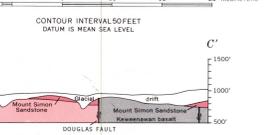
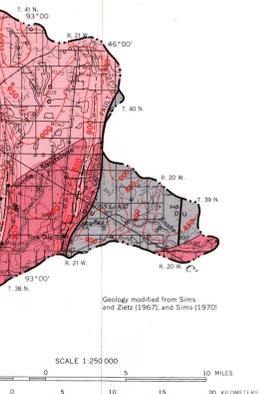


GEOLOGY AND GROUND WATER

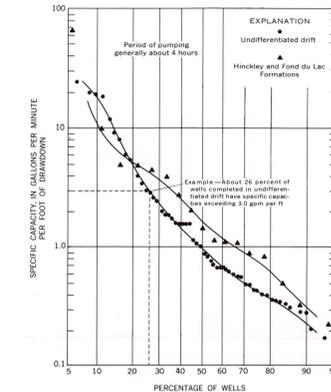
AGE	GEOLOGIC UNIT	WATER-YIELDING CAPABILITY
CAMBRIAN	Mount Simon Sandstone	Known to yield small supply; largely undefined in watershed but well is believed to be thin and discontinuous
	Hinckley and Fond du Lac Formations of Winchel (1886 and 1895) (mainly sandstone)	Commonly yields 300 gpm to single well; yield is dependent on texture, concentration and amount of aquifer open to well. Maximum capability unknown
PRECAMBRIAN	Keweenaw basalt	Often yields small supply. Amount of water obtainable is dependent on degree of jointing and fracturing
	McGrath Gneiss of Woytki (1949)	May yield small supply from joints and fractures. Often unreliable source
	Warman Quartz Monzonite of Woytki (1949)	May yield small supply from joints and fractures. Often unreliable source



DIVERSE BEDROCK TYPES, WHOSE SURFACES SLOPE GENERALLY SOUTHEASTWARD, UNDERLIE THE SNAKE RIVER WATERSHED

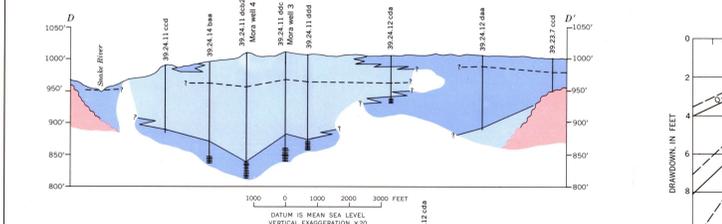
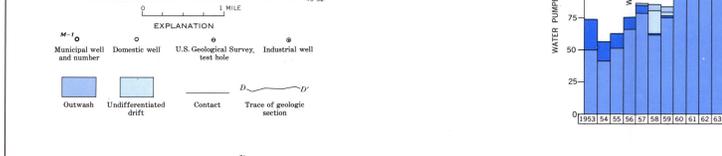
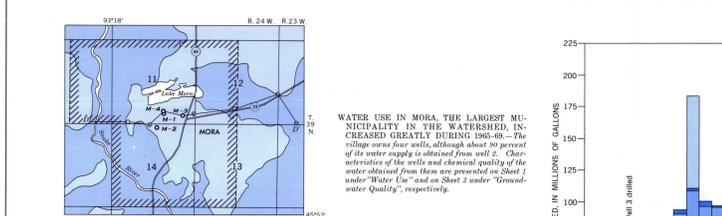
Granitic rocks, in the northwestern half of the watershed, are the Warman Quartz Monzonite and the McGrath Gneiss of Woytki (1949). They are part of a regional bedrock high that extends beyond the watershed boundary. Numerous small outcrops are found throughout this area. The Hinckley and Fond du Lac Formations of Winchel (1886 and 1895), and (revised in this area, underlie most of the southeastern half of the watershed. The Hinckley Formation is principally a medium- to coarse-grained, poorly sorted, quartz sandstone that dips southeast at about 10-15 degrees (Atwater and Clement, 1935). Underlying the Hinckley are coarse-grained arkosic sandstones and conglomerates of the Fond du Lac Formation. Interbedded shales occur in each of the above units. The eastern extent of the Hinckley-Fond du Lac unit is limited by the Douglas fault. Although drill-

hole data verifies 100-500 feet of sandstone near the fault, it has been estimated (Welch, 1941) that several thousand feet may be present. Keweenaw basalt, with interbedded sediments, occurs east of the Douglas fault. Sandstone of possible Cambrian age has been reported in a few wells along the southern boundary of the watershed and at an outcrop near the mouth of the Snake River (Thiel, 1947). Present well control does not permit accurate delineation of its extent. The relatively low resistance to erosion of the Hinckley and Fond du Lac Formations is an important factor in the shaping of bedrock topography. The surface of these units is generally lower than that of the more resistant granitic and basaltic rocks. Several bedrock valleys, probably fashioned by glacial melt waters, occur principally in the sandstone. Present well control is sufficient for only fractional delineation of the valleys.



