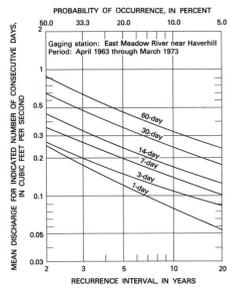
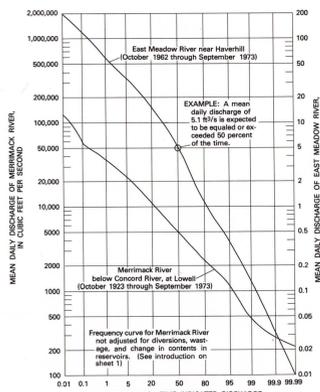


SURFACE WATER AVAILABILITY

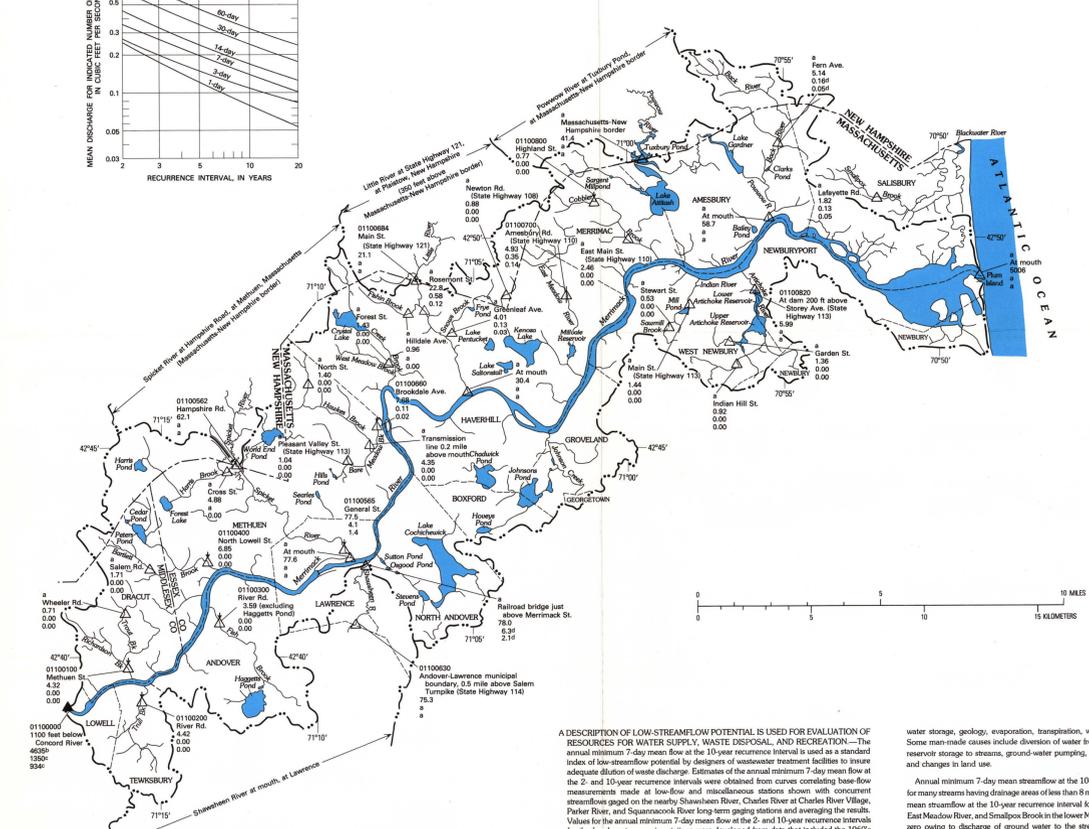
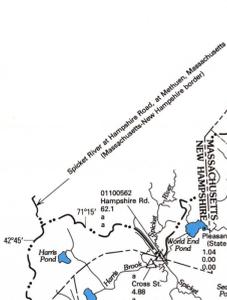
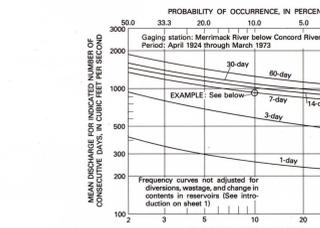
LOW STREAMFLOW

