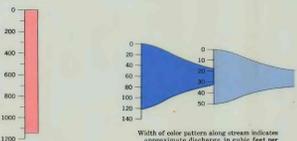


BASE FLOW AND WATER RECREATION

EXPLANATION



MORaine

Till and sand

Geologic contact

Low flow measurement site

Physiographic boundary

Watershed boundary

LAKE AGASSIZ PLAIN

Sand, silt, clay and peat underlain by till

Sandy areas consistently above state of river

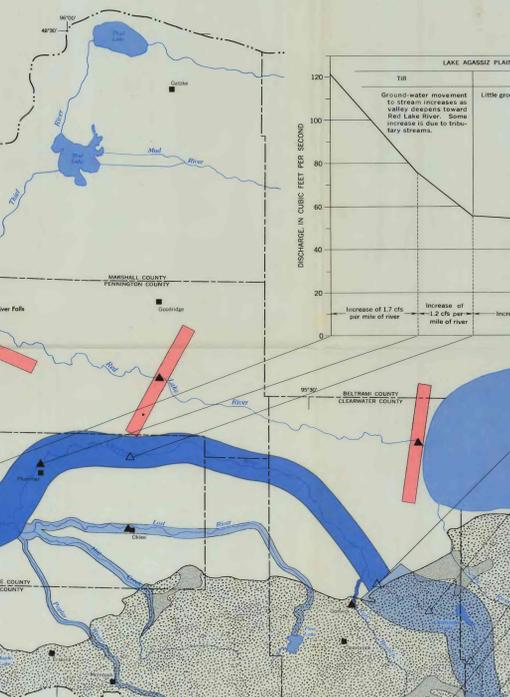
State of river (see physical setting map sheet 1)

Geologic contact

Watershed boundary

Physiographic boundary

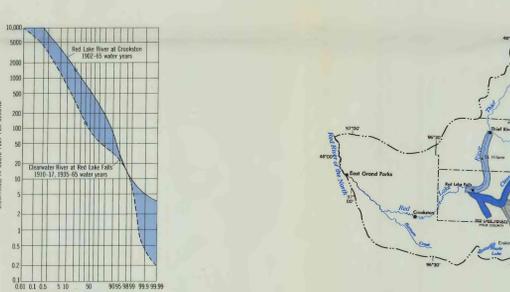
Watershed boundary



LAKE AGASSIZ PLAIN	MORaine	PHYSIOGRAPHY
Till and sand	Till	Geologic material
Ground water movement to stream increases as valley deepens toward Red Lake River. Some increase is due to tributary streams.	Much ground water movement to stream in sand and gravel area.	Relation of ground water to stream as indicated by water-table contours.
Little ground-water movement to stream.	Much ground-water movement to stream.	
Increase of 1.7 cfs per mile of river	Increase of 1.2 cfs per mile of river	Increase of 0.4 cfs per mile of river
Increase of 1.5 cfs per mile of river	Increase of 1.5 cfs per mile of river	Increase of 1.5 cfs per mile of river

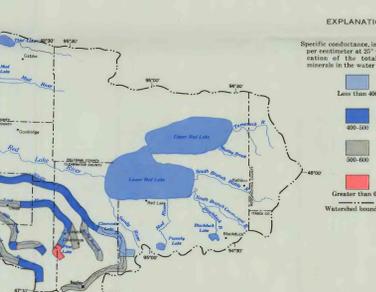
BASE FLOW OF THE CLEARWATER RIVER AND TRIBUTARIES IS DIRECTLY RELATED TO SURFICIAL GEOLOGY AND TOPOGRAPHY

Although the sharp increase in streamflow in the headwaters of the Clearwater River occurs where the stream flows through permeable outwash sand and gravel, the river continues to pick up ground-water discharge, but at a lower rate, throughout the till part of the moraine. The discharge remains lively content as it flows west in the Lake Agassiz plain until another influx of discharge occurs along the deeply incised channel in the lower reach of the river. Similarly, most ground-water discharge into the tributaries of the Clearwater River occurs within the moraine areas with only a little additional flow entering the streams in the Lake Agassiz plain. Although there was no surface runoff from precipitation during the base flow investigation, the flow of the Red Lake River was above the annual normal (at 30 percent on the flow duration curve) because of release of stored water in Upper and Lower Red Lake. The pattern of increase in base flow may change during extended dry periods.



