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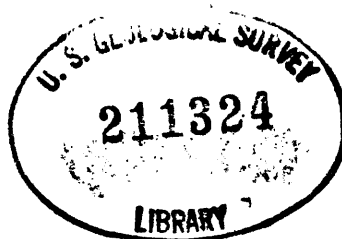
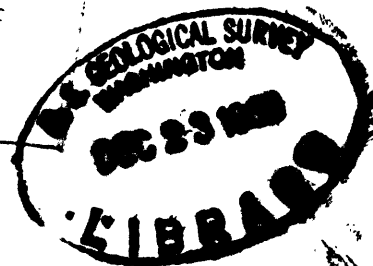
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GEOCHEMICAL PROSPECTING FOR ZINC, LEAD, COPPER, AND
SILVER, LANCASTER VALLEY, SOUTHEASTERN PENNSYLVANIA

By

Jacob Freedman 1911-

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Geochemical prospecting for zinc, lead, copper, and silver,

Lancaster Valley, southeastern Pennsylvania

by

Jacob Freedman

Introduction

Zinc, lead and copper anomalous patterns occur in the soils, rocks, and stream sediments in the vicinity of two old mines and a prospect in the Lancaster Valley, southeastern Pennsylvania. The latter has the maturely dissected rolling terrane of the Piedmont Province, is underlain dominantly by carbonates disrupted by imbricate thrusting and folded into Alpine-like nappes. The poorly exposed rocks in the heavily farmed area are early Paleozoic in age, with local Triassic sedimentary rocks and intrusive diabasic dikes.

Bamford zinc mine

The Bamford area is underlain by Cambrian undifferentiated Chickies-Antietam-Harpers formation, Vintage Dolomite, Kinzers shale and limestone and Ledger Dolomite. These have been thrust northward, folded into a major anticline plunging east and numerous minor folds asymmetric to the north, and further disrupted by high angle faults.

Sphalerite, the dominant mineral, and galena have been oxidized in the shallow zone to smithsonite, hemimorphite, cerussite, anglesite, and aurichalcite (?). These minerals occur in two known and possible other fill-breccias in steeply dipping beds 12 and 14-18 feet thick, averaging 12 percent sphalerite, and in a number of prospects. Anomalous values of soils, rocks and stream sediments range to 30,000 ppm zinc, and 1,000 ppm each of lead and copper. The geochemical dispersion pattern is similar to that of plunging folds suggesting stratigraphic control of ore deposition and extends for 35,000 linear feet with anomalous values at or above 300 ppm Zn, and 13,000 linear feet, at or above 500 ppm Zn. With a threshold of 50 ppm for lead, 219 (22.60 percent) values of 969 samples were equal to or higher than threshold. With a threshold of 50 ppm for copper, 29 (6.67 percent) values of 435 samples are equal to or greater than threshold.

A sample of sphalerite contained 420 ppm Ag equivalent to 12.264 oz. of Ag per ton. Past production has been stated as 357 tons of metallic zinc, but it is possible that as much as 5,000 tons of zinc was produced from over 60,000 tons of ore from depths down to 110 feet.

Pequea Silver Mine

The Pequea area is underlain by asymmetrical, gently doubly plunging folds of the Cambrian Antietam quartzite and schist, the Vintage dolomite, and the Ordovician (?) Conestoga Formation. The basal black carbonaceous phyllite of the Conestoga Formation either disconformably overlies or is a deeper sea facies of the Vintage Formation. The yield by flowage of the former acted as a barrier to solutions in fractures of the latter, which responded to deformation in a brittle manner. Quartz veins with dominantly galena, some chalcopyrite and sphalerite filled fractures and bedding planes. The minerals were oxidized near the surface to cerussite, anglesite, calamine, aurichalcite (?) and vauquelinite (?).

The geochemical distribution pattern in the Pequea area is characterized by isolated anomalous areas. An area in the vicinity of the Pequea silver mine about 7,000' by 5,000' is the largest of the few anomalies that seem to have areal extent. Values range up to 1,700 ppm Pb and 860 ppm Zn in this dominantly lead area. Although thresholds based on frequency curves were chosen at 80 ppm Pb and 130 ppm Zn, contours on geochemical maps nearest these values are 100 ppm Pb and 200 ppm Zn.

Previous studies have indicated as much as 250 ounces of Ag per ton. Two samples of pure galena from the Pequea silver mine gave values of 7,000 ppm Ag (168 oz./ton), and 20,000 ppm (600 oz./ton). Mining in the past proceeded only to a depth of about 75 feet. Past production and reserves can only be estimated.

