Och

SCALE 1:25,000

PRELIMINARY GEOLOGIC MAP OF ELLIPSE II-6-1 AND VICINITY

 $\mathbf{B}\mathbf{y}$

M.J. Grolier

Mercator Projection

23°34′

Ray material

Band, 1-3 km wide, or material slightly bright-

er than adjacent material. This band is the western part of a wider band (5-10 km) mapped in site II P-6 at the scale of 1:100,000

Probably a discontinuous and thin layer of ejecta over mare units m_2 and Cm_2 . This ray material is traceable to the crater Theophilus 320 km south of the site.

The slightly greater brightness of the material perhaps is due to a greater surficial rough-

ness than in surrounding mare. This surficial roughness may be caused by blocks smaller than

face, hundreds of miles from the primary impact crater, is not always clear; ray material

cratering effect on the mare surface to be neg-

Cch

Chain crater material,

Materials of shallow Cc1 and Cc2 craters, gen-

Probably materials of secondary impact craters.
The chain-forming craters mostly are subdued

The thickness of the surficial fragmental layer

probably is thicker than on the adjacent mare.

Some of the crater chains are oriented northeastward. Structural control along one di-rection of the lunar grid perhaps could be

Ccci Cccr

Crater cluster material,

Ccci, crater materials, undivided. Material of shallow contiguous craters, generally 125 meters or more in diameter, and of irregular

outline. Clusters consist of subdued Cc1 largest cluster in the northwestern part of

the map. Patterned ground is conspicuous.
Cccr, crater rim material. Relatively smooth

material, probably consisting of crater

degrees. Superposed craters are less than 75 meters in diameter. Blocks less than 1

meter in diameter may occur near the rim

ejecta blankets; some material may be of second-ary impact origin.

Interpreted engineering properties
Surficial fragmental layer probably thicker than

of the rim, and also because the originally fragmental ejecta was churned by subsequent impacts. For the same reasons, the material of

the surficial fragmental layer on the smaller

landing and trafficability hazards.

ent on the dynamics of impact.

irregular craters may be more finely divided

than on the mare. In the largest irregular cra-

ter, blocks and patterned ground may constitute

The irregular outline of some cluster-forming

craters may be due to deformation along the main

ar mare ridge.

Moon (Ellipse area). Geol. 1:25,000. 1967.

Sheet 2-6-1,

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C. 1

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surficial fragmental layer.

directions of the lunar grid, rather than depend-

md

Mare dome material

Material of subcircular mare dome, about 200

meters in diameter and slightly elongated to the northwest. The material forms a slight

apparently is not directly related to a line-

Material of small volcanic plug, mantled with

Interpreted engineering properties
Surficial fragmental material derived from dome

Very small mare domes either are primary de-

positional (volcanic) features or they might result from differential erosion of mare mate-

fragments of ordinary mare material.

material perhaps is denser and more felsic than

bulge in the surface of mare unit ma. and

in any of the mare units, because of faster mass movement of debris down crater walls and slopes

Materials in impact craters and associated

and of early Copernican age.

checked in the field.

most commonly northeastward. Blocks larger

erally more than 125 meters in diameter, alined

Interpreted engineering properties

stitute trafficability hazards.

Scientific interest

Cm

Mare material

Very gently undulating material, restricted to two subcircular areas, in the northeastern and west-central parts of the map. Cc3 and older

whereas ejecta of Cc4 and younger craters extend onto it. Structural control of this sub-

circular area is suggested by its being wedged between two low mare ridges; lineaments trend

northeastward across it. There is a lack of craters larger than 75 meters in diameter.

Volcanic material at the top of a "volcanic

sink," possibly resembling terrestrial calderas or very large feeder dikes. The apparent struc-

tural control, perhaps along the northeast com-ponent of the lunar grid, suggests a nonrandom

tion of the subcircular area is that it represents a crater, older than the latest episode

of volcanic activity on the mare, and partly filled with mare material. The area in the

northeastern part of the map is perhaps thinly mantled with ray material from the crater

Fragmental material at the surface probably extends to a depth of 4-6 meters. The surface

generally slopes less than 5°. The density of craters 5-75 meters in diameter is less than

The low density of craters larger than 75 me-

ters in diameter may be related to late emplace-

ment of the mare material, or to a faster rate

of crater erosion related to subsidence of

material in caldera or feeder dike.

ter are likely to occur in scattered areas mantled with Theophilus ejecta.

