

\*Age of epoch boundaries from time scale of Berggren and others (1995). \*\*<sup>40</sup>Ar/<sup>39</sup>Ar ages shown in Ma are from table 2

**Artificial fill (Holocene)**—Unconsolidated soil, sand, gravel, and rock used to construct elevated road and railroad beds and levees on Columbia River floodplain and sandy dredge spoils removed from Columbia River shipping channel.

**Alluvium (Holocene and Pleistocene)**—Unconsolidated, poorly sorted to well-sorted, massive to laminated, commonly cross-stratified sand, silt, and minor gravel of Columbia River floodplain and local accumulations of poorly to moderately sorted, crudely stratified sand and gravel in valley bottoms of tributary streams. Locally include fine-grained lacustrine, aeolian, and organic-rich marsh deposits as well as unimpounded areas of artificial fill (af).

**Landslide deposits (Holocene and Pleistocene)**—Diamictics of bedrock and (or) surficial material transported downslope on mass. Mapped slides predominantly deposited failures ranging from rock or less coherent slumps to internally draped rock-avalanche deposits; small unimpounded debris-flow deposits are widespread in terrain characterized by thick soil horizons. Landslides particularly abundant in areas underlain by poorly consolidated sedimentary rocks of the Pittsburg Bluff Formation (Tpb) and clay-rich tuffaceous rocks (units Tu and Tt). Small unimpounded slides are also common along steep slopes of terraces composed of poorly consolidated, fine-grained flood deposits (unit Qa) or beds of the Sandy River Mudstone and Troutdale Formation (units Tr and Tf).

**Catastrophic flood deposits (Pleistocene)**—Sediments deposited by colossal glacial outburst floods or shallower inundations by repeated failures of ice dam at Pleistocene Lake Missoua in Montana, some of which released as much as 2184 km<sup>3</sup> of floodwater (Breitz, 1925, 1969; Allison, 1978; Baker and Baker, 1982; Waitt, 1985, 1994, 1996; Allen and others, 1986; O'Connor and Baker, 1992). Many floods, hydraulically dammed by the relatively narrow constriction of Columbia River valley in northern part of quadrangle, temporarily ponded in Portland Basin, backed up into tributary valleys, and deposited suspended sediment load (Trimbé, 1963). Radiocarbon ages and tephrochronologic data indicate most recent floods occurred between about 15.5 and 12.7 ka (Waitt, 1994) although there is evidence for older catastrophic flooding events (Patton and Baker, 1978; McDonald and Buacac, 1988; Zaffa and others, 2000; Björnsdottir and others, 2001). In this quadrangle divided into:

**Gravel facies**—Unconsolidated deposits of massive, poorly sorted, pebbly to cobbly gravel. Forms 1.5-m-thick veneer on terrace surface between Columbia City and Deer Island and smaller deposit north of Martin Island. Diverse and variable clast populations; well-sorted to subangular clasts, generally 530 cm diameter, include Columbia River Basalt Group, Tertiary and Quaternary volcanic rocks from the Cascade Range, pre-Tertiary granitic and metamorphic rocks, and quartzite. South of Columbia City, deposit rests on Grande Ronde Basalt and consists of subangular to angular basalt clasts with little matrix, presumably eroded from bedrock outcrop at Santa Helens immediately east of atly upstream. Deposit near Martin Bluff contains abundant rounded clasts probably reworked from older sedimentary units (Tr, Tf and Tc). Unconformably overlies bedded silt and sand along Smith Road west of Columbia City, suggesting deposition from a late flood on older eroded slackwater deposits.

**Silt and sand facies**—Unconsolidated deposits of rhythmically bedded silt and fine sand that form terraces with surface elevations of 200 to 250 ft (60 to 75 m) between Tide Creek and Columbia City. Locally exceeds 30 m in thickness. Bluish gray in fresh exposures but rapidly oxidizes to light brown. Composed largely of quartz, feldspar, and coarsening muscovite, indicating deposition by Columbia River rather than by local sources. Thin unimpounded deposits of similar character locally mantle topography to elevations as high as 400 ft (120 m). Interpreted as slackwater sediment that settled out of temporarily ponded floodwaters. Although silt and sand deposits near Columbia City are older than gravel facies (unit Qa), those found elsewhere may be in part coeval with gravel, reflecting different hydraulic environments in a single flood.

