

GROUNDWATER RESOURCES

Ground water in the Great Smoky Mountains National Park comes from rain and snow. A part of the precipitation, that falls on the steep mountain slopes, seeps into the ground, and then moves downward through a complex system of openings in the rocks underlying the park and eventually reaches the water table. There is always a reservoir of ground water in the park, and it is constantly being recharged by rain and snow. The water table is usually 10 to 20 feet below the surface of the land.

The rate and amount of water that will yield from the number, size, and interconnection of the openings in the rocks. These factors differ tremendously from place to place. In some places, the water table is only a few hundred feet apart from the surface. In other places, it is 100 to 200 feet below the surface. The amount of water that will yield from a well is determined by the size and interconnection of the openings in the rocks.

The abundance of water-filled openings in the rock depends partly on the type and nature of the rock. In some rocks, the water-filled openings are abundant, and in other rocks they are scarce. The most favorable rocks for developing ground-water supplies are those that have a high percentage of water-filled openings. These rocks are usually of igneous or metamorphic origin, and they are usually of the granitic or gneissic type. The most favorable rocks for developing ground-water supplies are those that have a high percentage of water-filled openings. These rocks are usually of igneous or metamorphic origin, and they are usually of the granitic or gneissic type.

For the purposes of this report, the rock materials of the park can be divided into two broad types, weathered material and unweathered material. The weathered material is that which has been broken down by the action of water and air, and it is usually of a soft, friable nature. The unweathered material is that which has not been broken down, and it is usually of a hard, crystalline nature. The weathered material is usually of a granitic or gneissic type, and the unweathered material is usually of a quartzite or schist type.



Figure 1.—Weathered material, unweathered material, and water table. The water table is shown as a dashed line. The weathered material is shown as a wavy line, and the unweathered material is shown as a solid line. Arrows indicate the downward movement of water from the surface into the ground.

Opening in the rock materials of the Great Smoky Mountains National Park are of two general types, weathered material and unweathered material. The weathered material is that which has been broken down by the action of water and air, and it is usually of a soft, friable nature. The unweathered material is that which has not been broken down, and it is usually of a hard, crystalline nature. The weathered material is usually of a granitic or gneissic type, and the unweathered material is usually of a quartzite or schist type.

When water is present from precipitation, the weathered materials have a high capacity to absorb water. This capacity is enhanced in the upper foot by the presence of pores. Because of addition of water to these pores, the water occurs nearly everywhere in the park, regardless of altitude or slope.

Recharge water moves down into the rock materials until it reaches the zone of saturation, or water table. The water table is the surface of the water in the ground, and it is usually 10 to 20 feet below the surface of the land. The water table is usually 10 to 20 feet below the surface of the land.

Once the recharge water reaches the water table, it is usually of a soft, friable nature. The water table is usually 10 to 20 feet below the surface of the land. The water table is usually 10 to 20 feet below the surface of the land.

The volume of water stored in fractures decreases with depth. The volume of water stored in fractures decreases with depth. The volume of water stored in fractures decreases with depth.

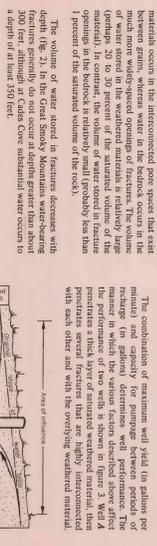


Figure 2.—Water table in fractures. The water table is shown as a dashed line. The fractures are shown as lines of varying thickness and orientation. Arrows indicate the downward movement of water from the surface into the fractures.

When the water table is high, the water table is usually 10 to 20 feet below the surface of the land. The water table is usually 10 to 20 feet below the surface of the land.

FACTORS AFFECTING AVAILABILITY OF GROUND WATER

Weathered material in the Great Smoky Mountains National Park is of two general types, weathered material and unweathered material. The weathered material is that which has been broken down by the action of water and air, and it is usually of a soft, friable nature. The unweathered material is that which has not been broken down, and it is usually of a hard, crystalline nature. The weathered material is usually of a granitic or gneissic type, and the unweathered material is usually of a quartzite or schist type.

Groundwater studies conducted in the mountain and plateau areas of the Great Smoky Mountains National Park have shown that the yield of drilled bedrock wells differs widely in regularity and amount. This is not surprising in view of the irregular nature of the rock materials. The yield of a well is determined by the size and interconnection of the openings in the rocks.

