



GENERAL FEATURES OF THE DISTRICT

By Carl Fries, Jr.

The rocks in the Black Range tin district, as mapped, are divided into six units. The four lower units are predominantly of volcanic origin and comprise lava flows, tuff, and breccia, which range in composition from basaltic to rhyolitic. The sum of their maximum thicknesses exceeds 8,000 feet, but the aggregate thickness at any one place was probably considerably less than that. For these reasons the units are designated as volcanic units. The two upper units are dominantly water-laid deposits derived from the volcanic rocks. Their aggregate thickness in the area mapped is more than 900 feet, and in places it may be as much as 1,000 feet.

The units are separated from each other by pronounced crustal unconformities, and there are many minor unconformities within the units. The rocks are partly altered but are not metamorphosed in the usual sense of the word. Although in large areas they appear to have been slightly tilted, they are not folded, and at most places where they dip at angles of less than 30° their dips are clearly initial, conforming roughly with the eroded slopes on which the beds were deposited or the lava flows poured out. The steep dip of the sheeting in the lava at certain places merely indicates the direction of flow at those places just before the lava solidified. Some small faults have been found, and others are doubtless concealed by alluvium or earlier gravel. The largest displacement observed along any fault is about 250 feet. The present basinlike structure of the separate areas within this district is probably a result of faulting.

The oldest rocks exposed in the tin district are probably underlain by Paleozoic sedimentary rocks, for the base sandstone, of Permian age, crops out west of Chloride, a few miles east of the area mapped. The volcanic rocks are undoubtedly Tertiary in large part, and they may all be Tertiary. Part of the clastic unit (Qt) also may be Tertiary, as this unit seems to correspond to what is known as the Olla conglomerate, of late Tertiary and Quaternary age, in western New Mexico and eastern Arizona.

The six units are described below in order of age.

Lower basic volcanic series

The rocks of the lower basic volcanic series (T₁) are dominantly basaltic and andesitic lavas, with some flows of dacite and latite. The lavas are interbedded with tuff, breccia, and water-laid clastic sediments that also have the composition of basic to intermediate volcanic rocks. Most of the flows and the layers of breccia are dark gray, green, blue-green, or purplish, but some are light-colored. The base of the series does not crop out in this district, but the part exposed has a maximum thickness of about 1,000 feet. The entire series is altered, and its most altered rocks contain abundant pyrite.

Felsitic volcanic series

The rocks of the felsitic volcanic series (T₂) are mainly lavas, which are generally whitish or pinkish in color and largely felsitic in texture. Most of the lava flows are fine-grained porphyritic rhyolite, but a few are andesite and latite. Tuff, breccia, and water-laid clastic sediments are interbedded with the flows in a few places. All the clastic and pyroclastic rocks are well cemented, but they generally appear fresher than the light-colored tuff and breccia of the lower basic volcanic series (T₁). The thickness of the series in the tin district ranges from 10 to about 700 feet.

Rhyolitic volcanic series

The rhyolitic volcanic series (T₃) is almost entirely confined to the west side of the Continental Divide. It occupies large areas widely distributed through the tin district, but it is best exposed along Diamond Creek. The maximum thickness of the series is now about 1,000 feet, but it must once have been much greater, for rhyolite debris is abundant in the overlying clastic unit (Qt). In a few places the rhyolitic series is absent and the clastic unit lies directly on the felsitic rocks (T₂).

The series comprises tuff, breccia, some sandstone and conglomerate, thick flows of porphyritic rhyolite (Tr), and thin flows of latite and andesite (T₃). The porphyritic rhyolite, which contains most of the tin-bearing veins in the district, is dominant in the lower part of the series; the latite and andesite rhyolite are dominant in the upper part. The rhyolite flows are light gray to reddish brown, and the latite is dark gray. The tuff, breccia, and clastic sediments are mostly light-colored, largely pale yellow or reddish, and are so much alike that it has not been possible to map them separately. They are interbedded with the flows and pinch out in many places, as along Diamond Creek. The clastic rocks are comparatively fresh, and unlike those of the felsitic series (T₂) they are poorly consolidated.

The porphyritic rhyolite (Tr) is fairly uniform in general appearance throughout the area mapped, but it shows minor differences in texture and in the relative abundance of the constituent minerals. These differences, together with the diverse attitudes of the flow lavas, are believed to indicate that the lava was extruded from several centers. Certain areas of rhyolite in which the sheeting is persistently steep are thought to mark the location of vents. One of the vents may be under the rhyolite that crops out along Diamond Creek, another between Seventyfour Creek and Silver Creek, and a third in the large area between Alexander Canyon and Squaw Creek. The rhyolite between Squaw Creek and Corduroy Canyon presumably has a separate source; those near the north edge of the map and those in the vicinity of Harbottle Creek probably were extruded from vents beneath them. No conclusion is offered regarding the source of the thin flat-lying flows that are exposed near Indian Peak between Corduroy and Harbottle Canyons; these flows overlie the rhyolite flows farther north and farther south.

