

GEOLOGICAL SURVEY CIRCULAR 610

Gold in Igneous, Sedimentary, And Metamorphic Rocks

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By Robert S. Jones

GEOLOGICAL SURVEY CIRCULAR 610

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CONTENTS,

Page

TABLES

Page

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ABSTRACT

Gold analyses were compiled for about 50 rock types of which more than 50 percent are igneous, about 30 percent are sedimentary, and the rest are metamorphic. The analyses for gold made before 1955 tend to be higher than the newer analyses, so separate lists are given. At present, minute amounts ofgold in rocks seem to be best determined by neutron-activation methods of analysis, and the following data are based on such analyses:

Igneous rocks contain as little as 0.2 and as much as 73 ppb (parts per billion) gold. The average, calculated on the basis of 50 percent granite and 50 percent basalt, is 3.0 ppb. Averages (in parts per billion) for individual plutonic rock types are: granite, 2.8; syenite, 2.5; diorite, 3.5; gabbro, 5.4; anddunite, 8.2. Averages (in parts per billion) for individual volcanic rock types are: rhyolite, 12; trachyte, 6.5; andesite, 5.2; and basalt, 3.2.

Sedimentary rocks contain as little as 0. 3 to as much as 41 ppb gold. The average of 5.0 ppb is based on the average of 7.5 ppb for sandstone, 3.9 ppb for shale, and 3.5 ppb for limestone.

In metamorphic rocks the range of gold content is from 0.86 to 22.4 ppb with an average of 4. 3. The average (in parts per billion) for gneiss is 1.8; for schist, 5.0; for quartzite, 4.8; and for marble, 3.1.

In relatively nonauriferous regions, the compiled gold analyses of rocks show only a few parts per billion, whereas in relatively auriferous regions the analyses tend to be one to two magnitudes higher.

INTRODUCTION

This report was prepared as background material for the Heavy Metals program of the U.S. Geological Survey, an intensified program of research on new sources of heavy metals, particularly gold. Data on the occurrence of gold in igneous, sedimentary, and metamorphic rocks are summarized. The number of gold analyses made on common types of rocks is meager compared with analyses of other elements, and many of the data are of questionable quality.

ACCURACY OF ANALYSES

Analyses of the gold content of rocks made prior to 1955 were made by fire-assay, spectrographic, and wet chemical methods. These methods are less sensitive than those used at present, and the results obtained seem to be higher than those given by the newer methods. The gold beads produced by fire assay were measured by gravimetric, colorimetric, or volumetric

m ethods until Hagen (1954) introduced the spectrographic method. Since 1954 a few analyses have been by methods similar to Hagen's, but most have been made by the neutron-activation method. Analyses presented in this report are grouped in tables according to date; analyses made before 1955 are'shown in tables 2, 6, and 10, and those made since 1954 are shown in the remaining tables.

Analyses of the U.S. Geological Survey's standard rock samples G-1 and W -1 (Fairbairn and others, 1951) by several methods are given in table 1. The values for G-1 range from 0.002 to 0.0112 ppm (part per million), and those for W-1 range from 0.0048 to 0.010 ppm. All an-alyses are by neutron-activation methods except for those of Das S arm a, Sen, an d Chowdhury (1965) , which were made by fire assay followed by spectrographic analysis, and those of Brown and Wolstenholme (1964), which were made by mass spectrometer.

GOLD IN IGNEOUS ROCKS

Gold analyses of igneous rocks are given in tables 2 and 3. The older analyses listed in table 2 for igneous rocks were made by fire assay, and comparison with some of the recent results indicates that the older analyses are very much higher. A major source of error may be gold impurities in blanks, especially in the test lead used in assaying, causing high values. Hillebrand (in Emmons, 1886) and Andrew (1910) warned that it was impossible, at that time, to obtain test lead that did not contain appreciable amounts of gold.

The analyses by Hagen (1954) were made by fire assay followed by spectrographic analysis. He did considerable study on losses that might occur during the assay and applied corrections for them but apparently did not consider the possibility of gold contamination in the reagents. Hagen's values are consistently higher than those given for similar rocks by neutron activation. No comparisons of identical samples are available, but the following comparison may be made for rocks from the same localities.

