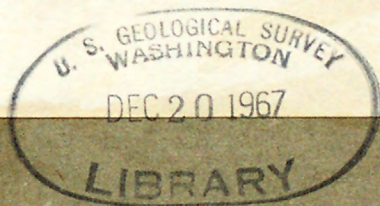


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1. Preliminary geologic map of the western quadrangle, Hampshire and Berkshire Counties, Massachusetts, by G. William Holmes. 1 map, scale 1:24,000, 1 table. Massachusetts Dept. of Public Works, 100 Nashua St., Boston, Mass. 02114; and USGS, 80 Broad St., Boston, Mass. 02110. Material from which copy can be made at private expense is available at the Broad St. address.

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3. Marine phosphorite deposits--economic considerations, by James B. Cathcart. 21 p., 1 table.

4. Phosphate exploration in California, by James B. Cathcart. 19 p., 1 fig., 1 table.

5. Florida type phosphorite deposits--prospects for production, by James B. Cathcart. 27 p., 1 fig., 1 table.

6. Distribution and geologic habitat of phosphate and associated deposit types, by Robert J. Ritz. 35 p., 1 fig., 1 table.



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3. Marine phosphorite deposits--economic considerations, by James B. Cathcart. 21 p., 1 table.
4. Phosphate exploration in Colombia--a case history, by James B. Cathcart. 19 p., 1 fig., 1 table.
5. Florida type phosphorite deposits--origin and techniques for prospecting, by James B. Cathcart. 27 p., 1 fig., 2 tables.
6. Distribution and geologic habitat of marine halite and associated potash deposits, by Robert J. Hite. 35 p., 11 figs.

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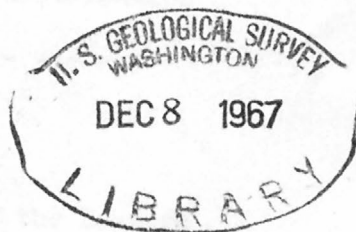
PHOSPHATE EXPLORATION IN COLOMBIA--A CASE HISTORY

*achilder*  
By James B. Cathcart *1917-*

U.S. Geological Survey, Denver, Colorado, U.S.A.

Open-file report

1967



Report prepared for presentation at ECAFE Seminar on Raw Materials,

Bangkok, Thailand, December 4-11, 1967



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1 Phosphate exploration in Colombia--a case history

2 by James B. Cathcart

3 Abstract

4 Exploration for phosphate rock was successfully carried out in  
5- less than 4 months in Colombia. Exploration was based on the theory  
6 that phosphorites are deposited in a miogeosyncline, adjacent to a  
7 foreland or craton, where deposition of clastic material is at a  
8 minimum, and where upwelling currents deposit a suite of rocks  
9 characterized by chert, black shale, carbonate rock, and phosphorite.  
10-

11 Previous work had demonstrated that phosphate was present in  
12 that it  
13 Colombia, was found only in rocks of Cretaceous age, and that the  
14 best phosphate was in rocks of Late Cretaceous age in the Eastern  
15 Cordillera.

16 Field work showed that phosphorite deposits were present only  
17 in the Eastern Cordillera (in the miogeosyncline) in rocks of Late  
18 Cretaceous age. Rocks of Late Cretaceous age in the Central and  
19 Western Cordillera are in the unfavorable eugeosynclinal facies and  
20 brief reconnaissance did not disclose any phosphorite.

21 Emphasis during field work was toward the delineation of litho-  
22 facies within the miogeosyncline, and the field investigation pointed  
23 to the area in Norte de Santander where the best phosphorite was  
24 found. Phosphate deposits are present in both the geosynclinal and  
25 the platform facies in Upper Cretaceous rocks, and are spread through  
a large area in the Eastern Cordillera--from Huila in the south to

Norte de Santander in the north--a distance of more than 600 kilo-  
meters.



