INTRODUCTION The Big Sandy quadrangle includes an area of about 200 square miles n the northeastern part of Chouteau County in north-central Montana. he quadrangle is in the Missouri Plateau section of the northern Great lains physiographic province. Deposits of Pleistocene drift mantle the region, overlying sedimentary rocks of Cretaceous age. Many igneous dikes and sills crop out, and several faults are exposed. Known economic resources include sandand gravel and minor amounts of coal and bentonite.

The Missouri River flows through the southwestern corner of the quadrangle. Tributary streams have dissected areas along their lower courses into badland topography. A broad shallow valley, the partly filled pre-glacial course of the Missouri River, trends northeasterly across the area, separating the gently rolling upland plains of the northwestern and southeastern parts of the quadrangle.

Big and Little Sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the mort by the property of the sandy Creeks are the sandy of the sa Big and Little Sandy Creeks are the major tributary streams in the area. Big Sandy Creek with headwaters in the Bearpaw Mountains east of the quadrangle, enters the area from the east and flows in a northerly course to join the Milk River. Little Sandy Creek rises in the southwestern foothills of the Mountains and joins the Missouri River in the southern part of the guadrangle. western foothills of the Mountains and joins the Missouri River in the southern part of the quadrangle.

Local relief is from 50 to 200 feet, maximum relief is about 500 feet. The surface ranges in altitude from about 2,500 feet along the Missouri River, to about 3,000 feet on the upland areas.

The climate is characterized by low rainfall, dry atmosphere, hot summers, cold winters, and a large proportion of sunny days. Annual precipitation at Big Sandy averages 12.15 inches. The heaviest rainfall is in late spring and early summer; occasional heavy snows are recorded during the winter. Maximum and minimum recorded temperatures are 105° and -48°. During the summer months the daily range is great and the nights are usually cool. The frost-free period dates from late May to early September, averaging about 100 days. The area is subject to strong winds which are more prevalent in the early spring. Hail storms occur Judith River formation

south of Big Sandy. A well-developed network of improved roads covers most parts of the quadrangle. Bedrock Formations Bedrock formations, all of Cretaceous age. which crop out in the quadrangle include the Colorado shale, Telegraph Creek formation. Eagle sandstone, Claggett shale, Judith River formation, Bearpaw shale, Fox Hills sandstone, and Hell Creek formation. The Colorado shale, Claggett shale, and Bearpaw shale are gray to brown marine shales, and the Telegraph Creek formation, Eagle sandstone, Judith River formation, Fox Hills sandstone, and Hell Creek formation consist of interbedded sandstone, shale, carbonaceous shale, and clay. The formations, locally interrupted by folding and faulting, dip northeast a few feet per mile.

winds which are more prevalent in the early spring. Hail storms occur

Wheat farming and cattle raising are the chief industries in the area. Big Sandy, a town of 800 population, is the local marketing center. The area is served by the Great Northern Railway and by U. S. Highway No. 87. State Highway No. 116 is paved for a distance of six miles

ocally during the summer.

by folding and faulting, dip northeast a few feet per mile. Colorado shale The Colorado shale (Kc) underlies the entire quadrangle and consists of about 1,800 feet of dark grayish-blue marine shale, of which only the upper 400 to 500 feet crop out. The formation is predominantly shale but contains thin beds of bentonite, sandstone, and sandy shale, as well as numerous calcareous concretions. The concretionary layers and bentonite beds are more prevalent in the upper two-thirds of the formation; the sandstone beds in the lower one-third. Most of the concretions are spheroidal masses of rusty-brown clay-ironstone or dark-gray limestone one to three feet in diameter that are generally arranged along bedding planes. Many of the concretions contain invertebrate fossils including Scaphites, Baculites, Inoceramus, and Ostrea. Calcite filled shrinkage cracks are common. Cracks and fissures in the shale contain veins of