²Vincent and Crocket (1960b).
³Baedecker (1967).

Assuming that the samples were similar in gold content, it would seem that Hagen's results are high by a factor of 20 to 30.

Gold seems to be heterogeneously distributed to a marked degree in some samples. Clifton, Hubert, and Phillips (1967), for example, found that the gold content of splits from a single sample of Oregon beach sand differed by as much as 29 ppm $\langle 0, 1 \rangle$ ppm to 29 ppm). De Silva (quoted in Baedecker, 1967) reported large discrepancies in neutron-activation analyses for gold even after the samples had been pulverized to achieve greater homogeneity. He divided one irradiated aliquot into four parts and found that one part contained 90 percent of the total activity of the sample. This lack of homogeneity, rather than analytical errors, is probably the cause of the discrepancies in the analyses of some samples, such as the differences (noted by Baedecker, 1967) in analyses of U.S. Geological Survey standard rocks G-1 and W-1 (table 1) and in four analyses of a diabase dike which showed 3.3, 140, 3.4, and 6.8 ppb gold (Shcherbakov and Perezhogin, 1963).

Clarke and Washington (1924) estimated the gold content of igneous rocks to be in the range 1 to 10 ppb, and this order of magnitude has been substantiated by later workers. The average amount of gold in igneous rocks has been estimated by Berg (1932) to be 6 ppb, by Noddack and Noddack (1934) to be 4 ppb, and by Mason (1952), Rankama and Sahama (1950), Polanski (1948), Anderson (1945), Goldschmidt (1937), and Fersman (1933) to be 5 ppb. Goldschmidt (1954) computed a mean value of 2 ppb gold based on the gold to silver ratio of 20:1; this ratio is based on the proportion of these two elements in important ores. Horn and Adams (1966) , with the aid of a computer and assumptions based partly on a modification of Goldschmidt's (1933) material balance, have derived a figure of 3.57 ppb of gold in igneous rocks, which is only slightly higher than the value of 3.0 ppb obtained by averaging the values for granite and basalt in table 5.

Shcherbakov and Perezhogin (1964) reported that, in the Altai-Sayan folded belt in the U.S.S.R., gold is more abundant in extrusive rocks than in their intrusive equivalents. For example, 14 rhyolites and trachytes contain an average of 5.4 ppb gold, compared with the average of 3.8 ppb gold in 33 granites and eight syenites; 19 intrusive basalts average 10 ppb gold compared with the average of 6.4 ppb gold in 14 Altai-Sayan gabbroids. Boyle (1960) reported that the Prosperous Lake granite in the Yellowknife district of Canada contains 10 ppb gold compared with 100 ppb gold in a quartz-feldspar porphyry.

The gold content of the Skaergaard intrusives in Greenland were studied by Vincent and Crocker (1960a), who concluded that they varied little even though the intrusives have undergone strong magmatic differentiation. The initial magma, represented by a chilled border, contained 4. 6 ppb gold. Most of the other rocks and minerals contained between half and twice this amount. However, Crocker, Vincent, and Wager (1958) believed that the gold content of the layered gabbros

Table 1.-Gold content of standard granite G-1 and standard d1abase W-1

[Analyses by neutron-activation methods unless specified]

Gold content (ppm)	Reference		
1 Granite G-l	2 Diabase W-1		
0.006	0.010	Vincent and Smales (1956).	
.0045	.0084	Vincent and Crocket $(1960a)$.	
	.0049	Hamaguchi and others (1961).	
.0046	.005	Shcherbakov and Perezhogin (1963) .	
3 < 0.08	3 < 0.08	Brown and Wolstenholme (1964).	
.0020	.0048	Baedecker and Ehmann (1965).	
4.0112	.0053	Das Sarma, Sen, and Chowdhury (1965).	
5 .0057	5.0064		

¹U.S. Geol. Survey standard granite from Weşterly, R.I.

U.S. Geol. Survey standard diabase from Cegterville, Va.

Analysis by mass spectrometer; not included in average. 4 Analysis by fire-assay and spectrographic

methods.

