

Number	Age	Paleontologist	Collector
	Carbonate-quartzi	te, shallow-water as	semblage
	Picacho	Colorado Formation	
1-12- 80C	Wordian	Douglass	No11
1-12- 80D	Leonardian	Douglass	Nd11
JN-1	Leonardian	Douglass	No11
29 428 -PC	late Wolfcampian	Wardlaw, Stamm	Ketner
29 429 -PC	late Wolfcampian	Wardlaw, Stamm	Ketner
29 430 -PC	Morrowan-Atokan	Wardlaw, Stamm	Ketner
29 431- PC	late Morrowan	Wardlaw, Stamm	Ketner
82-FP-15	Chesterian	Wardlaw	F.G. Poole,
1-12- 80A	Probably Mississippian	Sando	No11

Tinajas Group						
2 325	Middle or Late Devonian	Repetski, Harris	Ketner			
2309	late Canadian or early Whiterockian	Repetski	Ketner			
2331	late Canadian or early Whiterockian	Repetskí	Ketner			
2 307	early Canadian	Repetski	Ketner			

Siliceous, deep-water assemblage					
Vuelta Colorada Formation					
82-FP- 12	early middle Leonardian	Wardlaw	F.G. Poole,		
82-FP- 226	early middle Leonardian	Wardlaw	F.G. Poole, Noll		
82-FP- 4	early middle Leonardian	Wardlaw	F.G. Poole, Noll		
55J 28 894- PC	Atokan-Desmoinesian late Morrowan	Murchey Harris	Murchey, Ketner Harris, Wardlaw		

	Leonardian		MOTT
55J	Atokan-Desmoinesian	Murchey	Murchey, Ketner
28 894- PC	late Morrowan	Harris	Harris, Wardlaw
	or Atokan		Ketner
29 426- PC	late Morrowan	Wardlaw, Stamm	Ketner
.9 427 -PC	late Morrowan	Wardlaw, Stamm	Ketner
	Gua	yacan Group	
29 409 -PC	late Osagean	Denkler, Harris	Ketner
29 432 -PC	middle Osagean	Wardlaw, Stamm	Ketner
29 433 -PC	middle Osagean	Wardlaw, Stamm	Ketner
2-FP- 148	Osagean	Wardlaw	F.G. Poole, Noll
8 892- PC	early Osagean	Harris	Harris, Wardlaw, Ketner
28 893- PC	early Osagean	Harris	Harris, Wardlaw, Ketner
28 891- PC	late Kinderhookian	Harris	Harris, Wardlaw, Ketner
32-FP- 149	late Kinderhookian	Wardlaw	F.G. Poole, Noll
32-FP- 230	late Kinderhookian	Wardlaw	F.G. Poole,
29 408 -PC	middle Kinderhookian	Denkler, Harris	Ketner
29 410- PC	Early Mississippian	Denkler, Harris	Ketner
110 54- SD	latest Famennian	Denkler, Harris	Ketner
10 55 -SD	latest Famennian	Denkler, Harris	Ketner
5-12- 79 P	Famennian	Dutro	Nol1
3-9- 80D	Famennian	Dutro	No11
8-9- 80E	Famennian	Dutro	Nol1
10 792 -SD	latest Frasnian to earliest Famennian	Harris	Harris, Wardlaw, Ketner
10 793 -SD	latest Frasnian to earliest Famennian	Harris	Harris, Wardlaw, Ketner
82-FP- 9	late Frasnian	Wardlaw	F. G. Poole.