The Telegraph Creek formation (Kt), a transition zone between the Colorado shale and the Eagle sandstone, consists of alternating beds of white to buff sandstone and dark grayish-blue shale. In places the beds reach a thickness of 150 feet or more, and in other areas are locally absent. The formation is generally exposed in gentle to steep weathered The Telegraph Creek formation was named by Thom (1922, p. 38) from exposures on the Crow Indian Reservation southeast of Billings, Montana. Eagle sandstone

Telegraph Creek formation

The Eagle sandstone underlies most of the quadrangle and is exposed along the Missouri River and in several faulted areas. The upper and middle members (Ke) together consist of 125 to 150 feet of alternating beds of gray to buff sand, sandstone, shale, carbonaceous shale, and coal. The upper member is locally a thick massive sandstone, the middle member is predominantly shale. Two thin beds of coal are present near the base of the middle member. A conglomerate six inches to two feet thick of polished chert and quartzite pebbles in a matrix of fine sand and clay occurs at or near the top of the upper member. The pebbles are black, dark brown, dark blue, and dark green, round to oval, and range from one-eighth to two inches in diameter. In some localities the pebbles occur along bedding planes in the upper sandstone member, and also in the basal part of the overlying Claggett shale. The Virgelle sandstone member (Kev) comprises the lower 80 to 100 The Virgelle sandstone member (Kev) comprises the lower 80 to 100 feet of the formation and is a hard, massive, white to buff, medium-grained cross-bedded sandstone containing sandy limonitic concretions. The upper and middle members generally form a subdued type of badland topography. The Virgelle sandstone member forms steep cliffs. The formation was named by Weed (1899, p. 2) from exposures at the mouth of Eagle Creek six miles south of the quadrangle. Bowen (1915, p. 7) applied the name Virgelle sandstone member to the lower massive sandstone at exposures near the town of Virgelle in the southwestern corner of the quadrangle. The Eagle sandstone is conformably overlain by the Claggett shale.

The Claggett shale (Kcl) underlies the northern and central parts of the quadrangle. The formation consists chiefly of brownish-gray marine shale ranging in thickness from 450 to 500 feet. The upper 150 feet

beds of bentonite as much as 18 inches thick. Rusty-yellow to dark-gray spheroidal calcareous concretions are common in the middle part of the formation. These concretions are generally one to four feet in diameter: elicite-filled shrinkage cracks are common: cone-in-cone structure oc-irs along upper and lower surfaces; many contain specimens of Baculites. nite (gypsum) crystals occur in creeks and fissures in the shale. Badland topography is formed on both steep and gentle slopes of the Claggett shale. On steep slopes slumping and landsliding are common.

The formation was named by Stanton and Hatcher (1905, p. 13) from exposures near old Fort Claggett, at the mouth of the Judith River, 30 miles southeast of the Big Sandy quadrangle.

The Claggett shale is conformably overlain by the Judith River formation. The contact was originally placed at the top of a massive, light-brown ferruginous marine sandstone, which overlies sandy yellowishgray marine shale. In this report, however, the contact is placed at the base of this massive sandstone because the latter is lithologically more similar to the Judith Biver formation than it is to the the latter. similar to the Judith River formation than it is to the typical shale of

The Judith River formation (Kjr), which underlies the eastern and northwestern parts of the quadrangle, consists of about 500 feet of alternating beds of light-gray to buff sandstone, and gray to buff shale. The sandstone is generally fine-grained and thin-bedded to massive; the shale is in many places carbonaceous and contains a few thin beds of coal. The lower 200 feet includes more coal, carbonaceous shale, and dark-colored deposits than the upper 300 feet, which consists predominantly of light-colored sandstone and shale. Light-brown ferruginous sandstone beds are a marked feature of the basal 100 feet.

Vertebrate fossils, mainly dinosaur remains, are most common in the the upper beds; invertebrate fossils, chiefly oyster shells, are found in the carbonaceous zones in the lower beds, and at the top of the formation. The bivalve Tancredia americana occurs in the massive ferruginous sandstone at the base of the formation. Large quantities of silicified wood and plant fragments are also found in the formation.