5 percent. Blocks less than 2 meters in diame-

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Crater materials NOTE: Crater materials occupying larger areas, including deposits outside the rim crest, are outlined by geologic contacts; materials occupying smaller areas are assigned numbers only. Characteristics

8

8, crater materials, undivided. Rim crest dirim less than that of surroundings; crater rim crest sharp.

7 Cc7 7, crater materials, undivided. Rim crest diray material; floor material hummocky; cra ter density on rim less than that of surroundings; crater rim crest slightly sub Cc₇, crater materials, undivided. Rim crest diameter 125-400 meters. Includes bright ray material; no blocks resolvable in rim or wall deposits; structure abundant in floor; crater rim crest slightly subdued.

6 Cc 6 Ccr6

6, crater materials, undivided. Rim crest di-ameter 75-125 meters. Includes bright halo material; crater density on rim less than that of surroundings; crater rim crest slightly subdued.

Cc, crater materials, undivided. Rim crest diameter 125-400 meters. Includes bright halo material; blocks moderately abundant

in wall material; strong structure occurs in floor; crater density on rim less than that of surroundings; crater rim crest slightly subdued. Ccr₆, crater rim material. Wider than 125 meters. Abundant blocks throughout for largest craters, blocks occur around rim crest of craters less than 150 meters in diameter; crater density less than that

5 Cc5 Ccr5 5, crater materials, undivided. Rim crest di crater density on rim slightly less than that of surroundings; crater rim crest Cc₅, crater materials, undivided. Rim crest diameter 125-400 meters. Unit not bright; blocks present in wall material; slightly density on rim slightly less than that of surroundings; crater rim slightly subdued but well defined and raised.

Ccr5, crater rim material. Wider than 125 meters. No blocks present; crater density less than that of surroundings. 4 Cc4 Ccr4

4, crater materials, undivided. Rim crest di-ameter 75-125 meters. Unit not bright; crater density on rim same as that of surroundings; associated crater has shape of shallow bowl; crater rim crest slightly raised.

Cc₄, crater materials, undivided. Rim crest diameter 125-400 meters. Unit not bright; no blocks or few blocks present in and around most craters, large block fields in others: concentric structure very associated crater has conical shape or shape of shallow bowl; crater rim crest Ccr4, crater rim material. Wider than 125 meters. Conspicuous block fields present in rim deposits of some craters; crater density slightly less than that of sur-

3 Cc3 Ccr3 3, crater materials, undivided. Rim crest di-ameter 75-125 meters. Associated crater is gentle pan-shaped depression.

Cc₃, crater materials, undivided. Rim crest
diameter 125-400 meters. Few blocks
present in rim deposits, abundant blocks in wall; patterned ground occurs in wall; crater density in rim same as that of surroundings associated crater has shape of shallow bowl; crater rim crest sub-Rim crest diameter 400-800 meters. Few blocks present in floor and wall material; concentric structure slight in rim; crater density on rim same as that of surroundings; associated crater has shape of shallow bowl.

Ccr₃, crater rim material. Wider than 125 me

ters. No block fields present except in higher part of rim; crater density about

the same as that of surroundings. In southwestern part of map, rim material of two craters is wider than the crater

2 Cc₂ Ccr₂ 2, crater materials, undivided. Rim crest di-ameter 75-125 meters. Associated crater is very gentle depression.

Cc₂, crater materials, undivided. Rim crest diameter 125-400 meters. Blocks present in wall material; patterned ground in wall; associated crater is very gentle depression. Rim crest diameter 400-800 meters. Blocks present on rim crest; patterned ground occurs in wall and floor material; crater density on rim same as that of surroundings; associated crater has shape of shallow bowl.

rim material; crater density same as that of surroundings. Cc 1 Ccr1 Cc₁, crater materials, undivided. Rim crest diameter 125-400 meters. No blocks present in wall material; patterned ground occurs in wall and floor; crater density on rim same as that of surroundings; associated crater is pan-shaped depression. Rim crest diameter 400-800 meters. Blocks

Ccr₂, crater rim material. Wider than 125 meters; occurs only around craters more than 400 meters in diameter. No block

fields present except in higher part of

ground occurs in wall and floor; rater density on rim same as that of surroundings; associated crater is pan-shaped depression. Ccr₁, crater rim material. Few blocks present in upper part of rim; crater density same as that of surroundings. Poorly sorted fragments in and around primary and secondary impact craters; as in deposits

of terrestrial impact craters, no more than 2 percent of the fragments are likely to be shock compressed or to contain high-pressure and high-temperature minerals resulting from deposits are tongue shaped (Cca craters in the northwest and southwest parts of the map) or where the rim deposit is much wider than the part of the map). Morphology of craters is progressively modified by mass wasting, tectonism, and superposition of craters and ejecta with increasing age (Cc1, oldest; 8, youngest).

