

MAP A. BERYLLIUM IN THE NONMAGNETIC FRACTION OF HEAVY-MINERAL CONCENTRATES

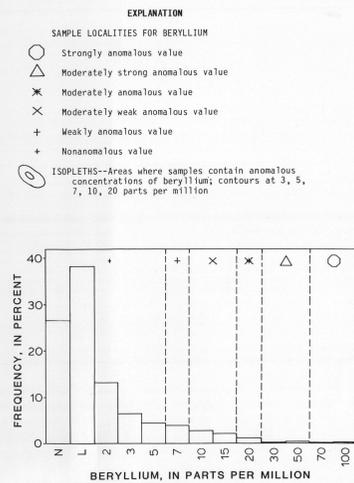
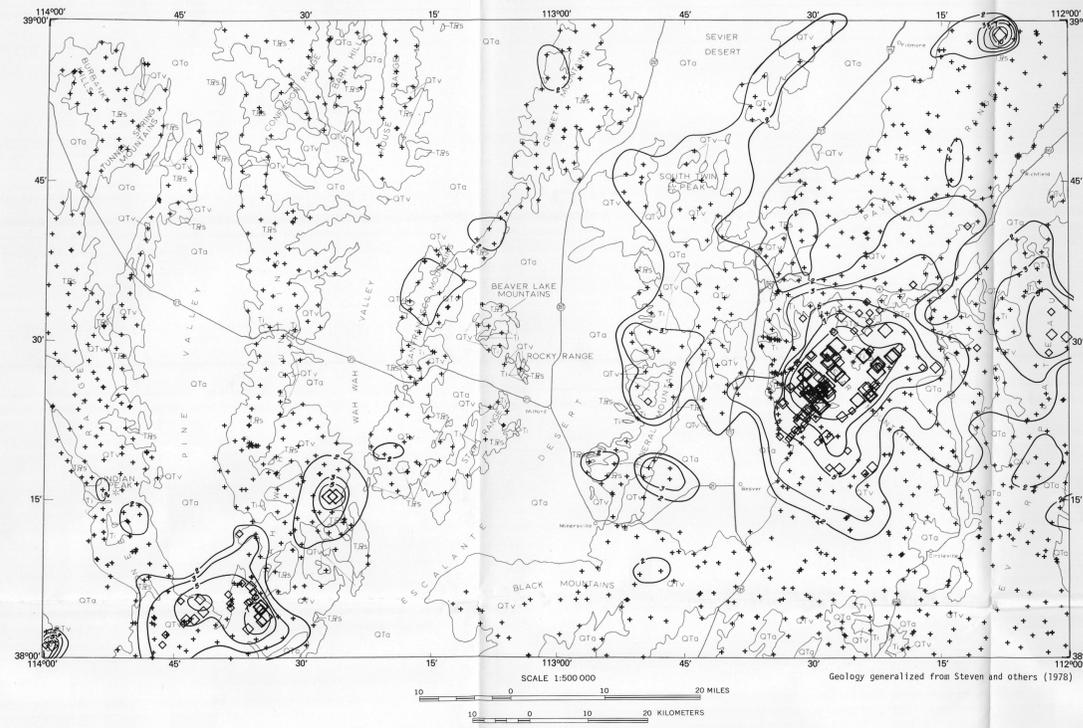


Figure 1.—Histogram showing concentrations of beryllium in the nonmagnetic fraction of heavy-mineral concentrates from the Richfield 1° x 2° quadrangle, Utah. Number of samples, 1,456; N, not detected at 2 parts per million (ppm); L, detected but less than 2 ppm.

- LIST OF MAP UNITS**
- QTa SURFICIAL DEPOSITS, UNDIVIDED (QUATERNARY AND TERTIARY)
  - QTV VOLCANIC ROCKS, UNDIVIDED (QUATERNARY AND TERTIARY)
  - T1 IGNEOUS ROCKS, UNDIVIDED (TERTIARY)
  - TP2s SEDIMENTARY ROCKS, UNDIVIDED (TERTIARY TO PALEOZOIC)
  - CONTACT