The sandstone, shale, and clay of the formation generally weather into hard barren surfaces and erode to badland topography.

into hard barren surfaces and erode to badland topography.

The Judith River formation was named by Hayden (1871, p. 97) from exposures near the mouth of the Judith River. The formation is conformably overlain by the Bearpaw shale. Bearpaw shale

The Bearpaw shale (Kbs) which outcrops in a downfaulted area five The Bearpaw shale (Kbs) which outcrops in a downfaulted area five miles southeast of the town of Big Sandy, consists of about 700 feet of brownish-gray to bluish-gray marine shale, of which only about the upper 300 to 400 feet is exposed. Bentonite beds, ranging in thickness from a fraction of an inch to 10 inches, occur throughout the formation. Selenite (gypsum) crystals occur in cracks and fissures in the shale. Gray and dark-red calcareous concretions are numerous, particularly in the middle part of the Bearpaw shale. Many of the concretions are fossiliferous and contain specimens of Baculites, Acanthoscaphites, and Placenticeras. Fossils of marine reptiles are also reported.

In the Big Sandy quadrangle the Bearpaw shale crops out in steep slopes, many of which are covered by slumps and landslides. slopes, many of which are covered by slumps and landslides.

The formation was named by Stanton and Hatcher (1905, p. 14) from exposures south of the Bearpaw Mountains.

The Bearpaw shale grades into the Fox Hills sandstone through a transition zone of alternating thin-bedded gray to buff shale and sandstone beds. Fox Hills sandstone

Scaphites, Baculites, Inoceramus, and Ostrea. Calcite filled shrinkage cracks are common. Cracks and fissures in the shale contain veins of selenite (gypsum).

The formation is generally exposed in steep weathered slopes that are subject to landsliding and slumping, particularly along bentonite beds. It is conformably underlain by the Kootenai formation which is not exposed in the quadrangle.

The Colorado shale was named by Hayden (1876, p. 45) from outcrops east of the Front Range in Colorado.

The Fox Hills sandstone (Kfh) also crops out in the faulted area five miles southeast of Big Sandy. The formation consists of 100 to 200 feet of light-brown to light-gray sand and sandstone interbedded with tan and gray shale and carbonaceous shale. The sandstone in the lower to 100 feet is thin-bedded to massive and locally cross-bedded. Limonite concretions are common throughout the sandstone beds. Invertebrate brackish water fossils including Anomia and Corbula were found in carbonaceous shale beds near the top of the formation. The Fox Hills sandstone crops out as steep slopes and numerous sandstone ledges. The formation was named by Meek and Hayden (1862, p. 419) from exposures in Fox Ridge in north-central South Dakota. The Fox Hills sandstone is overlain by the Hell Creek formation. Hell Creek formation

> The Hell Creek formation (Khc), also exposed in the downfaulted area southeast of Big Sandy, is composed of thin-bedded, gray, green, pink, and buff clay, shale, and sandstone. Small, nodular claystone concretions are abundant in the middle part. A black bentonitic clay bed, 10 to 15 feet thick, is a distinctive marker bed in the lower 100 feet. The formation ranges from 400 to 500 feet in thickness, but only the lower 150 to 200 feet are exposed in the Big Sandy quadrangle. Fragments of dinosaur bones are found in the lower part of the formation.
>
> In the Big Sandy quadrangle the formation has been eroded into badland topography. and topography.
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> The Hell Creek formation was named by Barnum Brown (1907, p. 829-835) from exposures along Hell Creek south of the Fort Peck reservoir in northeastern Montana.

> SURFICIAL FORMATIONS Most of the quadrangle is covered by unconsolidated surficial deposits resulting from glaciation and stream action during the Pleistocene and Recent periods. These deposits are separated into three groups:
> (1) ground moraine deposited directly from glacial ice masses, (2) outwash deposits laid down by meltwater derived from the ice masses, and 3) deposits of material reworked by stream action, slopewash, or wind

Outwash deposits Silt, sand, and gravel classified as terraces (Qt), kames (Qk), eskers (Qe), silt deposits (Qst) and sand deposits (Qs), rest on the ground moraine in many parts of the quadrangle. The kames and eskers were deposited by glacial meltwater at or near the ice margin. The terraces were deposited along the Missouri River. The sand deposits are an outwash plant deposited by missouri River. races were deposited along the Missouri Hiver. The sand deposits are an outwash plain deposited by running water immediately in front of the ice margin, and the silt deposits represent sediments laid down in temporary lakes formed in front of the ice margin.

