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UNITED STATES
DEPARTMENT OF THE INTERIOR
GEOLOGICAL SURVEY

MODEL ANALYSIS OF THE
IMPACT ON GROUND-WATER
CONDITIONS OF THE
MUSKEGON COUNTY
WASTEWATER DISPOSAL
SYSTEM, MICHIGAN



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Prepared in Cooperation with the Geology Division,
Michigan Department of Natural Resources
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MODEL ANALYSIS OF THE IMPACT ON GROUND-WATER

CONDITIONS OF THE MUSKEGON COUNTY WASTEWATER DISPOSAL SYSTEM, MICHIGAN

By Michael G. McDonald and William B. Fleck

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Okemos, Michigan

January 1978

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FACTORS FOR CONVERTING ENGLISH UNITS TO INTERNATIONAL

SYSTEM (SI) UNITS

<u>Multiply English Units</u>	<u>By</u>	<u>To Obtain SI Units</u>
inches (in)	2.54	centimeters (cm)
feet (ft)	0.3048	meters (m)
miles (mi)	1.609	kilometers (km)
acres	0.4047	hectares (ha)
square miles (mi ²)	2.590	square kilometers (km ²)
cubic feet per second	.02832	cubic meters per second (m ³ /s)
gallons per day per square foot ((gal/d)/ft ²)	0.0407	meters per day (m/d)
square feet per day (ft ² /d)	0.0929	square meters per day (m ² /d)

Glossary

Aquifer.--A formation, group of formations, or part of a formation that contains sufficient saturated permeable material to yield significant quantities of water to wells and springs.

Artesian water.--Ground water that is under sufficient pressure to rise above the level at which it is encountered by a well, but which does not necessarily rise to or above the surface of the ground.

Base flow.--The discharge entering stream channels as effluent from the ground-water or other delayed sources; sustained or fair weather flow of streams.

Bedrock.--In this report, designates the consolidated rock underlying the glacial deposits.

Confining bed.--A body of relatively impermeable material stratigraphically adjacent to one or more aquifers.

Evapotranspiration.--Water withdrawn from a land area by direct evaporation from water surfaces and moist soil and by plant transpiration, no attempt being made to distinguish between the two.

Head.--The height of the water surface in a cased well at a point in an aquifer. Also called static head.

Hydraulic conductivity.--Ability of a porous medium to transmit water.

Hydrograph.--A graph showing stage, flow, levels or other property of water with respect to time.

Potentiometric surface.--A surface which represents the levels to which water will rise in tightly cased wells.

Recharge.--Comprises the processes by which water is absorbed and is added to the zone of saturation. Also, the quantity of water added to the zone of saturation.

Specific yield.--The ratio of the volume of water which the saturated medium will yield by gravity to the volume of the porous medium.

Transmissivity.--The ability of aquifer material to transmit water. It is equal to the product of hydraulic conductivity and thickness.

Water table.--The upper surface of the zone of saturation, except where the surface is formed by an impermeable body.

Water-table aquifer.--One in which the upper surface of the body of water is a water table.

MODEL ANALYSIS OF THE IMPACT ON GROUND-WATER CONDITIONS OF THE
MUSKEGON COUNTY WASTEWATER DISPOSAL SYSTEM, MICHIGAN

By

Michael G. McDonald and William B. Fleck

Okemos, Michigan

ABSTRACT

A digital model was developed to study the impact on ground-water conditions of the Muskegon County wastewater disposal system. At the disposal site, wastewater is stored in two 850-acre (344-ha) lagoons and then spray-irrigated on crop land. About 70 miles (105 km) of drainage tile, which underlies the irrigated land, has caused the water table to be lowered substantially. The decline in water levels has been partially offset by irrigation and leakage from the lagoons; at some places the water table is higher than it was prior to construction. Predictive simulations by the model were used to study the effects of varying tile drainage, amount of irrigation water applied, lagoon leakage, and natural ground-water recharge. If the effectiveness of the tile to collect drainage is reduced by 75 percent, large areas within the wastewater site would become waterlogged. However, the effect outside of the wastewater site would be negligible.

INTRODUCTION

This report summarizes the results of an investigation of ground-water conditions from 1974 to 1976 at the Muskegon County, Michigan spray-irrigation wastewater disposal facility. The principle purpose of the investigation was to develop a mathematical model that, when used on a digital computer, would simulate the hydrologic system. Figure 1 shows the location and boundaries of the modeled area and the location of the wastewater facility. Geologic data for development and calibration of the model were obtained during the course of the study from a network of observation wells in and surrounding the treatment facility.

As additional information becomes available through the operation of the system, and as data on ground-water fluctuations accumulates, an improvement in the predictive capability of the model will be possible.

COOPERATION

The investigation was undertaken in cooperation with the Geology Division, Michigan Department of Natural Resources. In addition to sharing in the cost of the study, the Geology Division made available well records and other data from their files and provided assistance in field work at the treatment site.

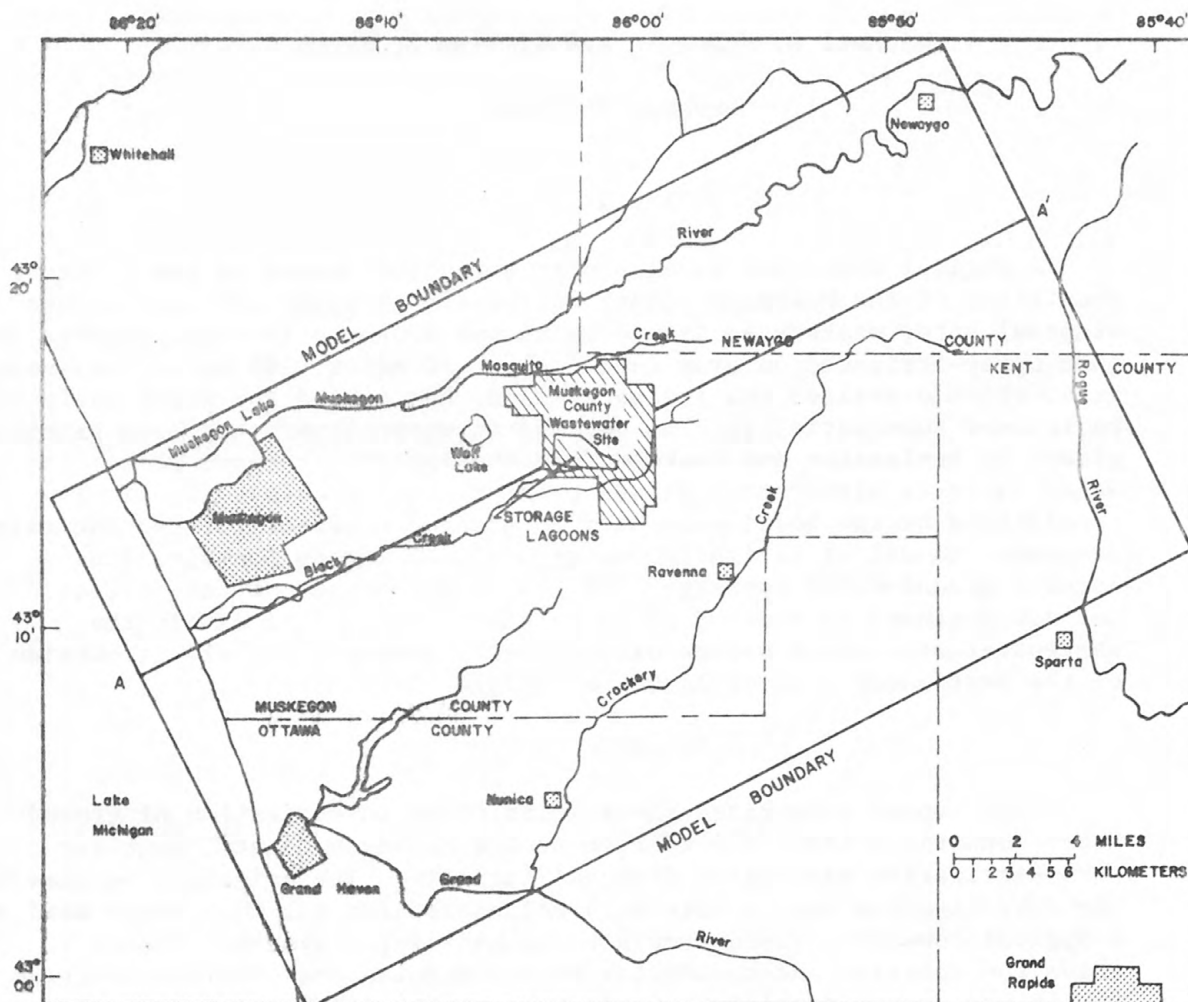


Figure 1. Location of study area. Line A-A' is cross section shown in figure 3.

ACKNOWLEDGEMENTS

The authors are indebted to all individuals who provided assistance in the course of this study, and are especially indebted to Dr. Y. A. Demirjian, Director of the Muskegon County Wastewater Management System, and Gordon D. Bennett, Hydrologist, U.S. Geological Survey. Their cooperation and personal assistance facilitated work at all stages of model development.

DESCRIPTION OF WASTEWATER DISPOSAL SYSTEM

The wastewater system has a design capacity of 42 M gal/d (1.8 m³/s) (Bauer Engineering, Inc., 1973). Five municipalities and two industries currently pipe 29 M gal/d (1.3 m³/s) to the site. Wastewater is aerated, chlorinated, and then used to irrigate corn. The treated water is sprayed on the land for 8 months; during the winter it is stored in two large lagoons, each having a surface area of 850 acres (344 ha). The bottoms of the lagoons are about 10 ft (3 m) above the preconstruction altitude of the water table. Leakage from the lagoons is intercepted by ditches that surround about 90 percent of the perimeter of lagoon area. Water in these seepage ditches is either pumped back to the lagoons or to adjacent streams. From the lagoons, water is pumped to center pivot rigs capable of irrigating crops at a rate of 3.5 in (8 cm) per week. There are 54 circles irrigating an area of 5400 acres (2200 ha) (fig. 2). Most circles are drained by corrugated polyethylene tiles having diameters of 6 to 10 in (15 to 25 cm). The tile is perforated with 0.06- by 1.5-in (0.2- by 3.8-cm) slots and is encased in a 0.02 in (0.5 mm) mesh fiberglass fabric. Drainage tile lines are generally set 5 to 8 feet (1.5 to 2.5 m) below land surface and are spaced at 500 ft (150 m) intervals. They are generally just below the water table, except in the northwest corner of the site, where the tile lines are few and above the water. Total length of the tile lines is 70 mi (105 km). Drainage tiles are connected to concrete collector tiles (fig. 2) that discharge to drainage ditches. In addition to tile, there are 30 wells (designated as pumped wells on fig. 2) that are used to control ground-water levels. Seven of the wells along the northwest edge of the lagoons are pumped to reduce ground-water mounding caused by lagoon leakage. The rest of the wells are used to control mounding caused by irrigation. This system of tiles, discharge wells, seepage ditches, and drainage ditches, was designed to lower ground-water levels in the project area.

GEOLOGY

Glacial Deposits

The western half of the modeled area (fig. 1) is underlain by sediments deposited in Pleistocene lakes (fig. 3) (Martin, 1955). The composition of the sediments varies from tight clay to fine gravel. Generally, the upper 20 to 80 ft (6 to 24 m) is well-sorted stratified sand, interlayered with fine gravel and silt. Underlying the sand and gravel are beds of relatively impermeable silt and clay. Topography in the western half is flat and poorly drained; recharge to the ground-water body is high.

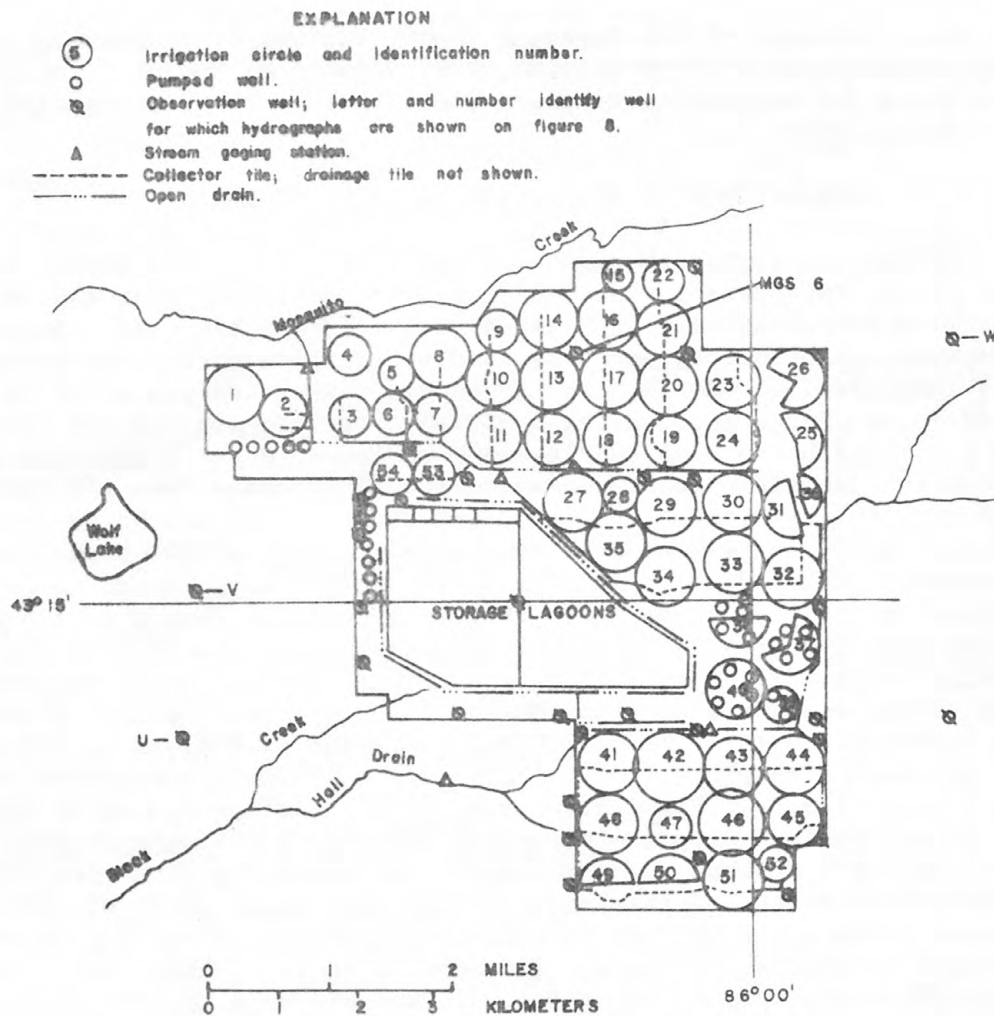


Figure 2. Muskegon County wastewater disposal system.

Most of the eastern part of the area is underlain by morainic deposits, poorly sorted silt and clay containing lenses of sand and gravel. Thin sand and gravel beds occur along some stream valleys. The thickness of the unconsolidated deposits ranges from 100 to 500 ft (30 to 165 m). Under the wastewater site these deposits are approximately 275 ft (85 m) thick. Undulating topography, more extensive stream development, and low recharge to the ground-water body characterize the eastern part.

Bedrock Deposits

Underlying the Pleistocene deposits are consolidated sedimentary rocks of Mississippian age (figs. 3, 4). They include the Bayport Limestone, Michigan Formation, Marshall Formation, and Coldwater Shale (Martin, 1936). The Bayport Limestone is predominantly limestone,

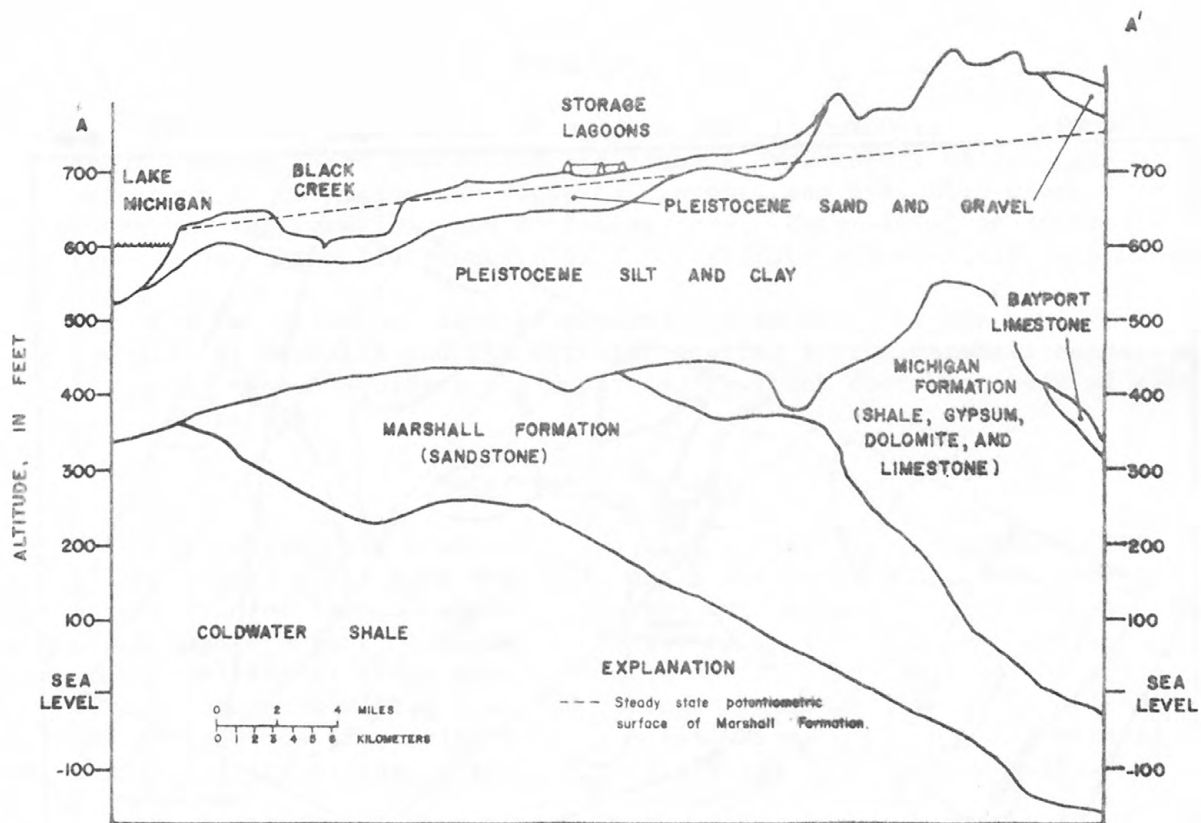


Figure 3. Cross section of study area showing geologic formations. Location of section A-A' shown in figure 1.

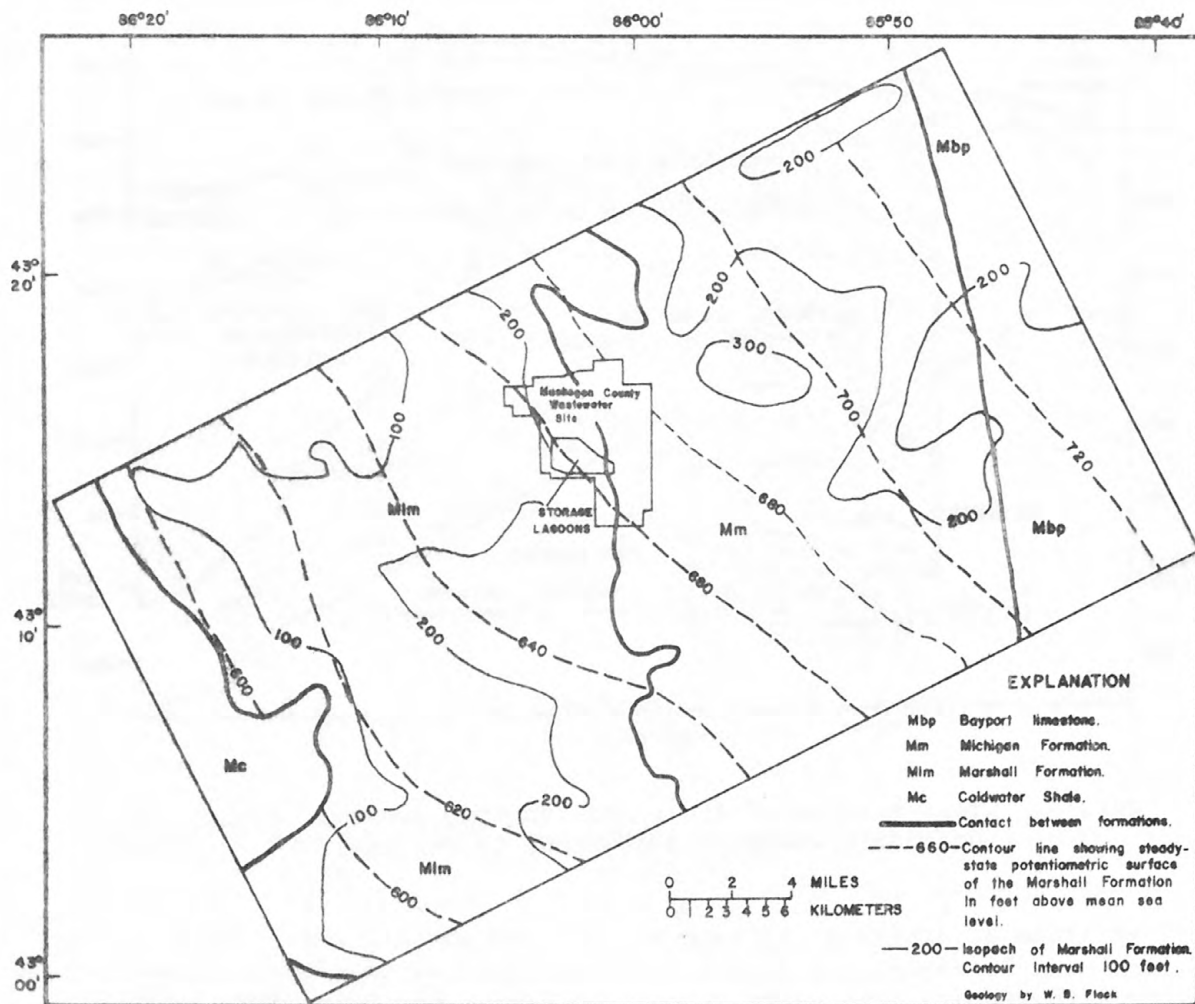


Figure 4. Subcrop formations, thickness, and potentiometric surface of Marshall Formation.

sandstone, and shale. The Michigan Formation is composed of relatively impermeable shale, gypsum, dolomite, and limestone. These two formations and the overlying silt and clay beds range in thickness from 80 ft (24 m) on the west to about 800 ft (245 m) on the east; they form a confining layer over the Marshall Formation. The Marshall is composed of highly permeable gray, pink, and red sandstones (deWitt, 1960, Dorr and Eschman, 1970). This formation ranges in thickness from 0 in the western part of the study area to as much as 300 ft (90 m) in the eastern part. The Coldwater Shale, a thick impermeable shale sequence, underlies and confines the Marshall Formation.

