier, cohesive lahars commonly originate as debris avalanches (Janda and others, 1981; Scott, 1988; Scott and others, 1992). Noncohesive lahars have a variety of origins, and some of them are generated by erosion of enough volcanoclastic material that flow concentrations exceed 60 to 80 percent sediment by weight. The higher concentrations are characteristic of coarser material. Some of the initiating flows are water flows generated by lake breakouts (Scott, 1988), pyroclastic flows that have eroded glacier ice (Scott and others, 1992) as did the September 16–17 Crater Peak lahar, and wet snow avalanches similar to the initiation of the June 27 lahar (Waitt and others, 1983; Pierson and Scott, 1985; Waitt, 1989). However, flow dilution (the incorporation of water in the stream channel that is overridden by debris flow) may play an important role in the transformations from debris flow to hyperconcentrated streamflow and from hyperconcentrated streamflow to sediment-laden streamflow (Pierson and Scott, 1985). The Bench at Crater Peak, where both the June 27 and September 16–17 lahars changed from debris flow to hyperconcentrated flow and from hyperconcentrated flow to sediment-laden streamflow, was not a stream channel. These transformations took place on a broad, vegetated plain (fig. 11), and it was not likely that any water was overridden or incorporated into the flow. Similarly, the Nevada and California debris-flow transformations occurred on alluvial fans, and these flows did not override significant streamflow.

Comparison of the August 18 and September 16–17 Crater Peak lahars confirms an observation of the 1989–90 lahars from Redoubt Volcano that ice is most effectively incorporated into pyroclastic flow when its surface is steep and rough. Most of the ice removed from Drift Glacier on Redoubt Volcano was from the steep, narrow canyon on the north breach of the volcano. Where pyroclastic material was deposited on the relatively flatter piedmont lobe of Drift Glacier, little melting occurred (Trabant and Meyer, 1992; Trabant and others, 1994). The August 18 pyroclastic flows were deposited primarily on smooth areas of Kidazgeni Glacier and no lahars were generated. The September 16–17 pyroclastic flows traversed steep, highly crevassed areas of the glacier and generated lahars.

OTHER HYDROLOGIC EFFECTS OF THE 1992 CRATER PEAK ERUPTIONS

Increased meltwater runoff can be expected from all glacier areas where ash deposits are less than about 24 mm in thickness because of decreased albedo, whereas ash thickness of more than a few centimeters can be expected to reduce meltwater runoff because the low heat conductivity of the materials insulates the ice from diurnal temperature fluctuations (Driedger, 1981). The only known ash deposit with thicknesses more than a few centimeters is on the Kidazgeni Glacier and Straight glacier (*informal name*, fig. 1). Prior to the 1992 eruption, thick, ablation-reducing ejecta from the 1953 eruption of Crater Peak mantled Kidazgeni Glacier below the junction of the east and west arms as well as the lower few kilometers of Straight Glacier.

The effects of the 1992 eruption on Barrier Glacier (fig. 1) and its debris-covered piedmont lobe, which dams the Chakachatna River to form Chakachamna Lake, are not expected to be significant. However, these effects need to be considered because release of Chakachamna Lake would be a significant hydrologic hazard. The flow of Barrier Glacier is not likely to have been affected by a change in the heat flow to the bed of the glacier, because the glacier is more than 5 km from Crater Peak, the active vent. Furthermore, almost no ash was deposited on the glacier, and no pyroclastic flows occurred on or near the glacier. The ice-flow speed of the piedmont lobe fluctuated in the past (G.C. Giles, U.S. Geological Survey, written commun., 1967); however, the single flood release from Chakachamna Lake probably was caused by lateral erosion of the ice dam during rainfall-generated high flow (Lamke, 1972).

Fallout tephra deposits from the three Crater Peak eruptions in 1992 blanketed much of the heaviest populated areas in south-central Alaska (see Neal and others, this volume). After the August 18 eruption deposited several millimeters of ash on Anchorage, both air and water quality were affected. Water use was greater than at any previous time, as businesses, homeowners, and the municipality washed away the ash. (Charley L. Bryant, Anchorage Water and Wastewater Utility, Municipality of Anchorage, written commun., 1992)

CONCLUSIONS

The interaction between snow and glacier ice and the pyroclastic flows and lahars produced by the 1992 eruptions of Mount Spurr volcano confirm the conclusions drawn from similar occurrences during the

1989–90 eruptions of Redoubt Volcano (Trabant and others, 1994) and the 1985 eruption of Nevado del Ruiz Volcano (Pierson and others, 1990). First, flows generated by mechanical disruption and entrainment of snow and glacier ice by pyroclastic flows are generally larger and more likely to be debris flows than when hot pyroclastic material is deposited more passively over flat surfaces. Second, bulking of lahars by erosion of debris downstream of the initiation point can produce greater hazards than those created by the initial flows. The lahars contained extremely low proportions of clay-size material and came to rest on relatively steep slopes, a behavior more similar to debris flows observed in Nevada and California than to flows observed at other volcanoes.

During the August 18 eruption, when pyroclastic material was, for the most part, deposited passively on glacier ice, small meltwater flows were produced. During the September 16–17 eruption, more energetic pyroclastic flows traversed a larger and steeper area of the glacier that was broken by crevasses. These flows eroded and melted enough glacier ice to produce large floods downstream. This observation indicates that the largest potential for incorporation of snow and glacier ice by volcanogenic flows is at places where ice surfaces are highly crevassed, steep, and exposed to large volumes of volcanogenic materials.

During the June 27 eruption, the eruption triggered snow, ice, and rock avalanches, but significant amounts of snow and ice were not melted by interaction with hot volcanoclastic material. In many ways, the June 27 lahar was similar to the March 19, 1982, lahar at Mount St. Helens, which began as a snow avalanche, and bulked by eroding volcanoclastic material both on the flank of the volcano and on the 1980 debris-avalanche deposit. The June 27 Crater Peak lahar bulked on the flanks of the volcano, but it was deposited at the base of the volcano where it entered the Chakachatna River valley.

The character of the lahars generated during the 1992 eruptions changed rapidly as they flowed down the flanks of the volcano and into the Chakachatna River valley. Where abundant, readily erodible sediment existed along the channel margins, such as along Crater Peak creek and Bench gully, the lahars incorporated additional sediment and sometimes bulked to debris flows. However, the flows contained very little clay-sized sediment, and most of the gravel in the flows was deposited on fans having slopes between 0.04 and 0.09 m/m, which is much steeper than depositional areas of lahars having more clay-sized sediment. Because the flows deposited most of their sediment on relatively steep slopes, the hazards resulting directly from the lahars did not extend far beyond the base of the volcano.

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June 29, 1993, Outburst Flood from Kidazgeni Glacier, Mount Spurr Volcano, Alaska

By Christopher J. Nye, William R. Hammond, Guy C. Tytgat, and Joseph M. Dorava

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ABSTRACT

A long, high-amplitude seismic signal caused by an outburst flood from the Kidazgeni Glacier occurred during the late morning of June 29, 1993. This was the longest continuous seismic signal recorded on the Spurr seismic network since the tremor episodes of October 2, 1992, and the earthquake swarm of November 9–10, 1992. The event was significant because visual analysis of helicorder records generated considerable concern that it might have been tremor premonitory to an eruption and because poor weather made direct observation of the Crater Peak vent of Mount Spurr impossible. The seismic signal was different from tremor both in the amplitude ratios between stations and in frequency content. This was not fully, or equally, appreciated by all those involved in the emergency response, and there was protracted discussion about whether or not the Alaska Volcano Observatory (AVO) should raise the level-of-concern color code. Fortunately, a field crew was in the area during the increased seismicity, and they observed an outburst flood from the terminus of the Kidazgeni Glacier. This flood was the source of the seismicity.

INTRODUCTION

The Kidazgeni Glacier is a valley glacier that drains the southern part of the Mount Spurr amphitheater. The glacier is 1.5 to 3 km wide and about 7.5 km long (fig. 1). It averages 125 m in thickness

and covers an area of about 21 square kilometers (R.S. March, oral commun.). It issues from the amphitheater on the east side of Crater Peak, the active vent of Mount Spurr volcano and the source of the 1992 eruptions (Eichelberger and others, this volume). It flows steeply from an altitude of about 2,135 m (7,000 ft) to about 900 m (3,000 ft) then spreads into a broad, flat lobe that ends at about 700 m (2,300 ft) elevation (fig. 1).

The Alaska Volcano Observatory currently maintains a local network of about 10 seismometers in the Mount Spurr area. Between 1981 and 1988, only 3 stations were maintained, and during that time the station closest to the terminus of the Kidazgeni Glacier was station CRP, 4 km to the north. From 1988 through 1992 the network density was increased. Station CKN, 1 km south of the Kidazgeni Glacier was installed in late 1991 (Power and others, this volume).

This report summarizes the local seismicity and event chronology of a small outburst flood that issued from the terminus of the Kidazgeni Glacier. The event was well recorded on local seismometers. Similar floods or seismic signals had not previously been noted at Mount Spurr, but they have occurred at other snow- and ice-covered volcanoes (see Brantley and Power, 1985; Driedger and Fountain, 1989).

ACKNOWLEDGMENTS

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SEISMICITY

Seismicity associated with the flood began at 9:20 a.m. ADT (Alaska daylight time, Universal Time minus 8 hours) on June 29, 1993, with an abrupt increase in amplitude of about two times the background level at station CKN (figs. 1, 2, 3). Within about 3 min-

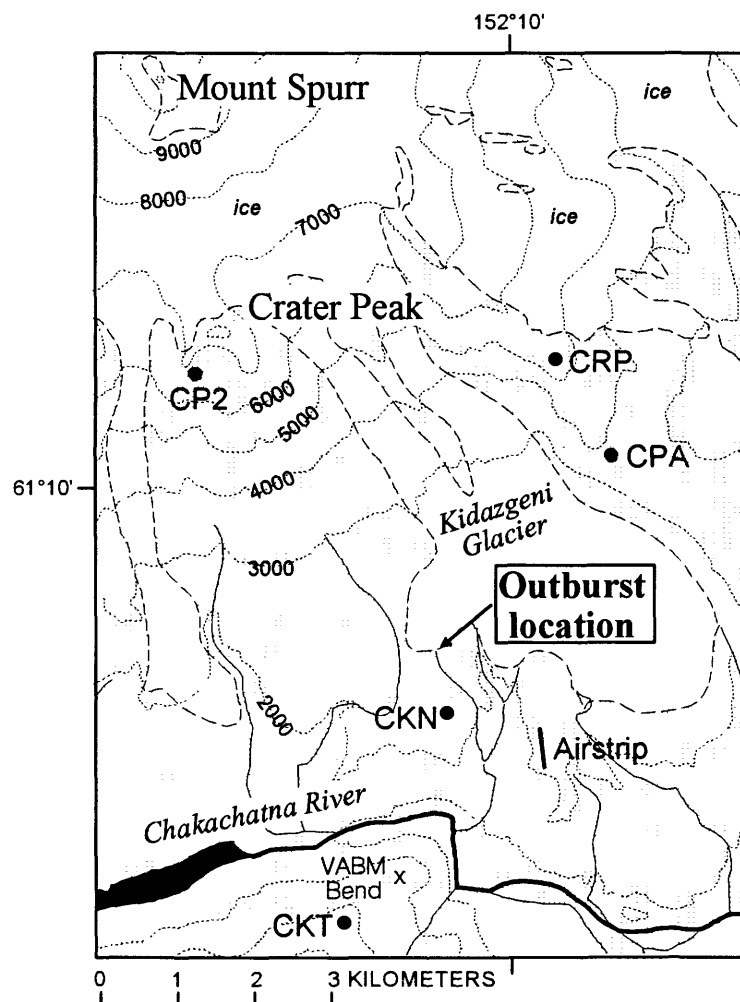


Figure 1. Geography of the Kidazgeni Glacier area of Mount Spurr volcano, Alaska, showing outburst flood location, location of seismic stations (CP2, CRP, CPA, CKN, and CKT), and geographic features referred to in the text. Contour interval 1,000 feet.

utes the seismic signal had begun clipping (some part of the seismometer and telemetry system was saturated). Seven minutes after the initial onset, the amplitude abruptly decayed to nearly background level, where it remained for about 8 minutes. At 9:36 a.m. ADT, a second stronger and longer seismic pulse began (figs. 2, 3). By about 10:11 a.m. ADT, seismicity at CKN was back to two times background level. It remained at this level for about another 9 minutes, when there was a third seismic pulse similar in duration to the initial pulse. Unlike the initial pulse, which stopped abruptly, the third pulse decayed over a period of about 10 minutes. The background seismicity level remained slightly above the pre-event level for the next day or so. The observed frequency of the seismic signal peaked at 4 to 6 Hz, and it had a broad high-frequency tail extending out to about 16 Hz (fig. 4). This frequency content was distinct from the June

26, 1992, pre-eruptive tremor as recorded at station CKN, which had lower frequencies (2–4 Hz) and did not have much energy above 5 Hz. Following the main seismic event on June 29, an increased number of discrete, short duration events were recorded primarily at station CKN. These persisted for a few days (figs. 2, 5).