These deposits are chiefly light-brown or buff, exhibit crude to well-developed stratification, and with the exception of the silt deposits, are generally less than 20 feet thick. The silt deposits are locally more than 100 feet thick.

Reworked deposits Alluvium (Qal), colluvium (Qco), and alluvium-colluvium (Qac) consists of interbedded light-brown to gray clay, silt, sand, and gravel. Alluvium is generally confined to valley bottoms adjacent to streams, and colluvium mantles the sides of most major stream valleys and areas downslope from large bedrock outcrops. Alluvium-colluvium fills the pre-glacial valley of the Missouri, south, east, and northeast of Big Kettle deposits (Qkd), consisting of dark gray clay, silt, and sand, floor the bottoms of undrained depressions on the ground moraine.

Sand dunes (Qsd) consist of sand deposits (Qs) reworked by wind

IGNEOUS ROCKS Igneous dikes and sills of Tertiary age crop out in the eastern and northern parts of the quadrangle. They are shonkinite, a dark igneous rock composed of hornblende, olivine, pyroxene, biotite, and orthoclase. Most of the dikes and sills are more resistant to weathering than the confining sedimentary rocks and protrude a few feet above the surrounding land surface. Other than a slight baking effect at the contact, the igneous rocks have caused little alteration of the rocks into which they were intruded. The dikes and sills are believed to have been intruded after the faulting of the sedimentary strata, as indicated by the fact that in adjacent areas, many dikes cut across fault planes without displacement.

The Cretaceous sedimentary rocks of the region are essentially flatlying with a slight tilt to the northeast locally interrupted by faulting and folding. The faults are part of the thrust-faulted belt encircling the Bearpaw Mountains and have been the subject of a special study by All of the faults mapped in the quadrangle, with the exception of those in the downfaulted area five miles southeast of Big Sandy, are of the reverse type. Vertical displacement along the faults ranges from about 100 to 500 feet. The strike and dip of the fault plane generally parallels the strike and dip of the beds on the upthrown side of the fault. Within a quarter of a mile from a fault, steeply dipping strata on the upthrown side generally flatten to the regional dip on the upthrown side generally flatten to the regional dip.

Reeves (1929, pp. 172-185) believes that the thrust faults are shallow and do not extend below the upper part of the Colorado shale. His reasons are summarized as follows: (1) none of the faults expose strata below the middle of the Colorado shale; (2) thrust faults of this type are absent where strata below the Colorado shale are at the surface; (3) the lowest beds exposed along faults have about the correliance. the lowest beds exposed along faults have about the same inclination as the fault plane, suggesting that the faulting began as a movement along a bedding plane and passed upward along a curving surface; (4) subsurface data indicate that below the thrust faults the underlying Kootenai formation is in a parameter. mation is in a normal position and not at a higher level on the upthrust side of the hypothetically extended fault; (5) mathematical calculation of the depth of faulting gives an average of 900 feet below the Eagle sandstone. During the present investigations no evidence was found that did not agree with Reeves' conclusions concerning the depth of the fault-The faults mapped at the downfaulted area five miles southeast of Big Sandy are of the normal type. Displacement along the fault planes ranges from about 100 to 300 feet. Additional faults may exist under cover of glacial drift in the

ECONOMIC GEOLOGY Extensive resources of sand and gravel are located in the glacial outwash deposits. Minor amounts of coal, bentonite, and riprap have been found. Five shallow wells have been drilled in an unsuccessful Sand and gravel

Kame, esker, and terrace deposits are sources of large quantities of sand and gravel suitable for road material. Material near the surface of the deposits generally is carbonate coated, but at greater depths clean material suitable for concrete aggregate is available. The outwash sand deposits covering about six square miles in the southwest part of the quadrangle. quadrangle are a source of large quantities of fine to medium, well-

Scattered glacial boulders, consisting chiefly of quartzite, granite, and basic igneous rocks ranging from three inches to two feet in diameter, are a limited source of riprap. The more resistant parts of the shonkinite dikes and sills are another possible source of material.