HYDROLOGY

The region studied has an area of 625 mi² (1620 km²). Logs of water wells drilled since 1969 and logs of exploratory oil and gas wells were used to determine the location, extent, and hydraulic properties of water-bearing materials and confining beds. Water-level measurements recorded on the well logs were used to estimate steady-state conditions.

The two principal aquifer systems are the water-table aquifer in the glacial deposits and the artesian aquifer in the Marshall Formation (fig. 3). These aquifers are separated by thick confining beds of silt, clay, and shale.

Water-table Aquifer

The water-table aquifer in the glacial deposits is composed principally of sand containing some silt and gravel. The thickness of the aquifer ranges from 0, where it pinches out against the morainic deposits in the eastern part, to about 80 ft (24 m) in the west. Hydraulic conductivities of these materials were assigned on the basis of well-performance data within the study area and a general knowledge of the characteristics of glacial materials elsewhere. Initial estimates of transmissivity of the aquifer were determined from logs of wells.

Initial transmissivity values were later modified during model calibration. In the model, the water-table aquifer is represented by two layers; values contoured in figure 5 were the sum of the transmissivity values for the two layers.

The regional gradient of the water table is toward Lake Michigan. Water levels range from 820 ft (250 m) above mean sea level on the east side of the study area to 580 ft (177 m) along the shore of Lake Michigan. Water discharges from the aquifer regionally to Lake Michigan and locally to streams. Along the north edge of the wastewater facility, the water table slopes downward about 80 ft/mi (15 m/km) northward toward Mosquito Creek.

Recharge to the water-table aquifer was calculated to be 8 in/yr (20 cm/yr), based on long-term precipitation, streamflow, and groundwater levels. This value agrees well with those of Allen, Miller, and Wood (1972) and Walton (1970) for similar geohydrologic conditions in

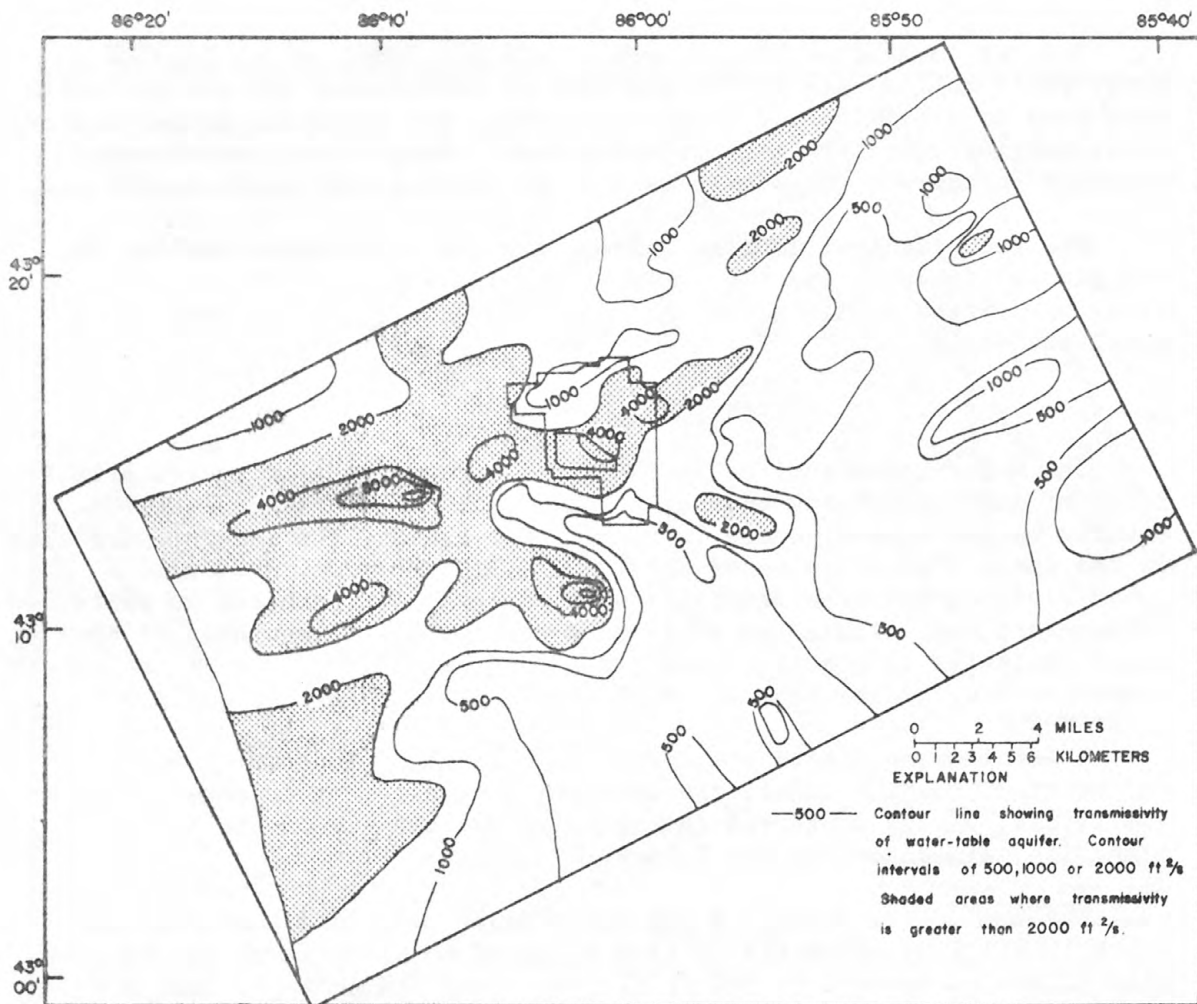


Figure 5. Transmissivity of the water-table aquifer.

southwestern Michigan and Illinois. Precipitation at the wastewater site, and recharge for 1974-75, are shown on table 1. Recharge rates were higher than average in 1974 and 1975 because of intensive storms.

Artesian Aquifer

The Marshall Formation is an artesian aquifer. Hydraulic conductivity values, determined from 7 pumping tests and 33 specific-capacity analyses (Brown and others, 1963, Walton, 1962), ranged from 200 to 20,000 (gal/d)/ft² (8 to 800 m/d); the median value was approximately 2000 (gal/d)/ft² (80 m/d). The thickness of the Marshall Formation was determined from 750 oil and gas logs. In the eastern half of the study area, where the formation underlies the Michigan Formation, the thickness ranges from 60 to 300 ft (100 m); the average thickness is about 240 ft (73 m). In the western half, the thickness gradually decreases westward from about 230 to 0 (70 m to 0) along the contact with the Coldwater Shale (fig. 3). Transmissivity of the Marshall Formation, determined by multiplying the formation thickness by the median hydraulic conductivity, ranges from 0 along the west border of the study area, where the Coldwater Shale is present, to 86,000 ft²/d (8000 m²/d) just east of the wastewater site.

In most of the study area and to the east, the Marshall Formation is recharged by downward flow from the water-table aquifer. Water levels in the artesian aquifer are lower than the overlying water table; at the east edge of the study area, this difference is as much as 100 ft (30 m). Flow within the Marshall Formation is to the west. The slope of the potentiometric surface is 4 to 10 ft/mi (0.8 to 1.9 m/km) (fig. 4). At the shore of Lake Michigan, the potentiometric head of the Marshall Formation is about 590 ft (180 m) above sea level; whereas, the lake level is 580 ft (177 m) above sea level. The vertical flow pattern here changes from downward recharge to upward discharge toward the lake.

Confining Unit

The confining beds of silt, clay, and shale that separate the water-table and artesian aquifers range in thickness from 80 ft (24 m) in the west to about 800 ft (245 m) in the east. Vertical hydraulic conductivity of confining beds, 0.10 (gal/d)/ft² (0.004 m/d), was estimated from previous investigations of similar materials in southern Michigan (Allen, Fleck, and Hanson, 1972) and in Ohio (Norris, 1963).

MODEL DESCRIPTION

A ground-water model is a mathematical description of the movement of water in a geologic environment. When converted to a form that permits computer solution of equations containing hydrologic and hydraulic parameters, the response of the ground-water system to stress may be predicted.

Table 1. Monthly precipitation, ground-water discharge to streams, change in ground water in storage and ground-water recharge in inches, in the model area.

Negative values of change in ground water in storage indicate declining water levels. Negative values of ground-water recharge result from losses due to evapotranspiration.

	Precipitation ^{1/}		Ground-water discharge to streams ^{2/}		Change in ground-water storage ^{3/}		Ground-water recharge ^{4/}	
	1974	1975	1974	1975	1974	1975	1974	1975
January	3.5	2.0	0.9	0.8	2.0	1.2	2.9	2.0
February	1.4	2.0	0.9	0.7	-0.5	1.2	0.4	1.9
March	4.3	2.2	1.4	1.1	0.8	1.0	2.2	2.1
April	2.7	2.0	1.6	1.2	1.2	0	2.8	1.2
May	5.8	2.0	1.5	0.9	-0.5	-1.9	1.0	-1.0
June	3.4	4.7	1.0	0.8	-2.5	-0.2	-1.5	0.6
July	1.1	1.6	0.6	0.6	-2.5	-2.2	-1.9	-1.6
August	3.6	9.3	0.5	0.5	-0.2	1.4	0.3	1.9
September	0.5	1.0	0.4	1.2	-0.8	1.2	-0.4	2.4
October	1.2	0.8	0.6	0.8	0.5	-0.2	1.1	0.6
November	2.6	3.6	0.7	1.0	1.5	0.7	2.2	1.7
December	1.3	2.5	0.7	1.4	0.2	1.1	0.9	2.5
TOTAL	31.4	33.7	10.8	11.0	-0.8	3.3	10.0	14.3

^{1/} Data from rain gage at wastewater disposal site.

^{2/} Average values calculated from streamflow hydrographs.

^{3/} Average from seven observation wells using specific yield of 0.2.

^{4/} Sum of ground-water discharge and change in ground-water storage.

Basic ground-water flow model

The differential equation describing flow in a porous saturated medium is

$$\frac{\partial}{\partial x} K_x \frac{\partial h}{\partial x} + \frac{\partial}{\partial y} K_y \frac{\partial h}{\partial y} + \frac{\partial}{\partial z} K_z \frac{\partial h}{\partial z} = S_s \frac{\partial h}{\partial t} + w(x, y, z, t), \quad (1)$$

where h is the head at time t .

K_x , K_y , and K_z are principal components of the hydraulic conductivity tensor aligned with the coordinate axes,

S_s is the specific storage.

and w is a source term for inflow or withdrawal per unit volume of an aquifer. In most cases equation 1 cannot be solved analytically. However, it can be replaced with an approximating finite difference equation and thus solved numerically.

The area being modeled is divided into discrete cells by a rectangular grid. The hydraulic properties of the material in each cell are assumed to be homogeneous, and the potentiometric head is calculated for a node at the center of each cell.

The computer program incorporating the techniques used to approximate a solution to equation 4 was developed by Trescott (1975). Modifications were made by Larson (written commun., March 1976) to simulate tile drainage and river leakage. A typical nonequilibrium simulation with this program will consist of several recharge periods, during each of which recharge to the top layer remains constant. The recharge periods are subdivided into time steps. The length of the time step (Δt in equation 5) affects the accuracy of the approximation to equation 3. As the time steps get shorter, the accuracy of approximation improves, but computation time increases. Similarly, as the cell dimensions become smaller, the accuracy of approximation improves, but again the computation time increases. Thus, the length of the time step and size of the cells must be selected to give acceptable results in a reasonable processing time.

A listing of the computer program with instructions for setting up the data deck is contained in the appendix. More complete documentation is contained in the user's manual by Trescott (1975). The program was written in Fortran for use on IBM 360/370 computer.¹ For this application approximately 600,000 bytes of core storage were required.

¹ The use of the brand name in this report is for identification purposes only and does not imply endorsement by the U.S. Geological Survey.

Drainage Analysis

A separate analysis was made to determine the effects of the drainage tiles. In this analysis, a cross-sectional model, based on simulation techniques described by Prickett and Lonquist (1971), was utilized. The model simulated flow in a vertical plane perpendicular to a single drainage tile. The vertical depth of the cross section was assumed to be 33 ft (10 m), which is representative of the thickness of the water-table aquifer; the width was taken as half the drain spacing or 250 ft (75 m). To represent the drainage tile, a single node at one side of the mesh was held at a head equal to its elevation. This node was a square, 0.5 ft (0.15 m) on a side; inflow was permitted through one side and through the bottom. The simulation represented half of the flow field to a drain. The hydraulic conductivity of the drain node was reduced to a fraction of that elsewhere to represent the hydraulic resistance of the drain pipe and the surrounding fiberglass net. Also, modifications were made in the model to allow the water-table boundary to move.

During each simulation, the model was set to represent a full saturated rectangular cross section, and recharge was applied to the uppermost node. A nonequilibrium simulation was conducted until steady-state conditions were achieved. As the nodes became dewatered during the simulation process, they were removed from the system by lowering the water-table boundary, and calculations were repeated to achieve a solution corresponding to the new boundary position. Except for the drain node and the recharge nodes, all boundaries were treated as zero-flow boundaries.

Model runs were made using values of lateral hydraulic conductivity that are in the range found at the Muskegon site. Also, several different ratios of lateral to vertical hydraulic conductivity were used. Results showed that the flow into both halves of the drainage tile could be expressed approximately as a function of hydraulic conductivity and average head above the drain by the relation

$$Q = K_L (h - V)G \text{ when } h > V, \quad (2)$$

where Q is the flow into a unit length of drainage tile,

h is the average water-table elevation in the area drained by the tile,

V is the elevation of the tile,

K_L is the lateral hydraulic conductivity,

and G is a factor that varies with the anisotropy and with the hydraulic conductivity of the drain node.

When a vertical hydraulic conductivity equal to one-tenth the lateral hydraulic conductivity was used and when the conductivity in the drain node was set at one-tenth of the lateral conductivity elsewhere, G was found to be approximately 0.1. By varying the anisotropy and the conductivity of the drain node, values of G from 0.05 to 0.40 were obtained. These values were subsequently used in checking the sensitivity of the three-dimensional model under various drainage conditions.

The Muskegon Model

The simulated area was divided by a rectilinear grid into 3 water-bearing layers each having 44 rows and 48 columns, as shown in figure 6. Layer 1, the lowermost layer, represents the Marshall Formation. Layer 2 represents all the water-table aquifer except the upper 20 ft (6 m); layer 3 represents the upper 20 ft (6 m). The water-table aquifer was divided into two layers so that drains and streams could be simulated in the uppermost layer as being shallow, rather than deep and affecting the full thickness of the aquifer. The confining unit between the Marshall Formation and the water-table aquifer was not treated as a separate layer; its effect was incorporated into the vertical hydraulic conductivities of layers 1 and 2.

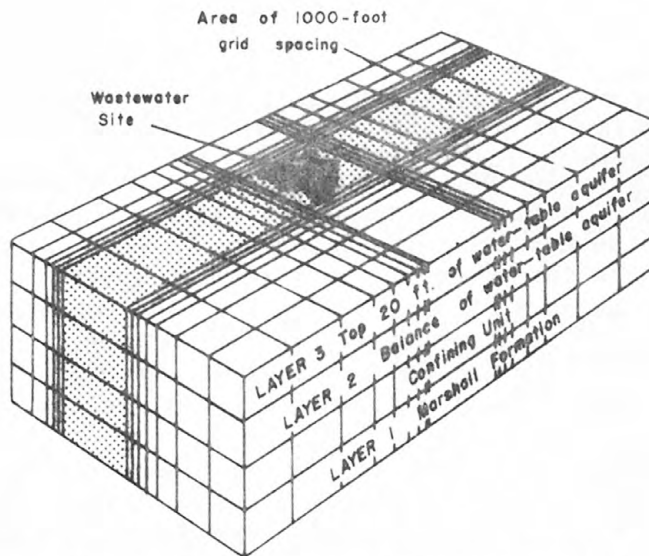


Figure 6. Grid spacing used in finite-difference model. Horizontal spacing ranges from 1,000 ft (305 m) at wastewater site to 15,000 ft (4500 m) at edge of model.

The horizontal grid was designed so that the smallest cells (1000 ft (300 m) on a side) represented the wastewater site (shaded area in figure 6). The grid spacing increased by a factor of 1.5 to the boundaries of the model. Moving the boundaries far beyond the simulated wastewater site greatly diminishes the sensitivity of the model at that site to the effects of the boundaries. All boundaries were treated as being the constant-flow type except the western boundary of the top layer, which was assumed to have a constant head equal to the elevation of Lake Michigan.

Storage effects were simulated only for the uppermost layer. Compressive storage in the unconfined aquifer below the water table was not simulated, nor was artesian storage in the Marshall Sandstone. Compressive storage in the confining unit above the Marshall was also neglected. It was assumed that these compressive effects would be negligible in comparison to the water-table storage in the uppermost layer. Thus, the model grid below the uppermost layer was purely transmissive in character.

The effects of streams on the ground-water system were simulated by the method described by Prickett and Lonquist (1971, p. 33). The method is based on assumptions that a streambed layer separates the stream from the aquifer and that seepage from the stream to the aquifer becomes constant when the water level in the aquifer falls below the bottom of the streambed. Under these assumptions, the rate of flow through the streambed is expressed by the equation.

$$Q_r = K' A (M-p)/b \quad (3)$$

where Q_r is the flow between the stream and the aquifer (it is positive when the flow is from the stream to the aquifer),

K' is the hydraulic conductivity,

b is the thickness of the streambed layer,

A is the area of streambed within the model cell,

M is the elevation of the stream surface,

p is the head in the top layer or the elevation of the bottom of the streambed layer, whichever is greater.

In the Muskegon model, K' was taken as one-tenth of the local lateral hydraulic conductivity in layer 3 of the model, and b was taken as 5 ft (1.5 m). Other values of these parameters were also tried during calibration.

Recharge to the water-table aquifer from the lagoons was treated as an evenly distributed constant seepage. In the initial calibrations, the seepage rate through the bottom of the lagoons was assumed to be 21 ft³/s (0.50 m³/s). This estimate was based on pumpage records from the seepage ditches around the lagoons. Approximately 23 ft³/s (0.66 m³/s) is pumped from these ditches, and it was assumed that 80 percent of this, or 19 ft³/s (0.53 m³/s), represents seepage from the lagoons. The balance represents drainage from other sources--that is, from surrounding irrigation circles or from regional ground-water flow. Also, because the seepage ditches do not completely enclose the lagoons, it was assumed that only 90 percent of the seepage from the lagoons was intercepted by the ditches. A total seepage of 21 ft³/s (0.58 m³/s) was, therefore, used.

If this estimate of lagoon seepage is accurate, it implies that the vertical hydraulic conductivity beneath the lagoons is considerably lower than that of the streambed or that of the water-table zone. This may reflect the effect of organic matter deposited in the lagoons, even though the system has been operated only since 1974.

The seepage ditches around the lagoons were simulated as streams, using equation 3; the streambed conductivity was assumed to be one-tenth of the local lateral hydraulic conductivity, and the streambed layer thickness was assumed to be 5 ft (1.5 m). Seepage into the ditches in computer simulations was found to be in close agreement with field data.

Data Base

Data required for the model included hydraulic properties of the hydrogeologic units, hydraulic properties of rivers and drainage tiles, and constant recharge and discharge sources. Heads and baseflow calculated by the model were compared with measured water levels and streamflow for calibration.

Initial estimates of the transmissivity of each cell in layers 1 and 2, and horizontal hydraulic conductivity of each cell in layer 3 were estimated from lithologic data. Specific yield of layer 3 was set to 0.2. Storage coefficient for layers 1 and 2 was set to 0. The vertical hydraulic conductivity between the bottom two layers at each horizontal grid location was assumed to be equal to the vertical hydraulic conductivity of the confining unit divided by the thickness of the confining unit at the grid location. The vertical hydraulic conductivity between the top two layers at each grid location was initially assumed to be equal to one-tenth of the horizontal hydraulic conductivity of the cell in the top layer at that grid location.

Engineering blueprints were used to estimate length and average elevation of drainage tiles. For each cell, the length and average elevations of all tiles in the cell were used.

The width, length, and elevation of rivers in each cell of the upper layer were determined from topographic maps. Thickness of the streambed layer, as noted previously, was assumed to be 5 ft (1.5 m). The vertical hydraulic conductivity of the streambed in each cell was initially assumed to be one-tenth of the horizontal hydraulic conductivity of the cell. A stream leakage factor was calculated independently of the model program from width, length, hydraulic conductivity, and streambed thickness. The leakage factor and the stream elevation were used in model computations.

In calculating irrigation rates, the design capacities and number of hours of operating time of the irrigation rigs were considered. Total recharge for each cell in the model area was calculated from lagoon leakage, amount of irrigation, and natural recharge.

Water level data from two sources were available. A long-term average water table map was derived from levels reported in driller's well records. Well hydrographs were drawn from records of water levels, which were measured twice a month in 90 observation wells during the study period.

All available geologic and hydrologic data were used. When field data were not available, regional estimates were made, as explained earlier. As more data became available they can be used to refine the model.

