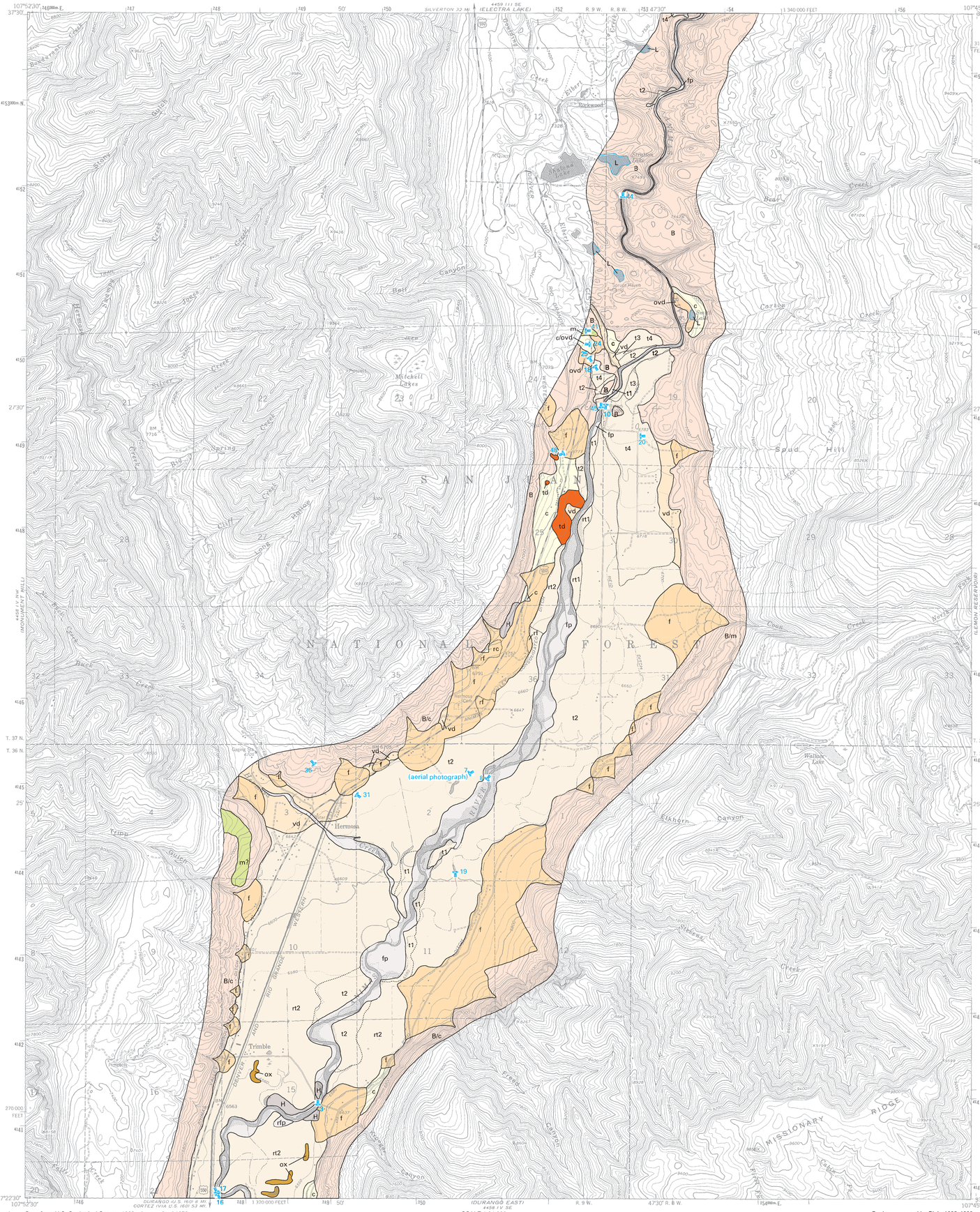


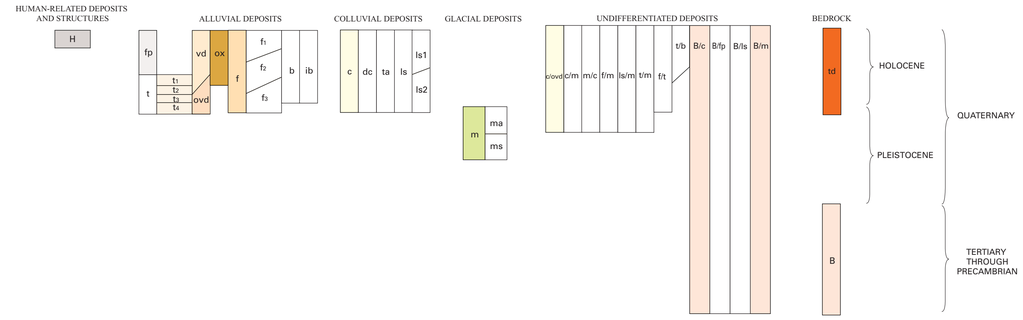
Base from U.S. Geological Survey, 1963; photorevised 1979
Projection: Universal Transverse Mercator, zone 13.
10,000-foot grid based on Colorado coordinate system, south zone.
1927 North American datum

SCALE 1:24 000

CONTOUR INTERVAL 40 FEET
NATIONAL GEODETIC VERTICAL DATUM OF 1929

Geology mapped by Blair, 1998–1999.
Digitally compiled by Yager, 1998–1999.

