

### INTRODUCTION

With the exception of the St. Louis area, south-central Missouri is primarily an agricultural and recreational area. The more productive farms, which produce chiefly livestock products and feed crops, are located on the broad, rolling uplands and along the narrow stream valleys of the region. The many large springs, clear streams, scenic caves, and rugged uplands offer numerous opportunities for outdoor activities such as camping, fishing, hunting, and hiking. Public access to the area is provided by Clark National Forest (800,000 acres or 243,000 ha), Mark Twain National Forest (600,000 acres or 193,000 ha), four U.S. Army Corps of Engineers reservoirs, and several State parks. In addition, the eastern part of the area has developed into a leading producer of lead, zinc, and iron ore.

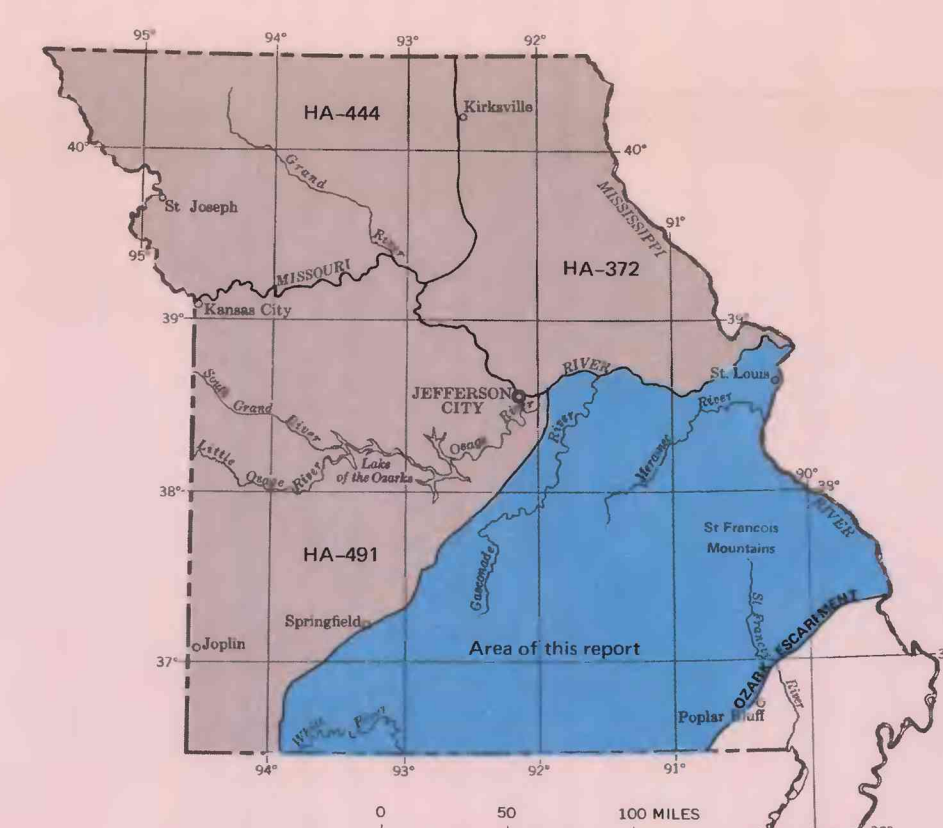
Although the population of several rural counties in the area have declined in recent years, significant population increases have occurred in the vicinity of the two principal population centers, St. Louis in the northwest and Springfield in the southwest. Population increases have also occurred along the interstate highway connecting these two urban areas, in the vicinity of the Corps of Engineers reservoirs, and in the mining area. During 1960-70, for example, the population of Jefferson, St. Louis, and Tynes counties increased by 58.2, 35.3, and 27.2 percent, respectively. Future population increases are expected to occur as a result of continued urban expansion, increased recreational use of land and water resources, and additional development of the mining industry.

Knowledge of the water resources of south-central Missouri is necessary in planning the use, development, and preservation of the area's resources. The purpose of this atlas is to present a general summary of information concerning the occurrence, availability, use, and quality of water in the area and to offer alternative considerations, and information sources for use in developing the area's water resources.