Calibration

A model is calibrated by repetitively running the computer program using available hydraulic data. The results of each simulation are then used to refine estimates of hydraulic parameters for subsequent simulations. Developing the Muskegon model involved a steady state calibration and a transient calibration. In each of these calibrations, it was observed that head differences between the upper and middle layers of the grid were negligible - that is, that heads in the upper 20 ft (6 m) of the water-table aquifer were essentially equal to those in the lower part of the aquifer. In presenting each simulation result, therefore, a single map of head in the water-table aquifer or changes in head has been utilized, rather than presenting separate maps for the two model layers representing the unconfined aquifer. Results are not shown for the artesian aquifer, inasmuch as changes in the potentiometric surface were small in all simulations.

During the steady-state calibration a specific yield of 0 was set in all cells of the uppermost layer, thereby eliminating the time dependence in equation 1. The purpose of steady-state calibration was to match the long-term average water table before construction of the wastewater facility with the water table calculated by the model. A uniform recharge of 8 in/yr (20 cm/yr) was used in computations.

A series of steady-state simulations was made to determine the sensitivity of the model to variations in the hydraulic parameters. The simulations indicated that the model was insensitive to changes in vertical hydraulic conductivity or stream leakage. It was, however, very sensitive to changes in horizontal hydraulic conductivity and transmissivity. Best results were obtained when these parameters were reduced to half their initial values. The comparison between computed and measured water-table positions after steady state calibration is shown in figure 7. Along the north edge of the area, steep ground-water gradients caused difficulty in matching water levels.

Transient calibration was used to test assumptions regarding the specific yield and the hydraulic characteristics of the wastewater facility and to further refine other parameters. The period January 1974 to December 1975 was used for transient calibration. This period included the first growing season during which the facility operated at full capacity, as well as some earlier period of operation at part capacity. Twenty-four monthly recharge periods were simulated. Initial heads used in the transient calibration were those calculated in the steady state model.

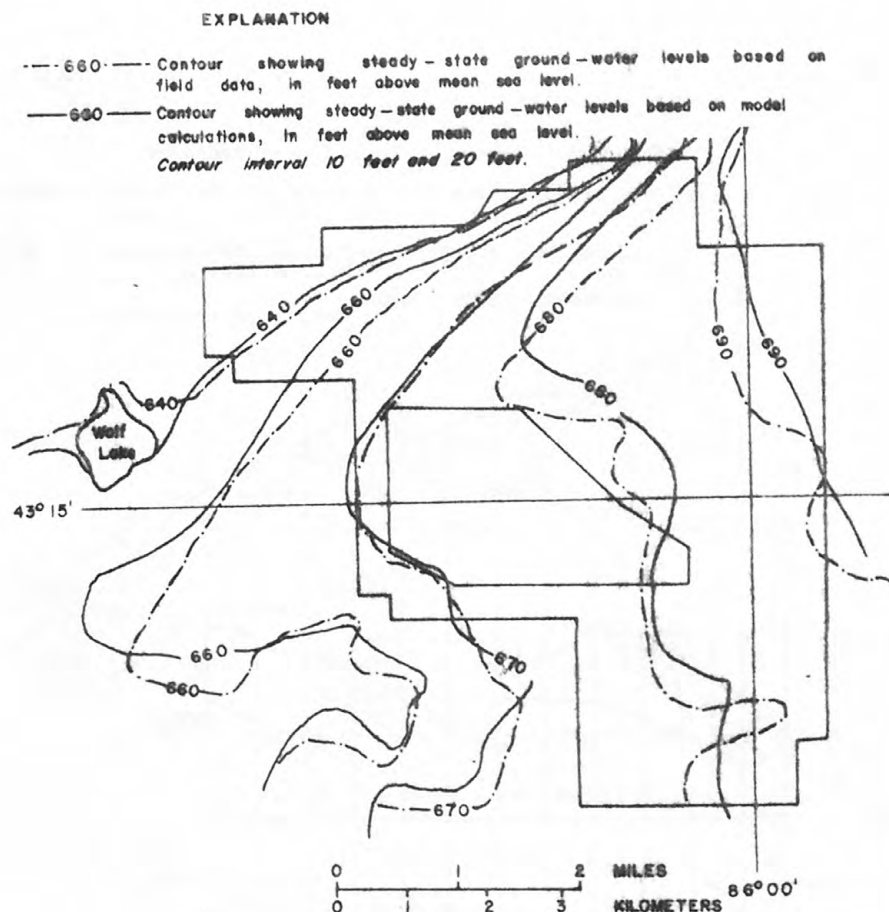


Figure 7. Steady-state water table from field measurements and from model calculations.

Only minor changes were made in hydraulic conductivity and transmissivity during transient calibration. The model was not particularly sensitive to changes in specific yield, and the original value of 0.2 was ultimately retained because it gave the best results. Stream location and stream surface elevation proved to be important factors. For this reason an effort was made to include all streams regardless of size and to establish stream surface elevation as accurately as possible.

Figure 8 shows four representative hydrographs after transient calibration, illustrating the final match between computed and observed water-level trends. Figure 9 compares measured water levels in the water-table aquifer during July 1975 with those calculated by the model during transient calibration. Figure 10 shows the comparison between the amount of observed and computed baseflow from the vicinity of the wastewater site during the period of transient calibration.

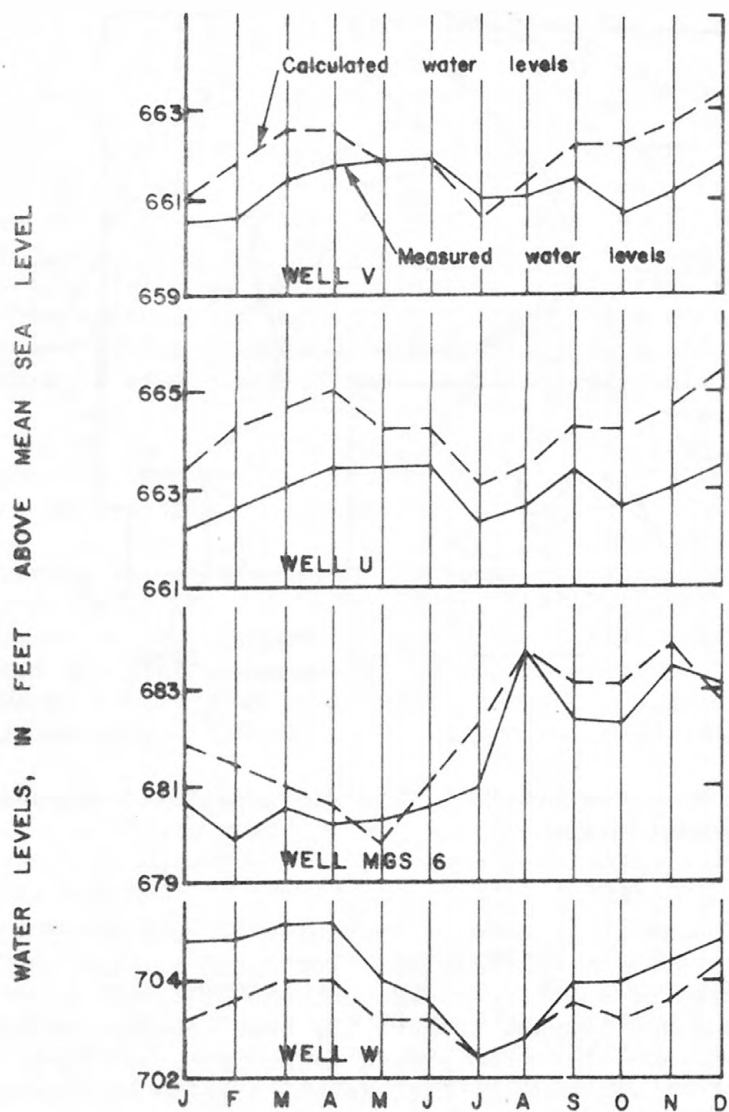


Figure 8. Hydrographs of water levels in selected wells during 1975 (from field measurements and from model calculations). Location of wells shown in figure 2.

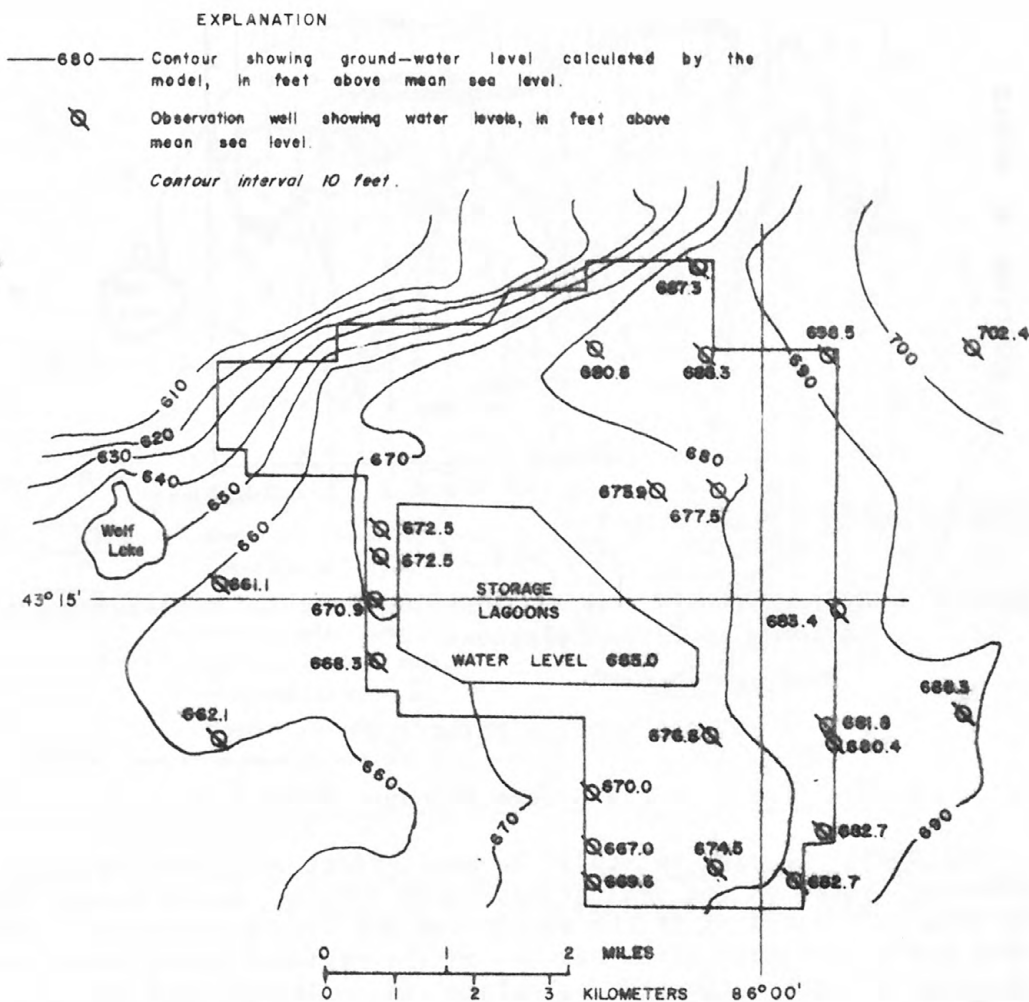


Figure 9. Altitude of water table during July 1975 from field measurements and from model calculations.

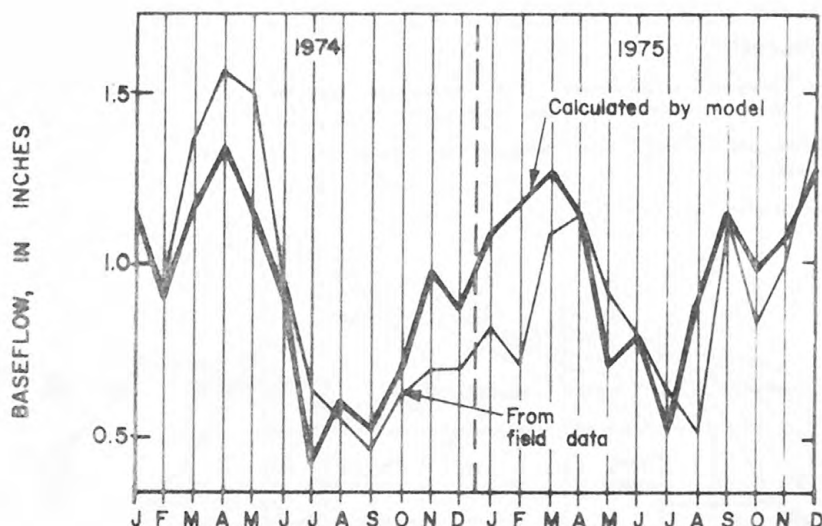


Figure 10. Hydrographs of monthly baseflow from field measurements and from model calculations.

INITIAL MODEL APPLICATIONS

Effects of the System through 1975

The model was used to study the past effect of operating the wastewater system on regional ground-water levels. Water levels for 1975 were calculated as if the system had not been constructed. The impact the wastewater system has had on the regional water table was determined by subtracting the calculated water levels from the actual levels. Figure 11 shows this impact. In most of the site, the effect of the facility has been to lower ground-water levels. However, water levels have risen in the northwestern part of the area, where there are relatively few drains.

Predictive Simulations

Several predictive simulations were made assuming a variety of operational conditions. In the first simulation, it was assumed that irrigation would be maintained at a uniform rate of 1.5 in (4 cm) per week (the average rate for 1975) in all irrigation circles, that lagoons would leak steadily at a rate of $21 \text{ ft}^3/\text{s}$ ($0.58 \text{ m}^3/\text{s}$), and that drainage-tile performance would again be described by equation 8, with G taken as 0.1. In this and in all subsequent predictive simulations, the December 1975 water-level configuration was used as the starting surface. Steady state was usually attained after 3 or 4 years of operation. The simulations were nevertheless carried for several additional years. The results for 10 years of operation are presented in each case as the steady-state condition.

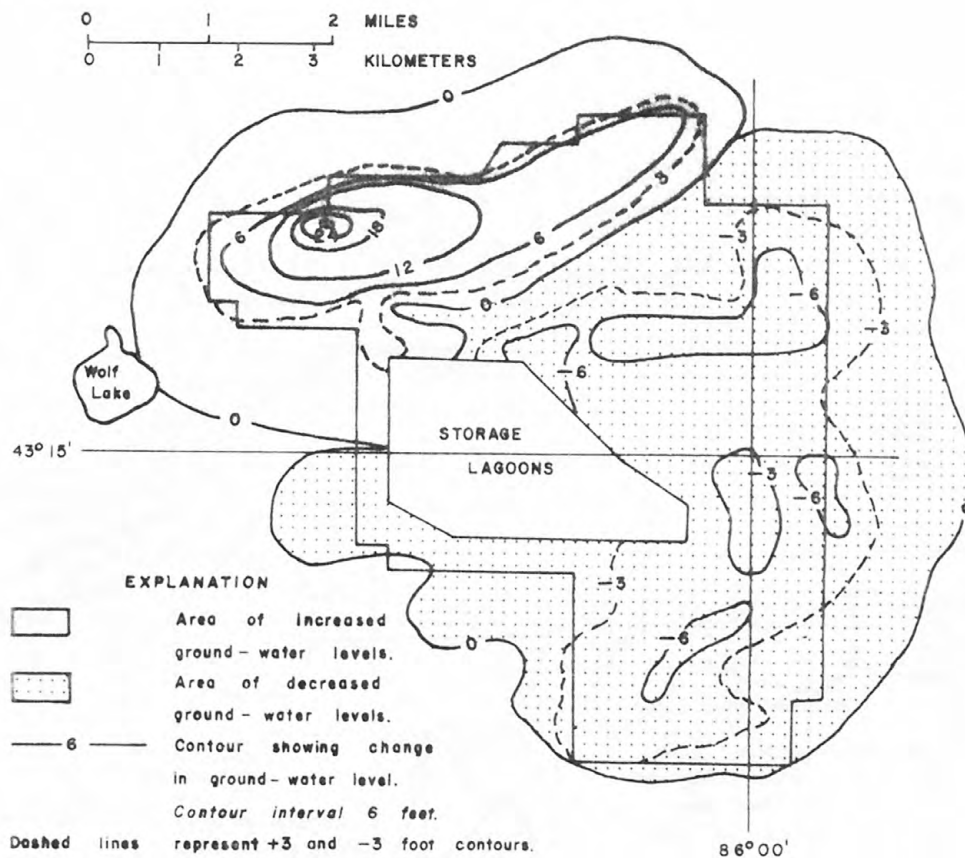


Figure 11. Changes in water level in the water-table aquifer caused by operation of the wastewater system, July 1975.

The results of the initial predictive simulation are shown in figure 12. The effects are primarily at the wastewater site. For a few areas within the site, the computed water levels are above land surface. These areas are shown on figure 12 as being waterlogged; that is, the water table in these areas is approximately at land surface.

The results shown in figure 12 do not imply that waterlogging is a necessary consequence of the waste disposal operation, but, rather, that irrigation at a rate of 1.5 in (4 cm) per week over the entire area would probably cause such problems. In practice, irrigation rates will vary from one part of the system to another and will be managed so as to avoid waterlogging, or alternatively, the number of drainage tiles could be increased to avoid waterlogging.

A second predictive simulation was made in which the irrigation rate was maintained at a uniform rate of 3 in (8 cm) per week (the design irrigation rate). Performance of drainage tiles and the lagoon seepage were the same as in the first predictive simulation. The 1975 water-level surface was again taken as the initial condition, and the

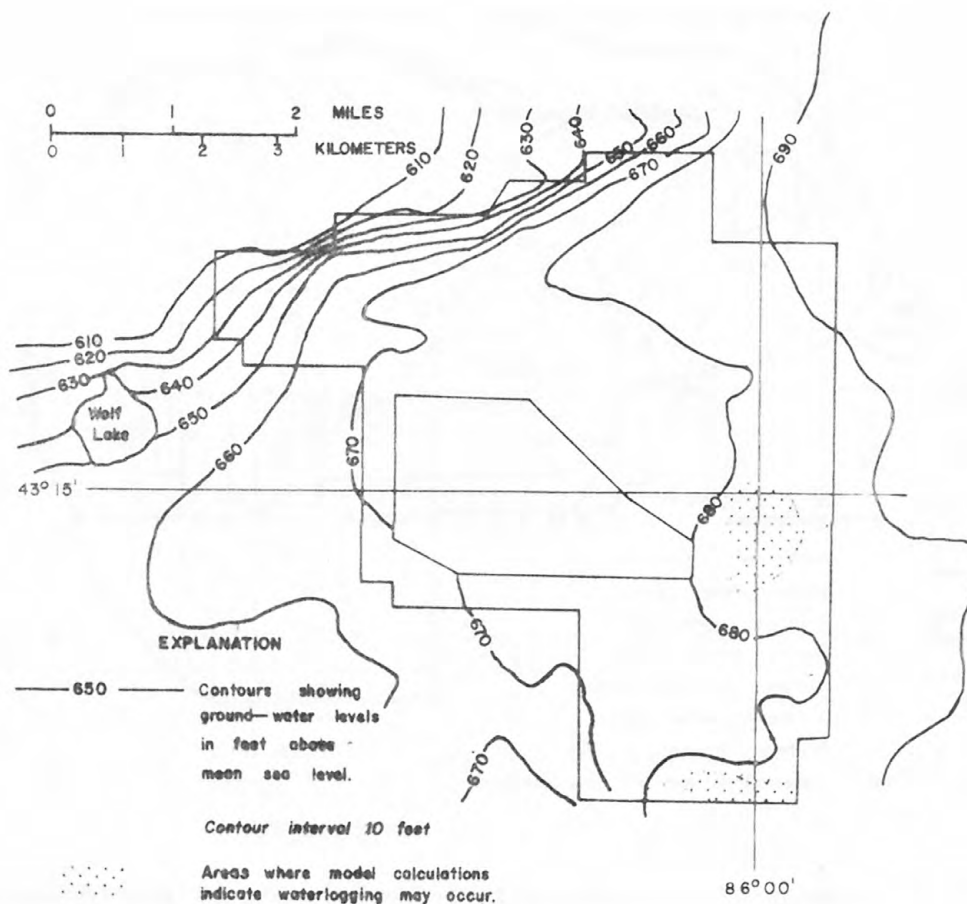


Figure 12. Steady-state water table from model calculations with 1.5 in (4 cm) per week irrigation, normal lagoon leakage, and tile seepage.

simulation was continued to steady state. Figure 13 shows contours of the water-table elevation after steady state was reached. The increase in the irrigation rate from 1.5 to 3.0 in (4 to 8 cm) per week caused the waterlogged area to increase from 350 acres to 1000 acres (150 ha to 400 ha). Figure 14 shows the rise in water level that may be expected with an increase in irrigation rate from 1.5 in (4 cm) per week to 3.0 in (8 cm) per week. The figure shows the difference between water level elevations in figures 12 and 13.