Interpreted engineering properties
All block fields constitute landing and trafficability hazards. Most 8, 7, Cc7, 6, Cc6, 5, and Cc5 craters are likely to have scattered blocks in and around them, up to 1 meter in diameter, and therefore, these craters are trafficability hazards. Depth to cohesive substrate in large subdued craters, Cc3, Cc2, and Ccl, and on the rims of all craters is higher than on the mare.

The rim material of many craters largely consists of fragments of mare material excavated from the mare substratum at the time of impact The analysis of ejecta samples may help elucidate the lithology of the mare substratum, even though the latter cannot be sampled in place because of trafficability hazards down the /walls and floors of craters. Shock-compressed rock fragments in the ejecta probably disintegrate faster than substratum fragments; the around any one crater perhaps may be used as an index of crater age.

> m₄ Mare material

Relatively level material with well-developed patterned ground. Less than 5 percent of the terrain is covered by craters, 5-75 meters in cover about 10 percent of the unit.

Volcanic flows mantled with fragmental layer. Interpreted engineering properties Fragmental layer at the surface extends to a depth of 6-8 meters.

Volcanic flows probably are among the youngest in site II P-6, and do not extend much beyond the area mapped at the 1:25,000 scale. Yet, patterned ground is conspicuous on an apparently gently sloping mare, which suggests a possible time of emplacement as old as late Eratos-

m₂

Mare material data data of the map. Low mare ridges trend northeastward across this unit; they are parallel or subparallel to narrow troughs and low scarps, which account for most of the relief on the mare. The swells and troughs of the patterned ground in and around shallow depressions on the mare contribute to the microrelief; patterned ground is less conspicuous than in unit m4. Less than 5 percent of the terrain in covered by craters, 5-75 meters in diameter. The density of craters in this diameter range is similar in the other mare units. Craters larger than 125 meters are more abundant than in the other three mare units; they cover about 15 percent of the terrain. Lineaments are present and are more abundant than in the other

Volcanic flows, covered in part with ejected debris from Theophilus, and ejecta from super-posed craters. Small indistinct scarps may mark the location of flow fronts.

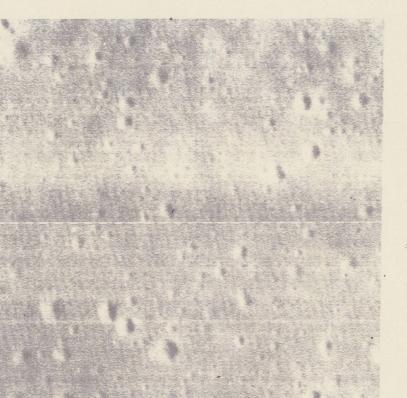
Interpreted engineering properties
Fragmental material at the surface probably than in other areas of the map. Blocks less than 2 meters in diameter are likely to occur in scattered areas mantled with Theophilus

Scientific interest Typical mare material, generally concealed under surficial fragmental layer, even along low scarps and low mare ridges. It is likely to be exposed in place only in the walls of some Cc6 and younger craters.

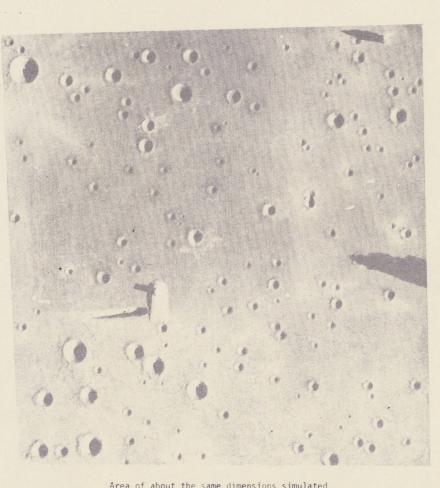
Contact Dashed where uncertain. (Cc_4) Concealed contact Dots show limit of buried unit, where topographically expressed. Buried unit indicated by symbol in (3) Buried small crâter Lineament Very gentle scarp or gentle narrow trough in mare.

