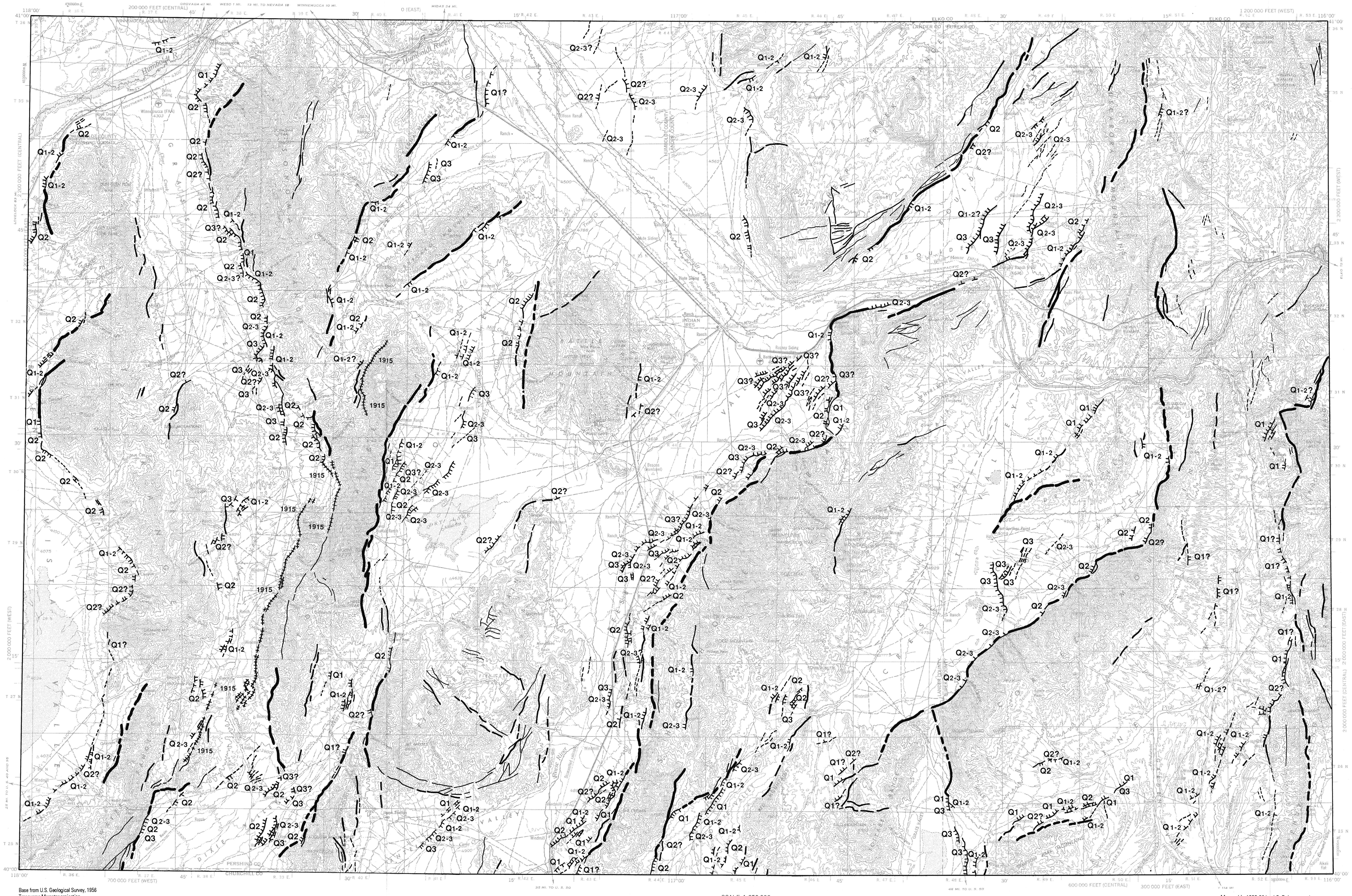
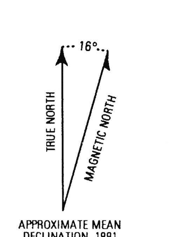


