

# DESCRIPTION OF THE DANVILLE QUADRANGLE.

## GENERAL RELATIONS.

The Danville quadrangle receives its name from Danville, Illinois, the county seat of Vermilion County. It embraces the country lying between parallels  $40^{\circ}$  and  $40^{\circ} 15'$  and meridians  $87^{\circ} 30'$  and  $87^{\circ} 45'$  and has an area of about 228 square miles. It is situated mainly in Vermilion County, Illinois, but extends about  $1\frac{1}{2}$  miles into Indiana and embraces parts of Warren and Vermilion counties in that State. This quadrangle is in a region of comparatively low altitude considering the fact that it is distant several hundred miles from the seaboard, the highest point being scarcely 760 feet and the lowest about 485 feet above tide. It is within <sup>Geographic position.</sup> the Mississippi Basin, about 200 miles east of the Mississippi River. It is situated in the great Illinois-Indiana coal field, near its northeastern margin, and contributes liberally to its output.

## TOPOGRAPHY.

*Drainage.*—The streams in this quadrangle belong to the Wabash River system, which drains to the Ohio, and this in turn to the <sup>Rivers.</sup> Mississippi. The Vermilion River is the avenue of discharge of the entire quadrangle, with the exception of a few square miles along the southern and eastern borders. This river is formed by the union of three streams, known as North Fork, Middle Fork, and Salt Fork. The name Vermilion River is commonly applied only to the portion below the junction of Middle and Salt forks.

*Relief.*—The surface of this quadrangle exhibits three distinct types of topography: prairies, ridges rising above the general level, and river valleys cut in the plain.

The prairies are particularly prominent in the region south and west of Vermilion River, and are also well developed east of that <sup>Prairies.</sup> stream as far north as Danville. A small area of similar character is found in the vicinity of Batesburg and Hillery, and the northwest corner of the quadrangle extends a short distance into the broad expanse of prairie which lies mainly north of this quadrangle.

Crossing this territory from northeast to southwest is a low, broad ridge, which has an elevation of about 90 feet above the prairie in <sup>Ridges.</sup> the vicinity of Danville. As seen from the south this ridge is rather prominent, but on its northern flank the level of the prairie is only slightly below the crest of the ridge.

Crossing indifferently the prairies and the ridge are deep valleys carved by Vermilion River and its principal tributaries. In some <sup>Valleys.</sup> respects these are the most conspicuous features of the topography; they have destroyed much valuable farming land of the prairies through the sharp backward cutting of the minor streams, and they have been decided barriers to easy transportation; but, on the other hand, they have opened excellent geologic sections, in which are shown the beds of coal that make this region so important economically. The valleys are generally narrow, but frequently swell out into broad amphitheaters a mile or more in width. They vary in depth from 50 to 150 feet, and their walls are generally steep and sometimes precipitous.

## GEOLOGY.

### OUTLINE OF GEOLOGIC HISTORY.

#### SEDIMENTARY ROCKS.

The rocks which underlie this quadrangle seldom appear at the surface; they are generally so deeply covered with clay and sand that their presence is not appreciated until they are reached by the drill in prospecting for mineral deposits or for water. South of the latitude of Danville, however, the rocks may be seen at many points in the bluff along the streams, in almost perpendicular cliffs of shale or shaly sandstone. Coal beds are of common occurrence in such exposures, and even under the heavily drift-covered plain they have been revealed by the

drill in a number of places. The entire rock series exposed at the surface and extending some distance below belongs to that portion of the geologic column known as the Carboniferous system. Beneath the coal-bearing rocks the drill reveals the presence of heavy beds of limestone, but they are not seen at the surface in this quadrangle.

The coal-bearing rocks occupy a broad, shallow syncline, the center of which is some distance southwest of Danville. As a result of this structure the rocks in this quadrangle have a very gentle southwesterly dip, toward the center of the basin.

#### EVENTS OF MESOZOIC AND EARLY CENOZOIC TIME.

After the deposition of the Carboniferous rocks, which must have occurred in some body of water, the crust of the earth was raised in <sup>Emergence of the land area.</sup> the Appalachian region and this area became dry land. In this condition it was subjected to the varying vicissitudes of a land surface for many geologic periods, but there is little evidence now existing to show the changes through which it passed.

Before the advent of the great ice sheet which modified the larger part of the surface of Indiana and Illinois, the surface of the Danville quadrangle was reduced to a gently rolling country, with a relief of less than 200 feet. The valleys were rather broad and the slopes were gentle, the entire topography strongly resembling that of southern Indiana, beyond the limit of glacial ice.

#### PLEISTOCENE.

The mantle of drift which covers the indurated rocks of Canada and the northern part of the United States bears evidence of much complexity in its history. Instead of a single deposit formed by one ice advance, it is now known that several distinct drift sheets are present which represent a succession of ice advances. The intervals of deglaciation or disappearance of ice between these ice advances are made apparent by the presence of soils and of beds of peat and marl and other effects of life, and also by the weathering of certain zones now buried in the midst of the drift deposits; while the sheets themselves differ markedly in extent and in physical constitution. In fact, the criteria for separating the drift deposits into distinct sheets are fully as reliable as the criteria for separating the indurated rocks into distinct formations. How far to the north the ice melted back, or, in other words, how extensive was the deglaciation, has not yet been fully determined for any of the intervals, but the ice is known to have withdrawn in each case nearly if not quite to the northern border of the United States, while in the most prolonged intervals it may have nearly or quite disappeared from Canada. For reasons not yet fully understood, the lines of flow of successive ice fields are in some cases far from coincident, so that in places the later ice field extended beyond the limits of its predecessor, while in other places it fell short of its predecessor's limits. These features alone would suggest that there had been more than one ice invasion, but their testimony is strengthened by the corroborative evidence of the buried soils and weathered zones.

The deposits of the later ice invasions have so greatly concealed those of preceding ones that the knowledge of the older deposits is somewhat meager and is confined largely to their outlying portions. Enough is known, however, to make clear certain distinctive characteristics in each drift sheet. These distinctive characteristics have been found helpful in drawing the line between contiguous drift sheets where the soil and weathered zone of the older drift surface had been removed by the later advance of the ice. These differences are in some cases merely differences in texture or in color, but often there is found a striking contrast in the kinds of rock present in the two drift sheets, a certain rock being rare or wanting in one sheet but abundant in the other.

The divisions of the drift series just mentioned are stratigraphic and are thus similar to the usual

divisions of the indurated rocks. But there are present in the drift deposits peculiar features not found in other deposits, and these have been made a basis for subdivisions. The features referred to are moraines or massive ridges of <sup>Moraines.</sup> drift, built up by the ice at its margin. These ridges recur at frequent intervals in passing north from the extreme edge of a given sheet of drift, and mark places of halting, and perhaps of readvance, which interrupted the melting away of the ice field. These morainic ridges have in some cases been formed in rapid succession and constitute a morainic system. Indeed, in Illinois there is a decided tendency to such grouping of the morainic ridges, as will appear below.

The sheets of drift formed by each of the ice invasions, the soils and weathered zones formed between the drift sheets, and the moraines and morainic systems of each drift sheet have received geographic names from localities where they are well displayed, in conformity with the prevailing custom of naming the indurated rock formations. The following outline includes the several drift sheets and deglaciation intervals recognized in North America.

#### OUTLINE DESCRIPTION OF THE DRIFT SHEETS AND INTERVALS.

*Stage 1.*—The oldest drift sheet (pre-Kansan), including perhaps the Albertan of Dawson. The limits of this sheet may be inside those of the Kansan in Iowa and neighboring States.

*Stage 2.*—The Aftonian soil and weathered zone. This soil was first recognized and brought to notice at Afton, Iowa, where it separates the Kansan drift from the pre-Kansan. Its distribution is mainly west of the Mississippi River.

*Stage 3.*—The Kansan drift sheet. This drift sheet is found mainly west of the Mississippi River. Its western limit is in Kansas and Nebraska and its southern limit in central Missouri. Possibly a sheet of similar age occurs at the base of the drift in northern Illinois, but the evidence is far from decisive.

*Stage 4.*—The Yarmouth soil and weathered zone. In western Illinois and eastern Iowa the Kansan drift sheet is overlapped by a later sheet, the Illinoian, which covers a soil and weathered zone that had been formed on the older till sheet. The evidence of this interval was first discovered at Yarmouth, in southeastern Iowa, from which the name has been derived. This stage of weathering and soil formation appears to have antedated the first glaciation of the region of which the Danville quadrangle is a part.

*Stage 5.*—The Illinoian drift sheet. This sheet is so named because of its extensive development in Illinois outside the limits of the Wisconsin and Iowan drift sheets. It extends apparently to the glacial boundary in western Ohio, Kentucky, Indiana, and Illinois, and forms the eastern border of the driftless area in southern Wisconsin and northwestern Illinois. It extends but a few miles west of the Mississippi River. In the Danville quadrangle it apparently constitutes the basal portion of the drift.

*Stage 6.*—The Sangamon soil and weathered zone. This soil and weathered zone separates the Illinoian drift sheet from the next sheet of drift, the Iowan, and its associated loess. The soil is a conspicuous feature (long ago described by Worthen) in the Sangamon Valley of Illinois, from which it takes its name. It is also a well-known feature of the drift of eastern Illinois and western Indiana in the vicinity of the Danville quadrangle.

*Stage 7.*—The Iowan drift sheet and associated silt deposits. This sheet, which is so named because of its wide extent in Iowa outside the Wisconsin drift, is elsewhere largely concealed beneath that drift. It is exposed to view outside the Wisconsin sheet, however, in a few counties in northwestern Illinois and south-central Wisconsin. Associated with it is a silt deposit widely known as loess, which extends far beyond the limits of the Iowan drift, overlapping the outlying portions of earlier drift sheets and extending down the Mississippi Valley to the Gulf of Mexico. The Iowan drift probably falls short a few miles of reaching the Danville quadrangle, but its associated silt or loess should underlie the Wisconsin drift in this quadrangle unless removed by erosion and the Wisconsin glaciation.

*Stage 8.*—The Peorian soil and weathered zone. This interval of deglaciation receives its name from Peoria, Illinois, near which city its weathering is well displayed. In addition to the presence of the soil and weathered zone, evidence of this interval is found in the great dissimilarity of the outline of the borders of the two drift sheets which it separates and in a change in the attitude of the land, by which drainage conditions became greatly improved.

*Stage 9.*—The Wisconsin drift, with its several morainic systems. This drift sheet is so named because of its excellent development in the State of Wisconsin, where it was first separated from earlier drift sheets by Prof. T. C. Chamberlin.

It extends southward into Illinois as far as Shelbyville and Paris, or 30 to 35 miles beyond the southern limit of the Danville quadrangle.

In each of the ice invasions above outlined, the portion of the great ice field which covered the quadrangle under discussion is <sup>Illinoian</sup> <sup>glacial lobe.</sup> known as the Illinois glacial lobe. The axis of movement for this lobe appears to have been along the basin of Lake Michigan, from which it extended southwestward into Illinois and spread into neighboring portions of Wisconsin and Indiana. It even pushed a few miles beyond the Mississippi into Iowa at the culmination of

the Illinoian stage of glaciation. Its border on the east is at the driftless portion of southern Indiana, which extends a few miles north of Bloomington, Indiana. Its limits on the west are at the driftless area of southwestern Wisconsin and northwestern Illinois, south of which they are marked by a marginal ridge or moraine which lies within a few miles of the Mississippi River from near Clinton, Iowa, down to St. Louis, Missouri. The Danville quadrangle is situated east of the middle of this lobe.

In the Illinoian and Iowan drift sheets of the Illinois glacial lobe moraines are comparatively insignificant features, the greater part of the drift surface being exceedingly flat, but in the Wisconsin drift several prominent morainic systems were developed, which form a rudely concentric series in northeastern Illinois and western Indiana. These, given in regular order of development, from the outermost moraine to the border of Lake Michigan, are as follows:

#### MORAINIC SYSTEMS OF THE WISCONSIN DRIFT.

1. The Shelbyville morainic system. This constitutes the outermost system of the Wisconsin series in the Illinois lobe and receives its name from the city of Shelbyville, which stands at the southwestern point of the morainic loop. In this system is included the Cerro Gordo moraine, which coalesces with the main member for a few miles east of Shelbyville, but is elsewhere distinct.

2. The Champaign morainic system. This system includes a series of small drift ridges which are separated into three distinct members in southeastern Champaign and neighboring portions of Vermilion, Edgar, and Douglas counties, Illinois. From the city of Champaign northwestward there is but a single ridge, and this apparently passes under the Bloomington system a few miles east of Bloomington. The line of the inner ridge of this system crosses the southern part of the Danville quadrangle, but is very ill defined in that portion of its course. From the Wabash Valley eastward also the ridges are feebly developed.

3. The Bloomington morainic system. This is the most prominent morainic system in the Illinois lobe. It usually becomes separated into two ridges, thus making a system of four well-defined ridges. The city of Bloomington stands on the outermost ridge of the system. The northern part of the Danville quadrangle contains the double ridge of the outer member. It is coalesced into single ridge in the western part of the quadrangle, but in the eastern portion the double ridge is well displayed, the valley of Stony Creek being between the two crests. A few miles north of the Danville quadrangle the double ridge of the inner member is well displayed, the village of Hooperston and the headwater portion of North Vermilion River being between the two crests.

4. The Marseilles morainic system. This consists of a bulky ridge, which only in places assumes a double or multiple crest. It is crossed by Illinois River at the village of Marseilles, from which the name is derived. It encircles the Morris Basin, a low-lying tract at the head of Illinois River.

5. The Kankakee morainic system and associated boulder belts. The ridges and boulder belts of this system lie just outside the Valparaiso system from Kankakee northward, but diverge rather widely from it in passing east from Kankakee, their course being south of east, while the Valparaiso bears north of east into Indiana.

6. The Valparaiso morainic system. This system received its name from Valparaiso, Indiana, which stands on the crest of its main member. It is a very bulky system, forming the elevated tract which lies east of Fox River and north of the Kankakee in northeastern Illinois and northwestern Indiana.

7. The Lake Border morainic system. As the name implies, this system of morainic ridges sweeps around the head of Lake Michigan.

The intervals between the morainic systems of this series were apparently much briefer than those between the drift sheets above outlined. There may have been some readvance <sup>Time intervals between moraines.</sup> of the ice in connection with the development of each morainic system, but as no decisive evidence of weathering occurs at the level of the base of the later moraines it is doubtful if the oscillations of the ice front were of great consequence. There is perhaps an exception in the case of the Kankakee ridges and boulder belts.

They seem to have been formed by an ice lobe whose axis of movement was decidedly out of harmony with that of the Marseilles and Bloomington systems. The concealment of the western part of the Champaign morainic system beneath the Bloomington also suggests a shifting of the axis of movement and an interval of greater length than that which separated the Champaign from the Shelbyville. It should not be inferred, therefore, that the several systems had a strictly rhythmic development. On the contrary, the halts seem to have been irregular.

The moraines or marginal ridges outlined above constitute only a small part of the Wisconsin drift of the Illinois glacial lobe and are given special attention because of their significance in an interpretation of the glacial history of the

Danville quadrangle. It is probable that a sheet of drift was deposited over much of the territory between a given moraine and the Lake Michigan Basin at the time that the moraine was forming, for the ice carried only a meager portion of its drift to its extreme margin.

In the Illinois lobe the moraines are but the thickened margins of sheets of drift which become gradually thinner in passing toward the Lake Michigan Basin. This is well shown by an examination of the Shelbyville drift sheet. At the main ridge it has a thickness of about 100 feet, but upon passing northward to the border of the Champaign system it decreases to less than half that amount. In the Champaign system another ridged margin appears, north of which there is a tapering sheet combined with the Shelbyville sheet. These pass beneath the Bloomington morainic ridges and are supplemented north of these ridges by the tapering Bloomington drift sheet. In all probability, therefore, the plain in Iroquois County north of the Bloomington morainic system carries deposits made by the ice during the formation of the Shelbyville and Champaign morainic systems as well as the Bloomington, while the plain south of the Bloomington system, a part of which falls within the Danville quadrangle, carries deposits made during the Shelbyville and Champaign substages of glaciation.