> Barbs point downslope. Gentle mare scarp or mare ridge escarpment. Interpretation: May be fault or flow front.

> > **-----**



Area in ellipse II 6-1 approximately 262 meters by 262 meters.



Principal sources of geologic information: Lunar Orbiter II, IV, V photographs

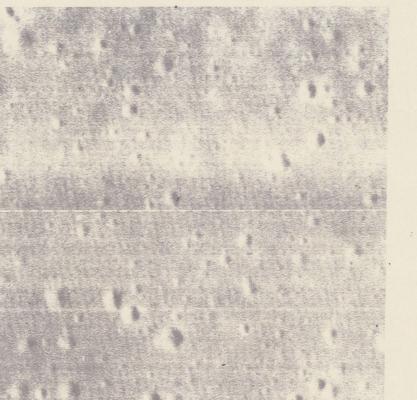
0°34′ 23°34′

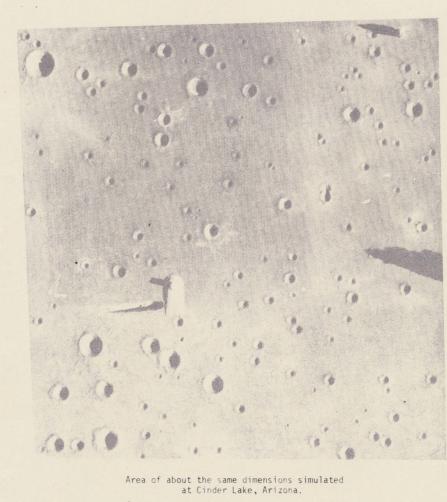
and 5819 taken at U.S. Naval Observatory, Flagstaff, Ariz.

(Langley Research Center, NASA, 1966, 1967); albedo.data from full-Moon plates 5818

Block field Area where blocks larger than 1 meter in diameter

Simulated cratered area Square area outlined by dots on the map was selected for simulation at Cinder Lake, about 10 miles northeast of Flagstaff, Ariz., because of its relatively low crater density and small diameter of its craters. The area in site II P-6 and the simulated area are shown below for comparison at the approximate scale of 1:2,540. Approximate North is at top of pictures.





General Geology Site II P-6, which is in the southwestern part of Mare Tranquillitatis, is a mare area crossed by ridges and rays. Ellipse II-6-1 is north of the equator in the southwestern quadrant of the site. The geologic map described here includes the ellipse and a contiguous band that is 1 km and Two of the four mare units mapped at the scale of 1:100,000 in site II P-6 (Grolier, 1967) occur in ellipse II-6-1. Within site II P-6, mare ridges are resolvable at telescopic resolution and are visible on Lunar Orbiter modextend into the ellipse. Part of a ray from Theophilus, a northwestward across the eastern part of the ellipse. Four mare units (m_1 , m_2 , m_3 , and m_4) are shown on the geologic map of site II P-6 at the 1:100,000 scale (Grolier, Two of these (m_0 and m_A) are shown on the 1:25.000 surficial fragmental layer. The other mare unit (Cm) is exposed in subcircular areas, which are surrounded disconbeen distinguished from mare units mg and ma at the 1:100,000 The subcircular areas are perhaps "volcanic sinks" mantled with a surficial fragmental layer but otherwise resembling small terrestrial calderas or very large feeder dikes. Alternatively, they may be remnants of old and rather large impact craters, partly filled with mare material (Cm) which may have been emplaced after the Cc_2 and Cc_3 craters formed but certainly before the Cc4 and Cc5 craters formed. (Rim material of Cc2 and Cc3 craters appears to be truncated by mare unit Cm whereas that of Cc4 and younger craters extends onto the mare unit.) This mare material (Cm) is relatively level, and its surface is pitted by craters that are mostly less than 75 meters in diameter. The materials of all craters more than 75 meters in diameter are indicated on the map, and relative ages are assigned on the basis of morphology and crater diameter, as shown in figure 1. All craters in the map area are of Copernican age: Cc1, oldest; Cc8, youngest. Most of the craters are probably due to impact; however, some may be of internal origin. Crater materials are subjected to cratering after their formation, and a surficial fragmental layer develops on them as on the mare. Impact craters probably consist of loose, fragmental mare material. Some small particles may of the mineral constituents of the fragments probably are