Other seismic stations recorded the seismicity associated with the flood, although the peaks in amplitude were shifted in time (fig. 3). These times reflect more than merely greater distance from the source. Instead, they record either a migrating source or the onset of the most energetic seismicity, which could only be sensed at more remote stations. An unequivocal choice between these two possibilities cannot be made because station CKN had clipped so the time of maximum seismic intensity there could not be determined. However, we favor the onset of more energetic seis-

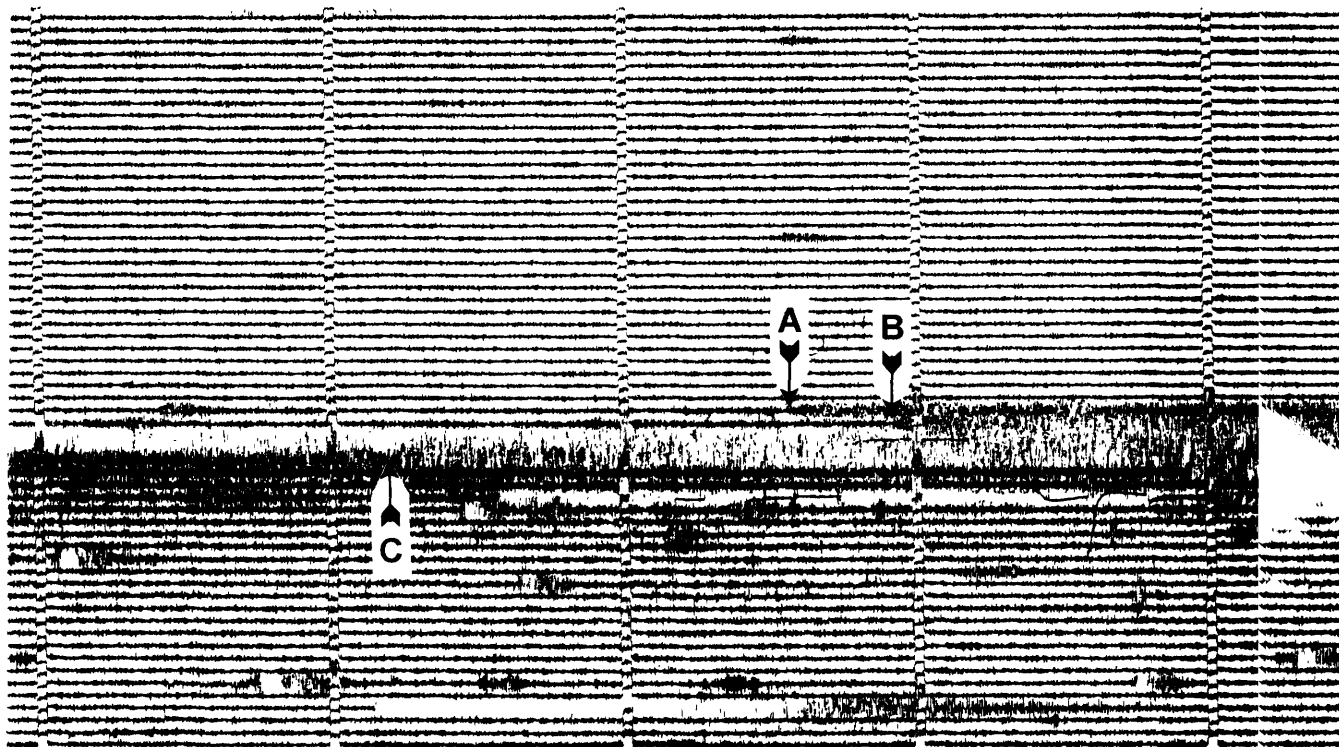


Figure 2. Part of the June 29, 1993, helicorder record for seismic station CKN, (Mount Spurr volcano, Alaska) showing the initial (A), second (B), and third (C) onsets of high-amplitude seismicity and post-outburst discrete events. The background noise preceding the outburst had been typical of station CKN since early June. Tick marks are 1 minute apart.

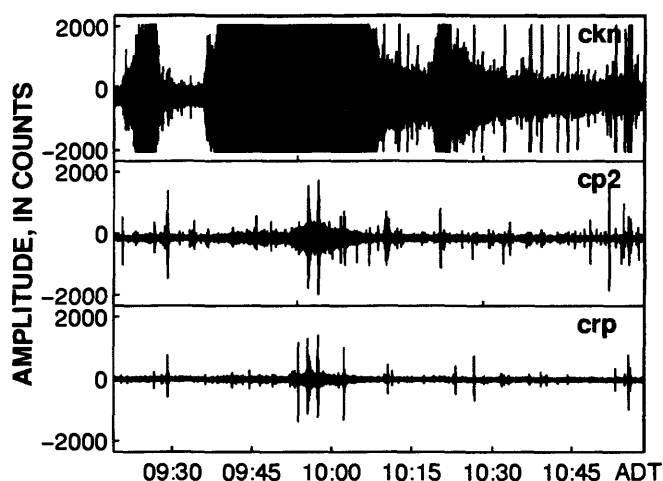


Figure 3. Digital seismic records of the outburst flood at seismic stations CKN (top), CP2 (middle), and CRP (lower) located on or near Mount Spurr volcano, Alaska. Note high intensity and early onset at station CKN.

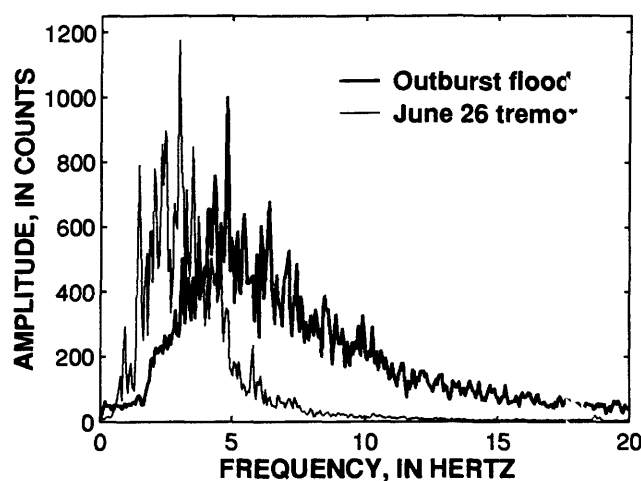


Figure 4. Frequency content of the June 29, 1993, outburst flood and June 26, 1992, pre-eruptive tremor as recorded at station CKN. Each spectrum is constructed from several 10-second segments, which have been stacked to reduce noise.

micity as a possibility. A plot of the envelope of seismicity (not shown) suggests that the maximum energy was released at the time the largest signal was recorded at stations CP2 and CRP.

Station CP2 showed an increase in seismic activity beginning at 9:38 a.m. and ending about 10:05 a.m. ADT. The frequency content was broadly distributed, extending from 2 to greater than 20 Hz with

no distinct spectral peaks (fig. 6). The signal-to-noise ratio was low, but the signal exceeded the noise for all but the very lowest and very highest frequencies. Station CRP showed an increase in seismic activity beginning at 9:52 a.m. ADT and ending about 10:05 a.m. ADT (fig. 3). The frequency content of the signal was narrow, peaking at 2 Hz, but had a broad, very low-amplitude tail extending to higher frequen-

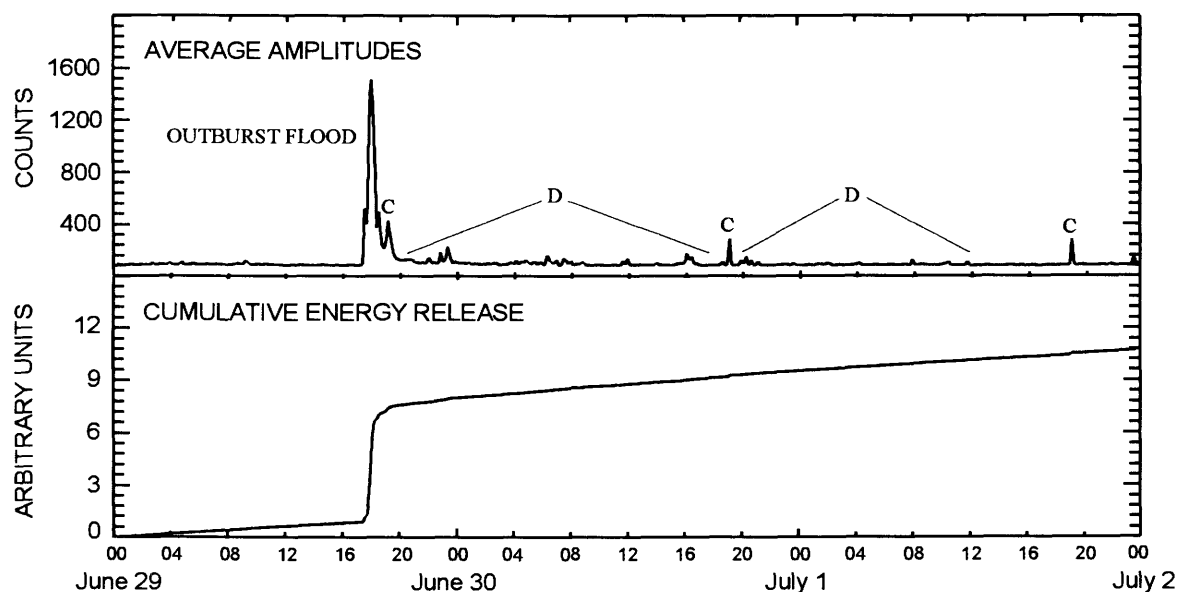


Figure 5. RSAM plots of the outburst flood and succeeding discrete events as recorded at seismic station CKN (located on Mount Spurr volcano, Alaska). In upper plot, discrete events are marked "D" and calibration pulses "C." Upper plot is the average amplitude in counts; lower plot is cumulative energy release in arbitrary units. The graph starts on June 29, 1993, and runs through July 1, 1993. Universal time (intermediate ticks on horizontal axis) expressed in 4-hour increments.

cies (fig. 7). The signal-to-noise ratio was low, however the signal exceeded noise below about 7 Hz. Note that the spectra from stations CP2 and CRP have amplitudes two orders of magnitude lower than those from station CKN. Some aspects of the frequency spectra, particularly high-frequency attenuation, may be artifacts of signal dampening associated with the low energy and greater distances between stations CP2 and CRP and the source and station CKN and the source. Additionally, the low signal-to-noise ratio for these two spectra may significantly reduce the meaningfulness of these data. The June 29 seismic activity at Mount Spurr volcano resembled tremor seen prior to past eruptions with the critical differences that the most energetic signal was not recorded at the station closest to the vent and the frequency content was higher at the station with the highest amplitude signal (CKN).