#### DESCRIPTIVE GEOLOGY.

##### ROCKS OF THE CARBONIFEROUS PERIOD.

*Relation of quadrangle to the coal field.*—All of the rocks appearing at the surface within the limits of this quadrangle belong to the Coal Measures, hence it is important to consider the relation which this area sustains to the coal field as a whole. The coal field, of which the Danville quadrangle forms only a small part, occupies portions of the States of Illinois, Indiana, and Kentucky. The territory embraced within it is shown by the line pattern in fig. 1, and the Danville

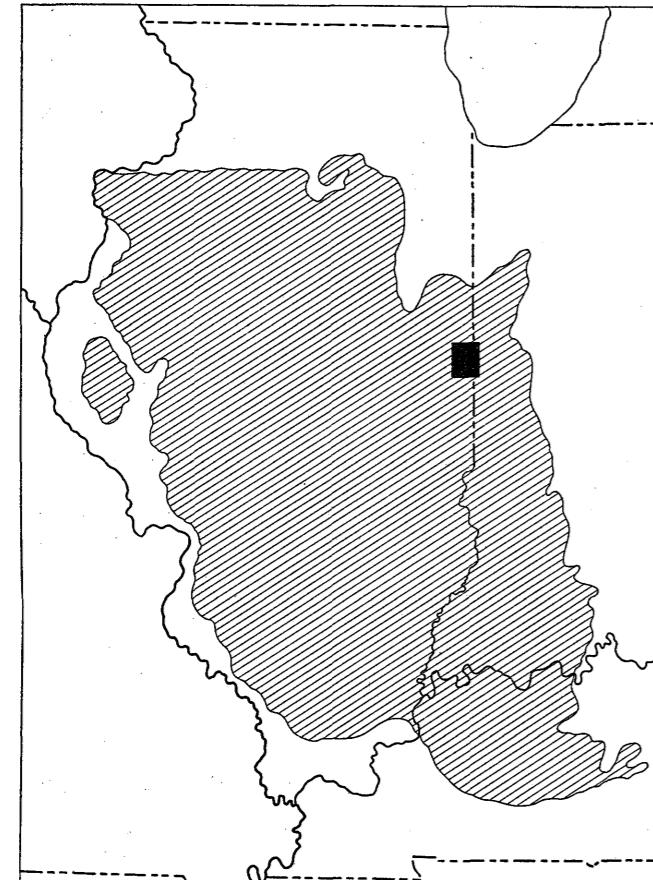


FIG. 1.—OUTLINE MAP SHOWING THE RELATION OF THE DANVILLE QUADRANGLE TO THE COAL FIELD.  
Coal field is represented by the shaded area.

quadrangle by the rectangle of solid black. It is thus seen that the quadrangle lies upon the northeastern border of the field, in a sort of projecting angle that extends into DuPage County, Indiana.

This coal field consists of a broad, shallow trough or synclinal depression, in which the coal bearing rocks occupy the central part, while the older, underlying rocks out-

crop around the margin. If the deposition of the rocks in and around this coal field had been continuous and free from disturbances of any kind, there would have been a gradual succession, from the center outward, of rocks lower and lower in the geologic scale. On all sides except the northern side this condition prevails—the Carboniferous rocks of the basin are succeeded by Devonian, and these in turn by rocks of Silurian age. On

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the level of the water, the erosion of this surface down to a nearly level plain, its subsidence below the level of the water, and, finally, the deposition of coal-bearing rocks upon it at the bottom of the sea, or of fresh-water lagoons. The southward extent of this overlap beneath the coal basin has not been determined, but, as will be described later, it extends, with little diminution in amount, at least as far south as the Danville quadrangle.

*Structure of the coal field.*—Although the coal bearing rocks lie in a trough-shaped depression in the older strata, the width of the trough is so great, when compared with the depth, that the rocks appear to lie horizontal. It is only by the comparison of sections in widely separated localities that the true dip of the rocks and the structure of the basin are apparent. In a small portion of the territory, like the Danville quadrangle, dips are generally so slight as to be unrecognizable, and that which appears to be dip in a coal bed is frequently not due to structure, but to change in thickness of some of the adjacent beds.

The two sections A-B and C-D on the Columnar Section sheet show some of the irregularities and also the dips of the strata in this region. The sections are constructed mainly from drill records, and hence are generally accurate as far as the coal beds are concerned.

The correlation of the coal outcropping at "Hanging Rock" on Vermilion River and on Coal Branch with the coal bed penetrated by the drill in several of the deep wells is a matter of some doubt, but the probabilities seem to be in its favor, hence it will be adopted provisionally. On this supposition the steepest dip in the quadrangle is between Coal Branch and the head of Trosper Branch, in a nearly westerly direction. Beyond the limit of this steep dip the rocks are nearly horizontal, the apparent dips being generally caused by the variations in thickness between certain well-marked strata.

*Character and thickness of the rocks.*—The strata exposed at the surface in the Danville quadrangle are prevailingly argillaceous or shaly. Occasional beds of sandstone and impure limestone are seen, but they are usually lenticular in form, and change in character from place to place, so that they can not be identified in different parts of the field. The geologic column of rocks, as interpreted by Prof. F. H. Bradley, was published in Volume IV of the Geological Survey of Illinois. This section is fairly accurate for the strata exposed along Vermilion River, but on account of the variability in thickness and character of the strata it does not agree with sections obtained in other parts of the field. The variation in thickness of the interval between the Danville and Grape Creek coal beds in the southern part of the quadrangle is well understood by all who have had occasion to search for these coal beds, and it seems highly probable that similar changes occur in other localities and affect different parts of the geologic column.

On account of the above-mentioned facts, it is impossible to give an exact general columnar section of the coal-bearing rocks which underlie the Danville quadrangle. The best sections of the strata from the Danville coal down to the bed showing on Coal Branch are probably those obtained by the diamond drill in the vicinity of Danville. Fig. 5 represents a bore hole made by the Ellsworth Coal Company 2 miles west of the city of Danville. The log of this well was published in Volume VII of the Geological Survey of Illinois, and from it the section is compiled. Fig. 4 represents the section reported by Mr. John Beatty from a well drilled in 1886 by the Diamond Prospecting Company in sec. 21, T. 19 N., R. 11 W., about  $\frac{1}{2}$  miles southeast of Danville. Fig. 3 is the section obtained in a deep drilling by the Himrod Coal Company at the Pawnee mine in 1894-95, as reported by Mr. H. H. Keefer, superintendent.

There is obvious similarity in the principal features of these sections, so that they may be correlated with considerable certainty, but the minor members vary so much in detail that it is difficult, if not impossible, to find any resemblances between them. The Danville and Grape Creek coal beds are easily identified in all of these sections. Below the latter there is a thickness of from 105 to 125 feet of

shale, more or less sandy, but in all cases without coal or carbonaceous matter. Following this heavy shale is a bed of black shale 7 or 8 feet in thickness, which is very pronounced in all of the sections. From 75 to 85 feet below the black shale occurs a prominent coal horizon, but the coal itself is so badly broken up with shale partitions as to be of little value. This coal bed is regarded as the equivalent of the one showing on Coal Branch and at "Hanging Rock" on Vermilion River. The interval between the black shale and the coal bed near the base of the section is variable in the extreme. In figs. 4 and 5 it is characterized by two small coal beds, but the Pawnee section (fig. 3) shows no indication of coal in this interval. It seems probable that this horizon may yield coal at many points, but the amount will probably be small and not valuable commercially. The lower coal bed is likewise uncertain and variable, and on the whole is of doubtful value.

The sections obtained by the drill show very little solid rock above the Danville coal bed, but in other localities shale to the thickness of 80 feet is exposed above this horizon. This is particularly well shown in the river bluffs about the junction of Salt and Middle forks of Vermilion River.

The information concerning the rocks below the limits of the detailed sections is much more meager, for it rests entirely upon two wells drilled about 1886 in the vicinity of Danville. These wells were drilled in search of oil, but as none was found, there has been no further prospecting in this field. Fig. 6 represents the well drilled at the city waterworks. Its section, recorded by the owners, is as follows:

##### Section of well at city waterworks, Danville.

	Thickness of stratum, in feet.	Depth, in feet.
1. Soil	10	10
2. Soapstone	285	295
3. Coarse sandstone	10	305
4. Soapstone	10	315
5. Sandstone	100	415
6. Soapstone	15	430
7. Gray sandstone	10	440
8. Blue sandy shale	80	520
9. Quartz, or pebble rock	10	530
10. Sandy shale	145	675
11. Hard gray limestone	30	705
12. Sandstone	30	735
13. Blue clay shale	30	765
14. Pebble, or flint rock	30	795
15. Hard blue shale	90	885
16. Gray sandstone	40	925
17. Hard blue shale	45	970
18. Light-green shale	30	1000
19. Black slate	75	1075
20. Limestone	74	1149

The limestone at the bottom of the well is labeled "Trenton" in the record preserved by the owners, hence it is probable that they believed they had reached the oil-bearing stratum and had found it barren.

Fig. 7 represents the section obtained in the well at Danville Junction. It was drilled for the same purpose and by the same company as the well just described. The record, furnished by the owners, is as follows:

##### Section of well at Danville Junction.

	Thickness of stratum, in feet.	Depth, in feet.
1. Glacial drift	175	175
2. Hard black slate and coal	6	181
3. Drab soapstone	20	201
4. Dark-blue soapstone	42	243
5. Coarse white sandstone	10	253
6. Coal	6	259
7. Blue clay, or soapstone	75	334
8. Hard, flinty rock	2	336
9. Dark-blue slate	35	371
10. Brown soapstone	20	391
11. Red clay	11	402
12. Soft white sandstone	68	470
13. Red clay	20	490
14. Coarse brown sandstone	27	517
15. Fine brown sandstone	40	537
16. Fine white sandstone	30	587
17. Dark-blue clay	73	660
18. Hard pebble rock	10	670
19. Fine white clay	36	706
20. Hard pebble rock	6	712
21. Dark-blue shale	96	808
22. Soft, light-blue shale	65	873
23. Soft, dark-blue shale	18	891
24. Red shale	62	953
25. Light-green shale	57	1010
26. Hard gray limestone	25	1085
27. Black slate	90	1125
28. Hard gray limestone	51	1176
29. Coarse, soft limestone	10	1186
30. White and dark limestone	160	1346
31. Soft white limestone	12	1358
32. Light and dark limestone	342	1700
33. White sandstone	35	1735
34. Clay shale	110	1845

25. Hard gray limestone	26	1871
26. Dark-blue limestone	65	1986
27. Hard blue shale	57	1998
28. Reddish limestone	15	2008

Fortunately for one who wishes to make comparisons, the driller noted a coal bed in this well at a depth of 259 feet. This is undoubtedly the same coal horizon that shows near the bottom of each of the detailed sections, and it serves to connect this section with the local features of the stratigraphy. At a depth of 1035 feet the bed of black shale was encountered which shows in the first well, and at 1125 feet the drill entered the limestone which was struck in the bottom of the first well and pronounced Trenton. This bed was pierced and it was found to have a thickness of 575 feet. Below this were alternating shales and limestones to the bottom of the well.

The interpretation of these sections is not possible without the assistance of sections from neighboring localities where the succession is unbroken and clearly understood. The nearest well section is that of the Lodi salt well. Fig. 8 shows this section as reported by Mr. E. T. Cox in the First Annual Report of the Geological Survey of Indiana, 1869. This section shows the same variable Carboniferous series as the Danville wells, but there is still no trace of the Carboniferous limestone. The same black shale was found at a depth of nearly 1000 feet, with the same limestone underlying it that was found previously, but the section is not complete and hence affords no standard for comparison.

A type section is obtained a little farther east, at Rockville, Parke County, Indiana. Fig. 9 shows the section at this point. To a depth of 350 feet the rocks consist of surface <sup>A typical section.</sup>

drift and Coal Measures. Below that, in typical development, is the Carboniferous limestone, with a thickness of 175 feet. Below the limestone there are about 600 feet of shales and sandstones belonging to the Waverly or Knobstone formation. This is followed in regular succession by the Devonian black shale, which is one of the best horizon markers in the oil field of Ohio and Indiana. Below this are the Niagara and associated limestones, having an aggregate thickness of 370 feet. The regular sequence continues, and the next shale bed, having a thickness of 470 feet, is undoubtedly the Hudson and Utica shales. Lastly the Trenton limestone appears, near the bottom of the well.

It is plainly apparent that the black shale which is present in all of the deep well sections here given is of Devonian age. With this fact established, it is also apparent that the underlying limestone is not Trenton, but that it belongs to that group of which, in this region, the Niagara limestone is the most prominent member. The shales and limestones lying below this bed belong to the Hudson-Utica group, but it is difficult to determine whether the reddish limestone noted at the base of the section is Trenton or not. A well at Champaign, 30 miles west of Danville, struck the undoubted Trenton at a depth of 663 feet. The alternating shale and limestone beds above this resemble strongly the beds of similar character in the well at Danville; therefore it seems probable that the last-mentioned well stopped a little short of the heavy Trenton limestone.

The rocks above the Devonian black shale contain no individual beds that are identifiable, hence the correlation of the strata in the various sections is extremely difficult to make. The Carboniferous limestone is well developed in the section at Rockville, but it is absent from the sections farther west. This has been accounted for on the supposition that it changed in character in that direction from a clearly marked bed of limestone to sandstone and shale. This explanation would, of course, account for its absence, but it is not in accord with other well-known facts. The Carboniferous limestone is well developed all around the coal basin except on the northern side, where the rocks of this horizon were removed long ago by erosion. If, then, this limestone changes to shales and sandstones in western Indiana, it is a local change, for the bed is of normal composition on the farther side of the coal basin. The unconformity by erosion and overlap in the northern part of the coal field has not been considered in this connection. In that

part of the basin the Coal Measures rest upon all beds from the Carboniferous limestone to the St. Peters sandstone, and presumably this condition extends underneath the basin itself. This change is in a westerly direction, and the disappearance of the limestone in western Indiana is in the same direction, consequently the supposition that both are due to the same cause is more probable than that which attributes the disappearance of the limestone to a change in character by which it remains unrecognized. The cutting out of this bed is accomplished between Rockville and Lodi, Indiana.

The disappearance of this limestone leaves the Coal Measures in contact with the Waverly or Knobstone formation, and the rocks of each are so similar in composition that, <sup>Divisions of the Coal Measures.</sup> in the absence of coal beds, it is almost impossible to separate them. In the recent survey of the coal field of Indiana, Mr. Ashley has arranged the Coal Measures into divisions which are separated by the horizons of well-known coal beds. Division 1 includes all of the rocks from the base of the series to the horizon of coal No. 2. There is considerable sandy material in this division and it is regarded as equivalent to the Millstone grit. In the well sections here represented it is difficult to locate this division, but it probably corresponds with a generally sandy formation 180 feet in thickness which was struck in the well at Danville Junction (fig. 7) at a depth of 400 feet. Connecting the sections, lines of possible correlation have been drawn from the base of this sandstone to the top of the Carboniferous limestone in the Rockville section. Its location in these sections does not agree with the determinations made by the Indiana geologists, but there is so much uncertainty on this point, even amongst those intimately acquainted with the field, that it is allowable to draw provisional lines of correlation as indicated.

Divisions 2, 3, 4, and 5 are grouped in this part of the field. They are represented in the section at Danville Junction by the interval between the sandstone and the coal bed 260 feet below the surface. This bed of coal is equivalent to the lowest coal in the detailed sections, figs. 3, 4, and 5.

Division 6 extends from this coal horizon to that of the Grape Creek coal, or coal No. 6, as it is generally known in this vicinity. The character of the rocks and the coals in this division has been discussed at some length, hence it is not necessary further to describe this division.

Division 7 includes all of the rocks lying between the Grape Creek and the Danville coal beds. This also has been fully described on a previous page.

Division 8 includes most of the rocks above the Danville coal bed in this region.

These divisions made in the Indiana field depend upon the identification of certain coal beds or coal horizons. Under the most favorable conditions this is a difficult problem, <sup>Difficulties of correlation.</sup> but in the coal field of which the Danville quadrangle is a part the conditions are far from the best for the identification of coals. In the first place there is no parallelism between the coal-bearing strata and those upon which they rest. The sediments forming the Coal Measures were laid down upon a more or less uneven floor, and on that irregular surface coal swamps developed wherever conditions permitted, but presumably the swamps were relatively small and of local occurrence. Under such conditions of local development of the lower coal deposits and of the lack of harmony between the coal-bearing series and that upon which they rest, the identification of the various coal beds which occur normally low in the series becomes most troublesome, if indeed it is not impossible.

If the coal beds were usually accompanied by roof shales bearing fossil plants, as they are in some localities, it might be possible to correlate widely separated outcrops of coal, or at least to determine to what general part of the series certain beds belong. The series, however, is generally barren of fossil plants and consequently the latter can not be depended upon for correlation purposes. The roof shales of the Grape Creek coal bed are an exception to the general rule, for they contain an abundant and distinctive flora. Mr. David White has examined small collections

from the Himrod and Kellyville No. 2 mines and his report upon them may be found on a subsequent page.

The Illinois part of the coal field is so deeply covered with glacial drift that it is generally impossible to trace the strata along their outcrops, hence most of the information regarding them must be obtained from prospect borings and mine shafts. In the limited territory embraced in the Danville quadrangle this is a comparatively simple matter, for the productive territory has been thoroughly prospected, but in other parts of the field prospect holes and mines are so far apart that the principal beds in their sections can be correlated only in an arbitrary manner.