Upper basic volcanic series

The upper basic volcanic series (T₄) consists mainly of basaltic lava flows, with which thin beds of rhyolitic tuff and water-laid sandstone and conglomerate are interbedded in a few places. The basalt in Pelona Mountain, where it is at least 1,500 feet thick, and in Beaver Creek Canyon, the upper 1,500 feet of Black Mountain, which is on the west side of Beaver Creek and west of the area mapped, appears to consist wholly of basalt. The basalt was probably extruded in part from beneath Pelona Mountain and in part from beneath Black Mountain.

Clastic unit

Clastic beds are exposed over the greater part of the area mapped. The clastic unit (Qt) comprises beds of three distinct types, which are not differentiated on the map. Thin-bedded tuff occurs near the base in a few places, notably along Beaver Creek north of Harbottle. The greater part of the unit consists of water-laid, poorly sorted, poorly cemented gravel, containing a large percentage of tuff eroded from the volcanic series. The uppermost beds, which are mostly unconformably deposited, consist mainly of water-worn fragments of volcanic rock, with very little tuff.

All of these water-laid, water-deposited clastic sediments were undoubtedly deposited during the same general period of time. The areas in which the beds were deposited were probably, as the unit contains no fine-grained lake-laid sediments, as are generally formed in closed basins. There are a few crustal unconformities within the unit. Some faults cut the lower beds but apparently do not affect the upper beds, indicating that the faulting occurred within the period during which this unit was being deposited.

The clastic unit contains a little cassiterite (tin oxide) but not enough to make it workable as placer ground. Before the present valleys and gullies were formed, the rocks in the area shown on the map were eroded, for the most part, to a nearly flat surface with a gentle slope to the west. A few hills, however, such as Bolter Peak and the Indian Peaks, rise above the surface. In some places the upland plain has a cover of gravel a few inches to about 3 feet thick, which has not been shown on the map. The age of this gravel is between that of the clastic unit (Qt) and that of the alluvium (Al).

Alluvium

Alluvium (Al) covers the bottoms of all the valleys in the district but is shown on the map only where it is more than about 4 feet thick. The greatest thickness of alluvium observed is about 25 feet, but it may be as much as 50 feet in a few places. The material is generally in character from place to place, but in general it is fairly coarse, and boulders as much as 3 feet in diameter are common at some localities. Nearly all the placer tin mined in the district has come from this recently deposited alluvium.

Mineral deposits

Tin-bearing veins.—Nearly all the tin-bearing veins known to date in this district are in the porphyritic rhyolite (Tr) of the rhyolitic volcanic series (T₃). A few veins of limonite have been found in the felsitic volcanics (T₂), for example north of Soles Creek, but they contain little tin. The veins in the porphyritic rhyolite are mainly stringers that consist predominantly of cassiterite (tin oxide) and cassiterite (tin oxide); the proportion of cassiterite ranges from a few percent to as much as 80 percent, the average being about 20 percent. The veins also contain a little magnetite, together with other quantities of the alluvial minerals cassiterite and stannite and of quartz, calcite, and pyrite. The veins are all small, most of them ranging in size from 1/2 inch to 2 feet long to tabular boulders an inch thick and 30 feet long. A few of the larger veins include lenses 2 to 3 inches thick and 1 to 4 feet long.

Tin placers.—Placer deposits of tin occur sparingly in the gravelly alluvium of the clastic unit (Qt). For example, in the vicinity of Bolter Peak, larger quantities are found in the alluvial deposits (all) of the stream valleys and in the residual soil and slope wash at higher levels than the stream deposits. The placers are mostly near the outcrops of the tin-bearing veins in the rhyolite, as on Squaw Creek and Harbottle Creek, but some of them are near the outcrops of tin-bearing gravels in the clastic unit, from which they have been reconstructed.

Other possible mineral deposits.—Some beds of kaolinized tuff (T₁) that occur north of Turkey Run, near the west edge of the area mapped, have been prospecting as a possible source of kaolin, and two small excavations have been made nearby in rock that is said to be gold-bearing.

Suggestions for prospecting

Search for tin-bearing veins may safely be confined to the porphyritic rhyolite. Placer tin, being derived primarily from these veins, is most likely to be found in slope wash on porphyritic rhyolite, and in alluvium near outcrops of that rock. Flat erosion surfaces of rhyolite and shallow draws on those surfaces appear to offer the greatest promise. The known placers probably be of the same character. The work done thus far has indicated that the grade of the placer gravel decreases as its thickness increases. The clastic unit and the deepest alluvial deposits have not been explored to great extent, but such exploration would be costly and seems unlikely to result in the finding of tin in commercial quantities. Small tin-bearing veins will probably be uncovered in mining placer deposits in the residual soil and slope wash on the rhyolite, but if these veins are not large enough at the surface to warrant mining it does not seem advisable to follow them far downward.

1/ Fries, Carl, Jr., Tin deposits of the Black Range, Catron and Sierra Counties, N. Mex.: U. S. Geol. Survey Bull. 922-N, p. 364, 1940.