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CRETACEOUS STRATIGRAPHY IN A NORTHEAST-TRENDING TRANSECT,  
NORTHERN UTAH TO SOUTH-CENTRAL SOUTH DAKOTA

By

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## INTRODUCTION

Sedimentary rocks of Cretaceous age at localities along Transect B-B' in Utah, Wyoming, and South Dakota (fig. 1) reflect intermittent deposition and erosion as well as regional and local subsidence and uplift. These former siliciclastic, calcareous, pyroclastic, and carbonaceous sediments accumulated in and adjacent to an epicontinental seaway that occasionally extended from northern Utah eastward to central Iowa and from the Gulf of Mexico northward to the Arctic Ocean. Along the transect, Cretaceous time (65.4-142 Ma; Obradovich, 1993) is represented by eight sequences of beds, which range in age from 65.4 Ma to about 120 Ma (Maastrichtian to Aptian), and by many unconformities.

The Cretaceous stratigraphy for Transect B-B' and three other east-trending transects in the Rocky Mountains and Great Plains was summarized by Dyman and others (1994) for the Western Interior Cretaceous (WIK) Project in the Global Sedimentary Geology Program of the International Union of Geological Sciences. Transect A-A' of the project trends irregularly from southwestern Montana to southwestern Minnesota. Transect C-C' extends from central Utah to northeastern Colorado. The southernmost transect, D-D', trends irregularly from east-central Arizona to westernmost Oklahoma. Dyman and others (1995) subsequently presented a more elaborate description of the rocks in Transect A-A'. The information originally compiled for Transect B-B' (Dyman and others, 1994) has been supplemented and refined for this publication.

Cretaceous rocks in this transect are depicted on lithostratigraphic and chronostratigraphic cross sections (figs. 2 and 3, respectively) as groups, formations, members, and lithologic units. The rocks of marine origin are classified as sandstone, mudrocks, or carbonate rocks, each of which includes lesser amounts of other rock types. Descriptions of these strata were obtained from studies of outcrops and borehole-records at the following localities, listed from southwest to northeast: Coalville, Utah; Church Butte, Wyoming; Dutch John, Utah; Red Desert, Shotgun Butte, Kaycee, and Red Bird, Wyoming; and White River, South Dakota. The distance from Coalville to White River, along a straight line, is about 580 mi (930 km). Near the southwestern end of the transect, in the thrust belt and the adjoining foredeep, beds of conglomerate, sandstone, shale, bentonite, and coal were deposited in continental and nearshore marine environments along the western margin of the epeiric sea. Laterally equivalent strata at the northeastern end of the transect on the craton consist of sandstone, shale, limestone, and bentonite that accumulated mainly in shallow-water marine environments. The sequences at the eight localities range in thickness from as much as 18,000 ft (5,500 m) at the west end of the transect in Utah, to about 3,000 ft (915 m) at the east end of the transect in South Dakota.

The geochronological framework for Cretaceous strata in the Rocky Mountains and Great Plains (figs. 3 and 4) is based on the identification and correlation of molluscan fossils (Cobban and others, 1994) and palynomorphs (Nichols, 1994) and the determination of radiometric ages (Obradovich, 1993) of associated beds. However, radiometric ages have been established for only 20 of the 73 marine-fossil zones listed in fig. 4; the absolute ages of 53 of the zones have been estimated by extrapolation. The sequences of beds at the localities in the transect (figs. 2 and 3) were correlated mainly by means of distinctive lithologic units, including conspicuous beds of bentonite, and marine molluscan fossils. Chronohorizons depicted on the cross sections indicate the established and estimated stratigraphic locations of marine fossils that represent five of the fossil zones listed on fig. 4. Fossils characteristic of these chronohorizons are widespread and are commonly associated with recognizable stratigraphic units. Fossils typical of other zones commonly are less abundant and are associated with less widely correlative beds. The datum for fig. 2 is fossil-zone 9, *Baculites reesidei* Elias of late Campanian age.

