[illegible]

Massasa Petrel. The austral petrel in the area, deserves more detailed comment. It has been raised as a rim rock by pyroclastic deposits ejected from the caldera (Reynolds and others, 1980). Lava flows, or intercalations of the two; the bands and flaments of dark material that radiate from Massaso are probably lava flows, but pyroclastic rocks may also be present (Reynolds and others, 1980; Strom and Schneider, 1982). Eruptions and effusions of rim, dark, and peripheral materials produced an initial ashfall that was followed by a series of smaller eruptions that produced additional surface layers of the local slumping of the walls. Sulfur-rich materials were deposited on the floor during the waning stages. The relief from the rim to the lower floor of Massaso suggests materials of considerable strength (Covacek and Carr, 1980) and does not support a postulated near-surface ocean of sulfur overlain by an upper crust, possibly kilometers thick, of porous sulfur and liquid sulfur dioxide (Smith and others, 1979a). If this sulfur ocean had existed, the whole pattern would have collapsed because of the high temperatures, and the sulfur would have been converted to hydrogen sulfide gas. A hot sulfur pool at the center of the caldera that a hot saline plume, as postulated by McEwen and Soderblom (1983), is at or near the surface in most places in the Massasa area.

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STRUCTU

Structural deformation, abetted by high temperatures, probably produced the vermicular ridged material. The southeastern part of this unit, the smooth facies, may be interpreted to be an uplifted diamond-shaped block of plateau and crustal materials with deformed margins. Its ridged facies, deformed by thermal creep, unconformably overlies ridged and ineated plateau materials. The elevated planar surface of the smooth facies is tilted to the northwest and, at places outside the map area where one would expect to find a reversal of slope and more ridges and lobate fronts, one finds instead large down-faulted slices, scarps about 1 to 3 km high, and hummocky surfaces of slumped blocks resting on plateaus.

COMPOSITIONS OF GEOLOGIC UNITS

Colors and ultraviolet reflectances of *a* are consistent with a surface composed of other varieties of alltopores of sulfur admixed with sulfur dioxide in variable amounts (Soderbom and others, 1980). Other factors, such as the temperature of the sulfur, grain size, and surface roughness, may later have alternative conclusions, but the photometry of the orange regions is consistent with sulfur (Grady and Fevejer, 1986). Although we have no spectroscopic evidence for the presence of sulfur dioxide, it is likely that some of the sulfur was oxidized during the eruption. The dark gray material at the base of the steeply sloping walls in the order of 2.2 km. For a local heat flow in excess of 1 to 2 W/m² (Pearl and Stoney, 1982), which might be expected for Massaw, sulfur would melt at depths about 60 m, where the silicates would be stable (Clow and Can, 1986). Marginal stability of the patina slope is implied by the low thermal conductivity of the sulfur.

The light-colored, rounded, and elongated material unit also suggest silicate-rich materials. Laboratory tests (Kane and others, 1987) indicate that the strength of a material increases with increasing grain size (3 to 20 cm), estimated pore volume (10 to 20%), an acceleration of gravity of 1.8 m/s², and assumed density of 2,000 kg/m³ imply yield strengths on the order of 10 to 100 psf, if the materials are plastic. Here again, pure sulfur appears improbable and silicates, or possibly sulfur-silicate mixtures are more likely.

appear to be more probable (Moore and others, 1978).
The mineralogical composition of the units that sustain high local relief in scarp and ridges—the lined plateau and mountain materials. These units are spectrally similar and mostly light grayish yellow green to greenish gray; the moderate grayish orange of parts of the vermicular ridged mountain unit could be caused by a local sulfur-rich veneer. The composition of smooth plateau materials is similar to that of the lined plateau materials. The moderate grayish-orange lowland unit and Euboea Fluctus central material have little or no local relief and could be ponded sulfur and sulfur flows (Sagan, 1979), the similarly colored Euboea Fluctus marginal material could be sulfur and sulfur flows. The dark grayish brown of the lowland and ridge materials could be a mixture of sulfur and sulfur dioxide. Dark orange-brown units such as Maassov's dark and rim materials could be explained as mixtures of red and brown sulfur and sulfur dioxide, but the relief of the patera rim suggests that silicates may be dominant. White material could be sulfur and sulfur dioxide, but the relief of the patera rim could be brown sulfur or silicates admixed with considerable amounts of sulfur dioxide.

GEOLOGIC PROCESSES

In view of its active volcanism, the large observed heat flows, and probable tides (Poole and others, 1976), most surface modifications are likely due to a combination of magmatic-volcanic processes, tectonism, and mass wasting abetted by high temperatures. For example, the following scenario might account for formation of the southeastern vermicularly ridged mountains. Lavas flowing from a caldera at the present site of Cradine Plateau (fig. 3) across the lineated plateau unit and onto the adjacent low-lying plateau, which was covered by a thin layer of ash, would have been impeded by the subjacent crustal materials, elongate to the northeast, was uplifted by interactions of tidal forces in the crust and magma at depth. Initial tilt of the fault block was to the northwest. Because of a large local heat flow beneath the tilted block, isotherms within it corresponding to temperatures required for plastic flow were asymmetrical (nearer to the surface on the northwest side than on the southeast side). This would have permitted the formation of the vermicular ridges. The ridges would have been formed by the upward pumping, instead of plastic flow, occurred on the higher, cooler, northwest side. This scenario, which is similar to that of Whitford-Stark (1982), allows the margins of the mountain to rest upon plateaus. A similar explanation might apply to the other occurrences of

