



The availability map should be used with the understanding that the estimated yields are for fully penetrating, properly screened and developed wells of adequate diameter. The map is intended as a general guide in the location of major aquifers, not as a map to locate specific wells. Few, if any, aquifers are so uniform in their water-bearing properties that production wells may be drilled in them without preliminary exploratory drilling. If the map is used with this understanding of its limitations, it should serve as a useful tool in the future development of the ground-water resources of Walsh County.

GLACIAL-DRIFT AQUIFERS

Walsh County is almost entirely covered by glacial drift, which may be subdivided into two distinct types. The most common type is till, a nonsorted mixture of clay, silt, sand, gravel, and pebbles. Till was mainly deposited by active glaciers. It has a low permeability, and will normally yield only small quantities of ground water to wells. Glaciofluvial deposits, the least common type, are stratified deposits of sand and gravel that are sorted according to grain size. These materials were deposited by moving water, are normally quite permeable, and form the principal aquifers in Walsh County.

The glacial drift generally ranges in thickness from about 300 feet in the eastern part of the county to about 30 feet in the western part. Locally, along streams in the western part of the county, the glacial drift has been removed by erosion and the underlying bedrock is exposed. Water levels in the glacial drift and other surficial deposits of Walsh County range from an elevation of about 775 feet in the northeastern part of the county to about 1,610 feet at the ground-water divide near Fairdale (see water-table map).

The largest and most productive glacial-drift aquifer in Walsh County is located near the city of Fordville, and is locally known as the Fordville aquifer. The aquifer is part of the Elk Valley delta (Upham, 1895, p. 27) deposits and extends from T. 156 N., R. 56 W. south to the county line. It underlies an area of about 33 square miles. Water-level fluctuations in the Fordville aquifer from August 1967 to December 1969 are shown by the hydrographs for wells M and N. Water-level rises in well M, located about 100 feet from the North Branch of the Forest River near the north end of the Fordville aquifer, are mainly the result of recharge from snowmelt and spring rains. The hydrograph of well N shows that only small water-level fluctuations occur near the discharge area at Fordville.

Several less productive aquifers in Walsh County are associated with various glacial features. The linear beach-ridge aquifers near Pisk and east of Edinburg consist of silt and sand deposits of glacial Lake Agassiz. Test drilling has shown that at these and many other locations the beach deposits have, near the base, a thin water-saturated zone that will supply sufficient water for farm use and domestic supply. Short-term yields from these deposits may be more than 50 gpm (gallons per minute), but long-term yields may be much less. The water-level fluctuations in a beach ridge near Pisk are shown by the hydrograph for well O.

BEDROCK AQUIFERS

Four bedrock units supply water to wells in Walsh County. The areal extent and yields of these units are indicated on the ground-water availability map.

The Pierre Shale underlies the western third of the county, and yields from 1 to 5 gpm from fracture zones in the upper part of the formation.

The Dakota Sandstone underlies all but the northeast corner of the county. Wells tapping the Dakota Sandstone range in depth from about 160 feet in the eastern part of the area to greater than 1,000 feet in the western part. This formation has been reported to yield as much as 500 gpm in the northeast corner of the county (Paulson, 1970), which borders Walsh County on the north. Test drilling indicates that potential yields from the Dakota Sandstone in Walsh County would be quite variable due to the presence of interbedded shale lenses and thinning of the formation toward the east. However, it appears that moderate to large yields would be available at least locally in many parts of the county. Many of the wells developed in the Dakota Sandstone in eastern Walsh County flow at land surface. Aerial photographs and maps of the area are being used, along with well location and construction, to water from the Dakota Sandstone is saline and is generally unsatisfactory for domestic and industrial uses.

The Jurassic(?) and Paleozoic bedrock formations, consisting of siltstone, sandstone, and limestone, underlie the entire county and are directly beneath the glacial deposits in the northeastern part. The limited quantitative hydrologic data available concerning these units indicate that moderate to large yields may be possible at selected sites where fractured or cavernous limestone is present. Water from these units is very saline and is unsatisfactory for most uses.

CHEMICAL QUALITY OF WATER

Wells tapping glacial-drift aquifers in Walsh County yield water with a wide range of chemical quality, as shown by the chemical-analysis symbols on the water-availability map and by the table of selected chemical analyses (analyses B, D, F, I, and J).

