

EXPLANATION

Log data
Bedrock not penetrated
Published data
Bedrock not penetrated
Quantity of outcrop
Sandstone
Shale
Bedrock contour
Show altitude of bedrock surface.
Dashed where approximately located.
Contour interval 100 ft.
Datum is mean sea level.

INTRODUCTION

The bedrock is known (Henshaw, 1937) to be generally overlain by unconsolidated deposits consisting of glacial drift, alluvium, and loess. Loess of only one foot thickness underlies the surface in some places. The total thickness of the unconsolidated deposits is estimated to be 10 to 100 feet. The configuration of the underlying bedrock surface is the result of a complex system of erosion that was developed during the late Pleistocene and early Holocene, and is characterized by a period of preglacial erosion and during shorter, but more intense, glacial periods. The bedrock topography is the result of the interaction of these two periods of erosion.

The following conversion factors can be used to convert topographic units to the International System of Units (SI):

Foot (ft)	0.3048	meter (m)
Gallon per minute (gal/min)	0.0038	cubic meter per second (m ³ /s)

BEDROCK TOPOGRAPHY

Primary control for the map is topographic data and information from quarries, outcrops, and seismic refraction surveys. Published data (Cagle, 1937) provide additional control but are not as accurate as the topographic data. More detailed information concerning the control data is available in the files of the Iowa Geological Survey and the U.S. Geological Survey, Iowa City, Iowa.

The accuracy of the map is directly related to the density of control points; the greater the number of points in a given area, the more reliable the placement of the contour. Because of the lack of control points in eastern and north-central parts of the map area, dashed contours reflecting a lack of control are used occasionally. Much of the interpretation of the bedrock topography along the eastern edge is inferred from comparison with an adjacent bedrock topographic map (Cagle, 1937). Dashed contours are used in other parts of the map where it seems reasonable to continue a ridge or valley, but where no control point was available to confirm the contour.

The principal features of the bedrock surface are the deeply bedded bedrock channels and the generally level bedrock-floored stream valleys. Several bedrock channels are located in eastern and western parts of the map area; the most prominent of these is the Fremont Channel which extends northward through Fremont, Mills, and Pottawattamie counties. The bedrock channels characteristically are (1) relatively wide with gently sloping walls where they have been cut in easily erodible, dominantly glacial drift; (2) narrow and steep walled where they have been cut in more resistant, dominantly carbonate bedrock. Bedrock spurs from the divides between and adjacent to the bedrock channels; they have a maximum thickness of about 1200 ft.

The bedrock surface in the central part of the map has been sculptured primarily by preglacial erosion. The streams have deeply incised the bedrock surface and have extended their valleys by bedrock erosion. These bedrock-floored valleys tend to be narrow and steep walled. Cross-sections of the bedrock surface are situated over bedrock channels and, because they have not eroded into bedrock, relatively broad valleys have developed.

The bedrock channels in the map area are the bedrock remnants of ancient bedrock drainage courses that continue into Missouri. Henshaw and How (1933) applied the name Adams Valley, Grand River Valley, and Mayville Valley to bedrock valleys in Missouri that originate in or cross the map area. In the present report the names Adams and Mayville have been retained, but the name Fremont has been applied to Grand River. Furthermore, in the present report, the term "channel" is used instead of "valley" to conform to general usage in Iowa reports.

USES OF THE MAP

The bedrock map, when used in conjunction with land-use information, is a valuable tool in hydrologic, environmental, and geologic problems.

The bedrock map is an aid in locating supplies of ground water. The sources of these supplies are water-bearing sand and gravel in the deposits of glacial drift and alluvium that underlie the bedrock.

Detailed data show a relationship between the physical features of the bedrock surface and the occurrence of ground water in the underlying surficial deposits. The areas most favorable for the development of supplies are bedrock channels and the alluvial deposits of these channels that may or may not have reached the bedrock.

Detailed bedrock channels often are favorable sites for obtaining ground-water supplies at relatively shallow depths because they frequently contain extensive sand-and-gravel deposits. Recent hydrologic studies of the drainage basin of the Mayville Channel confirmed the presence of water-bearing sand-and-gravel deposits in parts of the Adams Valley (Cagle, 1937). The bedrock data available for the bedrock channels in other parts of the map area.

Although sand-and-gravel aquifers often are distributed irregularly and are not present at all places within the bedrock channels, they are more abundant in bedrock channels than in areas underlain by bedrock channels than in areas underlain by bedrock channels. The distribution of these aquifers is difficult to predict and unpredictable in their occurrence. Test drilling should precede the development of ground water from these aquifers. Yields from bedrock channels generally range from 1 to 15 gallons per minute from individual wells tapping bedrock sand-and-gravel aquifers. The bedrock surface is the result of a complex system of erosion that was developed during the late Pleistocene and early Holocene, and is characterized by a period of preglacial erosion and during shorter, but more intense, glacial periods. The bedrock topography is the result of the interaction of these two periods of erosion.