Exploitation of the sub-bituminous coal beds in the Judith River Ground moraine (Qgm) forms the surface of over half the quadrangle operations. The coal beds range from a few inches to about two and oneand underlies younger deposits in other large areas. Much of the ground moraine is composed of till, and unstratified mixture of buff clay, silt, and sand interspersed with gravel, cobbles, and boulders. The deposits are generally 5 to 30 feet thick, but in filled pre-glacial valleys are as much as 150 feet thick. Small areas in the south and south-central parts of the quadrangle have a swell and swale topography of low rounded hills and shallow undrained depressions. Long ridges of till, 10 to 30 feet high, from one-fourth to 1 mile in length in the southeastern part of the quadrangle has been classified as drumlinoidal ridges, formed by filling of long tunnels on the underside of the glacial ice.

Half feet in thickness. Small mines were operated in sec. 16, T. 26 N., R. 12 E. and in sec. 22, T. 28 N., R. 13 E., but are now abandoned.

The thinness and variable extent of the coal beds, plus thick overburden limit the possibilities for profitable commercial mining in the area. Measured sections of the beds are shown on the Coal Sections Chart and the locations are recorded on the map.

A brief account of the coal outcrops in the eastern part of the quadrangle is included in a report by Bowen (1914, pp. 77-80). Coal outcrops in the western part of the quadrangle are described by Pierce and Hunt (1937, pp. 261-263). half feet in thickness. Small mines were operated in sec. 16. T. 26 N.

Commercial development of bentonite beds in the Colorado and Claggett shales is governed by the thickness of the beds and the amount of overburden. Because these beds are less than 18 inches thick and generally covered by 50 to 100 feet of overburden they are not believed to be of commercial importance at present. Small quantities of bentonite for lining stock reservoirs can probably be obtained from outcrops in sec.

23, T. 26 N., R. 12 E., where beds in the Claggett shale are near the surface.

Oil and Gas

GEOLOGIC HISTORY The fine-grained sediments which compose the Colorado shale were deposited in a broad shallow sea which occupied the continental interior of the United States during much of the Cretaceous period. Intermittent Bowen, C. F., 1914, The Big Sandy coal field, Chouteau County, Montana: U. S. Geol. Survey Bull. 541-H, p. 356-378. olcanic eruptions in adjacent areas spread thin layers of volcanic ash-Near the end of the deposition of the Colorado shale, local variations in source of supply produced the alternating sand and silt beds of Brown, Barnum, 1907, The Hell Creek beds of the Upper Cretaceous of Mon-River formation accumulated in a continental environment.

Another broad shallow sea served as a depository for sediments of the Bearpaw shale

This deformation probably initiated the erosion which removed most of the younger sedimentary rocks down to the Judith River formation. The resulting badland topography was highly modified by the Pleistocene continental glaciers which covered the area, depositing a blanket of drift. Streams were diverted into new courses when their valleys were drift. Streams were diverted into new courses when their valleys were blocked by ice, and on retreat of the glaciers failed to reoccupy their old valleys which were filled with drift. The present course of the Missouri River originated in this manner, and the old pre-glacial, drift-filled valley can yet be traced northeastward across the quadrangle. Several ice sheets may have covered the region during the Pleistocene epoch, but efforts to differentiate their deposits have not been successful The present ground surface of the quadrangle was formed during the The oil and gas possibilities of this region have been discussed by Reeves (1924, pp. 103-112) and by Pierce and Hunt (1937, pp. 258-261).