Average of values except <0.08 ppm.

is constant and tends to be evenly distributed among the various silicate and oxide phases. They thought that the gold is present in uncharged randomly distributed metallic atoms in the magma rather than as charged ions or complexes. In the middle part of the intrusion they found evidence of an appreciable concentration of gold in the immiscible copper sulfide droplets in contrast to the lack of concentration of gold in the iron sulfide droplets which separated later.

The lowest average gold content for granites (1.2 ppb) has been reported by Anoshin and Potap'yev (1966). The granites are from widely separated massifs in the U.S.S.R. (the Kolyvan' massif in thenorthwesternpart of the Altai and the Khangilay -Shilinskii massif in east Transbaikal).

Shcherbakov and Perezhogin (1963) examined 28 granites and concluded that the gold content of massifs bearing gold deposits are similar to those of barren massifs. The gold content of the younger granite phases, which are considered the source of the gold mineralization in the region, are lower than the gold content of the older mafic rocks.

Mantei and Brownlow (1967) showed that the Marysville quartz diorite stock in Montana has an appreciably higher background value for gold than the rocks of the Altai-Sayan folded belt in Asia (Shcherbakov and Perezhogin, 1963) or those of the Skaergaard intrusives of Greenland (Vincent and Crocket, 1960a). The gold content is greatest in and near the periphery of the stock, where it averages 71 ppb; elsewhere it averages 10 ppb.

Shcherbakov and Perezhogin (1964) used data obtained from igneous rocks of the U.S.S.R. to compare the gold content with the content of iron, magnesium, copper, vanadium, lead, zinc, and some unspecified metals. It is not clear whether they analyzed the same rocks for these metals and for gold or whether they used values furnished by Zavaritskii (1955). The metals and oxides that show correlation with gold, in decreasing order of correlation, are: Cu, Fe203, MgO, V, and FeO. Shcherbakov and Perezhogin (1964) stated, "It should be noted that the correlation coefficient between the gold content and the content of another cation decreases as the bonding between the cation and oxygen becomes more strongly ionic." In addition they noted that the patterns of gold distribution are different in the intrusive and extrusive rocks so that when the data for the extrusive rocks are introduced, the correlations between gold and all the other components are much lower.

Table 4 gives the average gold content of igneous rocks based on all known published analyses made since 1954. The analyses represented by each average are probably too few to give reliable background values for the igneous rocks. For example, the average for nepheline syenites (1.8 ppb gold) is based on only two analyses, and the value 100 ppb for quartz-feldspar porphyry is a single analysis.

Table 5 shows the average gold content of various igneous rock types as determined by the neutronactivation method only. The value for granites is an average of 13 individual averages of granites from 13 localities in North America and the U.S.S.R.: Algoma District, Ontario; Westerly, R.I.; Wausau, Wis.; Stone Mountain, Ga.; Altai; Turochak, southeast Altai; Aksai, southeast Altai; Kuznetsk, Altai; northwest Altai; east Transbaikal; Tuva, Soviet Asia; Sayan (Mount Sukhaya); and Tashtuzek, southeast Altai. The average gold content of basalts (table 5) was derived from 11 individual averages representing basalts in Jefferson County, Colo.; Cooper Falls, Oreg.; East Pacific Rise; Mauna Loa, Hawaii; Altai, U.S.S.R.; central Altai, U.S.S.R.; Lintz, Germany; Morvern, Scotland; Giant's Causeway, Ireland; and two parts of the mid-Atlantic ridge.

The data for volcanic rocks show a direct relationship between gold and silica, whereas the relationship seems to be inverse for the plutonic rocks (table 5). The relatively large amount of gold in the rhyolites may be the result of selective sampling.

GOLD IN SEDIMENTARY ROCKS

Data on gold in sedimentary rocks are presented in tables 6 and 7. Table 6 gives the analyses made prior to 1955; table 7, the analyses made since 1954.

Reported gold content of sedimentary rocks ranges from at least 82 ppm for conglomeratic ores from the Black Hills, S. Oak., to as little as 0.0003 ppm for a sandstone from Mount Darwin, Tasmania.

Many of the conglomerates represented in table 6 are from ore-producing areas, and some may, in places, contain enough gold to be mined profitably. More than 500 fire-assay-atomic-absorption analyses have been made for gold in conglomerates in northwestern Wyoming (table 7). The average gold content of the Harebell Formation is estimated at 55 ppb, and that for the younger Pinyon Conglomerate is estimated at 80 ppb (Antweiler and Love, 1967).

