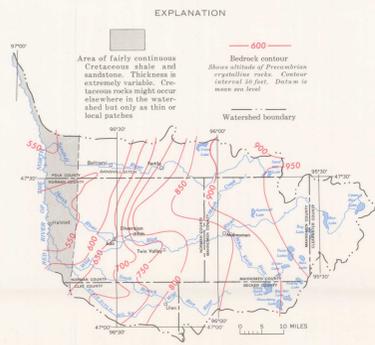


GROUND WATER



CONTOURS OF THE PRECAMBRIAN CRYSTALLINE ROCKS, WHICH FORM THE BASE OF THE GROUND-WATER RESERVOIR, SHOW RELIEF OF MORE THAN 400 FEET WITHIN THE WATERSHED. The general slope of this surface is toward the west. Bedrock valleys of low relief occur southwest of Ada. Cretaceous sediments occur mainly in the western part of the watershed. Thickness varies but is generally less than 50 feet and progressively thins toward the east. These deposits occur elsewhere in the watershed, but only as thin or discontinuous deposits.

The Cretaceous sediments consist largely of shale, but thin layers of fine to coarse sand are commonly found. Yields of wells tapping the sand lenses are usually less than 50 gallons per minute.

EXPLANATION

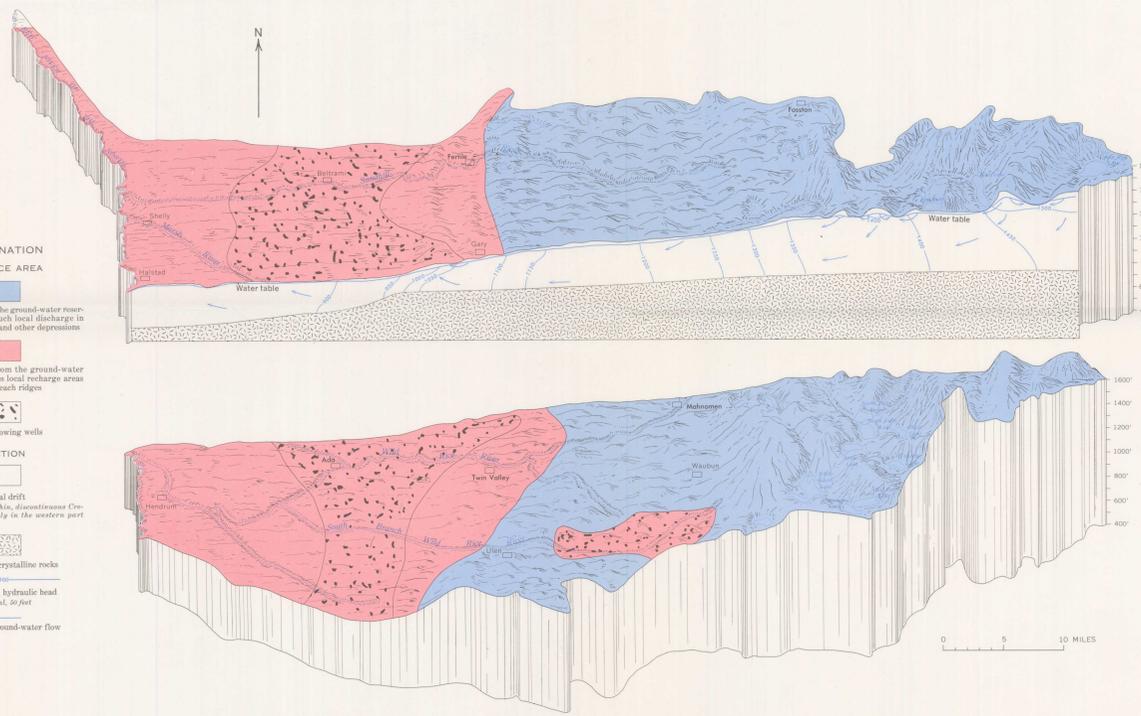
- Area of recharge to the ground-water reservoir. Includes local recharge areas in the areas of lakes and other depressions.
- Area of discharge from the ground-water reservoir. Includes local discharge areas in the vicinity of beach ridges.
- Area of flowing wells.

SECTION

- Glacial drift.
- Includes underlying thin, discontinuous Cretaceous rocks mainly in the western part of the watershed.
- Precambrian crystalline rocks.

EXPLANATION

- Line of equal hydraulic head. Interval, 50 feet.
- Direction of ground-water flow.

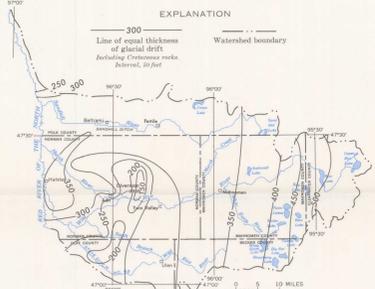


WATER ENTERS THE GROUND-WATER SYSTEM LARGELY IN THE ROLLING UPLANDS OF THE MORAINAL AREA.