From the consideration of all of these facts it seems almost impossible to extend into Illinois the system of classification used in the adjoining State, or in fact to use any system which depends upon the identification of the coal beds over a wide extent of territory.

In the work previously done in Illinois there was no attempt to subdivide the coal-bearing strata into formations, but the whole was regarded as one series—the Coal Measures. In this series the various beds of coal were designated by numbers, which began at the lowest coal and extended upward. According to this system the Grape Creek bed is known as No. 6 and the Danville bed as No. 7. Such a system is very simple and easily applied if the coal beds are regular throughout the basin or have distinctive features whereby they may be recognized in distant parts of the field. The objections against the extended application of this system are fully as strong as those against the Indiana system, for with the existence well established of an unconformity by erosion and overlap at the base of the series it is almost certain that there is irregularity and variability both in the thickness and in the geographic distribution of the coal beds, and hence the system will fail in its practical application.

The writer is not familiar enough with the entire coal basin to suggest a plan of geologic representation that will meet all the requirements, but from the data at present available it seems inadvisable to attempt the subdivision of the coal-bearing series into formations. The general plan of the United States Geological Survey is to divide strata into lithologic units to which geographic names are assigned. In this field the coal-bearing rocks have the appearance of being a lithologic unit, hence they will be considered as belonging to one formation, but the limited knowledge of that formation will not permit the assignment of a definite name at the present time. The coal-bearing rocks, therefore, will be considered simply as Coal Measures until a wider knowledge of the limits and character of the formation is attained.

*Fossil plants.*\*—The small collection of fossil plants enumerated below was obtained from the roof of the Grape Creek coal at the Pawnee shaft and at Kellyville mine No. 2. The matrix is a gray argillaceous shale, often breaking in conchooidal fracture. The very distinct and beautiful carbonized plant remains are often somewhat oblique to the bedding, thus indicating a relatively rapid process of silting. Many of the fragments bear evidence of driftage.

The species present in the collection made by Mr. Campbell are as follows:

- Pseudopeccopteris, indeterminable fragment.
- Peccopteris dentata Brongn.
- Peccopteris Strigosa Lx.?
- Peccopteris (Scolecopteris) ovatifolia n. sp.
- Alethopteris Gibsoni Lx.
- Callipteridium Grandini (Brongn.) Lx.
- Neuropteris rarinerbis Bunby.
- Neuropteris ovata Hoffm.
- Neuropteris plicata [Lx.]
- Linopteris obliqua (Bunby.) Pot.
- Odontopteris longifolia Lx.?
- Megaphyton insignis (Lx.) Gr'ey.
- Megaphyton Grand'Euryi Lx.
- Asterophyllites equisetiformis (Schloth.) Brongn.
- Annularia ramosa Weiss.
- Annularia sphenocephaloides (Zenk.) Gut.
- Sphenophyllum emarginatum Brongn.
- Sphenophyllum oblongifolium (Germ.) Ung.
- Macrostachya ferrilis (Lx.) D. W.
- Lepidodendron Tijou Lx. (?)
- Lepidophyllum hastatum Lx.
- Lepidophyllum missouriense D. W.
- Lepidocystis vesicularis Lx.
- Sigillaria campotœnia Wood.
- Trigonocarpum grande Lx.

\* By David White.

This florula is at once recognizable as belonging to the Productive Coal Measures. But two species, *Annularia ramosa* and *Lepidocystis vesicularis*, have been found in the uppermost portion of the Pottsville formation, and they are not characteristic of that horizon. On the other hand, the observed stratigraphic occurrence of several of the species strongly points to a position well up in the Coal Measures groups.

As compared with the floras of the older formations of the Appalachian Upper Carboniferous, the Grape Creek plants are conspicuously younger than those of the lower division of the Kanawha formation, while the presence of such species as *Peccopteris ovatifolia*, *Callipteridium Grandini*, *Alethopteris Gibsoni*, *Neuropteris plicata*, *Sphenophyllum oblongifolium*, and *Lepidophyllum hastatum*, clearly refer the flora to a stage higher than the upper portion of the Kanawha formation, or the Clarion group of the Allegheny series. In fact, an examination of the vertical range and of the association of the species in the Allegheny series, whose floral succession is better understood, shows that the flora in hand can, at the earliest, hardly antedate the Freeport group, the uppermost division of that series. Certain peculiar phases exhibited by several of the species are corroborative of a reference to this or even to a still higher group. Such are the very large and obtuse-pinnuled form of *Peccopteris dentata*, the dilated pinnules of *Neuropteris ovata*, the elongate *Odontopteris*, and the large form of *Annularia sphenocephaloides*.

Although the collection in hand is small, we may safely conclude that the horizon of the Grape Creek flora is not lower than the Freeport group on the one hand, while <sup>Correlation by fossil plants.</sup> granting that it may, on the other hand, eventually be found to lie at a somewhat higher stage in the as yet paleobotanically unknown Conemaugh series. Nothing more definite in the way of paleontological correlation of the flora is feasible until standard paleontological sections of the Conemaugh and Monongahela series for comparison shall have been made. As tending to oppose so high a reference as the Pittsburg coal, it may, however, be noted that the somewhat scant material in hand from near that horizon seems to present a more advanced development than that of our flora, while apparently lacking very many of the common Allegheny types, such as *Neuropteris rarinerbis*, *Sigillaria campotœnia*, and the large *Trigonocarpus*.

In the Interior or trans-Mississippi basin little is known of the floras above the Cherokee division of the Des Moines series. However, from such data as have been obtained it appears probable that the closest affinities of the flora under consideration will be found in the lower portion of the Missourian, or possibly in the uppermost Des Moines.

The composition of the Grape Creek flora indicates a stage in the lower Stephanian of the Old World. The latter division of the European Coal Measures appears, in the present stage of our knowledge of the fossil floras, to correspond to the Monongahela and Conemaugh series, together, perhaps, with the Freeport group of the Allegheny series of the eastern United States, and to the Missourian, with perhaps the upper portion of the Des Moines, of the Interior basin.

#### PRE-PLEISTOCENE CONDITIONS.

There was probably sufficient breaking down and erosion of the surface of the Carboniferous rocks before the Glacial period to give this region an appearance similar to <sup>A type of the pre-glacial conditions.</sup> that of the neighboring unglaciated portions of the Coal Measures in Indiana, the kinds of rocks, the general altitude, and the relation of the drainage lines being very similar in the two districts. The unglaciated district referred to presents low, irregular hills and ridges that rise usually 100 to 150 feet above the bordering valleys. The bluffs present very gradual slopes and the valley bottoms are broad. Streams which drain only a few square miles flow in valleys with bottoms one-fourth mile or more in width, while the large streams occupy valleys several miles wide. The Ohio on the border of southwestern Indiana has a valley 6 to 8 miles wide, with a rock floor about 75 feet below the present stream and bluffs that rise scarcely 100 feet above the

water. The valley of the Wabash in its lower course is even broader than that of the Ohio above the junction of these streams, and the bluffs also are low, with gradual or receding slopes. Similar features are found on the lower courses of all the tributaries of the Wabash in southwestern Indiana and southeastern Illinois which chance to fall within a district so thinly covered with drift that the pre-Glacial valleys and ridges are open to inspection.

The Danville quadrangle, being but a short distance back from the valley of a large pre-Glacial stream (the Wabash Valley), there should be concealed beneath its coating <sup>A drift-covered old surface.</sup> of drift a series of broad valleys connecting with the Wabash Valley, separated by low uplands rising perhaps 150 to 200 feet above the rock floor of the valley bottoms. The glacial deposits have so great thickness within the quadrangle that the region has been filled to a level higher than the tops of the old divides, so that the present surface offers no suggestion as to the position of either valleys or ridges, which can be determined only by the height of the outcropping ledges of rock along the river bluffs and by such wells as have been sunk deep enough to reach bed rock.

In a general way the rock surface is very regular in the upland south of the river. There is a variation of a little more than 30 feet, and the slopes are gentle, showing that <sup>Topography of the old surface.</sup> the topography of that time was well matured. The highest areas are situated on the south and west sides of the quadrangle. On the south side the rock floor has an altitude of nearly 630 feet above sea level. This altitude holds northward along the line of the Georgetown and Danville pike to within a mile of Vandercook. From this ridge the rock floor slopes in either direction, sharply toward the river on the east, and gently toward Catlin and the Wabash Railroad on the west. On the western edge of the quadrangle, near the junction of Salt and Middle forks, the surface of the rock is approximately at the same elevation as along the Georgetown and Danville pike. Between these two slightly elevated ridges there is a low, flat valley, which was probably once occupied by a small tributary branch of the ancient Vermilion River.

The narrowness and the rocky character of the valley of Vermilion River clearly indicate that the present course of the stream is of recent date and that it does not necessarily correspond with any drainage line which existed in pre-Pleistocene time. It is not known positively that any drilled wells have yet reached the bottom of the valley in which flowed the old Vermilion River, but its position is indicated by the great depth of drift found in the well at Danville Junction. The altitude of the rock floor at this point is 435 feet above the level of the sea. The well recently drilled at the Soldiers' Home reached solid rock at an altitude of 480 feet, and a well drilled a number of years ago near the depot at Grape Creek reached the rock floor at about 470 feet above sea level. These wells show a much lower rock surface than that which is found south of the river, and they are also slightly below the level of the rock floor northeast of Danville. Three wells in Newell Township reached the rock at an altitude of about 530 feet, one at 528 feet, one at 520 feet, one at 515 feet, and one at 480 feet. From these figures it will be seen that there is a rather deep valley in the vicinity of Danville Junction, and presumably this valley extends southeastward along the course of Stony Creek and Vermilion River to the mouth of Grape Creek. The sharp, rocky gorge which the river enters on leaving the great alluvial amphitheater at this point shows that the course of the old stream must have been different from that which the river occupies at present. The drill hole at Grape Creek station is located on the extreme western edge of the valley, and consequently it is doubtful whether it represents the full depth of the old channel in this vicinity. Unfortunately there has been no drilling in the center of this wide flat, and one can only surmise the depth to which the rocky strata have been eroded. The old valley turned at this point and presumably pursued an easterly direction beyond the margin of this quadrangle.

There is no evidence of reversal or any such decided change in the drainage of this region. It is simply a case of readjustment to new <sup>A readjusted drainage.</sup> conditions along practically the same lines that the streams occupied before the advent of the ice. During the time of maximum glaciation the streams were probably entirely arrested by the ice, but upon its retreat they formed along lines of least resistance, which in this case appear to have been nearly in their former courses. The ice front occupied a nearly east-west position, hence the southern parts of the valleys would be open first, and here the streams generally formed. When the ice had retreated to about the position of Danville the water from its melting front found a channel along the present course of the Vermilion, more than halfway up the rocky side of its old valley. Toward the south the high rock floor interposed a barrier to the pathway of the stream, but it pursued its course parallel to that highland and just at its foot. Doubtless the streams were at first located upon glacial drift only; as time advanced they eroded their channels and encountered the solid rock, but their pathway was selected and they persisted in it despite the solid rock through which they had to cut, even though only a short distance to the northeast there was a channel already cut through the solid rock to a greater depth than the one which they now occupy. The pre-Pleistocene river flowed east and probably joined the Wabash, and the modern stream has followed in nearly the same course.

North and west of Danville the old topography is not so easy to study, for the creeks have not yet cut down to solid rock except in <sup>An old valley of a tributary of the Vermilion.</sup> the immediate vicinity of Danville, and only a few wells have reached the original rock floor. There seems to be a general depression in this rocky floor along a line running directly northwest from Danville, for in sec. 16, T. 20 N., R. 12 W., solid rock was struck at an altitude of 470 feet, while in sec. 32, T. 21 N., R. 12 W., a well reached the rock at an altitude of 525 feet. It seems probable, therefore, that the western fork of the old river entered this territory near where the present Middle Fork enters, flowed southeast, and north of Danville joined another fork which occupied a slight depression along approximately the present line of North Fork.

This rather extensive drainage system in the northern part of the quadrangle is responsible for the lowness of the rocky floor in that region, and hence is indirectly responsible for the deep filling of drift that occurs there and for the absence of the coal beds which are so important in the territory south of the river. In the extreme northeastern part of the quadrangle the height of the rock floor is not known, since all of the deep wells in that section have failed to reach solid rock. It shows, however, along Wabash River in the vicinity of Covington at an altitude of about 500 feet, hence it seems probable that it is about the same in the vicinity of State Line.

The sudden termination toward the east of the productive coal, as shown on the economic map, is due to the erosion of the rocks and the <sup>Coal beds worn away.</sup> coal beds to a plane below the level of the latter. The outlines are only approximately correct, and hence do not represent the actual rugosities of the pre-Pleistocene topography. In Danville the coal beds are similarly cut away by early erosion, so that the beds which are prominent along the river bluff on the southwestern side of the city are not present in the northeastern part. In this locality the change is very abrupt and striking and its existence has been verified again and again by the many wells that have been sunk below the level of the coal beds without encountering solid rock.

#### PLEISTOCENE FEATURES.

**Illinoian drift.**—The Kansan and pre-Kansan drift sheets have not as yet been identified with certainty among the deposits of the Illinois glacial lobe. The oldest drift known to occur here is the Illinoian, and this extends beyond the Danville quadrangle about 150 miles, to the limits of glaciation in Illinois and Indiana. As already noted, it is overlain within this quadrangle by Wisconsin drift, which conceals it except in the valleys of the principal streams.

The Danville quadrangle does not show so clearly the succession of drift deposits as do the neighboring parts of Illinois. On the <sup>Buried soil rare here.</sup> north, east, and west of this quadrangle there are numerous instances of the occurrence of buried soils and fossiliferous silt, probably loess, between the Wisconsin and Illinoian drift sheets, but no exposures of these soils or silts were discovered within its limits, and there are very few instances reported of the penetration of buried soil in the wells. The fragmentary nature of the exhibit seems to indicate that the ice sheet at the Wisconsin stage removed much of the soil and silt. It also seems to have removed quite extensively the oxidized surface portion of the Illinoian drift sheet. This removal of the oxidized part of the lower sheet makes it somewhat difficult to determine the border lines between the Wisconsin and the Illinoian drift, either in the exposures along the streams or in the wells, for the two drifts present similar rock constituents and similar variations in color. But the Illinoian drift is to a marked degree indurated by cementation with lime, while the Wisconsin is usually soft and loose. The difference in induration seems to be, for this area, the best criterion for separating the drift sheets, and yet this will in places fail or yield unsatisfactory results.

In this as in other districts where both drifts are present, the Illinoian, being the older sheet, is largely found as a filling in the pre-Glacial valleys, and it has but a meager representation on the old uplands; while the Wisconsin drift, being deposited on a surface that had been rendered smooth, makes a more uniform sheet over uplands and valleys. Being of considerable depth throughout, it has the appearance of contributing the main part of the drift, for the old valleys lie largely below present stream level; yet it probably is but little more bulky than the Illinoian.

If we may judge from the neighboring districts, where the soils are well preserved between the Wisconsin and Illinoian drift, the thickness of the Wisconsin is generally 35 to <sup>Thickness of the drift sheets.</sup> 60 feet on the plains and 80 to 120 feet on the moraines; while the Illinoian sheet ranges from almost nothing on parts of the old uplands to nearly 300 feet in the deepest of the old valleys. In the immediate vicinity of the Danville quadrangle the thickest drift yet found is only about 250 feet, and this probably is more largely Illinoian than Wisconsin. An illustration of this is found in boring west of Eugene, Indiana, just outside the limits of the quadrangle, which failed to reach rock at a depth of 241 feet and is largely through a hard or partially cemented blue bouldered clay or till, thought to be Illinoian. Another boring of about the same depth east of Bismarck, Illinois, and just north of the limits of the quadrangle, passed through a bluish-black muck at about 60 feet, which probably marks a soil horizon at the top of the Illinoian drift. This muck is also found at about 90 feet beneath the moraine in the northeastern part of the quadrangle.

Throughout the quadrangle the Illinoian drift is composed largely of stony clay or till, there being probably less than 10 per cent of <sup>Composition of the Illinoian drift.</sup> assorted material (silt, sandy gravel, and cobble). In the till there is a liberal admixture of small stones with the clay, but boulders are not numerous. Their scarcity may be understood from the fact that, out of the hundreds of deep wells that have been driven or bored into the till, only a very few have struck boulders so large that they could not be pushed aside by the drill or auger. The exposures along streams also show only an occasional boulder.

