

## INTRODUCTION

Sedimentary rocks of Late Cambrian through Early Cretaceous age in Kansas are part of a regional flow system of hydraulically connected aquifers and confining units. Future demands for water require that these deeply buried rocks be studied to describe hydrologic properties and ground-water-flow conditions and to provide information that will serve as the basis for decisions concerning the protection and the management of the water resources contained therein. Toward this end, the U.S. Geological Survey, as part of a Central Midwest Regional Aquifer-System Analysis (CRMASA), began a 5-year hydrologic investigation of this regional flow system in Arkansas, Colorado, Kansas, Missouri, Nebraska, New Mexico, Oklahoma, South Dakota, and Texas (Jorgensen and Sigler, 1981).

This chapter is one of nine contained in Hydrologic Investigations Atlas HA-722, which present a description of the physical framework (Chapters B-F) and the geohydrology (Chapters G-I) of principal aquifers and confining systems in Upper Cambrian through Lower Cretaceous rocks in Kansas; the stratigraphic relations of these geohydrologic systems are discussed in detail in Chapter A (Wolf and others, 1990). This chapter (C) describes the geohydrology of the Great Plains aquifer system; the physical framework of the Great Plains aquifer system is presented in Chapter B (Sponholz and others, 1992).

The maps in this chapter are based on existing data from selected geophysical and lithologic logs, drill-stem tests, water-level measurements, water-quality analyses, and published maps of stratigraphically equivalent units. An index to the geohydrologic data compiled for the CRMASA in Kansas is presented in Sponholz and others (1987). For the most part, data used to construct the maps were collected over many years and do not reflect aquifer conditions for any specific time period.