FIELD OBSERVATIONS

During the time of increased seismicity, a field party consisting of personnel from AVO, the Cascades Volcano Observatory (CVO), and Carnegie Mellon University was attempting to reach Crater Peak for a reconnaissance in preparation for a robotic mission scheduled for 1994. They approached Crater Peak from the east at about 10:05 a.m. ADT. The cloud base was at about 1,200 m. The field party was informed of the sustained seismicity by AVO Anchorage and landed

at the airstrip 1.5 km east of station CKN (fig. 1) to obtain more information because internal radio communication in the helicopter was not working properly. At the airstrip the field crew was informed of the nature of the seismicity and the concern that it might be tremor premonitory to an eruption. The field crew moved to VABM Bend (fig. 1) by 10:30 a.m.

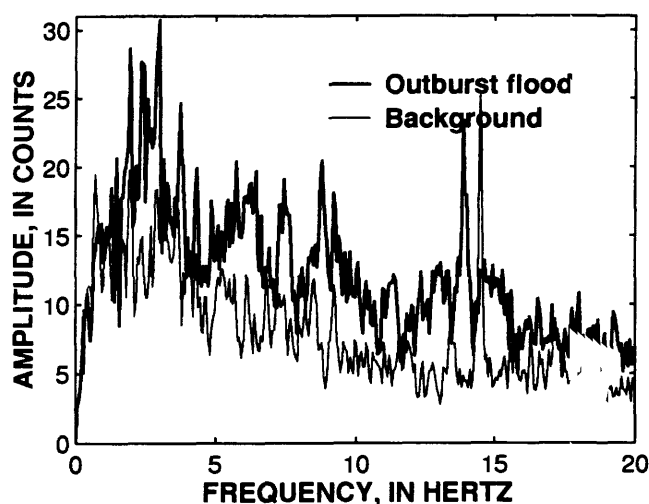


Figure 6. Frequency content of the June 29, 1993, outburst flood from Kidazgeni Glacier, Alaska, as recorded at seismic station CP2 on Mount Spurr volcano. Note extremely low amplitude compared to signal recorded at station CKN (see fig. 4).

ADT. When leaving the airstrip those members of the field crew familiar with the Mount Spurr area noted that the creek flowing between the airstrip and station CKN (informally called Kidazgeni Creek here) was in flood with dark brown water that was heavily laden with sediment. The water level was higher than had ever been observed previously by those in the field party. Where Kidazgeni Creek entered the Chakachatna River the flow was large enough that mixing was immediate across the entire width (80–100 m) of the Chakachatna River. Below the mouth of Kidazgeni Creek, the Chakachatna River was the same shade of dark brown as Kidazgeni Creek. Above Kidazgeni Creek the Chakachatna River was its typical milky blue-green. We estimate that the flow out of Kidazgeni Creek must have been a few tens to several tens of percent of the total flow of the Chakachatna River to so completely and immediately dominate the river's color.

The field crew remained at VABM Bend observing Kidazgeni Creek, waiting for a change in the seismicity, and waiting for the clouds to lift to allow access to Crater Peak. By 12:25 a.m. ADT the flow in Kidazgeni Creek had subsided enough to expose freshly cut and collapsing stream banks with cutbanks on the order of 1 m high. The field crew conducted an aerial reconnaissance of Kidazgeni Creek and lower Kidazgeni Glacier. Flow into the Chakachatna River was reduced to the point that the muddy brown flood water no longer dominated the Chakachatna River at the confluence with Kidazgeni Creek. Instead, flow from Kidazgeni Creek influenced about one quarter of the width

of the Chakachatna River, and it followed the northeastern bank downstream for a few hundred meters before complete mixing. After mixing with the reduced discharge from Kidazgeni Creek, the Chakachatna River was intermediate in color between the dark chocolate brown of Kidazgeni Creek and the milky blue-green of the upper Chakachatna River. The field crew also noted a new stream exit in the western terminus of the Kidazgeni Glacier (fig. 1) surrounded by newly broken ice blocks. It was from this new exit that the dark brown flood waters were issuing. A small supraglacial stream a few tens of meters to the west was flowing clear, and another subglacial stream several tens of meters to the east was the typical milky blue-green of glacial streams. The field crew proceeded up glacier to about the 1,500 m level looking for major changes to the surface of the glacier but saw none. There is no indication that the outburst was generated by rapid draining of a supraglacial lake. Instead the water storage must have been englacial or subglacial. The glacier in this area is heavily crevassed with a thick cover of debris and 1992 tephra. The field crew returned to the airstrip to observe Kidazgeni Creek and wait for better weather. A small rock slide was correlated with one of the discrete seismic events recorded by station CKN. By 3:50 a.m. ADT the ceiling had lifted to within a hundred meters of the low spot on Crater Peak rim and the field crew flew up along the southern flank of Crater Peak hoping for at least a view into the crater. The ceiling was solid and below the lip of the crater so the field crew returned to Anchorage.

Periodic field investigations of the Kidazgeni Glacier surface were made throughout the 1993 field season to search for evidence that a single subglacial cavity containing the water had drained. However, no surface collapse features were found.

DISCUSSION

On June 29, 1993, subglacial water burst from a small opening in the terminus of Kidazgeni Glacier and rushed down the steep slopes of Crater Peak to the Chakachatna River. This outburst flood was well recorded seismically and confirmed by field observations. Seismicity at the station nearest the outburst location (CKN, 1 km south) was distinct from volcanic tremor in its high frequency content (peaking at 5 Hz compared to 3 Hz tremor) and high-frequency tail extending to greater than 20 Hz (fig. 4). The distinct high-frequency signal was greatly attenuated by the time outburst seismicity was recorded at stations more distant (but still within 5 km of the creek). Similar broadband high-frequency seismicity was also observed during floods and lahars at Mount St. Helens, Redoubt

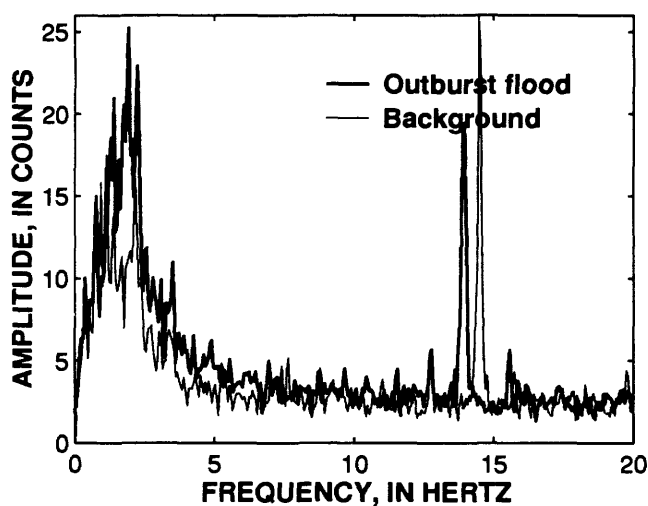


Figure 7. Frequency content of the June 29, 1993, outburst flood from Kidazgeni Glacier, Alaska, as recorded at seismic station CRP on Mount Spurr volcano. Note extremely low amplitude compared to signal at CKN (see fig. 4). Signal-to-noise ratio is small, but signal exceeds noise between about 1.5 and 7 Hz.

Volcano, and Mount Pinatubo (Brantley and Power, 1985; Dorava and Meyer, 1994; Bautista and others, 1991). The characteristic frequency of flood events has been recognized by developers of acoustic flow monitoring instruments, which have been used to remotely detect lahars (Hadley and LaHusen, 1991). The increase in discrete events after the June 29 outburst was also recorded most strongly at station CKN, and presumably it reflects settling of the glacial ice after withdrawal of englacial or subglacial water. It is not uncommon to have large quantities of water stored in or under glaciers, especially those on active volcanoes. Post and Mayo (1971) noted that "glaciers with no visible lakes may present unusual flood hazards" and that there is "a common association of glacier outburst floods with glacier-clad volcanoes."

The total discharge of the flood is not precisely known because no flow rate measurements were made during the event and no channel slope and cross-section measurements were made after the event. However, based on the complete and immediate mixing of Kidazgeni Creek and the Chakachatna River, field observers estimated that peak flow from Kidazgeni Creek was a few tens of percent of the total flow of the Chakachatna River. Average June, July, and August flow of the Chakachatna River was about 280 m³/s during the period 1959–1972 (U.S. Geological Survey, 1973). If the outburst flood of June 29, 1993, produced a sustained discharge of 20 to 30 percent of this flow during the 66 minutes of seismic shaking at station CKN, then the total flood volume would have been 220,000 to 330,000 m³. This may underestimate the total volume because field observers noted that Kidazgeni Creek flow remained at least somewhat elevated after seismicity at station CKN returned to near-background levels. The volume of this flood is about four orders of magnitude less than more commonly known outburst floods such as the Grimsvotn jökulhlaups in Iceland (Drewry, 1986). If this volume of water was released from a reservoir on the surface of the glacier it would require a pond about 100 m in diameter and 10 m deep. No such pond was seen on the surface of the Kidazgeni Glacier at any time during the years of periodic observations prior to the outburst, thus the outburst must have come from englacial or subglacial storage.

The magnitude of this outburst flood is difficult to compare with normal run-off from the glacier because of the scarcity of meteorological data, streamflow data, and comparable basin-characteristic information from similar glacier streams. The probable summer flow from the Kidazgeni Glacier was estimated by

dividing the average summer discharge of the Chakachatna River at the gauging station (280 m³/s; U.S. Geological Survey, 1973) by the 2,900 km² area of the basin providing that flow, resulting in a unit discharge of 0.0966 m³/s/km². We compare this to the discharge of two unglaciated basins 125 km to the northeast, which average 0.0366 m³/s/km² and conclude that about 60 percent of the flow of the Chakachatna River is contributed by the 30 percent of the basin covered by glaciers. This analysis suggests that the average discharge from glaciers in the Chakachamna Lake basin is 0.20 m³/s/km². The Kidazgeni Glacier is about 21 km², and thus should have an average runoff of about 4.3 m³/s, making the peak outburst flood discharge of 70 m³/s about 16 times the mean summer runoff. The estimated flood volume corresponds to about one day of average runoff for the entire glacier at 4.3 m³/s. We caution that there may be substantial errors in estimating both the flood volume and typical runoff from the Kidazgeni Glacier. We believe, however, that the figures presented here illustrate at least the order of magnitude of volumes involved.

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Effects of the 1992 Crater Peak Eruptions on Airports and Aviation Operations in the United States and Canada

By Thomas J. Casadevall and M. Dennis Krohn

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ABSTRACT

Crater Peak vent on the south side of Mount Spurr volcano, 67 nautical miles (125 km) west of Anchorage, Alaska, erupted on June 27, August 18, and September 16–17, 1992. Each eruption produced an ash column that rose to between 43,000 and 49,000 ft (13 and 15 km) above sea level. Although drifting ash clouds from the three eruptions disrupted air traffic over Alaska, western Canada, and the central conterminous United States, no aircraft were damaged by encounters with the ash clouds.

Ashfall from the August 18 eruption deposited from 1/16" to 1/8" (1 to 3 mm) of ash in Anchorage and caused Anchorage airports to close or curtail operations for several days. The cost of removing the ash from the airports, including the cost of protecting and cleaning aircraft caught on the ground in Anchorage is estimated at \$683,000. The net passenger revenues lost owing to flight cancellations while Anchorage airports were closed are estimated at \$276,000.

Because encounters with volcanic ash clouds from the 1989–90 eruptions of Redoubt Volcano caused serious damage to five passenger airplanes, there has been a high level of concern in the aviation community about volcanic ash and the threat it presents to aviation safety. The lack of encounters between airplanes and ash clouds from the Mount Spurr eruptions reflects better understanding and increased awareness about the hazards of volcanic ash. Since 1990, a number of improvements have been made to reduce hazards resulting from volcanic ash. These improvements have resulted from major initiatives by Federal and international regulatory agencies and by aviation associations and include: (1) improved warnings of eruptions provided by increased monitoring of Cook Inlet volcanoes, (2) improvements in the detection and tracking of volcanic ash clouds, (3) improved education of pilots and flight dispatchers, and (4) streamlining the ways in which information and warnings are communicated among the agencies concerned with aviation safety.