Six wells have been drilled in the quadrangle in search of oil and gas, but all have been abandoned as dry holes. A water well drilled in the NMXSE4 sec. 23, T. 27 N., R. 13 E., in 1952 obtained a moderate flow of dry gas at a depth of 400 feet. This is believed to flow from the upper part of the Eagle sandstone. The gas is utilized for heating and cooking purposes.

The present ground surface of the quadrangle was formed during the advance and subsequent retreat of the last glacier to pass over the region. The ground moraine consists chiefly of material deposited under recede, meltwater flowing over the ice and through cracks and crevasses deposited sand and gravel as kames, eskers, terraces, and outwash sand plains.

Since the disappearance of the ice sheet the ground surface has been somewhat modified in several areas. Erosion has formed badland topography near the Missouri River and its tributaries, and widespread deposits of alluvium and colluvium have accumulated.

REFERENCES

of swamps, lagoons, and shallow lakes.

The Claggett shale, including the beds of volcanic ash, was laid down in a broad shallow sea under conditions similar to the deposition of the depositio in a broad shallow sea under conditions similar to the deposition he Colorado shale.

The alternating beds of clay, sand, shale, and coal of the Judith formation accumulated in a continental environment.

Another broad shallow sea under conditions similar to the deposition morth-central Chouteau, western Hill, and eastern Liberty Counties, Montana: U. S. Geol. Survey Bull. 847-F, p. 225-270.

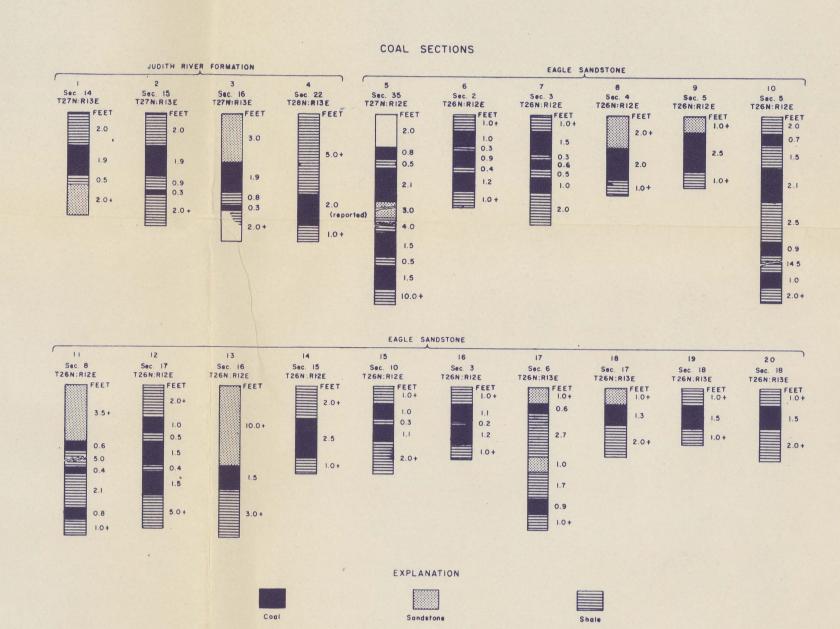
Reeves, F. M., 1924, Geology and possible oil and gas resources of the faulted area south of the Bearpaw Mountains, Montana: U. S. Geol. Overlying the Bearpaw shale the lower sands of the Fox Hills sand-stone were deposited under marine conditions. The upper part of the Fox Hills sandstone, and the overlying Hell Creek formation were deposited