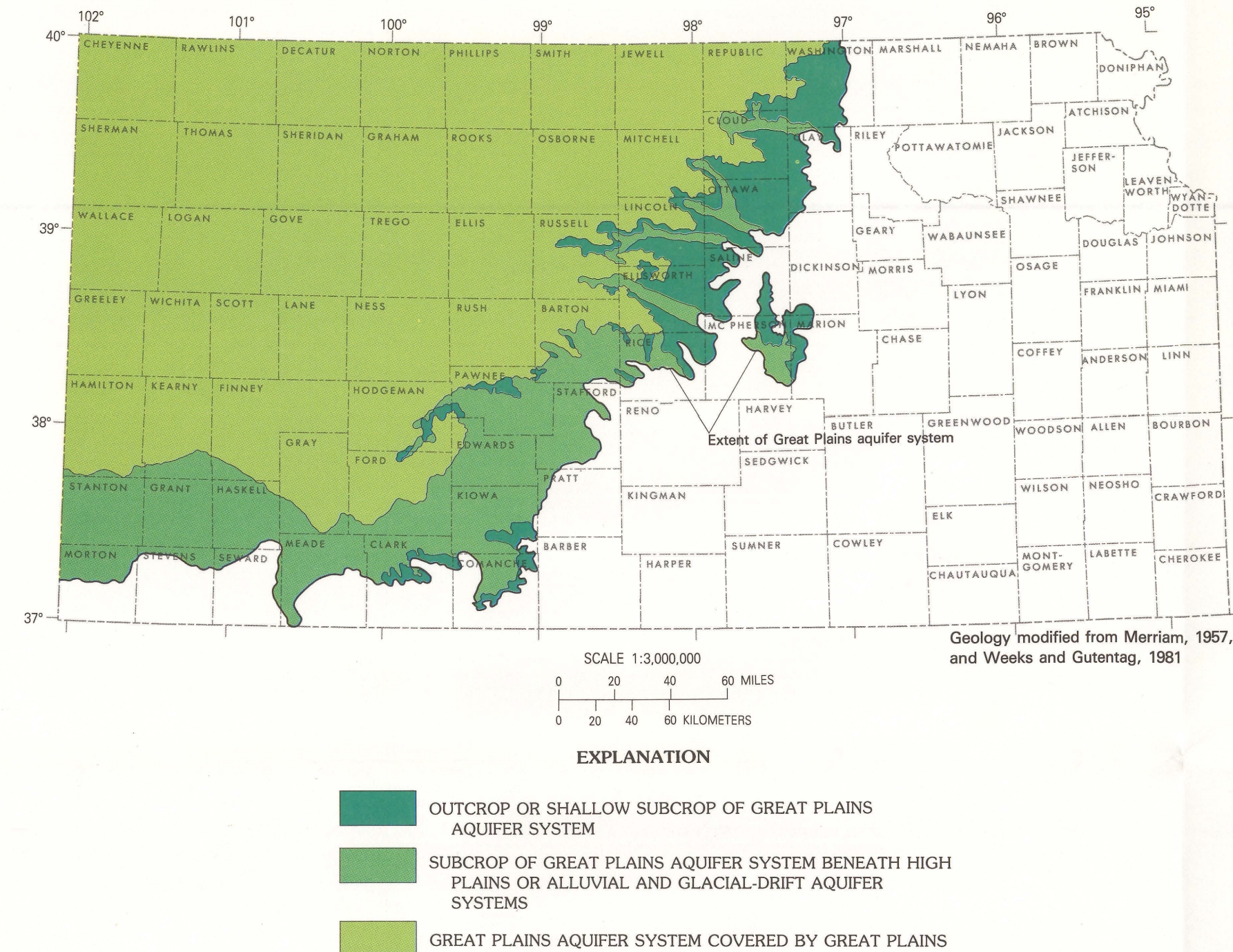


FIGURE 1.—Geohydrologic map showing extent, outcrop, and subcrop of Great Plains aquifer system.

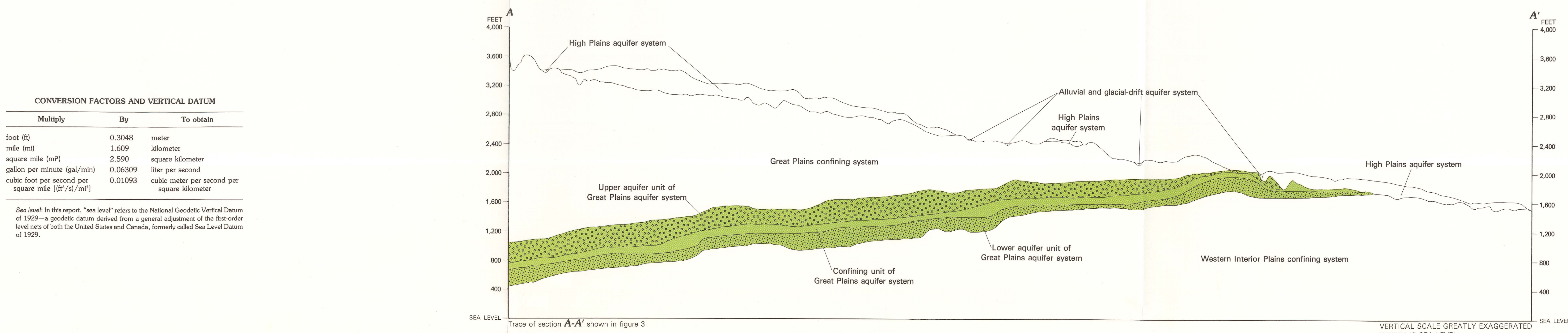


FIGURE 2.—Geohydrologic section showing relation of Great Plains aquifer system to other geohydrologic systems.

