

DISCUSSION

This sheet includes maps that show the interpreted thickness and the depth to base of uppermost Pleistocene and Holocene deposits in California's State Waters for the Offshore of Half Moon Bay map area (Map A, B), as well as for a larger area that extends about 91 km along the coast from Bolinas to Pescadero (Map C, D) to establish a regional context. This uppermost stratigraphic unit (blue shading in seismic-reflection profile of Fig. 1; see also, Figs. 1, 2, 3, 4, 6, 7, 9; 10 on sheet 8) is inferred to have been deposited during the post-Last Glacial Maximum (LGM) sea-level rise in the last about 21,000 years (see, for example, Pelletier and Fairbanks, 2006; Stanford and others, 2011). The unit commonly is characterized either by "acoustic transparency" or by parallel, low-amplitude, low- to high-frequency, continuous to moderately continuous, diffuse reflections (terminology from Michum and others, 1977). The acoustic transparency can be caused by extensive wave winnowing, which results in a uniform sediment grain size and the consequent lack of acoustic-impedance contrasts needed to produce seismic reflections. On the continental shelf, the contact with underlying units is a transgressive surface of erosion commonly marked by angularity, channeling, or a distinct upward change to lower amplitude, more diffuse reflections.

To make these maps, water bottom and depth to base of the post-LGM horizons were mapped from seismic-reflection profiles (Fig. 1; see also, sheet 8). The difference in the two horizons was exported for every shot point as XY coordinates (UTM zone 10) and two-way travel time (TWT). The thickness of the post-LGM unit (Map B, D) was determined by applying a sound velocity of 1,600 m/s to the TWT. The thickness points were interpolated to a preliminary continuous surface, overlaid with zero-thickness bedrock outcrops (see sheet 10), and contoured, following the methodology of Wong and others (2012).

The thickness data points are dense along tracklines (about 1 m apart) and sparse between tracklines (1 km apart), resulting in minor contouring artifacts. To incorporate the effect of a few rapid thickness changes along faults, to remove irregularities from interpolation, and to reflect other geologic information and complexity, minor manual editing of the preliminary thickness contours was undertaken. Contour modifications and regridding were repeated several times to produce the final sediment-thickness maps. Information for the depth to base of the post-LGM unit (Maps A, C) was generated by adding the sediment-thickness data to water depths determined by multibeam bathymetry (see sheet 1).

The thickness of the post-LGM unit in the Offshore of Half Moon Bay map area ranges from 0 to 12 m (Map B), and the depth to base of the unit ranges from less than 5 to 56 m (Map A). This relatively thin sediment cover suggests a lack of sediment "accommodation space" (Cattaneo, 2006), consistent with regional uplift expressed by the young, high topography of the Santa Cruz Mountains and Coast Ranges, as well as nearby uplifted marine terraces (Weber, 1990; Lajoie and others, 1991; Brabb and others, 1998). This uplift has been attributed to the northwest-

transpressive bend in the San Andreas Fault (see, for example, Zoback and others, 1999), which lies about 12 km northeast of the map area. The San Gregorio Fault Zone, which cuts through the Offshore of Half Moon Bay map area, consists of an east strand ("Seal Cove Fault" or "Covepoint Fault"), a west strand ("Frijoles Fault"), and related splay faults (Fig. 2). The thick sediment (about 12 m) in the map area is west of or east of the Frijoles Fault, and it results from offset and (or) gentle downwarping associated with that structure (for example, Fig. 1).

Five different "domains" of sediment thickness are recognized on the regional sediment-thickness map (Map D): (1) the Bolinas shelf, located west of the east strand of the San Gregorio Fault Zone, in the northwest part of the regional map (Map D); (2) the San Andreas graben, located between the San Gregorio Fault Zone and the Golden Gate Fault, east-southeast of the Bolinas shelf and both southwest and southeast of the Marin shelf; (3) the Marin shelf, located both northeast and northwest of the San Andreas graben and north of the San Francisco ebb-tidal delta paleovalley; (4) the northeast-trending San Francisco ebb-tidal delta paleovalley, located outside the Golden Gate at the mouth of San Francisco Bay, between the Marin shelf and San Andreas graben on the north and the Pacifica-Pescadero shelf on the south; and (5) the Pacifica-Pescadero shelf, which is located south of the San Francisco ebb-tidal delta paleovalley and extends south all the way to Pescadero Point (including all of the Offshore of Half Moon Bay map area).