## Introduction

Exploration for phosphate rock in Colombia was successfully carried out between August and December 1966. The program was sponsored by the U. S. Agency for International Development and was undertaken in collaboration with the Inventario Minero Nacional of Republic of Colombia.

This report is a summary of the exploration technique and the results of the exploration, and is based on the more complete and detailed report by Cathcart and Zambrano (1966).

Many geologists of the U. S. Geological Survey and the Inventario Minero Nacional of the Republic of Colombia assisted in the work. Without their complete and wholehearted cooperation, it would not have been possible to accomplish the work. I would like to thank Dr. Dario Suescun Gomez, Director, and Drs. Francisco Zambrano, Marino Arce and Pedro Mojica, geologists of the Inventario Minero Nacional, and E. M. Irving, D. E. Ward, D. H. McLaughlin, Jr., and C. M. Tschanz, geologists of the U. S. Geological Survey for their assistance.



Previous work

Exploration for phosphate rock in Colombia began as early as 1942 (Hubach, 1952), but with so little success that in a summary in 1960, Wokittel pointed out that there were no known economic deposits of phosphate in the country. In 1961, Bürgl made a reconnaissance study of the Eastern Cordillera which showed that most of the phosphate is in rocks of Late Cretaceous age and that the samples which contained the highest percentages of  $P_2O_5$  are overwhelmingly concentrated in rocks of the Senonian Stage (Bürgl and Botero, 1962). Slansky (1963) concluded that the Upper Cretaceous was the most favorable stratigraphic interval and that the western part of the Eastern Cordillera was the most favorable structural location for possibly economic phosphorite. From 1963 to 1966, work done by the U.S. Geological Survey and the Inventario Minero Nacional had shown that there were scattered outcrops of phosphorite in rocks of Late Cretaceous age from near Neiva in the Department of Huila in the southern part of the country to near Surata in the Department of Santander in the north. Phosphorite is also known to crop out in Táchira State in Venezuela in rocks of Late Cretaceous age.

## General theory of phosphorite exploration

The theory used in exploration for marine phosphorite deposits is based on the pioneer work of Kazakov (1937), who first proposed that phosphate is deposited in areas of oceanic upwelling. His concept was used by McKelvey and others (1953) to explain the deposition of the phosphorite of the Phosphoria Formation of the Northwestern United States, and was later further developed and expanded to define the lateral and vertical sequence of facies common where phosphate-rich waters upwell onto a shoaling bottom (McKelvey, 1963; McKelvey and others, 1959; Sheldon, 1964a, 1964b). The facies concept defines the unique sequence of rocks deposited by upwelling currents, from a black shale and phosphorite facies in the deeper parts of the basin to a limestone and chert facies containing some phosphate to a sandy facies on the shelf. The sandy facies also contains phosphorite beds, but they are usually not as high-grade nor as abundant or thick as the phosphorite beds in the deeper parts of the basin.

The very large amounts of phosphate in the marine phosphorite deposits indicates that the area of deposition must be open to oceanic circulation over long periods of time. Because deposition of this type of sediment is extremely slow, the site of deposition of the phosphorite must be a long distance from a source of clastic or volcanic sediment. If too much clastic material is dumped into the basin of deposition, the phosphate is so diluted that no economic concentration will occur.



1       The miogeosyncline is the only type of tectonic basin that  
2       meets the requirements of being open to oceanic circulation while  
3       being away from a source of clastic or volcanic sediment. The  
4       study of known economic marine phosphorites shows that most were  
5-      deposited in a miogeosynclinal environment.

#### 6                   General geology of Colombia

7       The dominant physiographic features of Colombia are the Andes  
8       Mountains, which form a broad, northerly-trending belt that extends  
9       the entire length of the country. In Colombia the Andes chain  
10-     consists of three high ranges--the Western, Central, and Eastern  
11     Cordilleras, separated by the Cauca and Magdalena River Valleys.  
12     East of the Eastern Cordillera is the plain of the Llanos--which  
13     includes the Llanos basin and the Guyana shield.