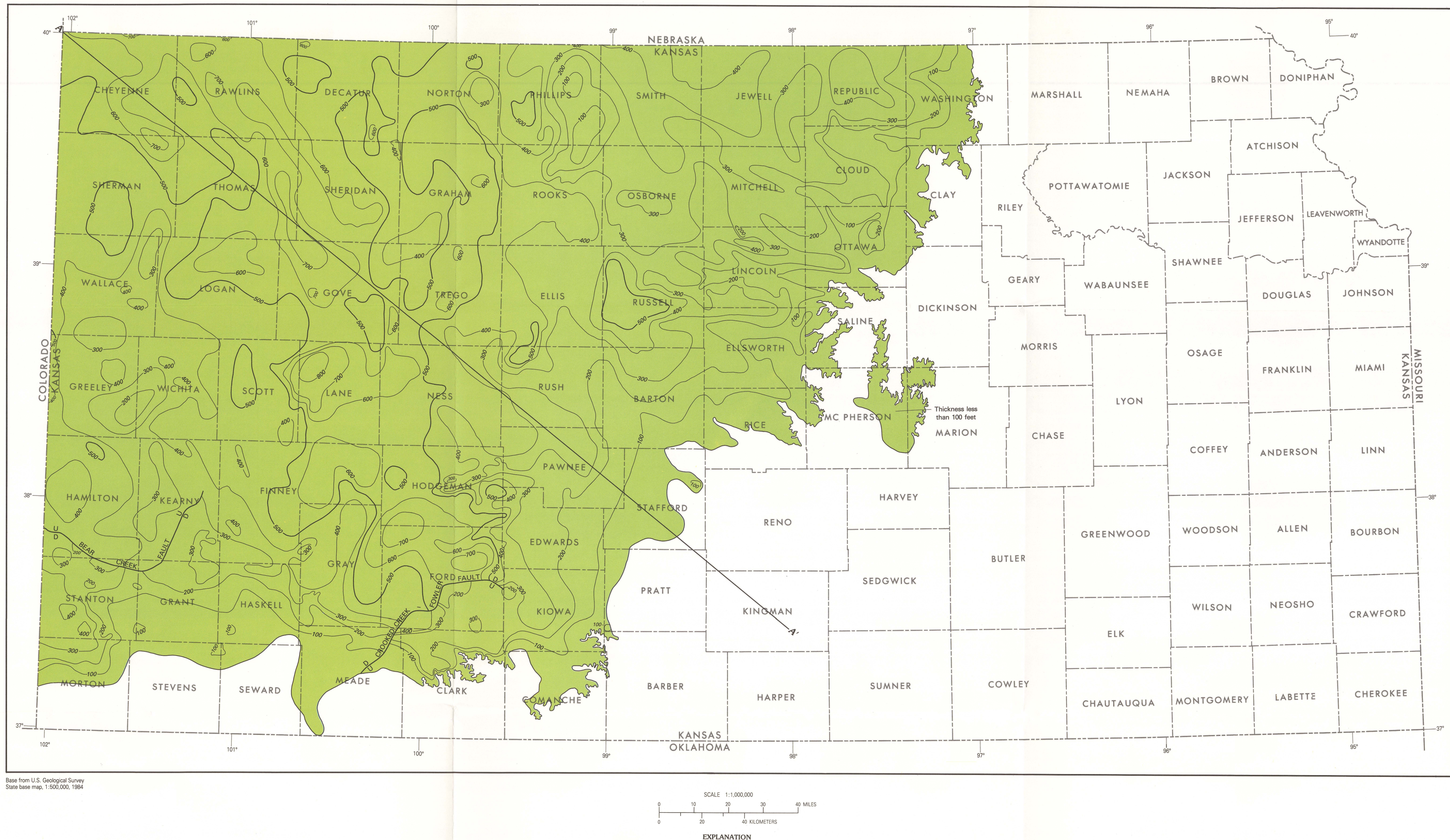


FIGURE 3.—Geohydrologic map showing areal extent and thickness of Great Plains aquifer system.

## DESCRIPTION OF PHYSICAL CHARACTERISTICS

The Great Plains aquifer system, which is generally known as the "Dakota aquifer" in hydrologic literature, is one of the most extensive aquifer systems in North America—from northern Canada to New Mexico in the United States. The aquifer system extends over a large part of western Kansas, as shown in figure 1, and consists mostly of Lower Cretaceous rocks. The aquifer system crops out along the eastern end of the southern margin of the area. The aquifer system generally is located toward the west and the south by the Great Plains confining system (Sponholz and others, 1992). The aquifer system is directly overlain by the High Plains aquifer system, which comprises deposits of Quaternary and Tertiary age, and the alluvial and glacial-drift aquifer system, which comprises Holocene, Pleistocene, and Miocene sediments. The Great Plains aquifer system is underlain everywhere in Kansas by the Western Interior Plains confining system, which is a thick sequence of Jurassic, Permian, Pennsylvanian, and Upper Mississippian rocks. The stratigraphic relation of the Great Plains aquifer system to overlying and underlying geohydrologic systems is illustrated in figure 2 and listed with equivalent stratigraphic units in table 1.

Rocks of the Great Plains aquifer system are an interbedded sequence of sandstone, shale, and siltstone. The thickness of the aquifer system ranges from a few feet at the exposed edges in the east and the south to about 800 ft in the western part of the State (fig. 3). In general, thickness of the aquifer system increases toward the north and the west. Depth to the top of the aquifer system below land surface also increases from a few feet in the outcrop area to about 2,500 ft in the northwestern corner of the State (fig. 4). The Great Plains aquifer system in Kansas is included within a sequence of rocks that consists of shale and siltstone with interbedded and lenticular sandstone. An examination of geophysical logs commonly indicates significant lithologic variability from site to site. The log from one site may indicate mostly thick beds of sandstone, whereas that from a nearby site may indicate mostly shale with a few thin beds of sandstone.