The five sediment-thickness domains have distinct geologic controls. The Bolinas and Pacifica-Pescadero shelves are uplifted and are relatively sediment poor (mean sediment thicknesses of 0.5 and 3.6 m, respectively). Thicker sediment accumulations (as much as 20 m) on the western margins of the Pacifica-Pescadero shelf (within California's State Waters) are associated with west-sideward slip on the west strand of the San Gregorio Fault Zone and with deposition on the outward, west-dipping Pigeon Point block (McCulloch, 1987; see also, Fig. 2) farther south, offshore of Pescadero Point. The San Andreas graben is a rapidly subsiding, fault-controlled sedimentary basin (Cooper, 1973; Ryan and others, 2008) that has sediment thicknesses of as much as 37 m; the Marin shelf forms the upland paleovalley that formed during the last sea-level lowstand, with sediment thicknesses of as much as 12 m along the trough axis. Although the southern part of the San Andreas graben may extend into the paleovalley, the north flank of the paleovalley is used here as the boundary when calculating sediment volumes for the five sediment-thickness domains (see table 7.1 in accompanying pamphlet). Subsidence in the San Francisco ebb-tidal delta paleovalley and the San Andreas graben can be partly attributed to the northward change in strike of both the San Andreas and San Gregorio Fault Zones offshore of San Francisco, which has resulted in the local change from contractional deformation to extensional deformation (Zoback and others, 1999).

Map E shows the regional pattern of major faults and of earthquake occurrence between 1967 and April 2014 that have influenced or measured magnitudes of 2.0 and greater. Fault locations, which

have been simplified, are compiled from our mapping within California's State Waters (see sheet 10) and from the U.S. Geological Survey's Quaternary fault and fold database (U.S. Geological Survey and California Geological Survey, 2010). Earthquake epicenters are from the Northern California Earthquake Data Center (2014), which is maintained by the U.S. Geological Survey and the University of California, Berkeley, Seismological Laboratory. Map E also shows the inferred location of the devastating 1906 California earthquake (M_s 8.1) (1906), thought to have occurred on the San Andreas Fault offshore of San Francisco (see, for example, Bolt, 1968; Lomas, 2005). Map E clearly shows that the largest number of earthquakes in the region occur within the broad San Andreas Fault Zone between Pacifica and Bolinas; events west of the east strand of the San Gregorio Fault Zone and east of the Golden Gate Fault are much less common. The Offshore of Half Moon Bay map area is notably devoid of seismicity during this time period.