A third predictive simulation assumed that the irrigation rate would be maintained at 3.0 in (8 cm) per week and that lagoon leakage would continue at $21 \text{ ft}^3/\text{s}$ ($0.58 \text{ m}^3/\text{s}$), but that the efficiency of the drainage tiles would be severely reduced by clogging. In the model, the factor G in equation 8 was reduced from 0.1 to 0.025, and thus the flow to the drains, for a given head differential, would be only 25 percent of that in earlier simulations. The 1975 water-table surface was again

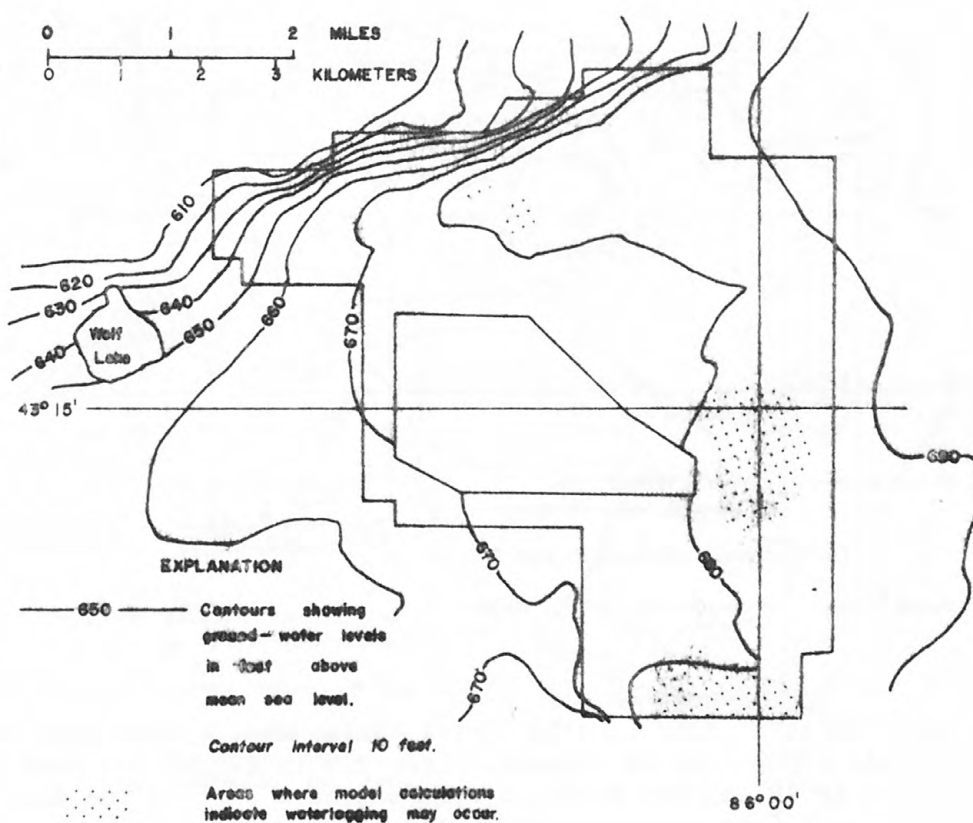


Figure 13. Steady-state water table from model calculations with 3 in (8 cm) per week irrigation, normal lagoon leakage, and tile seepage.

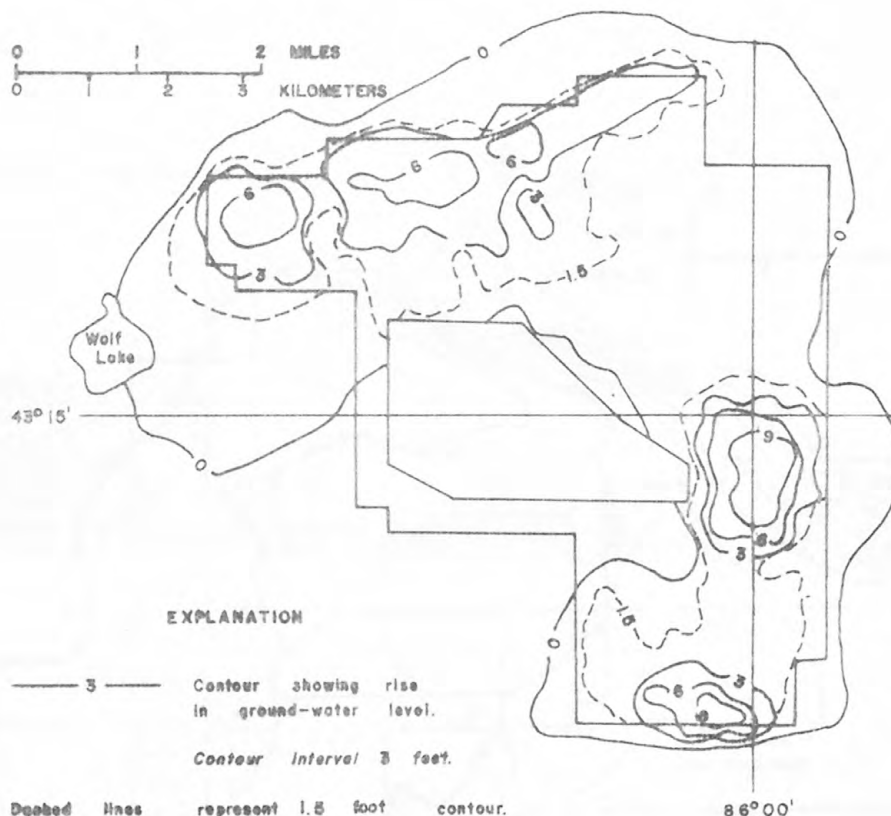


Figure 14. Predicted rise in water level in the water-table aquifer, if irrigation is increased from 1.5 in (4 cm) per week to 3 in (8 cm) per week. In the southern part of the site, the 1.5-in (4-cm)-per-week increase in irrigation rate causes a rise of as much as 9 ft (3 m).

taken as the starting condition, and calculations were continued until steady state was achieved. Figure 15 shows contours of the change in water level that would result from clogging of the drains. The contours represent the rise in water level, above the levels shown in figure 13, that would result if drainage efficiency were reduced but other factors remained as in the previous simulation. Nearly half the irrigated area becomes waterlogged. However, long before waterlogging became extensive, irrigation practices or drainage tiles could be modified.

A fourth predictive simulation was made keeping the rates of lagoon leakage, tile seepage, and irrigation the same as the second predictive simulation. However, the natural recharge was increased from 8 in/yr (20 cm/yr) to 12 in/yr (30 cm/yr). At the end of 10 years this simulation had not achieved steady state. Figure 16 shows the rise in water levels above the levels shown in figure 13 that would result from the increased recharge. The rise is negligible in most of the wastewater site, but as much as 6 ft (2 m) outside the site.

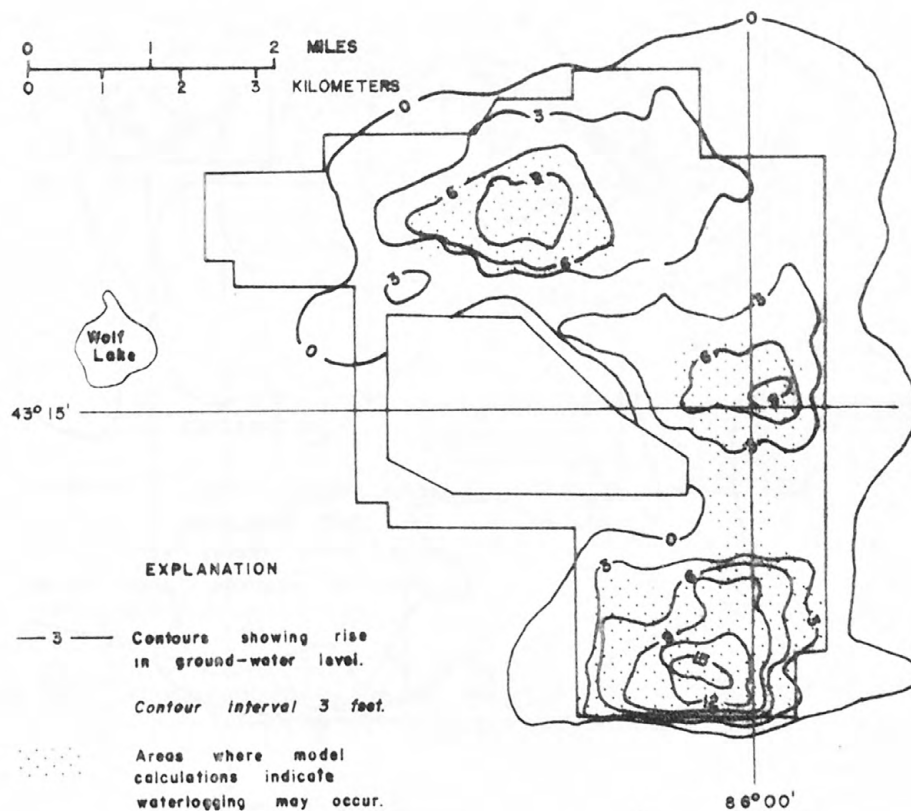


Figure 15. Predicted rise in water level in the water-table aquifer when seepage into drainage tile is reduced by 75 percent.

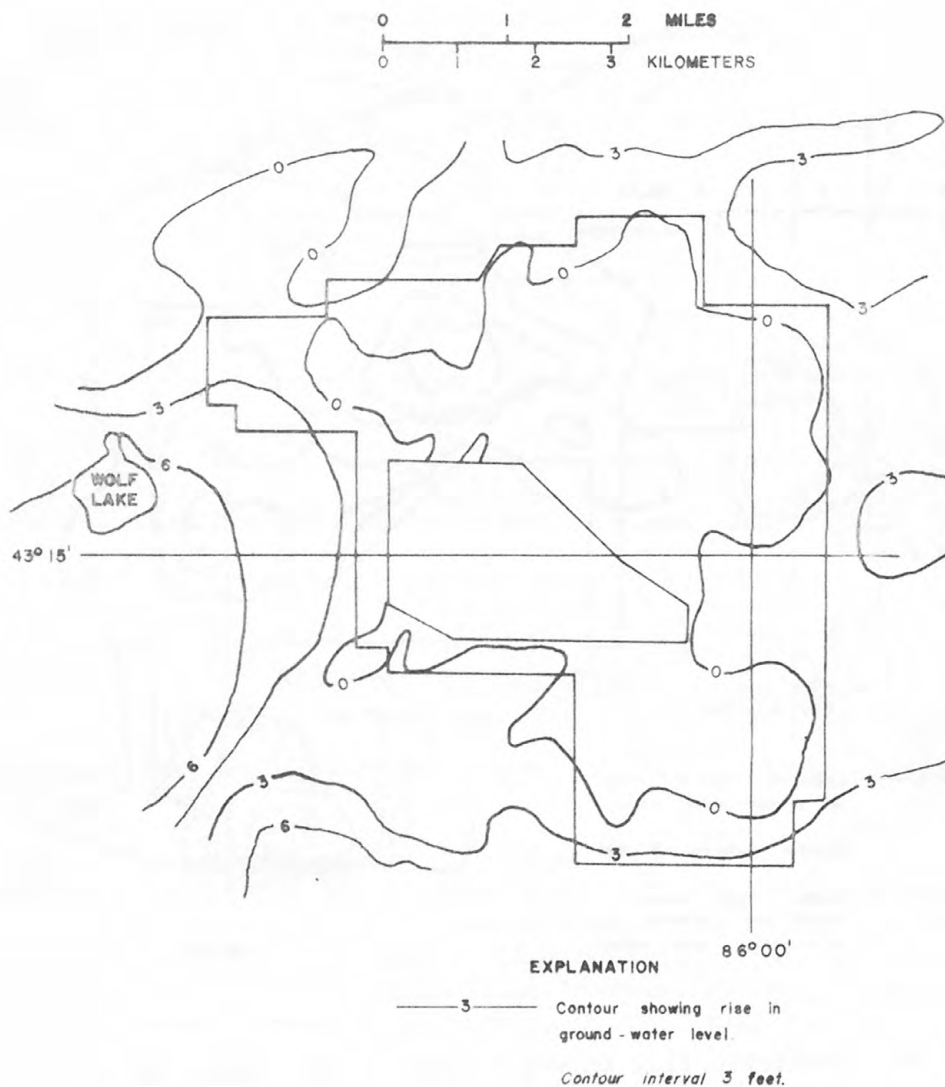


Figure 16. Predicted rise in water level in the water-table aquifer when natural recharge is increased from 8 in/yr (20 cm) to 12 in/yr (30 cm).

In the final predictive simulation, seepage from the lagoons was reduced to zero while the other factors remained unchanged. The purpose of this simulation was to determine the effect on water levels if the bottoms of the lagoons became effectively sealed with organic matter. The 1975 water-table was again taken as the starting condition, and the simulation was continued to steady state. Figure 17 shows contours of the changes in water level that would result from sealing the lagoons; that is, the changes caused by eliminating lagoon seepage while all other factors were maintained as in the second simulation. The changes in water level are relatively minor and are restricted to the immediate vicinity of the lagoons and a small area to the west that extends almost to Wolf Lake. There is virtually no reduction in the waterlogged areas from those shown in figure 13.

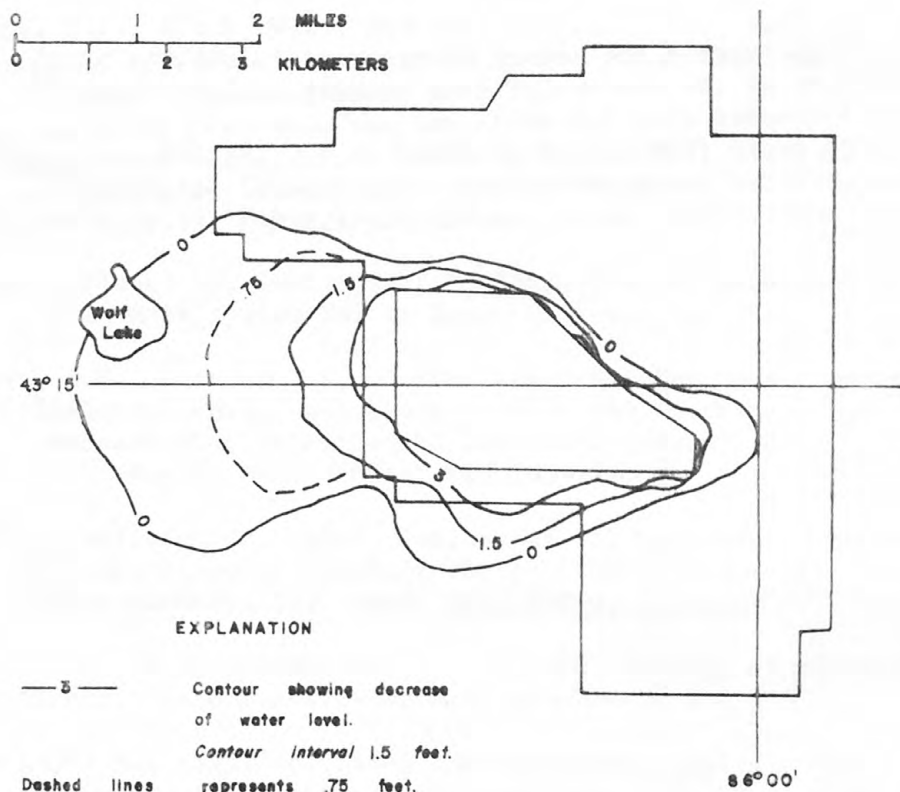


Figure 17. Predicted decline in water level in the water-table aquifer if the bottom of the storage lagoons become sealed.

SUMMARY

The ground-water system at the Muskegon wastewater site was simulated with a three-dimensional finite difference model. The model was calibrated from water-level records from 90 observation wells and from regional estimates of specific yield, recharge, and base flow. Data for simulating the necessary physical aspects of the wastewater system were taken from available engineering blueprints.

The calibrated model showed that, in general, the tile lines are effective in draining ground water from the site and that ground-water levels are lower than they would be under natural conditions. However, in the northwest part of the wastewater site, where tile lines are few and above the water table, water levels are as much as 24 ft (8 m) higher than they are at other locations within the site. Outside the site to the northwest, water levels are as much as 3 ft (1 m) higher than they would be under natural conditions.

Predictive simulations were made to determine the effects that varying operating conditions would have on the ground-water system. The simulations show, for example, that if the drainage tiles lost 75 percent

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APPENDIX

Model Program

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C -----MAN0010
C FINITE-DIFFERENCE MODEL FOR SIMULATION OF GROUND-WATER FLOW IN MAN0020
C THREE DIMENSIONS, SEPTEMBER, 1975 BY P.C. TRESCOTT, U. S. G. S. MAN0030
C WITH CONTRIBUTIONS TO MAIN, DATAI AND SOLVE BY S.P. LARSON MAN0040
C -----MAN0050
C MAN0060
C SPECIFICATIONS: MAN0070
C REAL *BYSTR MAN0080
C MAN0090

COMMON/FILS/ IFIL(14)
COMMON/IDDC/ IDD(80)
DIMENSION Y(106000), L(32), HEADNG(33), NAME(42), INFT(2,2), IOFT(
19,4), DUM(3) MAN0110
C MAN0120
C EQUIVALENCE (YSTR,Y(1)) MAN0130
C MAN0140
COMMON /INTEGR/ IO,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NMAN0150
1WEL,NUMT,IFINAL,IT,KT,IHEAD,IDRAW,IFLO,IERR,I2,J2,K2,IMAX,ITMX1,NCMAN0160
2H,IDK1,IDK2,IWATER,IGRE,IP,JP,IQ,JQ,IK,JK,K5,IPU1,IPU2,ITK MAN0170
COMMON /SPARAM/ TMAX,CDLT,DELT,ERR,TEST,SUM,SUMP,QR MAN0180
COMMON /RESMUN/ RESMO
COMMON /SARRAY/ ICHK(13),LEVEL1(9),LEVEL2(9) MAN0190
C MAN0200
DATA NAME/2*4H ,4H S,4HTART,4HING ,4HHEAD,4H ,4H STO,4HRAGMAN0210
1E,4H COE,4HFFIC,4HIENT,2*4H ,4H TR,4HANSM,4HISSI,4HVITY,5*4H MAN0220
2 ,4H TK,4H HY,4HDRAU,4HLIC ,4HCOND,4HUCTI,4HVITY,2*4H ,4HBOTMAN0230
3T,4HOM E,4HLEVA,4HTION,2*4H ,4H R,4HECHA,4HRGE ,4HRAE/ MAN0240
DATA INFT/4H(20F,4H4.0),4H(8F1,4H0.4)/ MAN0250
DATA IOFT/4H(1H0,4H,I2,,4H2X,2,4H0F6.,4H1/(5,4HX,20,4HF6.1,4H)) ,MAN0260
14H ,4H(1H0,4H,I5,,4H14F9,4H.5/(,4H1H ,4H5X,1,4H4F9.,4H5)) ,4H MAN0270
2 ,4H(1H0,4H,I5,,4H10E1,4H2.5/,4H(1H ,4H,5X,,4H10E1,4H2.5),4H) MAN0280
3,4H(1H0,4H,I5,,4H10E1,4H1.3/,4H(1H ,4H,5X,,4H10E1,4H1.3),4H) / MAN0290
C MAN0300
DEFINE FILE 2(14,2112,U,KKK)
C .....MAN0320
C MAN0330
C ---READ TITLE, PROGRAM SIZE AND OPTIONS--- MAN0340
C READ (5,200) HEADNG MAN0350
C WRITE (6,190) HEADNG MAN0360
197 READ(5,197) (IFIL(IRX),IRX=1,14)
FORMAT(14I4)
198 READ (5,198) (IDD(IRX),IRX=1,80)
FORMAT(80I1)
READ (5,497) RESMO
497 FORMAT (F10.0)
WRITE(6,498) RESMO
498 FORMAT ('ORESMO ',F10.3)
WRITE(6,191)
191 FORMAT('OVERSION G 1400 SEP 02 1976')
READ (5,160) IO,J0,K0,ITMAX,NCH,ND,NRIV
WRITE (6,180) IO,J0,K0,ITMAX,NCH,ND,NRIV
READ (5,210) IDRAW,IHEAD,IFLO,IDK1,IDK2,IWATER,IGRE,IPU1,IPU2,ITK MAN0390
WRITE (6,220) IDRAW,IHEAD,IFLO,IDK1,IDK2,IWATER,IGRE,IPU1,IPU2,ITKMAN0400
IERR=0 MAN0410
C MAN0420
C ---COMPUTE DIMENSIONS FOR ARRAYS--- MAN0430
C J1=J0-1 MAN0440
C I1=I0-1 MAN0450
C K1=K0-1 MAN0460
C I2=I0-2 MAN0470

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J2=J0-2	MAN0480
K2=K0-2	MAN0490
IMAX=MAX0(I0,J0)	MAN0500
NCD=MAX0(1,NCH)	MAN0510
ITMX1=ITMAX+1	MAN0520
ISIZ=I0*J0*K0	MAN0530
IK1=I0*J0	MAN0540
IK2=MAX0(1K1*K1,1)	MAN0550
ISUM=2*ISIZ+1	MAN0560
L(1)=1	MAN0570
DO 30 I=2,14	MAN0580
IF (I.NE.8) GO TO 20	MAN0590
L(8)=ISUM	MAN0600
ISUM=ISUM+IK2	MAN0610
IF (IK2.EQ.1) GO TO 10	MAN0620
IK=I0	MAN0630
JK=J0	MAN0640
K5=K1	MAN0650
GO TO 30	MAN0660
10 IK=1	MAN0670
JK=1	MAN0680
K5=1	MAN0690
GO TO 30	MAN0700
20 L(1)=ISUM	MAN0710
ISUM=ISUM+ISIZ	MAN0720
30 CONTINUE	MAN0730
L(15)=ISUM	MAN0740
ISUM=ISUM+J0	MAN0750
L(16)=ISUM	MAN0760
ISUM=ISUM+I0	MAN0770
L(17)=ISUM	MAN0780
ISUM=ISUM+K0	MAN0790
L(18)=ISUM	MAN0800
ISUM=ISUM+IMAX	MAN0810
L(19)=ISUM	MAN0820
ISUM=ISUM+K0*3	MAN0830
L(20)=ISUM	MAN0840
ISUM=ISUM+ITMX1	MAN0850
L(21)=ISUM	MAN0860
ISUM=ISUM+3*NCD	MAN0870
L(22)=ISUM	MAN0880
ISUM=ISUM+NCD	MAN0890
L(23)=ISUM	MAN0900
IF (IWATER.NE.ICHK(6)) GO TO 40	MAN0910
ISUM=ISUM+IK1	MAN0920
L(24)=ISUM	MAN0930
ISUM=ISUM+IK1	MAN0940
IP=I0	MAN0950
JP=J0	MAN0960
GO TO 50	MAN0970
40 ISUM=ISUM+1	MAN0980
L(24)=ISUM	MAN0990
ISUM=ISUM+1	MAN1000
IP=1	MAN1010
JP=1	MAN1020
50 L(25)=ISUM	MAN1030
IF (IQRF.NE.ICHK(7)) GO TO 60	MAN1040
ISUM=ISUM+IK1	MAN1050
IQ=I0	MAN1060
JQ=J0	MAN1070

