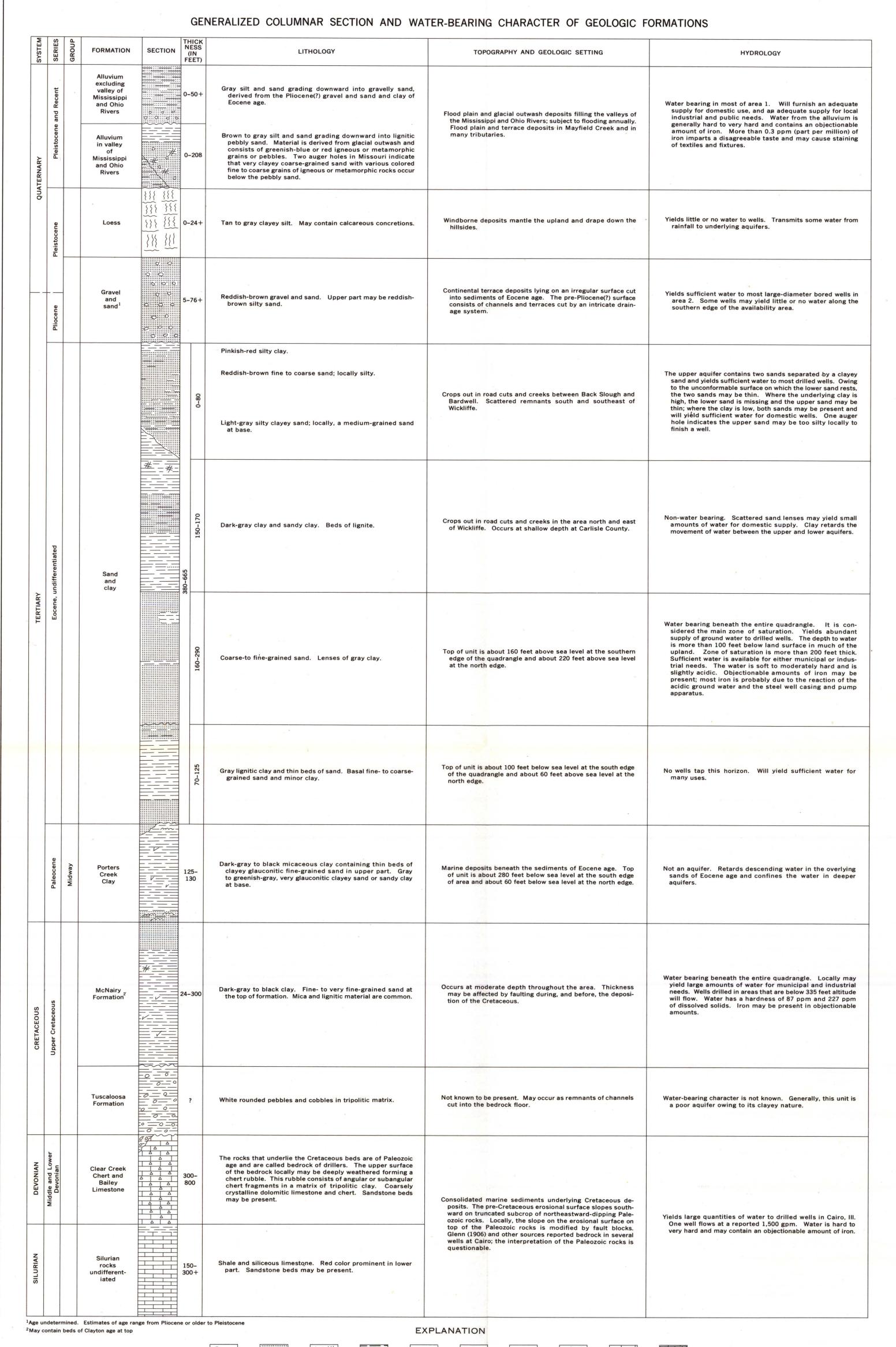
EXPLANATION



AVAILABILITY OF GROUND WATER IN THE PARTS OF THE WICKLIFFE AND WICKLIFFE NW QUADRANGLES IN JACKSON PURCHASE REGION, KENTUCKY

