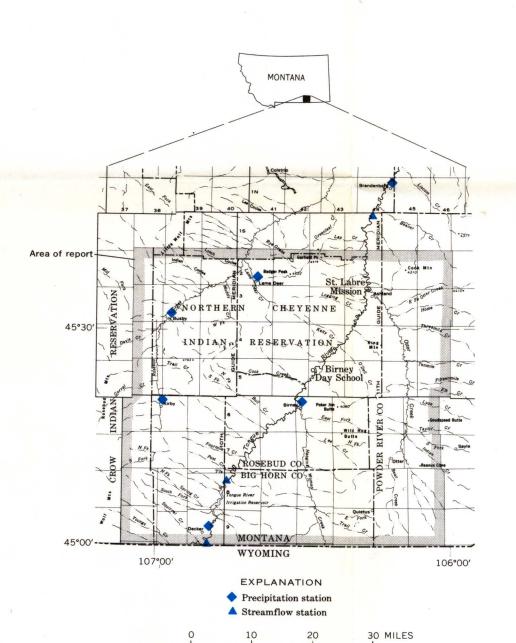


PURPOSE OF INVESTIGATION This investigation was made to evaluate the water resources of the Northern Cheyenne Indian Reservation and adjacent area in terms of sources, amounts available, and quality. Specific questions considered were (1) how are the water resources being used, (2) which rock formations yield water to wells and springs, (3) how much water is available from individual wells and springs and what is its quality, (4) what is the flow of the major streams, and what is the chemical quality of the surface water.

WATER USE Present water use is small as only about 2,500 people live on the reservation, and approximately 1,000 more live on ranches in the rest of the area. Busby, Lame Deer, Birney, Birney Day School, and St. Labre Mission use wells and community water systems. Private wells and springs supply domestic water to the rest of the area. Stock water is obtained from wells, springs, streams, and ponds. Extensive coal deposits in the area are being considered for industrial development. Abundant water is essential for conversion of coal to electricity, fluid fuel, or chemicals. Development of the coal deposits would significantly increase the use of water in the area.

GROUND-WATER RESOURCES

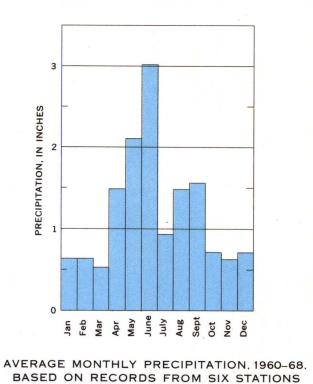
The Northern Cheyenne Indian Reservation and adjacent area (see index map) includes about 2,500 square miles of the unglaciated Missouri Plateau part of the Great Plains province in southeastern Montana (Fenneman, 1931, p. 63-65). The topography is rough, with long narrow ridges, extensive hillsides, and generally narrow flood plains.



INDEX MAP OF REPORT AREA SHOWING LOCATION

OF PRECIPITATION AND STREAMFLOW STATIONS

Precipitation provides soil moisture and recharge to ground water, and contributes to streamflow. Wells and springs depend on recharge from rainfall or snowmelt that temporarily exceeds soil moisture requirements. The average annual precipitation from 1960 through 1968 as measured at six stations in and near the study area was 14.47 inches (see



Exposed rocks include soil and weathered clinker, alluvial deposits of sand and gravel, the Wasatch Formation, and the Tongue River Member of the Fort Union Formation (see summary table of geologic formations and their water-bearing properties). The soil and weathered clinker, referred to as "mantle," cover most of the ridges and the gently sloping hillsides. Because of their position, mantle deposits normally do not contain water in quantities sufficient to satisfy stock Alluvial deposits of sand and gravel underlie the flood plains. In places, this alluvium is sufficiently permeable to

yield relatively large volumes of water to wells (see geohydrologic map). Test holes drilled in secs. 9 and 10, T. 3 S., R. 44 E., (see map of Ashland area) show the maximum thickness of the alluvium beneath the Tongue River flood plain to be 97 feet and the maximum saturated thickness about 60 feet. The alluvium in this area, as shown on the geologic section, is predominantly sand- and gravel-sized clinker fragments derived from the Tongue River Member of the Fort Union Formation. The Wasatch Formation overlies the Tongue River Member of the Fort Union Formation on the ridges and uplands along

the southern border of the project area. Of the Wasatch in

the project area, Baker (1929, p. 25) noted: "There is ***no

great lithologic difference between the Tongue River and Wasatch rocks *** " The Tongue River Member of the Fort Union Formation contains beds of sandstone, shale, shaly sandstone, and coal. In describing the member, Bass (1932, p. 37) commented: "In detail the beds in the Tongue River member are not continuous over wide areas, but groups of beds maintain certain distinguishing characters throughout distances of 50

