



PRELIMINARY GEOLOGIC MAP OF THE WHITE ROCK QUADRANGLE, RIO BLANCO AND MOFFAT COUNTIES, COLORADO

By
G. N. Pipirinos and Gene C. Rosenlund
1977

CORRELATION OF MAP UNITS

Q1s	Qa1	Holocene	QUATERNARY	
Tg3	Tg2	Eocene		TERTIARY
HA	L2	Paleocene		
L1	Tg1	Paleocene		
Tw	Tfs	Paleocene	CRETACEOUS	
Kw	Kw1	Upper Cretaceous		
Kw		Upper Cretaceous		

DESCRIPTION OF MAP UNITS

Q1s LANDSLIDE DEPOSITS (HOLOCENE)—Slump blocks developed in the Garden Gulch Member of the Green River Formation and associated with coal seams and shale in the Williams Fork Formation.

Qa1 ALLUVIAL DEPOSITS (HOLOCENE)—Mostly silt and sand as much as 40 ft (12 m) thick along Strawberry and Deep Channel Creeks.

GREEN RIVER FORMATION (EOCENE)

Anvil Points Member—Anvil Points units are westward thinning sequences made up dominantly of sandstone and lesser amounts of finer grained clastics. The units are dominantly fluvial in the east and grade westward into units entirely lacustrine.

Unit 3—Light gray to white, fine- to medium-grained, calcareous sandstone to thick-bedded sandstone forming ledges and cliffs. Locally differential cementation and weathering have produced irregularly shaped sandstone nodules. Unit crops out only in the SW sec. 18, T. 3 N., R. 95 W. and at the section line between sections 12 and 13, T. 3 N., R. 95 W. Thickness about 240 ft (73 m). Top is eroded.

Unit 2—Light gray to white, fine- to medium-grained, calcareous sandstone, medium to thick bedded; forms ledges and cliffs. The unit thickens from 150 ft (46 m) at the NW sec. 30, T. 3 N., R. 95 W. to 280 ft (85 m) near the center of sec. 18, T. 3 N., R. 95 W. It thickens northward to as much as 450 ft (137 m) near the center of sec. 12, T. 3 N., R. 95 W. Unit same as upper part of unit 2 (Tg2) of the Anvil Points as mapped in the Indian Valley quadrangle to the west (Pipirinos and Rosenlund, 1976) and approximately equivalent to the upper part (Tg2u) of the Anvil Points as mapped in the Buckskin Point quadrangle to the south (Pipirinos and Johnson, 1975).

Unit 1—Mostly sandstone, some siltstone and shale. The sandstone is light gray, fine to coarse grained, calcareous, medium bedded to massive, and forms cliffs, ledges and slopes. The sandstone beds are lenticular and commonly occupy channels cut in the underlying rocks in the southeastern part of the area. The siltstone and shale are gray and form slopes. The unit includes a few thin red beds near the base in the southeastern part of the area. The contact with the underlying Wasatch Formation is gradational, intertonguing and generally poorly exposed. Consequently a persistent chert pebble zone about 600-700 ft (183-213 m) below the top of this unit was arbitrarily mapped as the base of the Anvil Points along the foot of the east-facing escarpment that crosses the mapped area in a northwestern-southeastern direction. Lacustrine shale, siltstone and sandstone beds are scarce in the southern part but make up the upper half to one-third of the unit in the northwestern part of the quadrangle. In the southwestern part of the quadrangle, the chert pebble zone could not be found and the contact of the Anvil Points and the underlying Wasatch Formation is placed at the indefinite change from cliff-forming sandstone above to also-forming variegated claystone and siltstone below. As mapped, the thickness of unit 1 of the Anvil Points ranges from about 600 ft (183 m) to 800 ft (244 m) and averages about 700 ft (214 m). It is thickest in the southeastern part of the quadrangle and thins both westward and northward. It equals the lower part (Tg1) of the Anvil Points as mapped in the Indian Valley quadrangle to the west and the Buckskin Point quadrangle to the south (Pipirinos and Rosenlund, 1976; Pipirinos and Johnson, 1975). The top of unit 1 was arbitrarily placed at the top of the lowest persistent and mappable limestone bed (L1) of Garden Gulch Member. It is described at the end of the section on the Garden Gulch Member.