Lithologic and fossil evidence from rocks of the Great Plains aquifer system indicate marine and nonmarine sediments. Deposition probably occurred along the coastline of a sea that alternately advanced and retreated across the area (Kenne and Bayne, 1977). Sand probably was deposited in shallow seas and deltas, on beaches and nearshore bars, and along river channels. As the location of the coastline continued to move, beds of sand may have been partly or completely removed by planation and replaced by other beds of shale or sand. Deposition of this type would not provide definable boundaries that were consistent throughout large areas.

The Great Plains aquifer system was divided into the following regional geohydrologic subdivisions during the CRMASA (D.G. Jorgensen, U.S. Geological

Survey, written commun., 1985): (1) an upper sandstone aquifer unit named the "Haskell aquifer," (2) a middle, predominantly shale confining unit named the "Aftonian confining unit," and (3) a lower sandstone aquifer unit named the "Aftonian aquifer." The local names "upper aquifer unit," "confining unit," and "lower aquifer unit" are used in this report instead of the three CRMASA regional geohydrologic subdivisions. The relation between geohydrologic systems and equivalent stratigraphic units are shown in table 1.

### Upper Aquifer Unit

The upper aquifer unit of the Great Plains aquifer system in Kansas is stratigraphically equivalent to the Dakota Formation (table 1). This unit comprises shale and siltstone with an abundance of interbedded and lenticular sandstone. Thickness and character of the sandstone may differ significantly throughout the area. The fine to coarse sand may be loose and well cemented and may contain small to large percentages of silt and clay. Hydraulic conductivity, which is a measure of rock permeability, is estimated to range from 1 to 41 ft/d for the upper aquifer unit in southwestern Kansas (Sponholz, 1986; Kenne and Sponholz, 1988). Permeability of the sandstone in the upper aquifer unit mostly depends on the degree of cementation and the percentage of silt and clay. Although the sandstone in the upper aquifer unit may be in separate beds in some parts of the study area, generally there is sufficient interconnection between the beds to permit the movement of water. Fine-grained shale and siltstone, which have relatively low permeability, commonly retard local ground-water flow.

The upper aquifer unit in an area to the north of the west of an irregular line from Washington County in the northeast to Morton County in southwestern Kansas (fig. 5A). The upper aquifer unit crops out along stream valleys in the northeastern part of the area and mostly underlies the shale and siltstone of the Great Plains confining system elsewhere. In southwestern Kansas, however, the unit subsides beneath unconformable deposits of the High Plains aquifer system. Thickness of the upper aquifer unit ranges from a few feet in the outcrop and the outcrop area to about 600 to 500 ft in scattered areas mostly in northern and northwestern Kansas (Sponholz and others, 1992).

### Confining Unit

The confining unit of the Great Plains aquifer system in Kansas is stratigraphically equivalent to the Kiowa Shale (table 1). This unit predominantly consists of shale with some siltstone and thin beds of sandstone. The sandstone is most abundant in central and southwestern Kansas. In parts of north-central Kansas, sandstone constitutes as much as 20 percent of the rock. However, the thickness values do not indicate the degree of cementation or the percentage of silt and clay that may be included.

lower aquifer units. Where the percentage of sandstone increases relative to the silt and clay content, some water probably flows between the aquifer units as a result of the greater vertical flow through the confining unit. The confining unit is an area that mostly coincides with the area of the upper aquifer unit (fig. 5B). The confining unit extends beyond the upper aquifer unit in some areas along the eastern and the southern margins where it may subcrop beneath the High Plains or shallow aquifer systems but is missing in the northern part of Norton, Phillips, Smith, Jewell, Haskell, and Washington Counties. Thickness of the confining unit ranges from a few feet in the outcrop and subcrop areas to about 200 to 300 ft in north-central and south-central Kansas (Sponholz and others, 1992).