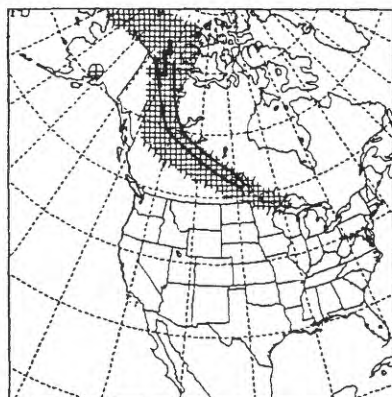
INTRODUCTION

Recent eruptions at Redoubt Volcano in Alaska and Pinatubo Volcano in the Philippines provided pilots, aircraft operations personnel, meteorologists, and volcanologists considerable experience in reacting to the presence of volcanic ash clouds and the resulting effects of ash on aviation safety (Casadevall, 1992; 1994a). The 1989–90 eruptions of Redoubt Volcano produced ash clouds which damaged five jet airliners (Casadevall, 1994b). The June 1991 eruptions of Pinatubo produced large ash clouds that were involved in at least 16 damaging encounters with jet airliners (Casadevall and Delos Reyes, 1991).

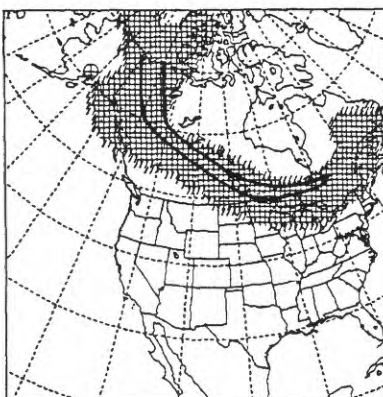
Mount Spurr (11,070 ft, 3,374 m) is one of three volcanoes in the Cook Inlet area of south-central Alaska that have erupted in the past decade. Eruptions of the Crater Peak vent (6,990 ft, 2,130 m) on the south side of Mount Spurr on June 27, August 18, and September 16–17, 1992, produced ash columns that rose to

43,000 to 49,000 ft (13 to 15 km) above sea level (Eichelberger and others, this volume). Upper-level winds carried ash from the three eruptions into airspace over south-central and eastern Alaska and western Canada. Ash from the eruptions in June and September was also carried over northern and central parts

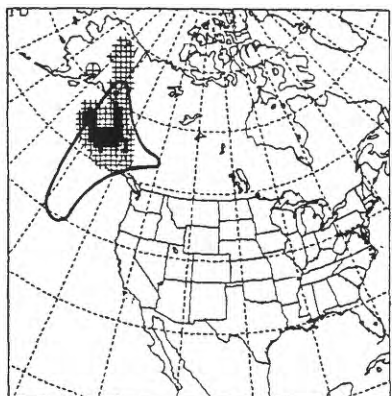
of the United States (fig. 1). This report evaluates the communication between agencies of information available at the time of the eruptions, describes the movement of the three Mount Spurr ash clouds, and examines the effects of the clouds on aviation operations.



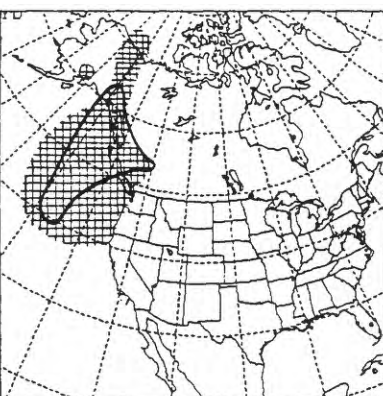
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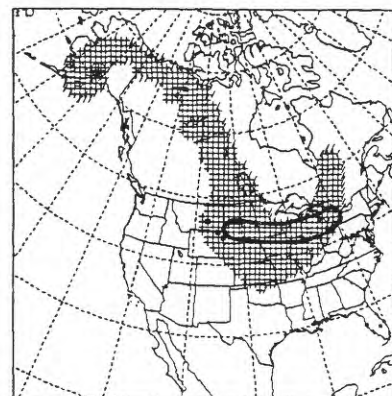
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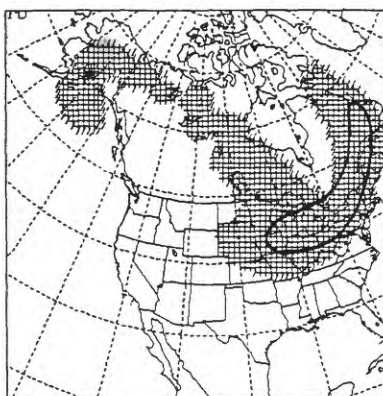
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FORECAST CYCLE 92/09/18-12Z

Figure 1. Comparative trajectories of the ash clouds from the June, August and September 1992 eruptions from Mount Spurr volcano, Alaska. Solid lines are ash cloud boundaries determined from composite satellite images. Patterned areas are cloud positions forecast by the volcanic ash forecast transport and dispersion (VAFTAD) model of Heffter and Stunder (1993). Modified from Heffter and Stunder (1993).

Prior to the 1992 eruptions, Mount Spurr last erupted on July 9, 1953, with a single explosive eruption that sent ash to a maximum estimated elevation of 70,000 ft (21.3 km) (Juhle and Coulter, 1955). Upper-level winds carried ash from that eruption over Anchorage, 67 nautical miles (125 km) east of the volcano, and from 1/8" to 1/4" (3 to 6 mm) of ash fell at Anchorage Airport and Elmendorf Air Base; this ashfall resulted in closing the airports for several days and disrupting civilian and military airport operations for several weeks (Juhle and Coulter, 1955; Wilcox, 1959). In addition to the airport closures, at least three F-94 military jets were damaged after flying into the eruption cloud. Ash also covered several C-124 transport aircraft (Globe Trotters) that were parked at Elmendorf; ash removal from these aircraft took 10 days (U.S. Air Force, 1955).

Whereas no damaging encounters between airplanes and ash clouds were reported for the 1992 Crater Peak eruptions, the ash clouds did disrupt air traffic over Alaska, Canada, and the northern and central United States. In addition, the removal of ash and the cleaning of airport facilities and airplanes required temporary suspension of operations at Anchorage airports for several days following the August 18 eruption.

COMMUNICATIONS

Prompt and direct communication of factual information about Mount Spurr's activity between the aviation community, meteorologists, and volcanologists was a key element in determining pilot response to the eruptions, just as it had been during the 1989-90 Redoubt eruptions (Brantley, 1990; Casadevall, 1994b). Beginning with the Redoubt eruptions, the main source of timely factual information about the activity of Alaskan volcanoes has been the volcano update issued by the Alaskan Volcano Observatory (AVO). The update is normally released weekly and may be supplemented with a level of concern color code that is based primarily on seismic activity and field observations. The color code uses one of four colors (Green, Yellow, Orange, or Red) to succinctly describe the status of the volcano when an eruption is expected or is in progress. Green refers to the normal or noneruptive state of the volcano. Yellow indicates that the volcano is restless with elevated seismic activity and the possibility of a plume of gas and steam rising several thousand feet above the volcano. Orange indicates that a small ash eruption is expected or confirmed with plumes not likely to rise above 25,000 ft (7.6 km) above sea level. Red indicates that a major explosive eruption is in progress or is expected within the next 24 hours with potentially hazardous ash plumes ex-

pected to exceed altitudes of 25,000 feet (7.6 km) or more. Changes in this information are issued at any time, 24 hours a day, as events at the volcano demand (Alaska Interagency Operating Plan, 1993).

During an eruption, updates are communicated immediately by telephone to the Anchorage Area Control Center, the Anchorage National Weather Service office, the Alaska Department of Emergency Services, and the U.S. Air Force at Elmendorf Air Base. Simultaneously, the update is sent by telephone facsimile to a wide user community including Anchorage International Airport and the airlines.

Concern that a future eruption of Mount Spurr would affect Anchorage in a manner similar to the 1953 eruption prompted scientists to initiate seismic monitoring of the volcano beginning in 1981. Since 1989, the AVO has maintained a network of as many as 10 seismometers on the volcano (Power and others, this volume). Beginning in August of 1991, intermittent swarms of small magnitude, shallow earthquakes were detected beneath Crater Peak. After a slightly larger swarm on February 5, 1992, AVO described the increased seismicity in the update of February 7, 1992 (table 1). Between February 7 and June 26, AVO issued 19 additional updates that described the activity of the volcano, including a special update issued on June 8. This special update contained analyses of increased seismicity and other monitoring data then available, such as rates of gas emission (Doukas and Gerlach, this volume) and physical and chemical changes to the small lake within Crater Peak vent (Keith and Thompson, this volume). On June 26, AVO used the color code Yellow to describe the status of the volcano. The 1992 eruptions of Mount Spurr began at 7:04 a.m. ADT (1504 UT) on June 27.

ACKNOWLEDGMENTS

We would like to express our appreciation for information about the eruptions provided by James Lynch and Grace Swanson of the NOAA-National Environmental Satellite, Data and Information Service, Synoptic Analysis Branch. For the meteorological analysis of the September eruption, we are grateful to David Helms (NASA-Johnson Space Center). For information about the effects on the Cleveland Air Traffic Center, we thank Thomas Howell (FAA). We also greatly appreciate discussions and the unpublished reports provided by Larry Michou (Anchorage International Airport), Captain Terry Spurgeon (Transport Canada), and Carl Shapiro (USGS). Reviews of this paper by John Pallister and Tom Miller of the USGS helped to improve the clarity of the presentation and are appreciated.

Table 1. Summary of activity from volcano updates issued by the Alaska Volcano Observatory concerning the June 27, 1992, eruption of Mount Spurr volcano.

Date	Time (ADT)	Update
2/07/92	--	First mention of increase in seismic activity at Mount Spurr. Minor increase in seismic activity had been detected previously, starting in August 1991.
3/13/92-6/05/92	--	Continuing mention of increased seismic activity made in weekly updates.
6/08/92	--	Extended update on Mount Spurr seismicity released.
6/09/92-6/26/92	--	Weekly updates mentioned increased seismicity.
6/26/92	4:30 p.m.	At 12:04 p.m. continuous tremor began. Second update issued at 4:30 p.m. on 6/26/92 describing increased seismicity. Color code Yellow declared.
6/27/92	8:00 a.m.	Eruption tremor began at 7:04 a.m. Visual verification of ash column to approximately 17,000 ft (5.2 km). Color code raised from Yellow to Orange.
6/27/92	9:15 a.m.	Ash column rose to greater than 30,000 ft (9 km). Color code raised from Orange to Red.
6/28/92	9:00 a.m.	Color code reduced from Red to Yellow.
6/28/92-7/07/92	--	Updates issued 6/28, 6/29, 6/30, 7/1, 7/2, 7/3, 7/6, 7/7. Color code remained at Yellow.
7/08/92	3:00 p.m.	Color code reduced to Green.
7/10/92	3:30 p.m.	Weekly updates resumed; color code remained at Green.

1992 CRATER PEAK ERUPTIONS

The three Crater Peak eruptions in 1992 each lasted from 3.5 to 4 hours, produced pyroclastic material of andesitic composition (Neal and others, this volume), and generated eruption columns that rose to between 43,000 and 49,000 ft (13 to 15 km) above sea level (Rose and others, this volume) (table 2). Each eruption also produced large clouds of volcanic ash and gas that affected air traffic and operations over a wide area of the United States and Canada (fig. 1).

JUNE 27, 1992, ERUPTION

The June 27, 1992, eruption of Mount Spurr's Crater Peak vent produced approximately 12 million cubic meters of andesitic tephra (Neal and others, this volume) and the maximum altitude of the eruption column, as detected by ground-based radar, was approximately 48,000 ft (14.5 km) (Rose and others, this volume). The forecast movement of the June 27 ash cloud during the first 24 hours was first communicated in the AVO updates issued on June 27 (table 2) and was based on analysis of the forecast upper-level winds. The projected path of the cloud indicated it would travel north of the volcano toward Denali National Park (Mount McKinley) and the Beaufort Sea and remain well clear of the Anchorage area. By June 28, upper-level winds changed direction toward the southeast and by June 29, the cloud had moved southeastward and was transported over western Canada toward the conterminous United States (Heffter and Stunder, 1993) (fig. 1).