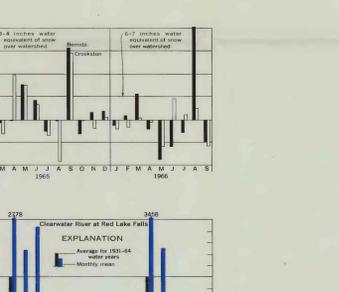
DISCHARGE OF THE CLEARWATER RIVER AT RED LAKE FALLS DURING THE BASE FLOW INVESTIGATION SEPTEMBER 19-22, 1966, AVERAGED 125 CUBIC FEET PER SECOND, WHICH IS 84 PERCENT ON THE FLOW DURATION CURVE.

The average discharge of the Red Lake River at Crookston for the same period was 1,110 cubic feet per second, or 69 percent on the flow duration curve.



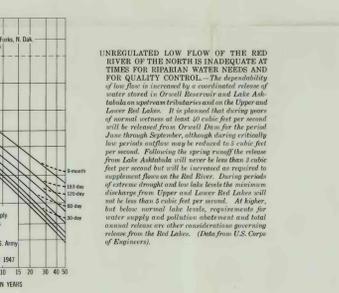
QUALITY OF SURFACE WATER IN THE WATERSHED DURING A TIME OF BASE FLOW IS MODIFIED BY WATER OF LOWER SPECIFIC CONDUCTANCE FROM LAKE STORAGE.

The conductivity is substantially reduced on the Clearwater River and tributaries in reaches below lake outflow. Water quality in streams that do not drain lake areas have higher specific conductance because of the influence of more highly mineralized ground water. The low conductivity of the Red Lake River reflects the large influence of the Red Lakes.



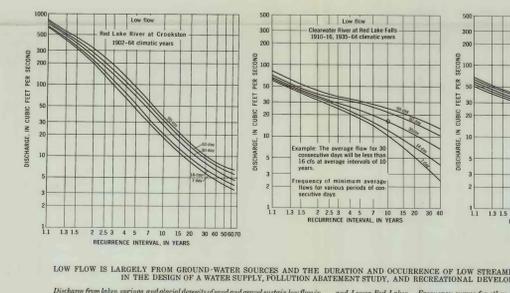
HIGH BASE FLOW IN THE CLEARWATER RIVER DURING THE BASE FLOW INVESTIGATION OF SEPTEMBER 19-22, 1966, WAS CAUSED BY EXCESSIVE PRECIPITATION DURING THE PREVIOUS MONTH OF AUGUST.

By September 1966 the precipitation totalled about 2 inches above the normal for the previous year and a half and prior to the 1966 spring leaving the water table of most of the ground in the upper part of the watershed elevated the normal by about 1 foot. Although discharge in the Clearwater River was above the normal during the spring the discharge precipitation is only normal or less during the summer. The discharge of the Clearwater River is water table specific; it responds rapidly to deficits or excessive precipitation. Therefore, the heavy August rains resulted in above average streamflow (as percent on the flow duration curve).



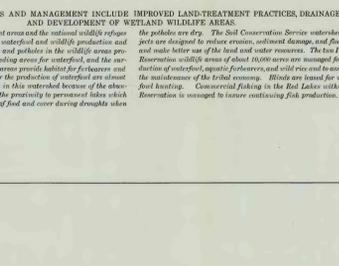
UNREGULATED LOW FLOW OF THE RED RIVER OF THE NORTH IS INADEQUATE AT TIMES FOR IRRIGATION WATER NEEDS AND FOR QUALITY CONTROL.

The dependability of low flow is increased by a coordinated release of water stored in Grand and Lake Agassiz. Although discharge in the Clearwater River is water table specific, it responds rapidly to deficits or excessive precipitation. Therefore, the heavy August rains resulted in above average streamflow (as percent on the flow duration curve).



LOW FLOW IS LARGELY FROM GROUND-WATER SOURCES AND THE DURATION AND OCCURRENCE OF LOW STREAMFLOW ARE CONSIDERED IN THE DESIGN OF A WATER SUPPLY, POLLUTION ABATEMENT STUDY, AND RECREATIONAL DEVELOPMENT.

