



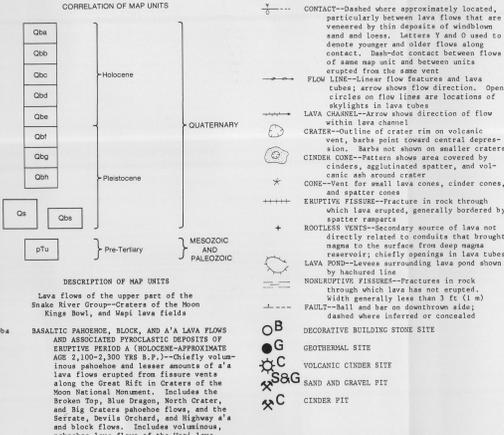
MAP A—GEOLOGIC MAP SHOWING THE LOCATION OF RESOURCES AND SITES INVESTIGATED IN RESOURCE EVALUATION

MINERAL RESOURCE POTENTIAL MAP OF THE GREAT RIFT INSTANT STUDY AREA, BLAINE, BUTTE, MINIDOKA, AND POWER COUNTIES, IDAHO

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DESCRIPTION OF MAP UNITS
Lava flows of the upper part of the Snake River Group—Craters of the Moon Kings Bowl, and Wapi Lava field.

Q4a BASALTIC PANDORRE, BLOC, AND A' LA LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD A (HOLOCENE)—APPROXIMATE AGE 2,100-2,300 YRS B.P.—Chiefly voluminous pahoehoe and lesser amounts of a' la lava flows erupted from fissure vents along the Great Rift in Craters of the Moon National Monument. Includes the Broken Top, Blue Dragon, North Crater, and Big Craters pahoehoe flows, and the Serate, Devil Orchard, and Highway A' la and block flows. Includes volcanism, pahoehoe lava flows of the Wapi Lava field and minor amounts of pahoehoe flow and cinders of the Kings Bowl lava field.

Q4b BASALTIC PANDORRE AND A' LA LAVA FLOWS OF ERUPTIVE PERIOD B (HOLOCENE)—APPROXIMATE AGE 2,300-500 YRS B.P.—Consists of pahoehoe lava flows erupted chiefly from fissure vents at Vermilion Clam and along the Great Rift for several miles northwest and southeast of Black Top Butte. Pahoehoe flows are typically shelly, slabby, and hummocky. Some flows appear dusty brown due to weathering. Highly vesicular glassy surfaces, particularly flow margins, are typically shelly pahoehoe. The fissures are parallel by spatter craters and thin blankets of spatter and ash. The vent areas at Devil Cauldron appear to have been lava lakes that underwent slight latest amounts of withdrawal of lava. The vent area for the Huckle-hole crater park flows presently shows only a small scintillate spire which probably represents a larger, buried vent complex. The source vent for the Rangefire flows may have been at Black Top Butte.

Q4c BASALTIC A' LA AND PANDORRE LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD C (HOLOCENE)—APPROXIMATE AGE 3,000-500 YRS B.P.—Consists of a later sequence of bulbous a' la lobes (Indian Wells north of the Kings Bowl and Sawtooth flows) with flow fronts as high as 80 ft (25 m), and an earlier sequence of pahoehoe a' la flows (South Echo, Sheep Trail Butte, Pleasure Butte, and Sentinel flows). The a' la flows are long flows, as much as 35 mi (55 km) with sinuosity. The flows contain small amounts of bedded cinder and spatter (fragments of cinder cones) as large as 300 ft (90 m) in longest dimension. Such a' la flows are covered by younger a' la and pahoehoe flows at their proximal ends, but flow directions and distribution of flows suggest that Big Cinder Butte cinder cone may be the source vent. The pahoehoe a' la flows were erupted from fissure vents on the Great Rift that extend from their Crater southeast to Sheep Trail Butte. The flows are mainly pahoehoe a' la, but their source vents and change to a' la lava several miles along their length.

Q4d A' LA LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD D (HOLOCENE)—APPROXIMATE AGE 4,500 YRS B.P.—Unit consists of bulbous a' la lobes with flow fronts as high as 80 ft (25 m). The a' la lobes have jagged, clinkery surfaces and contain the blocks of bedded cinder and spatter (fragments of a subventured cinder cone) as large as 300 ft (90 m) in longest dimension. All a' la flows of this unit are covered by younger flows at their proximal ends, but distribution and flow directions of the flows suggest that many or all of them were erupted from Silent Cinder cone.