GO TO 70	MAN1080
60 ISUM=ISUM+1	MAN1090
IQ=1	MAN1100
JQ=1	MAN1110
70 IF (ND.EQ.0) GO TO 75	
L(26)=ISUM	
ISUM=ISUM+IK1	
L(27)=ISUM	
ISUM=ISUM+ND	
L(28)=ISUM	
ISUM=ISUM+ND	
GO TO 76	
75 L(26)=ISUM	
ISUM=ISUM+1	
L(27)=ISUM	
ISUM=ISUM+1	
L(28)=ISUM	
ISUM=ISUM+1	
76 IF (NRIV.FQ.0) GO TO 77	
L(29)=ISUM	
ISUM=ISUM+IK1	
L(30)=ISUM	
ISUM=ISUM+NRIV	
L(31)=ISUM	
ISUM=ISUM+NRIV	
L(32)=ISUM	
ISUM=ISUM+NRIV	
GO TO 78	
77 L(29)=ISUM	
L(30)=ISUM+1	
L(31)=ISUM+2	
L(32)=ISUM+3	
ISUM=ISUM+4	
78 WRITE(6,170) ISUM	
C	MAN1130
---PASS INITIAL ADDRESSES OF ARRAYS TO SUBROUTINES---	MAN1140
CALL DATAI(Y(L(1)),Y(L(2)),Y(L(3)),Y(L(4)),Y(L(5)),Y(L(6)),Y(L(7)),	MAN1150
1,Y(L(8)),Y(L(9)),Y(L(15)),Y(L(16)),Y(L(17)),Y(L(19)),Y(L(23)),Y(L(MAN1160
224)),Y(L(25)),Y(L(26)),Y(L(27)),Y(L(28)),Y(L(29)),Y(L(30)),	
3 Y(L(31)),Y(L(32)))	
CALL STEP(Y(L(1)),Y(L(2)),Y(L(3)),Y(L(4)),Y(L(5)),Y(L(6)),Y(L(7)),	MAN1180
1Y(L(8)),Y(L(9)),Y(L(15)),Y(L(16)),Y(L(17)),Y(L(19)),Y(L(18)),Y(L(2	MAN1190
20)))	MAN1200
CALL SOLVE(Y(L(1)),Y(L(2)),Y(L(3)),Y(L(4)),Y(L(5)),Y(L(6)),Y(L(7)),	MAN1210
1,Y(L(8)),Y(L(9)),Y(L(15)),Y(L(16)),Y(L(17)),Y(L(19)),Y(L(10)),Y(L(MAN1220
21)),Y(L(12)),Y(L(13)),Y(L(14)),Y(L(20)),Y(L(25)),Y(L(23)),	
3 Y(L(26)),Y(L(27)),Y(L(28)),Y(L(29)),Y(L(30)),Y(L(31)),Y(L(32)),	
4 ND,NRIV)	
CALL COEF(Y(L(1)),Y(L(2)),Y(L(3)),Y(L(4)),Y(L(5)),Y(L(6)),Y(L(7)),	MAN1240
1Y(L(8)),Y(L(9)),Y(L(15)),Y(L(16)),Y(L(17)),Y(L(19)),Y(L(23)),Y(L(2	MAN1250
24)),Y(L(25)))	MAN1260
CALL CHECKI(Y(L(1)),Y(L(2)),Y(L(3)),Y(L(4)),Y(L(5)),Y(L(6)),Y(L(7)),	MAN1270
1),Y(L(8)),Y(L(9)),Y(L(15)),Y(L(16)),Y(L(17)),Y(L(19)),Y(L(21)),Y(L(MAN1280
2(22)),Y(L(25)),	
3 Y(L(26)),Y(L(27)),Y(L(28)),Y(L(29)),Y(L(30)),Y(L(31)),Y(L(32)),	
4 ND,NRIV)	
CALL PRNTAI(Y(L(1)),Y(L(2)),Y(L(4)),Y(L(5)),Y(L(9)),Y(L(15)),Y(L(MAN1300
16)))	MAN1310
C	MAN1320
---START COMPUTATIONS---	MAN1330

C	*****	MAN1340
C	---READ AND WRITE DATA FOR GROUPS II AND III---	MAN1350
	CALL DATAIN	MAN1360
	IRN=1	MAN1370
	NIJ=I0*J0	MAN1380
	DO 80 K=1,K0	MAN1390
	LOC=L(2)+(K-1)*NIJ	MAN1400
80	CALL ARRAY(Y(LOC),INFT(1,2),IOFT(1,1),NAME(1),IRN,DUM)	MAN1410
	DO 90 K=1,K0	MAN1420
	LOC=L(5)+(K-1)*NIJ	MAN1430
90	CALL ARRAY(Y(LOC),INFT(1,1),IOFT(1,2),NAME(7),IRN,DUM)	MAN1440
	DO 100 K=1,K0	MAN1450
	LOC=L(4)+(K-1)*NIJ	MAN1460
	L1=L(19)+K-1	MAN1470
	L2=L(19)+K0+K-1	MAN1480
	L3=L(19)+2*K0+K-1	MAN1490
	CALL ARRAY(Y(LOC),INFT(1,1),IOFT(1,2),NAME(13),IRN,DUM)	MAN1500
	Y(L1)=DUM(1)	MAN1510
	Y(L2)=DUM(2)	MAN1520
	Y(L3)=DUM(3)	MAN1530
100	WRITE (6,230) K,Y(L1),Y(L2),Y(L3)	MAN1540
	IF (ITK.NE.ICHK(10)) GO TO 120	MAN1550
	DO 110 K=1,K1	MAN1560
	LOC=L(8)+(K-1)*NIJ	MAN1570
110	CALL ARRAY(Y(LOC),INFT(1,1),IOFT(1,3),NAME(19),IRN,DUM)	MAN1580
120	K=K0	
	IF (IWATER.NE.ICHK(6)) GO TO 130	MAN1590
	CALL ARRAY(Y(L(23)),INFT(1,1),IOFT(1,4),NAME(25),IRN,DUM)	MAN1600
	CALL ARRAY(Y(L(24)),INFT(1,1),IOFT(1,1),NAME(31),IRN,DUM)	MAN1610
130	IF (IQRE.EQ.ICHK(7)) CALL ARRAY(Y(L(25)),INFT(1,1),IOFT(1,4),NAME(37),IRN,DUM)	MAN1620
	CALL M0AT	MAN1630
	IF (ND.NE.0) CALL D0AT(ND)	MAN1640
	IF (NRIV.NE.0) CALL D0AT2(NRIV)	
C		MAN1650
C	---COMPUTE TRANSMISSIVITY FOR UNCONFINED LAYER---	MAN1660
	IF (IWATER.EQ.ICHK(6)) CALL TRANS(1)	MAN1670
C		MAN1680
C	---COMPUTE T COEFFICIENTS---	MAN1690
	CALL TCOF	MAN1700
C		MAN1710
C	---COMPUTE ITERATION PARAMETERS---	MAN1720
	CALL ITER	MAN1730
C		MAN1740
C	---READ TIME PARAMETERS AND PUMPING DATA FOR A NEW PUMPING PERIOD---	MAN1750
140	CALL NEWPER	MAN1760
C		MAN1770
	KT=0	MAN1780
	IFINAL=0	MAN1790
C		MAN1800
C	---START NEW TIME STEP COMPUTATIONS---	MAN1810
150	CALL NEWSTP	MAN1820
C		MAN1830
C	---START NEW ITERATION IF MAXIMUM NO. ITERATIONS NOT EXCEEDED---	MAN1840
	CALL NEWITA	MAN1850
C		MAN1860
C	---PRINT OUTPUT AT DESIGNATED TIME STEPS---	MAN1870
	CALL OUTPUT	MAN1880
C		MAN1890
C	---LAST TIME STEP IN PUMPING PERIOD ?---	MAN1900

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C      IF (IFINAL.NE.1) GO TO 150                                MAN1910
C      ---CHECK FOR NEW PUMPING PERIOD---                        MAN1920
C      IF (KP.LT.NPER) GO TO 140                                MAN1930
C                                                                MAN1940
C      STOP                                                      MAN1950
C                                                                MAN1960
C                                                                MAN1970
C      ---FORMATS---                                           MAN1980
C                                                                MAN1990
C                                                                MAN2000
C                                                                MAN2010
C                                                                MAN2020
C      160 FORMAT (8I10)                                         MAN2030
C      170 FORMAT ('0',54X,'WORDS OF VECTOR Y USED =' ,I7)
C      180 FORMAT ('0',62X,'NUMBER OF ROWS =' ,I5/60X,'NUMBER OF COLUMNS =' ,I5MAN2040
C      1/61X,'NUMBER OF LAYERS =' ,I5//39X,'MAXIMUM PERMITTED NUMBER OF ITEMAN2050
C      2RATIONS =' ,I5//48X,'NUMBER OF CONSTANT HEAD NODES =' ,I5,
C      3 /,56X,'NUMBER OF DRAIN NODES =' ,I5,
C      4 /,56X,'NUMBER OF RIVER NODES =' ,I5)
C      190 FORMAT ('1',33A4)                                     MAN2070
C      200 FORMAT (20A4)                                         MAN2080
C      210 FORMAT (16(A4,1X))                                   MAN2090
C      220 FORMAT ('-SIMULATION OPTIONS: ' ,11(A4,4X))         MAN2100
C      230 FORMAT (1H0,44X,'DIRECTIONAL TRANSMISSIVITY MULTIPLICATION FACTORSMAN2110
C      1 FOR LAYER',I3,/,76X,'X =' ,G15.7/76X,'Y =' ,G15.7/76X,'Z =' ,G15.7) MAN2120
C      END                                                       MAN2130-
C      SUBROUTINE DATAI(PHI,STRT,OLD,T,S,TR,TC,TK,WELL,DELX,DELY,DELZ,FACDAT0010
C      1T,PERM,BOTTOM,QRE,LD,ELD,IDR,RH,RC,RB)
C      -----DAT0030
C      READ AND WRITE DATA                                     DAT0040
C      -----DAT0050
C      -----DAT0060
C      SPECIFICATIONS:                                         DAT0070
C      REAL *8PHI                                              DAT0080
C      REAL *8XLABEL,YLABEL,TITLE,XN1,MESUR                  DAT0090
C      REAL*4 LD
C                                                                DAT0100
C      DIMENSION PHI(I0,J0,K0), STRT(I0,J0,K0), OLD(I0,J0,K0), T(I0,J0,K0DAT0110
C      1), S(I0,J0,K0), TR(I0,J0,K0), TC(I0,J0,K0), TK(IK,JK,K5), WELL(I0,DAT0120
C      2J0,K0), DELX(J0), DELY(I0), DELZ(K0), FACT(K0,3), PERM(IP,JP), BOTDAT0130
C      3TOM(IP,JP), QRE(IQ,JQ), TF(3), A(I0,J0), IN(6), IOFT(9), INFT(2) DAT0140
C      4 ,ID(I0,J0),LD(1),ELD(1) ,IDR(I0,J0),RH(1),RC(1),RB(1)
C                                                                DAT0150
C      COMMON/FILS/ IFIL(14)
C      COMMON /INTEGR/ I0,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NDAT0160
C      1WEL,NUMT,IFINAL,IT,KT,IHEAD,IDRAW,IFLO,IERR,I2,J2,K2,IMAX,ITMX1,NCDAT0170
C      2H,IDK1,IDK2,IWATER,IQRE,IP,JP,IQ,JQ,IK,JK,K5,IPU1,IPU2,ITK
C      COMMON /SPARAM/ TMAX,CDLT,DELT,ERR,TEST,SUM,SUMP,QR
C      COMMON /SARRAY/ ICHK(13),LEVEL1(9),LEVEL2(9)
C      COMMON /CK/ ETFLXT,STORT,QRET,CHST,CHDT,FLUXT,PUMPT,CFLUXT,FLXNT
C      COMMON /PR/ XLABEL(3),YLABEL(6),TITLE(6),XN1,MESUR,PRNT(122),BLANKDAT0220
C      1(60),DIGIT(122),VF1(6),VF2(6),VF3(7),XSCALE,DINCH,SYM(17),XN(100),DAT0230
C      2YN(13),NA(4),N1,N2,N3,YSCALE,FACT1,FACT2
C      RETURN                                                    DAT0240
C                                                                DAT0250
C      .....DAT0260
C      *****DAT0270
C      ENTRY DATAIN
C      *****DAT0280
C                                                                DAT0290
C                                                                DAT0300
C      ---READ AND WRITE SCALAR PARAMETERS---
C      READ (5,330) NPER,KTH,ERR,LENGTH
C                                                                DAT0320

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	WRITE (6,340) NPER,KTH,ERR	DAT0330
	READ (5,460) XSCALE,YSCALE,DINCH,FACT1,(LEVEL1(I),I=1,9),FACT2,(LEDA0340	
	LEVEL2(I),I=1,9),MESUR	DAT0350
	IF (XSCALE.NE.0.) WRITE (6,470) XSCALE,YSCALE,MESUR,MESUR,DINCH,FA0360	
	ICT1,LEVEL1,FACT2,LEVEL2	DAT0370
C		DAT0380
C	---READ CUMULATIVE MASS BALANCE PARAMETERS---	DAT0390
	READ (5,450) SUM,SUMP,PUMPT,CFLUXT,QRET,CHST,CHDT,FLUXT,STORT,ETFL0400	
	IXT,FLXNT	DAT0410
	IF (IDK1.EQ.ICHK(4)) GO TO 20	DAT0420
	IF (IPUL.NE.ICHK(8)) GO TO 50	DAT0430
C		DAT0440
C	---READ INITIAL HEAD VALUES FROM CARDS---	DAT0450
	DO 10 K=1,K0	DAT0460
	DO 10 I=1,I0	DAT0470
	10 READ (5,360) (PHI(I,J,K),J=1,J0)	DAT0480
	GO TO 30	DAT0490
C		DAT0500
C	---READ INITIAL HEAD AND MASS BALANCE PARAMETERS FROM DISK---	DAT0510
	20 READ (4) PHI,SUM,SUMP,PUMPT,CFLUXT,QRET,CHST,CHDT,FLUXT,STORT,ETFL0520	
	IXT,FLXNT	DAT0530
	REWIND 4	DAT0540
	30 WRITE (6,430) SUM	DAT0550
	DO 40 K=1,K0	DAT0560
	WRITE (6,440) K	DAT0570
	DO 40 I=1,I0	DAT0580
	40 WRITE (6,350) I,(PHI(I,J,K),J=1,J0)	DAT0590
C		DAT0600
	50 DO 60 K=1,K0	DAT0610
	DO 60 I=1,I0	DAT0620
	DO 60 J=1,J0	DAT0630
	WELL(I,J,K)=0.	DAT0640
	TR(I,J,K)=0.	DAT0650
	TC(I,J,K)=0.	DAT0660
	IF (K.NE.K0) TK(I,J,K)=0.	DAT0670
	60 CONTINUE	DAT0680
	RETURN	DAT0690
C	*****	DAT0700
	ENTRY ARRAY(A,INFT,IOFT,IN,IRN,TF)	DAT0710
C	*****	DAT0720
	READ (5,330) FAC,IVAR,IPRN,TF,IRECS,IREC0	DAT0730
	IRLX=IFIL(IRN)	
	IC=4*IRECS+2*IVAR+IPRN+1	DAT0740
	GO TO (70,70,90,90,120,120), IC	DAT0750
	70 DO 80 I=1,I0	DAT0760
	DO 80 J=1,J0	DAT0770
	80 A(I,J)=FAC	DAT0780
	WRITE (6,280) IN,FAC,K	DAT0790
	GO TO 140	DAT0800
	90 IF (IC.EQ.3) WRITE (6,290) IN,K	DAT0810
	DO 110 I=1,I0	DAT0820
	READ(IRLX,INFT) (A(I,J),J=1,J0)	
	DO 100 J=1,J0	DAT0840
	100 A(I,J)=A(I,J)*FAC	DAT0850
	110 IF (IC.EQ.3) WRITE (6,IOFT) I,(A(I,J),J=1,J0)	DAT0860
	GO TO 140	DAT0870
	120 READ (2,IRN) A	DAT0880
	IF (IC.EQ.6) GO TO 140	DAT0890
	WRITE (6,290) IN,K	DAT0900
	DO 130 I=1,I0	DAT0910

130	WRITE (6,IOfT) I,(A(I,J),J=1,J0)	DAT0920
140	IF (IRECD.EQ.1) WRITE (2,IRN) A	DAT0930
	IRN=IRN+1	DAT0940
	WRITE (13) A	
	RETURN	DAT0950
C	*****	DAT0960
	ENTRY Mdat	DAT0970
C	*****	DAT0980
	DO 150 K=1,K0	DAT0990
	DO 150 I=1,I0	DAT1000
	DO 150 J=1,J0	DAT1010
	IF (I.EQ.1.OR.I.EQ.I0.OR.J.EQ.1.OR.J.EQ.J0) T(I,J,K)=0.	DAT1020
	IF (IDK1.NE.ICHK(4).AND.IPUL.NE.ICHK(8)) PHI(I,J,K)=STRT(I,J,K)	DAT1030
	IF (K.NE.K0.OR.IWATER.NE.ICHK(6)) GO TO 150	DAT1040
	IF (I.EQ.1.OR.I.EQ.I0.OR.J.EQ.1.OR.J.EQ.J0) PERM(I,J)=0.	DAT1050
150	CONTINUE	DAT1060
C DELX	DAT1070
	READ (5,330) FAC,IVAR,IPRN	DAT1080
	IF (IVAR.EQ.1) READ (5,330) (DELX(J),J=1,J0)	DAT1090
	DO 170 J=1,J0	DAT1100
	IF (IVAR.NE.1) GO TO 160	DAT1110
	DELX(J)=DELX(J)*FAC	DAT1120
	GO TO 170	DAT1130
160	DELX(J)=FAC	DAT1140
170	CONTINUE	DAT1150
	IF (IVAR.EQ.1.AND.IPRN.NE.1) WRITE (6,370) (DELX(J),J=1,J0)	DAT1160
	IF (IVAR.EQ.0) WRITE (6,300) FAC	DAT1170
C DELY	DAT1180
	READ (5,330) FAC,IVAR,IPRN	DAT1190
	IF (IVAR.EQ.1) READ (5,330) (DELY(I),I=1,I0)	DAT1200
	DO 190 I=1,I0	DAT1210
	IF (IVAR.NE.1) GO TO 180	DAT1220
	DELY(I)=DELY(I)*FAC	DAT1230
	GO TO 190	DAT1240
180	DELY(I)=FAC	DAT1250
190	CONTINUE	DAT1260
	IF (IVAR.EQ.1.AND.IPRN.NE.1) WRITE (6,380) (DELY(I),I=1,I0)	DAT1270
	IF (IVAR.EQ.0) WRITE (6,310) FAC	DAT1280
C DELZ	DAT1290
	READ (5,330) FAC,IVAR,IPRN	DAT1300
	IF (IVAR.EQ.1) READ (5,330) (DELZ(K),K=1,K0)	DAT1310
	DO 210 K=1,K0	DAT1320
	IF (IVAR.NE.1) GO TO 200	DAT1330
	DELZ(K)=DELZ(K)*FAC	DAT1340
	GO TO 210	DAT1350
200	DELZ(K)=FAC	DAT1360
210	CONTINUE	DAT1370
	IF (IVAR.EQ.1.AND.IPRN.NE.1) WRITE (6,390) (DELZ(K),K=1,K0)	DAT1380
	IF (IVAR.EQ.0) WRITE (6,320) FAC	DAT1390
C		DAT1400
C	---INITIALIZE VARIABLES---	DAT1410
	B=0.	DAT1420
	D=0.	DAT1430
	F=0.	DAT1440
	H=0.	DAT1450
	SU=0.	DAT1460
	Z=0.	DAT1470
	IF (XSCALE.NE.0.) CALL MAP	DAT1480
	RETURN	DAT1490
	ENTRY Ddat(ND)	

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      NK=1
      DO 500 I=1,I0
      READ(11,510) (ID(I,J),J=1,J0)
510  FORMAT (80I1)
      DO 500 J=1,J0
      IF (ID(I,J).EQ.0) GO TO 500
      ID(I,J)=NK
      NK=NK+1
500  CONTINUE
      NK=NK-1
      IF (NK.EQ.ND) GO TO 520
      PRINT 515,NK,ND
515  FORMAT ('  ERROR***NK.NE.ND      NK=',I5,5X,'ND=',I5)
      STOP
520  READ(11,330)FAC
      READ(11,530) (LD(I),I=1,ND)
530  FORMAT(40F2.0)
      PRINT 540,(LD(I),I=1,ND)
540  FORMAT(/,(20(1X,F5.0)))
      DO 550 I=1,I0
      DO 550 J=1,J0
      K=ID(I,J)
      IF (K.EQ.0) GO TO 550
      LD(K)=LD(K)*FAC*(I,J,K)/(DELX(J)*DELY(I))
550  CONTINUE
      READ(11,330)FAC
      READ(11,560) (ELD(I),I=1,ND)
      PRINT 540,(ELD(I),I=1,ND)
560  FORMAT (20F4.0)
      DO 570 I=1,ND
570  ELD(I)=FLD(I)*FAC
      RETURN
      ENTRY DDAT2(NRIV)
      NK=1
      DO 580 I=1,I0
      READ(12,510) (IDR(I,J),J=1,J0)
      DO 580 J=1,J0
      IF (IDR(I,J).EQ.0) GO TO 580
      IDR(I,J)=NK
      NK=NK+1
580  CONTINUE
      NK=NK-1
      IF (NK.EQ.NRIV) GO TO 600
      PRINT 585,NK,NRIV
585  FORMAT ('  ERROR***NK.NE.NRIV      NK=',I5,5X,'NRIV=',I5)
      STOP
600  READ(12,330)FAC
      READ(12,560) (RH(I),I=1,NRIV)
      PRINT 540,(RH(I),I=1,NRIV)
      DO 610 I=1,NRIV
610  RH(I)=RH(I)*FAC
      READ(12,330)FAC
      READ(12,560) (RB(I),I=1,NRIV)
      PRINT 540,(RB(I),I=1,NRIV)
      DO 620 I=1,NRIV
620  RB(I)=RB(I)*FAC
      READ(12,330)FAC
      READ(12,625) (RC(I),I=1,NRIV)
625  FORMAT(10F8.0)
      PRINT 635,(RC(I),I=1,NRIV)