Sand

Loess

Clay or

Gravel

The Wickliffe area has large supplies of water readily available, both on the surface and below the ground. Although surface water in the area is an important resource for man's use, groundwater resources in the area are significant also but have not been fully developed. Both sources of water can be further developed in the future for increasing farm and industrial needs. This atlas, one of a series that includes the entire Jackson Purchase region in western Kentucky, presents information about ground water in an area near Wickliffe and Bardwell, Kentucky. The water-availability map is a graphic representation of the occurrence and quality of ground water in the shallowest aquifer that may yield

sufficient water in adequate amounts for domestic use. The availability of ground water at a particular site may be determined from the pattern on the map and a study of a data on nearby wells. Chemical quality is shown by circular diagrams. At the present time only minor amounts of ground water are withdrawn for domestic, farm, and municipal uses. Several aquifers are available for development in the area, but two aquifers, the Pliocene(?) gravel and alluvium, are restricted to certain areas. Important aquifers are the chert and limestone of Paleozoic age, sands of Late Cretaceous age and of early Eocene age, and gravelly sand of Pleistocene age; it is doubt ful that 150,000 gpd (gallons per day) of ground water are withdrawn from these aquifers. Many farms still use cisterns for their source of supply; a large percentage of them catch the runoff of water from roofs. The geologic formations underlying the Wickliffe area, shown in the geologic and columnar

sections, are discussed according to their relative ages, the oldest first. The Paleozoic bedrock is an aquifer in the

Wickliffe area and may be a large potential source of ground water. At Cairo, Ill., a reported 1,200 foot well tapping the Paleozoic bedrock, flows about 1,500 gpm (gallons per minute) and

has a measured water level about 9 feet above land surface. The water has a hardness of 214 ppm (parts per million) and total dissolved solids of 564 ppm. Several other wells in Cairo may have tapped the Paleozoic bedrock below 600 feet and are reported to have flowed about 100 gpm. However, these wells were not maintained and were destroyed. Fresh water was obtained also from the bedrock in an oil-test hole near the head of Gray Creek in the Arlington quadrangle (south of report area) from a depth of 2,000 feet. The Tuscaloosa Formation has slight importance as an aquifer in the Wickliffe area. It oc-

Chert

Cherty

curs in thin irregular pockets on the Paleozoic bedrock surface and consists of gravel mixed with a tripolitic clay matrix. The McNairy Formation is a minor aquifer throughout the area. Better aquifers occur above the McNairy Formation. Logs of an old oil-test well at Wickliffe, dirlled about 1900, and of an oil-test hole at Cunningham in Blandville quadrangle (east of report area), drilled in 1966, indicate the McNairy Formation is a hard compact clay and has a thin sand at the top of the formation. The oil-test well at Wickliffe was left as a water well. It flows 1.2 gpm and has a static water level about 9 feet above land surface (333 feet altitude). L. C. Glenn (1906, p. 125-126) reported the well to be flowing about 5 or 6 gpm. He believed the flow came from the 600-foot level.

The water has a hardness of 87 ppm and total dissolved solids of 227 ppm. The temperature of the water is 57°F. The Porters Creek Clay of Paleocene age is not an aquifer. The clay is a barrier to the movement of ground water between the younger and older aquifers. The most important aquifers in the area are the sands of Eocene age. They will yield large

quantities of good-quality water. The aquifers are divided into three units and are referred to as the shallow aquifer, the lower aquifer, and the basal aquifer. The basal aquifer is unimportant at this time owing to its thinness and depth. The relation between the Eocene sands in Ballard and Carlisle Counties is uncertain. The base of the shallow aquifer crops out in the upland of Ballard County and slopes southward in Carlisle

County where it occurs below the level of the flood plain of the Mississippi River. The aquifer appears to rest on an irregular surface of clay. Large-diameter bored wells tap the water in the shallow aquifer perched above clay in Ballard County, while drilled wells tapping the shallow aquifer in Carlisle County are as deep as 160 feet. Yields in Carlisle County may be as much as 200 gpm. L. C. Glenn (1906, p. 125) reported dug wells, probably in the shallow aquifer, in Wickliffe between 40 and 50 feet deep.