**Mount St. Helens-derived alluvium (Pleistocene)**—Unconsolidated, moderately to poorly sorted, crudely bedded, coarse sand and sandy, clay-sappressed, pebbly and cobbly gravel that form small dissected terrace remnant about 1.5 km northwest of community of Deer Island, Oregon. As much as 20 m thick. Deposit dominated by Tertiary volcanic rocks but larger clasts include distinctive weathered, white to light-gray, coarsely porphyritic, quartz- and biotite-phyric dacite of the local produced during the earliest eruptions at Mount St. Helens volcano (Candell, 1987; Mullins, 1990); grains of plagioclase, quartz, and biotite, similar to phenocrysts in the dacite, dominate the sandy facies. Crossbedding indicates deposition from westerly-flowing currents. Terrace inferred to be distal remnant of large but ephemeral alluvial fan constructed at mouth of Lewis River following major eruptions during the Ape Canyon eruptive stage of Mount St. Helens, ca. 507 to 36 ka (Crandell, 1987; Scott, 1989) or possibly earlier (Hooper and Buacac, 1995; Whitlock and others, 2000; M.A. Clyne, oral commun., 1999).

**Old outwash deposits (Pleistocene)**—Unconsolidated, moderately well-sorted, moderately sorted to well-sorted silt and sand that unconformably overlies Neogene sedimentary units (Tr and Tf) on slope about 3 km north of Woodland. Contain well-rounded clasts of diverse volcanic rocks similar to Tertiary rocks of the Cascade Range but lacks quartzite and Columbia River Basalt Group clasts, indicating a Lewis River rather than Columbia River source. Clasts attributable to Mount St. Helens volcano or to Pleistocene basaltic centers in Lewis valley also absent. About 40 m thick. Interpreted as glacialofluvial deposits related to one or more periods of early or middle Pleistocene glaciation in Lewis River drainage (Mandoff, 1964; Crandell, 1987; R.C. Evans, unpub. mappings).

**Conglomerate (Pleistocene?) and Pliocene?**—Semi-consolidated, thick-bedded, poorly to moderately well sorted, clay-sappressed, commonly crossbedded, open-work and sand-matrix, pebbly to cobbly conglomerate that underlies virtually uninducted terrace that slopes uniformly from 90 ft (27 m) at Columbia City to 60 ft (18 m) at Deer Island. Less intensely weathered than typical Troutdale Formation conglomerate. Well-sorted to subrounded clasts are chiefly Tertiary volcanic rocks from the Cascade Range and Columbia Plateau; scarce pebbles of coarse-grained granitic rocks, quartzite, and siliceous siltstone. Matrix is typically weakly to moderately cemented by clay and ironite. Pronounced clay imbrication and foreset bedding indicate deposition from north-flowing currents. Deposit interpreted as erosional remnant preserved in lee of basalt outcrop at its south end.

**Troutdale Formation (Pliocene and Miocene?)**—Semi-consolidated, massive, poorly sorted to well-sorted pebbly and cobbly conglomerate and local subordinate friable sandstone. On Washington side of Columbia River, crop out in area about 2 to 4 km north of Woodland where as much as 80 m of conglomerate gradually overlies sand and silt of Sandy River Mudstone. In Oregon, unit is mapped between Merrill and Tide Creeks basally largely on water-well logs. Troutdale Formation underlies large area to the east and southeast of this quadrangle (Mandoff, 1964; Swanson and others, 1985; Brown, 2002; R.C. Evans, unpub. mappings). Clast population dominated by well-sorted cobbles of the Columbia River Basalt Group; typically includes some light-colored granitic and quartzofeldspathic metamorphic rocks and conspicuous quartzite clasts stained yellow to brown by iron oxides; weakly to moderately cemented by clay and limonite produced by weathering of basaltic clasts. Sandstone interbeds and matrix of conglomerate generally composed of micaceous arenaceous compositionally similar to Sandy River Mudstone. These lithologies indicate primary source areas east of the Cascade Range and deposition from ancestral Columbia River. Upper several meters of unit thoroughly decomposed to red, structureless, clayey silt; in such areas, presence of scattered quartzite pebbles on surface is only evidence of parent material. Higher mapped exposures found at elevation of about 550 ft (170 m) but original top of unit was apparently much higher, as indicated by scattered quartzite pebbles in soil at elevation as high as 1200 ft (365 m).

