

**INTRODUCTION**

Large gas accumulations are contained in low-permeability reservoir rocks of Late Cretaceous and early Tertiary age in the Greater Green River Basin of Wyoming, Colorado, and Utah (fig. 1). The vertical and horizontal limits of these gas-bearing reservoirs have been described by Law and others (1980), Spencer (1983), and Law (1984). These reports have shown that the gas-bearing rocks are coincident with anomalously high formation pressures (overpressures) and that the top of overpressuring occurs at a vitrinite reflectance ( $R_m$ ) mean random vitrinite reflectance) value of about 0.8 percent. Thus, the primary purpose of this study is to provide a thermal maturity map of the Greater Green River Basin showing the subsurface elevation of 0.8 percent  $R_m$ . Although thermal maturity studies are most often utilized in investigations of hydrocarbon generation, preservation, and migration, they also provide important information regarding the thermal and burial history of sedimentary basins, such as the studies in the northern part of the Greater Green River Basin by Lickus and others (1984), Kesser (1984), and Pollastro and Barker (1984).

**METHODS**

413 core and cuttings samples of coal, shale, siltstone, and sandstone were collected from 126 wells in the basin. The number of samples per well ranged from 1 to 47. Vitrinite reflectance measurements of coal samples were determined using reflected light under oil immersion. The kerogen from the noncoal lithologies was extracted by an acid maceration technique (King and others, 1963; Saby, 1970). The maceration process, for the most part, yielded good results and the data obtained from the analyses are considered reliable. Coal samples and thin organic-matter-rich lenses were not processed with the acid maceration technique.

Owing to the wide geographic and stratigraphic distribution of the samples, the basin was subdivided into four areas (fig. 2). Within each area, a linear regression line of depth (in feet) versus  $R_m$  was constructed, and the slope of the least squares regression line determined. In wells containing less than five samples distributed over a thin stratigraphic unit, the appropriate elevations of 0.8 percent  $R_m$  were calculated using the appropriate slope obtained from the regression line. In wells containing five or more samples distributed over a thick stratigraphic unit, the elevations of 0.8 percent  $R_m$  were calculated by constructing a separate linear regression plot for each well. Data from four wells were not used because the histograms were widely scattered and the calculated depths differed significantly from surrounding wells. The contours were drawn and were largely influenced by structure, especially in areas where the data were sparse.

**DISCUSSION**

Law (1984) has shown that gas generation and occurrence in the Greater Green River Basin is related to organic richness, subsurface temperature, overpressuring, and thermal maturity. Therefore, the dip of the 0.8 percent  $R_m$  contour would be best utilized in conjunction with other information such as the subsurface temperature map of the Greater Green River Basin by Law and Smith (1983).

Cross section A-A' (fig. 4), from the Church Buttes gas field in the southeastern part of the basin east to the Rawlins uplift, illustrates the general relationship of the 0.8 percent  $R_m$  isotherm to the structural dip. The isotherm is generally less than, or equal to, the structural dip. Stach and others (p. 43-44) (1975), and Barker and others (1983) have suggested that the degree to which isotherm (here isotherm) lines conform to structure provides information on the temporal relationships between thermal maturation and structural deformation. This line of reasoning would suggest that in these areas the development of thermal maturation was contemporaneous with structural deformation.

However, recent work by Law and others (1986) in the area of the Patrick Draw oil and gas field provides an alternative explanation. They suggest that the anomalously high level of thermal maturity in the Patrick Draw area is due to thermal alteration of organic matter by relatively hot fluids migrating upward along fault planes and (or) fracture systems. The timing of such a heating event may be independent of deformation and may occur at any time during the burial history of the affected rocks. Therefore, the conventional way of determining the relative timing of thermal maturation may be misleading in this area.

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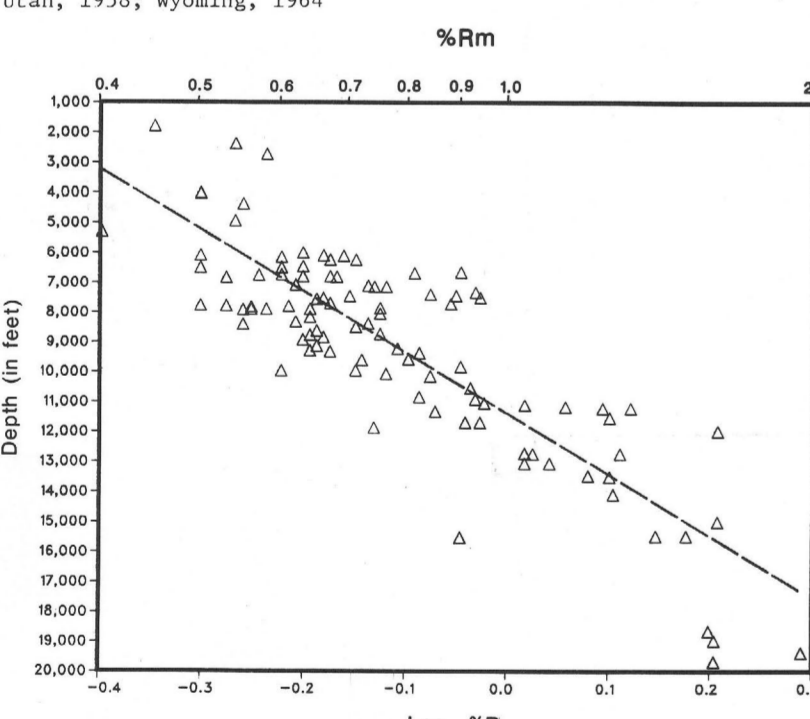