The Gap Prospect

The Gap mineral prospect is one mile N. 10° E. of the junction of U. S. Route 30 and State Route 41 at Gap. A small, abandoned quarry in Vintage dolomite lies about 100 feet west of an unnamed stream flowing north to Pequea Creek. Satterthwaite and Wright (1958) noted the presence of sulfides. The rock is a white, fine-grained, crystalline dolomite with numerous very thin phyllitic layers composed of muscovite, biotite and locally chlorite. It weathers to a distinctive light or dark ivory hue. Sphalerite, galena, chalcopryite and pyrite occur as scattered euhedral to subhedral grains to 0.1 to 0.2 mm. in diameter and veinlets to 5 mm. thick along the phyllitic cleavage planes. In both thin and polished sections, crystals have grown across the sheared, oriented micaceous minerals. This would imply that the metallic minerals were emplaced after the deformation that formed the cleavage.

The geologic map is by Curran and Gucwa (1967) and is based on distinctive lithologic units. It contrasts with the geologic maps of **Jonas and Stose** (1926) and Satterthwaite and Wright (1958). The geochemical map shows concentrations of zinc in parts per million. The values range from 42 to 4,230 ppm; two other anomalous values are 3,915 and 1,250 ppm. If the threshold for zinc is taken at 150 ppm, 28 or 18.7 percent of the 150 samples are anomalous. The average of all samples is 180 ppm; of all below threshold, 84.8 ppm; and of all anomalous values, 592.6 ppm.

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Stratigraphy of the Gap Area

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Triassic

Diabase dike

C k

Kinzers Formation: interbedded limestones and dolomite with local shale layers

C v

Vintage Formation: Interbedded limestones and dolomite

Vintage Formation: light blue, fine-grained dolomite with some phyllitic layers

C v s

Vintage Formation: sparkling white to dull, fine-grained dolomite with phyllitic layers and impregnated with sulfides

C v c

Vintage Formation: quartz-biotite schist with local mineralization

C a

Antietam Formation: Quartz-biotite schist

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Explanation

Xwc

Wissahickon schist

Oc

Conestoga limestone

Ev

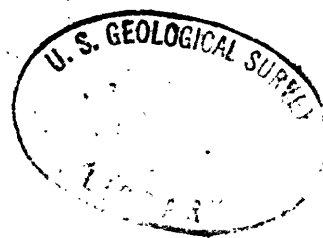
Vintage dolomite

ea

Antietam schist and quartzite,
including Harpers shale

—x—x—x—

Triassic diabase



Ages are omitted because of uncertainties. The above sequence is listed as it appears in the field.

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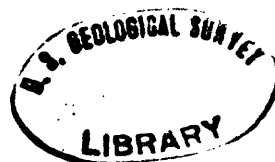
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References Cited

Curran, Steven, and Gucwa, John, 1967, Geology of the Gap area Lancaster County, Pennsylvania, .51p., unpublished senior thesis, Franklin and Marshall College.

Jonas, A. I., and Stose, G. W., 1926, Geology and mineral resources of the New Holland quadrangle, Pennsylvania: Pennsylvania Geol. Survey, 4th Ser., Topog. and Geol. Atlas, no. 178, 40p.