The sandstone and the overlying Hell Creek formation were deposited

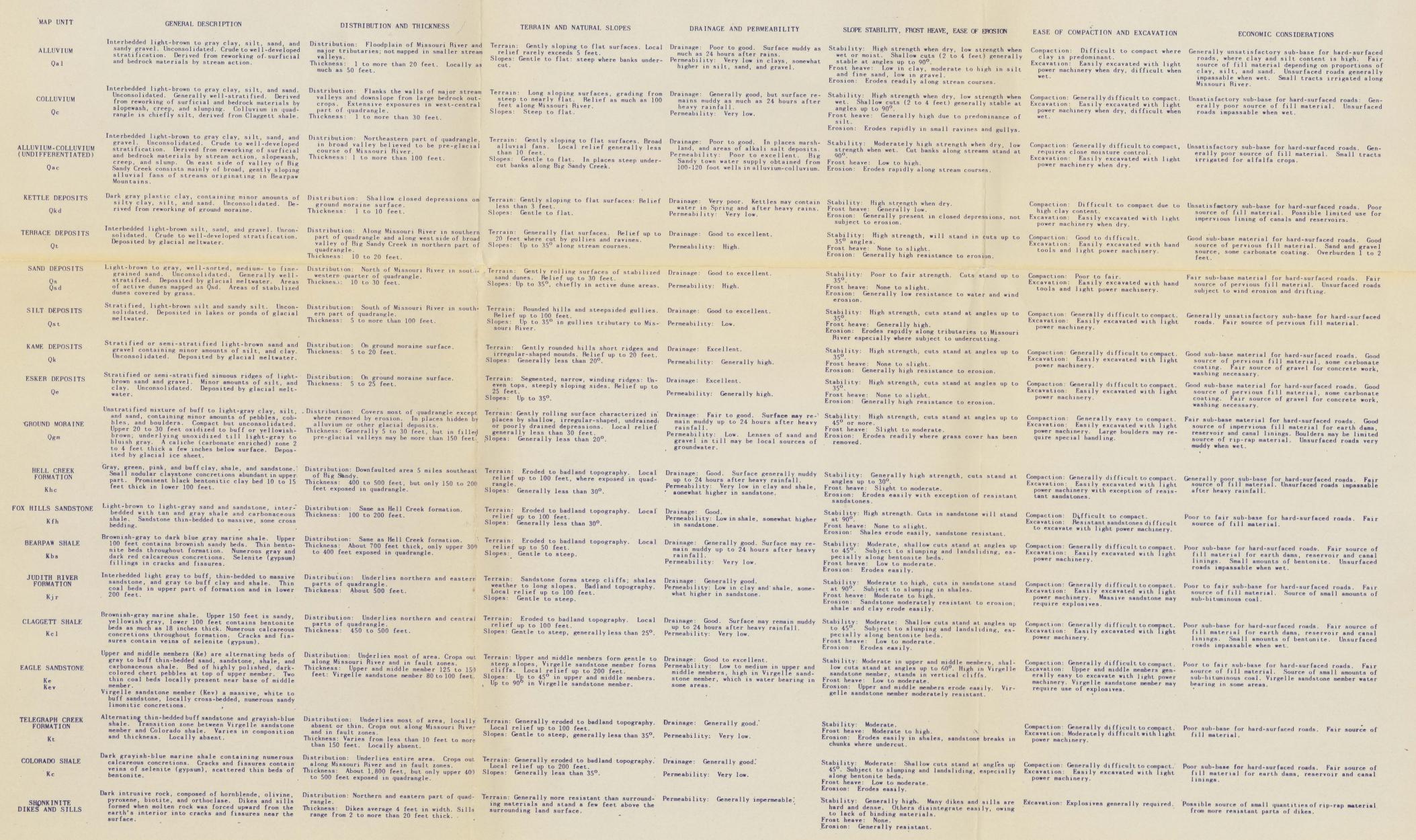
The sandstone and the overlying Hell Creek formation were deposited In a continental environment.

