

Figure 1. Location of study area and intermontane basins.

EXPLANATION

BASIN NAMES

1. Tobacco Valley	26. Beethoven Valley
2. North Fork Flathead River Valley	27. Upper Ruby Valley
3. Lake Creek Valley	28. Madison River Valley
4. Lake Creek Valley	29. Upper Snake Valley
5. Kallispell Valley	30. Root River Valley
6. Laramie River Valley	31. Cottonwood Valley
7. Laramie River Valley	32. Upper Madison River Valley
8. Laramie River Valley	33. Kootenai River Valley
9. Little Bitterroot Valley	34. Upper Snake River Valley
10. Missouri Valley	35. Snake River Valley
11. Snake Valley	36. Prairie River Valley
12. Snake Valley	37. Snake River Valley
13. Snake Valley	38. Rathbun Prairie area
14. Snake Valley	39. Clear Fork Snake River Valley
15. Snake Valley	40. St. Joe River Valley
16. Snake Valley	41. Lang Valley
17. Snake Valley	42. Lemhi Valley
18. Snake Valley	43. Round Valley (Payette)
19. Snake Valley	44. Round Valley (Challis)
20. Snake Valley	45. Palouse Valley
21. Snake Valley	46. Garden Valley
22. Snake Valley	47. Cottonwood Valley
23. Snake Valley	48. Meese Creek Valley
24. Snake Valley	49. Big Lost River Valley
25. Snake Valley	50. Snake River Valley
26. Snake Valley	51. Snake River Valley
27. Snake Valley	52. Snake River Valley
28. Snake Valley	53. Snake River Valley
29. Snake Valley	54. Snake River Valley

Figure 2. Generalized hydrogeology.

CONVERSION FACTORS, VERTICAL DATUM, AND ABBREVIATION

Multiply	By	To obtain
feet (ft)	0.3048	meter
feet squared per day (ft ² /d)	0.0292	meter squared per day
gallon per minute (gal/min)	0.003785	cubic meter per day
gallon per minute (gal/min)	0.003785	liter per second
gallon per minute per foot (gpm/ft)	0.207	liter per second per meter
square mile (mi ²)	2.59	square kilometer

Sea Level: In this report "sea level" refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929)—a geodetic datum derived from a general adjustment of the first-order level lines of the United States and Canada, formerly called the Mean Sea Level of 1929.

Abbreviation: million years m.y.

INTRODUCTION

The Regional Aquifer System (RAS) Program is a series of studies by the U.S. Geological Survey (USGS) to analyze regional ground-water systems that compose a major portion of the Nation's water supply (Ska, 1986). The Northern Rocky Mountain Intermontane Basins is one of the study regions in this national program. The main objective of the RAS studies are to (1) describe the ground-water systems as they exist today, (2) analyze the known changes that have led to the systems' present condition, (3) combine results of previous studies in a regional analysis, where possible, and (4) provide means by which effects of future ground-water development can be estimated.

The purpose of this study, which began in 1990, was to increase understanding of the hydrogeology of the intermontane basins of the Northern Rocky Mountains area. This report is Chapter A of a three-part series and describes the geologic history and the general hydrogeologic units of the Northern Rocky Mountains area. Chapter B (Briar and others, 1996) shows the general distribution of ground-water levels in the basin fill deposits. Chapter C (Clark and Dunbar, 1996) describes the quality of ground-water and surface water in the study area.

This atlas also provides a brief overview of the geologic history from early Precambrian time to the present and describes major tectonic events, depositional environments, and structural features in the study area. This atlas also describes the hydrology, areal extent of outcrops, and water-yielding properties of the hydrogeologic units, with an emphasis on Tertiary to Quaternary basin fill deposits.

LOCATION AND GENERAL FEATURES

The Northern Rocky Mountains Intermontane Basins study area encompasses about 77,000 sq mi in western Montana and central and northern Idaho (fig. 1). The study area extends from the eastern front of the Rocky Mountains in Montana westward to the basal plains of the Columbia Plateau in western Idaho. In the south, the study area extends from the Snake River Plain in Idaho northward to the United States-Canada border. The Continental Divide separates the study area into two major drainage systems—the Missouri River drainage to the east and the Columbia River drainage to the west. Major tributaries of the Missouri River drainage in the study area include the Bearhead, Ruby, Big Hole, Jefferson, Madison, and Gallatin Rivers. Major tributaries of the Columbia River drainage in the study area include the Kootenai, Blackfoot, Bitterroot, Flathead, Clark Fork/Powder/Oreille, Spokane, Salmon, Selway, Lostcha, South Fork Clearwater, and North Fork Clearwater Rivers.