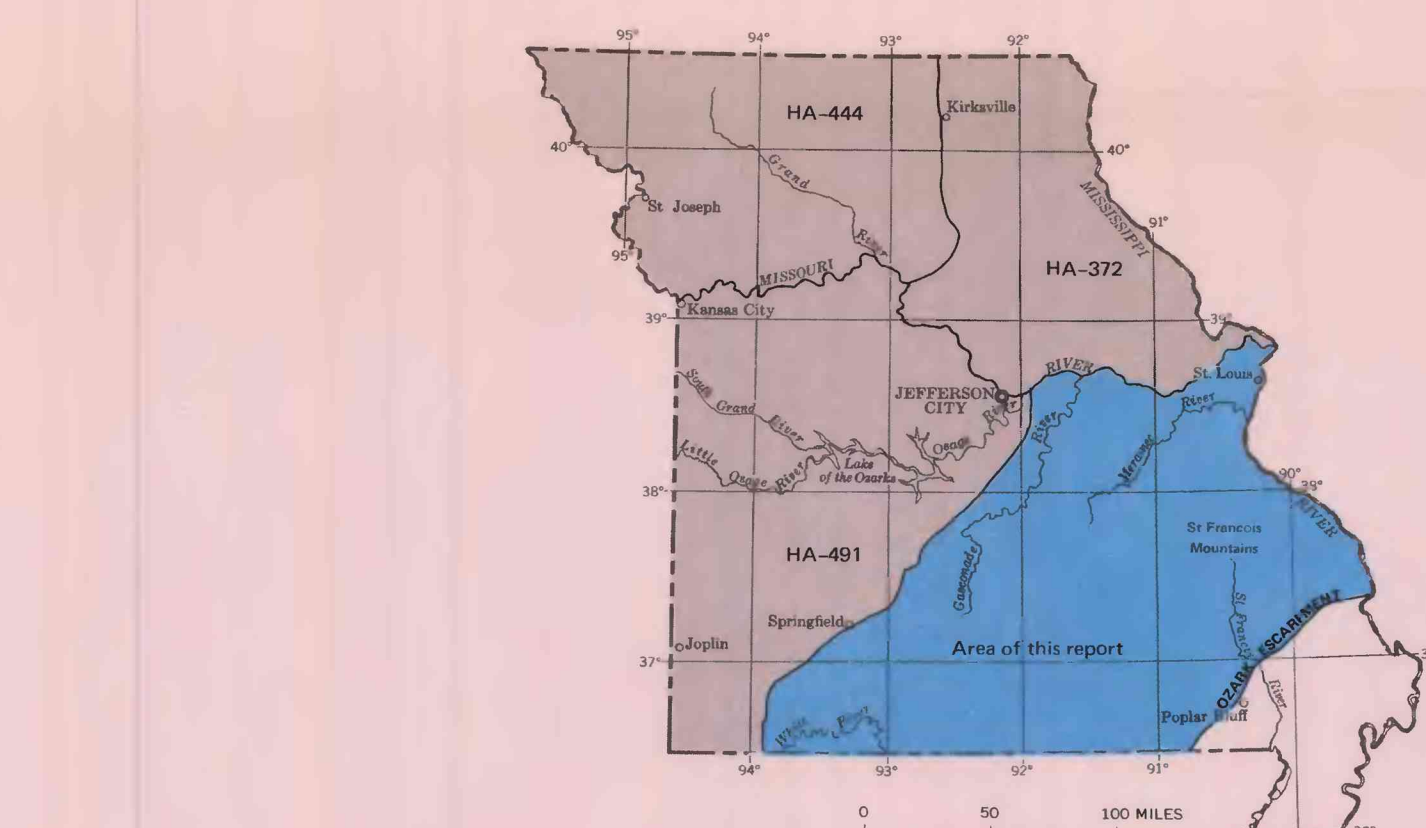
This atlas describes an area of approximately 23,000 mi<sup>2</sup> (59,600 km<sup>2</sup>) and includes all or parts of 28 counties. The area is bounded on the north by the southern edge of the Missouri River flood plain, on the east by the Mississippi River and the Flatwoods-Lowlands boundary (Clark Escarpment), on the south by the Missouri-Arkansas State line, and on the west by the western drainage divide of the Ozark and White River basins. The alluvial valley of the Missouri River is excluded from the present study. Water resources of the Missouri River valley are described in Emmett and Jeffrey (1969) and Gann and others (1971).

This investigation was made in cooperation with the Missouri Geological Survey Department of Natural Resources, (formerly Missouri Geological Survey and Water Resources). Special appreciation is expressed to the many individuals, public utilities, private industry, and representatives of other State and Federal agencies who furnished information and assistance for the study. The stratigraphic nomenclature used in this atlas follows that of the U.S. Geological Survey and varies from the current usage of the Missouri Geological Survey.

A table of conversion factors is included on this sheet for the benefit of those users desiring to convert the English units published herein to metric units.