MAP C. BERYLLIUM IN LESS-THAN-0.180mm-FRACTION OF STREAM SEDIMENTS

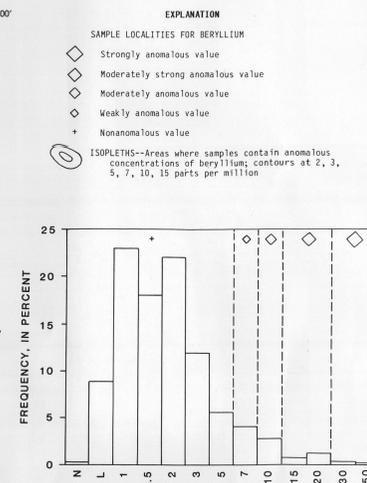
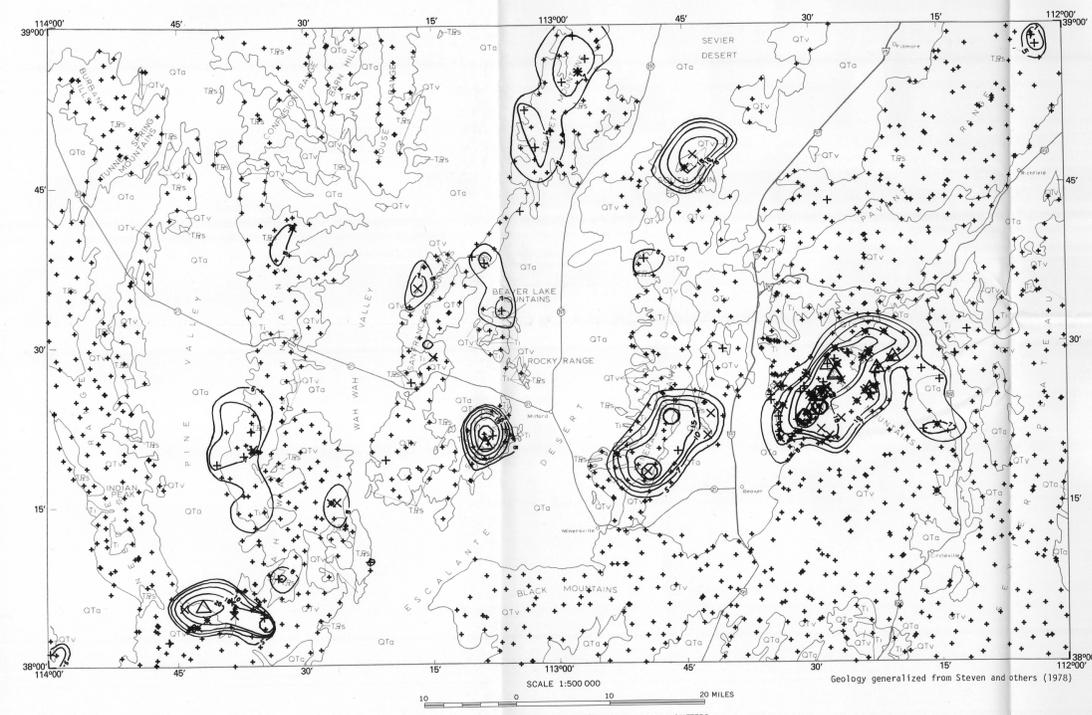


Figure 3.—Histogram showing concentrations of beryllium in the less-than-0.180-mm fraction of stream sediments from the Richfield 1° x 2° quadrangle, Utah. Number of samples, 1,445; N, not detected at 1 part per million (ppm); L, detected but less than 1 ppm.



MAP B. BERYLLIUM IN THE MAGNETIC FRACTION OF HEAVY-MINERAL CONCENTRATES

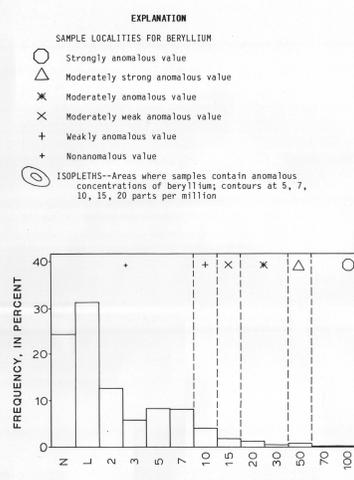


Figure 2.—Histogram showing concentrations of beryllium in the magnetic fraction of heavy-mineral concentrates from the Richfield 1° x 2° quadrangle, Utah. Number of samples, 1,570; N, not detected at 2 parts per million (ppm); L, detected but less than 2 ppm.

**INTRODUCTION**

This report is part of a series of maps of the Richfield 1° x 2° quadrangle, Utah, prepared under the Continuous United States Mineral Assessment Program. Other publications in this folio are listed in the selected references.

Located in west-central Utah, the Richfield quadrangle covers the eastern part of the Pliocene-Marysvale igneous and mineral belt, which extends from the vicinity of Pliocene in southeastern Nevada east-northeastward for 250 mi (155 mi) into central Utah. The western two-thirds of the Richfield quadrangle is in the Basin and Range province and the eastern third is in the High Plateaus of Utah, a subprovince of the Colorado Plateau.

Bedrock in the northern part of the Richfield quadrangle consists predominantly of latest Precambrian and Paleozoic sedimentary strata that were thrust eastward during the Sevier orogeny in Cretaceous time onto an autochthon of Mesozoic sedimentary rocks in the eastern part of the quadrangle. The southern part of the quadrangle is largely underlain by Oligocene and younger volcanic rocks and related intrusions. Extensional tectonism in late Cenozoic time broke the bedrock terrane into a series of north-trending fault blocks; the uplifted mountain areas were deeply eroded and the resulting debris deposited in the adjacent basins. Most of the mineral deposits in the Pliocene-Marysvale mineral belt were formed during igneous activity in middle and late Cenozoic time.