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      DO 630 I=1,NRIV
630  WC(I)=RC(I)*FAC
635  FORMAT(/(14(1X,F8.0)))
      RETURN
C      .....DAT1500
C      ---READ TIME PARAMETERS AND PUMPING DATA FOR A NEW PUMPING PERIOD---DAT1510
C      *****DAT1520
C      ENTRY NEWPER DAT1530
C      *****DAT1540
C      DAT1550

      READ (5,330) KP,KPM1,NWEL,TMAX,NUMT,CDLT,DELT,IRECH
      IF (IRECH.EQ.1) READ(3) GRE

C      DAT1570
C      ---COMPUTE ACTUAL DELT AND NUMT--- DAT1580
      DT=DELT/24. DAT1590
      TM=0.0 DAT1600
      DO 220 I=1,NUMT DAT1610
      DT=CDLT*DT DAT1620
      TM=TM+DT DAT1630
      IF (TM.GE.TMAX) GO TO 230 DAT1640
220  CONTINUE DAT1650
      GO TO 240 DAT1660
230  DELT=TMAX/TM*DELT DAT1670
      NUMT=I DAT1680
240  WRITE (6,400) KP,TMAX,NUMT,DELT,CDLT DAT1690
      DELT=DELT*3600. DAT1700
      TMAX=TMAX*86400. DAT1710
      SUMP=0.0 DAT1720
C      DAT1730
C      ---READ AND WRITE WELL PUMPING RATES--- DAT1740
      WRITE (6,410) NWEL DAT1750
      IF (NWEL.EQ.0) GO TO 260 DAT1760
      DO 250 J=1,NWEL DAT1770
      READ (5,330) K,I,J,WELL(I,J,K) DAT1780
      WRITE (6,420) K,I,J,WELL(I,J,K) DAT1790
250  WELL(I,J,K)=WELL(I,J,K)/(DELT(J)*DELT(I)) DAT1800
260  RETURN DAT1810
C      DAT1820
C      ---FORMATS--- DAT1830
C      DAT1840
C      DAT1850
C      DAT1860
280  FORMAT (1H0,52X,6A4,' =',G15.7,' FOR LAYER',I3) DAT1870
290  FORMAT (1H1,45X,6A4,' MATRIX, LAYER',I3/46X,41(' -')) DAT1880
300  FORMAT ('0',72X,'DELX =',G15.7) DAT1890
310  FORMAT ('0',72X,'DELY =',G15.7) DAT1900
320  FORMAT ('0',72X,'DELZ =',G15.7) DAT1910
330  FORMAT (A610.0) DAT1920
340  FORMAT ('0',51X,'NUMBER OF PUMPING PERIODS =',I5/49X,'TIME STEPS BDAT1930
      1ETWEEN PRINTOUTS =',I5//51X,'ERROR CRITERIA FOR CLOSURE =',G15.7//) DAT1940
350  FORMAT ('0',12,2X,20F6.1/(5X,20F6.1)) DAT1950
360  FORMAT (A610.4) DAT1960
370  FORMAT (1H1,46X,40HGRID SPACING IN PROTOTYPE IN X DIRECTION/47X,40) DAT1970
      1(' -')//('0',12F10.0)) DAT1980
380  FORMAT (1H-,46X,40HGRID SPACING IN PROTOTYPE IN Y DIRECTION/47X,40) DAT1990
      1(' -')//('0',12F10.0)) DAT2000
390  FORMAT (1H-,46X,40HGRID SPACING IN PROTOTYPE IN Z DIRECTION/47X,40) DAT2010
      1(' -')//('0',12F10.0)) DAT2020
400  FORMAT ('-',50X,'PUMPING PERIOD NO.',I4,':',F10.2,' DAYS'/51X,38(' DAT2030
      1-')//53X,'NUMBER OF TIME STEPS=',I6//59X,'DELT IN HOURS =',F10.3//) DAT2040

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253X,'MULTIPLIER FOR DELT ='F10.3)                                DAT2050
410 FORMAT ('-',63X,I4,' WFLLS'/65X,9('-'//50X,'K',9X,'I',9X,'J' PUDAT2060
    IMPING RATE'/)                                                DAT2070
420 FORMAT (41X,3I10.2F13.2)                                       DAT2080
430 FORMAT ('-',40X,' CONTINUATION - HEAD AFTER ',G20.7,' SEC PUMPING DAT2090
    1'/42X,5X('-'//)                                              DAT2100
440 FORMAT ('1',55X,'INITIAL HEAD MATRIX, LAYER',I3/56X,30('-'//) DAT2110
450 FORMAT (4G20.10)                                               DAT2120
460 FORMAT (3G10.0,2(G10.0,9I1.1X),A8)                            DAT2130
470 FORMAT ('0',30X,'ON ALPHAMERIC MAP:'/40X,'MULTIPLICATION FACTOR FODAT2140
    1R X DIMENSION ='G15.7/40X,'MULTIPLICATION FACTOR FOR Y DIMENSION DAT2150
    2='G15.7/55X,'MAP SCALE IN UNITS OF ',A11/50X,'NUMBER OF ',A8,' P DAT2160
    3FR INCH ='G15.7/43X,'MULTIPLICATION FACTOR FOR DRAWDOWN ='G15.7,DAT2170
    4' PRINTED FOR LAYERS',9I2/47X,'MULTIPLICATION FACTOR FOR HEAD ='G DAT2180
    515.7,' PRINTED FOR LAYERS',9I2)                               DAT2190
    END                                                            DAT2200-
    SURROUTINE CHECKI(PHI,STRT,OLD,T,S,TR,TC,TK,WELL,DELX,DELY,DELZ,FACHK 10
    1CT,JFLO,FLOW,QRE,IO,FLD,FLD,IDR,RH,RC,RB,IDRAIN,IRIV)
C -----CHK 30
C COMPUTE A VOLUMETRIC BALANCE                                     CHK 40
C -----CHK 50
C                                                                    CHK 60
C SPECIFICATIONS:                                                 CHK 70
C REAL *8PHI                                                       CHK 80
C                                                                    CHK 90
C DIMENSION PHI(I0,J0,K0), STRT(I0,J0,K0), OLD(I0,J0,K0), T(I0,J0,K0)CHK 100
    1), S(I0,J0,K0), TR(I0,J0,K0), TC(I0,J0,K0), TK(IK,JK,K5), WELL(I0,CHK 110
    2J0,K0), DELX(J0), DELY(I0), DELZ(K0), FACT(K0,3), JFLO(NCH,3), FLOCHK 120
    3W(NCH), QRE(I0,J0)                                           CHK 130
    4,ID(I0,J0),FLD(1),FLD(1),IDR(I0,J0),RH(1),RC(1),RB(1)
C                                                                    CHK 140
C COMMON /INTEGR/ IO,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NCHK 150
    1WEL,NUMT,IFINAL,IT,KT,IHEAD,IDRAW,IFLO,IERR,I2,J2,K2,IMAX,ITMX1,NCCHK 160
    2H,IDK1,INK2,IWATER,IGRE,IP,JP,IQ,JQ,IK,K5,IPU1,IPU2,ITK
C COMMON /SPARAM/ TMAX,CDLT,DELT,ERR,TEST,SUM,SUMP,QR             CHK 170
C COMMON /SARRAY/ ICHK(13),LEVEL1(9),LEVEL2(9)                   CHK 180
C COMMON /CK/ ETFLXT,STORT,QRET,CHST,CHDT,FLUXT,PUMPT,CFLUXT,FLXNT CHK 190
    RETURN                                                         CHK 200
C .....CHK 210
C *****CHK 220
C ENTRY CHECK                                                       CHK 230
C *****CHK 240
C ---INITIALIZE VARIABLES---                                       CHK 250
C PUMP=0.                                                           CHK 260
C STOR=0.                                                           CHK 270
C FLUXS=0.0                                                         CHK 280
C CHD1=0.0                                                          CHK 290
C CHD2=0.0                                                          CHK 300
C QREFLX=0.                                                         CHK 310
C CFLUX=0.                                                         CHK 320
C FLUX=0.                                                           CHK 330
C ETFLUX=0.                                                         CHK 340
C FLXN=0.0                                                         CHK 350
C II=0                                                             CHK 360
C .....CHK 370
C .....CHK 380
C .....CHK 390
C ---COMPUTE RATES,STORAGE AND PUMPAGE FOR THIS STEP---          CHK 400
C DO 220 K=1,K0                                                     CHK 410
C DO 220 I=2,I1                                                     CHK 420
C DO 220 J=2,J1                                                     CHK 430

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	IF (T(I,J,K).EQ.0.) GO TO 220	CHK 440
	AREA=DELX(J)*DELY(I)	CHK 450
	IF (S(I,J,K).GE.0.) GO TO 180	CHK 460
C		CHK 470
C	---COMPUTE FLOW RATES TO AND FROM CONSTANT HEAD BOUNDARIES---	CHK 480
	II=II+1	CHK 490
	FLOW(II)=0.	CHK 500
	JFLO(II,1)=K	CHK 510
	JFLO(II,2)=I	CHK 520
	JFLO(II,3)=J	CHK 530
	IF (S(I,J-1,K).LT.0..OR.T(I,J-1,K).EQ.0.) GO TO 30	CHK 540
	X=(PHI(I,J,K)-PHI(I,J-1,K))*TR(I,J-1,K)*DELY(I)	CHK 550
	FLOW(II)=FLOW(II)+X	CHK 560
	IF (X) 10,30,20	CHK 570
10	CHD1=CHD1+X	CHK 580
	GO TO 30	CHK 590
20	CHD2=CHD2+X	CHK 600
30	IF (S(I,J+1,K).LT.0..OR.T(I,J+1,K).EQ.0.) GO TO 60	CHK 610
	X=(PHI(I,J,K)-PHI(I,J+1,K))*DELY(I)*TR(I,J,K)	CHK 620
	FLOW(II)=FLOW(II)+X	CHK 630
	IF (X) 40,60,50	CHK 640
40	CHD1=CHD1+X	CHK 650
	GO TO 60	CHK 660
50	CHD2=CHD2+X	CHK 670
60	IF (K.EQ.1) GO TO 90	CHK 680
	IF (S(I,J,K-1).LT.0..OR.T(I,J,K-1).EQ.0.) GO TO 90	CHK 690
	X=(PHI(I,J,K)-PHI(I,J,K-1))*TK(I,J,K-1)*AREA*2./(DELZ(K)+DELZ(K-1))	CHK 700
	1) FLOW(II)=FLOW(II)+X	CHK 710
	IF (X) 70,90,80	CHK 720
70	CHD1=CHD1+X	CHK 730
	GO TO 90	CHK 740
80	CHD2=CHD2+X	CHK 750
90	IF (K.EQ.K0) GO TO 120	CHK 760
	IF (S(I,J,K+1).LT.0..OR.T(I,J,K+1).EQ.0.) GO TO 120	CHK 770
	X=(PHI(I,J,K)-PHI(I,J,K+1))*TK(I,J,K)*AREA*2./(DELZ(K)+DELZ(K+1))	CHK 780
	FLOW(II)=FLOW(II)+X	CHK 790
	IF (X) 100,120,110	CHK 800
100	CHD1=CHD1+X	CHK 810
	GO TO 120	CHK 820
110	CHD2=CHD2+X	CHK 830
120	IF (S(I-1,J,K).LT.0..OR.T(I-1,J,K).EQ.0.) GO TO 150	CHK 840
	X=(PHI(I,J,K)-PHI(I-1,J,K))*TC(I-1,J,K)*DELX(J)	CHK 850
	FLOW(II)=FLOW(II)+X	CHK 860
	IF (X) 130,150,140	CHK 870
130	CHD1=CHD1+X	CHK 880
	GO TO 150	CHK 890
140	CHD2=CHD2+X	CHK 900
150	IF (S(I+1,J,K).LT.0..OR.T(I+1,J,K).EQ.0.) GO TO 220	CHK 910
	X=(PHI(I,J,K)-PHI(I+1,J,K))*TC(I,J,K)*DELX(J)	CHK 920
	FLOW(II)=FLOW(II)+X	CHK 930
	IF (X) 160,220,170	CHK 940
160	CHD1=CHD1+X	CHK 950
	GO TO 220	CHK 960
170	CHD2=CHD2+X	CHK 970
	GO TO 220	CHK 980
C		CHK 990
C	---RECHARGE AND WELLS---	CHK1000
180	IF (K.EQ.K0.AND.IQRE.EQ.ICHECK(7)) GREFLX=GREFLX+QRE(I,J)*AREA	CHK1010
	IF (WELL(I,J,K)) 190,210,200	CHK1020
		CHK1030

190	PUMP=PUMP+WELL(I,J,K)*AREA	CHK1040
	GO TO 210	CHK1050
200	CFLUX=CFLUX+WELL(I,J,K)*AREA	CHK1060
C		CHK1070
	---COMPUTE VOLUME FROM STORAGE---	CHK1080
210	STOR=STOR+S(I,J,K)*(OLD(I,J,K)-PHI(I,J,K))*AREA	CHK1090
	IF(K.NE.K0.OR.IRIV.LE.0) GO TO 212	
C	COMPUTE LFAKAGE TO RIVER	
	ND=IDR(I,J)	
	IF(ND.EQ.0) GO TO 212	
	IF(PHI(I,J,K).GT.RB(ND)) GO TO 211	
	ETFLUX=RC(ND)*(RH(ND)-RB(ND))*AREA+ETFLUX	
	GO TO 212	
211	ETFLUX=RC(ND)*(RH(ND)-PHI(I,J,K))*AREA+ETFLUX	
212	IF(K.NE.K0.OR.IDRAIN.LE.0) GO TO 220	
C	COMPUTE LEAKAGE TO DRAIN	
	ND=ID(I,J)	
	IF(ND.EQ.0) GO TO 220	
	IF(ELD(ND).GT.PHI(I,J,K)) GO TO 220	
	FLUX=FLUX+FLD(ND)*AREA*(ELD(ND)-STRT(I,J,K))	
	FLXN=FLXN+FLD(ND)*AREA*(ELD(ND)-PHI(I,J,K))	
	FLUXS=FLXN	
220	CONTINUE	CHK1100
C	CHK1110
C		CHK1120
C	---COMPUTE CUMULATIVE VOLUMES, TOTALS, AND DIFFERENCES---	CHK1130
	FLXPT=0.0	CHK1140
	FLXNT=FLXNT-FLXN*DELT	
	ETFLXT=ETFLXT-ETFLUX*DELT	
	STORT=STORT+STOR	CHK1150
	STOR=STOR/DELT	CHK1160
	QRET=QRET+QREFLX*DELT	CHK1170
	CHDT=CHDT-CHD1*DELT	CHK1180
	CHST=CHST+CHD2*DELT	CHK1190
	PUMPT=PUMPT-PUMP*DELT	CHK1200
	TOTL1=STORT+QRET+CFLUXT+CHST+FLXPT	CHK1220
	CFLUXT=CFLUXT+CFLUX*DELT	CHK1210
	TOTL2=CHDT+PUMPT+ETFLXT+FLXNT	CHK1230
	SUMR=QREFLX+CFLUX+CHD2+CHD1+PUMP+ETFLUX+FLUXS+STOR	CHK1240
	DIFF=TOTL2-TOTL1	CHK1250
	PERCNT=0.0	CHK1260
	IF (TOTL2.EQ.0.) GO TO 230	CHK1270
	PERCNT=DIFF/TOTL2*100.	CHK1280
230	RETURN	CHK1290
C	CHK1300
C		CHK1310
C	---PRINT RESULTS---	CHK1320
C	*****	CHK1330
	ENTRY CWRITE	CHK1340
C	*****	CHK1350
C		CHK1360
	WRITE (6,260) STOR,QREFLX,STORT,CFLUX,QRET,PUMP,CFLUXT,ETFLUX,CHST	CHK1370
	1,FLXPT,CHD2,TOTL1,CHD1,FLUX,FLUXS,ETFLXT,CHDT,SUMR,PUMPT,FLXNT,TOTL	CHK1380
	2L2,DIFF,PERCNT	CHK1390
	IF (NCH.EQ.0) GO TO 240	CHK1400
	WRITE (6,270)	CHK1410
	WRITE (6,280) ((JFLO(I,J),J=1,3),FLOW(I),I=1,NCH)	CHK1420
C		CHK1430
C	---COMPUTE VERTICAL FLOW---	CHK1440
240	X=0.	CHK1450

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      Y=0.                                CHK1460
      IF (K0.EQ.1) RETURN                  CHK1470
      DO 250 I=2,I1                        CHK1480
      DO 250 J=2,J1                        CHK1490
      X=X+(PHI(I,J,1)-PHI(I,J,2))*TK(I,J,1)*DELX(J)*DELY(I)*2./(DELZ(1)+CHK1500
      1DELZ(2))                            CHK1510
250  Y=Y+(PHI(I,J,K1)-PHI(I,J,K0))*TK(I,J,K1)*DELX(J)*DELY(I)*2./(DELZ(CHK1520
      1K1)+DELZ(K0))                      CHK1530
      WRITE (6,290) Y,X                   CHK1540
      RETURN                               CHK1550
C                                          CHK1560
C      ---FORMATS---                      CHK1570
C                                          CHK1580
C      -----CHK1590
C                                          CHK1600
C                                          CHK1610
C                                          CHK1620
260  FORMAT ('0',10X,'CUMULATIVE MASS BALANCE:',16X,'L**3',23X,'RATES FCHK1630
      1OR THIS TIME STEP:',16X,'L**3/T',11X,24(' '),43X,25(' ')//20X,'SOUCHK1640
      2RCES:',69X,'STORAGE =',F20.4/20X,8(' '),68X,'RECHARGE =',F20.4/27XCHK1650
      3,'STORAGF =',F20.2,35X,'CONSTANT FLUX =',F20.4/26X,'RECHARGE =',F20CHK1660
      40.2,41X,'PUMPING =',F20.4/21X,'CONSTANT FLUX =',F20.2,30X,'      RI
      5VER LEAKAGE =',F20.4/21X,'CONSTANT HEAD =',F20.2,34X,'CONSTANT HEA
      6D:',27X,'LEAKAGE =',F20.2,46X,'IN =',F20.4/21X,'TOTAL SOURCES =',FCHK1690
      720.2,45X,'OUT =',F20.4/96X,'LEAKAGE:',20X,'DISCHARGES:',45X,'FROM CHK1700
      8PREVIOUS PUMPING PERIOD =',F20.4/20X,11(' '),68X,'TOTAL =',F20.4/1CHK1710
      96X,'      RIVER LEAKAGE =',F20.2/21X,'CONSTANT HEAD =',F20.2,36X,'S
      SUM OF RATES =',F20.4/19X'QUANTITY PUMPED =',F20.2/27X,'LEAKAGE =',CHK1730
      $F20.2/19X,'TOTAL DISCHARGE =',F20.2//17X,'DISCHARGE-SOURCES =',F20CHK1740
      $.2/15X,'PER CENT DIFFERENCE =',F20.2//)      CHK1750
270  FORMAT ('0FLOW RATES TO CONSTANT HEAD NODES:',' ',34(' ')//' ',3(9CHK1760
      1X,'K',4X,'I',4X,'J',5X,'RATE (L**3/T)')//' ',3(9X,'-',4X,'-',4X,'-'CHK1770
      2,5X,13(' '))//)      CHK1780
280  FORMAT ('/(1X,3(I10,2I5,G18.7)))      CHK1790
290  FORMAT ('0FLOW TO TOP LAYER =,G15.7,'      FLOW TO BOTTOM LAYER =',GCHK1800
      115.7,'      POSITIVE UPWARD')      CHK1810
      END                                  CHK1820-
      SUBROUTINE STEP(PHI,STRT,OLD,T,S,TR,TC,TK,WELL,DELX,DELY,DELZ,FACTSTP 10
      1,DDN,TEST3)                        STP 20
C      -----STP 30
C      INITIALIZE DATA FOR A NEW TIME STEP AND PRINT RESULTS      STP 40
C      -----STP 50
C                                          STP 60
C      SPECIFICATIONS:      STP 70
      REAL *8PHI      STP 80
      REAL *8XLABEL,YLABEL,TITLE,XN1,MFSUR      STP 90
C                                          STP 100
      DIMENSION PHI(I0,J0,K0), STRT(I0,J0,K0), OLD(I0,J0,K0), T(I0,J0,K0STP 110
      1), S(I0,J0,K0), TR(I0,J0,K0), TC(I0,J0,K0), TK(IK,JK,K5), WELL(I0,STP 120
      2J0,K0), DELX(J0), DELY(I0), DELZ(K0), FACT(K0,3), DDN(IMAX), TEST3STP 130
      3(ITMX1), ITT0(50)      STP 140
C                                          STP 150
      COMMON/INDC/ IND(80)
      COMMON /INTEGR/ I0,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NSTP 160
      1WEL,NUMT,IFINAL,IT,KT,IHEAD,IDPAW,IFLO,IERR,I2,J2,K2,IMAX,ITMX1,NCSTP 170
      2H,IDX1,IDX2,IWATER,IQHE,IP,JP,IQ,JQ,IK,JK,K5,IPU1,IPU2,ITK      STP 180
      COMMON /SPARAM/ TMAX,CDLT,DELT,ERR,TEST,SUM,SUMP,QR      STP 190
      COMMON /SARRAY/ ICHK(13),LEVEL1(9),LEVEL2(9)      STP 200
      COMMON /CK/ FTFLXT,STOPT,QRET,CHST,CHDT,FLTXT,PUMPT,CFLTXT,FLXNT      STP 210
      COMMON /PR/ XLABEL(3),YLABEL(6),TITLE(6),XN1,MESUR,PRNT(122),BLANKSTP 220