### PHYSICAL SETTING



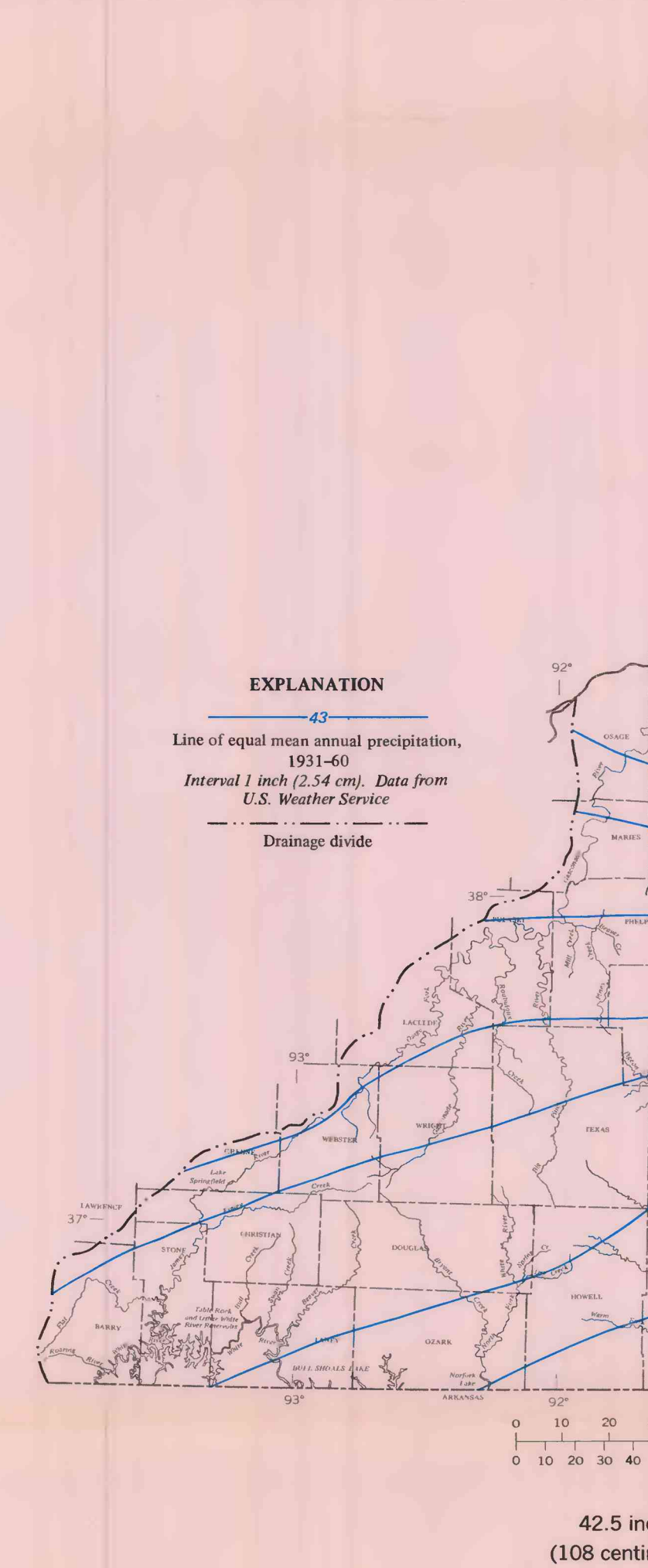
Map number	Lakes and reservoirs	Storage capacity (acre-feet)
1	Bethland	1,400,000
2	Atchigee	700,000
3	Meramec Park	1,000,000
4	LSE	30,000
5	Union	477,300
6	Current Bluff	33,000
7	Insudale	161,000
8	Pine Ford	260,000
9	Wappapello	613,000
10	Clearwater	613,000
11	Table Rock	3,602,000
12	Taneytown	30,300
13	Ball Shoals	5,408,000
14	Norfork	1,800,000
15	Water Valley	1,800,000

Map number	Active watershed projects	Drainage area (mi <sup>2</sup> )
16	Fountain Creek	156
17	Little Black	225
18	Pike Creek	145
19	Boone Spring	55
20	Fountain Creek	71