**Sandy River Mudstone (Pliocene? and Miocene?)**—Variable but generally fine-grained and poorly indurated sedimentary rocks including thin-bedded claystone, mudstone, and siliceous, laminar, lenticular, and cross-bedded sandstone and siltstone, and lenses of pebbly sandstone and pebble to small-cobble basalt-clast conglomerate, commonly limonite cemented. Underlies extensive upland surface as high as 500 ft (150 m) west of Columbia River, where unit is as much as 100 m thick. Distribution east of river patchy; a sandstone-dominated section as much as 120 m thick overlies Pliocene bedrock adjacent to Interstate Highway 5 about 2 to 4 km north of Woodland and grades upward into conglomerate of Troutdale Formation. Typical sandstone is light to medium gray and moderately friable where fresh, weathers to orange-brown color. Locally limonite-cemented; composed of subangular grains of quartz, plagioclase, orthoclase, a variety of volcanic and metamorphic rock fragments, muscovite, and minor biotite, ironmineral, magnetite, hornblende, and pyroxene. Lenticular conglomerates lithologically similar to Troutdale Formation. Features such as crossbedding and cut and fill structures indicate fluvial depositional environment. Variegated, clayey, buffaceous beds with locally abundant carbonaceous woody fragments unconformably overlie member of Sentinel Bluffs of the Grande Ronde Basalt (unit Tpb) southeast of Columbia City; observed only in quarry wall exposures immediately south of this quadrangle, hence distribution of unit in this area is uncertain; at least 20 m and possibly as much as 50 m thick. Lithology and stratigraphic position of these deposits suggest correlation with Sandy River Mudstone of Trimbé (1963), lower member of Troutdale Formation of Mandoff (1964), and Wilkes Formation of Roberts (1958); the strata outlying Grande Ronde Basalt near Columbia City may be equivalents of the Vanage interbeds (Brown and others, 1985; Brown, 1989; Brown and Tolan, 1990). Plant fossils from locally along top runs of Interstate Highway 5 near east edge of quadrangle are latest Miocene to early Pliocene in age (E.L. Sanborn, cited in Wilkinson and others, 1946).

**Basaltic-cobble conglomerate (Miocene?)**—Poorly indurated, poorly weathered, rounded, pebbles and cobbles of Grande Ronde basalt, accompanied by small amounts of class similar to Eocene volcanic rocks of the area. Lensoid body approximately 40 m thick that occupies block-inclined inner margin sandstone (unit Tpb) and directly underlies low-to-intermediate-MgO Grande Ronde Basalt flow (unit Tgr) on northeast flank of Maple Hill. Correlative with the Scappoose Formation of Van Ant and Kelly (1985).

**Pittsburg Bluff Formation (Oligocene)**—Poorly exposed and deeply weathered, massive, friable, well-bedded, fine-to coarse-grained, tuffaceous, arkosic, and lithic sandstone, siltstone, mudstone, tuff, and rare granitic and pebble conglomerate. Crops out only west of Columbia River. Approximately 200 m thick in this quadrangle but both base and top of unit are erosional unconformities. Thickness increases to the west (Wilkinson and others, 1946). Depositional environments predominantly shallow marine but unit contains rare lithic lenses (Vance and Geer, 1940; Wilkinson and others, 1946) indicating local erosion of alluvial fans. Unit contains abundant angular to subangular grains of porphyritic to aphyric, mafic to intermediate volcanic rocks, pumice, micromorphic rocks, monocrystalline quartz, plagioclase, and alkali feldspar, muscovite, biotite, green and brown hornblende, pyroxene, epidote, staurolite, zircon, and garnet. Exposures 3 to 5 km west of quadrangle contain a mollican fauna similar to that of the Pittsburg Bluff Formation (Wilkinson and others, 1946), which is of latest Eocene to early Oligocene age; poorly preserved molds and casts of gastropods and polychaetes collected from roadcut about 1.5 km southeast of Maple Hill are consistent with an Oligocene age and probably reflect deposition in shallow continental shelf setting (E.J. Moore, written commun., 1999). As mapped, upper part may include beds correlative with the late Oligocene to middle Miocene Scappoose Formation (Warren and Norbrink, 1946; Van Ant and Kelly, 1985).