Water from the sand and gravel deposits of the Fordville aquifer and from the deposits forming the beach ridges is normally low in dissolved solids (analyses B and I). The chemical quality of water in buried sand and gravel deposits is dependent, to a degree, upon their connection with the underlying bedrock and with the overlying surficial units. Where these buried deposits receive recharge from the surface and have only a limited connection with the bedrock, dissolved solids will be low, as shown by analysis J.

Where the buried sand and gravel deposits have a good hydraulic connection with the underlying bedrock, and ground-water movement is from the bedrock into the sand and gravel deposits, the amount of dissolved solids in the water is dependent upon the degree of mineralization of the water in the bedrock. This effect is illustrated by analyses D and F, where highly mineralized water from the underlying rocks of Jurassic(?) and Paleozoic ages has migrated into the buried sand and gravel deposits.

Water from the Dakota Sandstone in Walsh County (analyses A, C, and H) ranged in dissolved-solids content from 3,420 mg/l (milligrams per liter) to 4,900 mg/l; with a mean value for 23 analyses of 4,500 mg/l.

Water quality from the Pierre Shale (analyses K and L) varies widely within the county depending upon depth and location of the well. The wide range in water quality is due, apparently, to differences in sources of recharge and degree of fracturing within the shale. The mean value of dissolved solids for nine analyses of water from the Pierre Shale was 3,500 mg/l, with a range from 1,820 to 5,110 mg/l.

The most highly mineralized water in Walsh County is produced from wells tapping the rocks of Jurassic(?) and Paleozoic ages. At the present time, little use is made of this water due to the poor quality. Dissolved solids ranged from 10,500 mg/l (analysis G) to 23,900 mg/l (analysis E).

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Well	Depth	Aquifer	Date of collection	Temperature (°C)	Silica (SiO ₂)	Total iron (Fe)	Calcium (Ca)	Magnesium (Mg)	Sodium (Na)	Potassium (K)	Bicarbonate (HCO ₃)	Carbonate (CO ₃)	Sulfate (SO ₄)	Chloride (Cl)	Fluoride (F)	Nitrate (NO ₃)	Boron (B)	Dissolved solids		Hardness as CaCO ₃		Percent sodium	Specific conductance (micro-mhos at 25°C)	pH	Sodium adsorption (SAR)
																		Residue on evaporation at 180°C	Sum	Calcium meq/l	Non-carbonate				
A	1,042	K1PM	8-18-67	_____	7.0	3.3	21	9.1	1,700	15	676	0	1,010	1,610	2.0	17	4.6	4,900	4,730	90	0	97	7,610	8.2	78
B	35	QG03	5-24-68	6	24	1.8	78	21	19	3.8	297	0	83	6.4	2	0	39	401	384	283	39	13	597	8.2	5
C	295	K1PM	11-2-67	7	6.8	3.4	49	24	1,370	18	642	0	1,200	1,030	2.5	25	4.1	4,060	4,050	220	0	92	6,210	8.2	40
D	140	QG51	7-15-68	8	19	19	584	234	3,840	94	293	0	1,960	6,280	2.5	0	6.1	13,800	13,200	2,420	2,180	77	19,000	7.9	34
E	300	E	8-25-69	_____	18	11	1,390	583	6,180	55	85	0	1,650	12,800	4	0	1.9	23,900	22,700	5,870	5,880	69	34,000	6.39	35
F	203	QG1	6-21-69	7	24	7.4	366	138	3,680	51	361	0	1,390	5,730	7	0	4.1	12,300	11,600	1,880	1,880	84	18,900	7.4	42
G	322	E	6-14-68	8	14	5.4	245	114	3,610	44	475	0	1,020	5,210	1.9	17	5.0	10,500	10,500	1,080	690	87	16,000	7.3	48
H	246	K1PM	4-4-68	6	9.6	4.5	15	1,190	155	559	19	325	1,350	2	34	2.9	3420	3,290	1,760	0	93	5,410	8.4	39	
I	18	QG02	11-20-67	_____	26	82	88	22	12	2.9	301	0	64	5.3	3	17	44	376	387	309	62	8	598	8.0	3
J	103	QG51	8-21-69	_____	27	3.0	173	51	46	7.6	339	0	444	15	5	1	15	896	935	641	363	13	1,280	7.7	8
K	107	K3PD	7-17-68	6	22	12	24	6.6	608	9.1	809	10	218	390	2	2.5	3.6	1,820	1,690	87	0	93	2,720	8.3	28
L	95	K3PD	8-27-68	6	27	0	224	80	1,320	22	710	0	2,630	390	2.1	18	3.8	5,110	5,070	890	307	76	6,440	7.9	19

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