Q4e BASALTIC PANDORRE AND A' LA LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD E (HOLOCENE)—APPROXIMATE AGE 6,500 YRS B.P.—Pahoehoe lava erupted from Grassy Cone cinder cone forms a long, narrow, composite flow (Grassy Cone and Laidlaw Lava flows) located between the main and western lobes of the Craters of the Moon lava field. Parts of the Grassy Cone and Laidlaw Lava flows consist of a' la lava, particularly where pahoehoe lava flowed over steep slopes. A thin blanket of volcanic ash covers the proximal parts of the Grassy Cone flow within several miles (kilometers) of Grassy Cone. The Lava Point a' la flow is a bulbous lobe with flow fronts as high as 80 ft (25 m).

Q4f BASALTIC PANDORRE LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD F (HOLOCENE)—UPPER PLISTOCENE, APPROXIMATE AGE 10,000-11,000 YRS B.P.—Pahoehoe lava flows located along the southwest margin and in the central part (Fremonts, and Buttes-in-the-flow) of the main lobe of the Craters of the Moon lava field; source vents are unknown but are located on the Great Rift probably near the present site of Sheep Trail Butte. Pahoehoe flows located at the east margin of the Craters of the Moon lava field (Walker Reservoir flows) are assigned to eruptive period F based on their radiocarbon age of about 10,500 yrs. The source vent for the Walker Reservoir flows is on the Great Rift probably near the present site of Crescent Butte.

Q4g BASALTIC PANDORRE AND A' LA LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD G (UPPER PLISTOCENE, APPROXIMATE AGE 12,000-14,000 YRS B.P.).—Pahoehoe flows moved as much as 22 mi (35 km) east (Sunset) and as much as 35 mi (55 km) southeast (Crazy) from a source vent at Sunset Cone cinder cone. The Sunset flow is covered by a thin wheat, (3 ft (2 m), of volcanic ash within 1 mi (1.5 km) east (downwind) of Sunset Cone. Before the eruption of the flow from Sunset Cone, a' la flows with jagged, blocky, clinkery surfaces erupted from 4 or more vents near Lava Creek in the southern Pioneer Mountains. The flows from the Lava Creek vents moved as much as 22 mi (35 km) east onto the Snake River Plain and about 2 mi (3 km) north in a tributary of Dry Fork Creek. Pyroclastic deposits of unknown thickness mantle vent areas.

Q4h BASALTIC PANDORRE LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS OF ERUPTIVE PERIOD H (UPPER PLISTOCENE)—APPROXIMATE AGE ABOUT 15,000 YRS B.P.—Pahoehoe lava flows are located along the southwestern (Glema flow) and western margins (Near Den Lake flow) of main lobe of the Craters of the Moon lava field; their source vents are unknown. Unit also includes isolated exposures of pahoehoe lava east of the Great Rift near the Wichman and Two Point Butte cinder cones; isolated exposures of pahoehoe lava with no known source vents are located on the south margin (Brown flow) and on the east margin (Baseline flow) of the main lobe of the Craters of the Moon lava field.

Q4i SURFICIAL DEPOSITS (LATE PLISTOCENE)—Chiefly alluvial deposits along streams and colluvial deposits at the base of steep slopes. This deposit of eolian sand and silt are not mapped.

Q4j Basalt flows of the lower part of the Snake River Group

Q4k BASALT LAVA FLOWS AND ASSOCIATED PYROCLASTIC DEPOSITS (PLISTOCENE)—Chiefly pahoehoe lava flows, cinders, and ash, with minor amounts of a' la lava. Older units are covered by eolian deposits as much as 5 ft (1.5 m) thick. Younger flows have fresh, glassy surfaces and thin, (3 ft (1 m), localized mantles of eolian sediment. Pyroclastic deposits occur mainly near vent areas.

Q4l PRE-TERTIARY ROCKS OF THE PIONEER MOUNTAINS, MESOZOIC AND PALEOZOIC SEDIMENTARY AND IGNEOUS ROCKS, UNMAPPED

CONTACT—Dashed where approximately located, particularly between igneous units that are separated by thin deposits of windblown sand and loess. Letters Y and G used to denote younger and older flows along contact. Dash-dot contact between flows of same map unit and between units erupted from the same vent.

FLOW LINE—Lineer flow features and lava tubes; arrow shows flow direction. Open circles on flow lines are locations of skylights in lava tubes.

