

AVAILABILITY OF GROUND WATER IN THE HICKMAN QUADRANGLE, KENTUCKY-MISSOURI-TENNESSEE

An abundance of ground water for domestic, irrigation, and industrial supplies is available in the west-central part of Fulton County, Kentucky. This valuable resource has not been used and will furnish sufficient amounts of water for future industrial developments and public demands. This atlas, one of a series that includes the entire Jackson Purchase region in western Kentucky, presents information about ground water in an area near Hickman.

The most important aquifers are the alluvial deposits of Quaternary age and the sands of Eocene age. Less important aquifers are the gravel or silt of Pliocene(?) age, the sands of Cretaceous age, and the limestone and dolomite of Paleozoic age.

The occurrence of ground water in the Hickman area is indicated on a water-availability map which is a graphic representation of the shallowest aquifer that may yield water in adequate amounts for domestic use. The availability of ground water at a particular site can be determined from the pattern on the map and from a study of data on wells near the site. The chemical quality of the water is shown by circular diagrams.

The maximum thickness of the alluvium may be more than 200 feet, but only 165 feet of alluvium has been penetrated by test holes. Most driven wells tap an upper sand, while jetted wells tap an underlying pebbly sand. Driven and jetted wells are labeled as driven wells in the Mississippi River bottoms lands. The depth to the main zone of saturation in the Quaternary alluvium ranges from a few feet below land surface near the valley wall at Hickman to more than 25 feet below land surface near the Mississippi River at low-river stage. Not monthly throughout the year, the ground water flows from the valley wall toward the river. At high-river stage, the flow of ground water near the river is reversed and wells near the levees may flow. The bottom land north of Hickman is normally flooded each spring. The altitude of the

maximum known flood is 316 feet (Hickman gage). Very little ground water is pumped from the alluvium in this area. Most wells are equipped with small capacity piston pumps. The most productive wells yield about 15 gpm (gallons per minute), but larger yields for industrial demands can be obtained from wells designed and equipped for this purpose.

Because an abundant quantity of ground water is available from the alluvium, the sands of the Eocene formations in the bottom lands have not been tapped for a water supply. Logs of oil-test holes in the adjacent Bondurant quadrangle, indicate several Eocene aquifers as potential sources of ground water. Wells tapping the deeper sand in the Eocene deposits underlying the flood plain may flow.

A minor aquifer is the gravel or silt of Pliocene(?) age which is tapped by several jetted wells around Brownsville. A few wells north of Brownsville are still in use, but recently many wells are being drilled into the underlying Eocene aquifers. Many dug wells tap the fine-grained facies of the Pliocene(?) gravel southeast of Hickman. They generally yield sufficient water for domestic use, except during long dry periods. Many small springs occur at the contact of the gravel with the underlying Eocene clay.

The principal aquifers in the Hickman quadrangle are the sands of Eocene age. Sparse data indicate the Eocene sediments to be about 1,300 feet thick and to dip generally southward. The Eocene aquifers can be divided into a shallow sand, two deeper sands, and a basal sand. The shallow sand and the upper of the two deep sands, both of which occur above mean sea level are tapped for domestic and farm use. The lower of the deep sands is tapped by the municipal wells in Hickman. The bottoms of the city wells are about 300 feet below mean sea level, but more than 200 feet of sand may be present below the bottoms of the wells. The potential yields of the upper two sands may be more than 100 gpm. The yield of the lower of the two deep sands may be as much as 2,000 gpm. Hickman pumped an average of about 285,000 gallons per day in 1966. The basal sand is too deep for present development. The water table in the Eocene deposits slopes

from the east edge of the quadrangle toward the Mississippi River. The difference in head between each of the Eocene aquifers is not known. Wells tapping the deeper Eocene deposits in the bottom lands may flow. In addition, during floods on the Mississippi River, water levels in Eocene aquifers near the river may show a loading effect caused by the additional weight of the floodwater on the flood plain.

The Porters Creek Clay of Paleocene age, which underlies the Eocene sediments, is not an aquifer; instead, it retards the ground-water movement between the Eocene and underlying Cretaceous sediments. The clay, about 300 feet thick, occurs at an altitude of 900 feet below mean sea level at Hickman.