A black to bluish-grey lignitic clay separates the shallow aquifer from the more important lower aquifer and the clay has a reported thickness of 150 to 170 feet. No sand units in the clay are known in the Wickliffe area, but they may be present to the south and southeast. The clay retards the downward movement of water in the area, and numerous spring seeps are reported near Wickliffe. The clay, weathered white, underlies the alluvial deposits of the Mississippi River and Mayfield Creek and their tributaries and acts as a barrier to the upward movement of water from

the deeper aquifers. The lower and most important aquifer is a massive body of fine to coarse sand that underlies the clay. Many wells tap this lower aquifer and range in depth from 100 feet north of Wickliffe to 500 feet west of Bardwell. The potential yield of this aquifer is more than 1,000 gpm. The largest user of water from this aquifer is the town of Wickliffe; the average distribution for the town is slightly more than 20,000 gpd. There are three different water surfaces in the Eocene aquifers in the Wickliffe area. The difference in water levels between the shallow and lower aquifers ranges from 13 to 70 feet and increases from east to west. The water level in the shallow Eocene aquifer shows the effect of topographic situation; the topographic site does not affect the water level in the lower and basal

aquifers. Water levels in the shallow aquifer in-

dicate that the water table slopes away from the

Mississippi River toward Bardwell. The main

area of the shallow Eocene aquifer is in Carlisle

County. The altitude of the water surface is from

330 to 370 feet and is highest near the Mississippi

River. The water surface slopes toward the May-

Ohio River Bridge borings 52-105Tu field Creek drainage and discharges into the alluvial deposits. The water surface in the alluvium is a continuation from the Eocene and ranges from 330 feet to below 280 feet altitude near the river. The shallow Eocene aquifer is not believed to be present beneath the alluvium of the Mississippi River and Mayfield Creek. The water surface in the lower Eocene aquifer as shown by the blue contours on the water-availability map, slopes to the Mississippi River from about 325 feet altitude along the east edge of the quadrangle to about 290 feet altitude along the bluffs of the river. Deep wells tapping the lower Eocene aquifer in the bottom lands of the Mississippi River and Mayfield Creek may flow. Recharge to the lower aquifer is to the east in eastern Blandville and Lovelaceville quadrangles. No wells penetrate the basal Eocene aquifer at this time. Future wells penetrating the aquifer in the bottom lands of the Mississippi River and Mayfield Creek may flow. The Pliocene(?) gravel is a minor aquifer in the Wickliffe area. Several bored wells between Bardwell and Winford Junction tap the Pliocene (?) gravel at depths less than 50 feet. One well southeast of Wickliffe taps the Pliocene above lignitic clay of Eocene age. Yields are reported adequate for domestic use; however, the saturated thickness of the Pliocene may be too thin for a domestic supply in the southern part of availability area 2. The alluvium of Quaternary age is a potential aquifer of major industrial importance for the Wickliffe area. Saturated thickness of the alluvium ranges from 18 to 64 feet. There are no wells in use in the bottom lands of the Mississippi River or Mayfield Creek. Several dug or bored wells occur in the bottom lands of Shelton and Buckheart Creeks. Development of ground-water supplies in the area along the alluvial valley faces two problems. One is the annual spring flooding of the bottom lands; the second is the high iron and manganese content of the water from the alluvium that may limit the use of untreated water. Water of better quality may be obtained at greater depths from the Eocene sands. Present withdrawals of ground water from the Eocene and the alluvial aquifers in the area are insignificant compared to the vast quantity of water that is available for use. Large-capacity wells tapping the Eocene and alluvial aquifers would probably be capable of supplying all industrial, public, farm, and domestic needs in the forseeable future. Excess ground water drains from the perched Eocene aquifers, recharging the minor alluvial aquifers and maintaining the perennial flows of certain creeks in the area. The temperature of ground water in the Wickliffe area ranges from about 58° to 62°F and may be more useful in the summer as a coolant than Mississippi River water whose temperature ranges from 33° to 82°F at Cairo. Ground water contains little or no sediment and is generally uniform in quality throughout the year. The chemical quality of ground water is generally good except that the concentration of iron and maganese in the water from some aquifers is more than 0.3 ppm. More than this amount may cause staining of textiles and porcelain, imparts a disagreeable taste, and limits the use of untreated water for some applications. Water collected from wells near the Mississippi River has a wide range of hardness and total dissolved sol-The high iron content in water sampled from the Eocene aquifers is a special problem. It is believed that the content is not typical of the natural water from the aquifer, but that the iron results from corrosion of steel well casings by the slightly acidic water. Water samples from the Eocene aquifers collected from plastic-cased wells are noted to have a much lower iron content than those from steel-cased wells. It may also be possible that the iron content of the water from the Eocene aquifers may be increasing downdip. The following table shows the iron and manganese content, in parts per million, and the hydrogen-ion concentration, expressed as pH, of the water samples analyzed. ApH of 7.0 indicates neutrality of a solution. Values higher than 7.0 denote increasing alkalinity, values lower than 7.0 indicate increasing acidity. Corrosiveness of water generally increases as pH decreases. 2 3 4 5 6 7 8 9 10 0.12 | 0.51 | 0.19 | 0.91 | 1.3 | 0.06 | 0.01 | 0.42 | 0.13 | 2.3 | .32 | .00 | .02 | .01 | - | .00 | pH | 7.0 | 6.8 | 6.2 | 6.5 | 7.7 | 6.3 | 7.8 | 6.0 | 6.9 | 6.7 Analysis | 11 | 12 | 13 | 14 | 15 | 16 | 17 | 18 10 | 0.09 | 0.30 | 0.54 | 0.24 | 0.00 | 0.32 | 0.46 Manganese content .05 .01 .00 .02 .10 .11 .54 pH | 6.3 | 7.2 | 6.3 | 6.4 | 7.3 | 6.5 | 6.6 | 6.4 SELECTED REFERENCES Glenn, L. 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of the Jackson Purchase region, Kentucky:

U. S. Geol. Survey Hydrol. Inv. Atlas HA-13.

EXPLANATION Alluvium of Quaternary age Gravel of Pliocene(?) age Sand and clay of Eocene age, undif-Porters Creek Clay of Paleocene age McNairy Formation of Cretaceous Devonian and Silurian rocks, undif-

The water-availability areas on this map show the occurrence and availability of ground water in the shallowest aquifer that may yield adequate amounts of water for domestic supply in each area. As considered in this report an adequate domestic supply will deliver approximately 500 gallons per day from a well equipped with a power pump and pressure-distribution system. The shallowest aquifer is underlain by deeper aquifers whose geologic and waterbearing properties are described in the generalized columnar section Water in Quaternary alluvium The alluvium is water bearing in area 1 at depths commonly less than 25 feet, although during the rainy seasons the area is commonly flooded by the Mississippi and Ohio Rivers. No wells are in use. Saturated thickness of the alluvium ranges from 20 to 30 feet at Winford Junction. An increase in saturated thickness should be expected in the area of Island No. 1 and Cane Island. The alluvium in the upper reaches of the tributaries of Mayfield Creek may be water bearing. Water from the alluvium may contain objectionable amounts of iron and manganese and is hard to very hard. Ground water of better quality occurs in the underlying Eocene sand Water in the Pliocene(?) gravel The Pliocene (?) gravel is water bearing in area 2 northwest of Bardwell. Large diameter bored wells are less than 50 feet deep and are reported to yield sufficient water for domestic use. The southern edge of the availability area is uncertain Water in Eocene sand Diagonal ruling shows areas in Ballard County where the water level in wells is more than 100 feet below land surface Abundant quantities of ground water are available in the Eocene sands at depths ranging from about 20 to more than 600 feet. Water levels range from a few tens of feet to more than 150 feet in the upland. Wells tap two aquifers separated by a thick clay Wells in Carlisle County generally tap two hydrologically connected sands above the clay and may be as deep as 160 feet. Locally, these sands are missing or too thin to complete a well; a well must tap the sand below the clay Wells in Ballard County tap the sand below the clay and range in depth from 80 to 364 feet. Yields of more than 1,000 gallons per minute may be obtained. Several shallow large-diameter bored wells tap sand lenses in the above clay and are reported to yield sufficient water for domestic Water from wells cased with steel casing may contain an objectionable amount of iron. Water from wells cased with plastic casing normally does not contain much iron. Water is soft to moderately hard and contains a low concentration of dissolved solids Area boundary Oil-test well Figure below line is depth of test hole, in feet below land surface Saturated thickness of alluvial or Pliocene(?) gravel, in feet Depth to base of gravel, in feet below land -Geologic unit underlying gravel (see aquifer symbols below) Depth of test hole, in feet below land surface Water well D, Drilled or jetted well, generally steel or plastic casing with well screen n, Driven well, sand point on lower end B, Bored or dug well, generally 24-inch concrete tile casing open at the Aquifer (see below) - Altitude of measured water level, in feet above