The general physical features affecting the yield of wells are the thickness of the weathered material underlying the well site and the topographic position of the site. The importance of these features is usually determined by the size and interconnection of the openings in the rocks. The yield of a well is determined by the size and interconnection of the openings in the rocks.

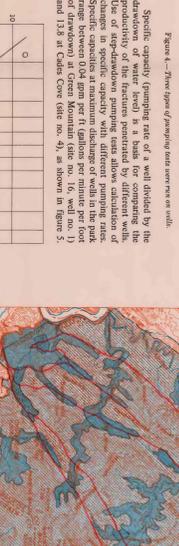


Figure 3.—Specific capacity of pumping wells in the Great Smoky Mountains National Park. The x-axis is labeled 'PERCENT OF SPECIFIC CAPACITY THAT EQUALS OR EXCEEDS NUMBER OF WELLS THAT EQUALS OR EXCEEDS VALUE' and the y-axis is labeled 'SPECIFIC CAPACITY, IN GALLONS PER MINUTE PER FOOT OF DRAINAGE'.

Another factor that may significantly increase well yields is the presence of quartz veins in the bedrock. Quartz veins are usually of a soft, friable nature, and they are usually of a granitic or gneissic type. The presence of quartz veins in the bedrock may significantly increase well yields.

Many park facilities such as the larger campgrounds, are situated on relatively broad ridges. But some smaller areas near ridges. The potential for developing more wells is greater in broad ridges than on narrow valleys. The potential for developing more wells is greater in broad ridges than on narrow valleys.

The rate at which water can be pumped from a well depends on (1) the number and size of the fractures between the well and the water table, and (2) the permeability of the rock materials. The rate at which water can be pumped from a well depends on (1) the number and size of the fractures between the well and the water table, and (2) the permeability of the rock materials.

The yield of a well drilled in the Great Smoky Mountains National Park ranges from less than 1 gpm to as much as 135 gpm (Fig. 7). The yield of a well drilled in the Great Smoky Mountains National Park ranges from less than 1 gpm to as much as 135 gpm.

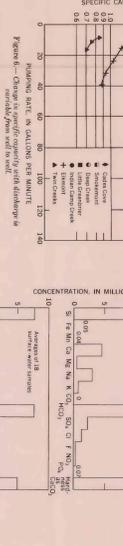


Figure 4.—Pumping rate of pumping wells in the Great Smoky Mountains National Park. The x-axis is labeled 'PERCENT OF WELLS THAT EQUALS OR EXCEEDS VALUE' and the y-axis is labeled 'PUMPING RATE, IN GALLONS PER MINUTE'.

The disintegrable content of both ground water and surface water is low because the rocks of the Great Smoky Mountains National Park are of a hard, crystalline nature. The disintegrable content of both ground water and surface water is low because the rocks of the Great Smoky Mountains National Park are of a hard, crystalline nature.

Water quality in the Great Smoky Mountains National Park is generally good. The water is usually of a soft, friable nature, and it is usually of a granitic or gneissic type. The water quality in the Great Smoky Mountains National Park is generally good.