**Wanupum Basalt**

**Freeman Springs Member**

**Basalt of Sand Hollow**—Small erosional remnant of dark gray, microvesicular, nearly aphyric to microporphic, spherulitically weathered basalt exposed on western flank of Maple Hill. Contains widely scattered, plagioclase phenocrysts and glomerocrysts as large as 1 cm and plagioclase and rare olivine microphenocrysts about 1 mm across in interstitial groundmass of plagioclase laths and angular olivine granules in abundant vitally opaque glass with skeletal Fe-Ti oxides; microvesicles partly filled with calcite or green, yellow, or orange clay. Distinguished from Grande Ronde Basalt flows by a higher TiO<sub>2</sub> content of about 3.0 wt percent (table 1). Chemistry, petrography, and magnetic polarity (J.T. Hagstrum, written commun., 2001) indicates correlation with informal basalt of Sand Hollow on Columbia Plateau (Brown and others, 1985; 1989; Hooper, 2000) erupted about 15.5 Ma (Reidel and others, 1989). Position topographically below Grande Ronde Basalt flows suggests this flow eroded after canyon eroded into the older flows.

**Grande Ronde Basalt**—Light- to dark-gray, vesicular to microvesicular, aphyric to microporphic to sparsely plagioclase-phyric lava flows of tholeiitic basalt and basaltic andesite. Unaltered nature contrast with pervasive zeolite-facies metamorphism of Paloque rocks. Where best visible in quarry-wall exposures, flows typically exhibit well-developed columnar and entablature jointing patterns and vesicular flow tops. Upper few meters commonly thoroughly weathered to red saprolite. In this quadrangle, consists of:

**Basaltic andesite**—Flows and flow breccia of dark gray to brown, porphyritic to seriate to aphyric basaltic andesite. Flows generally 3 to 6 m thick, with reddish, vesicular, zoolitoid top; flow interiors blocky jointed to platy. Flows locally separated by thin (ca. 0.5 m thick), massive, red-brown palisades developed on fine-grained tuffaceous interbeds. Typical flows contain phenocrysts of plagioclase (10-35 percent; 1 to 2 mm long) and olivine (as much as 4 percent; 0.5 to 2 mm across; commonly partly resorbed and rimmed by fine-grained granitic pyroxene or magnetite; invariably replaced by smectite a hematite, carbonate, serpentine, quartz; commonly contains minute inclusions of chromian spinel); many flows also contain phenocrysts of argillite (as much as 2 percent; 0.5 to 1 mm across) and a few contain hypersthene (0.5 to 1 mm long); intergranular to trachytic

**Basalt flows in Deer Island quadrangle** approximately 200 m near Columbia City, thins northward. K-Ar and <sup>40</sup>Ar/<sup>39</sup>Ar age determinations from the Columbia Plateau indicate that Grande Ronde Basalt flows were emplaced between 16.5 to 15.5 Ma (Reidel and others, 1989). Based on field, chemical, and paleomagnetic properties, most Grande Ronde Basalt flows in this quadrangle can be assigned to informal members corresponding to those defined by Reidel and others (1989) on the Columbia Plateau (terminology follows that of Reidel, 1989).

**Member of Sentinel Bluffs**—Flows of flows characterized by relatively high MgO content (4.5 to 4.7 wt percent) and normal magnetic polarity. Underlies member of Sentinel Bluffs (unit Tgr) approximately 20 to 40 m thick. Chemical and paleomagnetic characteristics indicate primary correlation with Water unit of Reidel and others (1979; Brown and others, 1989; Wells and others, 1989), which is within the N<sub>2</sub> magnetotatigraphic unit of Swanson and others (1979) and was erupted about 15.6 Ma.

**Member of Winter Water**—Sparsely plagioclase-phyric to glomerophytic basalt with relatively low MgO content (3.1 to 3.5 wt percent) and normal magnetic polarity; plagioclase phenocrysts 1 to 3 mm across; groundmass aphyric to sparsely microporphic. Underlies member of Sentinel Bluffs (unit Tgr) approximately 20 to 40 m thick. Chemical and paleomagnetic characteristics indicate primary correlation with Water unit of Reidel and others (1979; Brown and others, 1989; Wells and others, 1989), which is within the N<sub>2</sub> magnetotatigraphic unit of Swanson and others (1979).