An inspection of the till reveals the fact that the boulders are very largely crystalline rocks of Canadian derivation, though limestones <sup>Sources of the till.</sup> from the Silurian formations of northern Illinois and Wisconsin are also found. The smaller stones embrace large numbers of fragments of limestone as well as of crystalline rocks. The clayey matrix seems to be derived largely from the shales of the Coal Measures, though it contains also pulverized limestone and many small fragments of crystalline rocks. The gravel and cobble deposits which are interbedded with the till contain about the same proportion of limestone and crystalline rocks as is found in the stones of smaller size in the till, as is to be expected,

since the gravel and cobble are derived from the same source as the till and differ from it chiefly in lacking the finer ingredients. The sand and silt beds, like the clayey matrix of the till, contain considerable pulverized limestone, as well as fragments of shale and of crystalline rocks.

The till shows marked variation in the color of its clayey matrix, though the predominating color, in both the Illinoian and the Wisconsin <sup>Colors of the sheet, till.</sup> drift, is blue-gray. The deposit showing the most conspicuous variation from the prevailing tint, and having a color which does not seem to be due to oxidation, is a pink till, which is frequently found below the blue-gray till or at the same horizon with it. The color is very similar to that of a Coal Measures shale which outcrops at many places along Salt Fork and the lower courses of Middle and North forks, and there is little doubt that the till owes this color to the presence of the shale in its matrix. There are several exposures along Hungry Hollow, northwest of Danville, where the pink shale and the pink till may be seen in close association. The most extensive exposure of this pink till noted is on a west tributary of North Fork, in sec. 7, T. 20 N., R. 11 W. For nearly one-eighth of a mile it forms the base of the bluff to a height of 10 or 15 feet above the bed of the ravine. Above it

there are about 30 feet of blue-gray till, separated from the pink till by a definite horizontal band of cobbles and boulders. As this pink till is commonly found either at the base of the drift or at considerable depth, it probably belongs to the earlier or Illinoian drift. But since the Illinoian also includes a large amount of blue-gray till, and as the pink till is found to pass horizontally into blue-gray till in many of the exposures, this color distinction should not be depended upon to separate the Illinoian from the Wisconsin drift.

Occasionally a nearly black clay is found at or near the base of the drift, which, like the pink till, seems to be derived from Coal Measures shale. An excellent exposure of <sup>A black clay.</sup> such a clay is found at a road grading on the southern bluff of Hungry Hollow, about a mile above its mouth; and a disintegrated black shale bearing a striking resemblance to this clay may be seen, *in situ*, in the same bluff only a few rods distant. The black clay embedded in the drift has several feet of brown till and sandy gravel beneath it, separating it from the underlying rock. There are lumps or masses of it several feet long and a few inches thick embedded at different levels up to a height of 15 feet or more above the rock. In this connection it may be remarked that there is an exposure on the northeastern bluff of Vermilion River, a short distance southeast of the southeast corner of the quadrangle, showing a mass of Coal Measures sandstone several feet thick and several square rods in extent embedded in the till. When first observed it was mistaken for rock *in situ*, but upon closer examination it was found to be underlain by the blue-gray till of the drift. These instances serve to show how largely the color and the constitution of the drift may be affected by the incorporation of local rock material into it.

Turning now to deposits which are associated with the till, it is found that they include beds of assorted material of various degrees of coarseness, from cobble down to fine silt or clay, and that they occur at all depths in the till. If these beds of associated material were at the surface their extent might be accurately determined, but since they are buried beneath heavy deposits of till their extent has not been determined, even approximately. It is known only that the wells and the banks or bluffs of streams show but a small amount of assorted material compared with the amount of till. The general presence of large supplies of water in the sand and gravel struck by wells in the midst of the drift seems to indicate that these deposits may have frequent connection with one another beneath the surface and that they are also connected with porous surface deposits which take in the water. In some parts of the glaciated area there are extensive sheets of water-bearing sand and gravel at definite horizons, but within the Danville quadrangle no definite horizons of water-bearing beds have been discovered. The most conspicuous one noted is in the vicinity of Catlin, and this occupies an area of only about a square mile.

**Iowan silt or loess.**—The loess or silt which covers the Sangamon soil outside the limits of the Wisconsin drift is traceable throughout the district bordering the Mississippi River. It is directly connected with the Iowan till of northwestern Illinois as well as that of eastern Iowa, and appears to be a dependency of the Iowan drift sheet. It varies greatly in thickness, if wide areas are considered, but is systematic <sup>Thickness of the Iowan silt.</sup> in that it shows a tendency to become thick along the great rivers and thin on the

<sup>Beds of sand or gravel.</sup>

<sup>Silt beds.</sup>

<sup>A weathered zone between drift and loess.</sup>

divides. On the borders of the Illinois and Mississippi rivers there are belts having a thickness of 30 to 50 feet, and on the Missouri River a much greater thickness is reached. But on the borders of the Wabash Valley it averages but 10 or 15 feet. On the divides between the streams it often becomes reduced to less than 5 feet. It is probable that the Danville quadrangle, being near the Wabash, had nearly the average thickness found along that valley, or about 10 feet. The deposit was therefore much less in amount than the Illinoian drift, and so thin that the advancing ice sheet might easily have incorporated nearly all of it in its own drift.

Mechanical analyses of specimens of this silt from the portion of Illinois immediately outside the border of the Wisconsin drift sheet have been made under the direction of Prof. Milton Whitney, of the Department of Agriculture. He finds that the grains seldom exceed a diameter of one-tenth of a millimeter, or the grade called fine sand, and most of the grains fall below one-twentieth of a millimeter. Where grains exceed one-tenth of a millimeter they are usually concretions of iron oxide. More than half of the deposit is made up of grains falling between one-twentieth and one-hundredth of a millimeter. The following table sets forth the results of three of these analyses.

*Mechanical analyses of the Iowan silt in eastern Illinois.*

Diameter, in millimeters.	Conventional name, etc.	Giant:		
		1-18 inches.	Near Greenup:	Moreau:
		Percent.	Percent.	Percent.
2-1	Fine gravel.....	0.00	0.30	0.00
1-.5	Coarse sand.....	0.00	1.05	0.08
.5-.25	Medium sand.....	0.02	3.42	0.77
.25-.1	Fine sand.....	0.30	3.30	0.11
.1-.05	Very fine sand.....	5.21	6.47	4.88
.05-.01	Silt.....	57.75	55.48	52.50
.01-.005	Fine silt.....	12.78	11.70	12.15
.005-.0001	Clay.....	20.36	14.90	22.10
	Total mineral matter.....	96.42	96.62	92.59
	Organic matter, water, loss.....	3.58	3.38	6.61
	Loss by direct ignition.....	6.01	3.11	5.73

Chemical analyses of the loess have been made by Prof. W. A. Noyes from two samples collected near Terre Haute, Indiana, only a few miles south of the limits of the Wisconsin drift, and these probably illustrate well its constitution near the Wabash River. The results of these analyses are given below. The first sample is from a point 10 inches and the second from a point 22 inches below the surface. The analyses were made for the Indiana Geological Survey and first appeared in its Twenty-first Annual Report.

*Analyses of Iowan silt from near Terre Haute, Indiana.*

Constituent.	No. 1.	No. 2.
	Percent.	Percent.
SiO <sub>2</sub> .....	79.77	72.87
Al <sub>2</sub> O <sub>3</sub> .....	9.95	11.25
Fe <sub>2</sub> O <sub>3</sub> .....	3.89	6.75
Ti O <sub>2</sub> .....	.70	.95
Ca O.....	.67	.69
MgO.....	.26	1.06
Na <sub>2</sub> O.....	1.08	.39
K <sub>2</sub> O.....	2.05	2.26
H <sub>2</sub> O.....	2.55	4.24
Total.....	100.42	100.46

*Peorian soil.*—The Peorian soil as displayed in the type locality near Peoria, Illinois, seems to mark a less prolonged interval than the Sangamon. The humus stain is meager and the leaching of the surface part of the loess is far less than that of the leaching of the Illinoian drift in the Sangamon inter-glacial stage. Whether the black earth or clay found in a few wells in the Danville quadrangle belongs to the Peorian or to the Sangamon can scarcely be decided. It is found, however, that in the neighboring districts where the records are more complete and the number of instances much greater, the best-defined buried soil is at the Sangamon horizon, the soil being directly overlain by till. This fact renders it probable that in the Danville quadrangle the Sangamon rather than the Peorian soil has been noted in well records.

*Wisconsin drift.*—It is the Wisconsin drift

that gives the Danville quadrangle its topographic expression as well as its agricultural conditions. By referring to the topographic map it will be seen that the portion south of the latitude of Danville is largely a plain, there being only a few small ridges and knolls dotting its surface. The topography of the portion north of Danville, on the other hand, is in large part undulating. In this region are the prominent ridges which form the outer or southern portion of the Bloomington morainic system. Two distinct ridges appear in the northeastern part of the quadrangle, which are separated by the narrow sag or plain through which Stony Creek flows. On the west side of North Vermilion River these unite in a single broad ridge. The margin of the ice appears to have held its position in this western part of the quadrangle, while it withdrew in the eastern part from the south to the north ridge.

The strong morainic belt and its bordering nearly featureless plains here displayed are illustrative of classes of drift topography that are widely prevalent from Illinois eastward through Indiana and Ohio. The plains between the moraines are usually dotted with as few drift knobs as occur on the plains in this quadrangle, while the moraines often have as great relief and as pronounced topographic expression as here exhibited. The most conspicuous exception is found in the Champaign morainic system, and the weakest portion of one of its members chances to fall within the Danville quadrangle. The inner and weakest member or ridge of this Champaign system is traceable from Champaign eastward into Vermilion County nearly to the western border of the Danville quadrangle, passing south of the villages of Sidney, Homer, and Fairmount and constituting the water parting between Salt Fork of Vermilion and the north tributaries of Little Vermilion River. Near Fairmount it breaks up into a range of knobs, one of the most prominent being "Blue Mound" in sec. 11, T. 18 N., R. 13 W. These knobs become so scattering and so low upon passing eastward into the Danville quadrangle that it would be misleading to call them a moraine. Yet they serve to connect the definite ridge that leads up nearly to the western border of the quadrangle with an equally definite one that sets in just south of the quadrangle, in sec. 27, T. 18 N., R. 11 W., and they constitute a faint representation of the position of the ice margin at a certain stage. As only a small number of these knobs exceed 10 feet in height, they are imperfectly represented on the topographic sheet. Enough of them are shown, however, to indicate that the belt runs eastward from the western border of the quadrangle to the vicinity of Kellyville and thence southeastward to the well-defined ridge that sets in just outside the quadrangle, in sec. 27, T. 18 N., R. 11 W. They are scattered over a strip 2 or 3 miles wide. The most prominent knob is "Sandusky Hill," immediately south of Catlin; it is 35 or 40 feet in height. Aside from this there are few that rise 20 feet above the plain.

There are a few knobs within the Danville quadrangle that are not included in either the Champaign or the Bloomington morainic system. Three such knobs are to be seen near Hillery, about 4 miles west of Danville; another is found about one-half mile southeast of Rileyburg, Indiana, and several occur between Vermilion River and the State line, 3 to 5 miles south of Rileyburg. These knobs may be classed as kames, since they include a large amount of sand and gravel with the Kames.

In this connection it may be remarked that some of the knobs near Catlin, which are included in the Champaign morainic system, contain a large amount of gravel and sand and may be properly classed as kames. This is the case with the most prominent one, "Sandusky Hill." The kames here, as elsewhere, may be found either on moraines or on till plains.

Upon examining the Bloomington morainic belt

more in detail it will be found that it presents an irregular, somewhat wavy surface, with oscillations of 10 to 20 feet or more in nearly every 40 acres of the surface. The forest-covered portion presents sharper knobs than the prairie portion, a feature that is prob-

ably due to forest protection from frost and fire as well as from the action of small streams. There may, however, have been an original difference in the sharpness of the knobs. There is a marked relief on the southern or outer border of the moraine. It will be noted that the plain immediately south stands about 650 feet above tide, while the moraine, within a mile north from the border of the plain, rises, as a rule, slightly above the 700-foot contour, and there attains about as high an altitude as at points farther back from the plain. There is a similar relief in the ridges on each side of the sag through which Stony Creek has its course. But the relief is not so conspicuous at the northern border of the morainic belt, there being in places a gradual merging into the plain, and nowhere is there so marked a difference in altitude between the moraine and the plain as on the southern border, the plain at the north being about 30 feet higher than the one on the south.

The Wisconsin drift of this quadrangle, like the Illinoian, is composed largely of stony clay (till), both in the plain and in the morainic or undulating portion. The most striking difference between it and the underlying Illinoian, as already indicated, is the degree of induration or cementation, the Wisconsin being as a rule much softer and easier to penetrate with auger or spade than the Illinoian. The till of the moraines seems to be a trifle more stony than that of the plains, a difference that seems attributable to the removal of a larger proportion of the silt from the morainic drift.

Surface boulders abound on both of the ridges of the Bloomington morainic belt for 2 or 3 miles east from North Vermilion River, but with this exception the surface of the portion of this morainic belt within the quadrangle is not very bouldery. Boulders also abound in a small area on the southwestern side of Vermilion River just below Danville, but are not numerous on other parts of the plain or non-morainic portion of the quadrangle. The surface boulders are almost wholly crystalline rocks of Canadian derivation, of which granite far outnumbers all other classes, but within the till there are occasional large blocks of limestone, which were probably derived from neighboring ledges in northeastern Illinois, only 50 to 75 miles distant. The smaller stones of this drift sheet, like those of the Illinoian, embrace large numbers of rocks of comparatively near derivation, as well as granites and other crystalline rocks of distant derivation.

Aside from the occasional kames or knobs of gravel and sand, the Wisconsin drift of this quadrangle carries numerous small pockets or local accumulations of gravel and sand in the midst of the till, which are filled with water and supply the shallow or dug wells. The amount of these deposits is, however, much less than that of the till. In a well 30 feet in depth there is ordinarily but a few inches of sand and gravel, the excavation being almost wholly till. The movement of glacial water in connection with the deposition of this portion of the Wisconsin sheet and with the melting away of the ice sheet seems to have been remarkably feeble. This inference is sustained by the nature of the outwash from the Bloomington morainic system, as appears below.

In this drift sheet oxidation has changed the color of the surface portion from blue-gray, the common tint, to yellow or brown. The depth to which this change extends is usually but 10 or 15 feet on the morainic ridges, and somewhat less on the prairies. The surface leaching has been sufficient to remove the calcareous rock flour to a depth scarcely half as great as the depth of oxidation, but it has not removed the limestone pebbles to any appreciable extent. In this respect the surface here is in contrast with the leached surface portion of the Illinoian drift, where, as above indicated, pebbles as well as rock flour have been leached away to a depth of several feet. This feature strongly supports the view that the time since the disappearance of the Wisconsin ice sheet from this region is brief than the time that elapsed between the withdrawal of the Illinoian ice sheet, with attendant exposure of the Illinoian drift to atmospheric

action, and the advent of the Iowan glaciation, when the soil and leached surface which had formed on the Illinoian drift became covered with the Iowan silt or loess.

*Surface silt.*—On the plain immediately south of the Bloomington morainic belt the till is capped by a deposit of silt or nearly pebbleless clay having a depth of several feet, which within a distance of 2 or 3 miles south of the border of the moraine decreases to but a foot or two, and in the southern part of the quadrangle is scarcely sufficient to conceal the surface boulders. The ridges of the Bloomington morainic belt within this quadrangle and the plain north of them carry scarcely any silt on their surfaces. This silt deposit has been referred by Professor Bradley to the loess,\* but it evidently is much younger than the main loess deposits, for the latter are older than the Wisconsin drift. It is probably of complex origin. The thickest portion being situated on the outer border of the Bloomington morainic belt, it is natural to infer that it is an outwash from the ice sheet during the formation of the moraine. The validity of this inference is not, however, beyond question. The thin portion can scarcely be accounted for in this way, since it lies higher than the heavy portion, and upon ground that seems likely to have been above the level of glacial waters that were flowing toward the main avenue of discharge, the Wabash River. That portion may, however, have been deposited as an outwash from the ice in its retreat across the southern part of the quadrangle before drainage lines had been opened, and hence somewhat earlier than the Bloomington morainic belt. A supplementary if not alternative agency is found in the wind, and this may still be adding to the thickness of the deposit, especially in very flat places, where erosion is of little effect.

*Wisconsin valley-gravel.*—Throughout much of the region covered by Wisconsin drift the valleys leading away from the moraines carry deposits of gravel and sand which were laid down by the dirt-laden streams that issued from the ice. The degree of coarseness of the material varies greatly from place to place, and is thought to indicate the strength of the stream current. The strength of the current would evidently vary with the volume of the stream and the rate of fall or gradient of its bed. On portions of the ice margin where the valleys had not been opened to receive the glacial drainage directly, the water would spread out over the plain and carry sediment to form an outwash apron until some valley was reached through which its waters could find a discharge.