The regional sampling program was designed to define broad geochemical patterns and trends which can be utilized along with geologic and geophysical data to assess the mineral resource potential of this quadrangle. These maps of the Richfield 1° x 2° quadrangle show the regional distributions of beryllium in two fractions of heavy-mineral concentrates and the less than 0.180 mm fraction of drainage sediments.

**COLLECTION OF SAMPLES**

Drainage sediment samples were collected throughout the Richfield quadrangle during the summer of 1978. The sample sites were located along small, normally unbranched or first-order stream drainages, which ranged from 1.7 to 3.3 mi (1-2 mi) in length and from 1.8 to 3.7 m (6-12 ft) in width. Sample density was 1 sample per 9 mi<sup>2</sup> (3 mi<sup>2</sup>) within the bedrock areas. Intermountain basins containing sediments were not sampled. Each sample is a composite of material collected at four or five places across and along the active channel. Generally about 4 kilograms (8.8 pounds) of bulk sediment were collected for reduction to a paned-concentrate and a less than 0.180 mm fraction sample. The geochemical sampling was carried out by W. R. Miller, J. B. McHugh, G. K. Lee, J. F. Guadagnoli, L. DiGuardia, J. D. Tucker, and R. E. Tucker.

**PREPARATION OF SAMPLES**

Samples of drainage sediments were first panned to eliminate most clay minerals and the common rock-forming minerals, such as quartz, feldspar, and calcite. Most of the drainages were dry so the samples were panned at a field laboratory. About half of the panned heavy-mineral concentrates were prepared and analyzed at a field laboratory set up at Milford, Utah, and the remaining at the Geological Survey laboratory in Golden, Colorado. The preparation and analyses were done by J. M. Motooka, J. B. McHugh, J. D. Tucker, R. E. Tucker, and J. F. Guadagnoli.

The panned concentrates from each sample were dried and sieved to minus-10-mesh (<1.0 mm), and the magnetic fraction was separated with a hand magnet. The remaining concentrate was separated using bromoform specific gravity of 2.86 into a light and a heavy fraction. The light fraction, which contained mainly minerals such as quartz, feldspar, and calcite was discarded. The remaining heavy-mineral fraction was separated electromagnetically by a Franz 1 isodynamic separator with forward and side angle settings of 15 degrees and an amperage setting of 0.2. The magnetic fraction at 0.2 amperes was discarded and the remaining fraction was further separated electromagnetically into a nonmagnetic and magnetic fraction at a setting of 0.6 amperes. These two fractions were hand ground to less than 149 micrometers (microns) and then analyzed.

**SELECTION OF REFERENCES**

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1985, Maps showing distribution of arsenic in heavy-mineral concentrates and less than 0.180 mm fraction of drainage sediments, Richfield 1° x 2° quadrangle, Utah: U.S. Geological Survey Miscellaneous Field Studies Map MF-1246-B, scale 1:500,000.

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**GENERAL IMPLICATIONS OF THE MAGNETIC, NONMAGNETIC, AND LESS THAN 0.180 MM FRACTIONS**

The nonmagnetic, magnetic and less than 0.180 mm fractions of drainage sediments consist of different mineral suites, whose geochemical implications with regard to potential mineral resources differ significantly. The nonmagnetic fraction contains accessory minerals, such as zircon, monazite, and primary and secondary ore minerals. Anomalous beryllium associated with the nonmagnetic fraction of heavy-mineral concentrates generally indicates surface or near-surface sources and occurs in primary minerals such as chrysoberyl, helvite, phenacite, and monazite. The magnetic fraction contains mafic-rock minerals (such as biotite, amphibole, pyroxene) and more importantly, both detrital and detrital iron and manganese oxides containing anomalous trace metals. Iron and manganese oxides commonly fill or coat fractures, are abundant along or near mineralized faults, and extend significant distances from related ore deposits. Anomalous trace-metal content of the magnetic fraction could, therefore, indicate possible buried deposits. The less than 0.180 mm fraction contains rock-forming minerals such as quartz and feldspar but also beryllium-bearing minerals which are too light to be present in the heavy-mineral concentrates such as beryl and bertrandite. The use of the three fractions aid in the interpretation of geochemical data and provides clues as to the geological environment, and the source of anomalous metals.

Reconnaissance geochemical surveys are valuable tools in mineral exploration, but they should be used in conjunction with data from other earth science disciplines. In particular, outlining exploration targets generally involves considerable additional, more detailed investigations.

MAPS SHOWING DISTRIBUTION OF BERYLLIUM IN HEAVY-MINERAL CONCENTRATES AND STREAM SEDIMENTS, RICHFIELD 1° X 2° QUADRANGLE, UTAH

By  
William R. Miller, Jerry M. Motooka,  
and John B. McHugh  
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