

FIGURE 1.—INDEX MAP

INTRODUCTION

This report presents results of a high-resolution seismic-reflection survey conducted by the U.S. Geological Survey to determine the subsurface characteristics of Penobscot Bay, Maine. Penobscot Bay is one of the major embayments along the U.S. Atlantic coast (fig. 1), extending more than 50 km from the Gulf of Maine to the mouth of the Penobscot River. The bay constitutes an area of about 1,100 km² and is divided into three main passages by numerous islands.

The bay and surrounding coastal area have had a complex history of glaciation and sea-level change since late Wisconsinan time. The last ice sheet advanced across the area about 20,000 years ago, eroding the bedrock surface, removing most older sediments, and isostatically depressing the crust (Schafer and Hartshorn, 1965; Stuyver and Borne, 1975; Thompson, 1979). Deglaciation followed between about 12,700-13,500 years ago and was accompanied by a marine transgression (depression of crustal depression) with ice and sea in contact (Stuyver and Borne, 1975; Thompson, 1979, 1982; Smith, 1982). Isostatic recovery of the crust then caused the coast to emerge during the immediate postglacial period, and emergence was completed by about 11,500 years ago (Stuyver and Borne, 1975; Thompson, 1979, 1982; Thompson and Smith, 1982). Finally, readvance of the Maine coast has accompanied the Holocene rise of sea level that began about 9,000 to 11,000 years ago (Bloom, 1963; Kaye and Harghorn, 1964; Schnitzer, 1974; Oldale and others, 1983).

In the maps presented here, we illustrate and briefly describe the sea-floor topography, bedrock morphology, and variable sediment thickness within Penobscot Bay that were either produced or modified during the period of geologic change since late Wisconsinan time.

ACKNOWLEDGMENT

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FIELD WORK

Numerous subbottom profiles were collected along 376 km of trackline within the study area during June 11-17, 1983; three were selected for interpretation in this report (figs. 2 and 3). The data were obtained by using a Doherty seismic system (800- to 1,500-Hz band pass, quarter-second sweep rate) that had a resolution of about 1 to 2 m. The locations and spacings of the cruise tracks were based on the bathymetry (fig. 2) and were designed to cover the entire bay. Navigational control for all tracklines was provided by loan-C and was supplemented by radar and visual fixes. The ship's speed during the survey typically was 7 km/h.

¹Use of trade names in this publication is for descriptive purposes only and does not constitute endorsement by the U.S. Geological Survey.

REFERENCES CITED

Bloom, A. L., 1963, Late Pleistocene fluctuations of sea level and postglacial crustal rebound in coastal Maine: *American Journal of Science*, v. 261, p. 882-979.

Hussey, A. M., chief compiler, 1967, Preliminary geologic map of Maine: Augusta, Maine, Maine Geological Survey, scale 1:500,000, 1 sheet.

Kaye, C. A., and Harghorn, E. S., 1964, Late Quaternary sea-level change and crustal rise at Boston, Massachusetts, with notes on the autostratigraphy of the area: *Geological Society of America Bulletin*, v. 75, p. 83-89.

Knebel, H. J., and Scanlon, K. M., in press, Sedimentary framework of Penobscot Bay, Maine: *Marine Geology*.

National Oceanic and Atmospheric Administration, 1982, Tide tables 1983, east coast of North and South America: Washington, D.C., U.S. Government Printing Office, 285 p.

National Ocean Survey, 1981, Bathymetric map, Penobscot Bay and approaches: Washington, D.C., National Ocean Survey chart 13302, scale 1:60,000, 1 sheet.

Oldale, R. N., Womack, L. E., and Whitney, A.B., 1983, Evidence for a postglacial low relative sea-level stand in the drowned delta of the Merrimack River, western Gulf of Maine: *Quaternary Research*, v. 15, p. 323-336.

Schafer, J. P., and Hartshorn, J. H., 1965, The Quaternary of New England, in Wright, H. E., Jr., and Fry, D. G., eds., *The Quaternary of the United States*: Princeton, N.J., University Press, p. 113-139.

Schnitzer, Detmar, 1974, Postglacial emergence of the Gulf of Maine: *Geological Society of America Bulletin*, v. 85, p. 691-694.

Smith, G. W., 1982, End moraines and the pattern of last ice retreat from central and south coastal Maine, in Larson, G. L., and Stone, R. D., eds., *Late Wisconsinan glaciation of New England*: Dubuque, Iowa, Kendall/Hunt Publishing company, p. 195-209.

Stuyver, H., and Borne, H. W., Jr., 1975, Late Quaternary marine invasion in Maine—its chronology and associated crustal movement: *Geological Society of America Bulletin*, v. 86, p. 89-104.