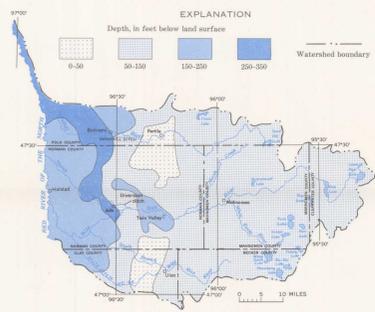
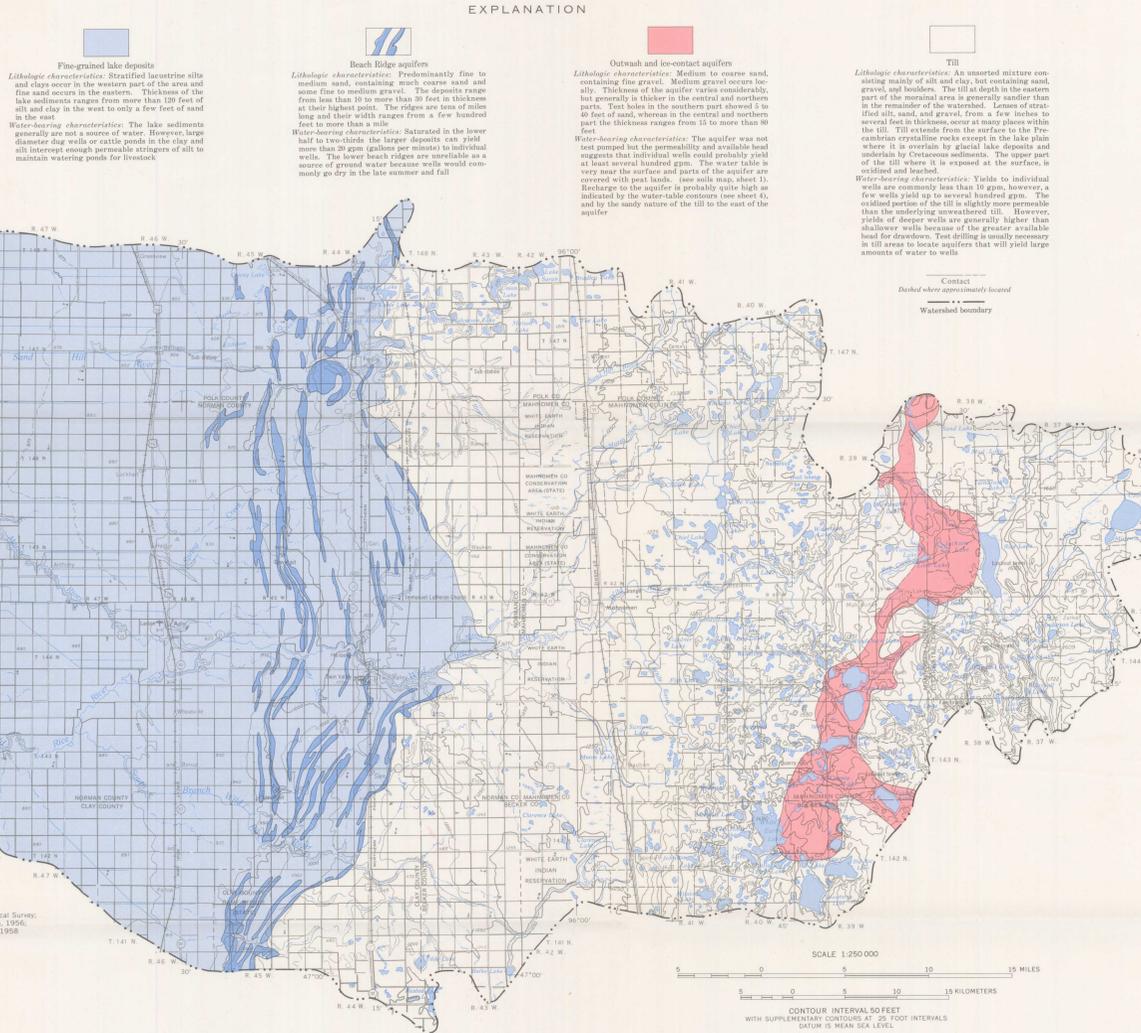
Most water in this area moves to and is discharged in adjacent depressions, commonly occupied by lakes, streams, or swamps. Some water, however, moves deeper into the ground-water reservoir where it becomes part of the regional flow system. Much of this water is discharged in the eastern and central part of the lake plain, as indicated by the zone of flowing wells, but the water deeper

in the reservoir moves upward in the vicinity of the Red River of the North. Within the regional discharge area, which is essentially the lake plain, small flow systems commonly are developed within the beach-ridge area where water enters the beach ridge and is discharged to nearby streams or adjacent swampy depressions.