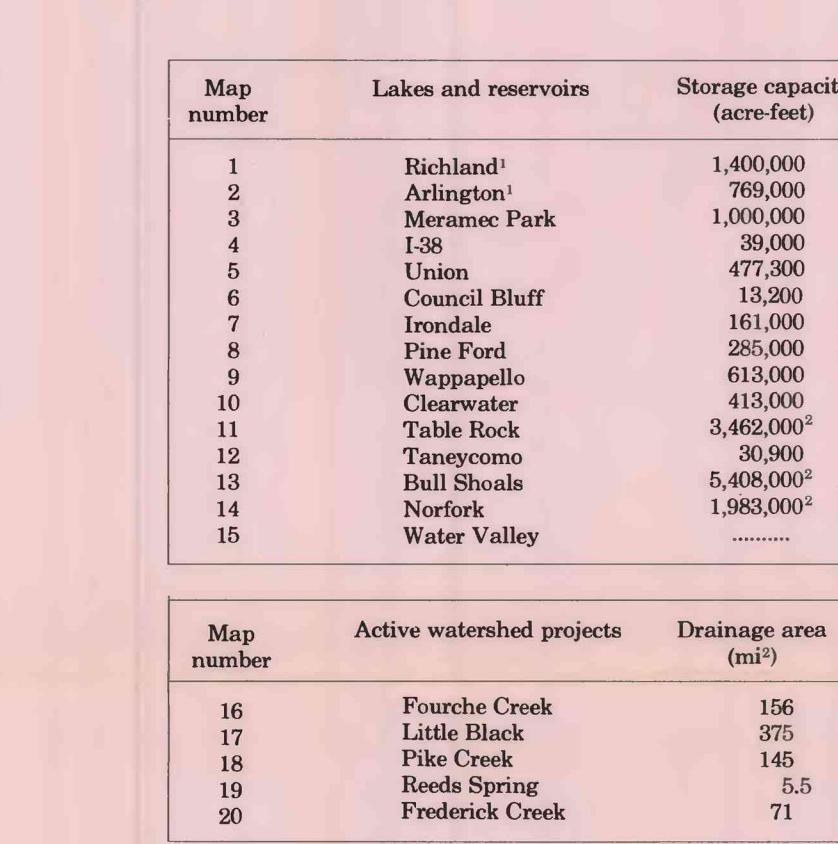
**FACTORS FOR CONVERTING ENGLISH UNITS TO METRIC UNITS**

Multiply English units by	To obtain metric units
Length	
Inches (in.)	2.54 centimeters (cm)
Feet (ft)	30.48 centimeters (cm)
Miles (mi)	1.609 kilometers (km)
Area	
Acre	0.405 square hectometers (ha) <sup>a</sup>
	0.00405 square kilometers (km <sup>2</sup> )
Square miles (mi <sup>2</sup> )	2.59 square kilometers (km <sup>2</sup> )
Volume	
Gallons (gal)	3.785 liters (l)
Million gallons (Mgal)	3.785 cubic meters (m <sup>3</sup> )
Acres-foot (acre-ft)	1.233 cubic hectometers (hm <sup>3</sup> )
	1.233 cubic kilometers (km <sup>3</sup> )
Flow	
Cubic feet per second (ft <sup>3</sup> /s)	0.0283 cubic meters per second (m <sup>3</sup> /s)
Gallons per minute (gal/min)	0.0631 cubic meters per second (m <sup>3</sup> /s)
Acres-foot (acre-ft)	6.309 cubic meters per second (m <sup>3</sup> /s)
Million gallons per day (Mgal/d)	0.228 cubic meters per second (m <sup>3</sup> /s)
Mass	
Tons per square mile (tons/mi <sup>2</sup> )	350 tonnes per square kilometer (t/km <sup>2</sup> )



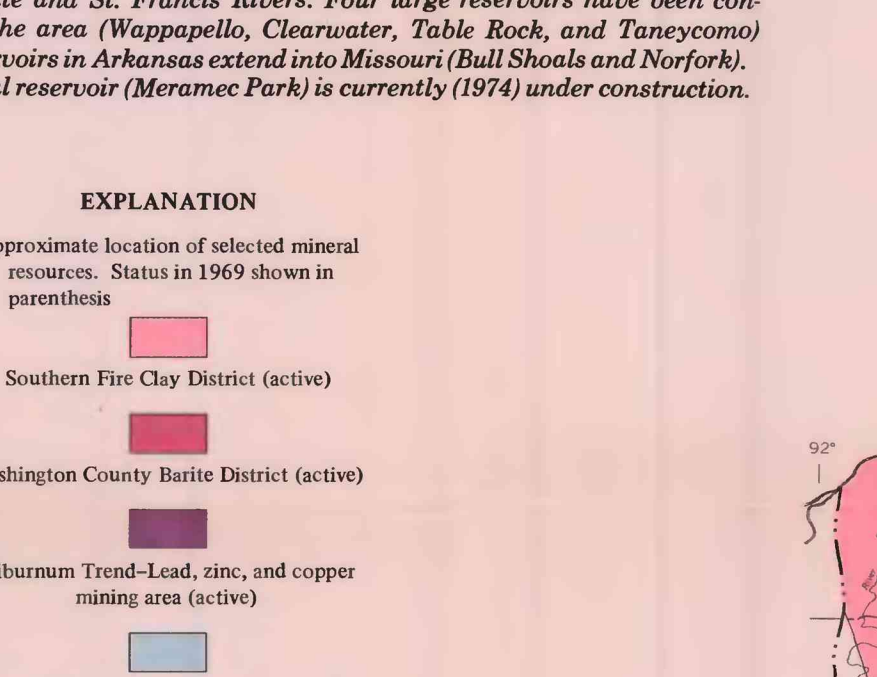
APPROXIMATELY 31 PERCENT OF THE MEAN ANNUAL PRECIPITATION BECOMES STREAMFLOW.—The remaining 69 percent is lost primarily to evapotranspiration. The amount of runoff from the area is unknown, but is estimated to be less than 1 percent of the mean annual precipitation. Underflow in the area is believed to be negligible. Although ground-water storage may vary significantly from year to year, the long-term change in storage is assumed to be negligible. Mean annual runoff and water loss values for several subbasins in the area deviate significantly from the values shown owing to interbasin subsurface flow. See map on sheet 3 for locations of basins that are known to contribute significant quantities of water to ground-water recharge.