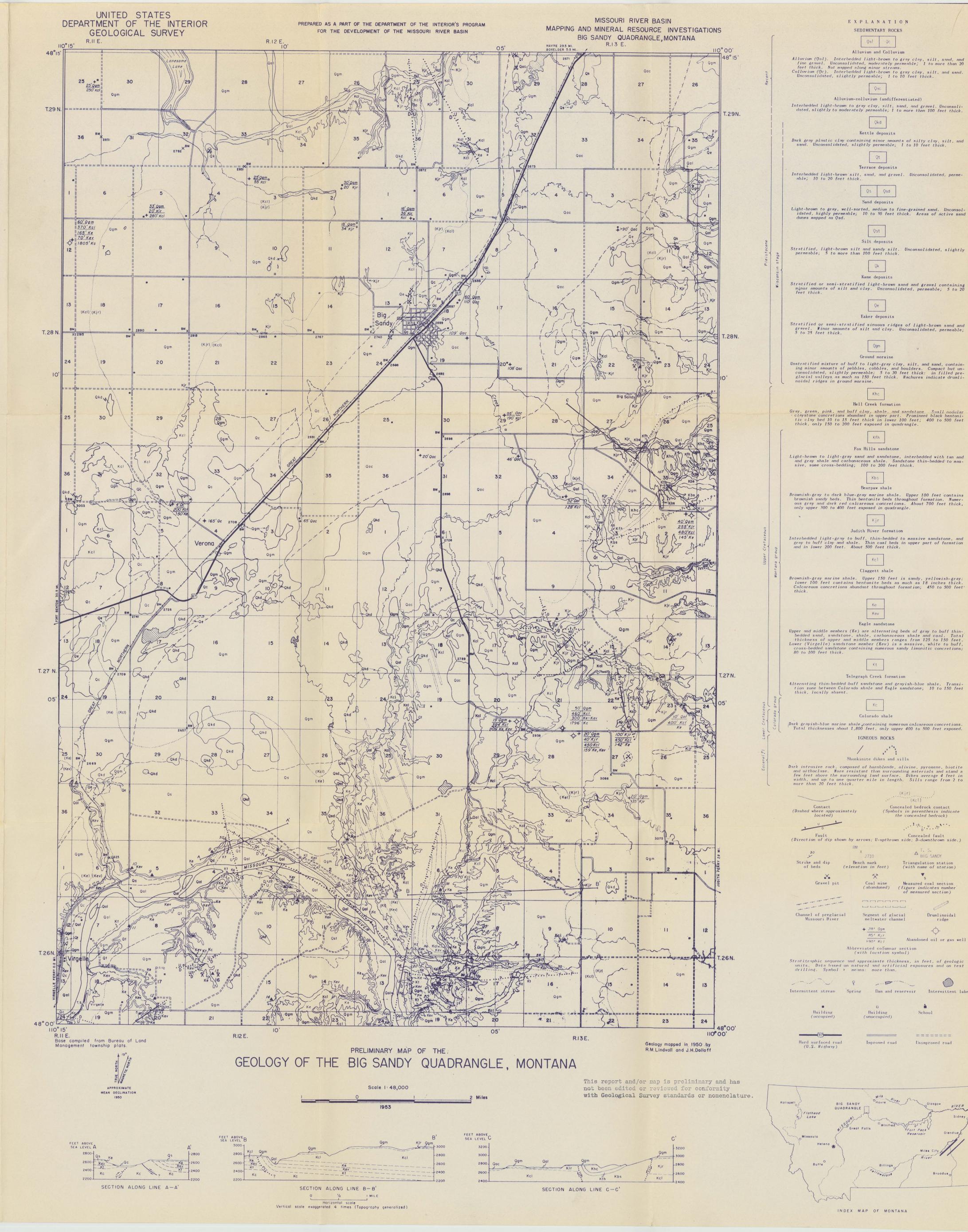
Still younger formations were probably deposited above the Hell
Creek formation within the quadrangle during the early part of the Tertiary period, but they have since been removed by erosion. Down-faulted blocks near the Bearpaw Mountain preserve parts of these formations.

During the middle part of the Tertiary period, forces accompanying the formation of the Bearpaw Mountains are thought to have caused the faulting of the strata in the Big Sandy quadrangle. The intrusion of the igneous dikes and sills also occurred at this time.



GENERALIZED DESCRIPTION OF ENGINEERING PROPERTIES OF MAP UNITS





Concealed bedrock contact (Symbols in parenthesis indicat the concealed bedrock

BIG SANDY

Abandoned oil or gas well