miles and more." The lithology of the Tongue River Member is indicated by the graphic logs of water wells shown on the geohydrologic map. Fractured "clinker" beds at the surface of the Wasatch Formation and Tongue River Member permit rapid infiltration of precipitation, which recharges underlying strata or emerges at springs or seeps. Bass (1932, p. 27) commented: "As thick coal beds ***are uncovered by

erosion they burn along their outcrops and for considerable

distances back from the crop lines, and the great heat produced ***bakes the [overlying] beds to a very resistant slag ***called clinker. The resulting clinker beds ***commonly are 20 to 40 feet thick." The Hell Creek Formation, which does not crop out in the area, yields usable water to wells. Water from the Hell Creek is under sufficient artesian pressure to flow from wells in the

valleys of Otter Creek and Tongue River. Other water-bearing formations that lie below the Hell Creek and that have been tested, contain water that is too saline to be used for most purposes. The Fox Hills Sandstone, which is an important aquifer in other places in southeastern Montana, is absent in OCCURRENCE OF GROUND WATER Wells and springs yield water for domestic or stock

supplies from the alluvium in stream valleys; from clinker beds, sandstone and coal beds in the Tongue River Member of the Fort Union Formation; and from sandstone beds in the Hell Creek Formation (see table and geohydrologic map). Wells that would yield more than 50 gpm (gallons per minute) would be limited to the alluvium along the perennial streams. A well in the SW¼NW¼ of sec. 21, T. 5 S., R. 42 E., reportedly yields 700 gpm from gravel along Cook Creek. Extensive saturated clinker beds would yield as much as 50 gpm to wells. Many wells and springs that yield 1 gpm or less from alluvium in the coulees have been developed for stock Shale beds in the Tongue River Member of the Fort Union

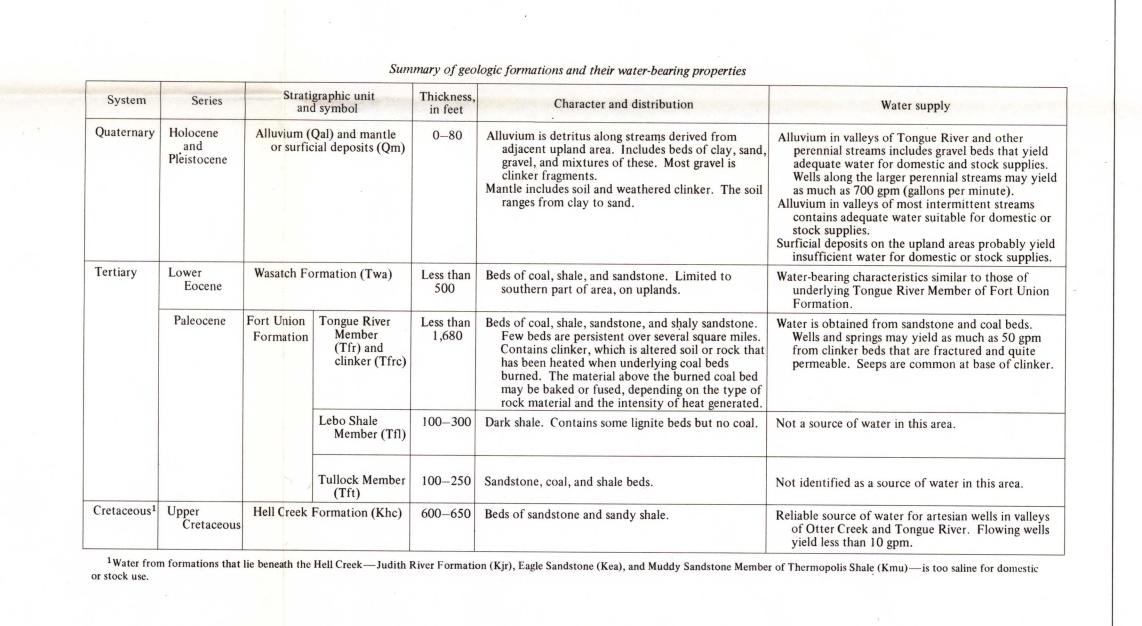
Formation restrict the downward percolation of water. Most springs in the area issue from hillsides at the contact between a water-bearing bed of sandstone or coal and an underlying shale bed. Springs are also found at or near the base of The potentiometric surface contours (see geohydrologic map) show the altitude of ground water in the alluvium and in the Tongue River Member of the Fort Union Formation. The contours are based on altitude of water in wells and the

altitude at which springs and seeps emerge along hillsides and

in valley bottoms. Where springs emerge at the top of a shale

bed, the rocks directly beneath the shale may not be water

bearing, and the water above the shale may be limited to a



WATER RESOURCES OF THE NORTHERN CHEYENNE INDIAN RESERVATION AND ADJACENT AREA, SOUTHEASTERN MONTANA

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