

FOLIO OF THE SEWARD AND BLYING SOUND QUADRANGLES, ALASKA

Local drainage basin

A source of gold within the same basin as occupied by a stream placer is demonstrable for some placer gold deposits. Placer gold from Bear Creek is associated with native silver that was more abundant than in any other placer deposit of the Turnagain Arm drainage system (Martin and others, 1915). The fineness of the gold is less (0.740 to 0.761, table 1) than that from any other placer south of Turnagain Arm (Becker, 1896, p. 81; Martin and others, 1915, p. 186), reflecting the greater abundance of silver. The placer gold from the Turnagain Arm drainage system of Tysdal, 1978) near the head of Bear Creek was staked originally as a silver lode (Tuck, 1933, p. 507), and "several" prospects reported valuable for their silver content" were located along Bear Creek (Martin and others, 1915, p. 179).

The placer gold of Crow Creek has the same physical characteristics as the gold of lode mines at the head of Crow Creek (Park, 1933, p. 404). Fineness of the placer gold commonly ranges between 0.720 and 0.730 (table 1), which corresponds with the 25 percent by weight of silver reported (Park, 1933, p. 408) for free gold of the lodes. Native silver was observed in the lodes (Martin and others, 1915, p. 134) and also in the placer deposits (Moffitt, 1906, p. 42). Quartz-diorite of the placers is like that which crops out near the head of Crow Creek, and pay gravel is associated everywhere with the quartz diorite (Park, 1933, p. 404). No source for quartz diorite is known in neighboring areas.

The placers of Lynx Creek yielded both gold and native copper. The only known vein of copper sulfide in the Tunnagain Arm drainage system is in the headwaters of Lynx Creek (Tuck, 1933; Tysdal, 1978) described by Paige and Knopf (1907, p. 124-125). No native copper was reported at this copper prospect, but native copper probably formed as an alteration product of the copper lode (Moffit, 1905). Pyrrhotite-impregnated diabase boulders occur in the stream gravels (Martin and others, 1915).

In the early part of the century, it was suggested by some miners who worked streams in the Turnagain Arm drainage system that glaciers transported gold-bearing rock into the stream drainages and the gold later was freed and reworked into placers (Morfit, 1906, p. 45). The idea probably came about because granite clasts in stream gravels on the south side of the arm have no source in the country rock of the same drainages. The argument is largely discounted in the preceding section, which shows local basin sources for some gold and a spatial association of lode mines and placer deposits. Nevertheless, some small amount of gold could have been glacially transported, and its possible sources are discussed.

Granitic clasts are in the gravels of Resurrection, Bear, Palmer, Sixmile, Gulch, and Canyon Creeks and East Fork. They increase in abundance northward from East Fork to Turnagain Arm, and the clasts of Resurrection Creek are larger than those to the east (Moffit, 1906, p. 24). Granitic clasts occur in gravels of Crow, California, and Winner Creeks north of Turnagain Arm, derived from intrusions at the head of Crow Creek. The clasts south of the arm probably came from the same source (Moffit, 1906), moved on glaciers across the arm, and were reworked by streams.

Glaciers flowed south from the Matanuska Valley (north of the quadrangle) along the western front of the Chugach Mountains and could have supplied gold-bearing rock. Eklutna moraines overlie bedrock on the western slopes of the mountains in the northwestern part of the Seward quadrangle (Karlstrom, 1964), and my observations show that they contain granite of types found in the Talkeetna Mountains and conglomerate of the Chickaloon Formation of the Matanuska Valley. The gravities in the streams on the south side of Turnagain Arm could be reworked from glacial debris transported into the glacial by ice that moved eastward up the arm. Granite clasts of the Eklutna moraines are pink, unlike the gray granites in the streams. Some of the clasts of the Turnagain gravels are described as "granite" in plate and sandstone of local origin (Moffitt, 1906), but a source in the Matanuska Valley or Talkeetna Mountains would have contributed rock types foreign to the Chugach terrane.

There is no immediate source for the granite south of Turnagain Arm. The nearest outcrop is near the Harding Icefield, immediately to the west of the Seward quadrangle, about 90 km south-southwest of Hope. Glaciers flow north and west from the icefield during Naptowne time but, according to maps of Karlstrom (1964), did not expand into the drainage system of Turnagain Arm. In addition, the number of granitic clasts in the gravels decreases southward, opposite to what should be observed. No outcrop source exists to the east within the reaches of westward-flowing glaciers, and the increased number and size of the clasts are opposite to what would be expected.