Topography in the northwestern part of the study area is varied. Low surface altitudes range from about 2,000 ft in the Kootenai River Valley in the northwestern part of the study area to more than 12,000 ft in the Lost River Range in the south-central part of the study area. In northwestern Montana and central Idaho, mountain ranges typically are separated by narrow, steep-sided valleys that have little or no basin fill deposits. In contrast, ranges of southwestern Montana and east-central and northern Idaho are separated by wide, relatively level valleys that are deeply filled with basin fill deposits. Valley floor altitudes range from about 2,000 ft in the Kootenai River Valley to about 7,000 ft in the Sawtooth Range in south-central Idaho.

CLIMATE

The climate is characterized by cold winters and mild summers. Annual precipitation ranges from about 8 in. for basins of east-central Idaho to about 100 in. for some mountainous parts of Montana. Most valleys receive about 10 to 30 in. of precipitation per year, with more than one-half falling in winter and spring. Large winter snowpacks in the mountains gradually release their water content as snowmelt that maintains streamflow well into summer.

Spring hydrographs in the study area include 15 generally north-south trending intermountain ranges (or valleys) (fig. 1). For the purpose of this study, "basin" refers to topographic as well as geologic structure. The perimeter of the basin are approximated from topographic geologic structures, extent of basin fill, and results of previous studies. The intermontane basin range in area from less than 10 mi² to more than 700 mi² and are filled with unconsolidated to consolidated Tertiary to Quaternary continental deposits. Intermontane basins comprise about 16 percent of the study area. All basins have through-flowing perennial streams with recent flood plains. In most southern basins, these flood plains are adjacent to older river terraces which grade into pediments or alluvial fans that meet mountain fronts with an abrupt change in slope. In northern basins, recent flood plains are adjacent to glacial deposits which extend to mountain fronts; in some areas, the glacial deposits reach an altitude of as much as 6,000 ft above sea level. Mountain fronts commonly coincide with faults or fault systems along which the basins have been down dropped relative to the mountains.

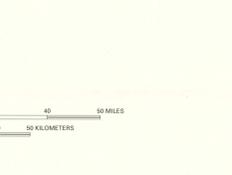
GEOLOGIC HISTORY

The early geologic history of the study area involved several episodes of tectonism and regional metamorphism, erosion, and deposition. Major tectonic events probably occurred about 3,000, 2,750, 1,600, 1,400-1,300, and 900 m.y. (James and Hodge, 1980; James, 1990) producing granitic plutons, gneiss, amphibolite, schist, dolomitic marble, quartzite, and banded iron (fig. 2A). The oldest sedimentary rocks in the study area might be as old as 1,700 m.y. (Ruppel, 1966). These Early to Middle Proterozoic rocks (Yms) form a thick sequence of argillite, siltstone, and sandstone with some volcanic tuffs, which were deposited as turbidites (Hahn and Hughes, 1984). In Middle Proterozoic time, central Idaho was uplifted to form a highland area that created two or more orogenic belts which persisted through the remainder of the Proterozoic. Sediments from this highland were deposited southward to a deep-ocean basin along the western margin of North America in west-central Idaho (Ruppel, 1966) and northward into a basin in western Montana now called the Belt Basin.

GENERALIZED HYDROGEOLOGY

The hydrogeologic units in the Northern Rocky Mountains RASA study area is subdivided into 10 hydrogeologic units. These subdivisions generally are based on major lithology, geologic age, and water-yielding properties. Hydrogeologic units in the study area consist of unconsolidated to consolidated basin fill deposits and consolidated bedrock. Basin fill deposits that are aquifers include Holocene and Pleistocene alluvial deposits (Qal), Pleistocene glacial deposits (Qgl), Quaternary and Tertiary unconsolidated deposits (Qtu), and Tertiary sedimentary deposits and rocks (Ts). Bedrock units that are aquifers in the study area include Quaternary through Cretaceous extensive rocks (Qtrk), Tertiary through Cretaceous intrusive rocks (Tki), Mesozoic clastic and carbonate rocks (Mcl), Paleozoic carbonate and clastic rocks (Pcl), Middle Proterozoic metamorphic rocks (Yms) and Cretaceous through Archean metamorphic rocks (Kcm). Clastic and carbonate rocks (Mcl) and metamorphic rocks (Kcm) can yield water locally, but generally their water-yielding properties are unknown.