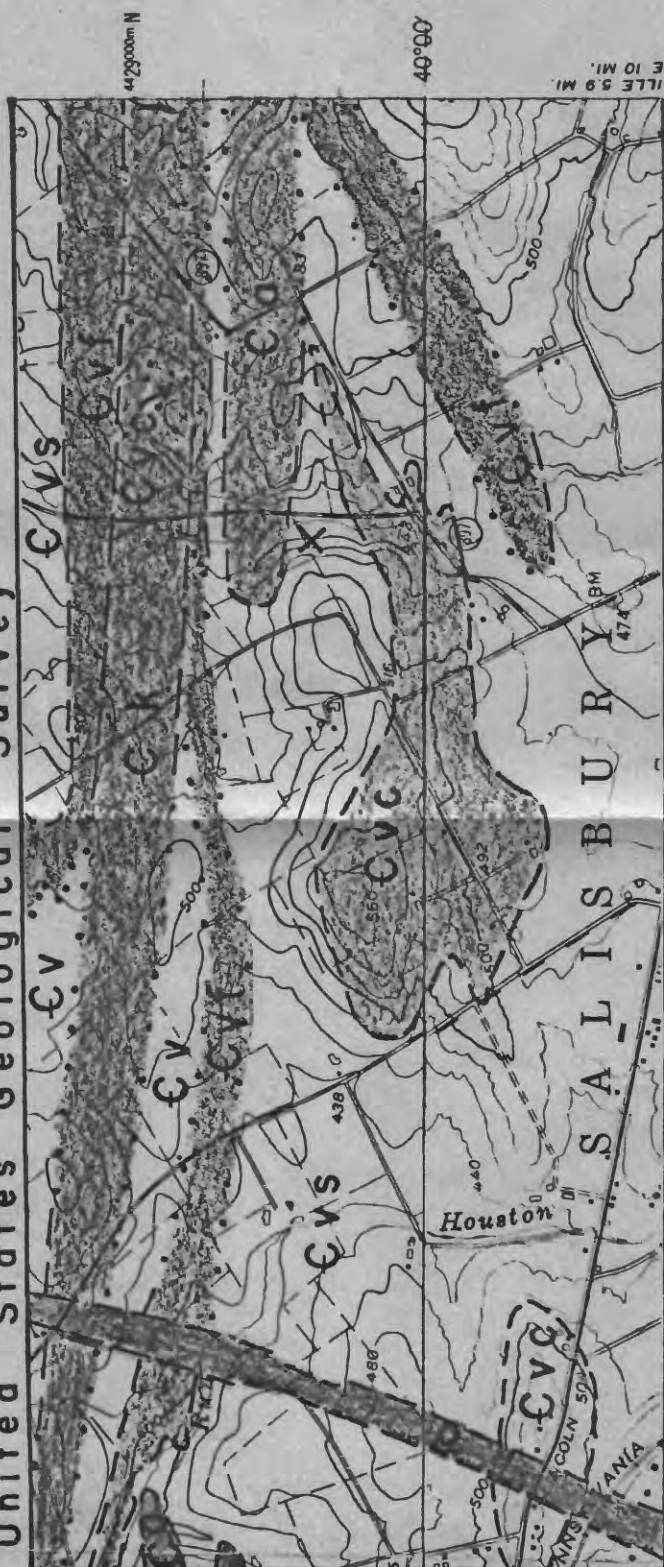
Satterthwaite, W. B., and Wright, R. E., 1958, Geologic report of the Gap area as covered by air photo ANG-4D-171, .35p., unpublished senior thesis, Franklin and Marshall College.



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Geology by
S. Curran and J. Gucwa

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Fig. 3a Geologic Map of the Gap Area, southeastern Penna.

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Fig. 3b. Geochemical Map for Zinc in the Gap Area
 Contoured in PPM