## STRATIGRAPHY

The Cretaceous sequences at the localities along Transect B-B' range from comparatively thick and dominantly nonmarine in northern Utah to relatively thin and dominantly marine in south-central South Dakota (figs. 1 and 2). Horizontal distances between the localities depicted on figures 1 and 2 during Cretaceous time have not been palinspastically restored. At Coalville, Utah, in the Sevier Orogenic Belt, the rocks are mostly siliciclastic and are as much as 18,000 ft (5,500 m) thick; they were deposited mainly in continental environments near the western shore of the epicontinental sea. About 82 mi (132 km) east-northeast of Coalville is Church Butte, situated on the Moxa Arch (figs. 1 and 2), where Cretaceous siliciclastic rocks are about 4,900 ft (1,500 m) thick. Most of the section preserved there accumulated in marine environments. At Red Desert, in the Cretaceous foreland, about 180 mi (290 km) east-northeast of Coalville, the strata are as much as 13,000 ft (4,000 m) thick and are composed of siliciclastic rocks and minor amounts of carbonate rocks that were deposited mostly in marine environments in the western part of the seaway. The Cretaceous sections at the localities north-northwest and northeast of Red Desert (figs. 1 and 2) thin progressively eastward and they consist of marine and continental rocks. At Shotgun Butte, Cretaceous strata are about