Tw WASATCH FORMATION (EOCENE AND PALEOCENE)—Mainly sandy claystone and siltstone containing lenses and channels of medium- to coarse-grained cross-bedded sandstone. The sandstone beds are thicker, more numerous, and more conspicuous than they are in the quadrangles adjacent to the west, southwest, and south. The upper part of the Wasatch yielded vertebrate fossils of early to middle Eocene age. These fossils were collected from outcrops in the southwestern part of the area stratigraphically about 750 ft (229 m) below the top of the formation. They were identified by L. C. Gazin of the U.S. National Museum. The lower part of the Wasatch as mapped includes rocks of Paleocene age. Fossil spores and pollen from a 5-ft (1.5-m) thick carbonaceous shale bed about 735 ft (224 m) above the base of the Wasatch were identified by K. L. Techow as being of late Paleocene age. The lower part of the Wasatch in the White Rock quadrangle is of the same age, lithology, and appearance as rocks assigned to the Fort Union Formation in central and southern Wyoming (Venne and Pipirinos, 1954; locs. 8, 9, 10; Pipirinos, 1955, p. 100; 1961, p. A10-A13; Denison and Pipirinos, 1969; Reynolds, 1968; 1971). It was tentatively included in and mapped with the Wasatch pending further investigation into the nature of the Fort Union-Wasatch contact and the feasibility of mapping it. In general, the claystone beds in the Eocene portion are variegated ochre, lavender, red, and green; the Paleocene claystone beds tend to be drab gray and white, and include a few thin rusty-brown to nearly black ferruginous zones, and at least one papery fissile brown carbonaceous shale bed already mentioned. The change from the variegated Eocene to the drab Paleocene part of the Wasatch occurs in an interval that generally is poorly exposed. If there is an erosional break between the two parts of the Wasatch in this area it is inconspicuous and difficult to find. The basal contact of the Wasatch is concealed throughout the quadrangle and whether it is unconformable

cannot be determined. The Wasatch as mapped ranges in thickness from about 2,300 ft (701 m) to 3,300 ft (1,006 m). It probably is about 4,500 ft (1,372 m) in the subsurface of the southwestern part of the mapped area based on the nearest control point, test hole 24-23 Federal; it was drilled in 1975 by Jack Grynborg and Associates in sec. 23, T. 2 N., R. 95 W. of the White River City quadrangle, about 0.5 mile (0.8 km) southwest of the southwest corner of the mapped area. The thickness of the Eocene part of the Wasatch averages about 2,500 ft (762 m) throughout the quadrangle.

UNIT 12 (L2) (PALEOCENE)

Tfs Sandstone member—Dominantly sandstone, light gray, which weathers buff, tan or brown, medium to coarse grained. Several conglomerate beds 1-4 ft (0.3-1.2 m) thick occur in the base in 40-70 ft (12-21 m) of the sandstone. Clasts are principally black, gray, and red chert and quartzite; pale-red, lavender, and pink volcanic rocks with white feldspar phenocrysts are common. The conglomerate clasts are locally as much as 6 in. (15 cm) in diameter, but generally average about 1 in. (2.5 cm). Member contains lesser amounts of brown-weathering iron-oxide concretions, limestone concretions and olive-gray shale, claystone and siltstone, carbonaceous and coaly shale and coal. The iron-oxide concretions are spheroidal and commonly as much as 3 ft (0.9 m) in diameter; the limestone concretions are spheroidal to ovoid spheroidal, as much as 2 ft (0.6 m) in diameter and weather rusty-yellow and buff. The member contains several carbonaceous zones one of which is about 27 ft (8 m) thick and a few coal beds one of which is at least 6 ft (1.8 m) thick. The coal occurs about 600 ft (183 m) above the base in the NE sec. 35, T. 3 N., R. 95 W. The sandstone beds of this member make prominent ridges and spectacular, steep bare rock walls. The intervening softer sequences make strike valleys. The member ranges in thickness from about 1,200 ft (366 m) to 1,400 ft (427 m) on the surface, and from about 800 ft (244 m) to 1,200 ft (366 m) in the subsurface in the near-east wells in the White River City and Indian Valley quadrangles. The conglomeratic beds in the lower part of the sandstone member may or may not be equivalent to the Ohio Creek Formation of Paleocene age. In the White Rock quadrangle, the conglomerate commonly contains volcanic clasts; according to D. L. Gaskill of the U.S. Geological Survey (unpub. data, 1976) however, volcanic clasts are absent in the Ohio Creek Formation in the type locality north of the West Elk Mountains in west-central Colorado. The base of the Fort Union Formation has long been thought to rest with erosional unconformity on underlying rocks in this region (Gazin, 1930, p. 42, 75). The contact is characterized by a change from coarse-grained rocks above to fine-grained rocks below.