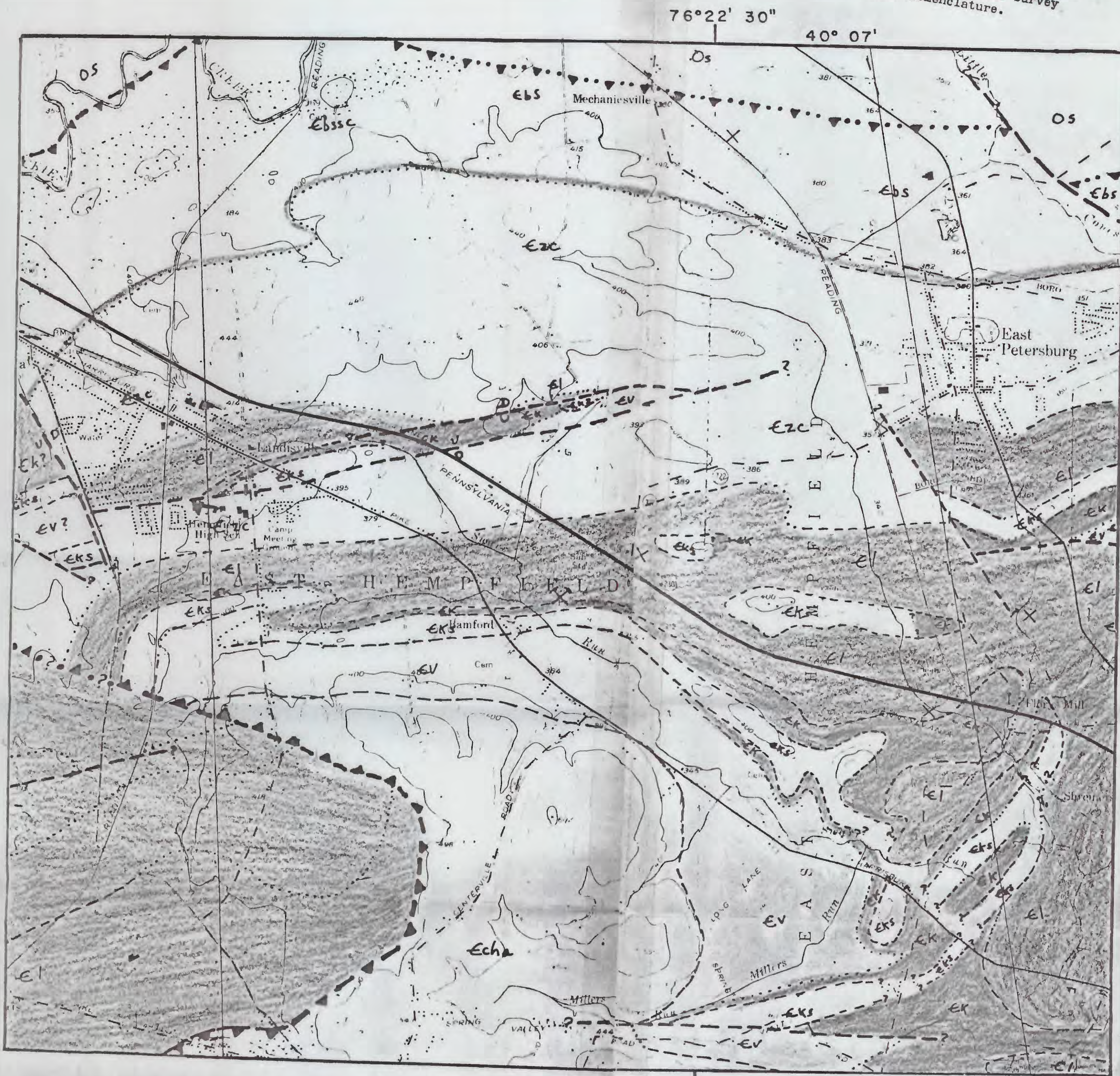
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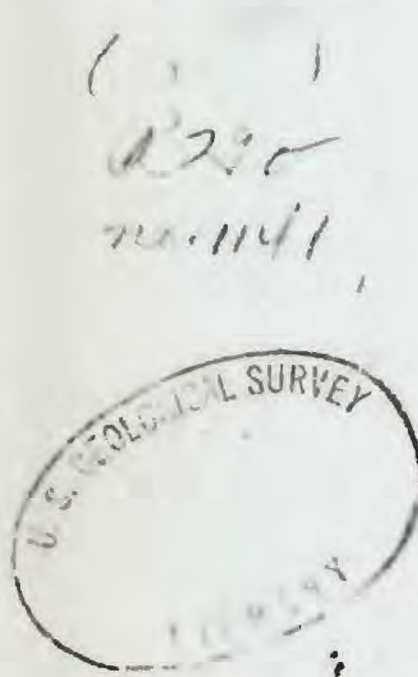
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Scale 1:24,000
Fig. 1a Geologic Map of the Sanford Area, Pa.

Geology by
H. Meisler & A. E. Becker



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EXPLANATION

CONESTOGA FORMATION
Ocs
Gray, finely to coarsely crystalline
limestone; schistose in part;
limestone conglomerate near base

STONEHENGE FORMATION
Os
Gray limestone contains shaly laminae and fossil-
iferous calcarenite beds.

RICHLAND FORMATION
Er
Gray, interbedded limestone and dolomite, contains
beds of fine conglomerate and calcarenite.

MILLBACH FORMATION
Em
White, pinkish-gray, and gray limestone, contains
scattered laminae and interbeds of dolomite

BUFFALO SPRINGS and SNITZ CREEK FORMATIONS
Esc, Ebs, Ebsc
Esc-Snitz Creek - Gray, argillaceous, silty, and sandy dolomite.
Ebs- Buffalo Springs - White, pinkish-gray, and gray interbedded
limestone and dolomite; scattered beds of sandstone.
Ebsc- Combined Buffalo Springs and Snitz Creek Formations, where Snitz
Creek lithology does not occur above the Buffalo Springs;
stippled where argillaceous, silty and sandy dolomite lens occurs
within the combined unit.

ZOOKS CORNER FORMATION
Ezc
Gray, very finely crystalline, dolomite; commonly
silty and sandy; contains some limestone; commonly
cross-laminated and ripple-marked.