Riva

Carter

Riva

Riva

No11

No11

11902

12136

FPR-1

JN-2

JN-3

latest

Ordovician

Ordovician

Ordovician

Middle Ordovician

Middle Ordovician

Middle Ordovician

No11

Ketner

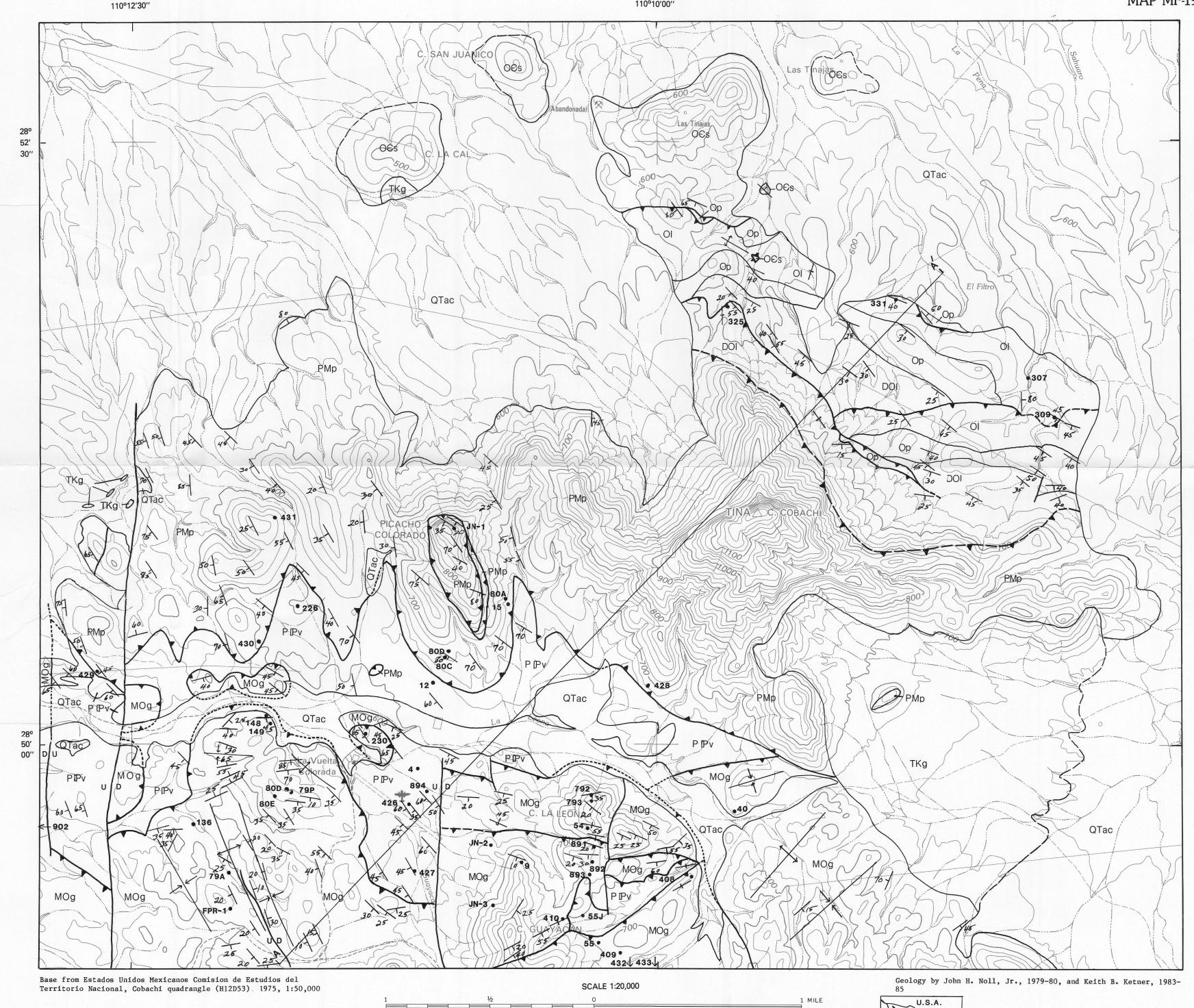
Ketner

No11

No11

No11

F. Peiffer-Rangin (1980)



DATUM IS MEAN SEA LEVEL -1000 METERS METERS 1000-900 -900 800--800 700 QTac -700 PPV DOI Op OI PIPV 500-TKg Op O€s 400--400 O€s -300 TKg 200 ABOVE SEA LEVEL ABOVE SEA LEVEL Contact -- Dashed where position uncertain

