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MAP SHOWING LOCATION OF STRIP MAP SEGMENTS

MAP SHOWING RECENTLY ACTIVE BREAKS ALONG THE SAN ANDREAS FAULT AND RELATED FAULTS BETWEEN THE NORTHERN GABILAN RANGE AND CHOLAME VALLEY, CALIFORNIA

PURPOSE OF THIS STRIP MAP

This strip map is one of several designed to show the location of the most recent movements on the San Andreas fault. The character and location of the most important tectonic and engineering geologic features and earthquakes and should be helpful to those concerned with land use and development or near the fault. Lines on the strip map mark displacements of the ground surface by rupture or creep along the San Andreas fault. Map users should keep in mind that these lines are primarily guides for locating fault breaks on the ground and are not necessarily located with the precision demanded by every engineering or land utilization need.

THE SAN ANDREAS FAULT AND FAULT ZONE

The San Andreas fault zone is a major flow or break in the earth's crust. It extends from the Gulf of California in northern Mexico northward through western California and under the Pacific Ocean off Point Arena in northern California, a distance of more than 600 miles. Movement within the fault zone has been distributed along many nearly parallel faults that differ in age and in amount of total displacement. This complex zone of movement—active over a span of millions of years and marked by displacements measurable in feet or hundreds of miles—varies in width from a few hundred feet to several miles. It is termed the San Andreas fault zone (or rift zone), as distinguished from the San Andreas fault, which consists of the trace of the most recent (in place, historic) movement (Noble, 1926, p. 415-417).

FIELD RECOGNITION OF MOST RECENT FAULTING

The most recent fault breaks are generally recognized by topographic discordance or a contrast in vegetation reflecting varying depth to groundwater or soil differences across the fault trace. The most common features are scarps, trenches, notches, ridges, offset streams, sag ponds, and other small undulating depressions, lines of springs or trees. These features have developed in complex and different ways, but they are all controlled by repeated movements along the fault, or by erosion along its trace. Horizontal and vertical displacements of a few inches or a few feet accumulate from successive displacements accompanied by earthquakes from slow tectonic creep between earthquakes, or from a combination of both. Whether they are caused by creep or by sudden movements, the displacements produce scarps and other topographic features that delineate the fault lines shown on this map. The brief notes along the fault traces on the map indicate especially clear examples of these features; these are not limited to spots indicated, but are present in some degree along the entire length of the mapped fault breaks. As the edges of opposing fault blocks slide by one another, topographic features are juxtaposed to form steep, sag ponds, or low ridges. Notches and trenches along the fault line commonly reflect increased erosion of the less resistant eroded and broken rock of the fault zone, or they may be down-dropped alluvial ridges between parallel breaks in the fault zone segments, which are 2 to 6 miles long, arranged in an almost but overlapping fashion as shown on the strip map. The segments are essentially parallel to the fault, which is relatively broad zone. Between Bitterwater Valley and Lewis Creek (Harradine Valley quadrangle), unusual drainage patterns and low escarpments suggest that cross faults are present. In the Coast Range of northern California, fault-block features may be identified by their characteristic sharp, vertical movements along the modern San Andreas fault are also known, but historically these have been small and localized as compared to right slip.

LOCATION OF THE FAULT BREAKS

The fault breaks were located by study of 1:250,000 scale aerial photographs flown on June 22, 1966, and were transferred to topographic maps by visual inspection. On this map some breaks were located on the ground mapping following an earthquake of moderate magnitude accompanied by significant movement in 1966.

Map users should consider a line on the map not as a precisely located boundary for field location of fault-block features. Where such features are large enough or distinctive enough to be shown by the contours of the topographic map, the fault break is accurately located within 100 feet laterally. Where these features are more subtle or where topographic maps show comparatively little detail, the mapped line may be as much as

500 feet from the actual fault break, and in areas of featureless topography, the accuracy of location may be even less. Consulting geologists, engineers, and others making specific and direct use of these strip maps will need to make ground surveys to confirm and refine the positions of these fault lines in relation to structures and land boundaries.

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as hazardous by builders, planners, engineers, public works agencies, utility companies, homeowners, landowners, developers, public works departments, school boards, civil defense officials—in short, anyone concerned with present structures, land use, or planned construction on or near these most recent fault breaks.

Sudden movement along the San Andreas fault in 1857 in northern California and again in 1906 in the northern California area produced disastrous earthquakes. In 1906, horizontal displacement across the fault was as much as 20 feet, and that in 1857 was probably comparable. Miscellaneous tectonic creep along the San Andreas fault in the Hollister area (Tucker, 1950) is slowly tearing apart a building that straddles an active break in the fault zone. Creep has displaced railroad tracks, a cistern, a water tunnel, and buildings along the Hayward fault zone (Radford and others, 1966). Within the area of this strip map, tectonic creep along recent fault breaks is shown by the deformation of fences and by failures in the surface of gravel roads (Brown and Walker, 1965). The ground fractures of the San Andreas fault during the 1966 Parkfield-Cholame earthquake followed almost exactly previously mapped fault breaks (Brown and others, 1967, p. 4). Studies of the San Andreas fault in the Carrizo Plain area by Wallace (1965) show that during approximately the last 10,000 to 20,000 years, displacements occurred over and over again along the same break of the San Andreas fault that was active in 1857. Simply summed up, these observations show that the line of most recent ground breakage has a good chance of breaking again during another major earthquake along this segment of the fault.