LEDGER FORMATION
El
Light-gray, coarsely crystalline, sparkling dolomite

KINZERS FORMATION
Eks, Ek
Eks - Gray, rusty weathering shale
Ek - Limestone commonly containing reticulated silty
and argillaceous laminae; some dark-gray
earthy dolomite.

VINTAGE FORMATION
Ev
Gray, very finely to coarsely crystalline dolomite

CHICKIES, HARPERS, and ANTIETAM FORMATIONS
Ech2
Quartzite, phyllite, slate and schist.

Contact
Dashed where inferred;
dotted where concealed

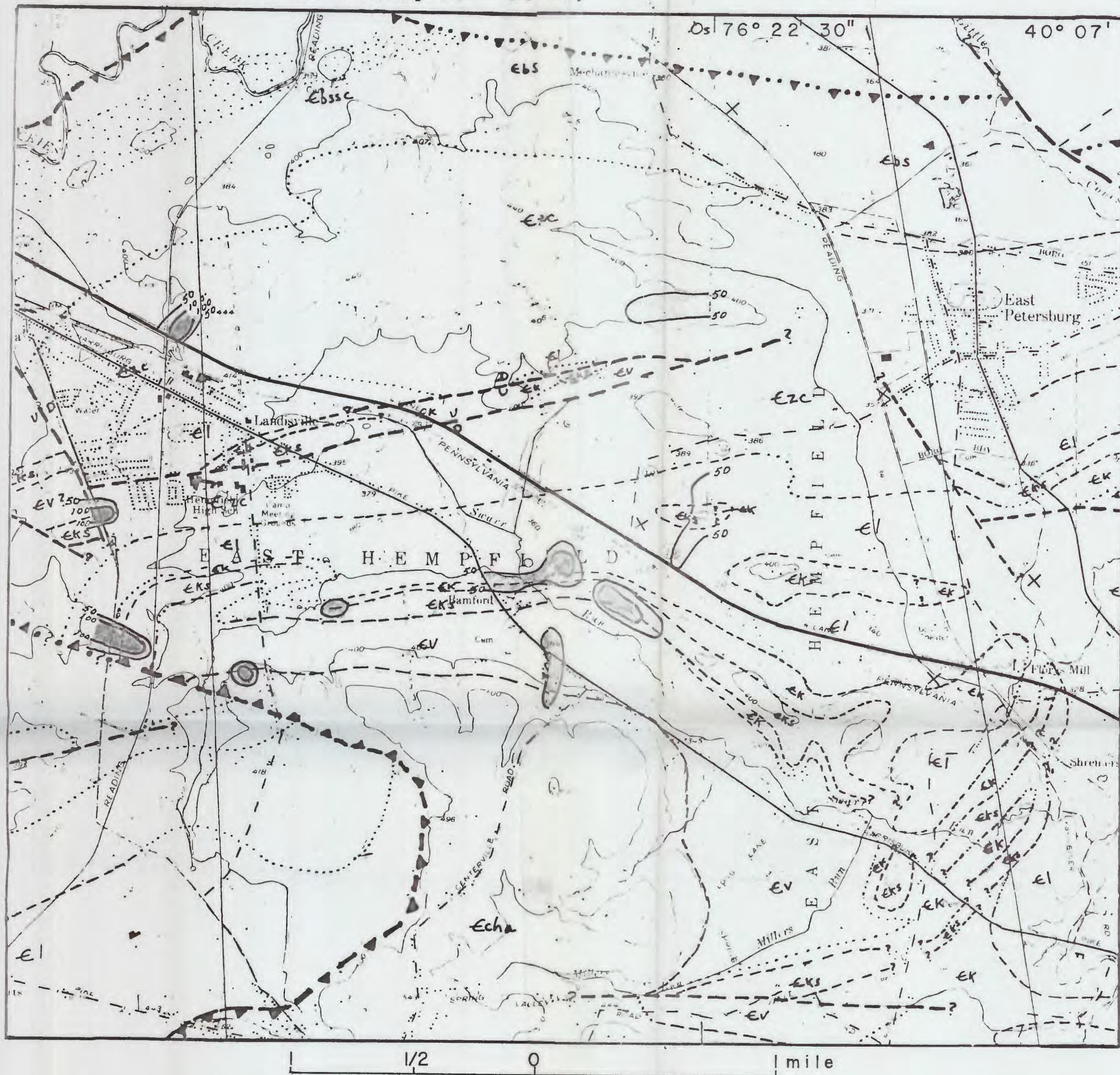
--- U ---
Inferred fault
U, upthrown side;
D, downthrown side