CONTOUR INTERVAL 20 METERS

CORRELATION OF MAP UNITS QUATERNARY AND TERTIARY TERTIARY AND CRETACEOUS PERMIAN Lower Permian Upper Pennsylvanian PPv PMp Middle PENNSYLVANIAN Pennsylvanian Pennsylvanian Mississippian - MISSISSIPPIAN Mississippian Upper Devonian DEVONIAN Middle Devonian MOg DO1 Lower Devonian SILURIAN Upper Ordovician Middle Op ORDOVICIAN Ordovician

DESCRIPTION OF MAP UNITS

Ordovician

- CAMBRIAN

QTac Alluvium and colluvium (Quaternary and Tertiary)--Unconsolidated sand and gravel; thick soil and talus

Lower

01

TKg Granitic intrusives (lowest Tertiary or Cretaceous)--Coarse-grained granodiorite and granite. Granodiorite near the western boundary of the map area was dated by the potassium-argon method. Hornblende yielded an age of 65.9 ± 1.4 my and biotite yielded an age of 66.7 ± 1.4 my (Noll, 1981). Granite at Cerro La Cal in the northwestern part of the map area was dated in U.S. Geological Survey laboratories by the fission track method. Zircons yielded a date of 137.3± 15.3 my (R.A. Zimmerman, 1987, written commun.)

CARBONATE-QUARTZITE, SHALLOW-WATER ASSEMBLAGE

Picacho Colorado Formation (Permian to Mississippian)--Thick-bedded limestone containing abundant shelly fauna; cliffy exposures tend to weather pink; some beds quartz-sandy and silty or dolomitic; sporadic chert nodules; altered to marble and calc-silicates near contact with granite. Named and first described by Noll (1981). Thickness probably very great but unknown owing to faults, subparallel to bedding, whose presence is revealed only by reversals in faunal sequence.

Tinajas Group (Devonian to Ordovician) -- Subdivided into three mappable units: Upper limestone (Devonian, Silurian?, and Upper Ordovician?) -- Thick sequence of thinly laminated black limestone that contains lenses of coarse, gray limestone in upper part and grades upward to thickbedded, coarse-grained, gray limestone; dark dolomite locally at base. No fossils have been obtained from the dolomite or the thinly laminated limestone, but a single Middle or Late Devonian conodont was extracted from thick-bedded, coarse-grained part of sequence (J.E. Repetski, oral commun., 1985). Thickness of upper limestone is at least 250 m, but the top, and

possibly base, are faults. Quartzite of Peña Blanca (Middle Ordovician) -- Mature fine- to medium-grained quartzite; upper, major part of unit is massive, pure-white, dense quartzite; basal beds are limy, bioturbated, discolored, and porous; constituent grains well sorted and well rounded; upper contact may be a fault or unconformity. Informally named quartzite of Peña Blanca and correlated with Middle Ordovician Eureka Quartzite by Ketner (1986). Thickness about 132 m.

Lower limestone (Lower and Middle? Ordovician) -- Thickbedded, gray, quartz-silty, coarse-grained limestone; sporadic chert nodules and lenses; sparse, poorly preserved shelly fauna in upper part; probably gradational with overlying quartzite but contact obscure; unit well dated by conodonts; exposed thickness about 200 m, but base not exposed

Skarn (Ordovician? and Cambrian?) -- Limestone strongly metamorphosed to marble and calc-silicates; base not exposed, top faulted

SILICEOUS, DEEP-WATER ASSEMBLAGE

Vuelta Colorada Formation (Permian and Pennsylvanian)--Mainly limy, pyritic siltstone; lesser argillite, bedded chert, nodular and bedded barite; weathers to distinctive red soil, but all units dark gray or black in fresh exposures; graded bedding, sporadic lenses of radiolarians. Named and first described by Noll (1981). Base and top faulted; exposed thickness unknown owing to faults, subparallel to bedding, whose