Discharge from lakes, springs, and glacial deposits of sand and gravel sustains low flow in the Clearwater River. To contrast, the Thief River has long periods of no flow during many years because of a lack of ground-water contribution to the stream channel. The discharge of Red Lake River at Crookston is the integrated flow from streams of both high and low flow and from the regulated discharge from Upper and Lower Red Lakes. Frequency curves for other sites would differ from those shown because of variations in surficial geology. The slope of the low flow frequency curves for sites at Crookston, Red Lake Falls and near Thief River Falls have been affected by the low flow during the drought of the 1960's.



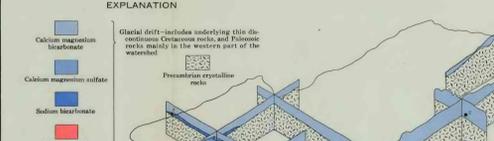
RIVER BASIN PLANS AND MANAGEMENT INCLUDE IMPROVED LAND-TREATMENT PRACTICES, DRAINAGE, AND DEVELOPMENT OF WETLAND WILDLIFE AREAS.

Name	Surface area (acres)	Length of shoreline (miles)	Depth (feet)	Fish and game classification	Sediment transparency (feet)	Total alkalinity (mg/l)	Vegetation	Composition of bottom	Shoreline description	Remarks		
Balm	908	6.59	33	12	Largemouth bass and panfish.	60 (green)	135	Submerged vegetation is sparse. Emergent plants include tules, cattail, and yellow water lily.	Sand	90 percent sand and 10 percent muskgrass.	4 resorts and 14 cottages	
Blackduck	2,749	—	28	—	Walleye and northern pike.	4.0	137	—	Sand and muskgrass, tall reed center and hardwood.	—	17 resorts and 95 cottages.	
Clearwater	977	—	65	—	Bass, panfish and walleye.	10.0	187	Pondweed and smartweed.	Muck.	—	1 resort and 36 cottages.	
Jula	484	5.75	45	—	Walleye.	16.0	137	Water lily, muskgrass, and pondweed.	Sand and muck.	Mixed hardwood and grass.	—	4 resorts and 10 cottages.
Lower Red	157,208	69	35	18	Walleye.	—	152	Open emergent vegetation along shoreline.	Sand and muck.	Flat, sand and bog.	—	Fished commercially by Indians in part of lake in reservation.
Upper Red	107,832	58	20	3	Walleye.	—	—	—	—	Flat bog.	—	—
Maple	1,695	—	13	10	Walleye.	—	237	Pondweed and bulrush.	Sand and gravel near shore and mud toward center.	20 percent open pasture and 70 percent wooded.	—	—
Pine	1,188	7.9	15	8	Northern pike.	6.5	—	Muskgrass and pondweed.	Muck.	Mud and sand, hardwood, marsh vegetation.	—	1 resort and 10 cottages.
Pupokoy	2,141	—	14	5.7	Marginal—some panfish.	Clear to bottom	103	25 percent open areas covered by standing emergent vegetation.	Peat and silt.	Moderate slope fringed with emergent and submerged vegetation.	—	—
Thief	7,100	—	45	3.2	Waterfowl and muskrat.	2.0	—	Pondweed and duckweed. Submerged vegetation over 30 percent of lake.	Soft muck.	Gentle slope, peat, emergent vegetation.	—	Managed for waterfowl production.