In the Danville quadrangle the Vermilion Valley afforded a line of escape to the Wabash for tributaries that headed at the ice margin, and certain beds of gravel and sand along it appear to have been deposited directly by these glacial streams. By referring to the geologic map it may be seen that at the outer border of the Bloomington morainic belt the gravel fills the interval between Vermilion River and Stony Creek, making a belt nearly 2 miles wide, which gradually narrows upon passing down the valley. The discharge seems to have been more vigorous here than along the remainder of the ice margin lying within the quadrangle. There was so little gravel discharged from Middle Fork that it has been largely removed by the later action of the stream. Possibly, as already indicated, the silt which lies along the outer border of the moraine between Middle and North forks, and also east of Danville, is largely the outwash from the ice. The flatness of the surface is such that vigorous drainage could scarcely be expected except at points where valleys connected directly with the ice margin.

It is evident that gravel deposition continued in North Fork and Middle Fork valleys down to the time when the morainic belt which crosses their headwaters was forming, for the gravel remains in conspicuous terraces up to the points where these streams emerge from this later morainic belt. Instances may be seen near the north edge of the Danville quadrangle. It seems not improbable that the streams flowing from this later moraine contributed also to the gravel

\* Geology of Illinois, Vol. IV, 1870, p. 242.

deposits found south of the outer belt of the Bloomington system, though perhaps only a minor part.

There is evidence that the deposition began while the ice stood at the outer belt. About 2 miles above Danville, in sec. 30, T. 20

N., R. 11 W., there are gravelly deposits along the western side of North Fork which seem to extend from the plain surface of the terrace back into the undulating surface of the moraine, after the habit of moraine-headed terraces, as if that had been the starting point of a glacial stream. There is also a large admixture of cobble and even larger stones with the gravel at the place, which suggests that the source of supply (the ice sheet) was close at hand. The remnants of the gravel filling farther up the valley are at a slightly lower altitude and are composed of finer material than that at this place.

The material of this terrace consists of a fine gravel which in places becomes sandy. The gravel is seldom cemented, though it contains a large amount of calcareous material in both coarse and fine particles. In a few places it reaches a depth of 30 feet, but is usually only 10 or 15 feet in depth. It commonly rests upon an eroded till surface, but occasionally extends to the Coal Measures. There is generally a bed of fine sand or silt covering this gravel to a depth of 1 to 4 feet, which seems referable to sluggish waters in the closing stages of deposition. Such a silt capping is a common feature on the Wisconsin valley-gravels.

*Drainage development.*—An inspection of the topographic map will make it evident that the drainage lines are in a youthful stage. There are perhaps 75 of the square-mile sections, or about one-third the number included in the quadrangle, in which no drainage lines are represented, and nearly every section in the quadrangle contains extensive uneroded tracts. In most cases the bluffs are abrupt and the original level of the drift surface is preserved almost to the brow of the bluff. It is estimated that an area of not more than 30 square miles of the 228 included in the quadrangle has been materially affected by erosion since the ice left it, and yet this quadrangle includes the lower course of an important tributary of the Wabash, where the streams have the advantage of rapid fall. There has been a very insignificant amount of erosion in the headwater portion of the same drainage system. The southern part of Cattin Township, in the southwestern part of the quadrangle, represents the practically unmodified condition which wide areas of the drift surface display in neighboring parts of Illinois and Indiana. This immature condition of the drainage system is, however, characteristic only of the region within the limits of the Wisconsin drift, the drainage systems on the Illinoian drift being far better developed.

Respecting the courses selected by the streams, it will be observed that Vermilion River has taken a somewhat direct course across the quadrangle, a remarkably level drift plain from Danville down to Wabash River. The courses of North Fork and Middle Fork in their passage through the Bloomington morainic belt were probably determined by sags or channels formed in connection with the deposition of the drift and the melting away of the ice. The meandering of the valley of Middle Fork seems to indicate that the stream was obliged to find a course among the knolls and ridges of the moraine, and is a feature which does not favor the view that a well-defined channel had been opened by subglacial waters. The comparatively small amount of glacial outwash at its point of emergence from the moraine leads to the same conclusion. North Fork, on the other hand, crosses the morainic belt in a somewhat direct course, such as a subglacial stream would naturally develop. The large amount of glacial outwash at the point where North Fork leaves the moraine may also sustain the interpretation of subglacial drainage. The westward course of Stony Creek through the northeastern part of the quadrangle was determined by the sag between the two morainic ridges. Its selection of the course southward through the outer moraine instead of westward between the moraines into North Fork seems to indicate that the ice left a suitable channel for it across the

moraine. Possibly some of the material in the gravel plain which heads at the moraine north of Danville was brought in through this channel. The courses of the small tributaries appear to have been governed by some detail of slope or by slight depressions in the drift surface, except perhaps a few very small streams which have worked back across a level plain from the valleys into which they discharge.

As the valley gravel along the Vermilion rests upon a terrace, there must have been stream excavation prior to the gravel deposition.

The depth of the excavation was nearly 50 feet in the vicinity of Danville, and seems to have been still greater toward the mouth of the river, while the width exceeds that of the modern valley, as may be seen by reference to the geologic map. Similar excavation is also found in valleys farther west that lead away from the Bloomington morainic belt, notably on tributaries of the Illinois River. In these valleys, as in the Vermilion, there are trains of valley gravel which head in the Bloomington morainic system. In each of these valleys the amount of erosion seems greater than streams having the present drainage area would have produced in the interval between the Champaign and Bloomington substages of glaciation, an interval so brief that aside from these valleys one can scarcely detect a difference in the stage of drainage development upon passing from the Champaign to the Bloomington drift sheet. The manner in which the excavation occurred is not yet clear. Possibly a stream adequate to effect this erosion was in operation between the Champaign and Bloomington substages in connection with the melting ice sheet. The water which flowed from the ice sheet may have dropped its burden of sediment before reaching the lower course of the Vermilion and taken on the power to erode in that part of the valley. The volume of water may thus have been far greater than that of the modern stream, and the work may have been accomplished in comparatively short time. The question is one, however, on which further light is needed.

It will be observed that the main drainage ways of the Danville quadrangle present marked irregularity in the size and form of their valleys, there being in places wide, low bottoms extending from bluff to bluff; in other places noticeable terrace remnants, with narrow, low bottoms; and in still other places only slight terraces and narrow, low bottoms. These irregularities appear to be due in part to variations in the texture of the formations in which the valleys have been excavated, the broad, low bottoms being found where there has been excavation in soft or easily eroded material, either drift or shale, and the narrow parts being largely confined to places where the bluffs are of somewhat resistant rocks.

The wide expanse of the Vermilion Valley opposite Grape Creek evidently marks the place where a pre-Glacial drainage line is entered; but the expansions on Salt Fork and on Middle Fork, near the west border of the quadrangle, seem to be due to weakness in the shale which forms a large part of the bluffs in this locality. The narrowness of Vermilion Valley below the Grape Creek expansion, and the narrowness of the trench cut in the terrace in the vicinity of Danville, are each attributable to the resistance offered by rock ledges; so also are the narrow gorges in the sharp bends of North Fork just above Danville.

The terraces are conspicuous chiefly where there is drift in the upper part and rock in the lower part of the bluffs, though they also occur where the bluffs are entirely of drift. The principal terrace on the main streams is capped with gravel, and it seems to bear a definite relation to the ice sheet at a time when it still stood at the Bloomington morainic belt. The most conspicuous remnant of this main terrace within the quadrangle lies between Vermilion River and Stony Creek, from Danville down to the point where the creek enters the river, the course of the creek being near the eastern side of the terrace, while the river follows nearly the western border.

#### RECENT FEATURES.

*Alluvium.*—The deposits previously considered are somewhat closely connected with glaciation, but the alluvium, as here considered, is entirely the

product of modern streams, subsequent to the final disappearance of the ice sheet. It is distributed along the flood plains or low bottoms of the principal streams and merges into the talus of the valley slopes and of the headwater portions of the tributaries. The alluvium is really an apron of talus which has been spread out by the stream on its flood plain. Having this derivation, it presents a mixture of fragments of rock of all sizes, from coarse blocks which have fallen from cliffs or rolled down steep ravines to the finest silt or clay, deposited where the stream is most sluggish. Since the streams of this quadrangle have a comparatively rapid fall they have carried much of the fine sediment beyond the limits of the quadrangle, leaving in the river bottoms a loose deposit of great agricultural richness. Much fine walnut timber has been obtained from these alluvial bottoms, and they are now largely under cultivation.

*Soils.*—In a glaciated region such as the Danville quadrangle, considerable variation in the character of the soils is to be expected. A mixture of several kinds of rocks of all degrees of coarseness could scarcely develop a uniform soil, such as results from the disintegration of but one rock formation. Any one class of soil is therefore liable to show considerable variation. As to classes, the soils may be separated into (1) silt or clay soil, (2) sandy soil, (3) gravelly soil, and (4) boulder-clay soil. Silt or clay soil is found not only where the boulder clay is covered with a silt deposit, but also along the bottoms of some of the valleys. Sandy soil is generally confined to the valleys or their immediate borders. Gravelly soil appears on the slopes of some of the kames, as well as on portions of the terraces and bottoms of the valleys. Boulder-clay soil characterizes much of the Bloomington morainic belt and the portions of the bordering plains where surface silt is very thin or is wanting.

The kind of vegetation which has occupied a region also seems to have considerable effect on the soil. The soils in a prairie region, being comparatively unprotected from frost action and from sunlight, are made

loose from the heaving that results from freezing and thawing, while those in the forest, being protected by a coating of leaves and kept in the shade, not only suffer less heaving and loosening, but even become "puddled" or more compact as a result of this protection. It is thought that this quadrangle affords a striking illustration of the effect of vegetation on the soil. The prairie portion has a comparatively open soil, into which humus has penetrated to a depth of several inches, giving it a rich black color, while the forest-covered portion, even though formed apparently from the same kind of deposit, is much more compact and has so little humus in its soil that it presents at the top an ashy or pale-gray color and changes to brown or yellow at much less depth than the prairie soil.

In respect to fertility the entire quadrangle, with the exception of a few gravelly or sandy strips, possesses a productive soil. The upland silt and the valley alluvium, as well as the boulder clay, seem to be sufficiently varied in rock constituents to furnish all that is needed on the mineralogical or chemical side. But in the matter of porosity, which is a very important factor in determining fertility, the prairie portions of these several classes of soil have the advantage, and command a higher price. A map showing the distribution of the forest and prairie is therefore a better guide to the fertility of the soil than the geologic map, which shows the several classes of drift found at the surface. It should be stated, however, that, owing to the tendency to heave by freezing and thawing, the soil of the prairie district is not so well adapted to the growth of winter wheat as that which was occupied by the forest. Fruits of the orchard and the garden and vegetables have been grown with success on the prairie as well as on the forest-covered portions. The growth is generally more rank on the prairie, but in amount and quality the fruit is not superior to that of the wooded portion. Indeed, it is thought that the rankness of growth has a disadvantage in the case of trees, by rendering them more liable to winter killing. In general it may be said that, for crops which mature in a single

season, the prairie is more productive than the wooded tracts, but for those which must be carried through the winter the latter is preferable.

*Vegetation.*—This quadrangle is situated on the border between the forest-covered region of the eastern United States and the great prairie of the interior. This border is irregular because of strips of forest which extend along streams for long distances into the prairie region. The quadrangle embraces such a strip of forest that extends northwestward from the Wabash along the Vermilion and its main tributaries about to the west border of Vermilion County, Illinois. The great prairie lies on either side of this strip of forest and extends eastward nearly to the Wabash River, in Warren and Vermilion counties, Indiana.

The distribution of forest and prairie within the quadrangle, as determined by the Government Land Survey, is as follows: The prairie covers about 42 square miles in the southwestern portion, about 13 square miles in the northwestern, and about 45 square miles in the northeastern, amounting to about 100 square miles, and leaving about 128 square miles of forest within the quadrangle. The greater part of the forest originally consisted of oaks of various kinds, but largely of white oak. There were, however, strips of walnut, poplar, and elm along the river valleys and also on the borders of the prairie. The part of the Bloomington morainic belt east of North Fork was largely occupied by walnut, but the part west of North Fork carried a mixture of oak with walnut, poplar, elm, etc. Only a small part of the walnut and poplar now remains, as most of the trees have been cut and converted into lumber. Much of the oak has been cleared for firewood and also for lumber, and a large part of the original forest area is now under cultivation.

#### MINERAL RESOURCES.

##### COAL.

The principal mineral resource of this quadrangle is coal. Three or more beds are known to underlie this territory, but it is probable that only two of these are of sufficient thickness to be classed as workable. These beds are locally known as the Danville or No. 7 coal and the Grape Creek or No. 6 coal. The areas of best development of these coal beds do not coincide, and it is a popular saying in the vicinity of Danville that where one is of good thickness and quality the other is too thin to mine or too impure to put upon the market.

*Danville coal.*—The Danville coal lies from 20 to 80 feet above the Grape Creek bed and its probable extent is shown on the Economic Geology sheet. Owing to the deep cutting of Vermilion River in pre-Pleistocene time, this coal bed is confined principally to the southwestern half of the quadrangle. In the southern part of the quadrangle the eastern limit of the coal is along a somewhat irregular line extending from a short distance east of Georgetown to Danville. This limit was determined by pre-Pleistocene erosion. Since that time the ice sheet has extended down from the north across this territory several times, grinding down the highest points, filling the inequalities, and leaving the surface with little resemblance to its former appearance. At present there is no indication on the surface of the eastward extent of the Danville coal, but its presence has been determined by numerous test holes drilled in search of the lower coal.

From the point where the Danville coal bed shows in the river bluffs, a short distance below the Wabash Railroad bridge, its outcrop may be traced continuously up the Salt Fork to near the mouth of Butler Branch, where it dips slightly beneath the level of the river, but it reappears in the bottom of the river in sec. 17, T. 19 N., R. 12 W. From this point it continues just above water level to the margin of the quadrangle. It is entirely beneath water level on Middle Fork, but it has been reached in a number of places by mines and prospect holes as far north as sec. 31, T. 21 N., R. 12 W.

At Danville this coal bed is exposed all along the river front in the southwestern quarter of the city, but in passing toward the northeast it soon

disappears, having been cut out in the formation of the old pre-Glacial valley of Vermilion River. This coal was once worked by stripping near the waterworks and also at the mouth of Hungry Hollow, but so far as known it extends no farther north along this stream. Little information is available regarding the occurrence and condition of this coal bed in the deeply drift-covered region between North and Middle forks. Its approximate northeastward limit is shown on the map, but the exact position of the boundary line in secs. 15, 22, 23, 26, 35, and 36, T. 20 N., R. 12 W., has never been ascertained. Coal has been reported from a number of deep wells in Blount and Newell townships, and while these beds can not be connected with any degree of certainty, it seems probable that they belong to this horizon. They will be described later in the discussion of the thickness of the Danville bed.

The Danville coal is thin along the southern margin of this quadrangle, consequently it has not been well prospected. Below Georgetown it is reported to have a thickness of about 3 feet 6 inches. West of this point and near the station of Riola it is said to be only 2 feet in thickness. From the evidence available it seems to be not over 3 feet 8 inches in thickness as far north as Westville. At the Pawnee mine it has a thickness of 5 feet, but the coal is impure and of little value. Professor Bradley reports this coal bed as showing at Lafferty's bank, on Grape Creek, with a thickness of over 5 feet. The exact location of this bank is not known, but an old opening was observed near the public road in sec. 32, T. 19 N., R. 11 W., that is undoubtedly upon this coal bed and is probably the one referred to by Professor Bradley. From this point north to Danville it holds a thickness of about 6 feet, as shown by several test holes and by the old mine at Vandercook. It is now almost impossible to see this coal in the vicinity of Danville. It has been extensively stripped in this region and now its outcrop is marked by heaps and ridges of refuse, but the coal itself is not visible. This coal has been mined and prospected along Salt Fork from Danville to Butler Branch. Throughout this distance its thickness is about 6 feet. A short distance above Danville an opening was observed which measures 6 feet 2 inches. Farther up the stream a cliff exposure shows the following section:

*Section of Danville coal on Salt Fork.*

	Feet. Inches.
Coal	5 6
Bone	1
Coal, blacksmith	9
Total	6 4

The Danville coal has been opened at two points on Butler Branch, where it is reported to have a thickness of 6 feet in each case. On Salt Fork near the western edge of the quadrangle it is reported to have a thickness of from 5 feet 6 inches to 6 feet. On Salt Fork a mile or two above this territory there are extensive stripings where the coal runs about 6 feet in thickness.