### MINERAL RESOURCES



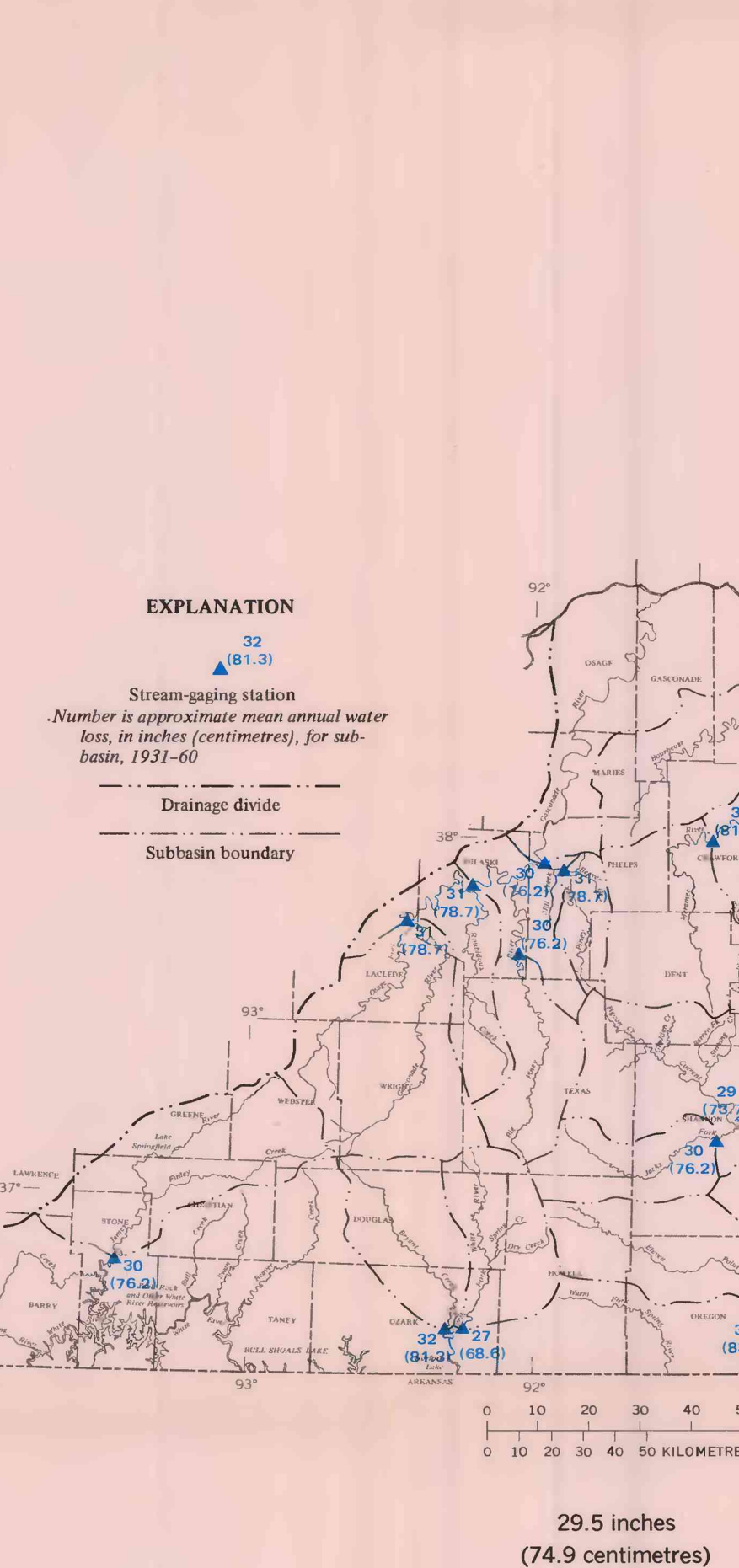
LAKES, RESERVOIRS, AND WATERSHED PROJECTS HELP TO STABILIZE STREAMFLOW, TO REDUCE DAMAGE FROM FLOODS AND EROSION, AND TO PROVIDE SITES FOR WATER-BASED RECREATION.—Approximately 339 lakes, mostly man-made, with surface areas in excess of 5 acres (2.2 ha) are known to exist in south-central Missouri in addition to numerous small farm ponds, abandoned quarries, and strip-mine pits. Several reservoirs and watershed projects have been authorized for future construction.

### TOPOGRAPHICALLY SOUTH-CENTRAL MISSOURI IS A MATURELY DISSECTED PLATEAU. The broad, rolling uplands are separated by intricate, meandering streams in deep, narrow valleys. Land surface altitudes range from 1,770 feet (539 m) in the St. Francois Mountains to 200 feet (61.0 m) on the flood plain of the Current River near the Missouri-Arkansas line. Principal streams in the northern part of the area flow northwest into the Missouri and Mississippi Rivers. Streams in the southern part flow south and southeast into the White and St. Francis Rivers. Four large reservoirs have been constructed in the area (Wappapello, Clearwater, Table Rock, and Taneytown) and two reservoirs in Arkansas extend into Missouri (Ball Shoals and Norfork). An additional reservoir (Meramec Park) is currently (1974) under construction.