**Member of Orley**—Aphyric, interstitial to intergranular basalt with relatively low MgO content (3.4 to 3.6 wt percent) and normal magnetic polarity. Underlies member of Sentinel Bluffs (unit Tgr) approximately 20 to 60 m thick. Chemical and paleomagnetic characteristics indicate primary correlation with Orley unit of Reidel and others (1989; Brown and others, 1989; Wells and others, 1989), which is within the N<sub>2</sub> magnetotatigraphic unit of Swanson and others (1979).

**Unassigned, reversely-magnetized, high-MgO basalt**—Two or more flows with generally intermediate MgO content (4.1 to 4.2 wt percent) and reversed magnetic polarity (J.T. Hagstrum, written commun., 1999, 2002). Forms bulk of Sentinel Hill locally contains rare plagioclase phenocrysts as large as 5 mm across in a microphyric, interstitial to intergranular groundmass. Relatively high CaO and Cr<sub>2</sub>O<sub>3</sub> contents indicate affinity with high-MgO Grande Ronde basalts (M.T. Brown and T.L. Tolan, oral commun., 2001). Does not correspond to any of the informal members described by Reidel and others (1989).

**Basaltic-cobble conglomerate (Miocene?)**—Poorly indurated, poorly weathered, rounded, pebbles and cobbles of Grande Ronde basalt, accompanied by small amounts of class similar to Eocene volcanic rocks of the area. Lensoid body approximately 40 m thick that occupies block-inclined inner margin sandstone (unit Tpb) and directly underlies low-to-intermediate-MgO Grande Ronde Basalt flow (unit Tgr) on northeast flank of Maple Hill. Correlative with the Scappoose Formation of Van Ant and Kelly (1985).

**Pittsburg Bluff Formation (Oligocene)**—Poorly exposed and deeply weathered, massive, friable, well-bedded, fine-to coarse-grained, tuffaceous, arkosic, and lithic sandstone, siltstone, mudstone, tuff, and rare granitic and pebble conglomerate. Crops out only west of Columbia River. Approximately 200 m thick in this quadrangle but both base and top of unit are erosional unconformities. Thickness increases to the west (Wilkinson and others, 1946). Depositional environments predominantly shallow marine but unit contains rare lithic lenses (Vance and Geer, 1940; Wilkinson and others, 1946) indicating local erosion of alluvial fans. Unit contains abundant angular to subangular grains of porphyritic to aphyric, mafic to intermediate volcanic rocks, pumice, micromorphic rocks, monocrystalline quartz, plagioclase, and alkali feldspar, muscovite, biotite, green and brown hornblende, pyroxene, epidote, staurolite, zircon, and garnet. Exposures 3 to 5 km west of quadrangle contain a mollican fauna similar to that of the Pittsburg Bluff Formation (Wilkinson and others, 1946), which is of latest Eocene to early Oligocene age; poorly preserved molds and casts of gastropods and polychaetes collected from roadcut about 1.5 km southeast of Maple Hill are consistent with an Oligocene age and probably reflect deposition in shallow continental shelf setting (E.J. Moore, written commun., 1999). As mapped, upper part may include beds correlative with the late Oligocene to middle Miocene Scappoose Formation (Warren and Norbrink, 1946; Van Ant and Kelly, 1985).

**Goble Volcanic (Eocene)**—Dominantly mafic lava flows and interbedded volcanoclastic beds west of Columbia River. In this quadrangle consists of:

**Tuff**—Heterogeneous unit composed of andesitic to rhyolitic tuff, pumiceous lapilli tuff, pumiceous and lithic tuff breccia; very poorly sorted, matrix-supported, relatively coarse-grained deposits predominate; inferred to be mostly of pyroclastic-flow and lahars origin. Exposed in roadcuts along U.S. Route 30 north of Tide Creek. Generally light colored, ranging from white to light gray to green; weathers buff to yellowish brown. Most tuff beds exhibit minimal compaction, and some contain unimpounded pumice blocks as large as 30 cm across. Angular lithic blocks in lithic breccia as large as 1 m across. Carbonized woody debris present in some beds, especially near their base. Phenocrysts rarely constitute more than 10 percent of juvenile material in tuff and include plagioclase, argillite, hypersthene, and Fe-Ti oxide but no quartz, hornblende, or biotite. Original glass is completely replaced by fine-grained smectite and (or) zeolites, of which clinoptilolite is the most common. <sup>40</sup>Ar/<sup>39</sup>Ar age of plagioclase separated from a crystal-rich lapilli tuff in northwestern part of quadrangle is 36.1±0.3 Ma (R.J. Fleck, written commun., 1999).