### Lower Aquifer Unit

The lower aquifer unit of the Great Plains aquifer system in Kansas is stratigraphically equivalent to the Cherokee Sandstone (table 1). This unit predominantly consists of a zone of conglomerate at the base and sandstone with lenses of sandy shale and siltstone. Fine to medium sand in the sand may be loose to well cemented and may contain small to large amounts of silt and clay. Permeability of the lower aquifer unit probably is similar to that of the upper aquifer unit. Leakage between the upper and the lower aquifer units is reported whenever they are separated by shale sequences of the confining unit. The lower aquifer unit occurs in an area that mostly coincides with the upper aquifer and the confining unit (fig. 5C). The lower aquifer unit extends beyond the overlying units mostly along the southeastern margin in Barton, Comanche, and Stafford Counties. The lower aquifer unit is missing beneath overlying units in all or part of the counties east of Norton, Osborne, Russell, Ellsworth, and Barton Counties. The lower aquifer unit is underlain everywhere by the Western Interior Plains confining system. Thickness of the lower aquifer unit generally increases northward and ranges from a few feet at the outcrop or outcrop to about 500 ft in northwestern Kansas (Sponholz and others, 1992).

### Thickness of Sandstone in Aquifer System

A generalized map (fig. 6), which is based on interpretations from geophysical logs, shows the total thickness of sandstone in Lower Cretaceous rocks. Sandstone thickness in the eastern one-half of the Great Plains aquifer system is commonly less than 100 ft and ranges from about 200 to 500 ft in the western one-half. The increased sandstone thickness suggests the increased quantity of water may be available from storage in the western part. However, the thickness values do not indicate the degree of cementation or the percentage of silt and clay that may be included.