14-     The Western and Central Cordilleras were part of a deep eugeo-  
15-     synclinal basin in Cretaceous time, the Llanos basin and the Guyana  
16     shield form the Craton of foreland of the Cretaceous miogeosyncline,  
17     and only the Eastern Cordillera contains rocks of the Cretaceous  
18     miogeosyncline (Weeks, 1947; Belding, 1955; Burgl, 1961; Campbell,  
19     1965).

Cretaceous rocks of the Eastern Cordillera

According to Bürgl (1961), the floor of the miogeosyncline of the eastern Andean province oscillated during Cretaceous time, leading to repeated marine transgressions, marked in the stratigraphic record by a variety of facies. The stratigraphy of the Cretaceous of the Eastern Cordillera is shown on the correlation chart (Table 1), modified slightly from Bürgl (1961).

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Table 1.--Near here

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Beds of phosphorite are present only in rocks of Late Cretaceous age--in the Guadalupe and La Luna Formations, and field work indicated that they are further restricted to the Galemba Member of the La Luna Formation and the stratigraphically equivalent upper and lower chert units of the Guadalupe Formation. Phosphatic beds are also present in the Villeta Formation and in the Pujamana Member of the La Luna Formation, but the beds contain only sparse phosphate grains and are not likely to be economic.

The Guadalupe Formation consists largely of clastic material--sandstone, siltstone, and siliceous claystone, and some chert, and phosphorite. Black and gray shale and some limestone are present toward the base of the formation. The La Luna Formation consists of black shale, limestone, chert, and phosphorite.



(after Bürgl, 1961)

European stages			Vicinity of Bogota		Santander		Norte de Santander		
Upper	Maestrichtian		Guadalupe	Lower Guaduas		Unir		Catatumbo	
				Tierna sandstone				Mito Juan	
				Dura sandstone				Colon Shale	
				Upper chert					
	Senonian	Campanian		Lower chert	La Luna Formation	La Luna Formation			
		Santonian							
		Coniacian							
	Turonian			La Frontera			Salada Member		
Cenomanian		Exogyra mermoti-Chipaque Lst.							
Lower	Albian		Villeta	Uno sandstone		Salto Limestone		Cogollo	
				Colombiceras horizon		Simiti Shale			
				Aptian		Tablazo Limestone			
	Barremian			Caqueza	Caqueza Sandstone		La Paja Formation		Yuruma
							Rosa Blanca		
							Tambor Formation		
	Neocomian	Hauterivian			Culebra Slate		Rio Negro		
		Valanginian							
Berriasian									
Upper Malin	Tithonian		Sanana schists						

1       Rocks of Cretaceous age underlying the Guadalupe Formation are  
2       dominantly black shale, fine-grained sandstone, and a few limestone  
3       beds--the Villeta and Ca'queza Formations. Black shale in the Villeta  
4       Formation is radioactive and thin beds of sandy limestone in the  
5       formation contain some phosphate (Burgl and Botero, 1962).

6       Cretaceous rocks underlying the La Luna Formation in the Department of  
7       consist of radioactive black shale, limestone, and a basal  
8       sandstone. In the Department of Norte de Santander, the Cogollo and Yuruma Formations  
9       that underlie the La Luna Formation consist of limestone and shale  
10      (table 1). None of the formations contain known phosphate.

11       The rocks of the uppermost part of the Guadalupe are sandstone.  
12      The overlying Guaduas Formation consists of claystone, minor  
13      sandstone, and some beds of coal. The Umir, Colón, Mito Juan and  
14      Catatumbo Formations, overlying the La Luna Formation in the northern  
15      part of the country consist of shale and fine grained sandstone and  
16      are not known to be phosphatic.