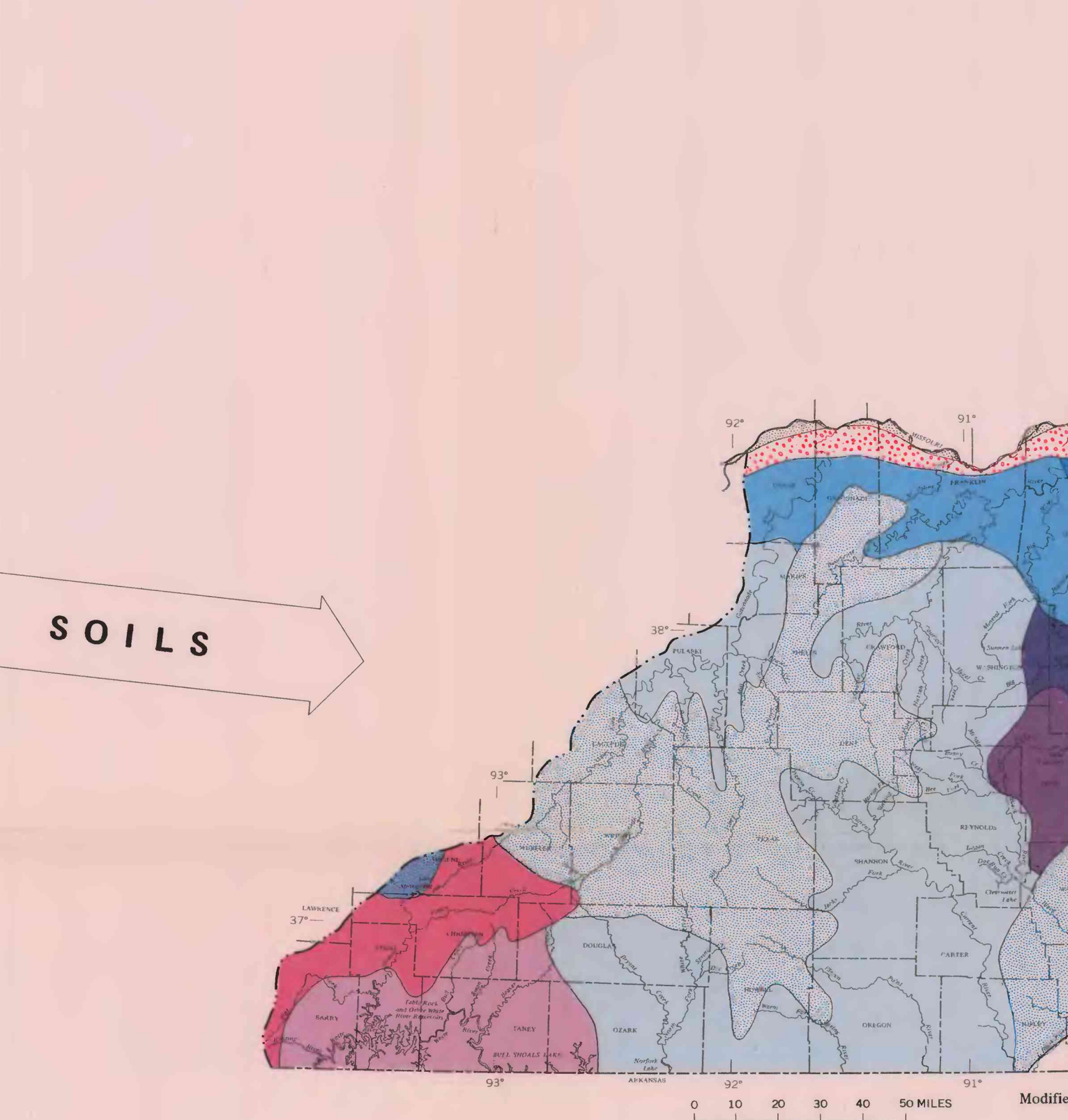
THE EXTRACTION AND PROCESSING OF MINERAL RESOURCES AFFECT THE WATER RESOURCES OF AN AREA.—Disseminating of mine workings may lower water levels in adjacent aquifers and increase the base flow of receiving streams. Tailings ponds may impermissibly impair flood waters and help to reduce downstream flood hazards. Failure of such impoundments during extreme runoff, however, may increase downstream flood hazards. Chemical reactions occurring in flooded mine shafts, strip pits, and tailings ponds may adversely affect the quality of water in streams and aquifers. Cooperation of mining interests, State and Federal agencies, universities, and the general public can eliminate most adverse effects.