Base from U.S. Geological Survey, 1956
Transverse Mercator projection

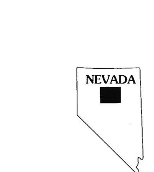


SCALE 1:250 000

5 0 5 10 15 20 25 MILES

5 0 5 10 15 20 25 KILOMETERS

CONTOUR INTERVAL 200 FEET
WITH SUPPLEMENTARY CONTOURS AT 100 FOOT INTERVALS
NATIONAL GEODETIC VERTICAL DATUM OF 1929



Mapped in 1988-89 by J.C. Dohrenwend
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INTRODUCTION

The State of Nevada occupies most of the central and western parts of the Great Basin, the largest tectonically active region within the Basin and Range geomorphic province of North America. The topography of this region is typified by generally north-northwest- to northeast-trending, subparallel mountain ranges separated by alluvial basins of similar plan form and orientation. This classic Basin and Range physiography is the product of at least two phases of middle to late Cenozoic extensional faulting (Zoback and others, 1981; Eaton, 1982; Stewart, 1983), and most of the basins and ranges of the region are at least partly bounded by late Cenozoic faults. The earlier phase of extension was marked by widespread shallow detachment faulting that began at approximately 35 Ma and locally continued into the time of the later phase (Eaton, 1982; Stewart, 1983). The later phase, which was dominated by high-angle, more deeply penetrating block faulting, may have begun locally at about 17 Ma and continues episodically to the present time (Christiansen and McKee, 1978; Eaton, 1982). This map, one of a series of 1° by 2° quadrangle maps showing young faults in Nevada, provides a generalized picture of the late Tertiary and Quaternary faulting that is associated with the latter extensional phase. These young faults are a primary determinant of the present configuration of ranges and basins within the quadrangle.

MAPPING PROCEDURE

Young faults are herein defined as those faults that have undergone latest Tertiary and (or) Quaternary movement. These faults are commonly marked by a variety of diagnostic constructional landforms and other surficial phenomena that can be readily identified and mapped on aerial photographs. These features include (1) scarps on latest Tertiary and (or) Quaternary surficial deposits, volcanic strata, or geomorphic surfaces (either erosional or depositional); (2) prominent alignments of linear drainageways, ridges and swales, active springs and (or) spring deposits, and linear discontinuities of structure, rock type, and vegetation; and (3) abrupt, steeply sloping range fronts with basal scarps, faceted spurs, 'wineglass valleys,' and elongate drainage basins with narrow valley floors (Thornbury, 1969; Bull, 1977; Bull and McFadden, 1977; Wallace, 1977, 1978).

National High Altitude Program (NHAP), 1:58,000-nominal-scale, color-infrared photographs were used for photogeologic interpretation. This mapping was transferred directly to 1/2° by 1° topographic quadrangle maps that were enlarged to the scale of the photographs. These maps were then reduced and compiled at 1:250,000 scale. This compilation was then compared with previous mapping of young faults within the quadrangle (Wallace, 1979) and significant differences between maps were resolved.

Following comparison and resolution with previous mapping, the final 1:250,000-scale compilation was digitized using a GTCO digitizing board connected to a Macintosh II minicomputer. The resulting vector file was converted to raster format (cell size = 200 m by 200 m) and analyzed to determine the approximate length and average orientation of each fault segment. These data are summarized in table 1 and figure 1.

General ages of surficial deposits and erosion surfaces cut by young faults were estimated using a variety of photogeologic and geomorphic criteria (table 2). These age estimates provide a general indication of the approximate timing of young faulting throughout the quadrangle. However, it should be emphasized that these data do not necessarily reflect the age of most recent surface rupture along any particular fault segment. Rather, they provide only very general (and commonly somewhat biased) age constraints on this surface faulting. Age estimates based on photogeologic analysis of surficial deposits and erosion surfaces are, at best, both tentative and imprecise. Moreover, the distribution of these deposits and surfaces is inherently biased by geomorphic process and environment. For example, in those areas of the Great Basin where range uplift rates are low to moderate, remnants of older geomorphic surfaces tend to be concentrated in proximal piedmont areas, whereas younger surficial deposits tend to accumulate on distal piedmonts and basin flats. Consequently, young faults located in intrabasin areas are more likely to offset younger surface deposits then are faults located along range fronts or in proximal piedmont areas. Therefore, inferences based on these data regarding the temporal distribution of young fault activity should be used with caution.