On June 27, shortly after notification of the start of the eruption by AVO, the National Oceanic and Atmospheric Administration (NOAA) and the Federal Aviation Administration (FAA) activated their volca-

nic hazards alert plan, which remained in effect for 5 days. The Synoptic Analysis Branch (SAB) of NOAA's National Environmental Satellite, Data, and Information Service (NESDIS) used imagery from the geostationary operational environmental satellite (GOES) and forecast trajectories for the ash clouds (Heffter and Stunder, 1993) to track the June 27 cloud through July 2 when it became indistinguishable from weather clouds over southwestern Canada (J. Lynch, NOAA, written commun., 1992). During this period, SAB issued 19 bulletins about the position and movement of the ash cloud to the Anchorage National Weather Service forecast office, to the National Meteorological Center, to the National Aviation weather advisory unit in Kansas City, to the FAA, and to AVO.

Analysis of images from the advanced very high resolution radiometer (AVHRR) aboard the NOAA-11 polar-orbiting satellite also detected the cloud over western Canada (Aviation Week and Space Technology, 1992). The sulfur dioxide in the eruption cloud was detected using the total ozone mapping spectrometer (TOMS) aboard the Nimbus 7 satellite and the cloud was tracked through July 3. The eruption cloud contained approximately 200,000 tons of sulfur dioxide gas (Bluth and others, this volume), a mass comparable to the 175,000 tons emitted by the December 15, 1989, eruption of Redoubt Volcano (Schnetzler and others, 1994).

No airplanes flew into the June 27 ash cloud. The only modification of commercial air-traffic patterns in Alaska owing to the eruption was the rerouting of flights from Anchorage to Nome and Anchorage to Kotzebue on June 27 to avoid the ash-contaminated airspace north of Mount Spurr. Several ski-equipped charter flights used to service climbers on Mount Denali were canceled on June 27 and 28 because of ashfall on the Denali snowfields. This situation improved by June 29 after fresh snow covered the ash.

On June 29, the pilot of a daytime domestic flight in Canada reported that he had passed through a 1,000-ft- (300-m)-thick brownish haze layer at 28,000 ft (8.5 km) altitude as his aircraft descended into Winnipeg (T. Spurgeon, Transport Canada, written commun., 1993). Fortunately, the cloud had been strongly diluted by the time it entered Canadian airspace and there were no reports of ashfall or damage to aircraft in Canada.

AUGUST 18, 1992, ERUPTION

The first indication of eruptive activity on August 18 was a pilot report received at 3:48 p.m. ADT (2348 UT) that ash was rising from 1,000 to 2,000 ft (300–600 m) above Crater Peak. There were no clear seismic precursors to the August eruption, and the eruptive activity reported by the pilot was related to weak tremor that began at 3:41 p.m. ADT (2341 UT) (McNutt and others, this volume). After several additional pilot reports of this low-level ash column and analysis of the weak tremor, AVO raised the color code alert from Green to Yellow at 4:25 p.m. ADT (0025 UT). At 4:42 p.m. ADT (0042 UT), strong eruption-related tremor began to record on the seismic station closest to the vent, and at 4:45 p.m. ADT (0045 UT) AVO telephoned authorities at Anchorage International Airport and Elmendorf Air Base to advise them of the tremor. Authorities at Merrill Field were notified by AVO through the State of Alaska Department of Emergency Services. Following further increase in tremor and notification by the National Weather Service that the Kenai radar had detected an ash column rising above the volcano to more than 35,000 ft (11 km), AVO issued an update at 4:48 p.m. ADT (0048 UT) and declared color code Red.

The maximum column height of 45,000 ft (13.7 km) was detected by radar at 4:55 p.m. ADT (0055 UT) (Rose and others, this volume). The eruption lasted

for more than 3 hours and eruption tremor began to decline by 8:10 p.m. ADT (0410 UT). AVO lowered the color code from Red to Orange at 9:30 p.m. ADT (0530 UT) on August 18, and further reduced the level to Yellow at 8:30 a.m. ADT (1630 UT) on August 19. During the August eruption, the NOAA SAB issued 21 bulletins that periodically updated the position of the ash cloud and included the trajectory forecasts of cloud position and movement (J. Lynch, NOAA, written commun., 1993).

The August eruption produced 14 million cubic meters of tephra (Neal and others, this volume). The ash column, measured by radar, rose as high as 45,000 ft (13.7 km) in altitude (Rose and others, this volume), and upper-level winds carried the ash eastward toward Anchorage (table 2). Approximately 1/16" to 1/8" (1 to 3 mm) of fine sand-size ash fell in the Anchorage area between 8:20 p.m. ADT (0420 UT) and 11:00 p.m. ADT (0700 UT); as a result, the three Anchorage airports were closed for several days (figs. 2, 3) (Casadevall, 1993).

SEPTEMBER 16–17, 1992, ERUPTION

Following the August eruption, AVO maintained a color code of Yellow because levels of seismic activity remained elevated. The update on September 9 drew attention to the first increase in seismic activity since the August eruption, and the update on September 11 stated that an eruption was possible within the next few days or weeks. On September 16, the first phase of a new eruption at 10:36 p.m. ADT (0636 UT) was indicated by increased seismic activity and by pilot reports from Kenai, Alaska (table 3). AVO immediately initiated its calldown and in response to the telephone notification of the eruption from AVO and pilot reports, the National Weather Service issued a notice of significant meteorological event (SIGMET) for the eruption at 10:40 p.m. ADT (0640 UT) (fig-

Table 2. Comparison of the 1992 Mount Spurr volcano eruptions to 1953 Mount Spurr volcano and 1989 Redoubt Volcano eruptions.

[DRE, dense rock equivalent; TOMS, total ozone mapping spectrometer; SAB, Synoptic Analysis Branch, NOAA National Environmental Satellite Data and Information Service (J. Lynch, written commun., 1993)]

Date of eruption	Maximum altitude of eruption column (km)	Tephra volume (DRE X 10 ⁶ m ³)	Sulfur dioxide (megatons)	Traceable on TOMS	Number of SAB bulletins issued	Pilot reports
Mount Spurr volcano						
June 27, 1992	14.5	12	200±60	to 7/3	19	--
August 18, 1992	13.7	14	400±120	to 8/26	21	--
September 17–18, 1992	13.9	15	230±70	to 9/21	18	to 9/21
Mount Spurr volcano						
July 9, 1953	<21.5 (pilot estimate)	not available	not available	not available	none issued	--
Redoubt Volcano						
December 15, 1989	>12 (pilot estimate)	not available	175	to 12/19	22	--



Figure 2. Road graders scrape ash from runways at Anchorage International Airport on August 19, 1992. Road scrapers could effectively move the ash only after it had been sprayed and dampened with water. However, too much water produced a slurry that could not be moved effectively. Dampened ash was scraped into furrows and then scooped up using front-end loaders. Some ash resuspended in the air would settle back onto runways and planes, and thus delayed aircraft cleanup. Photograph by Erik Hill courtesy of Anchorage Daily News.



Figure 3. Maintenance crew removing ash from wing of a Reeve Aleutian Airways Boeing 727-100 aircraft on August 19, 1992, at Anchorage International Airport. Photograph by Erik Hill courtesy of Anchorage Daily News.

ure 4). AVO increased the color code to Red in the volcano update issued at 10:45 p.m. ADT (0645 UT). The most intense phase of the eruption began shortly before midnight on September 17 and lasted for about 4 hours. At 4:00 p.m. ADT (0000 UT) on September 17, the color code was reduced to Orange where it remained until September 24 when it was further reduced to Yellow.

The September ash cloud was noteworthy in several ways. Although, the volume of magma erupted, and the basic dimensions of the eruption column and the ash cloud were similar to the two previous eruptions (table 2), the cloud traveled as a coherent mass for 5 days after the eruption. It remained visible during that time to ground observers, pilots, and even to astronauts aboard Space Shuttle mission STS 47. The cloud moved east, passed north of Anchorage and deposited about 1/16" (1 mm) of ash at Willow, Palmer, and the along Parks Highway in Alaska. It continued across western and south-central Canada, then down into the central United States by September 19 before moving back into eastern Canada and eastward over the North Atlantic by September 21 (fig. 1) (Schneider and others, this volume).

Table 3. Pilot reports (PIREPs) and ground observations of ash-cloud locations for the September 16–17, 1992, eruption of Mount Spurr volcano, Alaska.

[Numbers refer to reporting sites shown in fig. 5. AK, Alaska; Yk. Tr., Yukon Territory; Sask, Saskatchewan; Albta., Alberta; SD, South Dakota; ND, North Dakota; NE, Nebraska; IL, Illinois; IA, Iowa; MN, Minnesota; PA, Pennsylvania; OH, Ohio; Ont., Ontario; WI, Wisconsin; NY, New York; MI, Michigan; Lab., Labrador; N. Brs., New Brunswick; Latitude and longitudes given in degrees and tenths of degrees. Time given is Universal Coordinated Time (UT). Altitude given in feet above mean sea level, or for values greater than 18,000 ft, relative to a barometric pressure of 29.92 inches Hg. Single entries indicate the base of the cloud; second entry indicates top of the cloud. Pilot reports compiled by NOAA Synoptic Analysis Branch (G. Swanson, NOAA, written commun., 1993) and by Transport Canada Aviation (T. Spurgeon, Transport Canada, written commun., 1993)]

Number	Location	State/ Province	Lat.	Long.	Date	Time	Altitude/ Observation
1	Mount Spurr	AK	61.3	152.2	09/17	0636	
2	Kenai	AK	60.62	151.19	09/17	0637	ground observation
2	Kenai	AK	60.62	151.19	09/17	0646	ground observation
3	Anchorage	AK	61.15	150.21	09/17	0713	
2	Kenai	AK	60.62	151.19	09/17	0745	35,000 4,000
2	Kenai	AK	60.62	151.19	09/17	0746	ground observation
2	Kenai	AK	60.62	151.19	09/17	0757	
3	Fire Island	AK	61.15	150.20	09/17	0800	
2	Kenai	AK	60.62	151.19	09/17	0804	ground observation
3	Anchorage	AK	61.15	150.21	09/17	0821	19,000 29,000
3	Anchorage	AK	61.15	150.21	09/17	0835	
3	Anchorage	AK	61.15	150.21	09/17	0838	
3	Anchorage	AK	61.15	150.21	09/17	0846	10,000
2	Kenai	AK	60.62	151.19	09/17	0846	ground observation
3	Anchorage	AK	61.15	150.21	09/17	0853	21,000 37,000
3	Anchorage	AK	61.15	150.21	09/17	0900	31,000 32,000
4	Snowshoe Lake	AK	62.03	146.68	09/17	1447	ground observation
5	Gulkana	AK	62.15	145.45	09/17	1449	ground observation
3	Anchorage	AK	61.15	150.21	09/17	1452	ground observation
2	Kenai	AK	60.62	151.19	09/17	1455	
6	Talkeetna	AK	62.30	150.10	09/17	1457	ground observation
1	Mount Spurr	AK	61.3	152.2	09/17	1501	35,000
7	Watson Lake	Yk. Tr.	60.07	128.48	09/17	1620	35,000
8	White River	Yk. Tr.	61.35	139.00	09/17	1650	
9	Beaver Creek	Yk. Tr.	62.45	140.62	09/17	1656	ground observation
10	Lake Laberge	Yk. Tr.	61.17	135.17	09/17	1850	6500
11	Whitehorse	Yk. Tr.	60.72	135.03	09/17	2012	
12	Dawson	Yk. Tr.	64.07	139.42	09/17	2141	5,000
13	Carmacks	Yk. Tr.	62.10	136.32	09/17	2200*	ground observation
13	Carmacks	Yk. Tr.	62.10	136.32	09/17	2300	9,000
14	Palmer	AK	61.59	149.08	09/17	2352	ground observation
3	Anchorage	AK	61.15	150.21	09/18	0757	
15	Fairbanks	AK	64.80	148.01	09/18	0757	
16	Lumsden	Sask.	50.67	104.89	09/18	1250	29,000 38,000
17	Calgary	Albta.	51.12	113.88	09/18	1430	22,000 25,000
18	Pierre	SD	44.40	100.16	09/18	1513	33,000 35,000
19	Swift Current	Sask.	50.30	107.69	09/18	1520	29,000 37,000
18	Pierre	SD	44.40	100.16	09/18	1528	33,000 42,000
20	Empress	Albta.	50.56	110.01	09/18	1650	33,500
3	Anchorage	AK	61.15	150.21	09/18	1705	sulfur odor noted 7,500
3	Anchorage	AK	61.15	150.21	09/18	1715	sulfur odor noted 7,500
21	Minot	ND	48.26	101.29	09/18	1845	20,000
22	Scottsbluff	NE	41.90	103.48	09/18	2110	37,000
20	Empress	Albta.	50.56	110.01	09/18	2246	24,000 33,000
23	Northbrook	IL	42.22	87.95	09/18	2301	30,000
24	Sioux Falls	SD	43.65	96.78	09/18	2324	27,000
21	Minot	ND	48.26	101.29	09/18	2328	33,000
25	Cedar Rapids	IA	41.89	91.79	09/18	2334	
25	Cedar Rapids	IA	41.89	91.79	09/18	2334	27,000
26	Minneapolis	MN	44.88	93.29	09/19	0052	ground observation

Table 3. Pilot reports (PIREPs) and ground observations of ash-cloud locations for the September 16–17, 1992, eruption of Mount Spurr volcano, Alaska—Continued.