Where the Cretaceous sediments are continuous with those in North Dakota the hydraulic head in Cretaceous sediments is generally greater than that in the glacial drift. This results in water moving upward and mixing with water in the drift.



THICKNESS OF THE GLACIAL DRIFT IS BETWEEN 200 AND 500 FEET THROUGHOUT MOST OF THE WATERSHED. The extreme thickness in the eastern part is due to a large north-south trending moraine forming the eastern boundary of the area. Drift thickness is an important factor when considering availability of water from glacial deposits (see availability map, this sheet). Cretaceous sediments were included in the glacial drift in the compilation of drift thickness because of the difficulty in distinguishing the contact from drillers' logs. However, where the Cretaceous sediments have been definitely identified (see bedrock map) the thickness is extremely variable and probably is not more than 50 feet at most places.



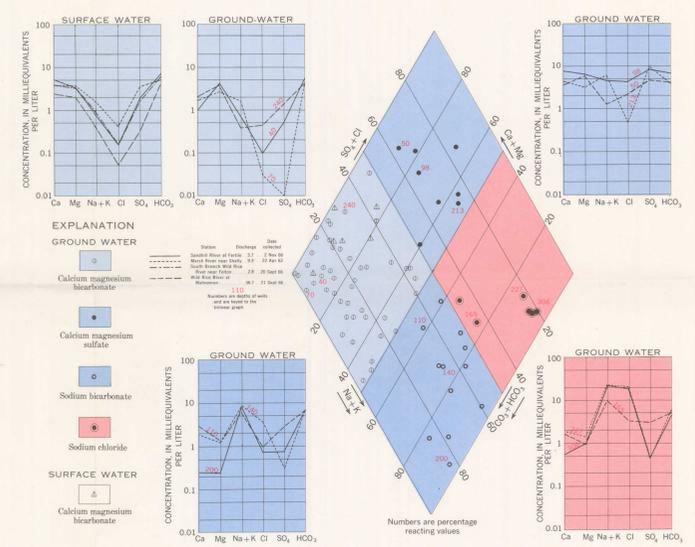
WATER FOR DOMESTIC AND FARM SUPPLY CAN USUALLY BE OBTAINED WITH A WELL LESS THAN 150 FEET DEEP IN MOST OF THE WATERSHED. However, it has been necessary to drill deeper than this locally. The areas where most wells are less than 150 feet deep are the part of the watershed where till is at or near the surface (for water-bearing characteristics of glacial till see availability map). Wells in the western part of the watershed are deeper because a considerable thickness of clay must be drilled through before reaching the glacial till, which contains permeable sand lenses. Wells less than 50 feet deep are most common in the areas of beach ridges.

Community	Well characteristics				Water quality				Remarks	
	Depth (feet)	Diameter (inches)	Operating rate (gallons per minute)	Depth to static water level (feet)	Specific capacity (gpm per ft)	Total dissolved solids (mg/l)	Hardness (mg/l)	Iron (mg/l)		
Ada	255	10	115	26	76	21	12	63	Chemistry well in town plots 600 gpm from same aquifer.	
Well 2	262	12	380	27	19	72	28	28		
Well 3	262	12	225	40	31	571	98	0.07	64	
Well 4	48	8	60	13	1.9	677	348	0.11	236	
Beju	95	8	45	12	1.9	677	348	0.11	236	
Cimax	120	6	50-100	20						
120	8	50-100	20							
Fertile	153	8	125	62	20	373	300	0.87	30	Abandoned well, 295 feet deep, owned by Northern Pacific Railway Company, yielded 80 gpm.
Halstad	297	8	125	25	16	1,350	207	1.1	207	
120	8	350	15			360	68	130		
Hendrum	300	8	60	11	1.3	170	23	350		
Mahomen	91	6	125	20	330	330	0.42	130		
143	6	75	20			360	68	130		
139	8	350	15			360	68	130		
Nielville	154	8	24	60		426	4.5	51		
258	6	75	29			120	2.9	41		
267	6	125	29			156	3	14		
140	4	50	20			678	156	3	14	
238	10	165	26			260	1.3			
240	10	75	26			260	1.3			
110	8	75	11			488	292	46	60	
111	6	125	11			488	292	46	60	

AVAILABILITY MAP SHOWS WHERE GROUND WATER CAN BE OBTAINED WITHIN THE WATERSHED

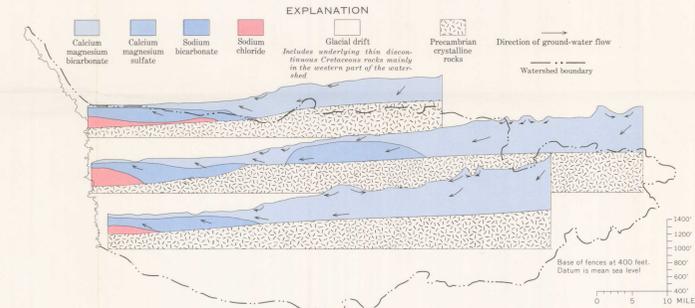
Ground water for domestic and stock use can be obtained at most places in the area. For larger yields the best potential is in the outwash and ice-contact sand and gravel in the eastern part of the watershed. Larger yields of ground water may be obtained in other parts of the area but test drilling is required to locate the sand and gravel bodies.