---▲---
Thrust fault
Sawteeth on upper plate
Dashed where inferred;
dotted where concealed

CAMBRIAN

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Fig. 1b Geochemical Map Showing Concentrations of Copper in the Bamford Area, Pa.
in PPM

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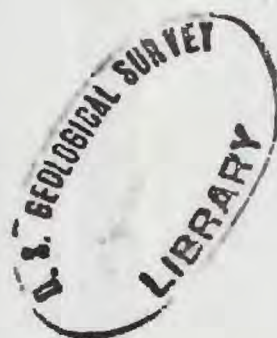
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Fig. 1c. Geochemical Map Showing Concentration of Lead in the Bamford Area, Pa.
in PPM

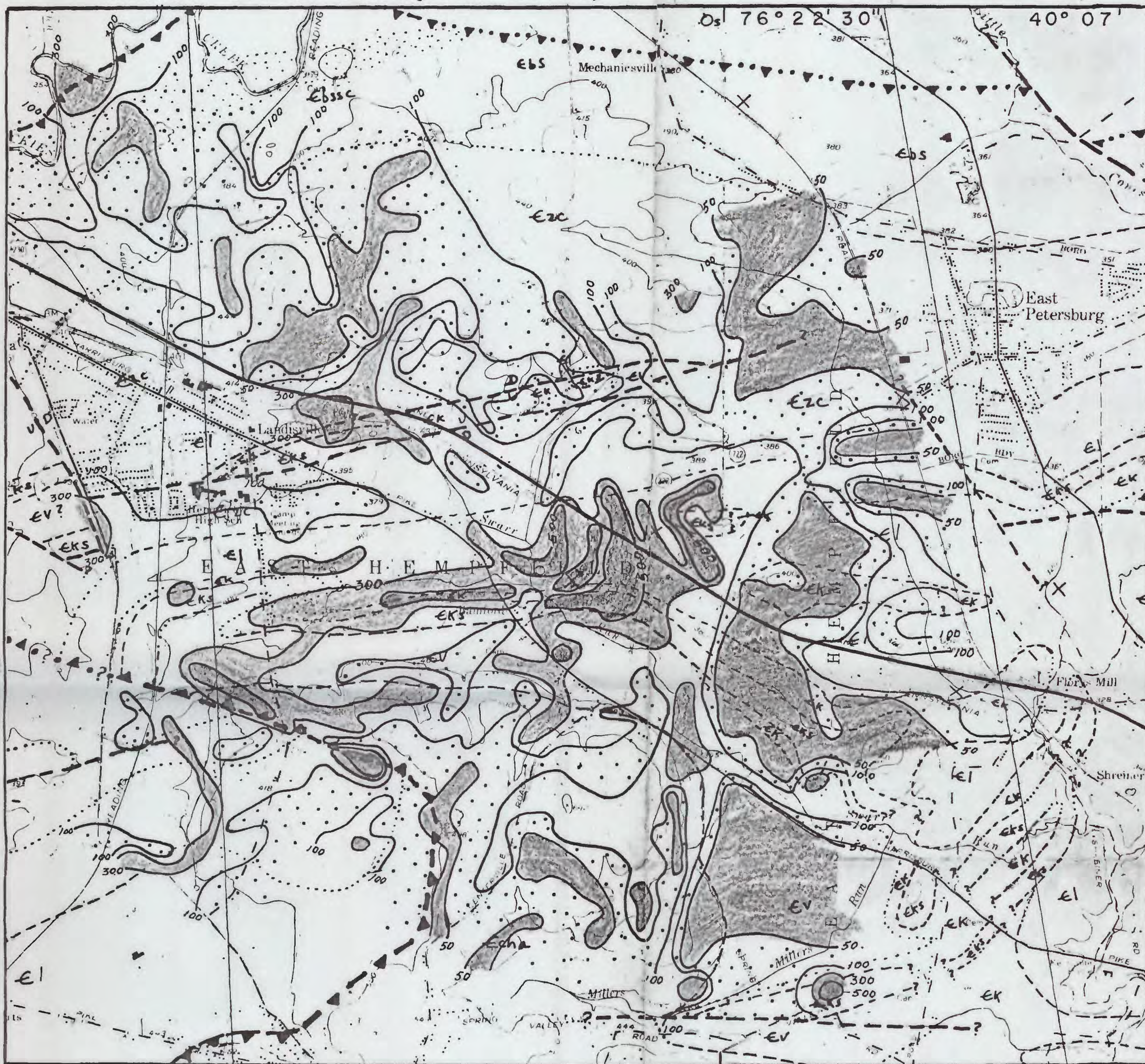
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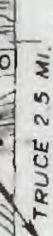
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Fig. 1d. Geochemical Map Showing Concentration of Zinc in the Bamford Area, Pa.
in PPM