CORRELATION OF HYDROGEOLOGIC UNITS



DESCRIPTION OF HYDROGEOLOGIC UNITS

MAJOR LITHOLOGY	AREAL EXTENT OF OUTCROPS	WATER-YIELDING PROPERTIES
Qal ALLUVIAL DEPOSITS (Holocene and Pleistocene)—Unconsolidated stream-laid gravel, sand, silt, and clay that is poorly to moderately well-sorted; includes some talus, cobblestone, boulders, and placer deposits, mine tailings, and Pleistocene glacial outwash.	Mostly present along present-day stream channels, flood plains, and low-level terraces near rivers in larger basins and as smaller deposits near perennial and ephemeral streams. In mountainous areas, Qal is found as narrow deposits that are typically vertically extensive. In most mountainous regions these deposits are not shown on figure 2.	Deposits yield abundant water to wells throughout the flood plains and low-level terraces near rivers in larger basins and as smaller deposits near perennial and ephemeral streams. In mountainous areas, Qal is found as narrow deposits that are typically vertically extensive. In most mountainous regions these deposits are not shown on figure 2.
Qgl GLACIAL DEPOSITS (Pleistocene)—Unconsolidated gravel, sand, silt, and clay that is poorly to moderately well-sorted; includes some talus, cobblestone, boulders, and placer deposits, mine tailings, and Pleistocene glacial outwash.	Extensive deposits are present in the northern half of the study area where the Cordilleran ice sheet and mountain glaciers covered a large part of the study area. Qgl is not shown on figure 2. Scattered deposits of Qgl are the result of meltwater from the study area near the basin margins or in mountainous areas, except for the Cordilleran Valley and the Broad Valley (Payette), which contain Qg through their extent.	Generally, a limited source of water to wells owing to the fine-grained material and poor pore connectivity. Yields are variable and might depend on interbedded gravel. Yields range from 0.2 to 1.500 gal/min/ft with a median of 0.2 gal/min/ft. Reported estimates of transmissivity include: 3.4 to 18 ft ² /d in the Little Bitterroot Valley (Shupe, 1988); 2.6 to 13 ft ² /d in the Upper Clark Fork Valley (McMurry and others, 1968); 2.0 to 3.0 ft ² /d in the Bitterroot Valley (McMurry and others, 1972); 1.3 to 10 ft ² /d in the Missouri Valley (McMurry and others, 1965).
Qtrk UNDIFFERENTIATED DEPOSITS (Quaternary and Tertiary)—Unconsolidated to consolidated gravel, sand, silt, and clay, and locally contains lignite and limestone. Includes colluvium and landslide deposits, extensive alluvial fan deposits within and outside of basins, alluvium on some alluvial fans, Tertiary deposits and rocks with a veneer of Quaternary deposits. Includes areas where Quaternary and Tertiary deposits are indistinct.	Extensive deposits are present in central Idaho where Qtrk covers most basins from the present-day stream channels, flood plains, and low-level terraces near streams and rivers to the mountain margins. In southwestern Montana, Qtrk is present mostly near the basin margins or as scattered residual remains where streams and rivers dissect these deposits.	Yields are variable and might depend on hydraulic interconnection with more permeable units and extent of interbedded gravel. Yields range from 0.2 to 3.420 gal/min/ft with a median of 0.2 gal/min/ft. Reported estimates of transmissivity include: 3.2 to 18 ft ² /d in the Little Bitterroot Valley (Shupe, 1988); 2.6 to 13 ft ² /d in the Upper Clark Fork Valley (McMurry and others, 1968); 2.0 to 3.0 ft ² /d in the Bitterroot Valley (McMurry and others, 1972); 1.3 to 10 ft ² /d in the Missouri Valley (McMurry and others, 1965).
Tki SEDIMENTARY DEPOSITS AND ROCKS (Upper Tertiary through middle to upper Eocene)—Unconsolidated to consolidated gravel, sand, silt, clay, ash, ash fall, and tuffs, and basaltic tuffs that is poorly to moderately well-sorted. Locally includes massive conglomerate, sandstone, siltstone, shale, and interbedded volcanic and volcanoclastic rocks, lignite, and limestone.	The most extensive outcrops, present in central Idaho and west-central Montana, are part of the Idaho Batholith. Dissected by the Snake River and other basins, outcrops between the Three Forks area, and northern Idaho near Clear Lake.	Yields are variable and depend on the occurrence and extent of fractures, faults, breccias, and permeable rocks. Yields are variable and might depend on interbedded gravel. Yields range from 0.2 to 3.420 gal/min/ft with a median of 0.2 gal/min/ft. Reported estimates of transmissivity include: 3.2 to 18 ft ² /d in the Little Bitterroot Valley (Shupe, 1988); 2.6 to 13 ft ² /d in the Upper Clark Fork Valley (McMurry and others, 1968); 2.0 to 3.0 ft ² /d in the Bitterroot Valley (McMurry and others, 1972); 1.3 to 10 ft ² /d in the Missouri Valley (McMurry and others, 1965).