At present, no one can accurately predict when movement on these faults will occur or which segments will move next, but it is certain that some will move again in the future. Movement will not necessarily be confined to mapped faults and fractures, and indeed surface fracturing could develop anywhere within the fault zone, and even on branching or otherwise related faults beyond the fault zone.

SPECIAL FEATURES OF THE SEGMENT FROM THE NORTHERN GABILAN RANGE TO CHOLAME VALLEY

Because of the nature of deformation along the San Andreas fault northwest of Middle Mountain (Parkfield and San Miguel quadrangles) is different from that to the southeast, the fault features in the mapped area are described separately.

North of Middle Mountain

The line of faulting in the San Andreas fault of small-scale maps—trends about N. 40° W. and is very nearly straight. This line is defined by fault segments, from 2 to 6 miles long, arranged in an almost but overlapping fashion as shown on the strip map. The segments are essentially parallel to the fault, which is relatively broad zone. Between Bitterwater Valley and Lewis Creek (Harradine Valley quadrangle), unusual drainage patterns and low escarpments suggest that cross faults are present. In the Coast Range of northern California, fault-block features may be identified by their characteristic sharp, vertical movements along the modern San Andreas fault are also known, but historically these have been small and localized as compared to right slip.

of Middle Mountain. Another linkage mechanism is indicated where a reach road crosses the ridge crest separating Panocho Rico Creek from the upper part of Black Canyon (Panocho Rico quadrangle); there a very short (about 600 feet) fault segment separates the opposed ends of two larger reaches and parallels them. The short nodal fault segment seems to provide dislocation continually by partly bridging the gap between the two opposed reaches, where the fault crosses Nelson Creek (San Miguel quadrangle).

The fault segments northwest of Middle Mountain vary considerably in arrangement and pattern, but when they are viewed along their trend, a significantly larger number of effects are seen to be stepped to the right. Evidence of current and historic movement along individual segments is abundant from Bitterwater Valley north to the vicinity of Palmdale. Very recent movement is shown by erosion fractures in the Airline Highway at several places where it is crossed by active fault segments; at most of these locations the highway has been patched many times. The patches vary in length, but all are localized in a narrow belt that coincides with the fault. Movement is also displayed by fences that cross the active segments of the fault. Most of these fences exhibit more than a foot of right slip at the fault; that is, the fence segments on the northeast side of the fault are displaced toward the southeast, because the other fences exhibit the greatest displacement. It is clear that fault movement has been repeated, or essentially continuous, for at least half a century. Both the fractured zones in the Airline Highway and the offset fence localities are shown on the strip map. More detailed information and additional interpretive discussion are afforded by Brown and Wallace (1968), who concluded that this part of the fault is moving at about 0.5 inch per year, and that the movement is essentially localized along the recently active fault breaks shown here.

From Bitterwater Valley south to Middle Mountain, evidence of current and historic fault movement is sparse. This is partly because the area is sparsely inhabited and only a few man-made features cross the fault, but it is also because all movement on steep slopes and narrow ledges tends to deform both the fault trace and any linear man-made features that cross the trace. Despite the scarcity of evidence, the similarity between fault-formed surface features here and those farther north suggest that historic and pre-historic movement rates are comparable.

Between Middle Mountain and the Coast Range, some segments are apparent at several places north of Middle Mountain. Broad zones of faulting characterized by two or more parallel breaks are evident along the north-west-trending ridge crest north of Blue Valley (San Benito quadrangle), along Mustang Ridge north of the Airline Highway about a quarter of a mile north of the James Ranch (San Benito quadrangle), and along the ridge crest between Pointed Camp Ridge and Lewis Creek (Harradine Valley quadrangle). Between Bitterwater Valley and Lewis Creek (Harradine Valley quadrangle), unusual drainage patterns and low escarpments suggest that cross faults are present. In the Coast Range of northern California, fault-block features may be identified by their characteristic sharp, vertical movements along the modern San Andreas fault are also known, but historically these have been small and localized as compared to right slip.

most of the S. to E. fault segments northwest of Middle Mountain are not joined in an obvious way to the neighboring segments, although presumably they are all part of the San Andreas fault system. The fault segments are separated by small gaps, and are typically no more than a few hundred feet apart where they overlap, but exceptions are numerous.

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Recovered triangulation stations (Meads and Small, 1961) and geodimeter measurements (Hoffman and others, 1967) show that the earth's crust southwest of the line of active breaks is clearly moving northeast relative to the crust on the northeast side. Most of the movement is concentrated along the mapped fault segments (Brown and Wallace, 1968), and surface displacements that to some extent are independent (even though they are apparently) continuously between apparently unjoined fault segments may be significant. Continuity between apparently unjoined fault segments may be significant. Continuity between apparently unjoined fault segments may be significant.

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LAND USE SIGNIFICANCE OF LOCATING RECENT FAULT BREAKS

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