Number	Location	State/ Province	Lat.	Long.	Date	Time	Altitude/ Observation
27	Ellwood City	PA	40.83	80.21	09/19	0116	35,000
28	Dayton	OH	39.90	84.22	09/19	0141	11,000
26	Minneapolis	MN	44.88	93.29	09/19	0154	ground observation
29	Windsor	Ont.	42.25	82.83	09/19	0200	27,000
30	Rochester	MN	43.78	92.60	09/19	0252	
26	Minneapolis	MN	44.88	93.29	09/19	0253	ground observation
31	Regina	Sask.	50.50	104.63	09/19	0320	17,300 20,500
26	Minneapolis	MN	44.88	93.29	09/19	0354	ground observation
1	Mount Spurr	AK	61.3	152.2	09/19	0444	steam observed 15,000
32	Milwaukee	WI	42.95	87.90	09/19	0653	ground observation
32	Milwaukee	WI	42.95	87.90	09/19	0950	ground observation
33	Dunkirk	NY	42.49	79.28	09/19	1220	28,000 33,000
34	Rosewood	OH	40.29	84.04	09/19	1250	23,000
35	Flint	MI	42.97	83.74	09/19	1315	28,000
36	Waterville	OH	41.45	83.64	09/19	1315	18,000 20,000
37	Buffalo	NY	42.93	78.65	09/19	1317	24,000 28,000
38	Dryer	OH	41.36	82.16	09/19	1437	18,000 21,000
39	Hayes Center	NE	40.45	100.92	09/19	1500	35,000 37,000
38	Dryer	OH	41.36	82.16	09/19	1538	18,000 23,000
40	Slate Run	PA	41.51	77.97	09/19	1543	21,000
41	Tidoute	PA	41.71	79.42	09/19	1543	18,000 33,000
42	Clarion	PA	41.15	79.46	09/19	1543	18,000 23,000
41	Tidoute	PA	41.71	79.42	09/19	1543	30,000 33,000
39	Hayes Center	NE	40.45	100.92	09/19	1543	31,500 37,600
43	McCook	NE	40.20	100.59	09/19	1553	28,000 33,000
44	Ottawa	Ont.	45.42	75.70	09/20	0123	19,000 23,000
45	S. Labrador	Lab.	52.90	62.33	09/20	2025	35,000
46	London	Ont.	42.98	81.25	09/20	2230	35,000 37,000
47	Saint John	N.Brs.	45.27	66.07	09/21	1123	30,000 35,000
48	Digby	Nova Scotia	44.55	66.79	09/21	1825	25,000

* report received at 2125 UT noting time of observation was 2200 UT.

COMMUNICATIONS ABOUT CLOUD MOVEMENT

Pilot reports and ground observations played a critical role in the tracking of the ash cloud from the September 16–17 eruption (table 3; fig. 5). The earliest altitude estimate of 15,000 ft (4.6 km) for the top of the main eruptive column was made by a pilot and was received at 0637 UT (table 3). A pilot report received at 0821 UT estimated the top of the eruption column at between 19,000 and 29,000 ft (5.8–8.8 km); this report corroborated the initial estimates of cloud tops at 28,000 ft (8.5 km), as determined from C-band radar at Kenai, Alaska (Rose and others, this volume). The ash cloud was tracked using images from the geostationary GOES satellite, images from the TOMS (Bluth and others, this volume), and from the AVHRR (Schneider and others, this volume) detectors aboard polar-orbiting satellites. Pilot reports (table 3; fig. 5), radiosonde analysis, and forecast upper-level wind data were incorporated from the beginning

of the eruption to indicate where the cloud would move (Heffter and Stunder, 1993). These observations were summarized in 18 volcanic hazard alert bulletins issued for the aviation community by NOAA's SAB (J. Lynch, NOAA, written commun., 1993). In addition to written descriptions of the ash cloud, these bulletins contained graphical information showing the location of the ash cloud.

For pilots in the air as well as those preparing for flight, the more succinct SIGMETs were the principal form of information about the ash cloud. These SIGMETs are in an abbreviated text format (fig. 4) and are derived from information in pilot reports, from AVO, from the volcano hazard alert bulletins, and from satellite images. For the September eruption of Crater Peak, the SIGMETs provide a rich source of data about the movement of the ash cloud. The initial SIGMET, ALFA 1, (fig. 4) was issued from Anchorage at 0648 UT on September 16, 12 minutes after the initial eruptive phase and 3 minutes after the AVO color code was changed to Red. That AVO was listed

WSPN1 PANC 170649

ANCA WS 170648

PAZA SIGMET ALFA 1 VALID 160640/161040 PANC-
ALASKA VOLCANO OBSERVATORY RPTS THAT MT SPURR VOLCANO AT
61.3N LAT/152.2 LONG HAS ERUPTED AT 1036Z.
TOPS INFORMATION NOT YET AVAILABLE. ANY ASH PLUME WL MOV NE-E.
A FOLLOW UP SIGMET TO BE ISSUED ASAP. NC.

WSPN1 PANC 170716

ANCA WS 170716 COR

PAZA SIGMET ALFA 1 VALID 160640/161040 PANC-
ALASKA VOLCANO OBSERVATORY RPTS THAT MT SPURR VOLCANO AT
61.3N LAT/152.2 W LONG HAS ERUPTED AT 0636Z.
TOPS INFORMATION NOT YET AVAILABLE. ANY ASH PLUME WL MOVE NE-E.
A FOLLOW UP SIGMET TO BE ISSUED ASAP. NC.

WSPN1 PANC 170844

ANCA WS 170841

PAZA SIGMET ALFA 2 VALID 160840/161240 PANC-
ALASKA VOLCANO OBSERVATORY RPTS THAT MT SPURR VOLCANO AT
61.3N LAT/152.2 W LONG ERUPTED AT 0636Z AND 0804Z. THE 0804Z
ERUPTION WAS MUCH STRONGER WITH AN ASH PLUME TO AT LEAST FL400
CONFIRMED BY WEATHER RADAR RHI. WINDS WL MOVE THE ASH TO THE ENE
THEN E. THUS VOLCANIC ASH IS PSBL BLW FL450 WI 90NM EITHER SIDE OF
A LN FM MT SPUR (61.3N 152.2W) TO ORT. INCRG.

WSPN1 PANC 171225

ANCA WS 171223

PAZA SIGMET ALFA 3 VALID 161240/161640 PANC-
AN ERUPTION OF MT SPURR VOLCANO AT 61.3N LAT/152.2 W LONG CONTINUED
FM 0804Z TILL NEARLY 1100Z WITH WEATHER RADAR INDICATING ASH TO FL 450
THRU THE PD. WINDS WL CONT TO MOVE THE ASH TO THE ENE THEN E.

VOLCANIC ASH IS PSBL BLW FL450 WI 60NM EITHER SIDE OF A LN
FM MT SPUR (61.3N 152.2W) TO GKN TO 62.0N 140.8W WITH THE MAX
CONCENTRATION XPCTD FM FL200-FL350. NC.

Figure 4. Notices of significant meteorological conditions (SIGMETs) issued by Anchorage National Weather Service Office for the September 16–17, 1992, eruption of Mount Spurr volcano, Alaska. Official record compiled by National Climatic Data Center, Asheville, North Carolina.

as the source of the information for the eruption indicates the good coordination between AVO and the National Weather Service during the eruption. No cloud-top information was given because the initial phase was not observed on radar. An error in the start time of eruption was corrected by reissuing SIGMET ALFA 1 at 0716 UT.

SIGMET ALFA 2 was issued at 0841 UT, 38 minutes after the onset of the main eruptive phase at 0803 UT (fig. 4), and it stayed in effect for the remainder of the eruption. Based on reports from the Kenai radar, SIGMET ALFA 2 informed pilots that

the top of the ash cloud had reached 45,000 feet. This SIGMET also contained a forecast of a 180-nautical-mile-wide track for the ash cloud based on the volcanic ash forecast model (Heffter and Stunder, 1993).

The first notification to pilots of the eruption's end at 1100 UT came from SIGMET ALFA 3 (fig. 4), issued at 1223 UT on September 16. All volcano SIGMET notices from Anchorage were canceled by a SIGMET issued at 0449 UT on September 18, stating that the last reported ash cloud sighting in Alaska was at 2200 UT on September 17.