#### 17               Paleogeography of the Late Cretaceous

18       One of the most important criteria in the search for phospho-  
19      rite deposits is the paleogeography during that period of geologic  
20      time when phosphorite was being deposited. The evidence in Colombia  
21      was very clear--that marine phosphate was deposited in Late  
22      Cretaceous time. Accordingly, a paleogeographic map of that time  
23      was prepared. Because data are not precise, the map is at best an  
24      approximation, but it is thought to be accurate within the limits of  
25      the scale. The position of the foreland or craton is based on the



1 work of Weeks (1947) and Campbell (1965), and the line between the  
2 eugeosyncline and the miogeosyncline is modified somewhat after  
3 Campbell (1965).  
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1 Available published information indicated the miogeosyncline  
2 in the latitude of Bogota was narrow and that it widened both  
3 to the north and to the south (fig. 1).  
4

5 

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Figure 1.--Paleogeographic and lithofacies map of the Late  
6 Cretaceous of Colombia, showing distribution of phosphorite.  
7 

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#### 8 Lithofacies

9 The facies change in the miogeosyncline, from clastic at the  
10 foreland to black shale-chert-phosphorite in the deeper waters is  
11 an important clue to the location of a possible economic phosphate  
12 deposit. Thus, each phosphorite locality was carefully examined to  
13 determine its position in the miogeosyncline, and its facies.  
14 Preliminary field examination indicated that there were at least  
15 two facies--a clastic or sandy facies represented by the Guadalupe  
16 Formation, and a limestone-black shale-chert facies represented by  
17 the La Luna Formation. Both facies contain beds of phosphorite.  
18 A series of traverses were made across the miogeosyncline to  
19 determine, more accurately, the position of the facies.

20 East of Bogota toward the foreland, the Guadalupe Formation  
21 consists of a thick section of light-colored clastic rocks that  
22 contain only a few thin beds of phosphatic sandstone. The section  
23 at Lake Tota, to the north, is similar. The phosphatic sandstone  
24 beds are about 0.2 to 0.3 meter thick, and contain a maximum of  
25 about 7 percent  $P_2O_5$ . The sandy facies, close to the foreland, is  
thick, and contains only minor phosphate, and none of economic interest.