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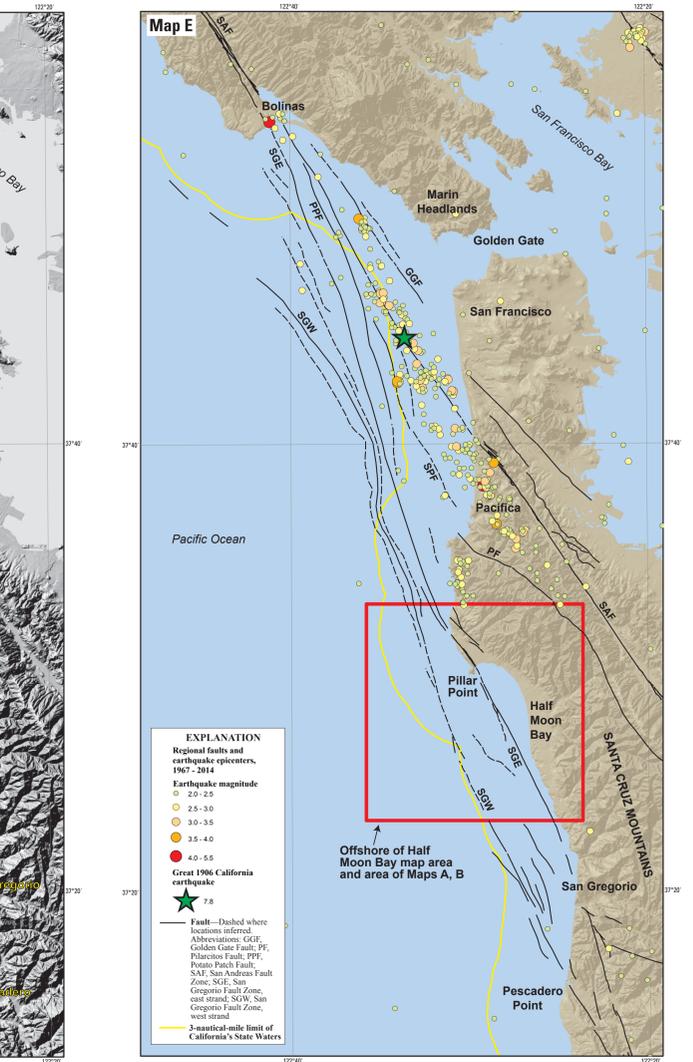
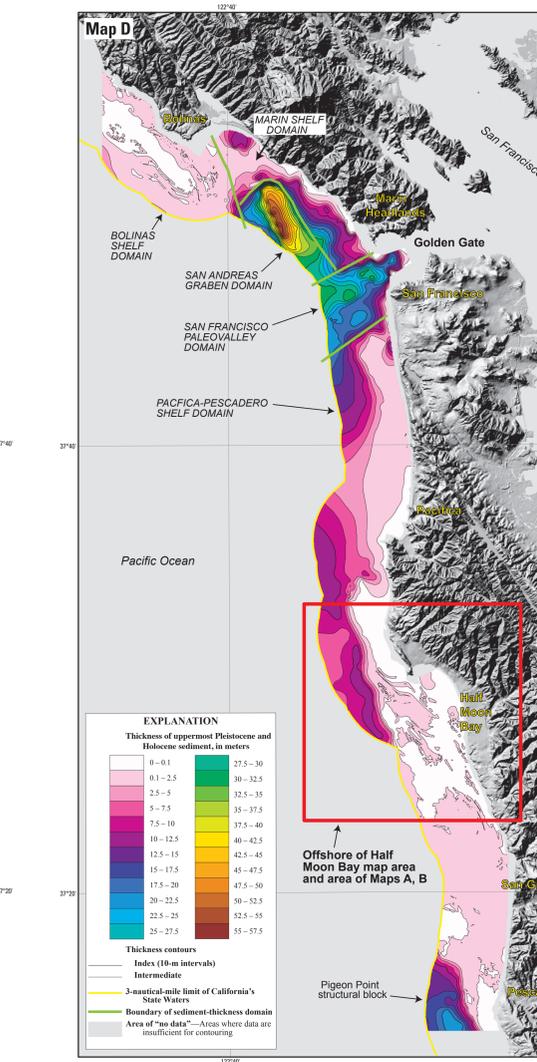
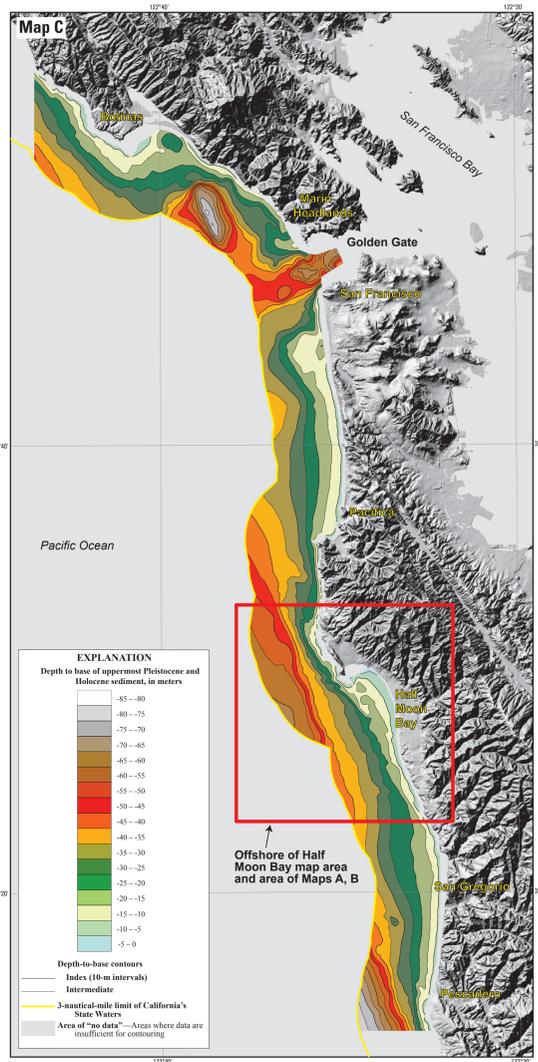
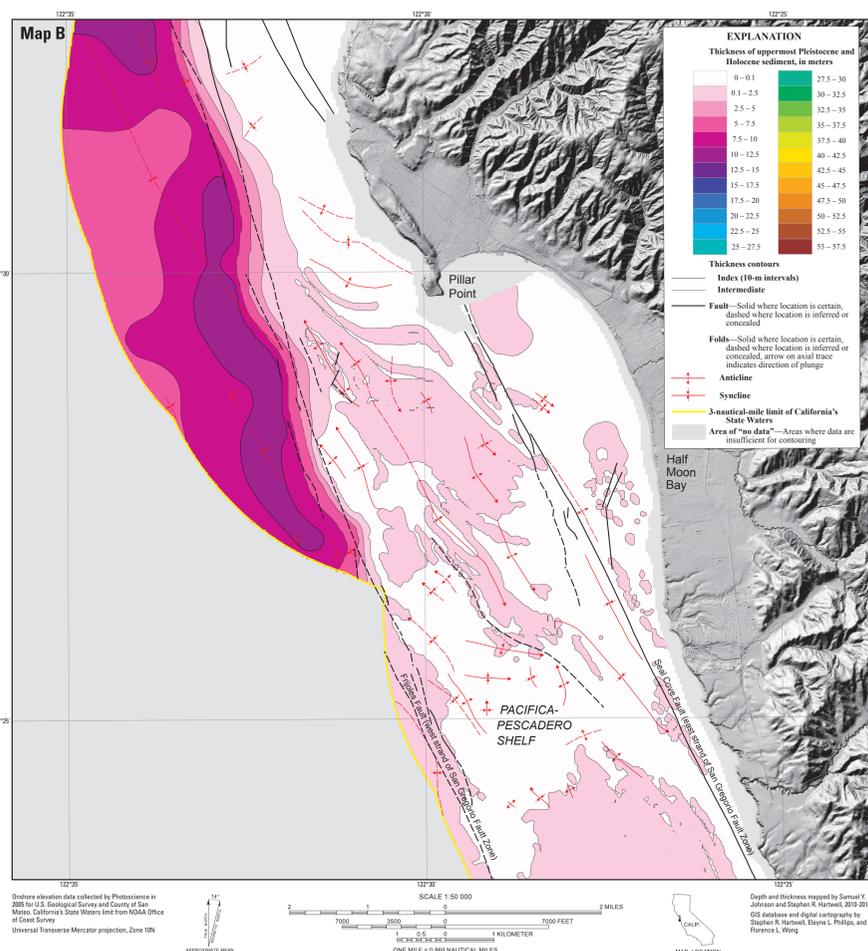
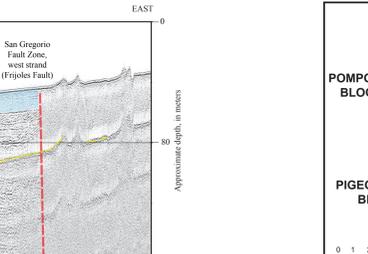
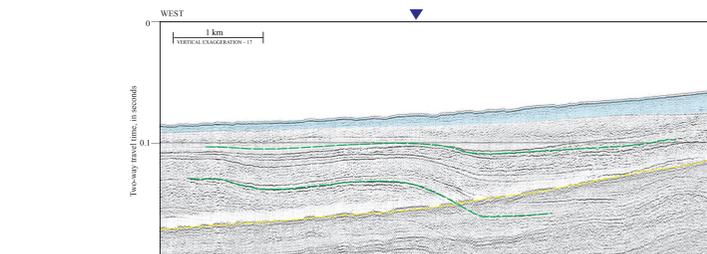
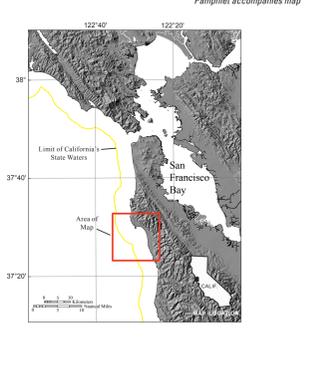
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Local (Offshore of Half Moon Bay Map Area) and Regional (Offshore from Bolinas to Pescadero) Shallow-Subsurface Geology and Structure, California

By
Samuel Y. Johnson, Stephen R. Hartwell, Ray W. Sliter, Janet T. Watt, Eyleyn L. Phillips, Stephanie L. Ross, and John L. Chin

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