FLOW DURATION



THE PERCENTAGE OF TIME FLOWS HAVE BEEN EQUALED OR EXCEEDED DURING PERIOD OF RECORD SHOWN.—The slope of the flow-duration curve for high and medium flows in the Merrimack River is moderated by flood control reservoirs, water-supply reservoirs, and lakes that store peak flows as well as other flow for later diversion and gradual release during periods of lower flow. The steep part of the curve, representing flows exceeded less than 0.1 percent of the time, is due to the peak flows of the March 1936 flood and the September 1938 hurricane, which occurred before construction of the five major flood-control dams in the upper Merrimack River basin. (See table 2 on sheet 1.)

The steep slope of the low-flow part of the duration curve (flows exceeded 95 percent of the time) for East Meadow River is caused by high rates of evapotranspiration from swamps. However, the gentle slope of the low-flow part of the duration curve (flows exceeded 99 percent of the time) for the Merrimack River is due to discharge from lake, reservoir, and ground-water storage.



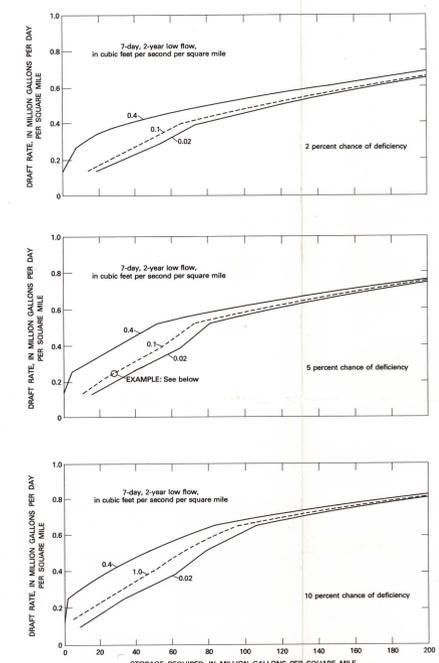
A DESCRIPTION OF LOW-STREAMFLOW POTENTIAL IS USED FOR EVALUATION OF RESOURCES FOR WATER SUPPLY, WASTE DISPOSAL, AND RECREATION.—The annual minimum 7-day mean flow at the 10-year recurrence interval is used as a standard index of low-streamflow potential by designers of wastewater treatment facilities to insure adequate dilution of waste discharge. Estimates of the annual minimum 7-day mean flow at the 2- and 10-year recurrence intervals were obtained from curves consisting base-flow measurements made at low-flow and miscellaneous stations shown with concurrent streamflows gaged on the nearby Shawheen River, Charles River at Charles River Village, Parker River, and Squamackoak River long-term gaging stations and averaging the results. Values for the annual minimum 7-day mean flow at the 2- and 10-year recurrence intervals for the four long-term gaging stations were developed from data that included the 1960's drought.

Streamflow per unit drainage area varies from stream to stream during low-flow periods as the result of both natural and man-made causes. Natural causes include nonuniformity of water storage, geology, evaporation, transpiration, water-table slope, and precipitation. Some man-made causes include diversion of water from streams, managed release from reservoir storage to streams, ground-water pumping, waste-water discharge into streams, and changes in land use.

Annual minimum 7-day mean streamflow at the 10-year recurrence interval will be zero for many stream having drainage areas of less than 100 mi²; however, annual minimum 7-day mean streamflow at the 10-year recurrence interval for Back River, Base Meadow Brook, East Meadow River, and Smallpox Brook in the lower Merrimack River basin are greater than zero owing to discharge of ground water to the streams from porous sand and gravel deposits. Scarce ground-water discharge from the upper part of the bedrock aquifer (consolidated rock) and from the extensive till layer (unconsolidated deposit) is mainly responsible for the zero values.