In addition to the limitations imposed by map and photo scales, one other factor also significantly constrains the resolution of the present map. The photography used in this analysis, which was acquired under high sun-angle conditions, is not well suited for the discrimination and mapping of subtle topographic features. Consequently, reexamination of any of the fault systems shown on this map, using larger scale and (or) lower sun-angle aerial photography, would very likely reveal a substantial number of additional young fault segments.

PATTERNS OF LATEST TERTIARY AND QUATERNARY FAULTING

Several factors significantly influence the preservation of fault-related landforms and, therefore, the apparent distribution of young faults as indicated by the distribution of these landforms can be significantly biased. These factors include (1) composition, induration, and structural integrity of the rock or sediment type(s) underlying fault scarps; (2) local geomorphic environment of the scarp or other fault related landform; (3) regional climatic conditions and paleoclimatic variations; and (4) magnitude and recurrence of fault movement (Wallace, 1977; Bucknam and Anderson, 1979; Nash, 1980, 1984; Hanks and others, 1984; Mayer, 1984; Pierce and Colman, 1986; Machette, 1986, 1988, 1989). Therefore, the distribution of young faults shown on this map provides, at best, only an approximate and somewhat biased picture of late Tertiary and Quaternary faulting within the quadrangle. Specifically, faults having a long history of recurrent movement, juxtaposing bedrock and alluvium, or cutting upper Cenozoic lava flows and (or) welded ash-flow tuffs tend to be overrepresented, whereas faults of pre-late Pleistocene age cutting unconsolidated surficial deposits and having either short histories of recurrent movement or long recurrence intervals tend to be underrepresented. Scarps developed on volcanic rocks may be preserved for periods of as much as 10 m.y. By comparison, scarps on the fluvially active parts of piedmont surfaces would likely be completely destroyed within a few thousand years at most, and even on inactive piedmont surfaces, fault scarps on unconsolidated alluvial fill are significantly rounded within 10,000 years (Wallace, 1977), and low scarps would be sufficiently degraded to be unrecognizable on standard aerial photography within a few hundred thousand years (Wallace, 1977; Hanks and others, 1984; Machette, 1989).

The Winnemucca quadrangle lies across the north end of the central Nevada seismic zone, one of the most tectonically active areas in the Great Basin (Wallace, 1977; Thenhaus and Barnhard, 1989). This zone trends generally northward into the southwestern part of the quadrangle where historic fault scarps, formed during the 1915 Pleasant Valley earthquake, score the west flanks of the Tobin Range and the Sou Hills. Young fault densities are uniformly high across most of the quadrangle, ranging between 0.11 and 0.12 km/km^2 (table 1 and fig. 1). Widespread range-bounding fault systems, displaying evidence of latest Pleistocene (10-30 ka) and (or) Holocene (0-10 ka) movement, bound the west or northwest flanks of essentially all of the major ranges within the quadrangle. Intrabasinal swarms of young faults, most notably within Dixie, Grass, Dry Lake, Antelope, Reese River, Boulder, and Crescent Valleys, also show evidence of latest Pleistocene and (or) Holocene activity. With the exception of the area northeast of the Humboldt River where the incidence of late Quaternary faulting is somewhat less than elsewhere in the quadrangle, no regional trends of changes in young fault density or orientation are apparent (table 1 and fig.1).