The analyses for gold in sandstones range from 0.3 ppb in a sandstone from Tasmania (table 7) to 446 ppb for a sandstone from Victoria, Australia (table 6). The average amount of gold detected in sandstones by neutron-activation analysis is 7.5 ppb (table 8). Siltstones are represented by two values (table 8), one of 5.5 ppb and the other 6.2 ppb (Shcherbakov and Perezhogin, 1963).

Amounts of gold reported in shales and lutites range from 2.3 ppb to 374 ppb (tables 6 and 7). The lutites, from shallow cores from the ocean floor, contain variable amounts of gold, but from a depth of 1 to 10.45 meters the gold content increases with depth from3.1 ppb to 17.2 ppb (DeGrazia and Haskin, 1964).

(Text continues on p. 15.)

Table 2.-Analyses of gold in igneous rocks made before 1955

[Some of these rocks may be metamorphic in origin. All analyses are by fire assay only unless specified]

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Table 2.--Analyses of gold in igneous rocks made before 1955--Continued

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by Lincoln (1911).

phase IV (0.0011 ppm).

Table 3.--Analyses of gold in igneous rocks made since 1954

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[All analyses are by neutron activation unless specified]

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Table 3.--Analyses of gold in igneous rocks made since 1954--Continued

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Rock	Location	Gold content (ppm)	Number оf samples	Remarks	Reference
	South Africa: Norite-------- Transvaal-----------------	0.0029	ı	Bushveld complex-----------------	DeGrazia and Haskin (1964) .
	Canada: Diabase ------- Yellowknife, Northwest Territories.	\sim 01	$---$	Late diabase dikes. Analyses by fire assay.	Boyle (1960).
	United States: Do-------- Centerville, Va-----------	.0065	ı	U.S. Geol. Survey standard rock W-1. Fire assay, spectro- graphic, and neutron activa- tion. Value is average of more than 10 analyses.	Vincent and Smales (1956), Vincent and Crocket (1960a), Hamaguchi and others (1961), Shcher- bakov and Perezhogin (1963), Baedecker and Ehmann (1965), and Das Sarma, Sen, and Chowd- hury (1965).
$Do-------$	U.S.S.R.: Altai---------------------	.0058	$\mathbf{2}^{\circ}$	Sarakokshinsk massif. Phase V.	Shcherbakov and Perezhogin
				Duplicate analyses.	(1963) .
Do--------	Nizhnyaya Tunguska River--	.0080	ı	Anakita massif, fine-grained part of the chilled border (duplicate analyses).	Shcherbakov and Perezhogin (1964) .
$Do-------$	Kuznetsk, Ala-Tau---------	.0384	$\mathbf{1}$	Cm_1 . Quadruplicate analyses, with values of 0.0033, 0.0034, 0.0068 , and 0.14 ppm.	Shcherbakov and Perezhogin (1963) .
	United States:				
Basalt--------	Jefferson County, Colo----	.0040	$\mathbf{1}$	Contains olivine-----------------	DeGrazia and Haskin (1964) .
$Do-------$	Cooper Falls, Bridal Veil quadrangle, Oregon.	.0006	ı	U.S. Geol. Survey standard BCR-1 (triplicate analyses reported as 0.6 ppm). Iridium content reported as 0.5 ppm, probably is 0.0005 ppm.	Baedecker (1967).
	Do-------- East Pacific Rise----------- Hawaii:	.0004	1	Iridium content, 0.0002 ppm ------	Do.
	Do-------- Mauna Loa-----------------	.0026	ı	Tholeiitic; olivine present (duplicate analyses).	Vincent and Crocket $(1960b)$.
		.0059	ı	Flow of 1859. Tholeiitic---------	Baedecker (1967).
	Do-------- -----do---------------------	.0028	ı	Flow of 1868. Tholeiitic-picrite-	Do.
	Do-------- -----do--------------------- and the control of the components	.0039	1.	Flow of 1881 near Hilo. Tholei- itic, hypersthene present. Iridium content, 0.0021 ppm.	Do.
	U.S.S.R.: Do-------- Altai--------------------	.0006	ı	Duplicate analyses---------------	Shcherbakov and Perezhogin (1963) .
$Do-------$	Central Altai ------------	.0089	ı.	Duplicate analyses. Basalt flow--	Do.
$Do-------$	Altai-Sayan folded belt (mainly).	.010	19	These are all intrusive----------	Shcherbakov and Perezhogin (1964) .
Do-------- Germany:	Lintz, Rhenish Prussia---	.0026	\mathcal{L}		DeGrazia and Haskin (1964) .