LAVA CHANNEL—Arrow shows direction of flow within lava channel.

CRATER—Outline of crater rim on volcanic vents, bars point toward central depression. Bars not shown on smaller craters.

CINDER CONE—Pattern shows area covered by cinders, applications spatter, and volcanic ash around crater, and volcanic ash around crater.

CONE—Vent for small lava cones, cinder cones, and spatter cones.

ERUPTIVE FISSURE—Fracture in rock through which lava erupted, generally bordered by spatter reservoir.

ROOTLESS VENT—Secondary source of lava not directly related to conduct that brought magma to the surface from deep magma reservoir; chiefly fissures in lava tubes.

LAVA POOL—Lavae surrounding lava pool shown by hatched line.

NONERUPTIVE FISSURE—Fractures in rock through which lava has not erupted.

FAULT—Wall and bar on downthrown side; dashed where inferred or concealed.

DECORATIVE BUILDING STONE SITE

GEOHERMAL SITE

VOLCANIC CINDER SITE

SAND AND GRAVEL PIT

CINDER PIT

STUDIES RELATED TO WILDERNESS
Bureau of Land Management Wilderness Study Areas

The Federal Land Policy and Management Act (Public Law 96-379, October 21, 1978) requires the U.S. Geological Survey to conduct mineral surveys on certain areas to determine their resource potential. Results must be made available to the public and be submitted to the President and the Congress. This report presents the results of a geologic and geophysical survey of the Great Rift Instant Study Area (33.1, Blaine, Butte, Minidoka, and Power Counties, Idaho).

SUMMARY STATEMENT
Aeromagnetic, gravity, and magnetotelluric investigations of the study area have yielded data on the magnetic field, gravity, and magnetotelluric fields, and provided data for an assessment of potential resources in the area. Magnetic anomalies in the area are complex and result from both strongly magnetized surface and near-surface lava flows and magnetic rocks underlying the flows. The magnetic data suggest that a deeply buried intrusive body underlies the study area. Gravity anomalies are related to variations in the thickness of Cenozoic volcanic rocks and possibly to the inferred intrusive body. Magnetotelluric studies have provided data for interpreting the structure and possible rock types in the upper 6 mi (10 km) of the crust.

No mining districts exist in either the northern or southern parts of the Great Rift Instant Study Area. No mining claims have been filed in the northern part of the study area, but a group of eight claims, presumably for building stone, was located at Pillar Butte in 1980.

Identified resources in the study area are confined mainly to decorative building stone and volcanic cinders (figs. 2 and 3). Decorative building stone consists chiefly of slabs of pahoehoe lava, 1-4 in. (2.5-10 cm) thick, that occur at the margin of the study area (Kuntz and others, 1983).

INTRODUCTION
The Great Rift Instant Study Area is located in the Snake River Plain of south-central Idaho. The Snake River Plain is an area of geologically young (Pliocene, Pleistocene, Holocene) basaltic volcanism. The Great Rift Instant Study Area consists of two parts. The northern part comprises 28,800 acres (1,165 km²) partly surrounding the Craters of the Moon National Monument and encompassing a large part of the Craters of the Moon lava field (fig. 1). The southern part of the study area comprises 86,400 acres (350 km²) encompassing most of the Wapi Lava field (fig. 1).

Lava flows in both parts of the study area are mainly fresh, glassy-surfaced pahoehoe, but a' la flows and cinders at and near vents are present locally. The flow are as young as they are nearly free of soil and vegetation. The flows were derived from eruptive fissures and fissure-controlled cinder cones along the Great Rift. The Great Rift is 3.5 mi (5.6 km) long, 1.5 mi (2.4 km) wide (2-8 km) wide of cinder cones and eruptive vents, and 1.5 mi (2.4 km) wide of the southern Pioneer Mountains on the northern margin of the Craters of the Moon National Monument, to Pillar Butte near the southern margin of the Snake River Plain. The younger flows of the study area lie on older flows that are more weathered and typically covered by a variable thickness of eolian sand and alluvial sediments (Kuntz and others, 1983).