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Historic surface faulting--Surface ruptures associated with historic earthquakes. Number indicates year of earthquake:
1915 Pleasant Valley earthquakes, M = 7.6 ± 0.2 (Slemmons, 1957; Wallace, 1979)

Fault scarps on Quaternary surficial deposits or Quaternary erosional surfaces--Scarps on Quaternary deposits or Quaternary erosion surfaces other than those displacements associated with known historic earthquakes. Hachures indicate downslope direction of scarp. Symbol indicates approximate age of offset surficial deposit or erosion surface:
Q3 Holocene (0-10 ka)
Q2-3 Latest Pleistocene and (or) Holocene (0-30 ka)
Q2 Late Pleistocene (10-130 ka)
Q1-2 Early to middle and (or) late Pleistocene (0.01-1.5 Ma)
Q1 Early to middle Pleistocene (0.13-1.5 Ma)
Photogeologic and geomorphic criteria used to estimate ages are summarized in table 2

Fault-related lineaments on Quaternary surficial deposits or on Quaternary erosion surfaces--Alignments, in Quaternary surficial deposits or across Quaternary erosional surfaces, of one or more of the following features: linear reaches of stream channels, linear stream valleys, shallow linear swales, springs, vegetation discontinuities. Commonly associated with fault scarps on Quaternary deposits and (or) erosional surfaces. Age designations as above

Major range-front faults--Faults bounding tectonically active fronts of major mountain ranges. These range fronts are characterized by: fault juxtaposition of Quaternary alluvium against bedrock, fault scarps and lineaments on surficial deposits along or immediately adjacent to range front, a general absence of pediments, abrupt piedmont-hillslope transitions, steep bedrock slopes, faceted spurs, wineglass valleys, and subparallel systems of high-gradient, narrow, steep-sided canyons orthogonal to range front. Only mapped in areas along range front where fault scarps and (or) lineaments on Quaternary surficial deposits or on Quaternary erosion surfaces are absent

Faults juxtoposing Quaternary alluvium against bedrock (other than major range-front faults)--Morphologically similar to major range-front faults except that associated fault systems are significantly less extensive and fault scarps are substantially lower, shorter, and less continuous. Solid lines indicate locations where fault scarps are abrupt and well-defined; long dashes indicate locations where scarps are less well defined

Faults forming scarps and (or) prominent topographic lineaments on Tertiary volcanic or sedimentary rocks--These scarps have morphologies ranging from partly rounded and moderately dissected to undissected. Topographic lineaments are composed of alignments of one or more of the following landforms: abrupt scarps, linear hillside ridges, benches and trenches, linear reaches of stream channels and small stream valleys, ridge-crest saddles and cols, linear depressions and small closed basins. Many of these faults form closely spaced groups in areas underlain by Tertiary ash-flow tuffs and lava flows

Table 1. *Characteristics of young faults in the Winnemucca 1° by 2° quadrangle*
[See figure 1 for area boundaries]

	Northeast		Southeast		West	
Orientation	Length (km)	Percent	Length (km)	Percent	Length (km)	Percent
90.0 to 80°	5.6	1.4	4.2	0.5	- - -	- - -
79.9 to 70°	19.4	4.9	43.2	5.2	18.0	2.0
69.9 to 60°	28.0	7.1	40.8	4.9	- - -	- - -
59.9 to 50°	38.0	9.7	85.2	10.2	66.4	7.2
49.9 to 40°	64.4	16.4	75.8	9.1	55.8	6.1
39.9 to 30°	39.4	10.0	85.6	10.2	112.6	12.3
29.9 to 20°	52.6	13.4	81.8	9.8	134.2	14.6
19.9 to 10°	23.6	6.0	135.4	16.2	200.0	21.8
09.9 to 0.0°	44.2	11.2	141.6	16.9	120.2	13.1
359.9 to 350°	41.2	10.5	78.0	9.3	47.0	5.1
349.9 to 340°	15.0	3.8	40.8	4.9	62.6	6.8
339.9 to 330°	19.0	4.8	11.4	1.4	17.4	1.9
329.9 to 320°	- - -	- - -	2.0	0.2	28.8	3.1
319.9 to 310°	2.8	0.7	- - -	- - -	17.0	1.9
309.9 to 300°	- - -	- - -	- - -	- - -	11.0	1.2
299.9 to 290°	- - -	- - -	- - -	- - -	4.2	0.5
289.9 to 280°	- - -	- - -	2.0	0.2	19.2	2.1
279.9 to 270°	- - -	- - -	8.2	1.0	4.0	0.4
Total (km)	393.2	100.0	836.0	100.0	918.4	100.0
Area (km2)	3680	- - -	6750	- - -	8460	- - -
Density (km/km2)	0.107	- - -	0.124	- - -	0.109	- - -
Number	88	- - -	160	- - -	193	- - -
Density (no./km2)	0.024	- - -	0.024	- - -	0.023	- - -
Mean length (km)	4.5	- - -	5.2	- - -	4.8	- - -