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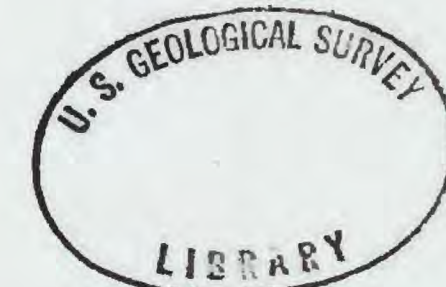


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Fig. 2a. Geologic Map of the "Thrust Belt" in vicinity of Pequea Silver Mine, southeaster Pennsylvania

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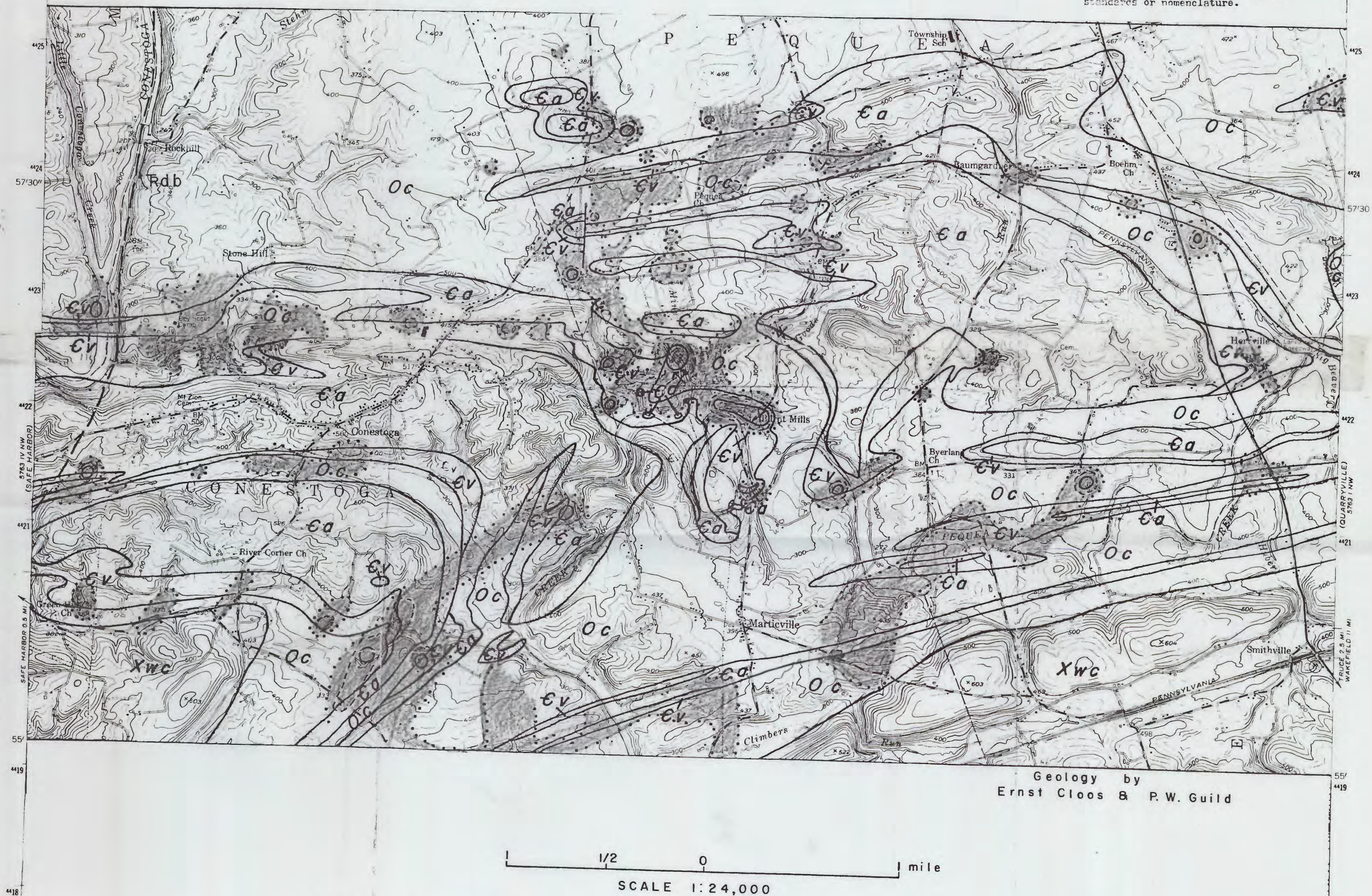