In passing up Middle Fork this coal bed dips rapidly to about 80 or 90 feet below water level in sec. 5, T. 19 N., R. 12 W. A little west of this section and beyond the limit of the quadrangle there is a mine in which the coal is reported to average about 5 feet 6 inches in thickness. From this point it rises to about 50 feet below water level in the northeast quarter of sec. 29, T. 20 N., R. 12 W. At this point it is reported to have a total thickness of about 6 feet, with a thin parting of slate, about one-fourth of an inch thick, above the blacksmith coal, which ranges in thickness from 5 to 6 inches. In the next section north this coal is reported to have a thickness of 6 feet and to lie 90 feet below water level. In sec. 16, T. 20 N., R. 12 W., it is reported with the same thickness and about 20 feet lower than at the last-described place. In this region the coal is very irregular in altitude; it seems to lie in rolls or waves. Beyond the last-mentioned locality there has been some coal reported, but it is doubtful whether it is the Danville coal bed.

Coal has been struck northeast of Danville in several wells which have penetrated the rock to depths of 20 or 30 feet. None of the wells have gone below the coal bed, hence it is difficult to say to what horizon it belongs. It ranges in

altitude from 450 to 490 feet above sea level. Its position is therefore intermediate between the Danville coal as it shows in the city and the coal bed reported from the deep well (see fig. 7) at Danville Junction. If it is the last-mentioned coal, the strata rise toward the northeast, and if it is the former or Danville coal, they dip in that direction. South of the river the rocks and associated coal beds are almost free from undulations, but upon passing up Middle Fork the Danville coal dips rapidly, so that in sec. 5, T. 19 N., R. 12 W., it is almost 100 feet lower than at the mouth of Butler Branch. It seems probable that this synclinal depression extends in an east-west direction; if it does, it is probable that it has depressed the Danville coal to the level indicated northeast of Danville. Little is known regarding the thickness and extent of the coal in this region. In every case it has been reported from near the upper surface of the rock floor. Since the pre-Glacial rock surface is somewhat irregular, it seems probable that this coal bed has suffered extensive erosion and will be found only in isolated localities of restricted area. For this reason successful mining is doubtful; at least mining should not be attempted without thorough prospecting in advance.

*Grape Creek coal.*—This is the most important coal bed in the Danville quadrangle, but its best development is restricted to a rather small territory lying along the southwestern side of the old valley of Vermilion River. The eastern limit of this coal bed is sharply defined by the pre-Glacial erosion of the rock floor below the horizon of the coal. The present surface gives no indication of the limit of the coal, but this has been obtained with considerable certainty by developments in the ravines which cut deep enough to reveal solid rock, and by numerous prospect holes drilled in search of the coal. In the valley of Grape Creek and in general along Vermilion River from Grape Creek nearly to Danville this coal shows in outcrop in so many places that it may be said to show continuously. Beyond the point indicated the coal can be seen in a number of places, but it is generally thin and has received little attention. The area of workable coal can not be accurately delimited with the data at hand, but it is represented approximately by the heavy shading on the Economic Geology sheet.

Near Riola station this coal is reported to be 6 feet 4 inches thick. East of Georgetown it was found 84 feet below the Danville coal, and with a thickness of 5 feet 2 inches. In this quadrangle it holds an almost constant thickness of over 5 feet as far north as the Wabash Railroad. The reported measures in this territory range from 4 feet 9 inches to 9 feet. It is the prevailing opinion that the best coal lies in the eastern half of the territory described, and all but one of the large mines are located in that section. Many test borings in the western part of Georgetown Township and the southwest corner of Danville Township show a body of fine coal, generally over 6 feet thick, running with great regularity in thickness and composition throughout the area. The exact line upon which this coal thins below workable proportions is not known, but it passes south of Vandercook and north of Catlin. At the former locality the Grape Creek coal is reported to have shown a thickness of 2½ feet in the deep shaft which was sunk there several years ago. South of this place it thickens rapidly to nearly 9 feet in a distance of 1½ miles. At Catlin its average thickness is reported to be about 6 feet 6 inches. North of this point no section of this coal bed has been obtained until Butler Branch is reached, within 1 mile of Salt Fork, where it has been penetrated by the drill 22 feet below the Danville coal and 2 feet in thickness. Here again is a reduction in thickness of over 4 feet in about 1½ miles. The exact limit of workable coal can not now be determined, but the heavy shading on the Economic Geology sheet shows approximately the area underlain by coal of workable thickness.

The westward dip of the rocks carries this coal nearer and nearer the waters of Salt Fork west of Danville, until, near the eastern edge of sec. 12, T. 19 N., R. 12 W., it dips below the level of the stream, but it reappears in sec. 11 and extends for nearly half a mile at about the level of the water. Above this point it is not known at the

surface in this quadrangle. In thickness and section the coal is variable along this stream. About due north of Vandercook it was seen in one outcrop 40 inches in thickness, and in another it gives the following section:

<i>Section of Grape Creek coal on Salt Fork.</i>	
	Feet. Inches.
Dirty coal	1
Coal	1 8
Shale	2
Coal	2 5½
Total	5 3½

On North Fork this coal bed continues thin as far as is known. Its normal section is about 25 inches, as is shown in an outcrop directly west of the waterworks. Generally over the northern half of the quadrangle there is no information regarding the thickness and extent of the coal in this region. In every case it has been reported from near the upper surface of the rock floor. Since the pre-Glacial rock surface is somewhat irregular, it seems probable that this coal bed has suffered extensive erosion and will be found only in isolated localities of restricted area. For this reason successful mining is doubtful; at least mining should not be attempted without thorough prospecting in advance.

*Coal mines.*—There are several very extensive coal mines in this quadrangle. In the early days of the development of this field the Danville coal supplied the bulk of the output, but in recent years the production from this coal bed has diminished and the Grape Creek bed has decided pre-eminence in coal production in Vermilion County.

According to the report of the State Inspector of Mines for 1898, Mine No. 2 of the Kellyville Coal and Coke Company, located at Kellyville, had an output for that year of 248,872 tons, ranking second in the State in point of production; and Mine No. 3 of the same company, located at Westville, is credited with a production of 235,220 tons and is given fourth rank in the list of State mines. The Pawnee mine, located in sec. 5, T. 18 N., R. 11 W., and the Himrod mine, in sec. 9, same township and range, are the property of the Himrod Coal Company, which stood second only to the Kellyville Company in the coal production of Vermilion County in 1898. Other large operators on the Grape Creek coal bed are the Westville Coal Company, located 1 mile southwest of Westville; the Catlin Coal Company, operating at the town of the same name; and the Brookside Coal and Mining Company, on Grape Creek.

The largest mining plant in operation on the Danville coal bed is that of the Economy Coal and Mining Company, 2½ miles west of the city of Danville. On Salt Fork just above the western limit of this quadrangle is a large mine on the Danville coal, and there is also one near the west line of sec. 5, T. 19 N., R. 12 W. In sec. 29, T. 20 N., R. 12 W., a mine equipped with steam hoisting apparatus is operating on the same coal, but it has no railroad connection, hence its output is limited to the demands of the surrounding region.

In addition to the above-described mines, which have a large output and are equipped with steam hoisting apparatus, there are a great number of small mines on both coal beds, worked by hand or horse power. In the early development of this region most of the coal was gained by stripping the soil and rock from above its outcrop. In the vicinity of Danville almost the entire line of exposure of the coal has been thus extensively worked. The most available stripings have been exhausted, and now the amount of cover is generally too great to warrant its removal, but a great many small shafts and slopes have been opened for the purpose of supplying the local trade. Most of such mines are transient enterprises, working this year and abandoned next; hence their location upon the map is not of very great importance.

Several small mines have been operated on the lowermost coal of the series on Coal Branch, but the bed is so badly broken by partings of black slate as to have little importance, especially in the immediate vicinity of the fine body of coal of the Grape Creek bed.

CLAY.

The glacial clays of this quadrangle have been extensively used in the manufacture of drain tile. The level prairie lands of this region are very much improved by underground drainage, and there has been a strong demand for tile for this purpose. These clays, which abound almost

everywhere, have been found to make excellent porous tile, and plants for its manufacture have been established at a number of places.

At Danville a company is engaged in the manufacture of vitrified brick, the raw material for which is obtained from a bed of sandy shale that outcrops at the foot of the bluff on the point between Salt and North forks of Vermilion River.

GRAVEL.

The surface clay and soil of this region form very bad roads in wet weather unless they are macadamized with broken stone or gravel. The former can be obtained here only by shipment, hence the latter is the main dependence.

The gravel in the terraces of the Vermilion Valley in and below Danville is generally rather fine for road ballast, though several pits have been opened for procuring that material. Those now drawn from are mainly on the east side of Stony Creek near the mouth of Lick Creek. In some places the sandier parts of the deposits have been screened for plasterers' use. An extensive pit of this sort is opened just north of the mouth of Lick Creek.

A gravel suitable for road ballast has been found on the south side of the east-flowing portion of Vermilion River above Danville, one large pit being opened about a mile northwest of Vandercook and another opposite Batestown.

The only kame that has been opened extensively for gravel is "Sandusky Hill," south of Catlin. The material is a rather sandy gravel. Slight excavations in other kames have shown a material too sandy for road ballast.

UNDERGROUND WATER.

*Shallow dug or open wells.*—The most common wells within this quadrangle are the open or dug wells. They usually have a diameter of 3 to 5 feet and a depth of 10 to 35 feet. In many cases they strike a bed of water-bearing sand or gravel which has sufficient extent to afford an adequate supply of water for household or farm use in the driest seasons. But in parts of the quadrangle the wells find only weak veins of water, either in boulders clay or in sand or other assorted material, and consequently are not to be depended upon in the dry seasons. The wells which penetrate only boulders clay are naturally much weaker than those which strike beds of sand or gravel. Their supply comes from the absorption or collection of the ground water in the immediate vicinity of the excavation, while wells which strike sand or gravel are likely to have wide absorption areas.

Throughout much of the southern half of the quadrangle the shallow wells are sufficiently strong to meet the needs of the residents, and deeper wells are comparatively rare. But in the northern half, especially along the Bloomington moraine belt, they have been to a great extent supplanted by tubular wells 50 to 200 feet in depth. The subjoined table will illustrate these statements.

In the city of Danville there are two classes of shallow wells, each of which is somewhat different from the wells on the uplands. One class is located in the gravel terrace between Stony Creek and Vermilion River and penetrates sand and gravel from top to bottom, while the other class is located in Stony Creek Valley and penetrates till before entering sand and gravel. They are each 20 to 30 feet in depth.

The water in the shallow wells is ordinarily sufficiently free from contamination to be wholesome and safe for household use. It is also sufficiently free from sulphur, salt, and other objectionable matter to be of pleasant taste, but it is, as a rule, rather hard. These wells are liable to contamination when they are situated in barnyards, or when, as is often the case, the housewife empties the kitchen slops near them. In some cases the wells receive these slops without much filtering. Perhaps the greatest danger from contamination is found in the city of Danville. The wells which obtain water at the base of the gravel on the terrace are still largely used in the southeastern part of the city, but in the northern and western parts they have been supplanted, to a great degree, by the waterworks supply, which is pumped from the North Fork of Vermilion River. These abandoned wells have in many cases been converted into cesspools, and it seems probable that

these are drained through the part of the city where wells are still in use. Although there are strong grounds for suspecting contamination of the wells in use, it was not learned that any sickness had resulted from such contamination.

By many citizens of Danville the present waterworks supply is also considered unsafe, because Vermilion River becomes very low in dry seasons, so that the water must be pumped from nearly stagnant pools. This has led to an investigation of available supplies from wells, and attention has naturally turned to the wells in Stony Creek Valley, just noted, some of which are very strong. An objection urged to the use of these wells is the presence of a large cemetery within half a mile up the creek from the best wells, though on the opposite bluff. At the cemetery there are gravel deposits 10 or 15 feet in depth, from the base of which springs issue into Stony Creek Valley. But as a compact clay sets in above the level of the creek bed and extends some distance below it, separating the surface gravel from the water-bearing gravel in which these wells are found, the latter may not be contaminated by cemetery drainage. It is probable also that nearly all of the cemetery drainage escapes along the creek, so on that account, even if there were no clay cover for the lower gravel, there would be only a slight chance of contamination from that source. The chances certainly seem to be very much less than of the contamination of wells in the southern part of the city by the cesspools in the northern part. It is unfortunate that the cemetery site should have been located on the upstream side of the city, where it would suggest a source of contamination, for perhaps no other adequate supply from wells will be found within a convenient distance.

*Tubular wells.*—The supply from this class of wells depends scarcely at all upon the ground around the well mouth. They are usually so deep as to be independent of local or even of seasonal rains. Commonly their supply is derived from deposits of gravel or sand which are interbedded with the sheets of till, but in some cases they extend into the underlying Coal Measures before striking water. They are supposed to be fed from surface outcrops of the water-filled beds in localities not far distant, or, in the case of wells in rock, through joints or other openings in the overlying drift deposits. They usually show a marked rise above the level at which water is struck, and in few cases they overflow.

Since these wells are apparently supplied from water which has been absorbed in the region near that in which they are located, it might be supposed that they would be more seriously affected by seasonal variation in the rainfall. They seem, however, to be in a zone of saturation, so far beneath the surface that no droughts can reach them, unless greater aridity should become prevalent. The water may be in many cases several months, or even years, in passing from the surface of the ground to the beds in which the wells are found. This seems to be indicated by the amount of mineral matter which the waters carry in solution. In most of the wells the water, besides being hard, is somewhat chalybeate, slightly sulphurous, and slightly saline. It is often called "mineral water" to distinguish it from the water of the shallow wells, in which hardness is the only mineral property that is noticeable. An analysis of water from one of the tubular wells at Potomac, Illinois, a few miles north of the Danville quadrangle, made at the laboratory of the University of Illinois, shows that chlorine in the form of chlorides forms 29 parts per million of the water. This is pronounced by the analyst to be in no way due to contamination, and is, in all probability, taken up, like the lime and iron and sulphur, from the stony material of the water-bearing beds.

The water from one tubular well located within the Danville quadrangle has been shipped for medicinal purposes. The well, known as the "Henrietta Springs," is on a farm about one-half mile south of Snider post-office. Its depth is only 75 feet, and the water does not overflow. So far as known, no analysis has been made, but the water is strongly chalybeate. The present owner, J. Marbles, reports that no sale has been made for some time, as the well is inconveniently

located for shipping water. The well may probably be duplicated at many points in that region, and its notoriety is due to advertising the beneficial effects which certain invalids have thought resulted from the use of the water. The water from these tubular wells appears to be as wholesome as that from the shallow wells, and the occurrence of this strong supply of water in parts of the quadrangle where shallow wells are difficult to obtain is a great boon to the residents.

Dr. Arthur W. Palmer, Professor of Chemistry at the University of Illinois, has made several hundred analyses of water from such wells as appear in the Danville quadrangle, and reports as follows concerning the water of shallow and tubular wells.\*

In appearance and in palatability the two classes of water present marked differences. The waters from shallow wells are well aerated, and are clear, sparkling, cool, and of agreeable taste; those from the deeper wells, on the other hand, contain little or no oxygen, possess in many cases a disagreeable taste due to the presence of marsh gas, accompanied occasionally by minute quantities of sulphuretted hydrogen, and are either turbid or become turbid quickly on exposure to air, owing to the oxidation of the iron carbonate which they contain and the consequent precipitation of insoluble ferric compounds. The precipitating particles are often so minute as to be at first indistinguishable except from the color which they impart to the water, and after a short exposure to the air the water becomes opalescent, then decidedly turbid; finally a brown deposit similar to iron rust is produced, and after this has separated the water becomes clear and colorless. Although these unpleasant characteristics of the deep drift waters give rise to much prejudice and objection to their general use for drink, nevertheless, from the sanitary standpoint, they are usually to be preferred to the clear and palatable waters of the shallow wells, since the evidence of numerous analyses shows that they are far less subject to pollution with refuse animal matter than are the latter.

The tubular wells which obtain a supply from the Coal Measures are comparatively few and the water is generally not so hard as in the wells obtained from the glacial drift. The strongest wells of this class of which information was obtained are situated in secs. 21, 22, and 23, T. 20 N., R. 11 W., about 3 miles northeast of Danville,

in Stony Creek Valley at Danville, which, as indicated by the table below, shows no sign of exhaustion when pumped at the rate of 750 gallons a minute.

*Flowing wells.*—Flowing wells from the drift are obtained in large numbers in neighboring parts of eastern Illinois, but only a few have been obtained within the limits of this quadrangle. The village of Myersville, situated in the North Vermilion Valley, about a mile north of the north boundary, has several wells 145 to 156 feet in depth, which overflow, in some cases at a rate of 12 to 15 gallons per minute from a pipe with orifice only  $1\frac{1}{4}$  inches in diameter. At the village of Potomac, a few miles from the northwest corner of the quadrangle, there are about 100 flowing wells, some of which will flow 60 gallons a minute, and it is estimated that 100 more have been obtained on farms in the vicinity of that village. There are also flowing wells in southern Vermilion County, Illinois, only a few miles south of the quadrangle.