Thompson, W. B., 1979, Surficial geology handbook for coastal Maine: Augusta, Maine, Maine Geological Survey, 49 p.

1982, Recession of the Late Wisconsinan ice sheet in coastal Maine, in Larson, G. L., and Stone, R. D., eds., *Late Wisconsinan glaciation of New England*: Dubuque, Iowa, Kendall/Hunt Publishing company, p. 211-228.

Thompson, W. B., and Smith, G. W., 1983, Pleistocene stratigraphy of the Augusta and Waldoboro areas, Maine: *Maine Geological Survey, Guidebook for the 46th Annual Meeting of the Friends of the Pleistocene*, 37 p.

KEY TO ACOUSTIC UNITS AND EXPLANATION FOR SEISMIC PROFILES

Qh—Marine sediments Qdu—Glacial drift
Qf—Fluvial deposits Pz—Bedrock
Qp—Glaucmarine detritus

Horizontal scales are approximate.
Vertical scales are based on an assumed compressional-wave velocity of 1,500 m/s.
V.E. = Vertical exaggeration.
Profile location shown in figure 2.

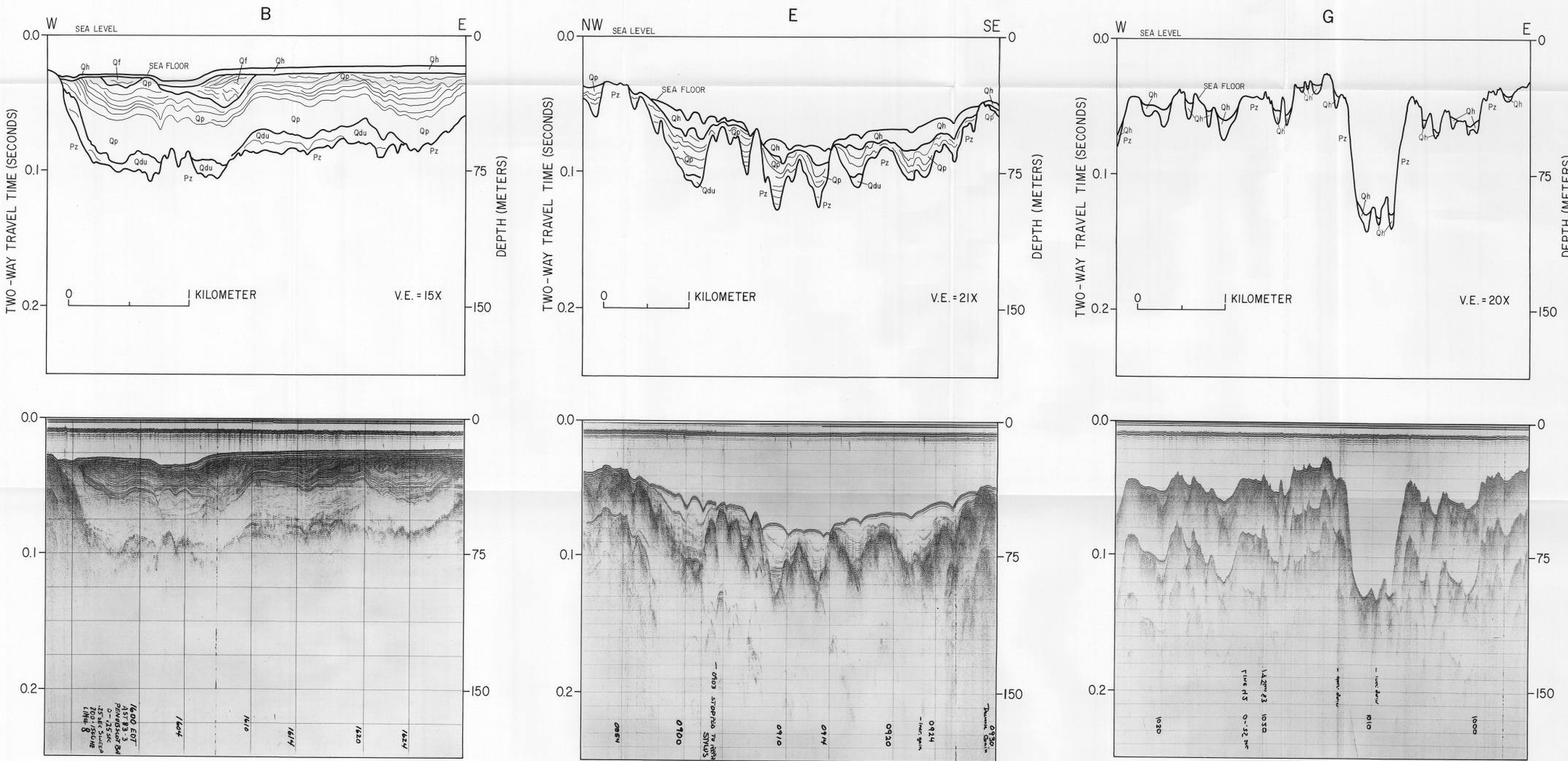


FIGURE 3.—SELECTED SEISMIC PROFILES AND INTERPRETIVE SECTIONS.

