

Base from U.S. Geological Survey, 1:62,500, Agness, Bone Mountain, Marial, Powers, 1954.

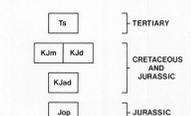
18ft  
MAGNETIC NORTH  
TRUE NORTH  
APPROXIMATE MEAN DECLINATION, 1982

SCALE 1:48 000  
0 1 2 3 MILES  
0 1 2 3 KILOMETERS  
CONTOUR INTERVAL 20 FEET  
DATUM IS MEAN SEA LEVEL



Geology generalized from Gray and others (1982).  
Aeromagnetic survey flown and compiled by Aero Services, 1978.

CORRELATION OF MAP UNITS



DESCRIPTION OF MAP UNITS

- T1 SEDIMENTARY ROCKS (TERTIARY)—As mapped, includes Tye Formation, Flourery Formation of Baldwin (1974), and Lounglass Formation of Baldwin (1974). Formations are composed of sandstone, conglomerate, siltstone, and mudstone.
- K1m MYRTLE GROUP (CRETACEOUS AND JURASSIC)—Includes Boys Creek Formation and Reddie Formation composed of siltstone, sandstone, and conglomerate.
- K1d DOBAN FORMATION (CRETACEOUS AND JURASSIC)—Includes a sedimentary unit of gneiss and metagneiss with local exposures of chert and a volcanic unit of andesitic to porphyritic basalt.
- K1e ANDESITIC TO DIABASIC ROCKS (CRETACEOUS AND JURASSIC)—Andesitic flows, agglomerates, tuff, hornblende dacite, quartz diorite, and minor hypabyssal rocks comprising an islandic sequence.
- J1p OPHIOLITIC ROCKS (JURASSIC)—Basaltic gneiss, sheeted diabase like, gabbro, serpentinite.

INTERPRETATION

The most striking feature of the aeromagnetic map is a linear north-south gradient (A) that crosses the volcanic andesites (B) associated with the Doban Formation to the east from high-amplitude, short-wavelength anomalies to the west. The gradient continues 3 mi (1.1 km) south of the map, closely following this contact, where it is coincident with the edge of a thin sheet of ultramafic rock thrust from the west over the Doban Formation. Modeling experiments suggest that the ultramafic sheet is the source of the short-wavelength, high-amplitude anomalies at this southern extremity (Blakely and others, 1982). Near the southern edge of the Wild Rogue Wilderness aeromagnetic map, short-wavelength, high-amplitude anomalies overlie a widespread andesitic unit that is unconformably overlain by Tertiary and Jurassic rocks (Papp and others, 1981). It seems likely, therefore, that many of the anomalies west of gradient A are caused by deep-seated mafic or ultramafic rocks buried at shallow depth below the sedimentary rocks. We cannot say from the magnetic data whether these rocks occur as a continuous sheet with variable magnetic properties or as a discontinuous sequence of tectonic wedges.

At about 124° 05' W., the north-south gradient broadens and abruptly shifts to a northeasterly trend (C). Most of the Wild Rogue Wilderness is covered by this broad gradient and associated positive anomaly (anomaly D). Geologic mapping (Gray and others, 1982) shows that this gradient and anomaly overlie a tectonic wedge of volcanic and intrusive rocks unconformably overlain by Jurassic and Tertiary sedimentary rocks on the west and by the Doban Formation on the east. In volcanic and intrusive rocks are usually more magnetic than sedimentary rocks, it is reasonable to assume that they are the source of anomaly D. However, the gradients that form anomaly D are too broad to be caused by rocks exposed at the surface. Moreover, the topographic features within the wedge should produce numerous short-wavelength magnetic anomalies if the surface rocks are the main magnetic source.

Figure 2 shows a simple modeling experiment to demonstrate that gradient C and anomaly D must be caused by subsurface rocks. Profile X-Y, directly across the wilderness, was extracted from the aeromagnetic data and a magnetic model was constructed assuming the volcanic rocks are uniformly magnetized and that the surrounding sedimentary rocks are nonmagnetic. Consequently, the upper surface of the model conforms to the topography directly below the profile and the east and west ends correspond to the volcanic and the sedimentary rocks, respectively. The model is shown in figure 2. The model incorporates two simplifying assumptions: the topography and anomaly are linear over the distance of the profile and the magnetic susceptibility of the volcanic rocks is negligible compared to the sedimentary rocks. The model shows that the magnetic profile produced by this model is displaced more than 1 mi (1.6 km) west and has a much broader shape than the calculated anomaly. The model also shows that the observed and calculated profiles would be even greater if the profile had been taken over one of the larger topographic features in the vicinity, such as Bone Mountain about 1 mi (1.6 km) north of profile X-Y. Hence, the magnetic wedge of subsurface volcanic and intrusive rocks of the Wild Rogue Wilderness is probably the source of anomaly D and gradient C.