presence is indicated by reversals in faunal sequence Guayacan Group (Mississippian to Ordovician) -- Subdivided into five units that locally can be mapped separately; unit named and first described by Noll (1981) Limestone, argillite, siltstone, chert (Lower Mississippian) -- Limestone is turbiditic, quartz-sandy, and chert-pebbly; argillite and siltstone are black; chert forms lenses in limestone and occurs as bedded chert sequences. Thickness probably great; top not

exposed in map area; base unconformable on Devonian to Ordovician strata Limestone, bedded chert, sandstone, conglomerate, bedded barite (Upper Devonian) -- Extremely heterogeneous unit that includes some beds of probable deep-water origin (bedded chert, graded sandstone) and some beds that may be of shallow-water origin (grain-supported conglomerate, cross-bedded calcarenite, crinoidal limestone). Thickness at one point 180 m, but top and base are unconformities

Chert (Silurian?) -- Massive white chert and thinly bedded, mainly light-colored chert; locally altered to strongly pigmented, porous gossan; thickness a few meters, but top is an unconformity Chert (Upper Ordovician) -- Black, bedded chert; beds generally more than 30 cm thick; sporadic ovoid nodules as much as 30 cm in diameter; dated by graptolites in

shaly partings; contact with overlying light-colored Silurian(?) chert is gradational; thickness about 25 m Shale (Middle Ordovician) -- Pyritic, black shale; sandstone bed a few meters thick near top; abundant graptolites; upper contact conformable; locally metamorphosed to phyllitic schist. Base not exposed; exposed thickness about 100 m

Anticline Syncline High-angle fault--Dotted where concealed, U on upthrown side; D on downthrown side Thrust fault--Dashed where position uncertain, dotted where concealed; sawteeth on upper plate Strike and dip of beds Inclined Vertical

STRUCTURE

Location and number of fossil collection

Overturned

Eight principal thrust plates were mapped in the Cerro Cobachi area. Plates 1, 2, and 8 comprise a relatively autochthonous carbonate and quartzite assemblage and plates 3--7comprise a relatively allochthonous siliceous assemblage. Each plate is internally complex and consists of folded and faulted sequences, some of which include significant unconformities. Each plate, however, is bounded by major thrust faults and consists of an originally coherent stratigraphic sequence.

One of the most important structures, a thrust fault that juxtaposed the two assemblages of lower Paleozoic rocks, is not exposed. However, it is inferred to be present, concealed beneath plate 8, because units of the siliceous assemblage exposed on the south side of plate 8 are less than 2 km from contemporaneous but contrasting units of the carbonate and quartzite assemblage exposed on the north side of it. Plate 1, the structurally lower

quartzite assemblage, is composed of marble and calc-silicates. Its internal structure is probably complex but remains unknown in the absence of paleontologic dating of beds. Plate 2, which consists of carbonate and quartzite, is sliced by thrust faults and locally folded. The axis of one well-defined fold trends northwest, and nearly all beds and faults strike

northwest and dip southwest. Plate 3, the structurally lowest plate of the siliceous assemblage, is composed entirely of the Permian and Pennsylvanian Vuelta Colorada Formation. Although the bedding consistently strikes northwest and dips southwest as if it were an uncomplicated homoclinal sequence, paleontologically dated beds are out of normal sequence, indicating the probable presence of concealed faults. No evidence of large-scale, tight folding was observed, but in several outcrops small-scale folds, breccia, and

mylonite are visible. Plate 4 consists of units assigned mainly to the siliceous Guayacan Group. Reconnaissance indicates that the Vuelta Colorada Formation may be present also near the southern border of the map area. The rocks of plate 4 are strongly folded about northeasttrending axes. These folds have yet to be defined by detailed mapping and measurements of bedding attitudes, but they are plainly visible on the ground and on aerial photographs. The structure of plate 5, composed of all units of the Guayacan Group, is basically simple, consisting of a series of broad, gentle folds whose axes strike northwest. Locally,

however, this pervasive pattern is interrupted by small, sharp folds, breccia zones, minor thrusts, and high-angle faults. Bedded chert seems especially prone to be locally tightly folded. The major folds in the western part of this plate have wave lengths of about 800 m and amplitudes of about 25 m. In the western part of plate 5, the Devonian beds tend to crop out on hilltops, whereas the Ordovician beds tend to be exposed in the valleys. A striking man-made landscape feature of this plate is the concentration of partially excavated Devonian barite deposits along the ridges.