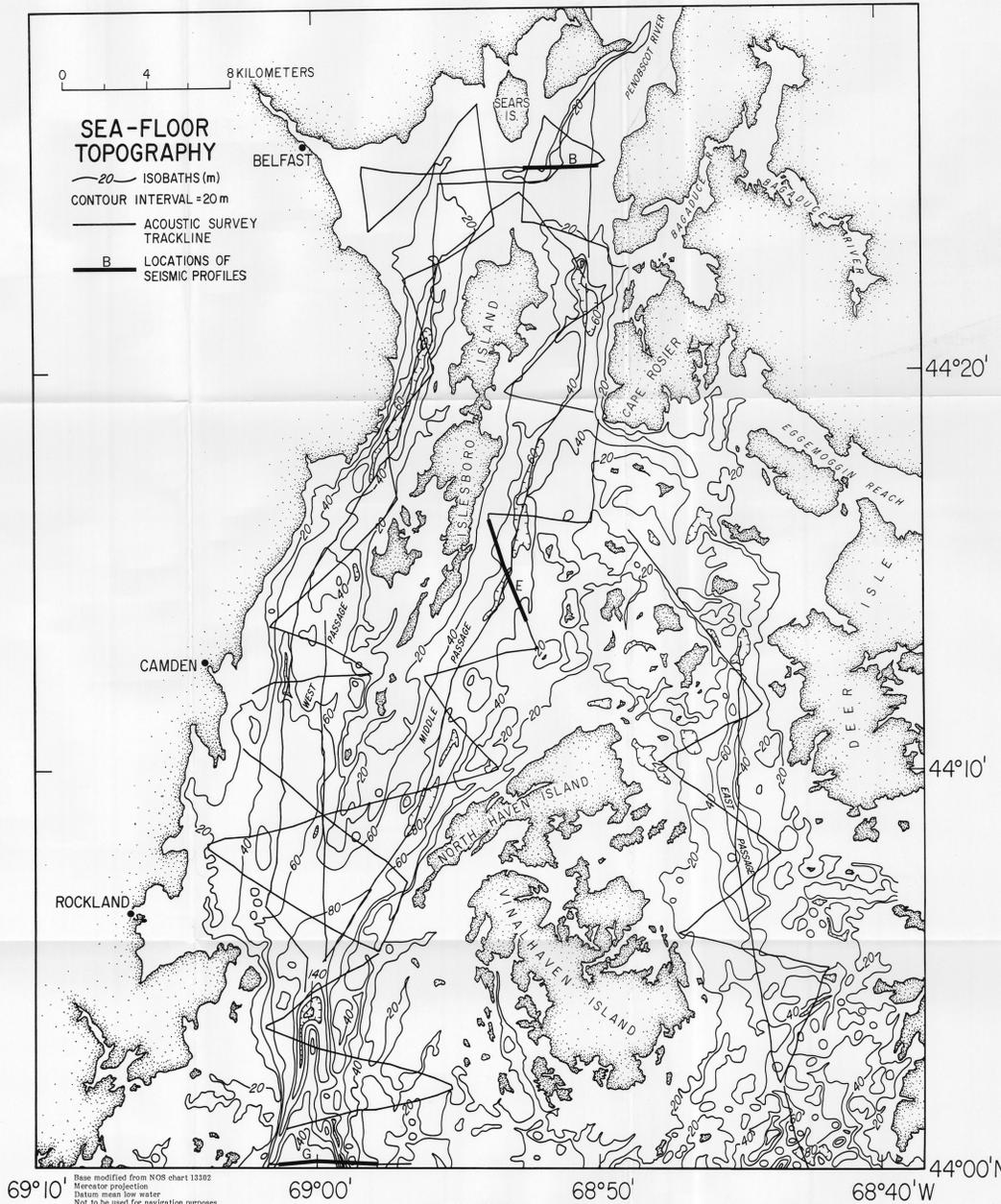


FIGURE 2.—SEA-FLOOR TOPOGRAPHY AND LOCATIONS OF ACOUSTIC SURVEY TRACKLINES AND SELECTED SEISMIC PROFILES.