Table 3.--Analyses of gold in igneous rocks made since 1954--Continued

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Table 3.--Analyses of gold in igneous rocks made since 1954-Continued

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Table 4.⁻⁻⁻Average gold content of igneous rocks based on all known published analyses made since 1954

[Data from table 3. Average Au: First figure in brackets is the number of areas from which the samples were taken, and the second figure is the number of neutron-activation analyses made on these samples]

1

¹A value of 6.2 ppb Au is obtained if the anomalous value of 140 ppb is not included (see table 3).

[Analyses by neutron activation only]

This value is increased to 5.7 if the anomalous figures (Shcherbakov and Perezhogin, 1964) of 39 and 16 ppb Au are used in the calculation. 2Value given by Shcherbakov and Perezhogin

(1964).

Leutwein (1951) made spectrographic analyses of 11 Thuringian alum- and silica-rich shales. The median amount of gold in these shales was 100 ppb. Pyrite separated from them contained no detectable gold, but gold was found in the hydrocarbon fraction. Seven sediments of Holocene age taken from the Baltic and North Seas, also analyzed by Leutwein, contained 500 to 2,500 ppb gold, which seems high. (See "sapropel," table 6).

Tischendorf (1959) found that black graptolitic shales near Tilkerode, Harz, Germany, contained 300to 1,300 ppb gold, whereas bleached shales from the same area contained from less than SO to 800 ppb (table 7). Six shales from the Harebell Form at ion in northwest Wyoming had an average of 72 ppb gold, and 18 shales from the Pinyon Conglomerate averaged 52 ppb gold (Antweiler and Love, 1967). Noddack and Noddack (1931) reported 10 ppm gold in a cupriferous shale from Germany. Gold values for clay range from 3.12 ppb (computer-derived geochemical balance, Horn and Adams, 1966) to 817 ppb.

Gold in lateritic ores of Ity, Ivory Coast, is finely dispersed in amounts up to 18.3 ppm according to available data. Bacteria may have played a role in the deposition of this gold (Pares and Martinet, 1964), as indicated by the results of the treatment of the ores and by the gold precipitated with cultures of microorganisms from streams and soils of the region. Experiments have shown that the initial solubility ofthe gold up to concentrations of 6 ppm is followed by slow reprecipitation.

Gold occurs in coal of all geological periods (Babička, 1943). Goldschmidt (1935) considered that the maximum factor of enrichment of gold in coal ash over the average gold content of the earth's crust is between 40 and lOOtimes. Thyssen-Bornemisza (1942) reported that some plants have the ability to concentrate gold.

Widely separated coals have yielded appreciable amounts of gold. Gold has been reported in coals from Pleasant Valley, Utah, and Kemmerer and Cambria, Wyo. (Gibson and Selvig, 1944; Jenney, 1903; Stone, 1912). Stone reported that samples from 31 carloads of coke (estimated one-third ash) from Cambria contained from 0.10 to 0.28 oz per ton (4 to 11 ppm) of gold, but more recent analyses have not confirmed these high figures. The split and boney coal from Cambria contained more gold than that with lower ash contents. The gold may be associated with pyrite and possibly derived from a slightly auriferous sandstone roof above the coal. (Text continues on p. 22.)