The only flowing wells from the drift within the limits of the quadrangle, so far as learned, are a few near its northern edge, in the valley of North Vermilion River, and some very shallow fountains in ravines on the farm of W. P. Smoot (sec. 23, T. 20 N., R. 12 W.), in the midst of the Bloomington morainic belt. A flowing well in the valley of Middle Vermilion River, in sec. 5, T. 20 N., R. 12 W., is scarcely one-fourth of a mile west of the border of the quadrangle. Only one of these wells exceeds 100 feet in depth, that of Allen Wood, in North Vermilion Valley. Three others are about 50 feet, and the remainder but 15 to 35 feet, in depth. Those which are but 15 to 35 feet in depth may be fed from the immediate vicinity of the wells; but the deeper ones, like the tubular wells above described, seem to be fed from more remote sources. Indeed, wells of this class apparently differ from the ordinary tubular

In parts of the drift-covered portion of eastern Illinois the variations in head may be readily explained by nearness to or remoteness from the gathering ground. Such is the case in the great field of flowing drift wells found in Iroquois and Ford counties, the largest field of this kind in the State of Illinois. The water-bearing beds slope with the surface of the ground, away from the moraine which constitutes the chief area of intake, and their head becomes gradually less upon receding from the moraine. This explanation may apply to a portion of the Danville quadrangle, especially that lying south of Salt Fork and west of Big Vermilion River. It is probable that that area is fed by waters which are absorbed in the somewhat elevated Champaign morainic belt in the southern part of Vermilion County, and that the course of underground drainage is northward, in harmony with the surface slope. Some loss of head would result through friction in the underground passage. But it is thought that the loss of head may be due in the main to the drainage into the main valleys, Salt Fork and Big Vermilion River, for both of these streams cut entirely through the drift deposits.

In the portion of the quadrangle east of the Big Vermilion River, the decrease in the head, though in the opposite direction from that west of the river, is like it in being toward the main drainage lines and suggests that it is due to the discharge of water into the Big Vermilion and Wabash valleys, both of which have been cut through the drift into the underlying rock.

In the northern part of the quadrangle there seems to be a slight drawing down of the head in the vicinity of North and Middle Vermilion rivers, and this, as in the southern half of the quadrangle, may be due to the escape of the underground water into these valleys. The apparent influence of these valleys in depressing the head is shown in the accompanying sketch, fig. 2, which crosses a portion

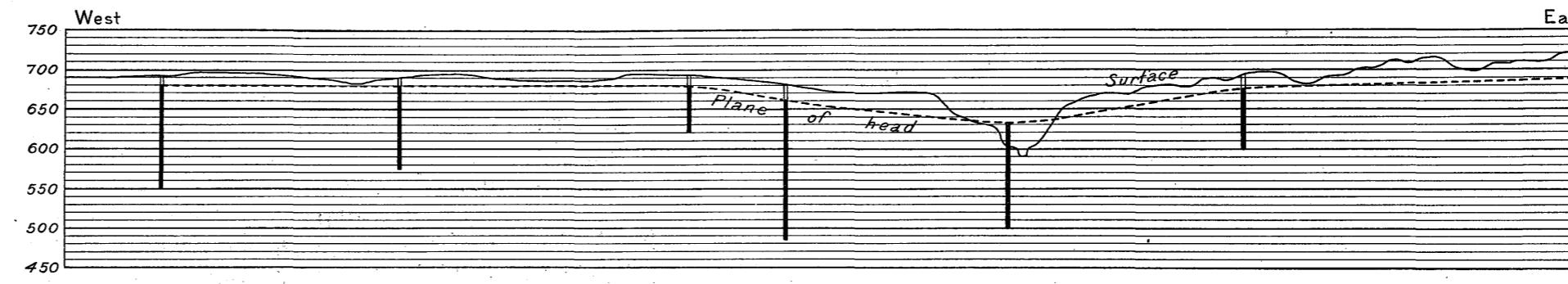


FIG. 2.—SECTION ACROSS A PORTION OF THE QUADRANGLE NEAR THE NORTH BORDER.  
The heavy vertical lines indicate the heights to which water rises in wells, and illustrate the depression of the plane of head in the vicinity of Vermilion River.  
Figures at the left refer to elevation (in feet) above sea. Horizontal scale same as map.

but weak ones are distributed over all parts of the quadrangle, as may be seen by reference to the subjoined table. The Coal Measures ordinarily contain, in their upper portions, rather compact shale instead of porous sandstone, and are therefore not a promising source of water supply.

A great majority of the tubular wells within this quadrangle are located on the Bloomington morainic belt, nearly every section on that belt now carrying one or more of them. There are also a few such wells south of the Bloomington belt, in Danville and its vicinity, and in the eastern part of the quadrangle, and a few in the southwestern part. Statistics concerning a large number of these wells are given in the table which follows.

Many of the tubular wells are reported to be very strong, but their capacity has seldom been measured. Some wells are decidedly weak, so that the owners are obliged to use care in preventing the windmills by which their water is usually drawn from causing a waste of the supply. Their weakness differs from that of the shallow wells in being constant instead of being confined to seasons of drought. While most of the wells are adequate to supply the needs of a small stock farm, it is doubtful if many can be found whose capacity would be sufficient to supply the city of Danville, or even such an institution as the Soldiers' Home, located near there. At the time work was begun on the Soldiers' Home several wells were sunk on the grounds, but, though they were pronounced strong by the well driller and would supply the needs of a farmer, they were entirely inadequate for the demands of the Home. They are far weaker than one of the shallow wells

wells simply in being located at a sufficiently low altitude to permit a flow. In most cases the head is no greater than that of wells on higher ground which do not overflow. The question of head, therefore, pertains to both classes of wells, and may be next considered.

By reference to the subjoined table it will be seen that the head shown by the wells presents considerable variation, though there are areas several square miles in size in which it is very uniform. In the northern part of the quadrangle, and also for several miles farther north, the water usually rises to about 680 feet above tide; but in the valleys of Middle and North Vermilion rivers it is markedly lower. In the Bloomington morainic belt there is much variability, the range in the head being from about 690 feet down to 580 feet or less. As in the plain north of the morainic belt, the head is generally lower near Middle and North Vermilion rivers than at points some distance back from these streams. But there are instances of very low head some miles back, in the midst of districts where wells show much greater head. On the nearly level tract south of this morainic belt the head in the part west of Big Vermilion River shows a general increase from north to south, being but 635 to 640 feet just south of the morainic belt, while it is 665 to 670 feet in the southern part of the quadrangle. A few miles farther south, on the slope of the main ridge of the Champaign morainic system, it reaches 680 feet. The head in the portion of the plain east of Big Vermilion River shows a decrease instead of an increase in passing from north to south, being about 635 feet near the southern border of the Bloomington morainic belt and scarcely 620 feet in the southeastern part of the quadrangle.

of the quadrangle from east to west near its northern border. The absorbing area for the northern part of the quadrangle probably extends from the Bloomington morainic belt northward, for, owing to the depth of Salt Fork Valley, the water absorbed in the southern part of the quadrangle can hardly be supposed to cross this valley into the northern part. Possibly this northern portion of the quadrangle lies also in the line of underground drainage from the somewhat elevated morainic belts in the northern part of Vermilion County. This suggestion does not, however, imply that the underground water beds have a southward dip or slope, but simply that, the nearest lines of escape for underground waters being southward, the accumulated waters may work toward the outlets, the hydrostatic pressure being sufficient to carry them over the inequalities found in the line of escape. The great strength of the flowing wells in the vicinity of Potomac, Illinois, at the southern border of one of these morainic belts, seems to support the view that the chief absorbing area is on the north rather than in the more distant morainic belt on the south. The gradual decrease in head in passing southward from Potomac down the Middle Vermilion Valley may be attributable to the southward escape of the underground water into that valley.

The very low head which is shown by some wells in the midst of the Bloomington morainic belt at points remote from the North and Middle Vermilion rivers may be difficult to explain. Defective work by the well driller is likely to account for some of this lack of head, for much depends upon the skill of the well driller in selecting the proper position for the pump strainer. It is not improbable, however, that the low head may in some cases be due to the peculiar arrange-

\* Chemical Survey of the Water Supplies of Illinois; Prel. Rept. to University of Illinois, 1897, pp. 52-53.

ment of the water-bearing beds. The fact that the low head is found in the morainic or jumbled-up portions of the drift to a far greater degree than in the plain or less disturbed portions, at least suggests that a disturbance or a lack of continuity of the water-bearing drift beds may cause the low head.

The question whether there is a difference of head displayed by water veins from different beds in the drift has been considered, but the data relative to the head are seldom sufficiently precise to throw light on the matter. The most precise data found relate to an area lying outside the quadrangle, in the Potomac flowing-well district. A well in Potomac, on the property of W. C. Mesner, struck a water-bearing bed at 85 to 96 feet which has a head 18 inches above the well mouth, and another at 127 feet which has a head 6 feet above the well mouth. It is reported by citizens of Potomac that several other wells show a similar

increase in passing from the upper to the lower vein of water. It is apparent that so slight an increase as is here shown would scarcely be detectable in wells which do not flow, and would be of little or no advantage unless a flow depended upon it.

The comparative head of wells terminating in the drift beds and those terminating in subjacent rock strata has also been considered, and it is found that the difference in head is usually very slight in neighboring wells of the two classes. The wells of moderate depth in rock, if near a deep drainage line, seem to be affected by it to about the same degree as wells that terminate in the drift. Referring to the table below, it will be seen that wells in sec. 5, T. 18 N., R. 10 W., and sec. 18, T. 18 N., R. 11 W., which are near stream valleys, show a much lower head than wells in sec. 13, T. 18 N., R. 12 W., and sec. 17, T. 19 N., R. 10 W., which are several miles distant from any

deep valley. It is probable that the wells of this class obtain much of their supply by percolation of water from the drift into the rock, and if so the head should be essentially the same as in wells that terminate in the drift.

Two deep borings have been made at Danville, one in North Vermilion Valley to a depth of 1250 feet, the other at Danville Junction near the corner of Collett and Wellington streets. A flow of brackish water was obtained from the one in North Vermilion Valley, but the one at Danville Junction, being on ground about 80 feet higher, did not obtain a flow. The precise head of either well was not ascertained, though it seems to be but little above the level of the well in the valley, 530 feet. No use is made of either of the borings.

*Tabulated well data.*—In the table which follows, but few logs of the wells are given, for the reason that nearly every well is reported to have been sunk mainly through boulder clay until the

water-bearing bed was struck. Where weak water veins are known to have been struck at higher levels than the main vein, note is made of them in the table. In nearly all cases the main vein is very near the bottom of the well. In many cases no information could be obtained as to the kind of material in which the water is found, but where wells are strong it may safely be assumed that it is in sand or gravel. Where the head or the depth is only approximately known that fact is indicated in the table. In many wells these were determined with considerable precision at the time the windmill was attached.

General geology by  
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Pleistocene geology by  
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December, 1900.

#### TABULATED DATA CONCERNING WELLS IN THE DANVILLE QUADRANGLE.

COMPILED BY FRANK LEVERETT.

[Approximate depths and altitudes indicated by the letter a.]

LOCATION, OWNER, ETC.	ALTI. TUDE.	DEPTH.	HEAD.	WATER-BEARING BED.	REMARKS.	LOCATION, OWNER, ETC.	ALTI. TUDE.	DEPTH.	HEAD.	WATER-BEARING BED.	REMARKS.
		Ft. above tide.	Feet.	Ft. above tide.					Ft. above tide.	Feet.	
T. 21, R. 10 W. Sec. 23, east part.....	720	120	680a	Gravel.	Reported by George Logan, well driller.	Same section, near center.....	690	80	?	Gravel below till.	Three good wells at 60 to 80 feet.
T. 21, R. 10 W. Sec. 24, west part.....	710	107	680a	Gravel.	Reported by George Logan, well driller.	T. 20, R. 11 W. Sec. 18, R. H. Conner.....	640	108	600	Gravel below till.	On gravel terrace, but penetrates till above water bed.
T. 21, R. 10 W. Sec. 26, Wm. Kizer.....	690	50	680a	Sand.	Reported by George Logan, well driller.	T. 20, R. 11 W. Sec. 19, John Phipps.....	690	140	600a	Gravel.	Water rises but little in the well.
Same section, Mrs. Mary Kizer.....	730	78	680a	?	Reported by owner.	Same section, E. H. Morrison.....	660	100	600a	Gravel.	A well near by, 28 feet deep, is much stronger.
Same section, W. J. Ashby.....	730	72	680a	?	Reported by owner.	T. 20, R. 11 W. Line of secs. 21 and 22, Philip Zorn and Joseph Phillips.....	690	210	650	Sandstone.	Rock struck at 175 feet; water is soft.
T. 21, R. 10 W. Sec. 27, north part.....	730	142	680a	?	Reported by tenant.	T. 20, R. 11 W. Line of secs. 23 and 23, E. W. and L. P. Adams.....	690	220	650a	Sandstone.	Rock struck at about 160 feet; water is soft.
T. 21, R. 11 W. Sec. 19, (Myersville and vicinity, several flowing wells). . . . .	630	145	650	Gravel between till.	Same section, Noah Young.....	680	175	660a	Sand.	Well at residence.	
	640	156	650		Same section, Noah Young.....	650	80	660a	Sand.	Well in field.	
T. 21, R. 11 W. Sec. 20, Noah Young.....	680	175	660a	Sand.	Same section, J. D. Campbell.....	680	148	660a	Sand.	Reported by J. D. Campbell.	
Same section, Noah Young.....	650	80	660a	Sand.	Same section, schoolhouse.....	680	117	660a	Sand.	Well was driven to 125 feet, but completed at 80 feet.	
Same section, Grant Wilson.....	680	188	660a	Sand.	Same section, Milton Berry.....	680	80	660a	Sand.	Well very strong.	
Bismarck, George Taylor.....	655	170	660	Sand between till.	Same section, Perry Makenson.....	680	75	650	Gravel.	Weak well.	
Bismarck, S. W. Claypool.....	674	181	660	Sand and gravel.	Same section, Newell Makenson.....	670	60	630	Gravel.	Reported by John Hoover.	
T. 21, R. 11 W. Sec. 26, W. Burnside.....	700	75	680	Sand and gravel.	T. 20, R. 11 W. Sec. 24, A. N. Leavenworth.....	710	130a	?	Gravel.	Another bored well on same farm is 85 feet and two dug wells 50 feet. The 110-foot well is the strongest.	
T. 21, R. 11 W. Sec. 27, F. Starr.....	690	73	670	Sand and gravel.	T. 20, R. 11 W. Sec. 26, A. Campbell.....	670	110	662	Gravel.	Blue till, 40 feet; pink till to water vein near bottom. Another boring 170 feet found no water; it passed through a black material (shale) at 160 to 170 feet.	
Same section, A. Cass.....	690	73	670	Sand.	T. 20, R. 11 W. Sec. 27, R. L. Howard.....	670	125	630	Sand.	Reported by owner.	
T. 21, R. 11 W. Sec. 28, Ira Ingram.....	685	90	673	Gravel.	Same section, C. P. Cozatt.....	700	95	692	Gravel.	Blue till, 40 feet; pink till to water vein near bottom. Another boring 170 feet found no water; it passed through a black material (shale) at 160 to 170 feet.	
T. 21, R. 11 W. Sec. 30, Allen Woods.....	615	110	633	Sand.	T. 20, R. 11 W. Sec. 28, Mrs. Sperry.....	700	117	635	Gravel.	Reported by C. P. Cozatt.	
Same section, Leonard farm.....	635	180	?	Sand.	Same section, Mrs. Sperry.....	700	315	?	Rock.	A boring for coal entered rock at 180 feet and found coal at 210 to 220 feet.	
Same section, John Furay.....	660	130	?	?	T. 20, R. 11 W. Sec. 29, E. Winter.....	685	180	585	Gravel.	Reported by owner.	
T. 21, R. 11 W. Sec. 33, J. I. Cunningham.....	690	70	670	?	Same section, E. Osborn.....	675	183	600	Gravel and top of rock.	One water vein at 175 feet, another at 180 feet, penetrated coal at bottom of boring 6 inches.	
Same section, Z. Deck.....	690	70	670	?	Same section, E. Rainier.....	700	176	580	Gravel.	Reported by owner.	
Same section, John Campbell.....	690	90	?	Sand.	Same section, J. W. Jones.....	665	120	620	Gravel.	Reported by owner.	
Same section, Daniel Cox.....	680	100	655	?	Same section, W. Myers.....	660	190	595	Gravel.	Terminated in gravel.	
T. 21, R. 11 W. Sec. 34, W. Nors.....	720	94	680a	?	Same section, P. Leavenworth.....	670	128a	?	Gravel.	Reported by W. Myers.	
Same section, George Huffman.....	710	77	676	?	Same section, Charles Keasler.....	675	128	?	Gravel.	Reported by owner.	
Same section, George Huffman.....	740	110	?	Sand.	Same section, Ira Weaver.....	670	65	?	Gravel.	Very strong well. Reported by owner.	
Same section, Mrs. Starr.....	710	70	?	?	T. 20, R. 11 W. Sec. 32, several wells.....	670a	100a	?	Gravel.	Several good wells at about 100 feet.	
T. 21, R. 12 W. Sec. 25, Prior Woods.....	630	44	635	Gravel below till.	T. 20, R. 11 W. Sec. 33, Jacob Hosch.....	700	98	?	Red sand.	Water rains from 90 to 100 feet, but some from 50 feet.	
Same section, Allen Woods.....	675	195	669	Gravel at 60 feet.	T. 20, R. 11 W. Sec. 34, Andrew Hosch.....	690	116	?	Gravel.	Yellow till, 15 feet; soft blue till, 70 feet; red sand, 8 feet. One-fourth mile south a well has a strong supply of water at 60 feet.	
Same section, Leonard farm.....	680	126	660	Gravel.	Same section, Charles Shean.....	690	65	?	Gravel.	Reported by Jacob Hosch.	
T. 21, R. 12 W. Sec. 26, Andrew White.....	690	72	678	Sand.	T. 20, R. 11 W. Sec. 35, S. T. Makenson.....	690	150	640	Gravel.	Reported by C. P. Cozatt.	
Same section, at schoolhouse.....	690	72	?	Sand.	T. 20, R. 11 W. Sec. 36, Jacob Hosch.....	700	98	?	Red sand.	The upper 80 feet mainly clay, remainder sand and gravel. Well sunk deep to obtain large supply.	
Same section, John White.....	690	65	?	Sand.	T. 20, R. 12 W. Sec. 1, Henry Swisher.....	690	150	?	?	Reported by owner.	
T. 21, R. 12 W. Sec. 29, V. Lane.....	690	100	?	Sand.	Same section, T. A. Hoskins.....	690	45a	655	?	Reported by owner.	
Same section, C. Lane.....	690	145	676	?	T. 20, R. 12 W. Sec. 5, Austin Juvinall.....	610	54	610	Sand.	Water barely rises to the surface.	
T. 21, R. 12 W. Sec. 32, B. Crawford.....	660	255	?	?	Same section, T. A. 10, schoolhouse.....	675	155	?	?	Reported by Henry Gray, well driller.	
Same section, F. Crawford.....	670	90	657	?	Same section, George Beaver.....	685	77	?	?	Reported by Henry Gray, well driller.	
T. 21, R. 12 W. Sec. 33, George Smith.....	690	115	680	?	Same section, Milton Fairchild.....	680	70	?	?	Reported by Henry Gray, well driller.	
T. 21, R. 12 W. Sec. 34, W. Henthorn.....	680	70	?	Sand.	T. 20, R. 12 W. Sec. 12, J. Johnson.....	600	65	?	Sand.	Boring was sunk to 165 feet without increasing supply.	
T. 21, R. 12 W. Sec. 36, Henry Woods.....	630	70	640	Gravel below till.	Same section, Henry Finley.....	650	70	?	Sand.	Reported by J. Johnson.	
Same section, John White.....	630	35	640	Gravel below till.	T. 20, R. 12 W. Sec. 13, John Fugot.....	680	200	?	Black sand.	Stopped in black sand. A well near by got good supply of water in gravel at 75 feet.	
Same section, John White.....	680	116	?	Fine sand.	Same section, F. Osborn.....	675	58	?	?	Good well completed at 58 feet after boring 70 feet.	
T. 20, R. 10 W. Sec. 17, State Line, public well.....	715	140a	660a	Fine sand.	Same section, O. B. Gravat.....	700	79	674	Gravel.	Reported by owner.	
Same section, John Hoover.....	720	144	660a	Fine sand.	T. 20, R. 12 W. Sec. 14, W. B. McKinley.....	700	200	?	?	A poor well, but a good one 40 feet deep was obtained near it.	
Same section, State Line tile factory.....	715	140a	660a	Sand.	Same section, G. L. Fairchild.....	730	118	620	Gravel.	Reported by owner.	
Same section, Mrs. Mary Daniels.....	710	90	660a	Sand.	Same section, Mrs. English.....	700	{ 240	?	Gravel.	Gas was struck near bottom in such amount that a new well only 75 feet deep was made near it.	
T. 20, R. 10 W. Sec. 18, Ephraim High.....	720	136	670a	Sand.	T. 20, R. 12 W. Sec. 15, E. Cunningham.....	725	134	?	?	Reported by Henry Gray, well driller.	
Same section, Mrs. J. M. Adams.....	720	135a	660	Sand.	Same section, E. Cunningham.....	700	75	?	?	A well 75 feet down has sulphur water, but one near by 58 feet deep has no sulphur.	
Same section											