Gravity, magnetic, and magnetotelluric methods have been employed to infer the structure and type of rock underlying the basaltic flows at the surface of the study area. A aeromagnetic investigation shows that magnetic anomalies are very complex in the study area and reflect 3 main factors: (1) local surface relief, (2) direction of magnetization, and (3) thickness and depth of sequence of flows with normal or reverse magnetization. A prominent magnetic anomaly that reflects a deeply buried source is a magnetic high in the area of the Craters of the Moon National Monument. Analysis of the anomaly by multibody inversion analysis (Bhattacharyya, 1980) indicates that the magnetic high is a highly magnetized body of rock with a top 0.5-2 mi (0.8-3.2 km) below sea level. The magnetic data suggest that a large intrusion exists across the Snake River Plain from Craters of the Moon National Monument to American Falls. The axis of the gravity low is roughly coincident with the magnetic high and the Great Rift in the northern part of the study area. Both the gravity low and magnetic high could be produced by a granitic intrusive body.

A magnetotelluric traverse across the northern part of the study area was used to infer the electrical resistivity characteristics. The uppermost layer, about 0.5 mi (0.8 km) thick, has a resistivity of 8-17 ohm-m, and is postulated to consist of basalt, dykeite flow and tuff, and sediments. The second layer, with resistivities of 200-400 ohm-m, and a thickness of 2-3 mi (3.2 km) is believed to be the basement complex of carbonate, igneous, or metamorphic rocks of or near the Crater type. The lowest layer has a resistivity of less than 10 ohm-m and is believed to be about 6.0 mi (9.6 km) thick (Kuntz and others, 1983).

MINERAL RESOURCE POTENTIAL
The mineral resource potential of the Great Rift Instant Study Area is discussed in two parts—the Craters of the Moon lava field to the north, and the Wapi Lava field to the south. Resources are classified as identified (measured, indicated, and inferred), and undiscovered (speculative and hypothetical), reflecting decreasing degrees of geological assurance (U.S. Geological Survey, 1980).

Identified resources in the study area are confined mainly to decorative building stone and volcanic cinders. The following criteria used to evaluate decorative stone potential were developed by examining a site near Black Butte, near Shoshone, Idaho, where similar material is currently being extracted. Slab pahoehoe of decorative stone must be a minimum of 1 in. (2.5 cm) thick and can be as thick as 4 in. (10 cm), although a greater thickness would probably be 2 in. (5 cm). These pieces should be at least 1 ft² (0.09 m²) in surface area and be flat to gently curved. These criteria favor material from pahoehoe flows. The degree of surface roughness ranges from levelled (smooth) (1/4 in. (0.5 cm) to medium-irregular (1/2 in. (1.25 cm) to (Champlon, 1973). Suitability criteria, such as type of texture, color, and presence of absence of lichens, which would lead to rejection of one place over another, are more difficult to evaluate without a detailed market study and are not considered in the overall resource estimates. However, some of the observed decorative stone sites contain material whose surficial features are unique, possibly enhancing desirability and marketability. Resources are based on surface material only, although this does not rule out the eventual use of material from underlying layers (where present) or fragments smaller or thinner than the dimensions stated above.

Because of the wide range in observed bulk densities of the pahoehoe (118.6 lb/ft³ to 217.4 lb/ft³, 1.9 g/cm³ to 3.5 g/cm³), tonnage represent an average of observed and inferred material. Slab pahoehoe, suitable for use as decorative building stone, was observed: (1) at the margins of flow and flow units, (2) in the vicinity of vent rims, and (3) in square pits. This material is more prevalent in the Wapi Lava field.

Volcanic cinders, prevalent in the Craters of the Moon lava field, occur low cinder cones within the lava field but especially along the Great Rift. This material has a variety of uses so that the various Craters of the Moon Lava field.

No known mining claims occur within the Craters of the Moon lava field, but sand and gravel and volcanic cinders have been removed from two sites adjacent to the eastern boundary of the lava field (fig. 2, map A). In the material, used locally as road metal, sand and gravel have been removed from a small pit near Tract Lake (DE/4 sec. 24, T. 1 S., R. 27 E., fig. 2, map A).