**Basaltic andesite**—Flows and flow breccia of dark gray to brown, porphyritic to seriate to aphyric basaltic andesite. Flows generally 3 to 6 m thick, with reddish, vesicular, zoolitoid top; flow interiors blocky jointed to platy. Flows locally separated by thin (ca. 0.5 m thick), massive, red-brown palisades developed on fine-grained tuffaceous interbeds. Typical flows contain phenocrysts of plagioclase (10-35 percent; 1 to 2 mm long) and olivine (as much as 4 percent; 0.5 to 2 mm across; commonly partly resorbed and rimmed by fine-grained granitic pyroxene or magnetite; invariably replaced by smectite a hematite, carbonate, serpentine, quartz; commonly contains minute inclusions of chromian spinel); many flows also contain phenocrysts of argillite (as much as 2 percent; 0.5 to 1 mm across) and a few contain hypersthene (0.5 to 1 mm long); intergranular to trachytic

groundmass of the same minerals plus Fe-Ti oxide and minor interstitial glass (mostly altered to smectitic clay or quartz or calcite).

**Dacite breccia (Eocene)**—Poorly sorted monolithologic breccia composed of angular clasts as large as 1 m across of black to red, locally flow-banded, dacite in matrix of cemented dacite. Clasts contain sparse small (0.5 mm across) phenocrysts of plagioclase and argillite in a felsic cryocrystalline groundmass. Forms small ridge east of Interstate Highway 5 northeast of Martin Island. Most likely a block-and-ash deposit generated by collapse of a dacite dome.

**Volcanoclastic sedimentary rocks (Eocene)**—Diverse assemblage of continental volcanoclastic rocks of inferred epiclastic origin. Consists of generally well bedded, well-sorted to poorly sorted siltstone, sandstone, conglomerate, and breccia, all composed of volcanic debris. Typically light green to olive green or greenish gray but also white, tan, brown, or maroon. Texturally and compositionally immature; most abundant clasts are volcanic rocks petrographically similar to interbedded lava flows and tuffs; other components include plagioclase, Fe-Ti oxide, and pyroxene crystals, pumice, vitric ash, and carbonized plant remains. Interpreted as debris eroded from older and peritotomogeneous volcanoes and deposited in streams and lakes in medial to distal volcanic settings. Indurated, but easily eroded to dust to pervasive replacement by soft clay minerals and zeolites, hence poorly exposed.

**Tuff (Eocene)**—Andesitic to aphyritic tuff, pumiceous lapilli tuff, and pumiceous and lithic tuff breccia lithologically similar to tuff of Goble Volcanic (unit Tgr). Well-exposed in cliffs along east bank of Columbia River near Martin Bluff, in nearby railroad cuts, and along Interstate Highway 5 northeast of Martin Island.

**Basaltic andesite breccia (Eocene)**—Massive, unsorted, monolithologic breccia composed of angular clasts as large as 1 m of dark-gray to reddish-gray, aphanitic to seriate, sparsely vesicular basaltic andesite in yellowish-gray matrix of cemented basaltic andesite. Predominantly clast supported but lower part of unit shows inverse grading and higher proportion of matrix. At least 50 m thick; base not exposed; top locally weathered to limonite-stained and bleached spherulite and unconformably overlain by Sandy River Mudstone. Clasts consist of plagioclase, argillite, and altered olivine 3-1 mm across in an intergranular to microlitic groundmass of plagioclase, pyroxene, and Fe-Ti oxides poikilocrystally replaced by equant olivine (ca. 5 percent) of hypopyroxene. Textural characteristics suggest unit is a debris-avalanche deposit like one formed in 1980 at Mount St. Helens. Wilkinson and others (1946) considered breccia to be interbedded with Troutdale Formation sandstone, but an <sup>40</sup>Ar/<sup>39</sup>Ar age of 37.1±0.3 Ma (R.J. Fleck, written commun., 1999) demonstrates it is a late Eocene bedrock unit.