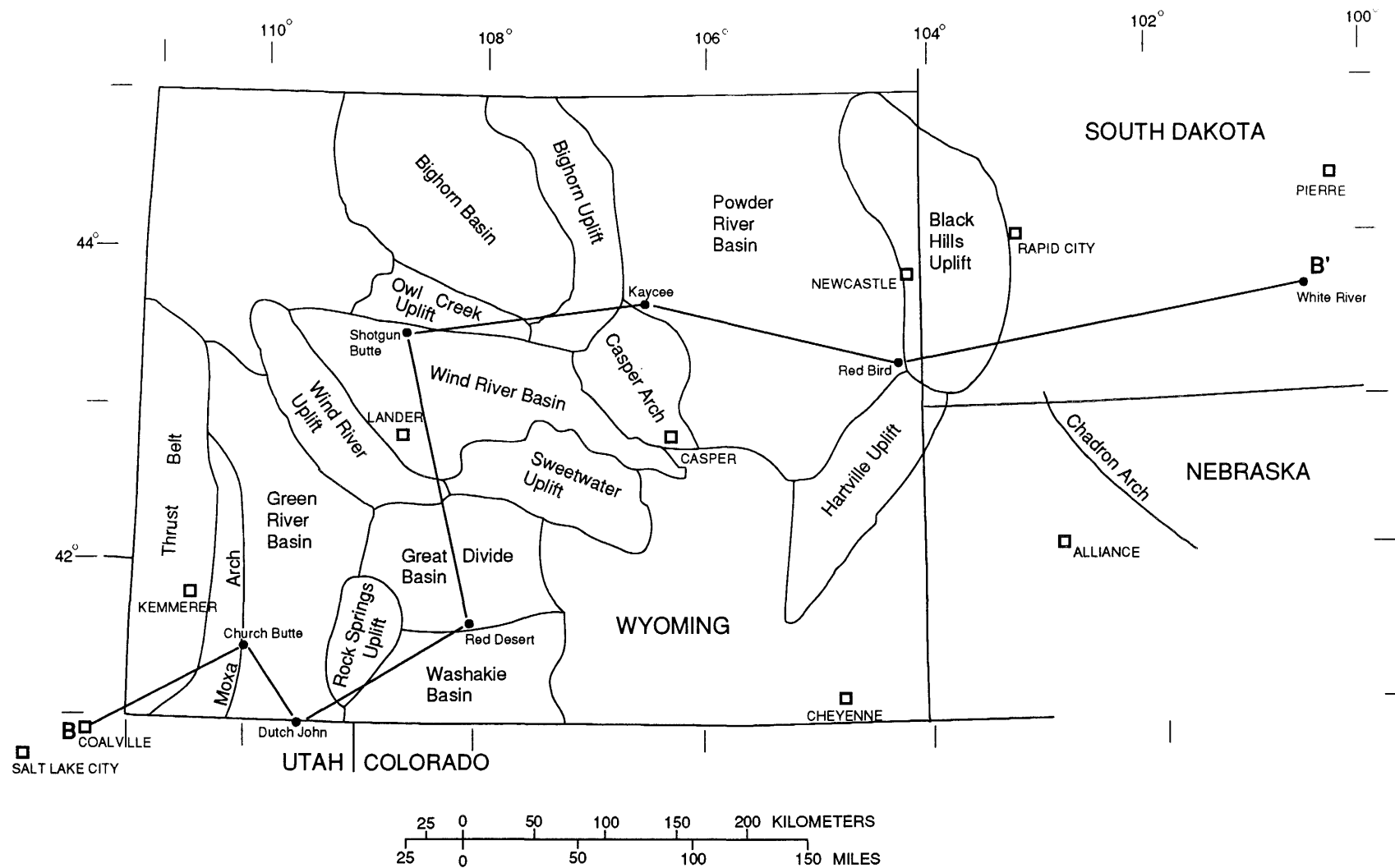


Figure 1. Map showing approximate location of cross section B-B', selected major structural features, and villages and towns in parts of Utah, Wyoming, South Dakota, and Nebraska.

STAGE	SUBSTAGE	FOSSIL ZONE NUMBER	FOSSIL ZONE
Maastrichtian (part)	upper	1	<i>Jeletzkytes nebrascensis</i>
		2	<i>Hoploscaphites nicolleti</i>
		3	<i>Hoploscaphites birkelundi</i>
	lower	4	<i>Baculites clinolobatus</i> *
		5	<i>Baculites grandis</i>
		6	<i>Baculites baculus</i>
		7	<i>Baculites eliasi</i>
Campanian	upper	8	<i>Baculites jenseni</i>
		9	<i>Baculites reesidei</i>
		10	<i>Baculites cuneatus</i>
		11	<i>Baculites compressus</i> *
		12	<i>Didymoceras cheyennense</i>
		13	<i>Exiteloceras jenneyi</i> *
		14	<i>Didymoceras stevensoni</i>
		15	<i>Didymoceras nebrascense</i> *
	middle	16	<i>Baculites scotti</i>
		17	<i>Baculites reduncus</i>
		18	<i>Baculites gregoryensis</i>
		19	<i>Baculites perplexus</i>
		20	<i>Baculites</i> sp. (smooth)
		21	<i>Baculites asperiformis</i>
		22	<i>Baculites maclearni</i>
		23	<i>Baculites obtusus</i> *
	lower	24	<i>Baculites</i> sp. (weak flank ribs)
		25	<i>Baculites</i> sp. (smooth)
		26	<i>Scaphites hippocrepsis III</i>
		27	<i>Scaphites hippocrepsis II</i> *
		28	<i>Scaphites hippocrepsis I</i>
		29	<i>Scaphites leei III</i>
Santonian	upper	30	<i>Desmoscaphites bassleri</i> *
		31	<i>Desmoscaphites erdmanni</i>
		32	<i>Clioscapites choteauensis</i>
	mid	33	<i>Clioscapites vermiformis</i>
	low	34	<i>Clioscapites saxitonianus</i> *
Coniacian	upper	35	<i>Scaphites depressus</i>
		36	<i>Scaphites ventricosus</i>
	mid	37	<i>Cremonoceras deformis</i> *
	low	38	<i>Forresteria peruana</i>

\*Fossil-zones dated radiometrically (Obradovich, 1993).

Figure 4. Ages and fossil-zones for marine strata of Cretaceous age in the Western Interior, U.S.A. (Cobban and others, 1994).

STAGE	SUBSTAGE	FOSSIL ZONE NUMBER	FOSSIL ZONE
Turonian	upper	39	<i>Prionocyclus germari</i>
		40	<i>Scaphites nigricollensis</i>
		41	<i>Scaphites whitfieldi</i>
		42	<i>Scaphites ferronensis</i>
		43	<i>Scaphites warreni</i>
	middle	44	<i>Prionocyclus macombi</i> *
		45	<i>Prionocyclus hyatti</i> *
		46	<i>Prionocyclus percarinatus</i>
		47	<i>Collignonicerus woollgari</i>
		48	<i>Mammites nodosoides</i>
Cenomanian	lower	49	<i>Vascoceras birchbyi</i> *
		50	<i>Pseudaspidoceras flexuosum</i> *
		51	<i>Watinoceras devonense</i>
	upper	52	<i>Nigericeras scotti</i>
		53	<i>Neocardioceras juddii</i> *
		54	<i>Burroceras clydense</i>
		55	<i>Euomphaloceras septemseriatum</i> *
		56	<i>Vascoceras diartianum</i>
		57	<i>Dunveganoceras conditum</i>
		58	<i>Dunveganoceras albertense</i>
		59	<i>Dunveganoceras problematicum</i>
		60	<i>Dunveganoceras pondi</i> *
		61	<i>Plesiacanthoceras wyomingense</i>
Albian (part)	middle	62	<i>Acanthoceras amphibolum</i> *
		63	<i>Acanthoceras bellense</i>
		64	<i>Acanthoceras muldoonense</i>
		65	<i>Acanthoceras granerosense</i>
		66	<i>Conlinoceras tarrantense</i> *
			(Gap in biostratigraphic record)
	lower	67	<i>Neogastrolites maclearni</i>
		68	<i>Neogastrolites americanus</i>
		69	<i>Neogastrolites muelleri</i>
		70	<i>Neogastrolites cornutus</i> *
		71	<i>Neogastrolites haasi</i> *
	upper	72	<i>Inoceramus bellvuensis</i>
		73	<i>Inoceramus comancheanus</i>