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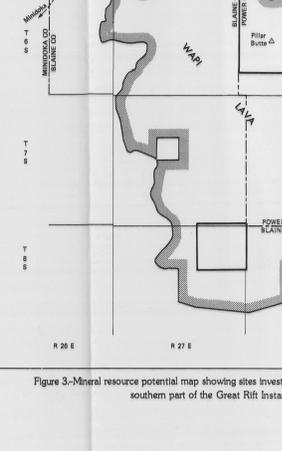


Figure 3—Mineral resource potential map showing sites investigated for decorative building stone in the southern part of the Great Rift Instant Study Area, Idaho

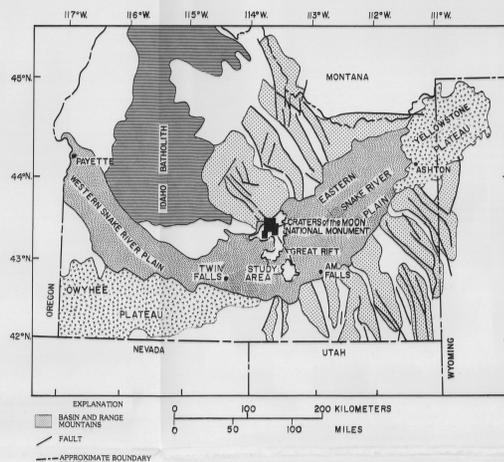


Figure 1—Index map showing the location of the Great Rift Instant Study Area, Idaho

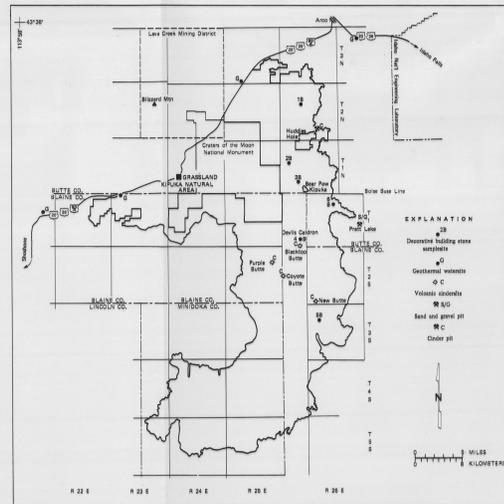


Figure 2—Mineral resource potential map showing the location of resources in the northern part of the Great Rift Instant Study Area, Idaho

Six sites (fig. 2, map A) were examined where decorative building stone was noted during aerial reconnaissance. Identified resources of this stone total 1,550,000 tons (15,740,000 cu yd); an additional 10,000 tons (10,100 cu yd) is inferred. The stone is mainly in but not confined to pahoehoe-type flows of eruptive period A (Kuntz and others, 1983) which cover an estimated 20 percent of the Craters of the Moon lava field. Among the types of pahoehoe observed, the most unique variety is found in the Blue Dragon flow, noted for its iridescent blue magnetization. A prominent magnetic anomaly that reflects a deeply buried source is a magnetic high in the area of the Craters of the Moon National Monument. Analysis of the anomaly by multibody inversion analysis (Bhattacharyya, 1980) indicates that the magnetic high is a highly magnetized body of rock with a top 0.5-2 mi (0.8-3.2 km) below sea level. The magnetic data suggest that a large intrusion exists across the Snake River Plain from Craters of the Moon National Monument to American Falls. The axis of the gravity low is roughly coincident with the magnetic high and the Great Rift in the northern part of the study area. Both the gravity low and magnetic high could be produced by a granitic intrusive body.

A magnetotelluric traverse across the northern part of the study area was used to infer the electrical resistivity characteristics. The uppermost layer, about 0.5 mi (0.8 km) thick, has a resistivity of 8-17 ohm-m, and is postulated to consist of basalt, dykeite flow and tuff, and sediments. The second layer, with resistivities of 200-400 ohm-m, and a thickness of 2-3 mi (3.2 km) is believed to be the basement complex of carbonate, igneous, or metamorphic rocks of or near the Crater type. The lowest layer has a resistivity of less than 10 ohm-m and is believed to be about 6.0 mi (9.6 km) thick (Kuntz and others, 1983).

The southern part of the study area contains no known mining districts. Bonanza Ber, with gold placers, lies approximately 5 mi (8 km) southeast of the southern tip of the Wapi Lava field. In 1980, during the field work was completed, a group of eight claims was located on Pillar Butte, presumably for building stone. Decorative building stone has been removed from several sites within and adjacent to the study area. Four areas (fig. 3, map A), representing approximately 16 percent of the Wapi Lava field, were investigated by aerial reconnaissance, aerial color photography, and on foot. These areas contain approximately 91,000 tons (93,000 cu yd) of indicated decorative building stone resources. A unique variety of decorative stone with a distinctive wire rope-like appearance, several uses of which had been extracted, geological assurance (U.S. Geological Survey, 1980).

Reconnaissance studies and examination of aerial photographs reveal that inferred surface resources may approach a half million tons of seable stone.

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