23°30′

Controlled base prepared by Army Map Service, Corps of

SITE LOCATION DIAGRAM KEYED TO LAC

37/38/39/40 41 42 43 44 43

55 56 57 58 59 60 61 62 63

73 74 75 76 77 78 79 80 81

91 92 93 94 95 96 97 98 99

70°W 50°W 30°W 10°W 10°E 30°E 50°E 70°E 32°S

* 0

Engineers, U.S Army, Washington, D.C. 20315

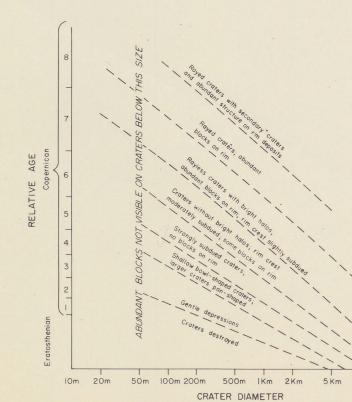


Figure 1.--Assumed relations among crater diameter (from rim Width of intervals on ordinate carries no implication as to lengths of intervals of lunar geologic time.

Some craters with a convex rim and rim deposits whose width equals or exceeds that of the central depression may be volcanic. Such craters occur in the extreme southwest corner of the map area. If they are volcanic, they probably Two sets of lineaments trend northeast and northwest across the mare; a third set, in the northwestern part of the map area, trends almost northward. Two very low mare ridges and a prominent escarpment trending almost eastward make up, with the lineaments, the structural elements of the ellipse.

The mare material in Mare Tranquillitatis is of Imbrian and Mare is and Wilhelm 1067. Notes of the entered in size in the mare material in the material in age (Morris and Wilhelms, 1967). None of the craters in site II P-6 mapped at the 1:100,000 scale are older than Eratosthenian. The oldest, Sabine D and E, might be older than the mare material which surround them, but there is no morphologic evidence within the site to assign with certainty a lower age limit to the main mare units. An age ranging from late Imbrian to early Eratosthenian is being suggested for Mare Tranquillitatis by two geologists (H. Å. Pohn and T. W. Offield, 1967, oral commun.) who are systematically analyzing the relative ages of large lunar craters.
Within and around ellipse II-6-1, the ejecta of craters more than 75 meters in diameter covers about 25 percent or more of the mare surface. Five to seven percent of the rest of the mare is covered by materials of craters in the 5-75 meter range, and Ranger and Surveyor photographs have shown that the cratering process is manifested by craters whose smallest size is limited only by photographic resolution.

Because all resolvable craters at the 1:25,000 map scale are Copernican, the surficial material in ellipse II-6-1--the regolith that mantles mare and crater materials -- is probably General alinement of craters on and along low mare ridges in the east-central and southern parts of the area suggests structural control, and many craters are alined along lineaments in the northeastern part of the area. However, more evidence is needed to state with certainty that these craters are structurally controlled.

Engineering Geology Regional relief in the map area probably does not exceed 200 meters; the terrain generally slopes gently eastward. Local relief occurs primarily in and around craters, and to a lesser degree between low mare ridges and troughs. The depth-to-diameter ratio of most fresh craters of impact origin is about 1:4. The largest fresh crater, 375 meters in diameter, is a Cc_6 crater in the southwest part of the map area. It is probably about 90 meters deep. Some subdue craters (Cc1-Cc3) are much larger and are probably more than 100 meters deep. The engineering properties of the lunar surface in ellipse II-6-1 can be evaluated from data returned by Surveyor V, which landed in the northwestern part of site II P-6, about 25 km northwest of the ellipse, in mare unit mo (Grolier, 1967) The spacecraft was not equipped with strain gages or with a surface sampler. However, the interaction of its footpads with the lunar surface was similar to that of Surveyor I (site III P-12) and Surveyor III (site III P-9), suggesting that the surficial fragmental layer is texturally similar to that consists of weakly cohesive, very slightly compressible, granular material whose engineering properties are as follows (Jaffe and others, 1967, p. 3, 5): Cohesion, 1.5 x 10^5 to 7 x 10^5 dynes per cm²; angle of internal friction, 35° to 40° ; and static bearing capacity, 2 x 10^5 to 6 x 10^5 dynes per cm². The surface roughness of the ellipse, in terms of rocks lying on or protruding through the surface, appears comparable to that of the Surveyor V landing site. Where mare units m_2 and Cm are modified by ray material from Theophilus, surface roughness may be greater than around Surveyor V. Mare unit Cm is less densely cratered than mare unit m₂ material at the The thickness of the surficial fragmental layer in the ellipse cannot be estimated by the method of Oberbeck and which internal structure is present. However, craters up to 20 meters in diameter within 200 meters of the Surveyor V spacecraft are surrounded with fields of blocks and boulders. Assuming a depth-to-diameter ratio of 1 to 4, a "ratio charblocks on the rims of these craters indicated that the depth to blocky or coherent material is locally not greater than about 5 meters" (Shoemaker and others, 1967, p. 645, 646). Probably, the older the mare material, the thicker the surfi cial fragmental layer, because of a more prolonged churning of the mare surface by successive impacts. The surficial fragmental layer probably is not much more than 5 to 8 meters thick, at least in the oldest mare unit m2, which includes the Surveyor V landing site (Grolier, 1967).