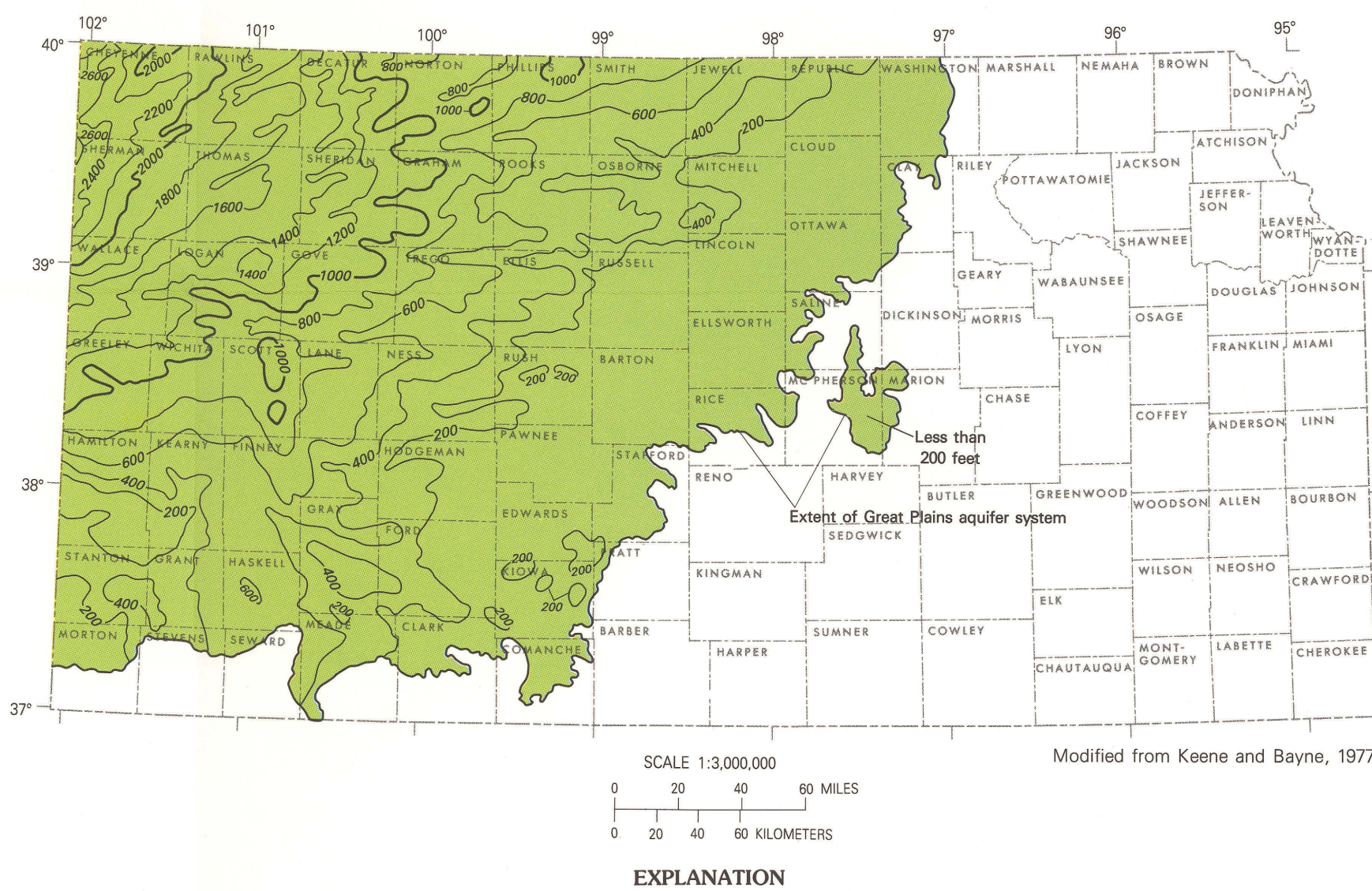


FIGURE 4.—Geohydrologic map showing depth to top of Great Plains aquifer system.

