



EXPLANATION

- Thermal-maturity units—R_r, reflectance value in oil; CAI, Conodont Color Alteration Index. See text for explanation.
- Undermature—R_r < 0.6, CAI < 1.0
- Mature I—R_r 0.6-1.3, CAI 1.0-2.0
- Mature II—R_r 1.3-2.0, CAI 2.0-3.0
- Overmature—R_r 2.0-3.6, CAI 3.0-4.0
- Supermature—R_r 3.6-5.0, CAI 4.0-4.5
- Igneous/Metamorphic—R_r > 5.0, CAI > 4.5
- No data
- Quaternary sediments (unconsolidated)
- Clastic cover
- Thermal-maturity unit contact—Dashed where inferred, dotted where certain
- High-angle fault—Dashed where inferred, dotted where concealed
- Thrust fault—Dashed where inferred, dotted where concealed
- Detachment fault—Dashed where inferred, dotted where concealed
- Contact as UMI unit boundary—Elevation of UMI thermal maturity map-unit boundary (0.6% vitrinite reflectance isograd). Values, meters below sea level.
- Vitrinite reflectance
- Conodont Color Alteration Index
- Selected oil well with vitrinite reflectance data—All elevations and vitrinite reflectance values are from regression lines drawn through the data; this line may be extrapolated slightly below the well bottom.
- Well name
- Bochard #1—Kelly bushing elevation (meters) above sea level
- R_r SURF(70%): 0.25—Vitrinite reflectance at surface (%)
- R_r T(200°C): 1.77—Vitrinite reflectance at total depth (%)
- UMI: 255 m—Hole bottom, meters below sea level
- MIMC: 252 m—Contact between thermal map units, meters below sea level
- Other oil well—Used to constrain contoured UMI thermal maturity map-unit boundary (0.6% vitrinite reflectance isograd) in Cook Inlet and North Slope regions
- City

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- Oil well
- Thrust fault
- Basement
- Boundary between thermal maturity units, dotted where inferred
- Other oil well—Used to constrain contoured UMI thermal maturity map-unit boundary (0.6% vitrinite reflectance isograd) in Cook Inlet and North Slope regions

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SYNOPSIS OF THERMAL MATURITY PATTERNS IN ALASKA

This map was produced from nearly 10,000 vitrinite-reflectance and CAI determinations from surface, offshore, and subsurface localities across the State. From these data, a number of generalizations can be made. First, thermal maturity is uniformly low in the Tertiary interior basins and in the Alaskan foreland and backarc basins. In contrast, thermal maturity is high in the Tertiary coastal basins and in the Yukon-Koyukuk and Colville Basins. The Yukon-Koyukuk and Colville Basins show elevated levels of thermal maturity at the surface, commonly with higher values in basin margins. This geometry suggests a pattern of greater uplift along basin margins, possibly reflecting tectonic readjustments as coastal lands are removed by erosion.

Johnson and others (1993) investigated thermal maturity relations in three sedimentary basins in the Colville-Cook Inlet and Kuskokwim basins. Their results indicate that thermal maturity is generally higher in the Tertiary coastal basins than in the Tertiary interior basins. This pattern may reflect the effects of the tectonic evolution of the principal geotectonic units of uplift in the central Brooks Range; greater uplift further to the south is indicated by the presence of proclinal facies and high-grade metamorphism. This pattern may also reflect the effects of the tectonic evolution of the principal geotectonic units of uplift in the central Brooks Range; greater uplift further to the south is indicated by the presence of proclinal facies and high-grade metamorphism. This pattern may also reflect the effects of the tectonic evolution of the principal geotectonic units of uplift in the central Brooks Range; greater uplift further to the south is indicated by the presence of proclinal facies and high-grade metamorphism.

Thermal maturity is a measure of the level of thermal alteration of organic matter in sedimentary rocks. It is expressed as different types of organic matter response differently to heat, thermal maturity is operationally defined differently for different substrates. Commonly, organic matter is measured in terms of vitrinite reflectance (R_r) and Conodont Color Alteration Index (CAI). The thermal maturity of rocks is related to the level of thermal alteration of organic matter in sedimentary rocks. It is expressed as different types of organic matter response differently to heat, thermal maturity is operationally defined differently for different substrates. Commonly, organic matter is measured in terms of vitrinite reflectance (R_r) and Conodont Color Alteration Index (CAI).

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SOURCES OF DATA

Data for this map have been compiled from the literature and field data, unpublished industry reports provided to the Alaska Division of Geological and Geophysical Survey, previous USGS investigations, data contributed by industry, academic and government colleagues, and by our own sampling and analysis where needed. Quantitative data used in the preparation of this map were restricted to vitrinite-reflectance and Conodont Color Alteration Index analyses because these types of data are the most common and best quantifiable thermal-maturity data available in Alaska. The database includes data from Johnson and others (1993), consisting of 3,716 vitrinite-reflectance determinations from 2,123 outcrop localities, 1,474 Conodont Color Alteration Index determinations from 1,306 outcrop localities, and 4,482 vitrinite-reflectance determinations from 217 wells.

Where vitrinite-reflectance and Conodont Color Alteration Index data were sparse or unavailable, other organic and inorganic thermal-maturity data from the literature cited below were used to help constrain thermal maturity. Such data include Rock-Eval pyrolysis (T_{max}), thermal alteration index (TAI), fluid inclusion, illite crystallinity, vitrinite reflectance, and fluid-inclusion data. In particular, the fluid-inclusion data from the Colville-Cook Inlet Basin (Bacon (1994)) generally correlate with the vitrinite-reflectance and Conodont Color Alteration Index data, and Duetz-Bacon (1994) used vitrinite-reflectance and Conodont Color Alteration Index data to constrain thermal maturity in the Yukon-Koyukuk Basin. The geologic base (metamorphic, plutonic, and unmetamorphosed rocks, surficial units, and faults) is modified from Duetz-Bacon (1994) and Beckman (1985).

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1996

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