mean sea level, in Carlisle County Water level in well; m, if measured; r, if reported. A plus sign (+) indicates the water level in feet above land surface; a minus sign (-) in feet Yield, in gallons per minute, or adequacy (see Depth of well, in feet below land surface AQUIFER SYMBOLS Alluvium of Quaternary age Perched water in sandy gravel of Pliocene(?) age Sand of Eocene age, undifferentiat Sand in the McNairy Formation of Cretaceous age Devonian rocks, undifferentiated YIELD OR ADEQUACY Yield, in gallons per minute - Well reported adequate for power pump for domestic and(or) stock supply Well reported adequate for hand pump or bailer Flowing well, measured flow in gallons per minute Destroyed or abandoned 300 ----Water-level contour Shows altitude of water level in the main zone of saturation. Dashed where approximately located. Contour interval 10 feet; datum is mean sea level. The approximate depth to water can be calculated by subtracting the altitude of the water level (as shown by the blue contours) from the altitude of the land surface. The difference will be the minimum depth of the well. Water-level measurements made in the late Chemical composition of dissolved solids Figure between circular diagram and well location refers to analysis number at end of text. Figure above line at center of circular diagram is carbonate hardness (calcium magnesium hardness as CaCO3), in parts per million, figure below line is dissolved solids in parts per million. Hardness of water is classified by the U.S. Geological Survey as follows: 0-60 ppm, soft; 61-120 ppm, moderately hard; 121-180 ppm, hard; and 181 ppm or more, very hard. Dissolved solids in partic analyses are computed from specific conductances and are only approximate values. Areas of the segments of each circle are proportional to the mineral component in the dissolved solids in the water. Percentages 89°07′30″ **5** Calcium and magnesium are shown as one segment in partial analyses All wells sampled contained less than 45 ppm of nitrate. Water con-Base from U.S. Geological Survey Wickliffe, 1951 MAP SHOWING AVAILABILITY OF GROUND WATER, LOCATION OF WELLS, AND QUALITY OF WATER Hydrology by T. W. Lambert, 1966 taining more than 45 ppm of nitrate may cause a type of methemogloand Wickliffe NW, 1951 binemia in infants ("blue baby" disease), sometimes fatal, and should not be used in infants' formulas. The Carter Coordinate System letters and numbers used SCALE 1:24 000 to designate five-minute divisions of latitude and longitude are shown along the margins; tick marks indicate one-minute divisions 1 72 0 KENTUCKY QUADRANGLE LOCATION HHHHH CONTOUR INTERVAL 10 FEET DOTTED LINES REPRESENT HALF-INTERVAL CONTOURS DATUM IS MEAN SEA LEVEL High river stage (Wickliffe gage) /Piezometric surface of confined of 327.5 feet, (1937) water in pre-Tertiary rocks Sand and minor clay beds Average low water, 1962 Qal Silt Piezometric surface of confined water in Eocene sediments \ Pebbly sand Sand and minor clay beds Sand and minor clay beds Clay and basal sand Clay and basal sand Base of sand in upper part of McNairy Formation

GENERALIZED GEOLOGIC SECTION ALONG A SOUTHEAST-TRENDING LINE FROM NEAR THE OHIO RIVER BRIDGE TO WICKLIFFE AND SOUTH-EAST TO THE VALLEY OF STOVALL CREEK

AVAILABILITY OF GROUND WATER IN THE PARTS OF THE WICKLIFFE AND WICKLIFFE NW QUADRANGLES IN JACKSON PURCHASE REGION, KENTUCKY

ATLAS HA-185

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