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	1(60),DIGIT(122),VF1(6),VF2(6),VF3(7),XSCALE,DINCH,SYM(17),XN(100),STP	230
	2YN(13),NA(4),N1,N2,N3,YSCALE,FACT1,FACT2	STP 240
	RETURN	STP 250
C	*****	STP 260
C	*****	STP 270
	ENTRY NEWSTP	STP 280
C	*****	STP 290
	KT=KT+1	STP 300
	IT=0	STP 310
	DO 10 K=1,K0	STP 320
	DO 10 I=1,I0	STP 330
	DO 10 J=1,J0	STP 340
10	OLD(I,J,K)=PHI(I,J,K)	STP 350
	DELT=CDLT*DELT	STP 360
	SUM=SUM+DELT	STP 370
	SUMP=SUMP+DELT	STP 380
	DAYS=SUMP/86400.	STP 390
	YRSP=DAYS/365.	STP 400
	HRS=SUMP/3600.	STP 410
	SMIN=HRS*60.	STP 420
	DAYS=HRS/24.	STP 430
	YRS=DAYS/365.	STP 440
	RETURN	STP 450
C		STP 460
C	---PRINT OUTPUT AT DESIGNATED TIME STEPS---	STP 470
C	*****	STP 480
	ENTRY OUTPUT	STP 490
C	*****	STP 500
	IF (KT.EQ.NUMT) IFINAL=1	STP 510
	ITTO(KT)=IT	STP 520
	IF (IT.LE.ITMAX) GO TO 20	STP 530
	IT=IT-1	STP 540
	ITTO(KT)=IT	STP 550
	IEHR=2	STP 560
C		STP 570
C	---IF MAXIMUM ITERATIONS EXCEEDED,WRITE RESULTS ON DISK OR CARDS---	STP 580
	IF (IDK2.EQ.ICHK(5)) WRITE (4) PHI,SUM,SUMP,PUMPT,CFLUXT,QRET,CHST	STP 590
	1,CHDT,FLUXT,STORT,ETFLXT,FLXNT	STP 600
	IF (IPU2.EQ.ICHK(9)) WRITE (7,230) SUM,SUMP,PUMPT,CFLUXT,QRET,CHST	STP 610
	1,CHDT,FLUXT,STORT,ETFLXT,FLXNT	STP 620
C		STP 630
20	IF (IFLO.EQ.ICHK(3)) CALL CHECK	STP 640
	IF (IERR.EQ.2) GO TO 30	STP 650
	IF (MOD(KT,KTH).NE.0.AND.IFINAL.NE.1) RETURN	STP 660
30	WRITE (6,210) KT,DELT,SUM,SMIN,HRS,DAYS,YRS,DAYS,SP,YRSP	STP 670
	IF (IFLO.EQ.ICHK(3)) CALL CWRITE	STP 680
	IT=IT+1	STP 690
	WRITE (6,180) (TEST3(J),J=1,IT)	STP 700
	I3=1	STP 701
	I5=0	STP 702
352	I5=I5+40	STP 703
	I4=MIN0(KT,I5)	STP 704
	WRITE (6,240) (I,I=I3,I4)	STP 710
	WRITE (6,260)	STP 720
	WRITE (6,250) (ITTO(I),I=I3,I4)	STP 730
	WRITE (6,260)	STP 740
	IF(KT.LE.I5) GO TO 353	STP 741
	I3=I3+40	STP 742
	GO TO 352	STP 743
C		STP 750

C	--- <td>STP 760</td>	STP 760
353	IF (XSCALE.EQ.0.) GO TO 70	STP 770
	IF (IDD(KP).EQ.0) GO TO 70	
	IF (FACT1.EQ.0.) GO TO 50	STP 780
	DO 40 IA=1,9	STP 790
	II=LEVEL1(IA)	STP 800
	IF (II.EQ.0) GO TO 50	STP 810
40	CALL PRNTA(1,II)	STP 820
50	IF (FACT2.EQ.0.) GO TO 70	STP 830
	DO 60 IA=1,9	STP 840
	II=LEVEL2(IA)	STP 850
	IF (II.EQ.0) GO TO 70	STP 860
60	CALL PRNTA(2,II)	STP 870
70	IF (IDRAW.NE.ICHK(1)) GO TO 100	STP 880
C		STP 890
C	--- <td>STP 900</td>	STP 900
	DO 90 K=1,K0	STP 910
	WRITE (6,200) K	STP 920
	DO 90 I=1,I0	STP 930
	DO 80 J=1,J0	STP 940
80	DDN(J)=STRT(I,J,K)-PHI(I,J,K)	STP 950
90	WRITE (6,170) I,(DDN(J),J=1,J0)	STP 960
100	CONTINUE	
	DO 103 K=1,K0	
103	WRITE (15) ((PHI(I,J,K),J=1,J0),I=1,I0)	
	IF (IMEAD.NE.ICHK(2)) GO TO 120	
C		STP 980
C	--- <td>STP 990</td>	STP 990
	IF (IDD(KP).EQ.0) GO TO 120	
	NKDKL=4-IDD(KP)	
	DO 110 K=NKDKL,K0	
	WRITE (6,190) K	STP1010
	DO 110 I=1,I0	STP1020
110	WRITE (6,170) I,(PHI(I,J,K),J=1,J0)	STP1030
C		STP1040
C	--- <td>STP1050</td>	STP1050
120	IF (IERR.EQ.2) GO TO 130	STP1060
	IF (KP.LT.NPEN.OR.IFINAL.NE.1) RETURN	STP1070
	IF (IDK2.EQ.ICHK(5)) WRITE (44) PHI,SUM,SUMP,PUMPT,CFLUXT,QRET,CHS	
	1T,CHDT,FLUXT,STORT,ETFLXT,FLXNT	
C		STP1100
C	--- <td>STP1110</td>	STP1110
130	IF (IPU2.NE.ICHK(9)) GO TO 160	STP1120
	IF (IERR.EQ.2) GO TO 140	STP1130
	WRITE (7,230) SUM,SUMP,PUMPT,CFLUXT,QRET,CHST,CHDT,FLUXT,STORT,ETFLXT,FLXNT	STP1140
	140 DO 150 K=1,K0	STP1150
	DO 150 I=1,I0	STP1160
150	WRITE (7,220) (PHI(I,J,K),J=1,J0)	STP1170
160	IF (IERR.EQ.2) STOP	STP1180
	RETURN	STP1190
C		STP1200
C	--- <td>STP1210</td>	STP1210
C		STP1220
C		STP1230
C		STP1240
170	FORMAT ('0',I4,1AF7.2/(5X,1BF7.2))	STP1250
180	FORMAT ('0MAXIMUM HEAD CHANGE FOR EACH ITERATION:'/' ' ,39('-'')/('0	STP1260
	1',10F12.4))	STP1270
190	FORMAT ('1',55X,'HEAD MATRIX, LAYER',I3/56X,21('-''))	STP1280


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200 FORMAT (10,55X,0 DRAWDOWN, LAYER,13/59X,18(0-0)) STP1300
210 FORMAT (1H1,40X,57(0-0)/45X,01,14X,0 TIME STEP NUMBER =,19,14X,0ISTP1300
1/45X,57(0-0)/50X,2HMSIZE OF TIME STEP IN SECONDS=,F14.2/55X,0TOSTP1310
2TAL SIMULATION TIME IN SECONDS=,F14.2/80X,8HMINUTES=,F14.2/82X,6HSTP1320
3HOURS=,F14.2/83X,5H DAYS=,F14.2/82X,0YEARS=,F14.2///45X,0DURATION STP1330
4OF CURRENT PUMPING PERIOD IN DAYS=,F14.2/82X,0YEARS=,F14.2//) STP1340
220 FORMAT (10F8.2) STP1350
230 FORMAT (4G20.10) STP1360
240 FORMAT (0TIME STEP 1,40I3) STP1370
250 FORMAT (0ITERATIONS:,40I3) STP1380
260 FORMAT (0,1U(0-0)) STP1390
END STP1400-
SUBROUTINE SOLVE(PHI,STRT,OLD,T,S,TR,TC,TK,WELL,DELX,DELY,DELZ,FACSP3 10
1T,EL,FL,GL,V,XI,TEST3,QRE,PERM,10,FLD,ELD,1DR,RH,RC,RB,1DRAIN,
2 IRIV)
C -----SP3 30
C SOLUTION BY THE STRONGLY IMPLICIT PROCEDURE SP3 40
C -----SP3 50
C SPECIFICATIONS: SP3 60
C REAL *8PHI,RHO,B,D,F,H,Z,SU,RHOP,W,WMIN,RH01,RH02,RH03,XPART,YPARTSP3 70
C REAL*8 UX,UXR SP3 80
1,ZPART,DMIN1,WMAX,XT,YT,ZT,DABS,DMAX1,DEN,TXM,TYM,TZM SP3 90
C REAL *8E,AL,AL,CL,A,C,G,WU,TU,U,DL,RES,SUPH,GLXI,ZPHI SP3 100
C DIMENSION PHI(1),STRT(1),OLD(1),T(1),S(1),TR(1),TC(1),TK(1)SP3 110
1,WELL(1),DELX(1),DELY(1),DELZ(1),FACT(K0,3),RHOP(20),TEST3(SP3 120
21),EL(1),FL(1),GL(1),V(1),XI(1),QRE(1) SP3 130
3,ID(1),FLD(1),ELD(1),PERM(1),1DR(1),RH(1),RC(1),RB(1) SP3 140
C COMMON /INTEGR/ 10,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NSP3 150
1WEL,NUMT,IFINAL,IT,KT,IHEAD,1DRAW,IFLO,IERR,I2,J2,K2,IMAX,ITMX1,NCSP3 160
2H,1DK1,1DK2,1WATER,1QRE,IP,JP,IQ,JQ,IK,JK,K5,IPU1,IPU2,ITK SP3 170
COMMON /SPARAM/ TMAX,CDLT,DELT,ERR,TEST,SUM,SUMP,QR SP3 180
COMMON /RESMUN/ RESMO SP3 190
COMMON /SARRAY/ ICHK(13),LEVEL1(9),LEVEL2(9) SP3 200
RETURN SP3 210
C .....SP3 220
C ***** SP3 230
C ENTRY ITER SP3 240
C ***** SP3 250
C ---COMPUTE AND PRINT ITERATION PARAMETERS--- SP3 260
C WRITE (6,240) SP3 270
C P2=LENGTH-1 SP3 300
C NT=10*J0*K0 SP3 310
C NIJ=10*J0 SP3 320
C PJ=-1. SP3 780
C READ(5,250) WMAX
250 FORMAT(F10.0)
DO 50 I=1,LENGTH SP3 790
PJ=PJ+1. SP3 800
50 RHOP(I)=1.D0-(1.D0-WMAX)**(PJ/P2) SP3 810
WRITE (6,230) LENGTH,(RHOP(J),J=1,LENGTH) SP3 820
RETURN SP3 830
C .....SP3 840
C SP3 850
C ---INITIALIZE DATA FOR A NEW ITERATION--- SP3 8 0
C 60 IT=IT+1 SP3 0
IF (IT.LE.ITMAX) GO TO 70 SP3 880
WRITE (6,220) SP3 890

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	CALL OUTPUT	SP3 900
70	IF (MOD(IT,LENGTH)) 80,80,90	SP3 910
C	*****	SP3 920
C	NTH=0	SP3 950
	ENTRY NEWITA	SP3 930
C	*****	SP3 940
90	NTH=NTH+1	SP3 960
	W=RHOP(NTH)	SP3 970
	TEST3(IT+1)=0.	SP3 980
	TEST=0.0	SP3 990
	HIG=0.	SP31000
	DO 100 I=1,NT	SP31010
	EL(I)=0.	SP31020
	FL(I)=0.	SP31030
	GL(I)=0.	SP31040
	G=CL*FL(NJB)	
	WU=CL*GL(NJB)	
	U=BL*GL(NIB)	
	IF (K.EQ.1) GO TO 140	
	AL=Z/(1.+W*(EL(NKB)+FL(NKB)))	
	A=AL*EL(NKB)	SP31540
	TU=AL*FL(NKB)	SP31550
	DL=E+W*(A+C+G+WU+TU+U)-CL*EL(NJB)-BL*FL(NIB)-AL*GL(NKB)	SP31560
	EL(N)=(F-W*(A+C))/DL	SP31570
	FL(N)=(H-W*(G+TU))/DL	SP31580
	GL(N)=(SU-W*(WU+U))/DL	SP31590
	SUPH=0.00	SP31600
	IF (K.NE.K0) SUPH=SU*PHI(NKA)	SP31610
	RES=-B*PHI(NIB)-D*PHI(NJB)-E*PHI(N)-F*PHI(NJA)-H*PHI(NIA)-SUPH-Z*P	SP31620
	WI(NKB)-WELL(N)-RHO*OLD(N)-QR	SP31630
	RES=RES*RESMO	
	V(N)=(RES-AL*V(NKB)-BL*V(NIB)-CL*V(NJB))/DL	SP31640
	GO TO 150	SP31650
140	DL=E+W*(C+G+WU+U)-CL*EL(NJB)-BL*FL(NIB)	SP31660
	EL(N)=(F-W*C)/DL	SP31670
	FL(N)=(H-W*G)/DL	SP31680
	GL(N)=(SU-W*(WU+U))/DL	SP31690
	SUPH=0.00	SP31700
	IF (K.NE.K0) SUPH=SU*PHI(NKA)	SP31710
	RES=-B*PHI(NIB)-D*PHI(NJB)-E*PHI(N)-F*PHI(NJA)-H*PHI(NIA)-SUPH-WEL	SP31720
	IL(N)-RHO*OLD(N)-QR	SP31730
	RES=RES*RESMO	
	V(N)=(RES-AL*V(NIB)-CL*V(NJB))/DL	SP31740
150	CONTINUE	SP31750
C		SP31760
C	---BACK SUBSTITUTE FOR VECTOR XI---	SP31770
	DO 160 K=1,K0	SP31780
	K3=K0-K+1	SP31790
	DO 160 I=1,I2	SP31800
	I3=I0-I	SP31810
	DO 160 J=1,J2	SP31820
	J3=J0-J	SP31830
	N=I3+(J3-1)*I0+(K3-1)*NIJ+I-I	SP31840
	IF (T(N).EQ.0..OR.S(N).LT.0.) GO TO 160	SP31850
	GLXI=0.00	SP31860
	IF (K3.NE.K0) GLXI=GL(N)*XI(N+NIJ)	SP31870
	XI(N)=V(N)-EL(N)*XI(N+I0)-FL(N)*XI(N+1)-GLXI	SP31880
C		SP31890
C	---COMPARE MAGNITUDE OF CHANGE WITH CLOSURE CRITERIA---	SP31900
	TCHK=ABS(XI(N))	SP31910

	IF (TCHK.GT.RIG) BIG=TCHK	SP31920
	PHI(N)=PHI(N)+XI(N)	SP31930
160	CONTINUE	SP31940
	IF (RIG.GT.FRH) TEST=1.	SP31950
	TEST3(IT+1)=BIG	SP31960
	IF (TEST.EQ.0.) RETURN	SP31970
	GO TO 60	SP31980
C	SP31990
170	DO 200 KK=1,K0	SP32000
	K=K0-KK+1	SP32010
	DO 200 II=1,I _c	SP32020
	I=I0-II	SP32030
	DO 200 J=2,J1	SP32040
	N=I+(J-1)*I0+(K-1)*NIJ	SP32050
	NIA=N+1	SP32060
	NIB=N-1	SP32070
	NJA=N+I0	SP32080
	NJB=N-I0	SP32090
	NKA=N+NIJ	SP32100
	NKB=N-NIJ	SP32110
C		SP32120
C	---SKIP COMPUTATIONS IF NODE OUTSIDE AQUIFER---	SP32130
	IF (T(N).EQ.0..OR.S(N).LT.0.) GO TO 200	SP32140
C		SP32150
C	---COMPUTE COEFFICIENTS---	SP32160
	D=TR(NJB)/DELX(J)	SP32170
	F=TR(N)/DELX(J)	SP32180
	R=TC(NIB)/DELY(I)	SP32190
	H=TC(N)/DELY(I)	SP32200
	SU=0.D0	SP32210
	Z=0.D0	SP32220
	IF (K.NE.1) Z=TK(NKB)/DELZ(K)	SP32230
	IF (K.NE.K0) SU=TK(N)/DELZ(K)	SP32240
	RHO=S(N)/DFLT	SP32250
	QR=0.	SP32260
	UXR=0.	
	UX=0.	
	IF (K.NE.K0) GO TO 180	SP32270
	IF (IQRE.EQ.ICHK(7)) QR=QRE(I+(J-1)*I0)	SP32280
	IF (IRIV.LE.0) GO TO 175	
	ND=IDR(I+(J-1)*I0)	
	IF (ND.EQ.0) GO TO 175	
	IF (PHI(N).GT.RB(ND)) GO TO 174	
	QR=QR+RC(ND)*(RH(ND)-RB(ND))	
	GO TO 175	
174	UXR=RC(ND)	
	QR=QR+RC(ND)*RH(ND)	
175	IF (IDRAIN.LE.0) GO TO 180	
	ND=ID(I+(J-1)*I0)	
	IF (ND.EQ.0) GO TO 180	
	IF (FLD(ND).GT.PHI(N)) GO TO 180	
	UX=FLD(ND)	
	QR=QR+FLD(ND)*ELD(ND)	
C		SP32290
C	---SIP REVERSE ALGORITHM---	SP32300
C	---FORWARD SUBSTITUTE, COMPUTING INTERMEDIATE VECTOR V---	SP32310
180	E=-R-H-D-F-SU-Z-RHO-UX-UXR	
	RL=H/(1.+W*(EL(NIA)+GL(NIA)))	SP32330
	CL=D/(1.+W*(FL(NJB)+GL(NJB)))	SP32340
	C=BL*EL(NIA)	SP32350

G=CL*FL(NJB)	SP32360
WU=CL*GL(NJB)	SP32370
U=HL*GL(NIA)	SP32380
IF (K.EQ.K0) GO TO 190	SP32390
AL=SU/(1.+W*(EL(NKA)*FL(NKA)))	SP32400
A=AL*EL(NKA)	SP32410
TU=AL*FL(NKA)	SP32420
DL=F+W*(C+G+A+WU+TU+U)-AL*GL(NKA)-BL*FL(NIA)-CL*EL(NJB)	SP32430
FL(N)=(F-W*(C+A))/DL	SP32440
FL(N)=(B-W*(G+TU))/DL	SP32450
GL(N)=(Z-W*(WU+U))/DL	SP32460
ZPHI=0.D0	SP32470
IF (K.NE.1) ZPHI=Z*PHI(NKB)	SP32480
RES=-B*PHI(NIB)-D*PHI(NJB)-E*PHI(N)-F*PHI(NJA)-H*PHI(NIA)-SU*PHI(N)	SP32490
1KA)-ZPHI-WELL(N)-RHO*OLD(N)-QR	SP32500
RES=RES+RESMO	
V(N)=(RES-AL*V(NKA)-RL*V(NIA)-CL*V(NJB))/DL	SP32510
GO TO 200	SP32520
190 DL=E+W*(C+G+WU+U)-BL*FL(NIA)-CL*EL(NJB)	SP32530
FL(N)=(F-W*C)/DL	SP32540
FL(N)=(B-W*G)/DL	SP32550
GL(N)=(Z-W*(WU+U))/DL	SP32560
ZPHI=0.D0	SP32570
IF (K.NE.1) ZPHI=Z*PHI(NKB)	SP32580
RES=-B*PHI(NIB)-D*PHI(NJB)-E*PHI(N)-F*PHI(NJA)-H*PHI(NIA)-ZPHI-WEL	SP32590
1L(N)-RHO*OLD(N)-QR	SP32600
RES=RES+RESMO	
V(N)=(RES-RL*V(NIA)-CL*V(NJB))/DL	SP32610
200 CONTINUE	SP32620
C	SP32630
C ---RACK SUBSTITUTE FOR VECTOR XI---	SP32640
DO 210 K=1,K0	SP32650
DO 210 I=2,I1	SP32660
DO 210 J=1,J2	SP32670
J3=J0-J	SP32680
N=I+(J3-1)*I0+(K-1)*NIJ	SP32690
IF (T(N).EQ.0..OR.S(N).LT.0.) GO TO 210	SP32700
GLXI=0.D0	SP32710
IF (K.NE.1) GLXI=GL(N)*XI(N-NIJ)	SP32720
XI(N)=V(N)-EL(N)*XI(N+I0)-FL(N)*XI(N-1)-GLXI	SP32730
C	SP32740
C ---COMPARE MAGNITUDE OF CHANGE WITH CLOSURE CRITERIA---	SP32750
TCHK=ABS(XI(N))	SP32760
IF (TCHK.GT.BIG) BIG=TCHK	SP32770
PHI(N)=PHI(N)+XI(N)	SP32780
210 CONTINUE	SP32790
IF (BIG.GT.ERR) TEST=1.	SP32800
TEST3(IT+1)=BIG	SP32810
IF (TEST.EQ.0.) RETURN	SP32820
GO TO 60	SP32830
C	SP32840
C	SP32850
C	SP32860
C ---FORMATS---	SP32870
C	SP32880
C	SP32890
220 FORMAT ('0FCEEDED PERMITTED NUMBER OF ITERATIONS'/' ',39(' '))	SP32900
230 FORMAT ('//1H0,I5,22H ITERATION PARAMETERS:6E15.7/(/28X,6E15.7/))	SP32910
240 FORMAT ('-',44X,'SOLUTION BY THE STRONGLY IMPLICIT PROCEDURE'/45X,SP32920	
143(' '))	SP32930