Kenai River drainage system

In the eastern part of the Kenai River drainage system, a minor amount of placer gold was recovered from the flood plain of the lower part of Falls Creek, the benches along the lower segment of the stream, and at the mouth of the canyon (Martin and others, 1915). Several mines that produced abundant gold occur in the eastern part of the drainage system, but the streams are short and generally contain only a small amount of gravel that has undergone little reworking. Glaciers retreated from the area not long ago, and unworked morainic material is present in some drainages (Johnson, 1912, p. 168).

In the northwestern part of the drainage system, some gold was produced from Quartz Creek, which flows across a bedrock floor. Lode mines that produced gold occur near Slate, Summit, and Colorado Creeks, which drain into Quartz Creek and probably supplied the gold. Bench gravels are thin, only 4 to 6 m thick, and the area had an extended period of time for reworking of the gravels.

Stream west of Kenai Lake yielded much of the gold recovered from the Kenai River drainage system, most of it from the stream bed. The stream bed is composed of gravel and sand, and is the source of much of the placer gold. Prospectors or mines can be found in the headwaters of Stetson and Cooper Creeks, but pan concentrates of stream sediments from several creeks in the general area gave gold values of 70 to more than 500 parts per million (Tripp and others, 1978), values commonly found in known gold field districts of the Seward quadrangle.

Gravels along the lower 1 km of Cooper Creek and downstream along the Kenai River are deltaic deposits, as much as 2 m thick, composed of glacial till (Martin and others, 1915). These sediments were not reworked, except for the volume of them removed during the 1964 earthquake. The stream bed is composed of gravel and sand, and is the source of much of the placer gold. Prospectors or mines can be found in the headwaters of Stetson and Cooper Creeks, but pan concentrates of stream sediments from several creeks in the general area gave gold values of 70 to more than 500 parts per million (Tripp and others, 1978), values commonly found in known gold field districts of the Seward quadrangle.

Resurrection River drainage system

The only record of placer gold production from the Resurrection River drainage system is a statement by Smith (1926, p. 12) that less than \$5,000 worth of gold (gold valued at \$20.67 per troy ounce) was recovered in 1924. Two lode gold

Aspects were located on the south side of the Resurrection River in 1900 (U.S. Bureau of Mines, 1973a, and KARDEX, 1976a), but no record of production is known. Gold values of 300 to more than 500 parts per million were detected in an concentrates of stream sediments from the western part of the drainage system (Tripp and others, 1978), and in the few quadrangles that have been sampled recently to determine the distribution of known and possible gold. The drainage system is of little or no historical interest; the placer deposits because the stream bed is young, from about the time the river was exposed only a relatively short time ago; the Resurrection River valley is clogged with glacialifluvial gravels; and the stream has not reworked gravels over an extended period of time. It does not flow across bedrock, and tributary

trans are generally do not contain such sedimentary debris.

Port Wells—Passage Canal drainage system

No significant amount of placer gold was produced from active stream or bench gravels in the Port Wells—Passage Canal drainage system even though gold-bearing quartz veins, as well as several local gold mines, exist throughout the area. Gold-bearing stream gravels were discovered on Billings Creek north of Whittier in 1896 (Johnson, 1914), and geochemical and spectrographic analyses of pan concentrates of stream sediments reveal gold in gravels throughout the entire drainage basin. However, because of the lack of record of placer production or significant activity in the drainage basin, and only a few claims were recorded, Ice still covers part of the area, and the remainder of the drainage system was examined only recently from beneath ice. The streams are short, many of them contain little gravel and are fed by deep floors, and the massive overlying ice prevents any significant placer activity.

BEACH PLACER DEPOSITS

No productive beach placer deposits are known in the map area. The beach placer at Sniper's Point (A) near the mouth of Sixmile Creek has not been worked (Johnson, 1912, p. 161; Martin and others, 1915, p. 187). The placer deposits of Harris Bay (B) and Johnstone Bay (C) were listed as active claims in 1975 (U.S. Bureau of Mines, 1973b, and KARDEX, 1976b), but nothing is known about production.

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