SPECIFIC CAPACITIES INDICATE A SIMILARITY IN THE WATER-YIELDING CAPABILITIES OF THE HINCKLEY AND FOND DU LAC FORMATIONS AND THE GLACIAL DRIFT.—Most water from the Hinckley and Fond du Lac Formations is obtained from wells open to poorly cemented sandstone zones. Poorly cemented zones, being more permeable than well-cemented zones, have much higher water-yielding capabilities. The thickness of the Hinckley and Fond du Lac Formations enhances the possibility of obtaining large yields. The heterogeneity of glacial drift accounts for considerable variation in its water-yielding capabilities. Large yields are possible where thick, permeable sand and gravel occurs within the drift. Specific capacity values for the granitic and basaltic rocks are generally less than 1.0 indicating low water-yielding capabilities. Porosity in these rocks is limited to fractures and joints. Because openings are usually poorly interconnected, recharge to wells is slow.

GROUND-WATER DEVELOPMENT AT MORA



LARGE-YIELD WELLS IN MORA ARE COMPLETED IN A BURIED SAND AND GRAVEL AQUIFER WITHIN A BEDROCK VALLEY. The geologic map and section show that the aquifer is overlain by undifferentiated drift of varying thickness. The aquifer is classified as leaky artesian. It is possible that the buried sand and gravel is stratigraphically and hydrologically continuous with surficial outwash to the east.

THE THICKNESS OF GLACIAL DRIFT IN THE WATERSHED IS RELATED TO THE UNDERLYING BEDROCK TOPOGRAPHY

Drift is commonly thinnest where it overlies bedrock highs and thickest over bedrock lows. The thickest drift occurs as long, narrow bands associated with bedrock valleys. Part of one such valley occurs immediately west of the Douglas fault. Its location may have been partially influenced by the presence of a preglacial fault escarpment. Glacial fill in another bedrock valley is the source of Mora's municipal water supply and is discussed in greater detail below.

EXPLANATION
100
Line of equal drift thickness interval, 50 feet
Bedrock outcrop
Drill hole, bottom in bedrock
Drill hole, bottom in drift
Approximate fault trace
U, upthrown side; D, downthrown side
Watershed boundary

EVALUATION OF SURFICIAL OUTWASH FOR WELL YIELDS



WHERE SATURATED THICKNESS IS SUFFICIENT SURFICIAL OUTWASH IS CAPABLE OF YIELDING IN EXCESS OF 500 GALLONS PER MINUTE TO LARGE DIAMETER WELLS.—Test drilling in outwash areas provided information as to textural variations and saturated thickness of the outwash. On the basis of assigned permeabilities and the determined saturated thicknesses, transmissivity values were estimated for each test hole. Estimates of maximum yields to wells were made assuming: 1) diameter of the pumped well is 1 foot; 2) the well is open to the full saturated thickness of the aquifer; 3) drawdown in the pumped well is two-thirds the original saturated thickness; 4) the period of pumping is 1 day; and 5) a uniform storage coefficient of 0.15 applies. Corrections for drawdown were not made. Locally, within the most productive areas, yields of as much as 1,000 gallons per minute are believed possible. Relating estimated yields to saturated thickness indicates the importance of having a large saturated thickness if a large yield is required. Rapid lateral and vertical natural changes over short distances, characteristic of glacial deposits, make the extension of yield production beyond actual test holes uncertain. The above interpretation indicates only the general magnitude of yields that might be obtained from wells completed in the surficial outwash.

AN AQUIFER TEST USING MORA VILLAGE WELLS INDICATES THAT THE WATER-YIELDING CAPABILITY OF THE AQUIFER IS VERY GOOD.—Transmissivity (T) of the aquifer was determined to be about 1.1×10^{-4} ft² per day and storage coefficient (S) about 0.15. The above yields of distance-drawdown curves can be used as an aid in planning well spacing and in planning for optimum development of the aquifer. Assumed well discharges (Q) and periods of pumping (t) are compatible with local pumping practices.

WATER RESOURCES OF THE SNAKE RIVER WATERSHED, EAST-CENTRAL MINNESOTA

By
G. F. Lindholm, J. O. Helgesen, W. L. Broussard, and D. W. Ericson
1974