- #### EXPLANATION
- ▲ Continuous-record gaging station
 - △ Discontinued gaging station
 - ▽ Low-flow measurement station
 - Miscellaneous-measurement station
 - Dam
 - 0110060 Station number, in downstream order. Number from U.S. Geological Survey's data-collection network
 - Brookdale Ave. Location on river course
 - 7.68 Drainage area, in square miles
 - Annual minimum 7-day mean flow, in cubic feet per second
 - 0.11 2-year recurrence interval
 - 0.02 10-year recurrence interval
 - Not determined.
 - 4050⁺ Total drainage area above gaging station; not adjusted for diversion reservoirs in drainage basin.
 - 1300⁺ Unadjusted for diversions, wastage, and change in contents in reservoirs above gaging station.
 - 2.1⁺ Withdrawal of ground water from drainage basin above station for municipal supply causes annual minimum 7-day mean flows at the 2- and 10-year recurrence intervals to be lower than predevelopment period.
 - Basin boundary

REGIONAL STORAGE ANALYSIS



REGIONAL DRAFT-STORAGE-FREQUENCY ANALYSIS USED IN CONJUNCTION WITH ANNUAL MINIMUM 7-DAY MEAN FLOWS AT THE 2-YEAR RECURRENT INTERVAL CAN PROVIDE INFORMATION USEFUL TO WATER MANAGERS AND PLANNERS IN COMPARATIVE STUDIES OR IN PRELIMINARY DESIGNS FOR RESERVOIR SITES.—These regional draft-storage analysis curves were prepared by Tasker (1977). They are based on a regional analysis of streamflow records for 12 gaging stations in eastern Massachusetts and Rhode Island, using the methods recommended by Riggs and Hardison (1973) to account for seasonal as well as year-to-year variations in streamflow.

Draft rate to gross reservoir outflow occurring at a uniform rate and is uncorrected for increased water losses resulting from enlarging the water surface area in the basin. Storage required is the usable volume of a reservoir required to maintain the indicated draft rate. When applying these curves to the evaluation of a specific reservoir site, the user should adjust the draft rates to account for expected evaporation and seepage losses, nonuniform draft rates, and low-flow augmentation requirements. Storage requirements should be adjusted for expected losses in reservoir capacity because of sedimentation.

The Committee on Rainfall and Yield of Drainage Areas of the New England Water Works Association (1969, p. 168-169) considers it economically unsound to develop storage capabilities much beyond 200 Mg/d in New England watersheds. The rationale for this decision can be seen from inspection of the draft-storage curves. Increases of draft rates require disproportionately large increases of storage for rates exceeding about 0.5 Mg/d.

EXAMPLE:—A stream has a 10.0 mi² drainage area and an annual minimum 7-day mean flow at the 2-year recurrence interval of 1.0 ft³/s or 0.11 (ft³/s)/mi² at a given site. To maintain a draft rate of 2.5 Mg/d (or 0.25 Mg/d/mi²), a storage capacity of 20 Mg/d or 250 Mg/d would be enough for an average of 95 out of 100 years (5 percent chance of deficiency), exclusive of losses.

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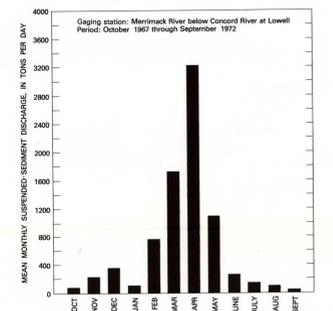
- Lowell (1969) Salem Depot (1974)
- Haverhill (1974) South Groveland (1974)

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SEDIMENT



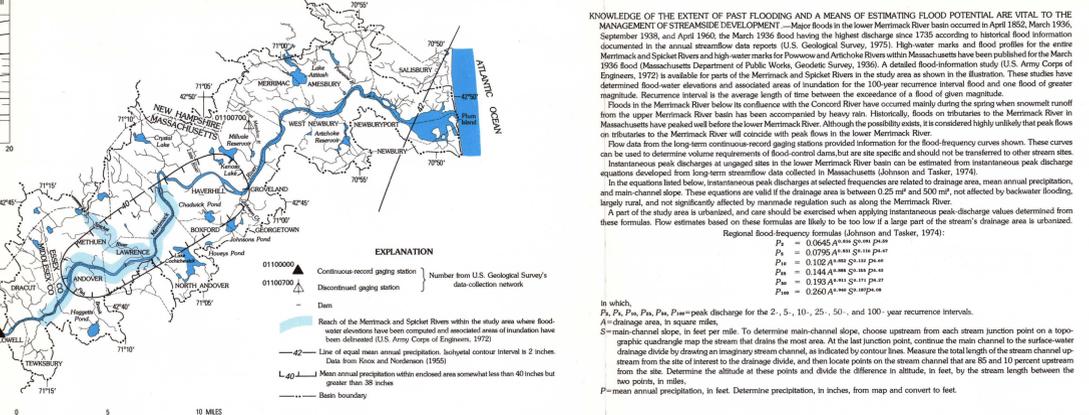
THE MERRIMACK RIVER TRANSPORTS APPROXIMATELY 75 PERCENT OF ITS YEARLY SUSPENDED-SEDIMENT LOAD DURING SPRING RUNOFF, MARCH THROUGH MAY.—Suspended sediment in streamwater is composed of clay, silt, and sand-size particles derived from the erosion of soils and unconsolidated deposits and maintained in suspension by the velocity of streamwater. The amount and particle size of suspended sediment varies with the stream's discharge and velocity as well as the amount and particle size of erodible material available.