### APPROXIMATE WATER BUDGET



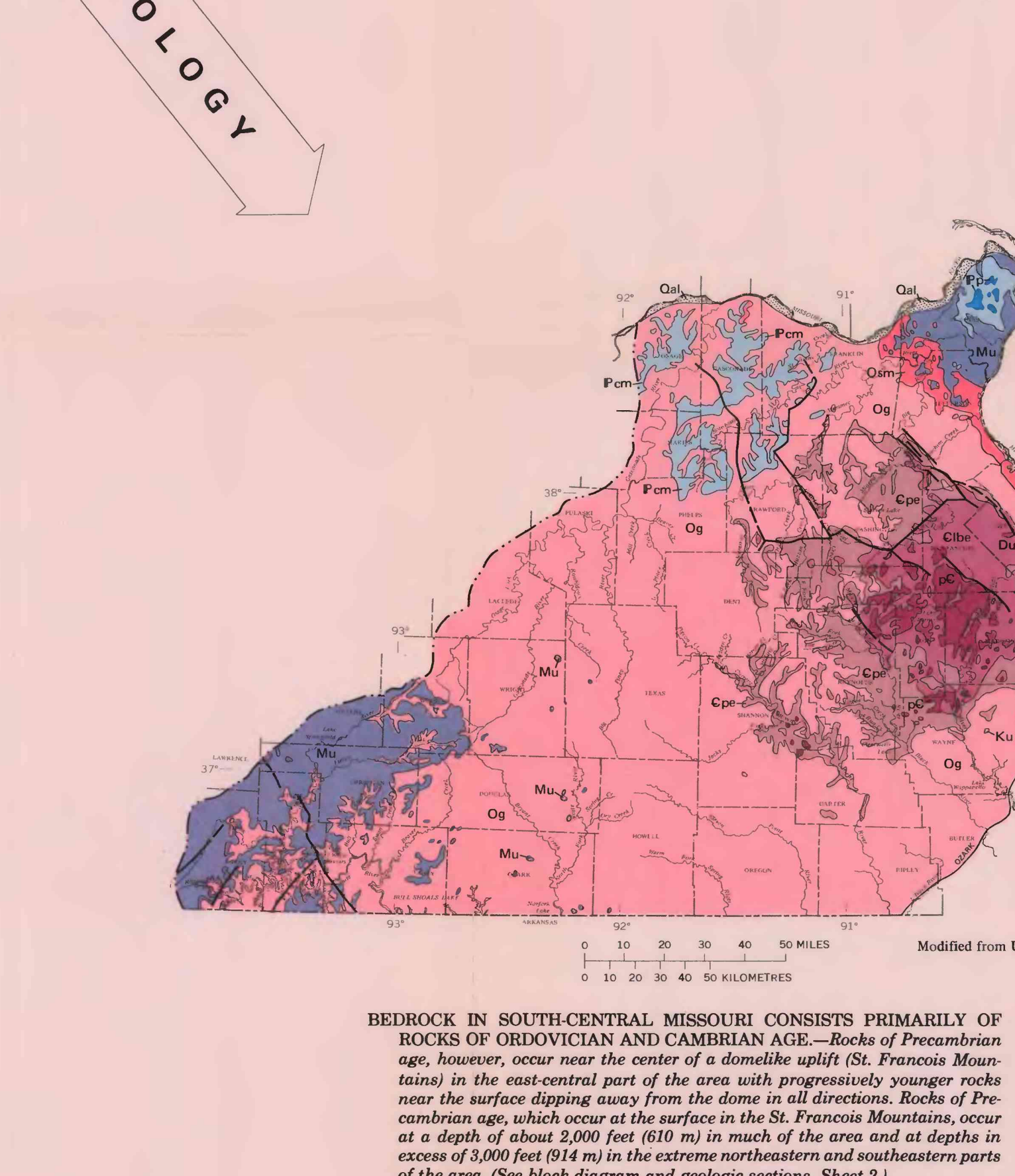
APPROXIMATELY 31 PERCENT OF THE MEAN ANNUAL PRECIPITATION BECOMES STREAMFLOW.—The remaining 69 percent is lost primarily to evapotranspiration. The amount of runoff from the area is unknown, but is estimated to be less than 1 percent of the mean annual precipitation. Underflow in the area is believed to be negligible. Although ground-water storage may vary significantly from year to year, the long-term change in storage is assumed to be negligible. Mean annual runoff and water loss values for several subbasins in the area deviate significantly from the values shown owing to interbasin subsurface flow. See map on sheet 3 for locations of basins that are known to contribute significant quantities of water to ground-water recharge.

### SOILS



SOILS DERIVED FROM CHERTY DOLOMITE, LIMESTONE, OR SANDSTONE OCCUPY ABOUT 75 PERCENT OF THE AREA.—Loess and loess-residual soils occupy the hill regions adjacent to the Missouri and Mississippi Rivers and soils derived from igneous and dolomitic rocks form the Ashe-Tilt Hugenotian and Tilt-Hugenotian soil associations. Estimates of permeability (the quality of a soil that enables water or air to move through it (Gilbert, 1971)) for most soils in the area range from moderate (1-5 inches or 2.5-12.7 cm of water percolation per hour) to high (5-10 inches or 12.7-25.4 cm of water percolation per hour). The Lebanon and associated soils, however, usually contain a dense subsoil, frequently 2 to 3 feet (0.6-0.9 m) below the surface which has low permeability (<1 inch or 2.54 cm of water percolation per hour).