To investigate the general depth and size of the source of gradient C and anomaly D, a model was constructed by trial and error that produces a profile in close agreement with the observed profile. The result is shown in figure 3. The calculated body is a relatively thin sheet that crops abruptly near the eastern edge of the volcanic wedge. Because of our simplifying assumptions and the non-uniqueness of magnetic modeling, the exact shape of the model is not of particular significance. We conclude only that the upper surface of the body is at shallow depth (1 to 2 km) below the surface. The model also shows that the magnetic profile produced by this model is displaced much less than 1 mi (1.6 km) west and has a much broader shape than the calculated anomaly. All of these conclusions are consistent with the geologic interpretations of Gray and others (1982).

Anomaly E is over the dacite of Saddle Peak and is the only substantial anomaly within the wilderness that is caused by rocks cropping out at the surface. Although this region of the dacite is sufficiently linear to produce a substantial anomaly, elsewhere the dacite appears only weakly magnetic. Mount Silver, for example, is a larger topographic feature than Saddle Peak, and is composed of the same dacitic lithology, but produces only a subdued magnetic anomaly. Dacite and andesite samples from the Wild Rogue Wilderness have high magnetic contents (Gray and others, 1982) that may explain anomaly E. Gray and others also note the presence of massive sulfide deposits in the Mount Silver region, and any associated hydrothermal alteration may have destroyed magnetite over a wide area. Apparently, anomaly E reflects a relatively unaltered part of the dacite or an near the topographic surface.

Anomaly F is located outside the wilderness but may have an impact on the mineral potential of the area. Its circular shape and broad gradients are typical of anomalies caused by magnetic plutons, in this case located several miles below the elevation of the aeromagnetic survey. The emplacement of a buried pluton often has profound effects on the mineralization of surrounding rocks, but there is no evidence of mineralization of the Jurassic sedimentary rocks in this region. We prefer to interpret anomaly F as being caused by an especially magnetic wedge of the same mafic or ultramafic unit that is postulated to cause gradients A and C and anomaly D. The latter conclusion is supported by the fact that anomaly F and D and gradient C appear connected.

Enough for this map region, the volcanic and intrusive rocks exposed at the surface are not the source of the major anomalies of the wilderness, but modeling experiments indicate that the source may be less than 1 km (0.6 mi) below the topographic surface. Considering the exceptionally linear extension of gradient C and its eastern outcrop at the southern edge of the map and to ultramafic outcrops 3 mi (1.1 km) south of the map, we conclude that gradients A and C, anomaly D, and many of the anomalies west of gradients A and C are caused by variations in the magnetization of a thin, widespread tectonic slice of mafic or ultramafic rocks. If this sheet has mineral potential, such as chromite, where it crops out to the south, then it is possible that similar potentials exist at relatively shallow depth in the Wild Rogue Wilderness.

REFERENCES

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- Gray, Floyd, Papp, Lisa, Morley, Barry, Douglas, L., and Donohoe, J., 1982, Geologic map of the Wild Rogue Wilderness, Coos, Curry, and Douglas counties, Oregon. U. S. Geological Survey Miscellaneous Field Studies Map MF-1281-A, scale 1:62,500.
- Papp, M. J., Gray, Floyd, Cannon, J. K., Foose, M. P., Lips, Bruce, Morley, Barry, Richmond, S. W., Swille, M. G., Till, Alison, and Ikonoski, W. P., 1981, Geologic map of the Kalmiopsis Wilderness, Oregon. U. S. Geological Survey Miscellaneous Field Studies Map MF-1240-A, scale 1:62,500.
- U.S. Geological Survey, 1979, Aeromagnetic map of the Redford area, Oregon. U. S. Geological Survey Open File Report 79-119, scale 1:250,000.

Figure 2. Magnetic model of Bone Mountain and vicinity assuming a buried source. See fig. 3 for further description.

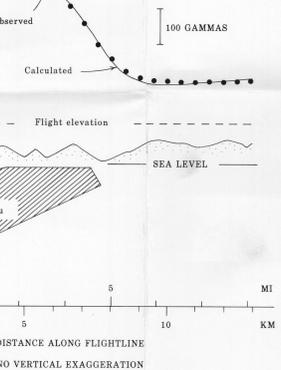


Figure 3. Magnetic model of Bone Mountain and vicinity assuming a buried source. See fig. 2 for further description.

MAP AND INTERPRETATION OF AEROMAGNETIC DATA FOR THE  
WILD ROGUE WILDERNESS, COOS AND CURRY COUNTIES, OREGON

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1983