In the lower Merrimack River, high suspended-sediment concentrations occur from snowmelt and intense rainfall in the upper part of the basin. Suspended sediment concentration is highly variable but usually varies directly with streamflow. Generally, greater suspended sediment concentrations occur during increasing streamflow than during decreasing streamflow of the same magnitude.

The occurrence of greater concentrations during increasing flow than during decreasing flow, the season of the year, amount and type of vegetative cover, land use, agricultural practices, intensity of rainfall, and construction are some of the factors explaining why a daily average streamflow of 10,100 ft³/s yielded 518 tons of suspended sediment load on March 26, 1972, whereas the same streamflow yielded only 164 tons of sediment on July 26, 1972 at the Merrimack River below the Concord River at Lowell gaging station.

At the Lowell gaging station, annual suspended-sediment loads during October 1967 through September 1972 ranged from 119,530 to 306,463 tons/yr and had an average annual suspended-sediment load of 244,231 tons/yr or 666 tons/d. Suspended-sediment data for the Lowell gaging station are published in the annual streamflow data reports (U.S. Geological Survey, 1969, 1970, 1971, 1973, and 1974).

HIGH STREAMFLOW



KNOWLEDGE OF THE EXTENT OF PAST FLOODING AND A MEANS OF ESTIMATING FLOOD POTENTIAL ARE VITAL TO THE MANAGEMENT OF STREAMFLOW DEVELOPMENT.—Major floods in the lower Merrimack River basin occurred in April 1952, March 1936, September 1938, and April 1960, the March 1936 flood having the highest discharge since 1735 according to historical flood information documented in the annual streamflow data reports (U.S. Geological Survey, 1975). High-water marks and flood profiles for the entire Merrimack and Spickett Rivers and high-water marks for the Merrimack and Spickett Rivers in Massachusetts have been published for the March 1936 flood (Massachusetts Department of Public Works, Geologic Survey, 1936). A detailed flood-information study (U.S. Army Corps of Engineers, 1972) is available for parts of the Merrimack and Spickett Rivers in the study area as shown in the illustration. These studies have determined flood-water elevations and associated areas of inundation for the 10-year recurrence interval flood and one flood of greater magnitude. Recurrence interval is the average length of time between the exceedance of a flood of given magnitude.

Floods in the Merrimack River below its confluence with the Concord River have occurred mainly during the spring when snowmelt runoff from the upper Merrimack River basin has been accompanied by heavy rain. Historically, floods on tributaries to the Merrimack River in Massachusetts have peaked well before the lower Merrimack River. Although the possibility exists, it is considered highly unlikely that peak flows on tributaries to the Merrimack River will coincide with peak flows in the lower Merrimack River.

Flow data from the long-term continuous-record gaging stations provided information for the flood-frequency curves shown. These curves can be used to determine volume requirements of flood-control dams, but are site specific and should not be transferred to other stream sites. Instantaneous peak discharges at ungaged sites in the lower Merrimack River basin can be estimated from instantaneous peak discharge equations developed from long-term streamflow data collected in Massachusetts (Johnson and Tasker, 1974).