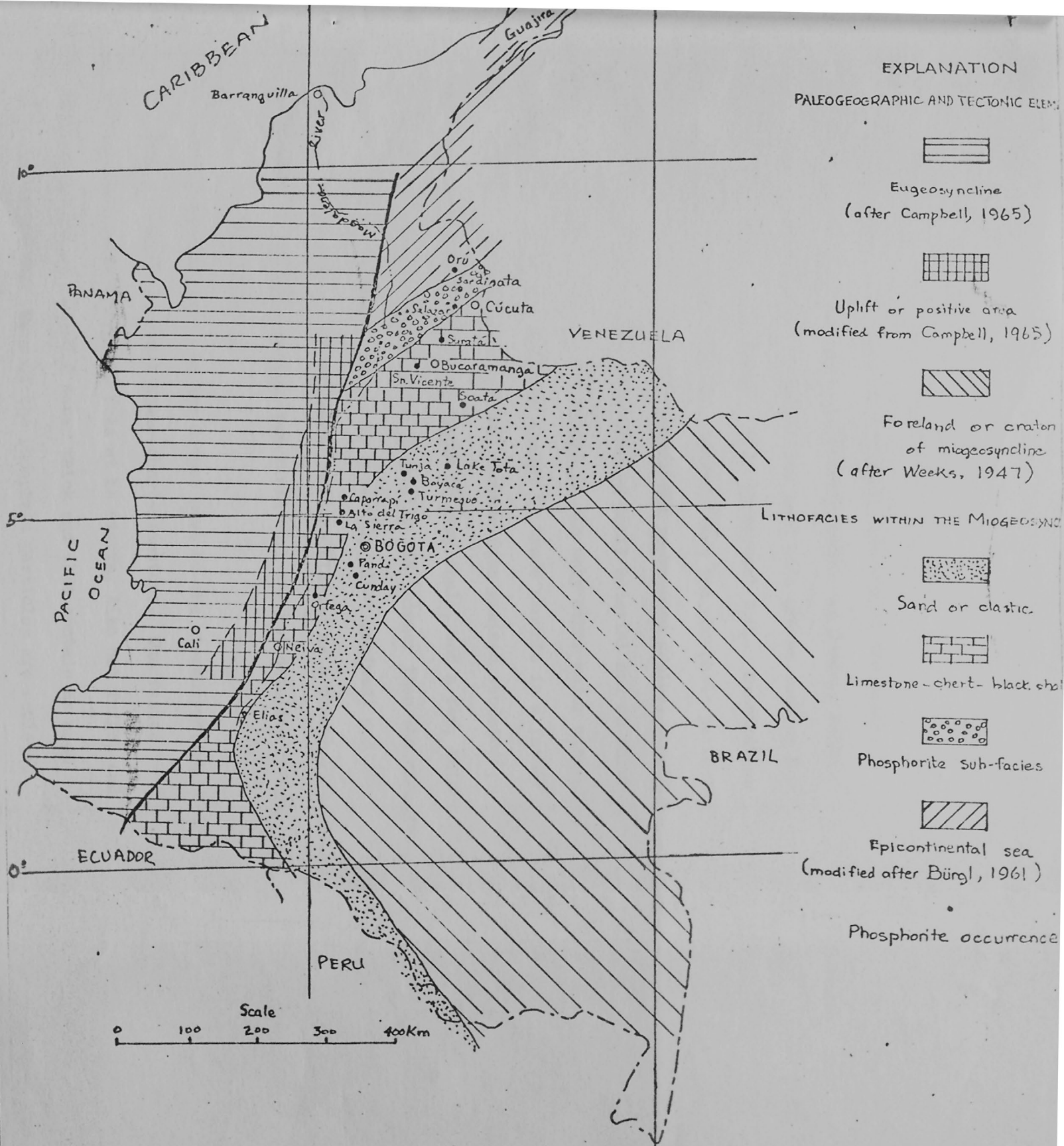


Figure 1. Paleogeographic and lithofacies map of the Late Cretaceous of Colombia, showing distribution of phosphorite.

1 To the west, away from the foreland, and toward the basin,  
2 the sand facies contains good phosphorite beds. At Cunday, Pandi,  
3 and Turmeque', for example, phosphorite sand beds are as much as 2  
4 meters in thickness and contain up to 30 percent  $P_2O_5$ . Examination  
5 of thin sections shows abundant evidence that the phosphate particles  
6 were reworked--i.e., there are broken and re-rounded grains, compound  
7 grains, broken and rounded phosphatized shell material and so forth.

8 Fine-grained (silt size) quartz is the only other major mineral  
9 phase. Calcite is not present, except in trace amounts in a few of  
10 the samples. The associated rocks include fine grained sandstone,  
11 chert or siliceous shale, some black or gray shale, and a few, thin  
12 beds of limestone. The sequence generally, is thinner than the clastic  
13 sequence close to the foreland, is <sup>mostly</sup> ~~generally~~ finer grained, and  
14 contains chert beds. The sequence could be classed as a phosphorite  
15 sub<sup>u</sup>facies of the sand facies, but field work was not detailed enough  
16 to map the sub<sup>u</sup>facies.

17 Still farther west, in the deeper part of the miogeosyncline  
18 the facies changes to black shale-chert-limestone. The facies  
19 contain calcareous phosphate beds at La Sierra, Alto del Trigo, and  
20 Caparrapi', that are 0.1-0.2 meter thick and that contain as much  
21 as 17 percent  $P_2O_5$ . Many of the phosphate particles are replaced  
22 foraminifers, and other phosphatized fossil material--including fish  
23 bones. Very fine-grained, silt-size quartz, is present in varying  
24 amounts, and calcite is the cementing material.  
25

1           The lateral sequence of facies in the latitude of Bogotá  
2 indicated that better deposits of phosphorite might be expected to  
3 the west--toward the Magdalena River. However, a traverse in this  
4 showed that near the Magdalena River rocks of Late Cretaceous age  
5-- consist of a thick series of coarse clastic material that must have  
6 been dumped into the basin from the west, probably from the positive  
7 area (fig. 1). Thus, the miogeosynclinal basin in Late Cretaceous  
8 time was too narrow in its central part for phosphorite deposits to  
9 form.