**Basaltic andesite (Eocene)**—Flows and flow breccia of dark gray to brown, porphyritic to seriate to aphyric basaltic andesite lithologically similar to flows of Goble Volcanic (unit Tgr); unit locally includes minor andesite flows and volcanoclastic rocks too small or poorly exposed to map.

**Xenolithic andesite (Eocene)**—Distinctive flow of platy, medium- to dark-gray, commonly microvesicular, sparsely porphyritic andesite containing scattered mafic xenoliths. Contains phenocrysts of plagioclase 1 to 2 percent, 0.1-1 mm long; argillite (ca. 5 percent; 0.5 mm across) and olivine (ca. 0.5 percent; 0.5-1 mm across; altered to smectite a hematite, with minute inclusions of chromian spinel) in a dark, Fe-Ti oxide-rich, commonly streaky, microgranular to cryocrystalline groundmass. Subsequent angular to oval xenoliths, 5-1 cm across, consist of coarse-grained, moderately recrystallized, porphyritic basaltic andesite consistent with about a 3 percent of flow.

**Contact**—Dashed where approximately located; short-dashed where inferred; dotted where concealed.

**Fault**—Dashed where inferred; dots where concealed. Ball and bar on downthrown side; arrows show relative horizontal movement.

**Basaltic andesite dike**

**Strike and dip of beds**

**Inclined**

**Horizontal**

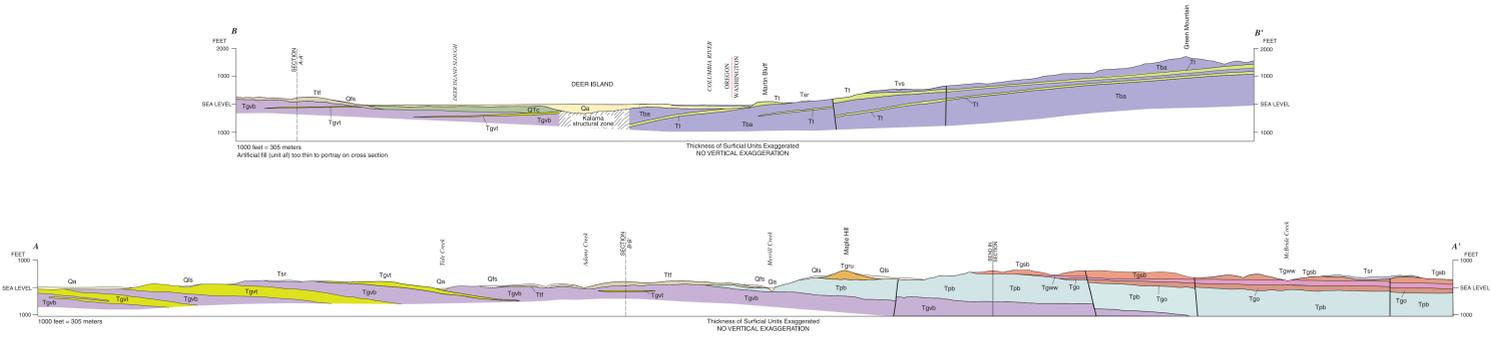
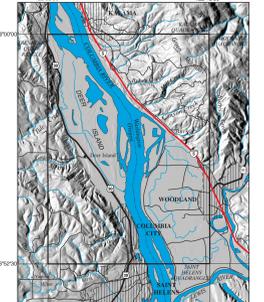
**Strike and dip of compaction foliation in pumiceous lapilli tuff**

**Strike and dip of platy parting in lava flows**

**Sample locality for chemical analysis—See table 1**

**Sample locality for radiometric age determination and age in Ma (1-σ error)**

**Megalof locality**



**GEOLOGIC MAP OF THE DEER ISLAND QUADRANGLE, COLUMBIA COUNTY, OREGON AND COWLITZ COUNTY, WASHINGTON**  
By Russell C. Everts  
2002

This map was prepared for an electronic poster display for the 2002 Geological Society of America meeting. It is not a final product and should not be used for any other purpose. The map is available in electronic form on the USGS website. For more information, contact the USGS at the address below.