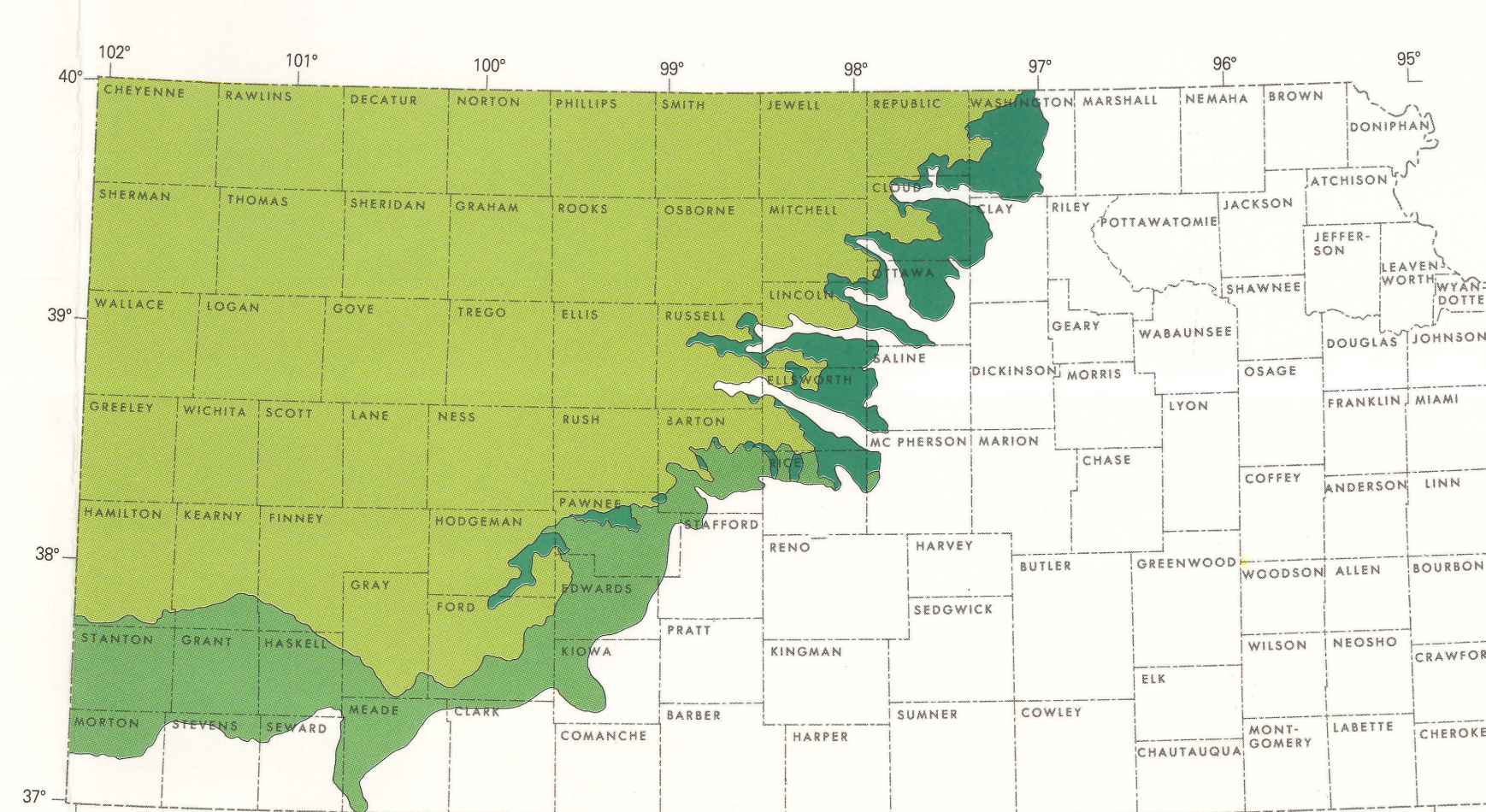


FIGURE 5.—Geohydrologic map showing areal extent of (A) upper aquifer unit, (B) confining unit, and (C) lower aquifer unit of Great Plains aquifer system.

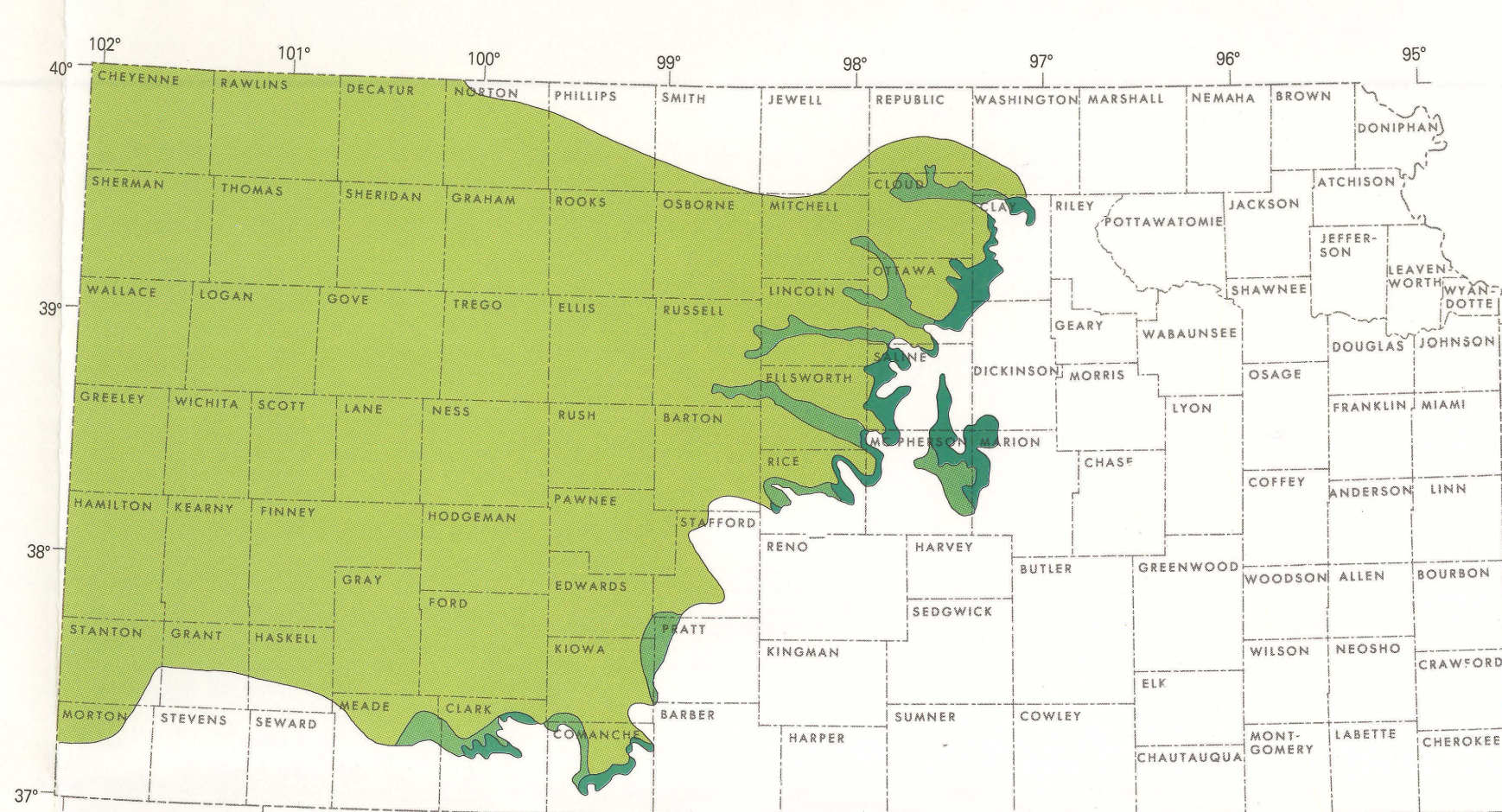


FIGURE 6.—Geologic map showing thickness of sandstone in Great Plains aquifer system.