Ttr EXTENSIVE ROCKS (Quaternary through Cretaceous)—Pyroclastic rocks, lava flows, ash-flow tuffs, welded tuffs, and breccias that range in composition from tholeiitic to basaltic. Includes water-lain, interbedded volcanoclastic rocks such as tuffaceous conglomerate, coarse-grained sandstone, and mudstone.	Most outcrops are found in four general areas: central Idaho where large outcrops extend to the southern edge of the study area, western and northwestern Montana where scattered rocks crop out, west-central Montana between Oroville and the Bunte Three Forks area, and northern Idaho near Clear Lake.	Yields are variable and depend on the occurrence and extent of fractures, faults, breccias, and permeable rocks. Yields are variable and might depend on interbedded gravel. Yields range from 0.2 to 3.420 gal/min/ft with a median of 0.2 gal/min/ft. Reported estimates of transmissivity include: 3.2 to 18 ft ² /d in the Little Bitterroot Valley (Shupe, 1988); 2.6 to 13 ft ² /d in the Upper Clark Fork Valley (McMurry and others, 1968); 2.0 to 3.0 ft ² /d in the Bitterroot Valley (McMurry and others, 1972); 1.3 to 10 ft ² /d in the Missouri Valley (McMurry and others, 1965).
Mcl CLASTIC AND CARBONATE ROCKS (Undifferentiated Mesozoic)—Primarily marine shales and siltstones, quartzite, siltstone, mudstone, and marine limestone.	The most extensive outcrops, present in central Idaho and west-central Montana, are part of the Idaho Batholith. Dissected by the Snake River and other basins, outcrops between the Three Forks area, and northern Idaho near Clear Lake.	Yields are variable and depend on the occurrence and extent of fractures, faults, breccias, and permeable rocks. Yields are variable and might depend on interbedded gravel. Yields range from 0.2 to 3.420 gal/min/ft with a median of 0.2 gal/min/ft. Reported estimates of transmissivity include: 3.2 to 18 ft ² /d in the Little Bitterroot Valley (Shupe, 1988); 2.6 to 13 ft ² /d in the Upper Clark Fork Valley (McMurry and others, 1968); 2.0 to 3.0 ft ² /d in the Bitterroot Valley (McMurry and others, 1972); 1.3 to 10 ft ² /d in the Missouri Valley (McMurry and others, 1965).
Pcl CARBONATE AND CLASTIC ROCKS (Undifferentiated Paleozoic)—Primarily marine shales and siltstones, quartzite, siltstone, mudstone, and marine limestone.	Most of these rocks are present as isolated outcrops east of longitude 117°15' in the troughs of the Snake River and other basins, and as scattered outcrops along the frontal and three-took and Dearborn Belt.	Yields are variable and depend on the occurrence and extent of fractures, faults, breccias, and permeable rocks. Yields are variable and might depend on interbedded gravel. Yields range from 0.2 to 3.420 gal/min/ft with a median of 0.2 gal/min/ft. Reported estimates of transmissivity include: 3.2 to 18 ft ² /d in the Little Bitterroot Valley (Shupe, 1988); 2.6 to 13 ft ² /d in the Upper Clark Fork Valley (McMurry and others, 1968); 2.0 to 3.0 ft ² /d in the Bitterroot Valley (McMurry and others, 1972); 1.3 to 10 ft ² /d in the Missouri Valley (McMurry and others, 1965).
Yms METASEDIMENTARY ROCKS (Undifferentiated Middle Proterozoic)—Primarily argillite, siltstone, and sandstone, and locally contains lignite and limestone. Includes local mud flows, dikes, and sills. Rocks have been subjected to low-grade metamorphism.	Metasedimentary rocks primarily crop out in the northern half of the study area. These rocks are also found in east-central Idaho, along the Snake River and other basins, and in the northern part of the Intermontane Range to about Hannock Pass in the Bitterroot Valley.	Metasedimentary rocks primarily crop out in the northern half of the study area. These rocks are also found in east-central Idaho, along the Snake River and other basins, and in the northern part of the Intermontane Range to about Hannock Pass in the Bitterroot Valley.
Kcm METAMORPHIC ROCKS (Undifferentiated Cretaceous through Archean)—Metamorphosed granitic plutons, dikes and sills, quartzite-feldspate, gneiss, amphibolite, schist, dolomitic marble, quartzite, and iron formation.	Metamorphic rocks primarily crop out in southwestern Montana between Yellowstone National Park, and Idaho, and in northern Idaho. Large outcrops are present in central Idaho and are associated with Tki of the Idaho Batholith. Scattered outcrops are also present in northern Idaho.	Metamorphic rocks primarily crop out in southwestern Montana between Yellowstone National Park, and Idaho, and in northern Idaho. Large outcrops are present in central Idaho and are associated with Tki of the Idaho Batholith. Scattered outcrops are also present in northern Idaho.

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GEOLOGIC HISTORY AND HYDROGEOLOGIC UNITS OF INTERMONTANE BASINS OF THE NORTHERN ROCKY MOUNTAINS, MONTANA AND IDAHO

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1996

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