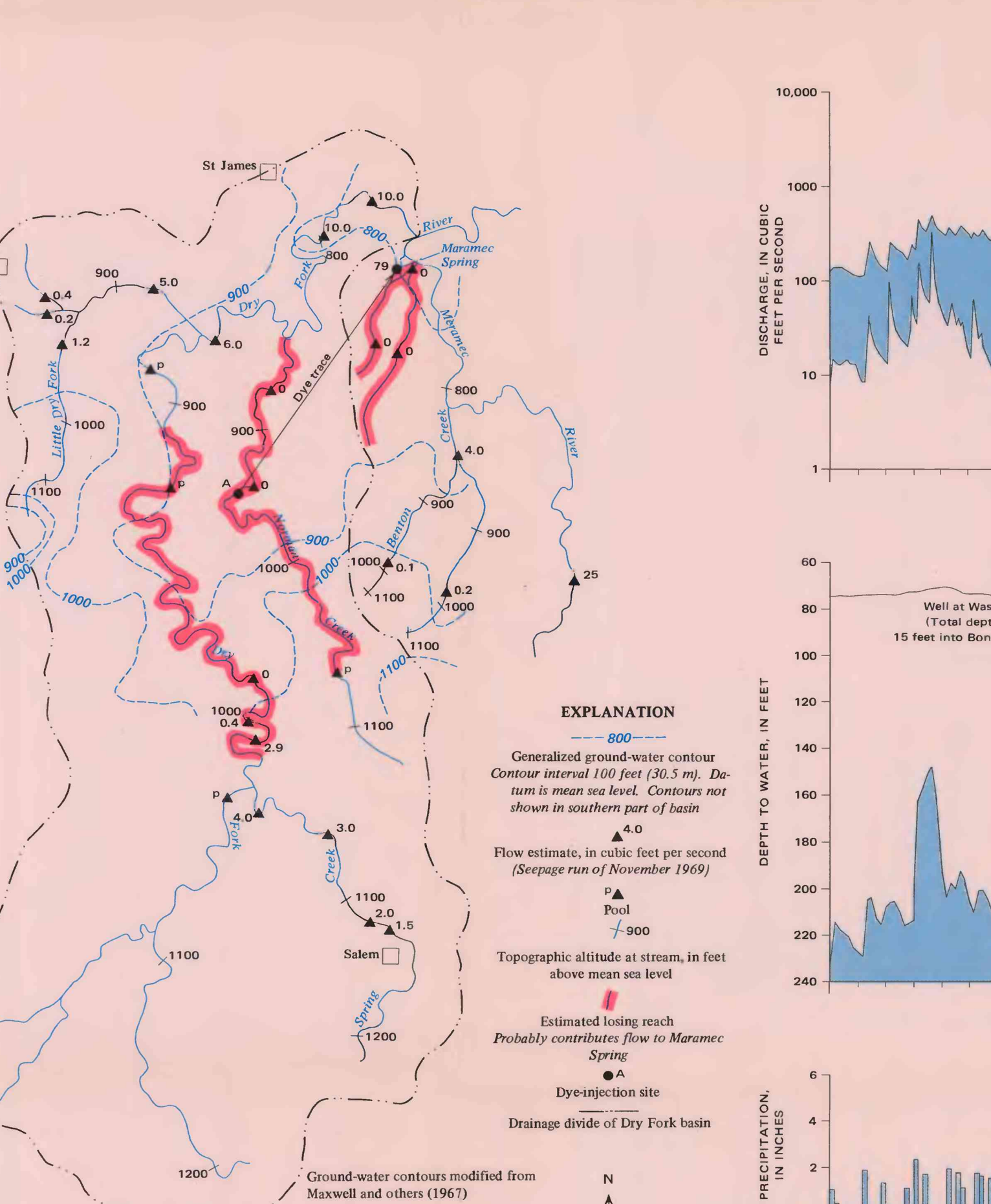
### GEOLOGY



BEDROCK IN SOUTH-CENTRAL MISSOURI CONSISTS PRIMARILY OF ROCKS OF ORDOVICIAN AND CAMBRIAN AGE.—Blocks of Precambrian age, however, occur near the center of a dome-like uplift (St. Francois Mountains) in the east-central part of the area with progressively younger rocks near the surface dipping away from the dome in all directions. Rocks of Precambrian age, which occur at the surface in the St. Francois Mountains, occur at a depth of about 2,000 feet (610 m) in much of the area and at depths in excess of 3,000 feet (914 m) in the extreme northeastern and southeastern parts of the area. (See block diagrams and geologic sections, Sheet 2.)

Rocks of Ordovician and Cambrian age are composed chiefly of dolomite with minor amounts of sandstone and limestone. Solution of these rocks has resulted in the occurrence of numerous sinkholes, losing streams, and large springs in much of the area. The presence of these features usually indicates relatively free exchange of flow between land surface and aquifers.

### HYDROLOGIC RELATIONSHIPS



LOSING STREAMS CONTRIBUTE WATER TO AQUIFERS AND SPRINGS.—Interpretation of seepage-run data, dye-tracing data, and ground-water contour maps aids in identifying losing streams and in delineating outflow areas of springs. Normally the headwaters of a stream are intermittent and after descending to lower altitudes, the stream becomes perennial. Where ground-water levels lie unusually deep, however, surface streams often lose much of their flow to subsurface drainage. Flow occurs throughout the length of these streams only following intense rainfalls of greater than about 2 to 4 inches (5.08 to 10.16 cm), which results in a cumulative flow that exceeds the infiltration capacity of the main-channel streambed. Ground-water levels in the losing reaches usually lie 50 to 100 feet (15.2 to 30.5 m) below the altitude of the streambed. Fluorescent dye injected in Norman Creek at Point A disappeared into the sand and gravel streambed and emerged at Meramec Spring 65-75 days later. (See dye-tracing data, Sheet 2.)

PRECIPITATION IS THE SOURCE OF WATER IN STREAMS, SPRINGS, AND AQUIFERS.—Following periods of heavy precipitation, significant quantities of water enter the ground through openings in the soil and rocks. Excess water is drained from the land surface by streams. Much of the water that enters the ground is returned to the atmosphere by evaporation from soils and transpiration by plants, but some water percolates downward to recharge aquifers and maintain the flow of springs. In deeply weathered carbonate terranes such as that which occurs in the West Plains area, water is conducted freely to aquifers and springs by leaching, sink holes, and solution openings. Ground-water levels in such areas respond rapidly to precipitation, as shown by the hydrograph for a well at West Plains, Missouri, which is conducted freely to aquifers and springs by leaching, sink holes, and solution openings. Ground-water levels in such areas respond rapidly to precipitation, as shown by the hydrograph for a well at West Plains, Missouri, which is conducted freely to aquifers and springs by leaching, sink holes, and solution openings. Large springs serve as outlets for the complex underground drainage system (hydrograph for Green Springs) and are usually fed both by local solution openings and regional aquifers. Where deep weathering of carbonate terranes is less pronounced as in the Washington area, local recharge is limited, ground-water levels respond less rapidly to precipitation, and large springs do not occur.