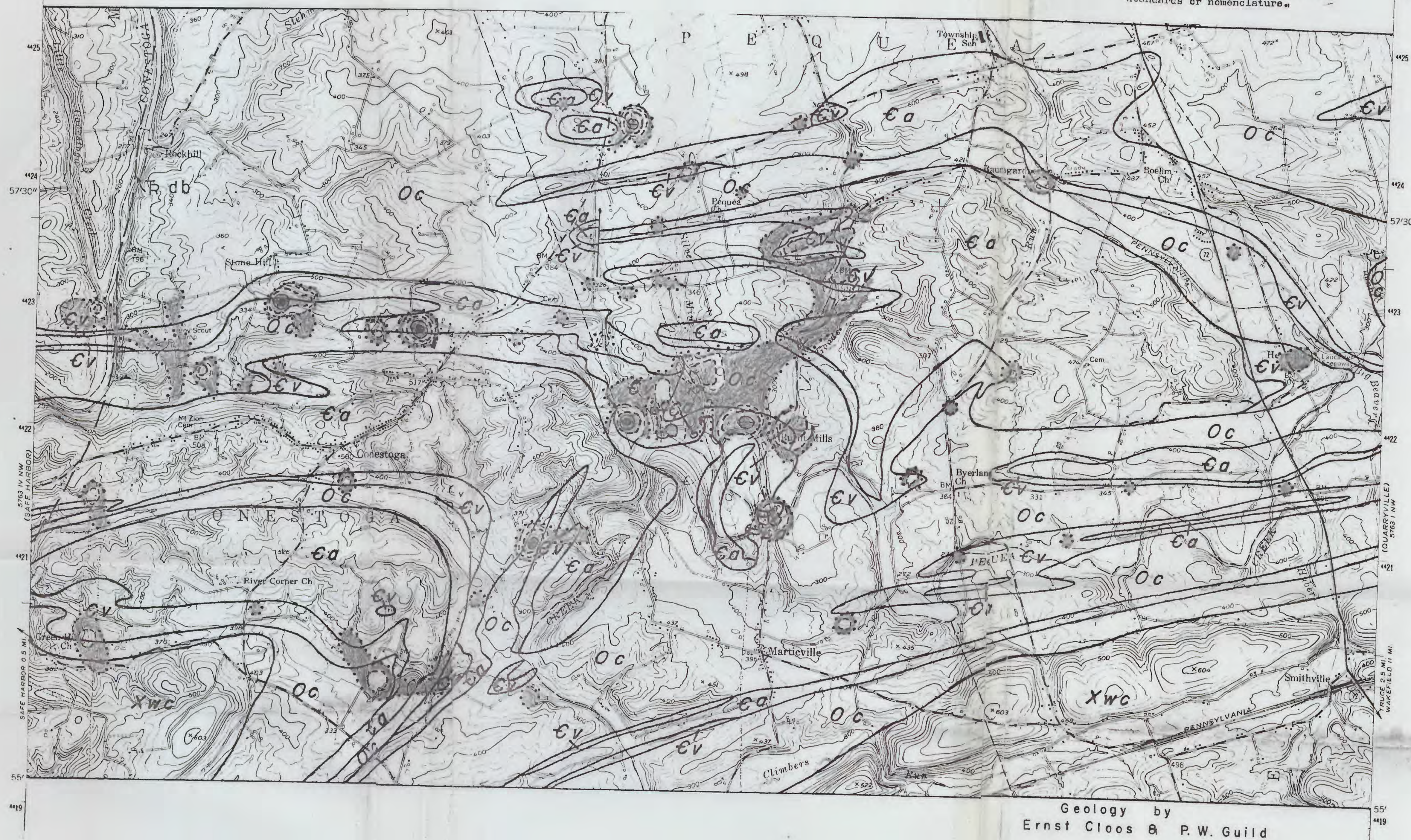
Fig. 2b. Geochemical Map Showing Concentration of Zinc in the "Thrust Belt"

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Fig. 2c. Geochemical Map Showing Concentration of Lead in the "Thrust Belt"

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