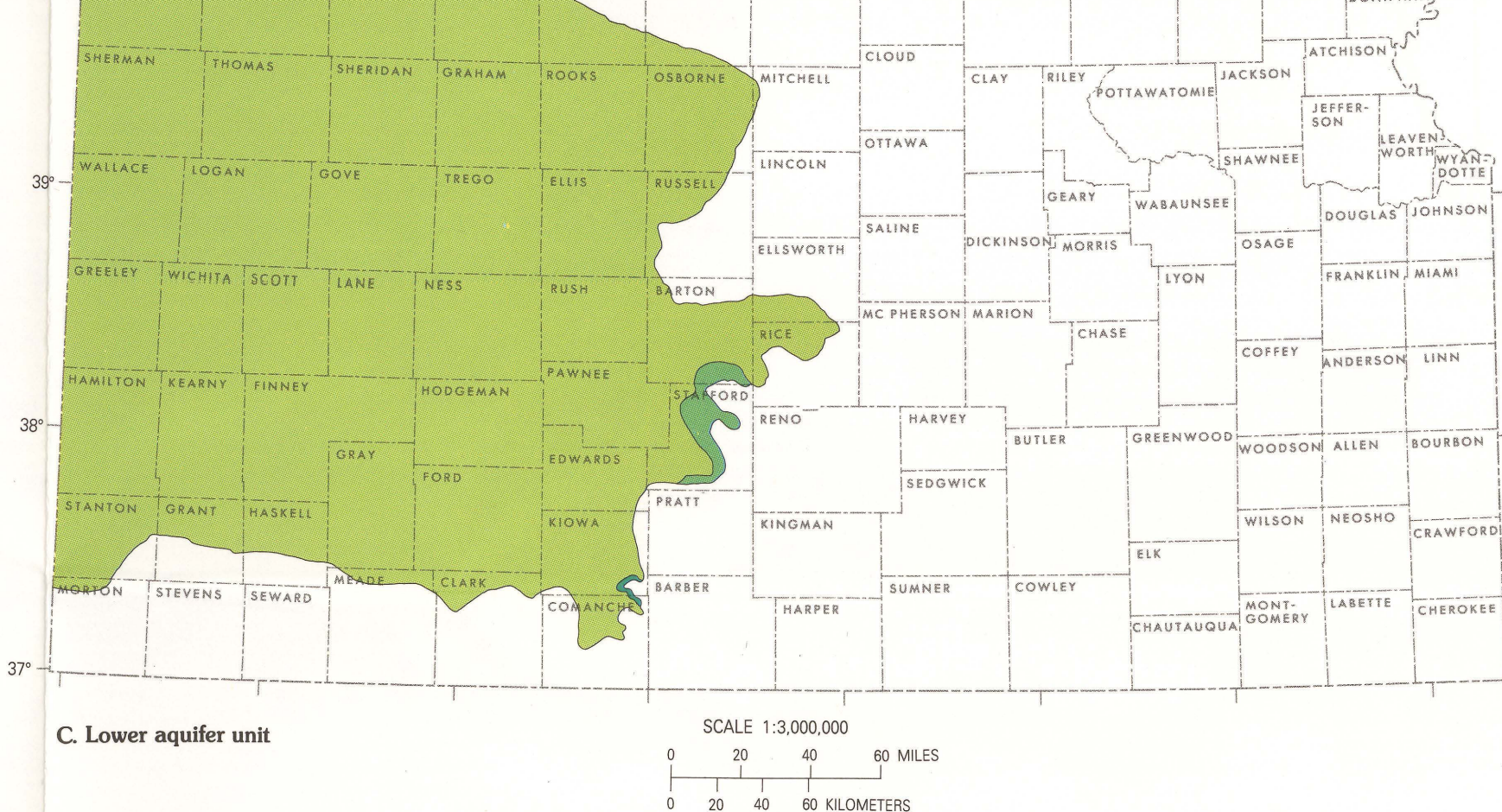


FIGURE 7.—Geologic map showing thickness of sandstone in Great Plains aquifer system.