Table 2. *General photogeologic and geomorphic criteria used to estimate general ages of piedmont surfaces*

Map unit	Age	Depth of dissection	Drainage net morphology	Interfluve morphology	Geomorphic relations	Typical geomorphic environments	General field criteria
Unit Q3	Holocene (0 to 10 ka)	Shallow to none; generally < 3 m	Predominantly radial from fan apex; channels typically poorly to moderately well defined	Typically poorly defined; bar and swale micro-topography common on most surfaces	Surfaces cut pluvial shorelines and (or) late Pleistocene glacial moraines	Distal piedmont surfaces; channels and terraces in proximal areas; (proximal surfaces along some highly active range fronts)	Unweathered to slightly weathered; very weak to weak soil development
Unit Q2	Late Pleistocene (10 to 130 ka)	Shallow to moderate; typically 2-6 m	Predominantly distributary; however, some channels head on piedmont; well-defined channels	Typically well defined; surfaces broad and flat with abrupt margins	Surfaces overlain by pluvial shorelines and (or) latest Pleistocene glacial moraines	Proximal to distal piedmont surfaces; some inset terraces	Weakly to moderately well developed soils; interlocking stone pavements
Unit Q1	Early to middle Pleistocene (0.13 to 1.5 Ma)	Moderate to deep; commonly >10 m	Predominantly subparallel; well-defined channels	Well defined; older interfluve surfaces (ballenas) commonly narrow and irregular	Surfaces overlain by pluvial shore-lines and (or) latest Pleistocene glacial moraines	Generally confined to intermediate and proximal piedmont areas	Moderately to very well developed soils; interlocking to highly degraded stone pavements

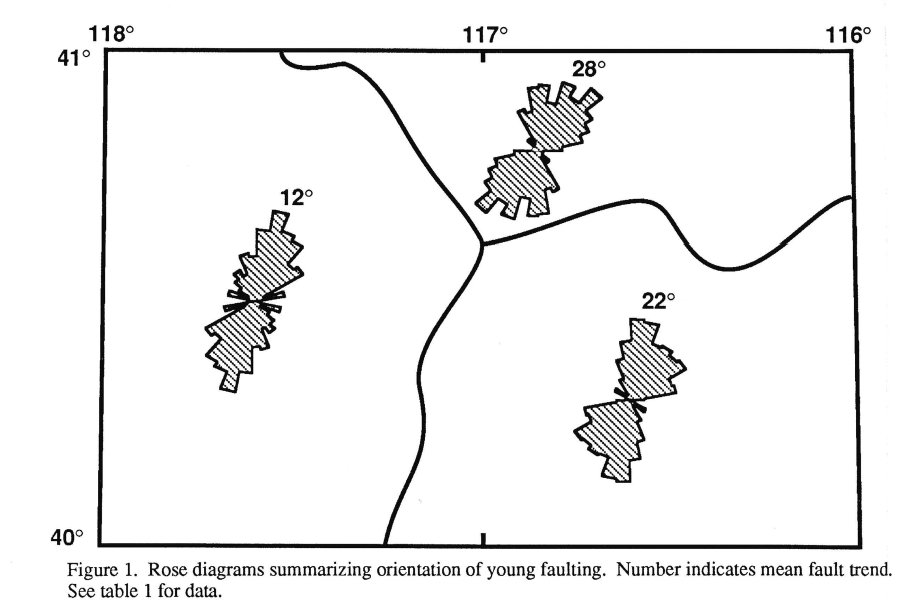

Figure 1. Rose diagrams summarizing orientation of young faulting. Number indicates mean fault trend. See table 1 for data.

RECONNAISSANCE PHOTOGEOLOGIC MAP OF YOUNG FAULTS IN THE WINNEMUCCA 1° BY 2° QUADRANGLE, NEVADA

By
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1991

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