## TABULATED DATA CONCERNING WELLS IN THE DANVILLE QUADRANGLE.—CONTINUED.

LOCATION, OWNER, ETC.	ALTI-TUDE.	DEPTH.	HEAD.	WATER-BEARING BED.	REMARKS.	LOCATION, OWNER, ETC.	ALTI-TUDE.	DEPTH.	HEAD.	WATER-BEARING BED.	REMARKS.
	Ft. above tide.	Feet.	Ft. above tide.				Ft. above tide.	Feet.	Ft. above tide.		
T. 20, R. 12 W. Sec. 25, O. G. Knight....	690	144	600a	Fine sand.	Sand is so fine as to clog pump strainer; several unsuccessful well borings on same farm.	Danville, College street south of Main street.	615	114	570	Sand.	Black soil struck at about 90 feet, below which was water-bearing sand. Weak well.
Same section, Johnson Knight.....	690	130	?	Sand.	Better well than preceding one. In sand below till. A well across road at Jacob Bailey's has similar section and head.	Danville, Columbus street south of Main street.	630	101	570	Sand.	Reported by owner, A. Bowman.
Same section, Jacob Mead.....	690	147	610	Sand.	Several tubular wells at 75 to 130 feet. Head is low.	Danville, Harrison street near Wabash R.R.	600	58	560	Sand.	Entirely sand and gravel.
T. 20, R. 12 W. Sec. 26, several wells....	{ 680 700	{ 75 130 }	610a	Sand and gravel.	A coal boring entered shale at about 200 feet.	Danville, P. Yager, Vermilion street near English street.	600	63	540a	Gravel.	Largely cemented gravel.
T. 20, R. 12 W. Sec. 27, E. Allen.....	700	{ 106 200	?	Rock.	A well was obtained in gravel at 106 feet. Wells on higher ground in same section are 60 to 90 feet in depth.	Danville, J. Foreman, Vermilion street near English street.	600	66	540a	Gravel.	Upper 15 feet clay; remainder gravel and sand.
Same section, Mr. Grimes's and other wells.....	675	{ 96 60	?	Gravel.	Some good wells in this section at 75 to 90 feet.	Danville, A. Hess, Oak street north of Fairchild street.	600	60a	540a	Gravel.	Clay 30 feet; remainder gravel.
T. 20, R. 12 W. Sec. 35, J. S. Lee.....	690	115	635	Gravel.	An exceptionally deep well for this section. Mainly through blue till. A well near by only 18 feet deep is strong.	Danville, A. J. Woodbury, corner of Oak and Clay streets.	600	53	540a	Gravel.	Water comes in below gravel at top of coal.
Same section, Mrs. Lizzie Hughes.....	705	150	?	Gravel.	Reported by owner; mainly through blue till.	Danville, N. W. part, R. H. Carnahan, James Knight, and James Clark.	600	24	.....	Top of coal.	In rock near bottom.
T. 20, R. 12 W. Sec. 36, F. Buy....	690	125	590	Gravel.	Strong well.	Danville, Mrs. James Partlow, north part of city.	630	158	?	Rock.	Several wells about 100 feet, mainly through till. A few good wells at 20 or 30 feet.
T. 19, R. 10 W. Sec. 6, Frank Shafer....	650	96	630a	Sand.	Strong well.	Danville, Roselawn addition.....	650	100a	?	Gravel.	Mainly through till.
Same section, Isaac Currant.....	650	85	630a	Sand.	Reported by owner; mainly through blue till.	Danville, N. W. part, F. Van Valkenburg.	630	35	?	Gravel.	Largely through gravel.
T. 19, R. 10 W. Sec. 7, Sam. Martin....	650	65	637	Sand.	Strong well; other good wells near at 30 feet.	Danville, N. W. part, J. Perry.	630	60	?	Gravel.	Rock struck near bottom.
Same section, T. O. Talbert.....	650	60	630a	Sand.	Soil and yellow till, 9 feet; gray till, 9 feet; dry red sand, 1 foot; hard, pebbly blue till, 40 feet; blue gravel with but little water, 5 feet; reddish clay, soapy, with few pebbles, 40 feet; ash or gray pebbly clay, 10 feet; blue pebbly clay, 20 feet; reddish sandy clay, 20 feet; soft sandstone, 8 feet; shale, 1 foot; coal, 6 feet; brecciated bed, 6 feet.	Danville, N. W. part, E. H. Palmer.	630	45	?	Rock.	Clay, 8 feet; gravel, 18 feet; till near bottom.
T. 19, R. 10 W. Sec. 17, R. Peterson....	650	182	638	Rock.	Deeper boring has little water; 60-foot well near by strong.	West Danville, Sanford Bailey.	610	28	?	Gravel.	Good wells at 10 to 20 feet in gravel or sand below till.
T. 19, R. 10 W. Sec. 18, W. Olehy.....	645	{ 140 60	635	Sand and gravel.	Reported by William Martin.	Batestown, several wells.....	640	20a	630	Gravel.	A well on a kame penetrated clay 4 feet and then sandy gravel to bottom.
Same section, Rileysburg elevator....	650	90	635a	?	Good well at about 16 feet.	One mile southwest of Batestown.....	655	15	640	Sandy gravel.	Wells usually enter water-bearing sand below till at slight depth.
T. 19, R. 10 W. Sec. 30, several wells....	650	16	640	Sand.	Much fine sand; till is weak. Other wells near by, 20 to 30 feet in depth, are strong.	Hillery and vicinity.....	650	15a	640	Sand.	Many wells obtain a good supply at 10 or 12 feet.
T. 19, R. 11 W. Sec. 1, schoolhouse....	650	175	?	Fine sand.	Reported by C. B. Brant.	Catlin.....	655	12	645	Sand and gravel.	Wells usually penetrate 15 or 20 feet of till before striking water-bearing sand and gravel.
T. 19, R. 11 W. Sec. 2, S. Hancock....	640	75	?	?	Water from blue silt at 92 feet; shale struck at 100 feet.	West of Catlin, several wells.....	660	25	645	Gravel.	Wells seldom exceed 25 feet, and obtain water in gravel or sand below till.
T. 19, R. 11 W. Sec. 10, J. Hough....	630	100	?	Silt.	Very strong well; springs in vicinity.	Northeast of Catlin, several wells.....	650	25	640	Sand or gravel.	Mainly till to 120 feet, coal at 120 to 126 feet.
T. 19, R. 11 W. Sec. 11, C. B. Brant....	620	12	?	Gravel.	Well rather weak.	T. 18, R. 10 W. Sec. 5, Milton Beau-champ.	620	135	550	Rock.	Water from shale below coal.
Same section, Charles Schoering....	640	60	610	Sand.	Reported by Robert Brasker, well driller.	T. 18, R. 10 W. Sec. 7, J. M. Jones....	635	60	618	Gravel.	Mainly through blue till.
T. 19, R. 11 W. Sec. 12, George Villars....	650	115	?	?	Well strong. Reported by owner.	T. 18, R. 11 W. Sec. 3, Mr. Boone....	550	39	?	Sand.	In Hawbuck Creek Valley, probably in line of a pre-Glacial valley.
Same section, George Walz.....	650	57	?	?	Well very weak; good well at neighboring farmhouse at 40 feet.	Kellyville, several wells.....	670	20	?	Sand.	Wells usually get water at 15 to 20 feet; shale entered at 40 feet.
T. 19, R. 11 W. Sec. 14, schoolhouse....	630	170	?	Fine sand.	Several wells obtain water in gravel below till about 50 feet. A boring 370 feet in depth obtained salt water.	Westville, several wells.....	675	15	?	Sand.	Wells 20 to 30 feet are usually strong.
Soldiers' Home, near Danville.....	635	{ 50a 370	615	Gravel.	Four wells stop at top of rock; water is hard and chalybeate. Capacity each 20 gallons per minute. A boring 310 feet deep and lower sand and pebbles quality of water than the shallower wells.	South Westville, several wells.....	680	25	?	Sand.	Rock entered at 90 feet. Surface sand 25 feet; remainder of drift mainly blue till.
Danville, Brewery.....	600	{ 78 315	560	Gravel.	Two strong wells from gravel near bottom. About 50 strong 3-inch wells 50 to 70 feet deep from gravel below till. Temperature 52° Fahr. Water rather hard.	T. 18, R. 11 W. Sec. 18, Mrs. Annie Sheets.	615	130	525	Sandstone.	On drift knoll. Well in sand and gravel entire depth.
Danville, Douglass Park.....	595	60	?	Gravel.	Very strong well from gravel below till. Water rather hard for boiler use.	T. 18, R. 11 W. Sec. 27, schoolhouse....	680	200	?	Rock.	Drift, 80 feet. Well rather weak.
Danville, starch works in Stony Creek Valley.....	575	{ 50 70	560a	Gravel.	A well 24 by 30 feet yields 750 gallons per minute without signs of lowering after 24 hours pumping. Is used by workmen at C. & E. I. shops. Water is from gravel below till and is hard.	Georgetown, several wells.....	670	{ 15 35	?	Gravel.	Wells seldom reach 35 feet before obtaining good supply.
Danville, Big Four Railway in Stony Creek Valley.....	575	30a	560a	Gravel.	Boring west of car shops. No water obtained; mainly blue till.	T. 18, R. 12 W. Sec. 3, Sandusky Hill..	708	100a	?	?	A well on top of hill was sunk to a depth of about 100 feet, mainly through sand and gravel.
Danville, C. & E. I. R. R. between Section street and Stony Creek.....	575	32	560a	Gravel.	Borings in Stony Creek Valley north of car shops obtained no water at depth of 40 feet.	T. 18, R. 12 W. Sec. 4, Marion Burroughs.	672	81	670	Sand.	Mainly through blue till; sand at bottom.
Danville, C. & E. I. prospect borings....	{ 610 575	{ 185 40	?	Clay.	Rock at bottom. Water from gravel below till.	Same section, R. O'Connell.	675	98	?	Shale.	Well weak; shale entered at 92 feet.
Danville, Bowman street north of Seminary street.....	620	70	.....	Gravel.	No rock reached. Water from sand below till. Some water at 60 feet.	T. 18, R. 12 W. Sec. 5, several wells....	675a	30a	?	Gravel.	Good wells usually obtained at less than 35 feet, but one well in south part of section is 69 feet.
Danville, corner of Anderson and Seminary streets.....	620	80	570	Sand.	Head exceptionally high.	T. 18, R. 12 W. Sec. 8, W. Church....	680	{ 40 100	?	Gravel.	A good well from gravel below till at 40 feet. Another boring 100 feet in depth did not obtain water. Shale was entered at 80 feet.
Danville, corner of Bowman and Seminary streets.....	620	59	585	Sand.	Entered rock at 104 feet; weak vein in fine sand at 67 feet. Drift mainly till.	T. 18, R. 12 W. Sec. 13, Dennis Clingan.	680	90	663	Rock.	A prospect boring for coal struck water in rock at 90 feet. Coal was found at about 210 feet.
Danville, corner of Kimball and Wellington streets.....	620	114	570	Rock.	Reported by owner; mainly through blue till; gravel near bottom.	T. 18, R. 12 W. Sec. 15, R. Clipson....	685	70	665	Gravel.	Mainly till; gravel near bottom.
						Same section, Charles Barley....	675	70	667	Gravel.	Good well from gravel below till; another boring 100 feet deep reached shale without obtaining water.
						T. 18, R. 12 W. Sec. 16, J. M. Douglass....	{ 696 690	{ 45 100	?	Rock.	Mainly till; sand near bottom.
						Same section, W. S. Douglass....	690	70	?	Sand.	Mainly till; gravel near bottom.
						T. 18, R. 12 W. Sec. 22, David Barley....	675	70	667	Gravel.	