Plate 6 is a small slice of the Pennsylvanian part of the siliceous Vuelta Colorada Formation. Bounding thrust faults strike northeast and dip southeast. Intense fracturing of this plate prevents determination of bedding attitudes. Plate 7 within the map area, and for a distance of about 2 $\ensuremath{\text{km}}$

2 KILOMETERS

south of the map boundary along Arroyo San Clemente, consists of latest Devonian and Mississippian rocks of the Guayacan Group that strike northeast-southwest and are nearly vertical. Plate 8, composed entirely of the Picacho Colorado Formation, is a homoclinal sequence internally sliced by faults subparallel to bedding. Although two of these are evident from direct observation, others are inferred to be present from the fact that paleontologically dated beds are out of normal sequence. Beds consistently strike northwest and dip southwest. No evidence of

tight folds was observed. In general, the northwest strike and southwest dip of faults and bedding suggest thrust faulting from southwest to northeast. Notable exceptions are plates 4, 6, and 7, which seem to have been thrust from the southeast.

Apparently plates 4, 5, and 8 lie on the same fault surface. Structure contours drawn to show the fault surface between these plates and the underlying plate 3 reveal a surface of low relief that dips very gently to the west. The structural position of plate 8 is anomalous in that it belongs to the carbonate assemblage but lies on essentially the same fault surface as do plates 4 and 5, which belong to the siliceous assemblage. It may be tempting to suggest that plate 8 originated in the Sierra Mazatan core complex, which is $15\ \mathrm{km}$ north of Cerro Cobachi (Anderson and others, 1980, p. 273) and was emplaced during crustal extension. However, this seems unlikely because the structural plates of the Cerro Cobachi area were assembled before intrusion of granitic plutons in Cretaceous or earliest Tertiary time, prior to the onset of middle Tertiary extensional faulting. We therefore conclude that plate 8 was thrust from the southwest and arrived at about the same time as plates 4 and 5 but that the latter came from farther outboard with respect to the craton.

The entire Cerro Cobachi Paleozoic terrane may have been transported along left-slip faults to its present location from localities far to the northwest (Silver and Anderson, 1974; Anderson and Schmidt, 1983; Stewart and others, 1984; Ketner, 1986). Detailed comparisons of the formations of Cerro Cobachi and those of southern California, Arizona, New Mexico, and West Texas are necessary to assess the validity of that hypothesis, and pertinent studies are in progress.

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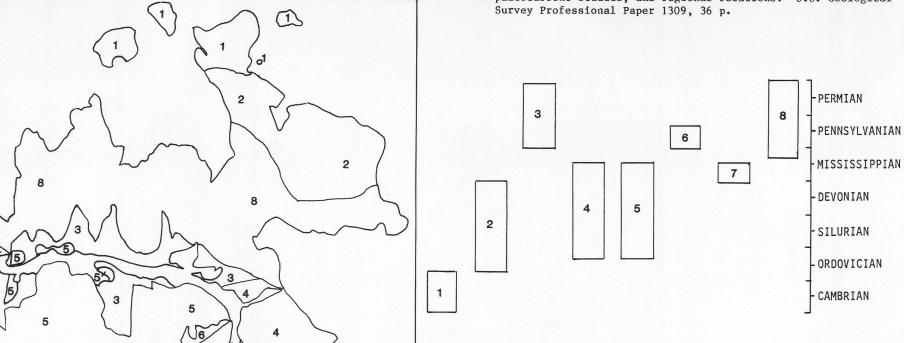
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Locations and stratigraphic ranges of structural plates. Structural positions of the plates range from 1 (lowest) to 8 (highest).

PRELIMINARY GEOLOGIC MAP OF THE CERRO COBACHI AREA, SONORA, MEXICO