CORRELATION OF MAP UNITS



DESCRIPTION OF MAP UNITS

[Unit descriptions do not differ between plates 1 through 10 with the exception of bedrock, which does differ based on lithologies and geologic formations represented in a mapped area. Prefix "r" indicates alluvial, colluvial, or glacial deposits disturbed by human activities. For example, rfp refers to reworked floodplain deposits from gravel mining operations. Unit symbols separated by a slash indicate that two units are preserved in mapped area; color in box denotes predominant unit. Uncolored boxes denote units not mapped on this plate]

HUMAN-RELATED DEPOSITS AND STRUCTURES

H — **Human works (upper Holocene)**—Includes human-caused deposits and structures larger than 10 m in length, such as fluvial mine tailings, ramps, buildings, open pits, and mine waste piles. Structures are either intact or in a state of decay and consist of wooden planks, stone or concrete pads with iron, steel, or aluminum sheets, beams, rebar, wire, and a variety of discarded machinery. Fluvial mine tailings are well sorted and consist of fine-grained sand, silt, and clay-size sediment. Mine waste piles consist of poorly sorted, sand- to boulder-size fragments with a silt matrix; minerals such as pyrite, galena, sphalerite, hematite, jarosite, and goetite are common. The sulfide and oxide minerals impart yellow, orange, and reddish color to these dumps. Human works were constructed between 1870 and the present. Natural processes such as snow avalanches, local flooding, and slope failure aided in the modification and destruction of historic human works. Some human works such as adits and mine dumps are point sources for acidic waters, which discharge elements such as Al, Fe, Cu, Zn, As, and Pb to streams

ALLUVIAL DEPOSITS

fp — **Floodplain deposits (upper Holocene)**—Unconsolidated sand and gravel, commonly with silt matrix; poorly sorted, subround to round gravels, commonly imbricated; boulders as much as 0.5 m in diameter common in upper Animas River watershed. Deposited in active stream channels and as overbank sediment on adjacent terraces less than 1 m above channel. Mostly deposited during spring floods and storm discharge events. Natural, fluvial processes throughout the watershed and, in places, human activities, such as between Eureka and Howardsville, have reworked many floodplain deposits since 1860. Red (iron oxide), white (aluminum hydroxide), and black (manganese oxide) staining of floodplain sediments in some reaches indicates high concentrations of dissolved metals. Thickness 1–10 m

t (t1, t2, t3, t4) — **Terrace deposits (Holocene to uppermost Pleistocene)**—Unconsolidated sand and gravel with silt matrix; poorly sorted, subround to round gravels, commonly imbricated; boulders as much as 0.5 m in diameter common in upper watershed reaches. In highly mineralized drainages, iron-oxide cements form ferricrete (iron-cemented conglomerate). Sediment is deposited during spring runoff and storm over-bank flow discharge events. Surfaces are narrow and flat, and parallel the modern stream channel. Where differentiated by their relative ages, youngest terraces are labeled t1, lie 1–2 m above the modern channel, and predate 1860 (Kirk Vincent, U.S. Geological Survey, written commun., 2001). Sequentially higher and older terraces are designated t2, t3, and t4. Higher terraces display stronger color and textural B soil horizons. Terrace height above the stream channel varies from 2 to 10 m. Terraces in the headwater regions of the Animas River watershed are not necessarily correlative with terraces in the lower Animas River basin. Thickness 1–10 m

vd — **Valley deposits (upper Holocene)**—Unconsolidated sand and gravel with a silty matrix and occasionally larger stones derived from valley sides. Deposits formed by fluvial and colluvial processes including periodic floods, small debris flows, and slope wash. Deposits often found on slightly concave valley floors near the heads of tributaries where no distinct fluvial terraces or alluvial fans can be found. Deposits are both historic and pre-historic in age. Thickness 1–5 m

ovd — **Old valley deposits (Holocene to uppermost Pleistocene)**—Unconsolidated sand and gravel with a silty matrix and occasionally larger stones derived from valley sides. Deposits formed by fluvial and colluvial processes including periodic floods, small debris flows, and slope wash. Same as valley deposits (vd) in mode of formation except unit ovd occupies older concave surfaces now perched above the modern valley floor. Thickness 1–5 m

ox — **Oxbow deposits (Holocene)**—Unconsolidated silt and sand, mixed with peat near the surface; discontinuous gravel lenses preserved at depth. Deposits occupy abandoned meanders that become oxbow lakes when flooded. Found in lower Animas Valley. Deposits are both historical and pre-historical in age. Older oxbow deposits commonly have thicker soil and peat horizons at the surface and eventually are filled with silt to become dry swales on the floodplain surface. Thickness 1–10 m