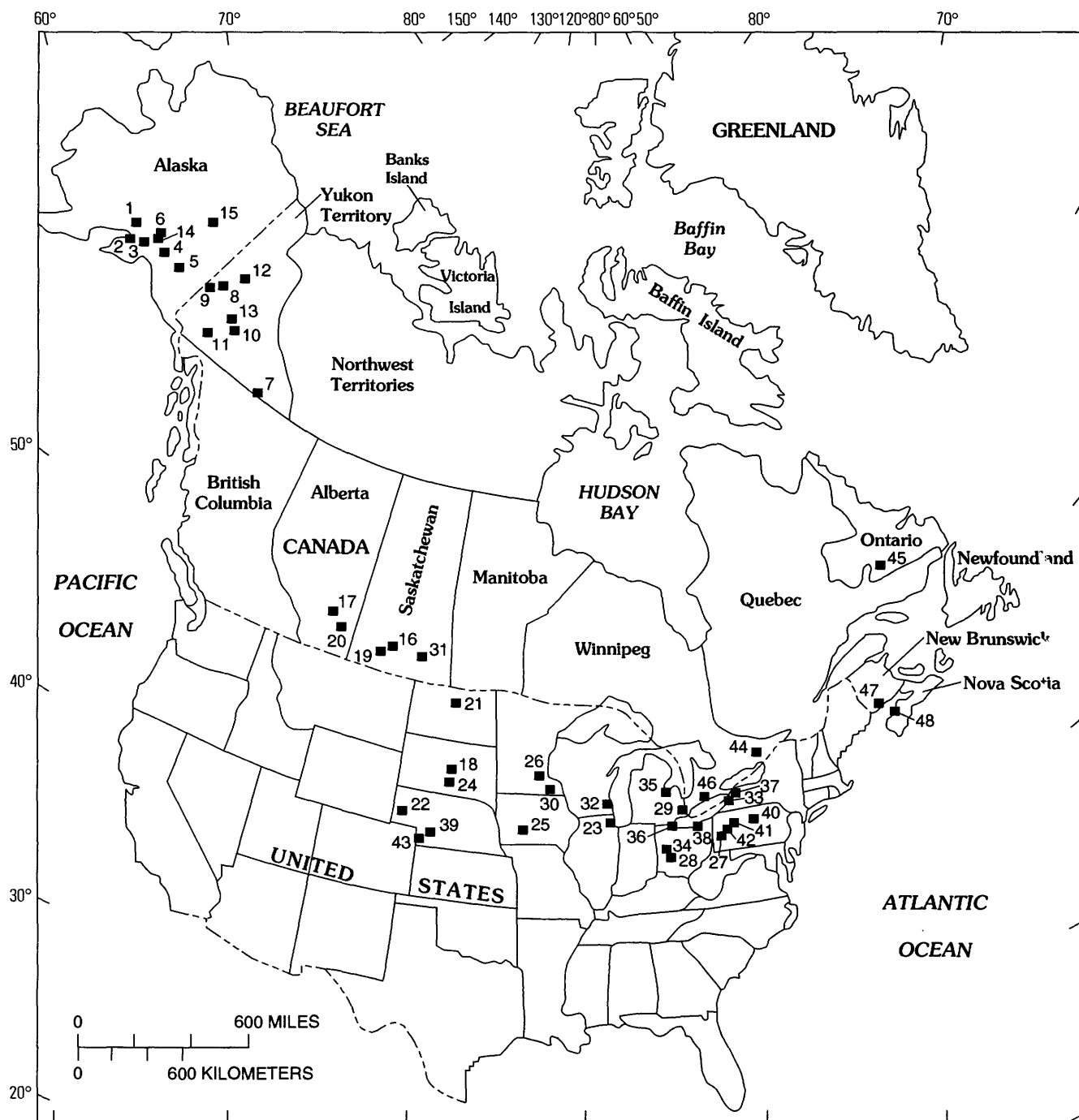


Figure 5. Positions of aircraft reporting visual observations of the September eruption cloud from Mount Spurr volcano, Alaska, as it traveled across North America. Numbers refer to locations of reporting sites listed in table 3. List compiled from data provided by G. Swanson (NOAA Synoptic Analysis Branch, written commun., 1993) and T. Spurgeon (T. Spurgeon, Transport Canada, written commun., 1993).

Satellite tracking of the cloud in Alaska and in the Yukon Territories was assisted by 24 reports from pilots and ground observers (table 3). At 0400 UT on September 18, after the cloud had been tracked into northwestern Alberta on satellite images, it became obscured by cirrus clouds (G. Swanson, NOAA, written commun., 1993). The next pilot sightings of the

cloud were from Lumsden, Saskatchewan, and Calgary, Alberta, at 1250 UT and 1430 UT on September 18. These reports were followed by observations of the cloud from several sites in South Dakota (table 3; fig. 5). This information was communicated to pilots through a series of SIGMETs issued from Canada and from the conterminous United States.

Eleven SIGMETs were issued by Canadian authorities starting at 1500 UT on September 17 for broad areas of British Columbia, Yukon Territories, Alberta, and Saskatchewan; these SIGMETs were based on computer model predictions for September 17 and 18. These SIGMETs stated that the ash cloud had weakened considerably, but they still advised caution. The Lumsden, Saskatchewan, pilot report (table 3) was issued as part of a SIGMET message from Edmonton at 1401 UT on September 18th, a gap of 14 hours from the last pilot report of a visual sighting at 0007 UT.

SIGMET XRAY 1, issued at 1445 UT on September 18 from both Salt Lake City and Chicago, was the first SIGMET issued in the conterminous United States, and was based on analysis of satellite imagery. This SIGMET was followed by SIGMET XRAY 2 issued at 1555 UT that was based on pilot reports over South Dakota that had been received at 1544 UT.

Twenty-five pilot reports of the cloud were recorded in the conterminous United States for September 18 and 19 (table 3), prior to the receipt of the first notification at the Cleveland Air Traffic Control Center at 1000 UT on Saturday September 19 (Howell, 1993). Fourteen additional pilot reports were recorded on September 19, including three from northeastern Nebraska. Additional pilot reports of the ash cloud continued through September 21 (table 3). The pilot reports agree well with satellite observations of ash cloud movement and validate the forecasts of cloud trajectory (Heffter and Stunder, 1993).

COHERENCY OF THE SEPTEMBER CLOUD

The remarkable coherency of the September ash cloud apparently resulted from meteorological conditions that persisted for several days after the eruption and which prevented the diffusion and dilution of the cloud. The ash cloud was injected into dense cirrostratus and altostratus clouds associated with a surface low-pressure system centered over the eastern Bering Sea. At the time of the eruption, the crest of an upper-level ridge was located at approximately 33,000 ft (10 km) over Mount Spurr. Winds at this level carried the ash east and northeast of the volcano (D. Helms, NASA, written commun., 1993).

GOES-7 satellite images showed the ash cloud as it passed beyond the layer of cirrostratus clouds over eastern Alaska and moved southeast in the polar-front jet stream toward Canada. Pilot reports documented the reappearance of the ash cloud during the morning of September 18 at about 1500 UT (figs. 5, 6; table 3). At that time, the cloud was observed forming one or more layers between 29,000 and 45,000 ft (8.8–13.7 km) in the area between western Saskatchewan and Lake Superior where an intensifying low pressure system was located.

On September 19, the ash cloud arrived in the airspace between the Great Lakes and the Mississippi Valley, trapped in the boundary of the polar jet stream. The low-pressure system that was located over Lake Superior on September 18 had moved northeast to-

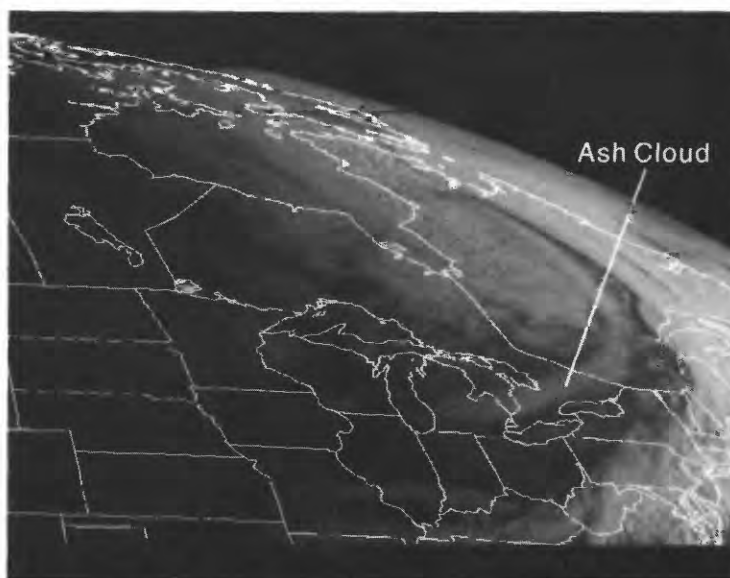


Figure 6. Image from GOES-7 satellite at dawn of September 19, 1992 (1200 UT) showing ash cloud from Mount Spurr volcano, Alaska. Cloud extends from eastern Michigan northeastward into the Canadian Maritime provinces.

ward Hudson Bay, Canada, and developed into a fully occluded system by September 19. During September 19, the ash cloud moved northeast over the St. Lawrence Seaway and ash particles were apparently concentrated because of the confluent nature of the upper-air flow (D. Helms, NASA, written commun., 1993). Through September 21, the cloud continued to move as a coherent mass visible on GOES-7 images as well as on AVHRR images (Schneider and others, this volume). Occasional pilot reports described the cloud over eastern Canada and even over the North Atlantic region (table 3; fig. 5).

EFFECTS ON AVIATION OPERATIONS

All three of the 1992 eruptions of Mount Spurr's Crater Peak vent temporarily disrupted air traffic over Alaska and western Canada. The ashfall from the August 18 eruption caused the temporary closing of the three Anchorage airports. The ash cloud from the September 16–17 eruption disrupted air traffic routing around the volcano, and several days later also disrupted traffic in the congested air corridors of the central United States, especially in the airspace managed by the Cleveland Air Traffic Center (ATC) (Howell, 1993).

AUGUST 18 ERUPTION AND ANCHORAGE AVIATION OPERATIONS

The August 18 eruption deposited approximately 1/16" to 1/8" (1 to 3 mm) of ash in the Anchorage area and resulted in the closing of the three Anchorage airports for several days. The recovery of Anchorage flight operations from the August 18 eruption required that time and money be spent to remove and clean ash from airport facilities and airplanes. The removal of ash from airport facilities and surfaces required a response effort similar to that used for snow removal. Details of the ash-removal operations at the Anchorage airports are reported in Casadevall (1993). In addition to costs associated with ash removal, the cancellation of flights resulted in the loss of revenue to both air carriers and the airport authority.

We tried to establish figures for the costs associated with the August 18 ashfall as one means to evaluate the potential benefits derived from monitoring and surveillance of the volcano as well as mitigation efforts. We report the costs associated with cleaning of the three Anchorage airports to return them to operational status, the costs associated with protecting and cleaning the planes on the ground, and the

costs to the air carriers of passenger revenue lost owing to cancellation of flights. Airport officials kept records of the costs associated with the cleanup of buildings and runway and taxiway surfaces at the three Anchorage airports (fig. 2). These costs related largely to employee compensation and minor costs for filters, spare parts, and maintenance of equipment used for cleanup activities and totaled approximately \$653,000. (L. Michou, Anchorage International Airport, in Casadevall, 1993).

Due to the lack of hanger space, about 30 passenger and cargo aircraft were caught unprotected on the ground at Anchorage International Airport and at Elmendorf. Prior to the ashfall, efforts were made to cover and protect these aircraft, and following the ashfall, the aircraft had to be cleaned (fig. 3) and inspected before they were returned to service. For wide-body jets, one manufacturer recommends a total of 30 person hours for preparing an airplane for ashfall (Casadevall, 1993). If a comparable time is spent cleaning an aircraft, the total time spent is approximately 60 person hours, which at a labor rate of \$20 per hour computes to a cost of approximately \$1,200 per wide-body airplane. For smaller passenger and cargo jets, we estimate approximately \$900 per airplane. Labor costs for protecting and cleaning the 20 wide-body and 10 smaller aircraft caught on the ground are estimated to total \$33,000.

The August 18 eruption and ensuing ash cloud caused cancellation of more than 100 domestic and international flights. Unfortunately, we have not been able to recover actual values associated with canceled operations such as the loss of landing fees, fuel-flowage fees, and concession fees. To estimate the revenue lost from these cancellations, we adopted the approach that was developed by Tuck and Huskey (1992) in an analysis of revenue loss attributable to the Redoubt Volcano eruption. In the Redoubt analysis, Tuck and Huskey (1992) calculated the value of \$46.03 per passenger for net revenues lost to the carriers through flight cancellation. Applying this value to the Mount Spurr situation, using an average of 2,000 passengers per day for the summer tourist season, times 3 days of lost revenue, yields an estimated revenue loss of \$276,000.

These costs are probably a conservative measure of several costs associated with disruptions to aviation operations due to the August 18 eruption. We have not included costs associated with loss of cargo service, loss of transit of international passengers, and loss of landing fees, fuel flowage fees, and concession fees. Secondary or indirect costs associated with the disruption, such as loss of revenues from car rentals, hotels, and meals were not considered.

SEPTEMBER 16–17, 1992, ERUPTION AND AIR TRAFFIC IN THE CLEVELAND AREA

Ash from the September 16–17 eruption caused a major disturbance on September 19 to air traffic managed by the Cleveland Air Traffic Center (ATC) (Howell, 1993). This center is responsible for en-route operations in an area bounded by Saginaw, Michigan; Syracuse, New York; Elkins, West Virginia; and Fort Wayne, Indiana. This airspace includes the major terminal areas of Detroit, Cleveland, and Pittsburgh.