\*Fossil-zones dated radiometrically (Obradovich, 1993).

Figure 4. (continued)

10,000 ft (3,050 m) thick. The rock types and hiatuses in the section at Shotgun Butte differ from those in the sections at Red Desert and Kaycee (figs. 2 and 3), primarily as a function of distances between the three localities and the paleoshorelines and the thrust belt to the west. Near White River, South Dakota, on the craton and about 580 miles (930 km) east-northeast of Coalville, Cretaceous strata are about 3,000 ft (915 m) thick and are composed of siliciclastic and carbonate rocks that accumulated on a shelf mostly in nearshore and offshore marine environments.

At least four regional marine transgressions are recorded by strata in the transect (fig. 3). The first and oldest major transgression is represented by the Skull Creek Shale and a laterally equivalent part of the Thermopolis Shale (fig. 3); the second by the Greenhorn Formation and laterally equivalent beds in the Frontier Formation; the third includes the Niobrara Formation and strata of the same age in the Cody, Baxter, and Hilliard Shales; and the fourth and youngest by the Lewis Shale near Kaycee and Red Desert (fig. 3).

Many of the stratigraphic units in the transect are separated from underlying or overlying units by unconformities (figs. 2 and 3), which were formed by erosion during one or more of the following: lowering of sea level, regional or local tectonic uplift, and changes in the sediment supply or the climate. Major regional unconformities are recognized at the base, within, and near the top of the Lower Cretaceous sections at most of the localities. The hiatus at the base of the Cretaceous along the transect (fig. 3) might correspond in age to one or more of the medium type-2 sequence boundaries illustrated by Haq and others (1987). The unconformity at the base of the Fall River Sandstone and laterally equivalent beds in the Cloverly Formation might be the same age as a major type-1 sequence boundary shown by Haq and others (1987). Van Wagoner and others (1990) concluded that the widespread unconformity at the base of the Muddy Sandstone Member of the Thermopolis Shale (figs. 2 and 3) is a type-1 sequence boundary. Upper Cretaceous strata of Turonian age generally contain one or two unconformities, which apparently are also type-1 sequence boundaries. In the eastern part of the transect near Red Bird and White River, an unconformity separates Turonian shale from Coniacian calcareous rocks (perhaps a minor sequence boundary), and a younger unconformity separates marine shales of early and middle Campanian ages. The unconformity between the Gammon Member and the overlying Sharon Springs Member of the Pierre Shale at Red Bird and White River is a type-1 sequence boundary (Van Wagoner and others, 1990). In the western part of the transect, an unconformity of middle Campanian age was found near Dutch John and an upper Campanian unconformity extends from Coalville eastward at least to Red Bird. The latter two unconformities could be type-1 sequence boundaries according to Haq and others (1987) and Van Wagoner and others (1990), respectively.

The Cretaceous System (65.4-142 Ma; Obradovich, 1993) is represented along the transect by sedimentary rocks and associated unconformities (figs. 2 and 3), although less conspicuous unconformities are not always depicted. In the region of Coalville, Church Butte, Dutch John, and Shotgun Butte, where Cretaceous deformation was substantial, less than 40 percent of Cretaceous time is represented by the strata; the remainder of the Cretaceous is represented by the hiatuses (fig. 3). The lithologic record of the System is least complete at Church Butte, on the Moxa Arch, where rocks probably represent only about 20 percent of Cretaceous time. This anomalous section indicates uplift and erosion in the area of the arch, perhaps beginning in early Turonian time and probably extending through most of late Campanian time. At Red Desert, Kaycee, Red Bird, and White River, about half of the System is recorded by strata. Upper parts of the Cretaceous sections at Kaycee and Red Bird, in contrast to the other localities, were not truncated by Tertiary erosion. Nevertheless, the most complete lithologic record of the Cretaceous along the transect is interpreted to be located on the craton near White River, where rocks represent slightly more than 50 percent of the period. The contrasting stratigraphic records of deposition at the eight localities apparently reflect differences in the amounts and effects of local tectonism as well as the effects of eustasy.