f (f1, f2, f3) — **Alluvial fan deposits (Holocene to uppermost Pleistocene)**—Unconsolidated, poorly sorted sand and gravel with silt matrix; cobbles and boulders common. Sediment supplied episodically to fan surfaces by streams and debris flows along confined, commonly braided channels from seasonal storm runoff events. Form as fan-shaped deposits (in map view) at mouths of streams where stream gradient and velocity decreases and where topography widens from confined channels to open valleys. Surface gradients are less than 20°. Fan development post-dates glaciation and therefore their growth was initiated after 15 ka. Where the relative fan age can be differentiated from cross-cutting geometry, the younger fan segments are designated as f1 and the older as f2 and f3. Thickness 1–20 m

COLLUVIAL DEPOSITS

c — **Colluvial deposits (upper Holocene to uppermost Pleistocene)**—Unconsolidated soil and sand- to cobble-size debris derived from valley walls. Processes include rockfall, sheetwash, and creep. Colluvium is discontinuous on the valley sides and produces irregular-shaped, hummocky surfaces. Growth has been continuous since glacial retreat. Thickness 1–5 m

GLACIAL DEPOSITS

m (ma, ms) — **Glacial till deposits (upper Pleistocene)**—Unconsolidated to compacted, poorly sorted, clay- to boulder-size sediment derived from ablation of glaciers at ice margins. Form hummocky end and lateral moraine ridges. Best displayed at north end of Durango and near Silverton. Deposits represent the last vestiges of glaciers between 12 and 15 ka. Carrara and others (1984) and Maher (1972) indicated ice-free conditions existed as early as 15 ka, but recent data indicate ice-free conditions at around 12 ka (Gillam, 1998). Till deposits north of Durango are subdivided into two relative ages. Animas City moraines (ma) are 12–35 ka and Spring Creek moraines (ms) are 85–160 ka (Gillam, 1998; Carroll and others, 1999). Thickness 1–20 m

UNDIFFERENTIATED DEPOSITS

c/ovd — **Colluvium and old valley deposits (upper Holocene to uppermost Pleistocene)**—See descriptions for colluvium (c) and old valley deposits (ovd)

B/c — **Bedrock and colluvial deposits (upper Holocene to Precambrian)**—See descriptions for bedrock (B) and colluvial deposits (c)

B/m — **Bedrock and glacial till deposits (upper Pleistocene to Precambrian)**—See descriptions for bedrock (B) and glacial till deposits (m)

BEDROCK

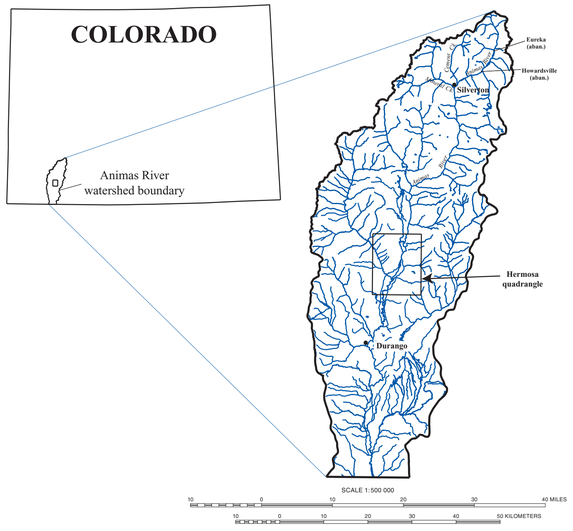
td — **Travertine deposits (Upper Holocene to Upper Pleistocene)**—Vuggy, thinly laminated calcareous tufa. Calcium carbonate deposited from super-saturated thermal spring waters associated with algae growth to produce pores and cavities. Deposits form tufa mounds along valley sides. Thermal springs are associated with fractured bedrock in lower Animas Valley. Water temperatures range from 32° to 44° Celsius (Barrett and Pearl, 1976). Growth has been continuous since last glacial retreat

B — **Bedrock (upper Paleozoic to Precambrian)**—Bedrock forms the exposed valley walls and occasional knobs on the valley floor with thin to no soil cover. Where soil cover is thin, vegetation cover may be present. Bedrock consists mainly of Paleozoic conglomerate, sandstone, shale, siltstone, mudstone, and limestone. Precambrian granite exposed at type locality of Bakers Bridge (Steven and others, 1974)

Contact—Dashed lines are contacts between geomorphic features in a map unit

L — Lake

35 — **Site of photograph**—Camera symbol is located where photograph was taken and camera symbol lens is approximately oriented in direction of photograph. Click on each photograph icon to view a linked .pdf-file photograph and site description



SURFICIAL GEOLOGY ALONG THE ANIMAS RIVER IN THE HERMOSA QUADRANGLE, ANIMAS RIVER WATERSHED, LA PLATA COUNTY, COLORADO

By
Robert W. Blair, Jr., and Douglas B. Yager
2002