EXPLANATION

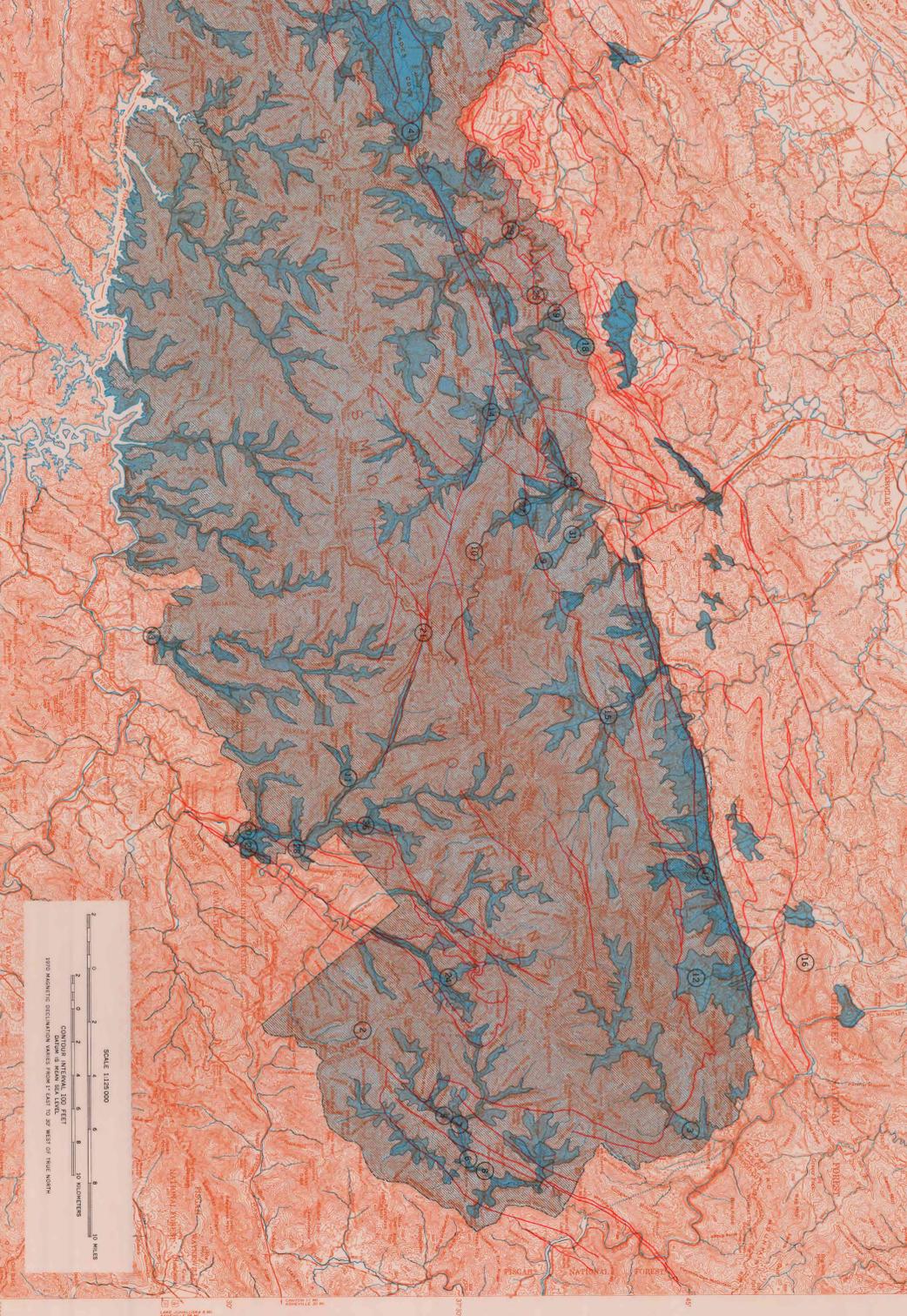


Figure 5.—Ground-water availability map. The map shows the distribution of groundwater availability across the Great Smoky Mountains National Park. The legend indicates the availability levels and other features shown on the map.

Specific capacity (pumping rate of a well divided by the drawdown) is a measure of the productivity of the fractures penetrated by different wells. The specific capacity of a well is a measure of the productivity of the fractures penetrated by different wells.

Use of step-drawdown pumping tests allows calculation of the specific capacity of a well. The specific capacity of a well is a measure of the productivity of the fractures penetrated by different wells.

Specific capacity (pumping rate of a well divided by the drawdown) is a measure of the productivity of the fractures penetrated by different wells. The specific capacity of a well is a measure of the productivity of the fractures penetrated by different wells.

Water samples were collected from several streams for analysis of calcium, magnesium, and total hardness. The water samples were collected from several streams for analysis of calcium, magnesium, and total hardness.

Selected references: A list of references used in the report, including works by King, P. B., and others. Selected references: A list of references used in the report, including works by King, P. B., and others.

Water quality: A section discussing the chemical quality of water in the park, including hardness and disintegrable content. Water quality: A section discussing the chemical quality of water in the park, including hardness and disintegrable content.