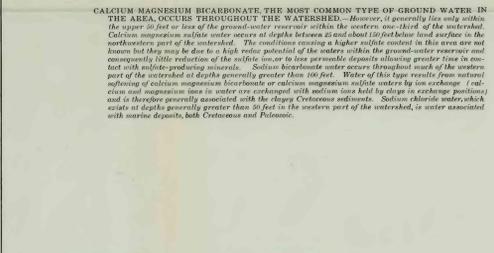
PHYSICAL AND ECOLOGICAL CHARACTERISTICS OF SELECTED LAKES

WATER QUALITY

EXPLANATION

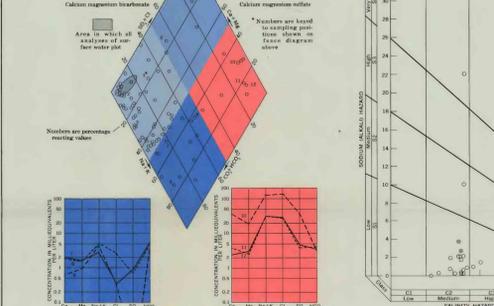


Glacial drift includes underlying this distribution. However, it generally lies only within the upper 30 feet or less of the ground-water reservoir within the western one-third of the watershed. Calcium magnesium bicarbonate water occurs at depths between 20 and about 100 feet below land surface in the northern part of the watershed. The conditions causing a higher sulfate content in this area are not known but they may be due to a high sulfate potential of the water within the ground-water reservoir and consequently little reduction of the sulfate ion or to less permeable deposits allowing greater flow in contact with sulfate-producing minerals. Sulfate bicarbonate water occurs throughout much of the western part of the watershed at depths generally greater than 100 feet. Water of this type results from natural softening of calcium magnesium bicarbonate or calcium magnesium sulfate waters by ion exchange (calcium and magnesium ions in water are exchanged with sodium ions held by clays in exchange position) and is therefore generally associated with the clayey Cretaceous sandstone. Sodium chloride water which exists at depths generally greater than 50 feet in the western part of the watershed, is water associated with marine deposits, both Cretaceous and Paleocene.



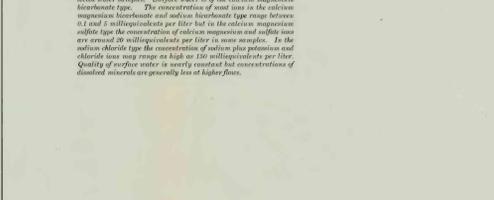
CHEMICAL QUALITY OF GROUND WATER VARIES WIDELY IN THE WATERSHED.

Water is classified into different types of chemical quality depending on the percentage distribution of the major ions in solution. The water type is controlled principally by the solubility of the minerals within the ground-water reservoir, the time water is in contact with these minerals, and the intermingling of water of different types. Some of the sodium bicarbonate water is also softening, which is due to release of calcium and magnesium ions, concentration and an increase in sodium ion concentration. The uppermost part shows the concentration of the water type for selected water samples. Surface water is of the calcium magnesium bicarbonate type. The concentration of most ions in the calcium magnesium bicarbonate and sodium bicarbonate type range between 1 and 2 milligrams per liter but in the calcium magnesium sulfate type the concentration of calcium magnesium and sulfate ions are around 10 milligrams per liter but in some samples. In the sodium chloride type the concentration of sodium plus potassium and chloride ions may range up to 100 milligrams per liter. Quality of surface water is nearly constant but concentrations of dissolved minerals are generally less at higher flows.



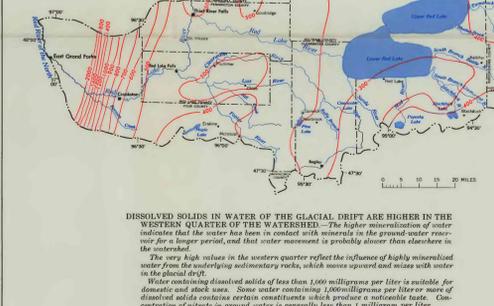
MOST GROUND WATER FOR IRRIGATION USE HAS A LOW SODIUM HAZARD BUT HIGH TO HIGH SALINITY HAZARD.

The sodium bicarbonate water is suitable for irrigation because of high sulfate content. Sodium bicarbonate is very high for the sodium chloride water.



DISSOLVED SOLIDS IN WATER OF THE GLACIAL DRIFT ARE HIGHER IN THE WESTERN QUARTER OF THE WATERSHED.

The higher mineralization of water indicates that the water has been in contact with minerals in the ground-water reservoir for a longer period, and that water movement is probably slower than elsewhere in the watershed. The very high values in the western quarter reflect the influence of highly mineralized water from the underlying Paleocene rocks, which moves upward and mixes with water in the glacial drift. Water containing dissolved solids of less than 1,000 milligrams per liter is suitable for domestic and stock uses. Some water containing 1,000 milligrams per liter more of dissolved solids contains constituents which produce a noticeable taste. Concentration of solids in ground water is generally less than 1 milligram per liter.



EXPLANATION

Line of equal hardness content of water in the glacial drift at depth greater than 20 feet below land surface. Based on 100 milligrams per liter and above.



MOST GROUND WATER IN THE WATERSHED IS VERY HARD—GREATER THAN 100 MILLIGRAMS PER LITER.

Water of extreme hardness occurs in the lake plain. Hardness is due mainly to the presence of calcium and magnesium ions. Mineral salts cause hardness and can form incrustation on well screens. Where hardness exceeds 100 milligrams per liter, incrustation of well screens may be rapid in small diameter wells that are heavily pumped.