Figure 3.—Sample regression plot showing data from the western Green River Basin area plotted on a graph of depth (in feet) versus  $\log_{10} R_m$ . Dashed line is least squares linear regression line.

**EXPLANATION**

CONTOUR—showing elevation in feet above mean sea level of 0.8 percent  $R_m$ . Dashed where appropriate. Markers show closed depressions. Contour interval 1,000 ft.

ANTICLINE—showing direction of plunge.

▲ WELL LOCATION AND CALCULATED ELEVATION OF 0.8%  $R_m$ .

▲ WELL FROM WHICH  $R_m$  WAS MEASURED FROM CORE.

● WELL FROM WHICH  $R_m$  WAS MEASURED FROM CUTTINGS.

○ WELL FROM WHICH  $R_m$  DATA WERE AVAILABLE BUT NOT USED IN COMPUTATION.

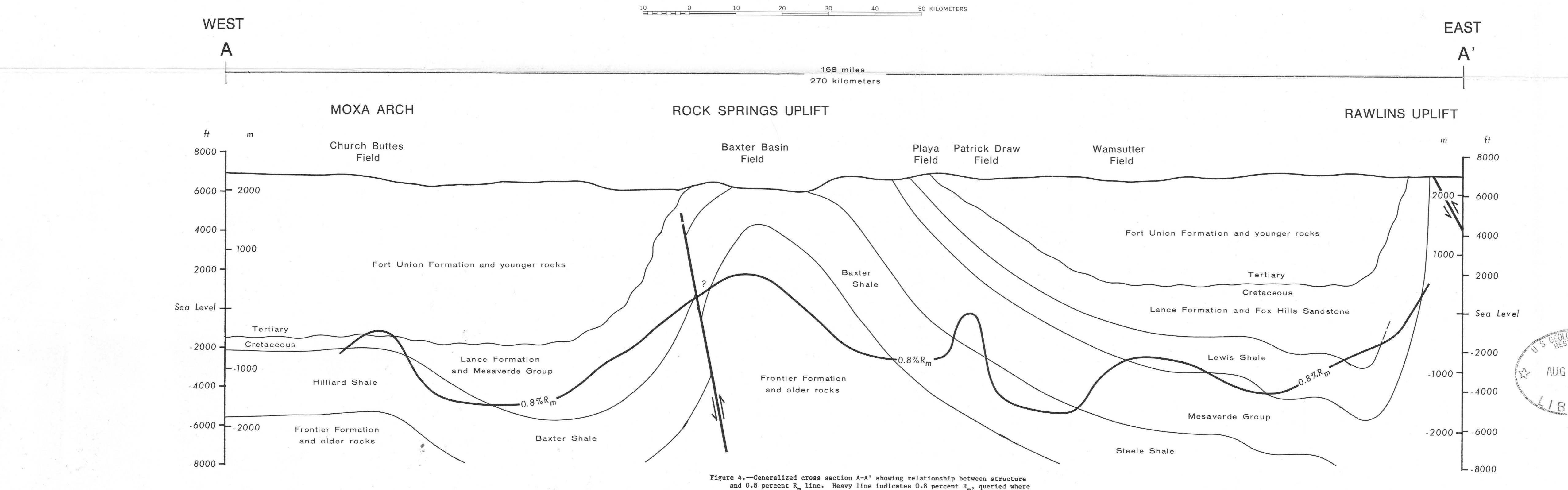


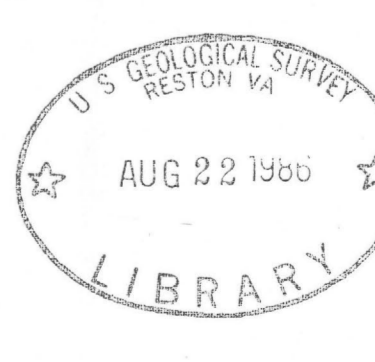
Figure 4.—Generalized cross section A-A' showing relationships between structure and 0.8 percent  $R_m$  isotherm. Heavy line indicates 0.8 percent  $R_m$  queried where uncertain; arrow indicate direction of relative motion along faults; irregular line indicates Tertiary-Cretaceous unconformity.

**THERMAL MATURITY MAP SHOWING SUBSURFACE ELEVATION OF 0.8 PERCENT VITRINITE REFLECTANCE IN THE GREATER GREEN RIVER BASIN OF WYOMING, COLORADO, AND UTAH**

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