L2—The L2 zone—The L2 zone comprises two light-gray to pinkish-gray ostracodal limestone beds, each of which grade downward into fine- to medium-grained light-gray calcareous sandstone beds. The two sandstone and limestone sequences are separated by about 20 ft (6 m) of greenish-gray fissile clay shale. The basal 5 ft (1.5 m) of the clay shale parting characteristically contains a thin zone of brownish stromatolites. The upper sandstone bed rarely exceeds 5 ft (1.5 m) in thickness, but the lower one frequently is as much as 25 ft (8 m) thick; both sequences are locally oolitic and pisolitic. In the southeastern part of the area, the limestones are absent, but the sandstone beds persist and usually make a conspicuous ledge zone. The top of the L2 zone is from 35 ft (10 m) to 100 ft (30 m) below the top of the HA zone. The line on map shows top of zone.

L1—The L1 zone—Commonly consists, in descending order, of sandy pisolitic limestone about 1-3 ft (0.3-0.9 m) thick, oolitic and ostracodal limestone 6 ft (1.8 m) thick, and light-gray fine- to medium-grained calcareous thin-bedded, current ripple-marked sandstone 6-35 ft (1.8-11 m) thick. The sequence makes a prominent cliff throughout the mapped area. In the southeastern part of the area, the L1 zone has an ostracodal limestone bed at the base 3-8 ft (1-2.4 m) thick separated from the L2 limestone zone by alternating beds of fine-grained sandstone, pisolite, stromatolitic layers and locally some greenish-gray fissile shale. The entire L1 zone is as much as 40 ft (12 m) thick. The top of the L1 zone is 155-220 ft (47-67 m) below the top of the L2 zone. The line on map shows top of zone, which was arbitrarily mapped as the top of Tg1.

Kw WASATCH FORMATION (EOCENE AND PALEOCENE)—Mainly sandy claystone and siltstone containing lenses and channels of medium- to coarse-grained cross-bedded sandstone. The sandstone beds are thicker, more numerous, and more conspicuous than they are in the quadrangles adjacent to the west, southwest, and south. The upper part of the Wasatch yielded vertebrate fossils of early to middle Eocene age. These fossils were collected from outcrops in the southwestern part of the area stratigraphically about 750 ft (229 m) below the top of the formation. They were identified by L. C. Gazin of the U.S. National Museum. The lower part of the Wasatch as mapped includes rocks of Paleocene age. Fossil spores and pollen from a 5-ft (1.5-m) thick carbonaceous shale bed about 735 ft (224 m) above the base of the Wasatch were identified by K. L. Techow as being of late Paleocene age. The lower part of the Wasatch in the White Rock quadrangle is of the same age, lithology, and appearance as rocks assigned to the Fort Union Formation in central and southern Wyoming (Venne and Pipirinos, 1954; locs. 8, 9, 10; Pipirinos, 1955, p. 100; 1961, p. A10-A13; Denison and Pipirinos, 1969; Reynolds, 1968; 1971). It was tentatively included in and mapped with the Wasatch pending further investigation into the nature of the Fort Union-Wasatch contact and the feasibility of mapping it. In general, the claystone beds in the Eocene portion are variegated ochre, lavender, red, and green; the Paleocene claystone beds tend to be drab gray and white, and include a few thin rusty-brown to nearly black ferruginous zones, and at least one papery fissile brown carbonaceous shale bed already mentioned. The change from the variegated Eocene to the drab Paleocene part of the Wasatch occurs in an interval that generally is poorly exposed. If there is an erosional break between the two parts of the Wasatch in this area it is inconspicuous and difficult to find. The basal contact of the Wasatch is concealed throughout the quadrangle and whether it is unconformable

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