One of the processes by which lunar craters become degraded is mass wasting. Some of the craters in ellipse II-6-1 are more than 100 meters deep and were probably even

deeper when first formed. A layer of debris derived from the walls and structural terraces accumulates on the floor

impacts. Thus, the combined colluvial layer and overlying

surficial fragmental layer on the crater floor is probably

much thicker within large subdued Cc1 and Cc2 craters than on the mare, perhaps by at least one order of magnitude.

The rim material of most craters probably consists of multiple layers (underlying mare surficial layer, impact or volcanic ejecta, and the surficial layer developed on the ejecta). When fresh, these layers are probably thicker than the surficial fragmental layer on the mare, but with time their thickness is reduced by mass wasting.

The chief engineering hazards are craters and block diameter are outlined on the map, but the number of craters from 5 to 75 meters alone exceeds the craters mapped by at least one order of magnitude (Natl. Aeronautics and Space Adm., 1967, p. 49-50). All block fields containing blocks larger than 2 meters in diameter are outlined on the map; they all are restricted to parts of craters or crater rims. Smaller blocks probably occur around many smaller craters also in the area covered by discontinuous ray material from trafficability hazards. Another such hazard lies in the possible presence of numerous small, fault-controlled scarps that cannot be resolved on Orbiter photographs.

A potential hazard to lunar construction, but perhaps not a landing or trafficability hazard, is permafrost. Water (1967, p. 55) recently suggested that a permafrost layer may occur 1 or 2 meters below the lunar surface. Lunar permafrost might form by the freezing of volcanic water or other endo-geneous volatile substances. The surficial fragmental layer in ellipse II-6-1 would be thick enough for the formation of

Scientific Possibilities

The primary scientific objective ought to be the sampling of the surficial fragmental layer and the rocks protruding through it or lying on it. The diversity in dimension, texture, and probably in composition of these rocks is shown on many of the Surveyor photographs. The youngest craters (Cc_6 to Cc_8), which are smaller than 50 meters in diameter and therefore not outlined on the map, will yield much information on impact debris and cohesive material derived from the mare substratum. The block fields and the blocks within the fields around small young craters probably are relatively small--a distinct advantage The rim materials of all craters Cc4 or older display the same density of superposed craters as the adjacent mare. The implication is that the steady state in the impacting process has not yet been reached for craters Cc_5 or younger. Therefore, the stability of rock aggregates and of rock compressed by shock could be estimated from the percentage of rock types systematically collected on Cc5 craters or younger. Shock-compressed rock derived from terrestrial impact craters is more fragile and disintegrates at a faster rate than coherent fragments derived from bedrock (D. Roddy, oral commun.). By analogy, shock-compressed rock on the Moon may also disinte-Study of high-pressure and high-temperature minerals produced by shock has shown that some are stable for very long periods of geologic time, even when no longer in equilibrium with their environment (E. C. T. Chao, oral commun.). Some minerals may be vitrified by shock; conceivably, different mineral glasses have different rates of devitrification. Sampling of crater materials of different ages may yield information on devitrification rates. Likewise, the outer skin of surficial lunar rocks may supply information on the residue Sampling of large blocks, 2 meters or more in diameter, if feasible, may provide stratigraphic evidence if they were excavated from a greater depth in the lunar substratum than ejecta of smaller size. Comparison of the morphology of big blocks on the rim of large subdued Cc1 and Cc2 craters (as in the northern part of the area) with the morphology of blocks on the rims of large young ${\tt Cc}_6$ craters (as in the southwest part of the area) may lead to better understanding of Material deep under overhangs or beneath large blocks Photography of the landscape traversed and of the rocks sampled is an integral part of the geologic investigation. REFERENCES

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