Table 1.—Results of first drilling and pumping (Location of each site as shown on ground-water availability map)

| Site No. | Site name and well number | Depth of well (ft) | Thickness of weathered material (ft) | Maximum drawdown (ft) | Specific capacity (gpm per foot) | Water level before pumping (ft below land surface) | Type of rock penetrated by well |
|----------|---------------------------|--------------------|--------------------------------------|-----------------------|----------------------------------|--|---------------------------------|
| 1 | Austin Creek No. 1 | 185 | 10 | 6 | 0.18 | 5 | Phyllite |
| 2 | Big Creek No. 1 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 3 | Big Creek No. 2 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 4 | Big Creek No. 3 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 5 | Big Creek No. 4 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 6 | Big Creek No. 5 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 7 | Big Creek No. 6 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 8 | Big Creek No. 7 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 9 | Big Creek No. 8 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 10 | Big Creek No. 9 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 11 | Big Creek No. 10 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 12 | Big Creek No. 11 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 13 | Big Creek No. 12 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 14 | Big Creek No. 13 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 15 | Big Creek No. 14 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 16 | Big Creek No. 15 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 17 | Big Creek No. 16 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 18 | Big Creek No. 17 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 19 | Big Creek No. 18 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 20 | Big Creek No. 19 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 21 | Big Creek No. 20 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 22 | Big Creek No. 21 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 23 | Big Creek No. 22 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 24 | Big Creek No. 23 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 25 | Big Creek No. 24 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 26 | Big Creek No. 25 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 27 | Big Creek No. 26 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 28 | Big Creek No. 27 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 29 | Big Creek No. 28 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 30 | Big Creek No. 29 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |
| 31 | Big Creek No. 30 | 200 | 10 | 6 | 0.18 | 5 | Phyllite |

Table 2.—Results of chemical analysis of ground-water samples (Concentrations in milligrams per liter; specific conductance in micromhos at 25°C; Fe and Mn in solution when analyzed. Analyzed by U.S. Geological Survey)

| Site No. | Site name | Date of collection | SO ₄ | Ca | Mg | Na | K | CO ₂ | HCO ₃ | SO ₂ | Cl | F | NO ₃ | Pb | Dis-solved solids (mg per liter) | Specific conductance (micromhos per centimeter) | Temperature (°C) |
|----------|--------------------|--------------------|-----------------|-----|------|------|------|-----------------|------------------|-----------------|------|------|-----------------|------|----------------------------------|---|------------------|
| 1 | Austin Creek No. 1 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 2 | Big Creek No. 1 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 3 | Big Creek No. 2 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 4 | Big Creek No. 3 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 5 | Big Creek No. 4 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 6 | Big Creek No. 5 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 7 | Big Creek No. 6 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 8 | Big Creek No. 7 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 9 | Big Creek No. 8 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 10 | Big Creek No. 9 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 11 | Big Creek No. 10 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 12 | Big Creek No. 11 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 13 | Big Creek No. 12 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 14 | Big Creek No. 13 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 15 | Big Creek No. 14 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 16 | Big Creek No. 15 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 17 | Big Creek No. 16 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 18 | Big Creek No. 17 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 19 | Big Creek No. 18 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 20 | Big Creek No. 19 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 21 | Big Creek No. 20 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 22 | Big Creek No. 21 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 23 | Big Creek No. 22 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 24 | Big Creek No. 23 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 25 | Big Creek No. 24 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 26 | Big Creek No. 25 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 27 | Big Creek No. 26 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 28 | Big Creek No. 27 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 29 | Big Creek No. 28 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 30 | Big Creek No. 29 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 31 | Big Creek No. 30 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |

Table 3.—Results of chemical analysis of ground-water samples (Concentrations in milligrams per liter; specific conductance in micromhos at 25°C; Fe and Mn in solution when analyzed. Analyzed by U.S. Geological Survey)

| Site No. | Site name | Date of collection | SO ₄ | Ca | Mg | Na | K | CO ₂ | HCO ₃ | SO ₂ | Cl | F | NO ₃ | Pb | Dis-solved solids (mg per liter) | Specific conductance (micromhos per centimeter) | Temperature (°C) |
|----------|--------------------|--------------------|-----------------|-----|------|------|------|-----------------|------------------|-----------------|------|------|-----------------|------|----------------------------------|---|------------------|
| 1 | Austin Creek No. 1 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 0.01 | 22 | 18 | 14 |
| 2 | Big Creek No. 1 | 6-28-48 | 13 | 1.0 | 0.01 | 0.01 | 0.01 | 0.01 | | | | | | | | | |