10--       The evidence that the miogeosyncline widened to the north and  
11 that the best area in which to look for phosphorite was north of  
12 Surata was confirmed by field work. Phosphate beds in the limestone-  
13 chert-black shale facies increase in thickness and in grade northward  
14 and westward. Thus, beds near Soatá are about 0.2 meter thick while  
15-- beds near San Vicente are as much as 1 meter thick and contain more  
16 abundant phosphate pellets than the beds near Soatá (fig. 1).

17       Still farther north, near Salazar and Sardinata, the  
18 phosphorite beds are as much as 2 meters in thickness and contain  
19 as much as 32 percent  $P_2O_5$ . The area is within the chert-limestone-  
20-- black shale facies, and the phosphorite beds are thick and high-  
21 grade. This area has been separated as a phosphorite subfacies of  
22 the limestone-chert-black shale facies.  
23  
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25--



1 North of Sardinata, near Oru, phosphate beds are present in  
2 rocks of Late Cretaceous age, but the beds are thin and low in  
3  $P_2O_5$  content. The phosphate beds are calcareous and are associated  
4 with thin beds of limestone, calcareous silty shale, and chert. The  
5- facies here may represent deposition in shallower water toward  
6 what Bürgl (1961) calls an epicontinental sea facies. The Upper  
7 Cretaceous (La Luna Formation) of the Guajira peninsula (fig. 1) is  
8 in Bürgl's epicontinental sea facies and consists of thin bedded  
9 gray calcareous shale, limestone, and very minor chert. No phospho-  
10- rite beds were found in this area, and the facies is certainly  
11 different from the facies in the deeper part of the miogeosyncline.

12 The miogeosyncline apparently widens to the south (fig. 1), but  
13 much of the area is covered with Tertiary sediments and little is  
14 known of the facies or their distribution. More phosphorite may  
15- remain to be discovered in this area. Much exploration remains to  
16 be done in Colombia--the paleographic and facies map (fig. 1) shows  
17 several unexplored areas and as exploration continues this map  
18 should be refined and made more complete.

#### 19 Field procedures

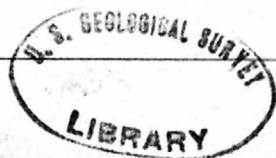
20- The procedure used to locate the phosphorite deposits in  
21 Colombia has been outlined in the previous pages, but it is thought  
22 that a brief, general recapitulation, stressing the actual field  
23 procedures, will be of value.  
24  
25-

1       The first step in a prospecting program is to locate a broad,  
2       generally favorable area in which a marine phosphorite deposit might  
3       be expected. This is the miogeosyncline. Previous geologic  
4       investigation might be expected to at least outline this type of  
5       broad structural environment, and indeed, this was true in Colombia.  
6       The second step is to determine if phosphorite was deposited in the  
7       area of the miogeosyncline, or if the unique suite of rocks  
8       associated with phosphorite--i.e., black shale-chert-carbonate rock--  
9       was deposited in the miogeosyncline. In Colombia, previous work  
10      (Bürgl and Botero, 1962; Slansky, 1963; and unpublished data of the  
11      U.S. Geological Survey and the Inventario Minero Nacional of Colombia)  
12      had clearly demonstrated the presence of phosphate and the associated  
13      suite of black shale-chert-carbonate rock in the miogeosyncline, and  
14      had also demonstrated that the phosphate was restricted to the  
15      Senonian Stage of the Late Cretaceous.