The first report of a volcanic ash cloud was received at Cleveland Center on September 19, 1992 at 1000 UT (6:00 a.m. EDT) as a SIGMET from the National Weather Service. The ash cloud was reported to extend from Milwaukee, Wisconsin to Sherbrooke, Quebec, a distance of 710 nautical miles (1,320 km), at an altitude of 30,000 ft (9.1 km) (figs. 5, 6). Sub-

sequent pilot reports through September 19 track the ash cloud moving steadily to the southeast. The last position reported for the cloud was over the Johnstown, Pennsylvania, area at about 1600 UT (table 3). The altitude of the cloud from pilot reports varied from 18,000 ft to 33,000 ft (5.5 to 10 km) and higher.

Because the precise altitude of the top of the ash cloud was not well known, aircraft reporting to the Cleveland center elected to descend below the cloud rather than to climb and fly over the cloud. When en-route aircraft were forced to descend below the base of the ash cloud at 18,000 ft (5.5 km), they entered the low-altitude sectors designed for managing the flow of arriving and departing aircraft. With this influx of aircraft from the high-altitude sectors, the low-altitude sectors quickly became saturated with air traffic (fig. 7). This influx caused numerous delays of as much as 30 minutes for aircraft departing terminals under

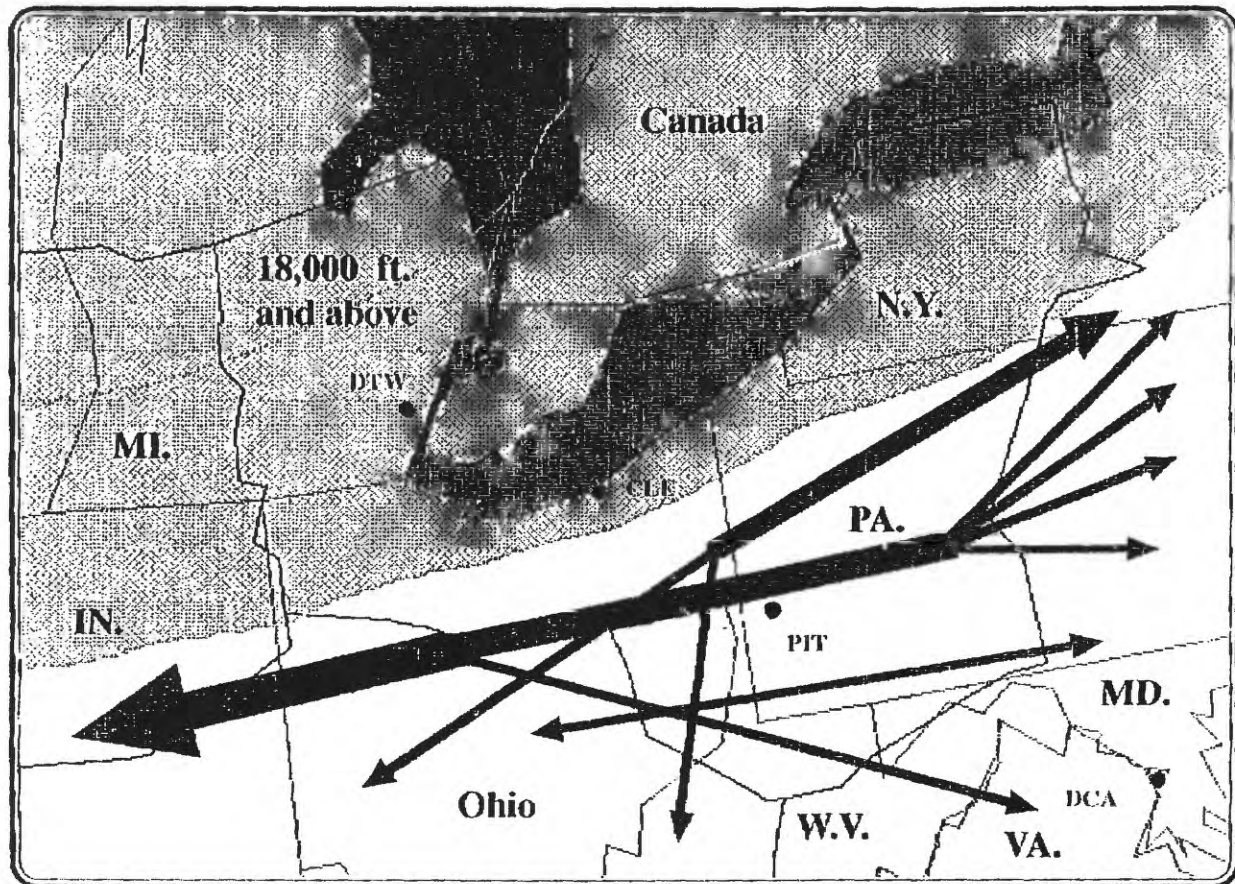


Figure 7. Schematic diagram of airspace around Cleveland Area Traffic Control Center, September 19, 1992 (1317 UT) showing compression of air routes as a result of the southward incursion of the Mount Spurr ash cloud (adapted from Howell, 1993). Dark shaded area shows outlines of Great Lakes. Light shaded area shows extent of ash cloud as interpreted from satellite imagery. Principal air routes are shown with heavy black line; arrows indicate direction of movement; line thickness indicates relative loads along the routes. DTW = Detroit; CLE = Cleveland; PIT = Pittsburgh; DCA = Washington/Dulles.

the control of the Cleveland ATC. As the ash cloud continued eastward, more pilots requested changes to their flight plans and additional disruptions to normal air-traffic flow took place. Air traffic operations in the airspace controlled by the Cleveland ATC returned to normal by 2300 UT on September 19 (Howell, 1993).

DISCUSSION

The three ash clouds produced by the Crater Peak vent of Mount Spurr in 1992 were each comparable in size and volume. The effects each cloud had on aviation operations were largely determined by wind dispersal patterns. The June eruption cloud had relatively little effect on air traffic because the southerly winds moved the cloud north away from busy air routes. The August eruption affected Anchorage airports because the west winds carried the ash into the Anchorage area. The September ash cloud affected the busiest en-route region in North America. Fortunately, by the time of the third eruption, the hazard notification system between the ground and air observers in and around Alaska had been used several times and was working efficiently (R. Hutcheon, NOAA, oral commun., 1993). However, in the conterminous United States, where the September ash cloud had not been expected, the air-traffic control system was forced to adjust to the cloud with additional disruption to service. Concerns about the safety aspects of this unique cloud, including its age and far-traveled nature, prompted the FAA to convene a special workshop on volcanic ash clouds held in April 1993 (Federal Aviation Administration, 1993b).

The eruption in September was distinctive because the meteorological conditions kept the ash cloud intact and visible to pilots along major air routes for about 100 hours after the eruption (table 3). The September cloud affected the air-traffic system of the conterminous United States as it passed through the heavily traveled north-central air corridor. However, no mechanical problems were recorded and the Cleveland ATC responded in an exemplary way to this major distortion of air routes. Flight delays were the only consequence of the September eruption.

The lack of damaging encounters between aircraft and ash clouds from the 1992 Crater Peak eruptions underscores the recent improvements in the ways that volcanic hazards are being addressed by personnel concerned with aviation operations. Many of the improvements in the hazard warning operations came following the 1989-90 Redoubt eruptions from national and international efforts (Casadevall, 1992, 1994a; ICAO, 1992; Casadevall and Oliveira, 1993; Hickson,

1994) to address the problem of aircraft encounters with volcanic ash. Specific advances include improved monitoring of Cook Inlet volcanoes by the Alaska Volcano Observatory (AVO, 1993); improved monitoring of meteorological conditions in the Cook Inlet region by the National Weather Service (Hufford, 1994); increased awareness and reporting of ash clouds by pilots (Fox, 1994); and refinement and utilization of computer-based forecasting of ash cloud trajectories and particle distribution (Heffter and others, 1990; Heffter and Stunder, 1993; Ellis and others, 1993; Tanaka, 1994).

Since the Redoubt eruptions, the Anchorage office of the National Weather Service has improved its ability to detect and track ash clouds (Hufford, 1994). These improvements include the installation in 1992 of a portable C-band radar on the Kenai Peninsula, which provided immediate detection of the Mount Spurr eruption columns (Rose and others, this volume). Installation of a NEXRAD (WSR-88D) weather radar system (Klazura and Imy, 1993) at Kenai in January 1994, offers several technical advances over conventional C-band weather radar and should enhance the ability to detect the onset of an ash-producing eruption from Cook Inlet volcanoes, the height of the eruption column, and the initial movement of the ash cloud away from the volcano. A newly installed wind profiler at Kenai combined with conventional radiosonde data and with wind data from aircraft overflying the area provide improved wind data for trajectory forecast models. Finally, the satellite downlink capabilities and digital-image processing have been upgraded both at the Anchorage National Weather Service office (Hufford, 1994) and at the University of Alaska in Fairbanks (Dean and Whiting, 1994) for operational application to detect and track ash clouds.

Moreover, the communication of information and warnings among agencies concerned with aviation safety has improved significantly since the 1989-90 eruptions of Redoubt Volcano (Casadevall and others, 1992; FAA, 1993a). In particular, the AVO volcano updates and color code and the volcano hazard bulletins issued by the SAB have been developed and tailored to the requirements of the aviation community in terms of timeliness and message content. In addition, agencies in Alaska have established inter-agency procedures and operational plans that specifically address the needs of the Alaskan theater of aviation operations (Alaska Interagency Operating Plan, 1993). Recent improvements have also been made in communications between remote field stations and aviation operations in Russia and Alaska (Miller and Kirianov, 1993) in a continuing effort to improve air safety in the North Pacific region.

Future research to mitigate ash-cloud hazards should include testing of the NEXRAD radar for detecting and tracking ash clouds, validation of remote-sensing data such as AVHRR and TOMS data, and validation of trajectory forecast models. This research would benefit from in-situ study of ash clouds, including use of unmanned vehicles to sample ash clouds (Riehle and others, 1994). More attention must be given to improve seismic monitoring of volcanoes in remote areas. Field studies are also needed to evaluate the explosivity of volcanoes, especially in remote regions such as the Kurile-Kamchatka-Aleutian volcanic arc, which parallels some of the world's busiest air routes.

SUMMARY

Ash clouds produced by the three eruptions of Mount Spurr's Crater Peak vent were comparable in size and volume to the ash clouds from the December 1989 eruptions of Redoubt Volcano (table 2), which damaged five commercial jets. Fortunately, the 1992 Crater Peak eruptions caused no damage to aircraft en-route. Like several of the Redoubt eruption clouds, the three Mount Spurr clouds also entered Canadian airspace. Two of the clouds entered airspace over the conterminous United States and significantly disrupted the flow of air traffic. This frequent incursion of ash clouds from eruptions of Cook Inlet volcanoes into Canadian and conterminous U.S. airspace highlights the need for timely communications about the threat to aviation safety from drifting ash clouds. In response to these events, U.S. and Canadian meteorologists, volcanologists, and aviation officials continue to develop procedures to provide timely warnings and advisories to pilots about drifting clouds of volcanic ash.

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APPENDIX

CLOCK TIME CONVERSION TABLE

ALASKA STANDARD TIME (AST) (UTC-9 hours)	ALASKA DAYLIGHT TIME (ADT) (UTC-8 hours)	UNIVERSAL TIME (UTC)
3:00 p.m.	4:00 p.m.	00:00
4:00	5:00	01:00
5:00	6:00	02:00
6:00	7:00	03:00
7:00	8:00	04:00
8:00	9:00	05:00
9:00	10:00	06:00
10:00	11:00	07:00
11:00	12:00 m	08:00
12:00 m	1:00 a.m.	09:00
1:00 a.m.	2:00	10:00
2:00	3:00	11:00
3:00	4:00	12:00
4:00	5:00	13:00
5:00	6:00	14:00
6:00	7:00	15:00
7:00	8:00	16:00
8:00	9:00	17:00
9:00	10:00	18:00
10:00	11:00	19:00
11:00	12:00 p.m.	20:00
12:00 p.m.	1:00	21:00
1:00	2:00	22:00
2:00	3:00	23:00
3:00	4:00	00:00

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