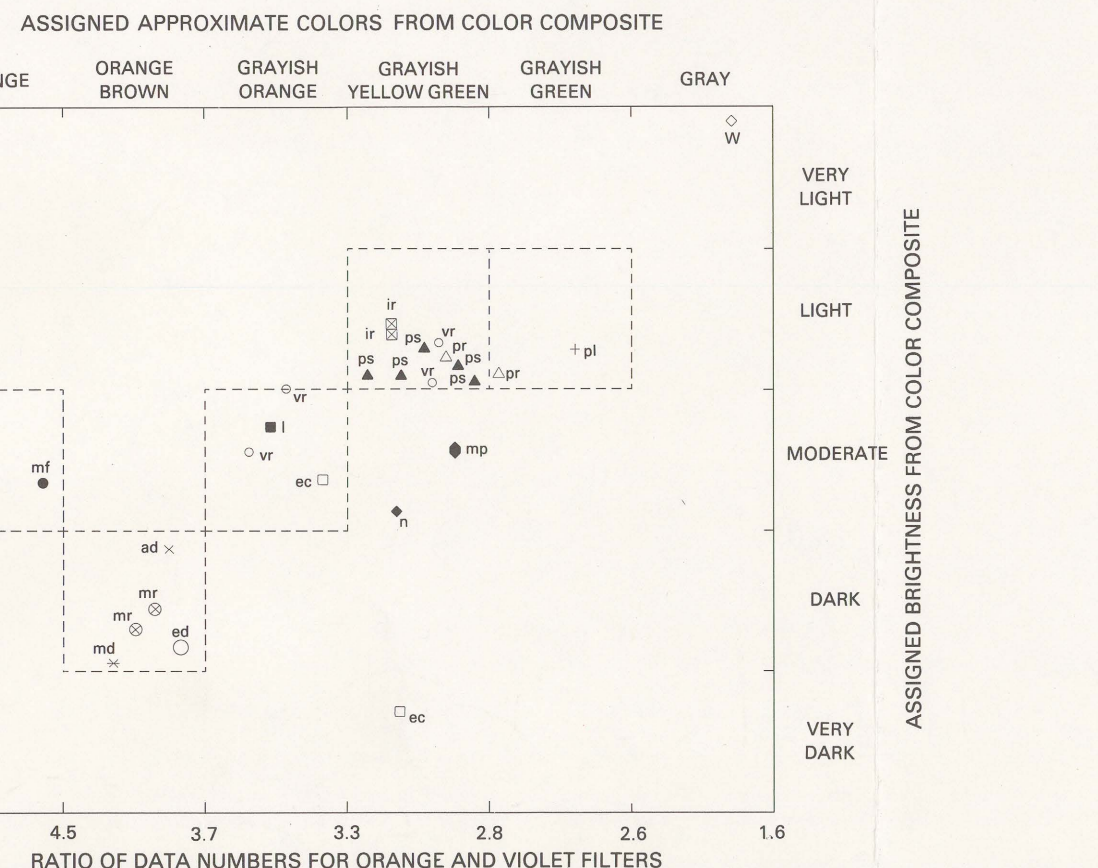


Figure 2—Assigned brightnesses and colors of geologic map units in Maasaw Patena map area. These assigned values are based on visual inspection of color composite (fig. 1A), work by Young (1984) that indicates an overall greenish cast for lo, and a rock-color chart (Godwin and others, 1948). Bin boundaries are based on the basis of color and brightness criteria. Plotted data points correspond to samples of map units indicated by symbols used on map. Samples of data numbers obtained by A.S. McEwen, U.S. Geological Survey, from composite of Voyager 1 frames 196, 198, and 200. (See fig. 1A and table 1).

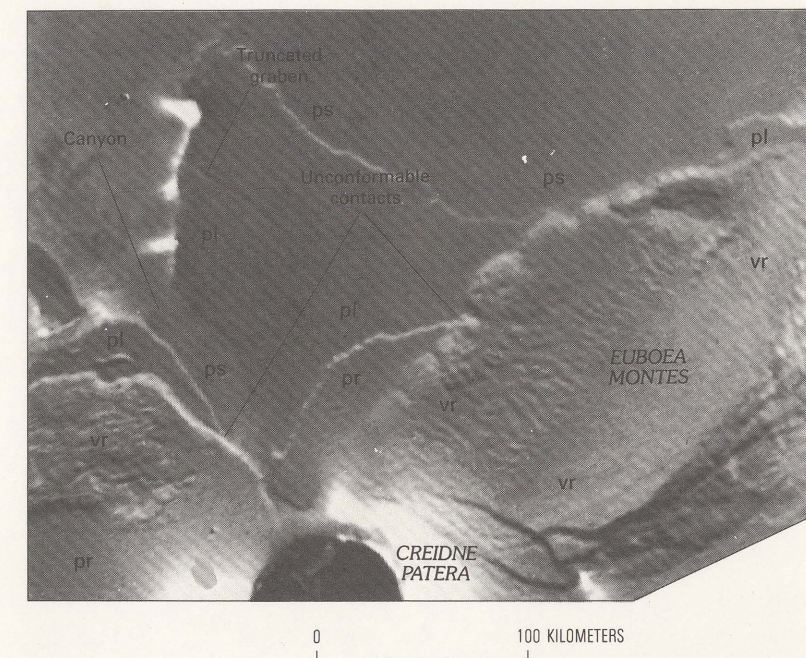


Figure 3.—Unconformable superposition of ridged plateau material (unit pr) on linedate plateau material (unit pl) and smooth plains material (unit ps) within and just south of Maasav Patara map area. Unconformable superposition of vermicular ridged mountain material (unit vr) on ridged (unit pr) and linedate (unit pl) plateau materials also shown (Voyager 1 frame 75).

GEOLOGIC MAP OF THE MAASAW PATERA AREA OF IO

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