WATER QUALITY



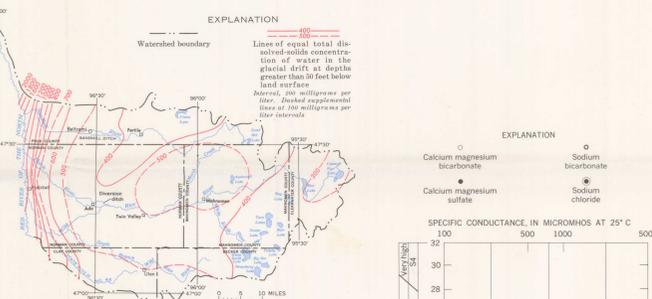
CHEMICAL QUALITY OF GROUND WATER VARIES WIDELY IN THE WATERSHED

Surface-water quality is fairly constant but concentrations of dissolved minerals are generally less at higher flows. Water can be classified into different types of chemical quality depending on the percentage distribution of the major ions in solution. The water type is controlled principally by the solubility of the minerals within the ground-water reservoir; the time water is in contact with these minerals, and the intermingling of water of different types. Some of the sodium bicarbonate water is due to softening, which is a decrease in calcium and magnesium ion concentration and an increase in sodium ion concentration. The logarithmic plots show the concentration of the major ions for selected water samples. The concentrations of most ions are between 0.1 and 10 milliequivalents per liter, but in the sodium chloride type the concentration of sodium plus potassium and chloride ions is about 20 milliequivalents per liter in some samples.



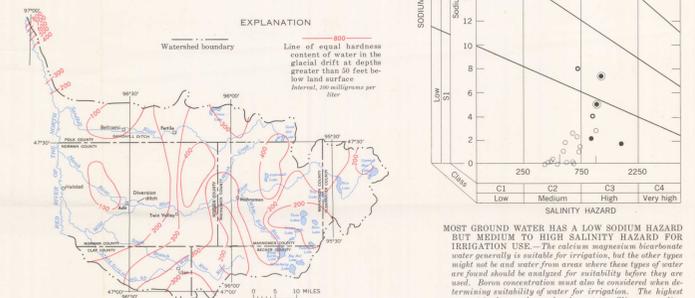
CALCIUM MAGNESIUM BICARBONATE, THE MOST COMMON TYPE OF GROUND WATER IN THE AREA, OCCURS MAINLY IN THE RECHARGE AREA

Sodium chloride water occurs in the Cretaceous sediments and in the glacial drift immediately adjacent to them. Ground water generally moves from east to west in the watershed and the mixing of the water from the eastern part of the area with water associated with the Cretaceous



CONCENTRATIONS OF DISSOLVED SOLIDS IN WATER OF THE GLACIAL DRIFT ARE HIGHER NEAR THE RED RIVER OF THE NORTH THAN IN THE EASTERN PART OF THE WATERSHED. The higher mineral content of water indicates that the water has been in contact with minerals in the ground-water reservoir for a long period, and that movement is probably slower than elsewhere in the watershed. The very high values in the extreme western edge reflect the influence of very poor quality water from the underlying sedimentary rocks which moves upward and mixes with water in the glacial drift.

Water containing dissolved solids of less than 1,000 milligrams per liter is suitable for domestic and stock uses. Some water containing 1,000 milligrams per liter or more of dissolved solids contains certain constituents which produce a noticeable taste. Concentration of nitrate in both ground water and surface water is generally less than 1 milligram per liter.



MOST GROUND WATER IN THE WILD RICE RIVER WATERSHED IS VERY HARD, BUT IN GENERAL THE HARDEST CONCENTRATION OF WATER IN THE LAKE PLAIN IS LOWER THAN IN THE MORAINAL AREA. The lower concentration of hardness in the lake plain is largely due to natural softening of ground water by base exchange. Hardness is due mainly to the presence of calcium and magnesium ions which react with soap to form insoluble products. Mineral salts of calcium and magnesium cause hardness and can form incrustation on well screens. Where hardness exceeds 200 milligrams per liter, incrustation of well screens may be rapid in small diameter wells that are heavily pumped.