ACKNOWLEDGMENTS

Particular recognition is given to the present and past members of the Iowa Geological Survey who over a period of many years have collected and analyzed data on bedrock topography, stream channels, and compiled other information necessary to the preparation of this map. Acknowledgment is also due to the Department of Geology, Iowa State University, who assisted in compiling much of the data. Acknowledgment is made to the many well-drilling contractors who have voluntarily collected drill cuttings and provided other well data.

SELECTED REFERENCES

Amey, M. F., 1916. Geology of Taylor County, Iowa. Geological Survey Annual Report, v. 21, p. 277-344.

Cagle, J. W., 1937. Bedrock topography of south-central Iowa. Iowa Geological Survey Bulletin, 18, 27 p.

Cagle, J. W., 1939. Geology of Page County, Iowa. Geological Survey Annual Report, v. 11, p. 297-460.

Cook, E. E., 1909. The nonconformity, topographic and correlation of the Pennsylvanian subdivisions in eastern Nebraska and western Iowa. Nebraska Geological Survey Bulletin, 16, 27 p.

Cook, E. E., and Upp, U. E., 1933. The Red Oak-Semert-Lewis nonconformity. Nebraska Geological Survey, Paper 3, 23 p.

Cook, E. E., and Reed, E. C., 1938. The Red Oak-Semert-Lewis nonconformity. Nebraska Geological Survey, Paper 12, 19 p.

Geological Society of Iowa, 1964. Guidebook Southwestern Iowa. Iowa Geological Survey, Iowa City, Iowa.

Geological Survey Annual Report, v. 27, p. 277-344.

Henshaw, G. E., and How, B. B., 1933. Map of bedrock topography of northwestern Missouri. In *Groundwater resources of Missouri*. Missouri Geological Survey and Water Resources.

Henshaw, J. B., and others, 1963. Preliminary interpretation of an aeromagnetic survey in eastern and northwestern Iowa. Iowa Geological Survey Open File Report, 29 p.

Henshaw, H. G., 1969. Geologic map of Iowa. Iowa Geological Survey, Iowa City, Iowa.

Henshaw, H. G., and others, 1960. Highway construction materials from the consolidated rocks of northwestern Iowa. Iowa Highway Research Bulletin, No. 15, 151 p.

Kent, C. R., 1898. Carboniferous stratigraphy of Montgomery County, Iowa. American Geologist, v. 21, p. 346-350.

Moore, R. C., 1948. Geology of Montgomery County, Iowa. Iowa Geological Survey Annual Report, v. 4, p. 381-451.

Murray, R. L., 1964. Geology of the Council Bluffs area, Nebraska-Iowa. U.S. Geological Survey Professional Paper 472, 72 p.

Moore, R. C., 1948. Classification of Pennsylvanian rocks in Iowa, Missouri, Nebraska, and northern Oklahoma. American Association of Petroleum Geologists Bulletin, v. 32, no. 11, 1937.

Nelson, W. H., and others, 1912. Underground water resources of Iowa. Iowa Geological Survey Annual Report, v. 21, p. 1100-1195.

1912. Underground water resources of Iowa. U.S. Geological Survey Water-Supply Paper 993, p. 807-865.

Palmer, M. C., 1971. Reservoir of oil exploration and potential in Iowa. Iowa Geological Survey Information Circular No. 2, 11 p.

Rubin, V., and others, 1967. Landscape evolution and soil formation in northwestern Iowa. Soil Conservation Service, U.S. Department of Agriculture Technical Bulletin, 1399, 282 p.

Sanderson, L. V., and others, 1971. Interrelationships of surface and subsurface geology in the Nebraska-Iowa area. Iowa State Water Resources Research Institute, Project No. A-630-IA, 179 p.

Shank, B., 1911. Pleistocene of the vicinity of Omaha, Nebraska. Council Bluffs, Iowa (Iowa Geological Society America Bulletin, v. 22, p. 730).

Turnell, J. J., and others, 1972. Hydrogeologic considerations in solid waste storage in Iowa. Iowa Geological Survey Public Information Circular No. 4, 59 p.

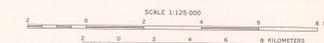
Udden, J. P., 1898. Geology of Montgomery County, Iowa. Geological Survey Annual Report, v. 11, p. 199-277.

Udden, J. P., 1900. Geology of Montgomery County, Iowa. Geological Survey Annual Report, v. 13, p. 123-183.

Wright, W. W., 1938. The Red Oak-Semert-Lewis nonconformity in Iowa. Iowa Geological Survey Annual Report, v. 36, p. 7-310.

1943. Geology of Adams County, Iowa. Iowa Geological Survey Annual Report, v. 37, p. 263-373.

BEDROCK TOPOGRAPHY OF SOUTHWEST IOWA
By
LYLE V. A. SENDLEIN AND JACK L. GILMORE
1980



Not from U.S. Geological Survey 1:250,000
National City, 1880 and Ocala, 1954