Sediments of Cretaceous age, below the Porters Creek Clay, are about 300 feet thick and rest on limestone and chert of Paleozoic age. No wells tap the Cretaceous near Hickman, but in southeastern Missouri, many municipal wells tap this strata for their water supply. Wells tapping this unit will flow if they are located in the flood plain. The top of the Paleozoic bedrock is about 1,500 feet below mean sea level at Hickman.

The quality of the water from the alluvium is considered to be adequate for most uses, although some objectionable constituents are present. Water is very hard and contains from 292 to 536 ppm (parts per million) of dissolved solids. Iron and manganese are present in objectionable amounts and may impart a disagreeable taste to the water. More than 0.3 ppm of iron and manganese may cause staining of porcelain and fixtures. The high iron and manganese content of the water from the alluvium may limit its use without treatment. Owing to the shallowness of the alluvial water table, shallow wells are susceptible to pollution. The bottoms of the city wells at Hickman tapping the deeper sands is slightly acidic and contains 4.5 ppm of iron, in parts per million, and the hydrogen-ion concentration, expressed as pH, of the water samples analyzed. A pH of 7.0 indicates neutrality of a solution. Values higher than 7.0 denote increasing alkalinity, values lower than 7.0 indicate increasing acidity. Corrosiveness of water generally increases as pH decreases.

EXPLANATION

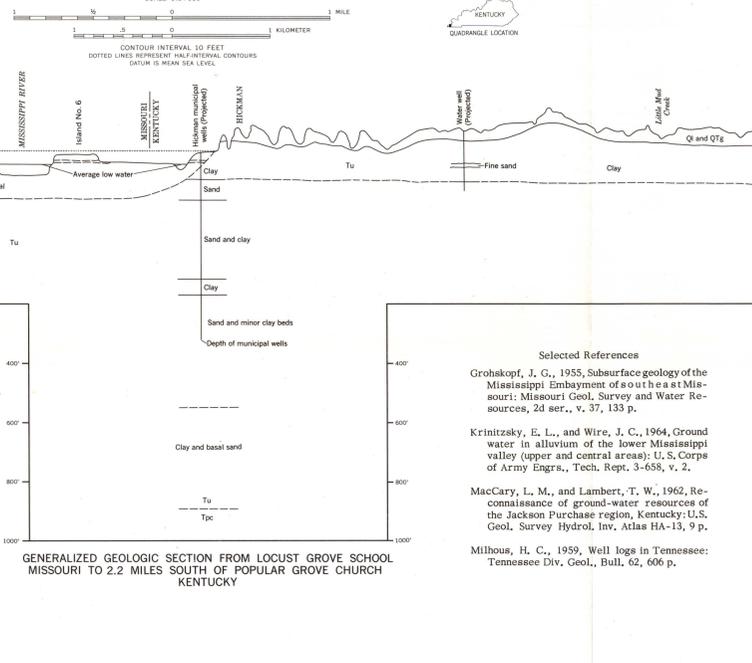
Gravel Sand Loess Clay Mica
Glaucous Carbonaceous material Limestone

1 2 3 4 5 6 7 8 9 10
Analysis number 0.08 0.27 0.31 0.9 0.88 0.87 0.68 1.84
Iron content 28 16 76 30 1.2 10
Manganese content 7.8 7.2 7.4 7.5 8.4 7.6 7.6 4.1 7.0 7.6
pH

11 12 13 14 15 16 17 18 19 20
Analysis number 0.39 1.1 7.6 5.4 4.7 1.9 0.18 0.96 4.5
Iron content 34 40 1.1 32 30 27 12
Manganese content 6.4 7.6 8.0 6.9 7.7 6.9 7.4 7.5 7.9
pH

21 22 23
Analysis number 0.05 6.0
Iron content 0.06 06
Manganese content 0.2 0.4 6.8
pH

* 1961; in 1968, 4.5 ppm of iron
† In solution and sediment



EXPLANATION

Q Quaternary
Q1 Alluvium of Quaternary age
Q2 Loess of Pleistocene age
Q3 Sandy gravel or silt of Pliocene(?) age
Tu Sand and clay of Eocene age
Tpc Porters Creek Clay of Paleocene age

Selected References

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