16       Thus, a study of the literature indicated that conditions were,  
17      overall, very favorable for the presence of an economic deposit of  
18      phosphorite in the Eastern Cordillera of Colombia. The value of  
19      the search of the literature cannot be overemphasized. This search,  
20      in Colombia, limited the area that could be expected to yield  
21      potentially economic deposits, and also restricted the stratigraphic interval  
22      that should be carefully examined.

1           Because marine phosphorites contain uranium, they are  
2 significantly higher in radioactivity than other sedimentary rocks.  
3 An examination was made of all available gamma-ray logs of oil wells  
4 in the country, and this examination showed that abnormal  
5- radioactivity was confined to black shales of the Villeta Formation,  
6 and to the black shale-chert-limestone facies of the La Luna and  
7 Guadalupe Formations. Field work showed that the high radioactivity  
8 of the Villeta Formation was caused by uranium in the black shales  
9 not by phosphorite, but showed that phosphorite caused high  
10- radioactivity in the La Luna and Guadalupe Formation. Therefore,  
11 the gamma-ray data also showed the stratigraphic interval that  
12 contained the phosphorite.

13           The next step was to make a lithofacies map of the area of the  
14 miogeosyncline to further restrict the area in which to look.  
15- Accordingly, the first step in the field work was to examine all of  
16 the previously discovered phosphate occurrences, and to plot these  
17 occurrences on a lithofacies map. In addition, general traverses  
18 were made across the miogeosyncline, to aid in constructing the  
19 lithofacies map. This step made it clear that the search should  
20- continue in the northern and southern parts of the country where the  
21 miogeosyncline widened, and where phosphorite might be present.  
22  
23  
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1       The general outcrop areas of rocks of Late Cretaceous age  
2 were known, and the actual discovery of the phosphorite was made  
3 by traversing the roads in the outcrop areas with a hand  
4 scintillation counter. Because marine phosphorites are very  
5- extensive deposits, a search of the road cuts in the area of  
6 outcrop will ordinarily show the presence of the phosphorite, even  
7 if the road net is not extensive.

8       Marine phosphorite beds contain enough uranium ~~so~~ that  
9 radioactivity anomalies will be detected with a portable  
10- scintillometer when the instrument is used in a car at slow speeds.  
11       Areas of anomalous radiation were carefully examined <sup>and</sup> sampled, and  
12 the samples were analyzed for  $P_2O_5$  in the laboratory. The samples  
13 were taken on the basis of hand lens examination of the rock.

14       Sample localities were plotted on the map <sup>and</sup> related to a  
15- lithofacies within the miogeosyncline, and the paleogeographic and  
16 lithofacies map was continually refined during the field work. The  
17 final product is shown in this report as figure 1. Some of the  
18 favorable areas that were not examined because of lack of time  
19 have since been investigated by geologists of the U. S. Geological  
20- Survey and the Inventario Minero Nacional of the Republic of Colom-  
21 bia, and additional phosphorite has been located. The work is  
22 continuing.  
23  
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1       The recognition of phosphate rock in the field may be difficult.  
2       for the inexperienced geologist. The best method of learning to  
3       recognize the material is by actual experience, but failing that,  
4       there are several chemical tests that can be employed. Ammonium  
5       molybdate solution used with nitric acid produces a yellow  
6       precipitate if the rock is phosphatic. The test is delicate and  
7       qualitative. A drop of dilute hydrochloric acid, will produce a  
8       white "bloom" on the rock, if phosphate is present, and phosphorite  
9       outcrops in areas where groundwaters are acid may have a white bloom.  
10      A semi-quantitative colorimetric test has been devised by Shapiro  
11      (1952).  $P_2O_5$  content can be estimated to <sup>within</sup> about 5 percent, but the  
12      test must be used with caution, because the very small sample volume  
13      used may give results that are not representative. Laboratory  
14      analysis of representative samples is necessary to obtain accurate  
15      grade data.

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