Table 6.--Analyses of gold in sedimentary rocks made before 1955

[Analyses are by fire assay unless specified]

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Table 6.--Analyses of gold in sedimentary rocks made before 1955--Continued

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 $\label{eq:2.1} \mathcal{L}(\mathcal{L}^{\text{max}}_{\mathcal{L}}(\mathcal{L}^{\text{max}}_{\mathcal{L}})) \leq \mathcal{L}(\mathcal{L}^{\text{max}}_{\mathcal{L}}(\mathcal{L}^{\text{max}}_{\mathcal{L}}))$

 $\label{eq:2.1} \frac{1}{2} \int_{\mathbb{R}^3} \frac{1}{\sqrt{2}} \, \mathrm{d} x \, \mathrm{d$

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Table 7.--Analyses of gold in sedimentary rocks made since 1954

[All analyses are by neutron activation unless specified]

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Bouska, Havlena, and Sulcek (1963) concluded that gold is bound predominantly to the organic part of coal. The coals they studied were found in fresh-water Cenomanian deposits in Bohemia and Moravia. Oeul (1958) found gold in the ash of the humic acid fractions of peat. He did not detect it in plant ashes, in the whole peat ash, or in any other fraction that he separated from the peat. In seven samples, Deul obtained from 1 to 10 ppm gold in the ash of the humic acid extract.

Goldschmidt and Peters (1933) gave the gold content of the ash of coal as 0.5 to 1 ppm and the gold content of clarain, a constituent of coal, as 0.5 ppm (table 6). Trace amounts of gold have been detected in the ash of crude oils from Poland (Kisielow and Gregorowicz, 1955).

Most carbonate rocks contain less gold than sandstones and shales (table 8). The amount of gold reported in carbonate rocks ranges from 0.8 ppb for a coral from the Florida Keys (table 7) to 256 ppb for a limestone from England (table 6). Two limestones from the Harebell Formation in northwest Wyoming had 20 ppb gold, and 11 limestones from the Pinyon Conglomerate in the same area contained 41 ppb gold (Antweiler and Love, 1967). Horn and Adams (1966) estimated the average gold content for carbonate rocks at 1.79 ppb.

Early analyses show that marl from the Keuper and Muschelkalk at Rancourt, France, contains a large amount of gold (ranging from 4 to 36 ppm) and also appreciable silver (table 6). The analyses for gold in gypsum and rock salt (table 6) are suspiciously high. An analysis for gold in anhydrite, however, seems reasonable (Lincoln, 1911). A manganese nodule obtained by the H. M. S. Challenger Expedition was found to have 0.2 ppm gold (Goldschmidt and Peters, 1932).

Of the sedimentary rocks, carbonates probably contain on the aver age the least gold, and sandstones contain the most (table 9). This conclusion is based on the neutron-activation analyses of Shcherbakov and Perezhogin (1963) and DeGrazia and Haskin (1964) and on the computer-derived geochemical balance of Horn and Adams (1966) .

GOLD IN METAMORPHIC ROCKS

Fewer metamorphic rocks have been analyzed for gold than either sedimentary or igneous rocks. The amounts detected range from 0.86 ppb in a marble (DeGrazia and Haskin, 1964; table 11), determined by neutron-activation methods, to 24 ppm in a schist (Levat, cited by Lincoln, 1911; table 10). Estimates of the average amounts of gold in various metamorphic rock types are given in table 12.

One analysis of a mica-garnet gneiss showed 1.8 ppb gold (DeGrazia and Haskin, 1964). Older analyses of gneisses gave 0.830 to 23 ppm gold (table 10).

Slates from gold mining regions in New Zealand and Australia had almost equal average amounts of gold: New Zealand, 111 ppb, and Australia 124 ppb (Don, 1898). Composite samples of slates and argillites from the Yellowknife district, Northwest Territories, Canada, averaged 10 ppb gold (Boyle, 1960). An argillite from Asia contained 8.3 ppb gold (Shcherbakov and Perezhogin, 1963).

Shcherbakov and Perezhogin (1963) detected 1.2 ppb gold in a phyllite. Chlorite schists from widely separated areas analyzed by the same workers contained about the same amount of gold (5.3 and 6.7 ppb). Don (1898) found an average of 89 ppb gold in 79 schists from New Zealand. Shear zones at Yellowknife, Northwest Territories, Canada, yielded 17 ppb gold in chlorite schist and 632 ppb gold in a carbonate-sericite schist (Boyle, 1960). Levat, cited by Lincoln (1911), reported large amounts of gold and silver in hornblende schists (table 10). Palei and others (1967) analyzed eight schists for gold. The amounts found range from 0.8 to 9 ppb and averaged 3.1 ppb. Hornfels from the Altai in the U.S.S.R. had a gold content ranging from 1.8 to-22.3 ppb.