Chronohorizons defined by biozones within the transect (figs. 2 and 3) delimit four chronostratigraphic units. The oldest chronohorizon (biozone 62) also serves as an upper boundary for strata of Aptian, Albian, and early and middle Cenomanian ages. The youngest chronohorizon (biozone 9) is the lower boundary for rocks of latest Campanian and Maastrichtian ages. Stratigraphic units of Aptian age through most of the middle Cenomanian, middle Cenomanian to middle Coniacian age (biozones 62 to 37), and late Coniacian and Santonian age (biozones 37 to 30) are thickest near Coalville, at the western end of the transect. However, the thickness of the rocks of late Coniacian and Santonian ages at Coalville is only slightly more than their thickness at Dutch John, which is about 93 mi (150 km) east of Coalville and 43 mi (69 km) southeast of Church Butte on the Moxa Arch. The unit, mainly of early Campanian age (biozones 30 to 23), and the overlying unit of middle and late Campanian ages (biozones 23 to 9) are thickest at Red Desert, which is about 180 mi (290 km) east-northeast of Coalville. Strata of latest Campanian and Maastrichtian ages have been truncated at six of the eight localities, which precludes identification of the thickest part of the unit. Nevertheless, the geographic location of the maximum thickness of each chronostratigraphic unit suggests that the main depocenter along the transect moved progressively eastward during Cretaceous time.

This movement might be related to the eastward movement of thrust plates in the thrust belt during the Cretaceous.

#### ACKNOWLEDGMENTS

The report by Dyman and others (1994) for the WIK was derived from information contributed by volunteers, including earth-scientists employed by governmental agencies and commercial organizations. Most of the contributors for Transect B-B' obtained data from unpublished records as well as from publications indicated in the attached list of references. Authors of the report for this transect supplied information for seven of the eight constituent localities. Strata at outcrops and in boreholes near Coalville were described by T.A. Ryer, The ARIES Group, Inc. Information concerning rocks in the subsurface near Church Butte was provided by R.E. Mueller, North American Resources Co. D.P. Stilwell, U.S. Bureau of Land Management, and A.C. Smith, Computational Geology, Inc., supplied data for beds at outcrops and in the subsurface near Dutch John. Information concerning strata in the subsurface near Red Desert was assembled by J.C. Dolson and W.B. Hanson, Amoco Production Co., and B.E. Law, U.S. Geological Survey. Outcrops at Shotgun Butte were described by W.R. Keefer (Keefer and Troyer, 1957), U.S. Geological Survey. Descriptions of outcropping beds near Kaycee were compiled by E.A. Merewether, U.S. Geological Survey. Information from outcrops in the vicinity of Red Bird was contributed by D.M. Wheeler, Ensign Oil & Gas, Inc. At the eastern end of the transect near White River, strata at outcrops were studied by J.R. Gill and W.A. Cobban (1966), and rocks in the subsurface were interpreted and depicted by D.D. Rice (1977).

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Molluscan fossils from outcrops of marine beds at the preceding localities (fig. 1) and many other localities in the Western Interior have been identified and assigned ages by W.A. Cobban (Cobban and Reeside, 1952; Cobban and others, 1994). Palynomorphs in nonmarine and nearshore-marine rocks at outcrops and in boreholes in western Wyoming and adjacent areas have been described and related to the molluscan fossils by Nichols and Jacobson (1982a), Nichols and others (1982b), and Nichols (1994). Absolute ages for Cretaceous strata in the Rocky Mountains and Great Plains were determined by J.D. Obradovich (1993, 1996) by means of radiometric analyses of sanidine-bearing bentonite from outcropping beds and drill-cores that are associated with fossiliferous beds. The geochronological information developed and provided by the earth-scientists named above was essential in the preparation of figures 2, 3, and 4.

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