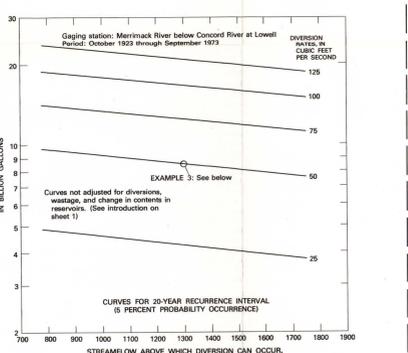
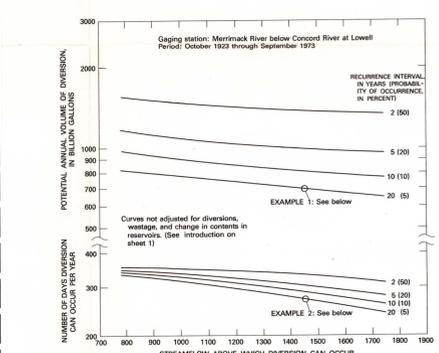
In the equations listed below, instantaneous peak discharges of selected frequencies are related to drainage area, mean annual precipitation, and main-channel slope. These equations are valid if the drainage area is between 0.25 mi² and 500 mi², not affected by backwater flooding, largely rural, and not significantly affected by manmade regulation such as along the Merrimack River.

A part of the study area is urbanized, and care should be exercised when applying instantaneous peak discharge values determined from these formulas. Flow estimates based on these formulas are likely to be too low if a large part of the stream's drainage area is urbanized.

Streamflow per unit drainage area varies from stream to stream during low-flow periods as the result of both natural and man-made causes. Natural causes include nonuniformity of water storage, geology, evaporation, transpiration, water-table slope, and precipitation. Some man-made causes include diversion of water from streams, managed release from reservoir storage to streams, ground-water pumping, waste-water discharge into streams, and changes in land use.

Annual minimum 7-day mean streamflow at the 10-year recurrence interval will be zero for many stream having drainage areas of less than 100 mi²; however, annual minimum 7-day mean streamflow at the 10-year recurrence interval for Back River, Base Meadow Brook, East Meadow River, and Smallpox Brook in the lower Merrimack River basin are greater than zero owing to discharge of ground water to the streams from porous sand and gravel deposits. Scarce ground-water discharge from the upper part of the bedrock aquifer (consolidated rock) and from the extensive till layer (unconsolidated deposit) is mainly responsible for the zero values.

VOLUME DIVERSION



DESIGNERS AND PLANNERS CAN DETERMINE THE ANNUAL VOLUME OF WATER AVAILABLE FOR DIVERSION FROM THE MERRIMACK RIVER BELOW THE CONCORD RIVER ONCE MINIMUM FLOW REQUIREMENTS DOWNSTREAM FROM THE DIVERSION STRUCTURE HAVE BEEN ESTABLISHED.—The set of curves for the Merrimack River below its confluence with the Concord River can be used to estimate the potential annual volume of diversion available for different minimum flow requirements (example 1), and to estimate how many days a year streamflow above a given flow will occur (example 2). The potential annual volume of diversion for diversion structures of different capacities can also be estimated (example 3). These relationships were prepared by a method developed by Collings (1968).

Summation of prior claims on river water: downstream from the diversion structure will yield minimum streamflow requirements, when available, to satisfy these demands. Examples of prior claims on river water are: industrial or municipal use, recreation, irrigation, fisheries resource, waste-water dilution, and limiting saltwater intrusion in an estuary.

EXAMPLE 1: If the streamflow greater than 1450 ft³/s will occur, the potential annual volume of diversion would be less than 700 billion gallons at intervals averaging 20 years in length, and the probability that the potential annual volume of diversion would be less than 700 billion gallons in any 1 year is 5 percent.

EXAMPLE 2: Streamflow greater than 1450 ft³/s will occur for less than 272 days in a year at intervals averaging 20 years in length, and the probability that streamflow greater than 1450 ft³/s will occur for less than 272 days in any 1 year is 5 percent.

EXAMPLE 3: If the capacity of a diversion structure is 50 ft³/s and if the full diversion rate is made whenever streamflow is greater than 1300 ft³/s, the potential annual volume of diversion will be less than 8.5 billion gallons at intervals averaging 20 years in length, and the probability that the potential annual volume of diversion will be less than 8.5 billion gallons in any 1 year is 5 percent.