Both the lowest and the highest gold values given in this report for metamorphic rocks as analyzed by neutron-activation methods are for marble: 0.9 ppb and 22.4 ppb (table 11). One quartzite sample from near Wausau, Wis., contained 7.3ppbgold, and another from near Baraboo, Wis., contained 2.4 ppb gold (DeGrazia and Haskin, 1964). These values are in the same order of magnitude as most values reported for gold in sandstones in tables 8 and 9.

Table 8.-Average gold content of sedimentary rocks

[Data from table 7. Average Au: First figure in brackets is the number of areas from which the samples were taken, and the second figure is the number of neutron-activation analyses made on these samples]

~Average including lutite and cla~, a.6 ppb. ³ 39 samples from the U.S.S.R.

Average including oolite, carbonate, and coral, 3.0 ppb.

Table 9.—Average gold content of sedimentary rocks, from recent data

[Data in parts per billion]

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Table 10.--Analyses of gold in metamorphic rocks made before 1955

[Analyses are probably by fire assay unless specified]

Gold Number of Rock Location content Remarks Reference (ppm) samples U.S.S.R.: Argillite----- Southeast Altai----------- 0.0083 1 Duplicate analyses--------------- Shcherbakov and Perezhogin (1963). Canada: Argillites Yellowknife, Northwest . 01 Analyses by fire assay----------- Boyle (1960) . $- - -$ Territories. and slates. U.S.S.R.: Yenisei Range------------- Duplicate analyses--------------- Phyllite------ .0012 1 Shcherbakov and Perezhogin (1963). Canada: Schist-------- Yellowknife, Northwest . 017 Boyle (1960) . $- - -$ Average of chlorite-schist Territories. phase of shear zones. Analyses by fire assay. Do-------- -----do--------------------- .632 Average of carbonate-sericite Do. schist phase of shear zones. Analyses by fire assay. D o-------- -----do--------------------.01 Knotted quartz-mica schist. Do. $- - -$ Analyses by fire assay. U.S.S.R.: Do-------- Northeast Altai----------- .0053 $\overline{1}$ Chlorite schists (duplicate Shcherbakov and Perezhogin analyses). (1963). Do-------- Kuznetsk, Ala-Tau--------- -----do-------------------------- .0067 1 Do. Do-------- Sultanuizdag (Uzbekistan)- .0031 Schists with two or more of the Pale1, Murovtsev, and 8 following minerals: quartz, Borozenets (1967). mica, chlorite, plagioclase, and unspecified carbonate mineral(s). Range, 0.0008 to 0.009 ppm Au. United States: Gneiss-------- New York, N.Y------------- Mica-garnet gneiss--------------- .0018 $-$ DeGrazia and Haskin (1964). U.S.S.R.: Hornfels------ Central Altai------------- .0070 Duplicate analyses--------------- Shcherbakov and Perezhogin 1 (1963). Do-------- -----do--------------------- .0223 Do. 1 Amphibole hornfels formed from basalt (duplicate analyses). Do-------- Southeast Altai----------- . 0018 1 Formed from shale (duplicate Do . analyses) . -----do-------------------------- Do-------- -----do---------------------. 0027 1 Do. United States: ----------------------------------DeGrazia and Haskin Quartzite----- Rib Mountain, Wausau, .0073 $---$ Wis. (1964). Do-------- Devil's Lake, Baraboo, .0024 -----------------------------------Do. $- - -$ Wis. United States: Marble-------- Rutland (Belden's Farm), . 00086 -----------------------------Do .

Table 11.-Analyses of gold in metamorphic rocks made since 1954

[All analyses by neutron activation unless specified]

tv *a-.*

Vt.

 $\label{eq:2.1} \frac{1}{2} \sum_{i=1}^n \frac{1}{2} \sum_{j=1}^n \frac{$

Table 12.-Average gold content of metamorphic rocks

[Data from table 11. Average Au: First figure in brackets is the number of areas from which the samples were taken, and the sec-ond figure is the number of neutron-activation analyses made on these samples]

¹If the anomalous value of 22.4 ppb Au (table 11) is included in the calculations, the gold content becomes 9.6 ppb.

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