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END
SUBROUTINE COEF(PHI,STRT,OLD,T,S,TR,TC,TK,WELL,DELX,DELY,DELZ,FACTCOF 10
1,PERM,BOTTOM,QRE)
C -----COF 30
C COMPUTE COEFFICIENTS
C -----COF 50
C SPECIFICATIONS:
C REAL *8PHI
C
C DIMENSION PHI(I0,J0,K0), STRT(I0,J0,K0), OLD(I0,J0,K0), T(I0,J0,K0)COF 100
1), S(I0,J0,K0), TR(I0,J0,K0), TC(I0,J0,K0), TK(IK,JK,K5), WELL(I0,COF 110
2J0,K0), DELX(J0), DELY(I0), DELZ(K0), FACT(K0,3), PERM(IP,JP), BOTCOF 120
3TOM(IP,JP), QRE(IQ,JQ)
C
C COMMON /INTFGR/ I0,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NCOF 150
1WEL,NUMT,IFINAL,IT,KT,IHEAD,IDRAW,IFLO,IFERR,I2,J2,K2,IMAX,ITMX1,NCCOF 160
2H,INDK1,INDK2,IWATER,IGRE,IP,JP,IQ,JQ,IK,JK,K5,IPU1,IPU2,ITK
C COMMON /SPARAM/ TMAX,CDLT,DELT,FPR,TEST,SUM,SUMP,QR
C COMMON /SAPRAY/ ICHK(13),LEVEL1(9),LEVFL2(9)
C RETURN
C
C -----COF 210
C ---COMPUTE TRANSMISSIVITY FOR UPPER HYDROLOGIC UNIT WHEN
C IT IS UNCONFINED---
C *****
C ENTRY TRANS(N3)
C *****
C
C DO 10 I=2,I1
C DO 10 J=2,J1
C IF (PERM(I,J).EQ.0.) GO TO 10
C T(I,J,K0)=PERM(I,J)*(PHI(I,J,K0)-BOTTOM(I,J))
C IF (T(I,J,K0).GT.0.) GO TO 10
C IF (WELL(I,J,K0).LT.0.) WRITE (6,60) I,J,K0
C IF (WELL(I,J,K0).GE.0.) WRITE (6,70) I,J,K0
C WRITE (6,180) KT,PERM(I,J),PHI(I,J,K0),BOTTOM(I,J)
180 FORMAT ('-',I5,E15.5,E15.7,E15.7)
C PERM(I,J)=0.
C T(I,J,K0)=0.
C TR(I,J-1,K0)=0.
C TR(I,J,K0)=0.
C TC(I,J,K0)=0.
C TC(I-1,J,K0)=0.
C IF (K0.NE.1) TK(I,J,K1)=0.
C PHI(I,J,K0)=1.D30
10 CONTINUE
C IF (N3.EQ.1) RETURN
C N1=K0
C N2=K0
C N4=K1
C GO TO 20
C
C ---COMPUTE T COEFFICIENTS---
C *****
C ENTRY TCOF
C *****
C N1=1
C N2=K0
C N4=1
20 DO 40 K=N1,N2
C DO 40 I=1,I1
C DO 40 J=1,J1

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      IF (T(I,J,K).EQ.0.) GO TO 40                                COF 580
      IF (T(I,J+1,K).EQ.0.) GO TO 30                             COF 590
      TR(I,J,K)=(2.*T(I,J+1,K)*T(I,J,K))/(T(I,J,K)*DELX(J+1)+T(I,J+1,K)*COF 600
      1DELX(J))*FACT(K,1)                                         COF 610
30  IF (T(I+1,J,K).EQ.0.) GO TO 40                                COF 620
      TC(I,J,K)=(2.*T(I+1,J,K)*T(I,J,K))/(T(I,J,K)*DELY(I+1)+T(I+1,J,K)*COF 630
      1DELY(I))*FACT(K,2)                                         COF 640
40  CONTINUE                                                       COF 650
      IF (K0.EQ.1.OR.ITK.EQ.ICLK(10)) RETURN                     COF 660
      DO 50 K=N4,K1                                                COF 670
      DO 50 I=2,I1                                                  COF 680
      DO 50 J=2,J1                                                  COF 690
      IF (T(I,J,K+1).EQ.0.) GO TO 50                               COF 700
      T1=T(I,J,K)*FACT(K,3)                                         COF 710
      T2=T(I,J,K+1)*FACT(K+1,3)                                     COF 720
      TK(I,J,K)=(2.*T2*T1)/(T1*DELZ(K+1)+T2*DELZ(K))             COF 730
50  CONTINUE                                                       COF 740
      RETURN                                                        COF 750
C                                                                    COF 760
C                                                                    COF 770
60  FORMAT ('-1,20(1*),'WELL',2I3,' IN LAYER',I3,' GOES DRY',20(1*))COF 780
70  FORMAT ('-1,20(1*),'NODE',2I3,' IN LAYER',I3,' GOES DRY',20(1*))COF 790
      END                                                            COF 800-
      SUBROUTINE PRNTAI(PHI,STRT,T,S,WELL,DELX,DELY)              PRN 10
C -----PRN 20
C PRINT MAPS OF DRAWDOWN AND HYDRAULIC HEAD                      PRN 30
C -----PRN 40
C                                                                    PRN 50
C SPECIFICATIONS:                                                PRN 60
      REAL *8PHI,Z,XLABEL,YLABEL,TITLE,XN1,MESUR                 PRN 70
      REAL *4K                                                    PRN 80
C                                                                    PRN 90
      DIMENSION PHI(I0,J0,K0), STRT(I0,J0,K0), S(I0,J0,K0), WELL(I0,J0,KPRN 100
10), DELX(J0), DELY(I0), T(I0,J0,K0)                             PRN 110
      DIMENSION SYM2(10)
C                                                                    PRN 120
      COMMON /INTEGR/ I0,J0,K0,I1,J1,K1,I,J,K,NPER,KTH,ITMAX,LENGTH,KP,NPRN 130
1WEL,NUMT,IFINAL,IT,KT,IHEAD,IDRAW,IFLO,IERR,I2,J2,K2,IMAX,ITMX1,NCPRN 140
2H,IDK1,IDK2,IWATER,IGRE,IP,JP,IQ,JQ,IK,JK,K5,IPU1,IPU2,ITK      PRN 150
      COMMON /PR/ XLABEL(3),YLABEL(6),TITLE(6),XN1,MESUR,PRNT(122),BLANKPRN 160
1(60),DIGIT(122),VF1(6),VF2(6),VF3(7),XSCALE,DINCH,SYM(17),XN(100),PRN 170
2YN(13),NA(4),N1,N2,N3,YSCALE,FACT1,FACT2                       PRN 180
      DATA SYM2/'A','B','C','D','E','F','G','H','I','J'/
      RETURN                                                       PRN 190
C .....PRN 200
C .....PRN 210
C ---INITIALIZE VARIABLES FOR PLOT---                             PRN 220
C *****PRN 230
      ENTRY MAP                                                    PRN 240
C *****PRN 250
      YDIM=0.                                                       PRN 260
      WIDTH=0.                                                      PRN 270
      DO 10 J=2,J1                                                  PRN 280
10  WIDTH=WIDTH+DELX(J)                                           PRN 290
      DO 20 I=2,I1                                                  PRN 300
20  YDIM=YDIM+DELY(I)                                             PRN 310
30  XSF=DINCH*XSCALE                                              PRN 320
      YSF=DINCH*YSCALE                                             PRN 330
      NYD=YDIM/YSF                                                 PRN 340
      IF (NYD*YSF.LE.YDIM-DELY(I1)/2.) NYD=NYD+1                PRN 350

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	IF (NYD.LE.12) GO TO 40	PRN 360
	DINCH=YDIM/(12.*YSCALE)	PRN 370
	WRITE (6,330) DINCH	PRN 380
	IF (YSCALE.LT.1.0) WRITE (6,340)	PRN 390
	GO TO 30	PRN 400
40	NXD=WIDTH/XSF	PRN 410
	IF (NXD*XSF.LE.WIDTH-DELX(J1)/2.) NXD=NXD+1	PRN 420
	N4=NXD*N1+1	PRN 430
	N5=NXD+1	PRN 440
	N6=NYD+1	PRN 450
	N8=N2*NYD+1	PRN 460
	NA(1)=N4/2-1	PRN 470
	NA(2)=N4/2	PRN 480
	NA(3)=N4/2+3	PRN 490
	NC=(N3-N8-10)/2	PRN 500
	ND=NC+N8	PRN 510
	NE=MAX0(N5,N6)	PRN 520
	VF1(3)=DIGIT(ND)	PRN 530
	VF2(3)=DIGIT(ND)	PRN 540
	VF3(3)=DIGIT(NC)	PRN 550
	XLABEL(3)=MESUR	PRN 560
	YLABEL(6)=MESUR	PRN 570
	DO 60 I=1,NE	PRN 580
	NNX=N5-I	PRN 590
	NNY=I-1	PRN 600
	IF (NNY.GE.N6) GO TO 50	PRN 610
	YN(I)=YSF*NNY/YSCALE	PRN 620
50	IF (NNX.LT.0) GO TO 60	PRN 630
	XN(I)=XSF*NNX/YSCALE	PRN 640
60	CONTINUE	PRN 650
	RETURN	PRN 660
C	PRN 670
C		PRN 680
C	*****	PRN 690
	ENTRY PRNTA(NG,LA)	PRN 700
C	*****	PRN 710
C	---VARIABLES INITIALIZED EACH TIME A PLOT IS REQUESTED---	PRN 720
	DIST=WIDTH-DELX(J1)/2.	PRN 730
	JJ=J1	PRN 740
	LL=1	PRN 750
	Z=NXD*XSF	PRN 760
	IF (NG.EQ.1) WRITE (6,300) (TITLE(I),I=1,3),LA	PRN 770
	IF (NG.EQ.2) WRITE (6,300) (TITLE(I),I=4,6),LA	PRN 780
	DO 290 I=1,N4	PRN 790
C		PRN 800
C	---LOCATE X AXES---	PRN 810
	IF (I.EQ.1.OR.I.EQ.N4) GO TO 70	PRN 820
	PRNT(1)=SYM(12)	PRN 830
	PRNT(N8)=SYM(12)	PRN 840
	IF ((I-1)/N1*N1.NE.I-1) GO TO 90	PRN 850
	PRNT(1)=SYM(14)	PRN 860
	PRNT(N8)=SYM(14)	PRN 870
	GO TO 90	PRN 880
C		PRN 890
C	---LOCATE Y AXES---	PRN 900
70	DO 80 J=1,N8	PRN 910
	IF ((J-1)/N2*N2.EQ.J-1) PRNT(J)=SYM(14)	PRN 920
80	IF ((J-1)/N2*N2.NE.J-1) PRNT(J)=SYM(13)	PRN 930
C		PRN 940
C	---COMPUTE LOCATION OF NODES AND DETERMINE APPROPRIATE SYMBOL---	PRN 950

90	IF (DIST.LT.0..OR.DIST.LT.Z-XN1*XSF) GO TO 240	PRN 960
	YLEN=DELY(2)/2.	PRN 970
	DO 220 L=2,11	PRN 980
	J=YLEN*N2/YSF+1.5	PRN 990
	IF (T(L,JJ,LA).EQ.0.) GO TO 160	PRN1000
	IF (S(L,JJ,LA).LT.0.) GO TO 210	PRN1010
	INDX3=0	PRN1020
	GO TO (100,110), NG	PRN1030
100	K=(STRY(L,JJ,LA)-PHI(L,JJ,LA))*FACT1	PRN1040
C	-TO CYCLE SYMBOLS FOR DRAWDOWN, REMOVE C FROM COL. 1 OF NEXT CARD-	PRN1050
C	K=AMOD(K,10.)	PRN1060
	GO TO 120	PRN1070
110	K=PHI(L,JJ,LA)*FACT2	PRN1080
120	IF (K) 130,160,140	PRN1090
130	N=-K	
	N=MOD(N,10)+1	
	PRNT(J)=SYM2(N)	
	GO TO 220	
140	N=K	
	N=MOD(N,10)+1	
	PRNT(J)=SYM(N)	
	GO TO 220	
160	PRNT(J)=SYM(15)	
	IF (T(L,JJ,LA).EQ.0.) PRNT(J)=SYM(14)	
	GO TO 220	
210	IF (S(L,JJ,LA).LT.0.) PRNT(J)=SYM(16)	PRN1370
220	YLEN=YLEN+(DELY(L)+DELY(L+1))/2.	PRN1380
230	DIST=DIST-(DEFLX(JJ)+DEFLX(JJ-1))/2.	PRN1390
	JJ=JJ-1	PRN1400
	IF (JJ.EQ.0) GO TO 240	PRN1410
	IF (DIST.GT.Z-XN1*XSF) GO TO 230	PRN1420
240	CONTINUE	PRN1430
C		PRN1440
C	---PRINT AXES,LABELS, AND SYMBOLS---	PRN1450
	IF (I-NA(LL).EQ.0) GO TO 260	PRN1460
	IF ((I-1)/N1*N1-(I-1)) 270,250,270	PRN1470
250	WRITE (6,VF1) (BLANK(J),J=1,NC),(PRNT(J),J=1,N8),XN(1+(I-1)/6)	PRN1480
	GO TO 280	PRN1490
260	WRITE (6,VF2) (BLANK(J),J=1,NC),(PRNT(J),J=1,N8),XLABEL(LL)	PRN1500
	LL=LL+1	PRN1510
	GO TO 280	PRN1520
270	WRITE (6,VF2) (BLANK(J),J=1,NC),(PRNT(J),J=1,N8)	PRN1530
C		PRN1540
C	---COMPUTE NEW VALUE FOR Z AND INITIALIZE PRNT---	PRN1550
280	Z=Z-2.*XN1*XSF	PRN1560
	DO 290 J=1,N8	PRN1570
290	PRNT(J)=SYM(15)	PRN1580
C		PRN1590
C	---NUMBER AND LABEL Y AXIS AND PRINT LEGEND---	PRN1600
	WRITE (6,VF3) (BLANK(J),J=1,NC),(YN(I),I=1,N6)	PRN1610
	WRITE (6,320) (YLABEL(I),I=1,6)	PRN1620
	IF (NG.EQ.1) WRITE (6,310) FACT1	PRN1630
	IF (NG.EQ.2) WRITE (6,310) FACT2	PRN1640
	RETURN	PRN1650
C		PRN1660
C	---FORMATS---	PRN1670
C		PRN1680
C	-----	PRN1690
C		PRN1700
C		PRN1710

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300 FORMAT ('1',49X,3A8,'LAYER',I4//) PRN1720
310 FORMAT ('0EXPLANATION'/' ' ,11(' - '))/' R = CONSTANT HEAD BOUNDARY'/'PRN1730
1' *** = VALUE EXCEEDED 3 FIGURES'/' MULTIPLICATION FACTOR =' ,F8.3)PRN1740
320 FORMAT ('0',39X,6A8) PRN1750
330 FORMAT ('0',25X,10(' * '),' TO FIT MAP WITHIN 12 INCHES, DINCH REVISPRN1760
1ED TO',G15.7,1X,10(' * ')) PRN1770
340 FORMAT ('0',45X,'NOTE: GENERALLY SCALE SHOULD BE > OR = 1.0') PRN1780
END PRN1790-
BLOCK DATA BLK 10
----- BLK 20
C BLK 20
C BLK 30
C SPECIFICATIONS: BLK 40
REAL *8XLABEL,YLABEL,TITLE,XN1,MFSUR BLK 50
C BLK 60
COMMON /SARRAY/ ICHK(13),LEVEL1(9),LEVEL2(9) BLK 70
COMMON /PR/ XLABEL(3),YLABEL(6),TITLE(6),XN1,MESUR,PRNT(122),BLANKBLK 80
1(60),DIGIT(122),VF1(6),VF2(6),VF3(7),XSCALE,DINCH,SYM(17),XN(100),BLK 90
2YN(13),NA(4),N1,N2,N3,YSCALE,FACT1,FACT2 BLK 100
C ***** BLK 110
C BLK 120
DATA ICHK/'DRAW','HEAD','MASS','DK1','DK2','WATE','RECH','PUN1','PBLK 130
1UN2','ITKR',3*0/ BLK 140
DATA SYM/'1','2','3','4','5','6','7','8','9','10','*', '|', '-', '+', BLK 150
1' , 'R', 'W' / BLK 160
DATA PRNT/122*' / ,N1,N2,N3,XN1/6,10,133,.8333333333D-1/,BLANK/60*'BLK 170
1' / ,NA(4)/1000/ BLK 180
DATA XLABEL/' X DIS- ',TANCE IN', ' MTLFS' / ,YLABEL/'DISTANCE', ' BLK 190
1FROM OR', 'IGIN IN ', 'Y DIRECT', 'ION, IN ', 'MILES' / ,TITLE/'PLOT BLK 200
2OF ', 'DRAWDOWN', ' ', 'PLOT OF ', 'HYDRAULI', 'C HEAD' / BLK 210
DATA DIGIT/'1','2','3','4','5','6','7','8','9','10','11','12','13'BLK 220
1,'14','15','16','17','18','19','20','21','22','23','24','25','26',BLK 230
2,'27','28','29','30','31','32','33','34','35','36','37','38','39','BLK 240
340','41','42','43','44','45','46','47','48','49','50','51','52','58BLK 250
43','54','55','56','57','58','59','60','61','62','63','64','65','668BLK 260
5','67','68','69','70','71','72','73','74','75','76','77','78','79BLK 270
6','80','81','82','83','84','85','86','87','88','89','90','91','92'BLK 280
7,'93','94','95','96','97','98','99','100','101','102','103','104'BLK 290
8,'105','106','107','108','109','110','111','112','113','114','115'BLK 300
9,'116','117','118','119','120','121','122' / BLK 310
DATA VF1/'(1H ',',',', ' ', 'A1,F', '10.2', ')') / BLK 320
DATA VF2/'(1H ',',',', ' ', 'A1,l', 'X,A8', ')') / BLK 330
DATA VF3/'(1H0',',',', ' ', 'A1,F', '3.1', ' ', '12F1', '0.2', ')') / BLK 340
C ***** BLK 350
END BLK 360-

```

Data Preparation

Basic model operation requires four input files.

1. The basic data deck describes the finite difference grid, input and output options, the hydraulic properties of the aquifers, initial conditions and pumpage rates.
2. The recharge file contains recharge rates for each node in the top layer of the model for each recharge period. This file is read at the beginning of each recharge period for which there is a "1" punched in column 80 of the recharge period card.
3. The river data deck contains arrays indicating the nodes on the top layer which have rivers and specifying the leakage factor, elevation of streambed and elevation of the bottom of the streambed layer.
4. The drainage tile deck contains arrays indicating the nodes in the top layer which have drainage tiles and specifying the length and elevation of tile in each node.

Basic Data Deck Instructions

These instructions have been copied, with minor modifications from model documentation by Trescott (1975).

Group I: Title, Simulation Options and Problem Dimensions.--This group of cards, which is read by the main program, contains data required to dimension the model. To specify an option on card 7 punch the characters underlined in the definition. For an option not used, that section of the card 7 can be left blank.

Note: Default typing of variables applies for all data input.

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
1	1-80	20A4	HEADING	Any title the user wishes to print on one line at the start of output.
2	1-52	13A4	"	
3	1-56	14I4	IFIL	Symbolic unit numbers of file from which arrays normally part of the Basic Data file are to be read. Normally set to 5.
4	1-80	80I1	IDD	Number of layers (starting from the top) for which the potentiometric heads are to

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
				be printed. The column number in which the number is punched corresponds to the recharge period for which heads will be printed.
5	1-10	F10.0	RESMO	Multiplier for "RES" field in subroutine SOLVE.
6	1-10	I10	IO	Number of rows
	11-20	I10	JO	Number of columns
	21-30	I10	KO	Number of layers
	31-40	I10	ITMAX	Maximum number of iterations per time step
	41-50	I10	NCH	Number of constant head nodes
	51-60	I10	ND	Number of nodes containing a drain
	61-70	I10	NRIV	Number of nodes containing a river
7	1-4	A4	IDRAW	<u>DRAW</u> to print drawdown
	6-9	A4	IHEAD	<u>HEAD</u> to print hydraulic head
	11-14	A4	IFLOW	<u>MASS</u> to compute a mass balance
	16-18	A3	IDK1	<u>DK1</u> to read initial head, elapsed time, and mass balance parameters from unit 4 on disk
	21-23	A3	IDK2	<u>DK2</u> to write computed head, elapsed time, and mass balance parameters on unit 44 (disk)
	26-29	A4	IWATER	<u>WATE</u> if the upper hydrologic unit is unconfined
	31-34	A4	IQRE	<u>RECH</u> for a constant recharge that may be a function of space
	36-39	A4	IPU1	<u>PUN1</u> to read initial head, elapsed time, and mass balance parameters from cards

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
	41-44	A4	IPU2	<u>PUN2</u> to punch computed head, elapsed time, and mass balance parameters on cards
	46-49	A4	ITK	<u>ITKR</u> to read the value of TK(I,J,K) for simulations in which confining layers are not represented by layers of nodes (TK (I,J,K)) = K_{zz}/b

Group II: Scalar parameters.--The parameters required in every problem are underlined. The other parameters are required as noted; when not required, their location on the card can be left blank. The G format is used to read E, F and I format data. Minimize mistakes by always right-justifying data in the field. If F format data do not contain significant figures to the right of the decimal point, the decimal point can be omitted.

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
1	1-10	G10.0	<u>NPER</u>	Number of recharge periods for the simulation
	11-20	G10.0	<u>KTH</u>	Number of time steps between printouts

Note: To print only the results for the final time step in a pumping period, make KTH greater than the expected number of time steps. The program always prints the results for the final time step.

	21-30	G10.0	<u>ERR</u>	Error criteria for closure (L)
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Note: When the head change at all nodes on subsequent iterations is less than this value (for example, 0.01 foot), the program has converged to a solution for the time step.

	31-40	G10.0	<u>LENGTH</u>	Number of iteration parameters
	1-10	G10.0	XSCALE	Factor to convert model length unit to unit in X direction on maps (e.g. to convert from feet to miles, XSCALE = 5280).

For no maps, card 2 is blank

CARD	COLUMN	FORMAT	VARIABLE	DEFINITION
	11-20	G10.0	YSCALE	Factor to convert model length unit to unit used in Y direction on maps.
	21-30	G10.0	DINCH	Number of map units per inch
	31-40	G10.0	FACT1	Factor to adjust value of drawdown printed*
	41-49	9I1	LEVEL(I)	Layers for which drawdown maps are to be printed. List the layers starting in column 41; the first zero entry terminates the printing of drawdown maps.
2	51-60	G10.0	FACT2	Factor to adjust value of head printed*
	61-69	9I1	LEVEL2(I)	Layers for which head maps are to be printed. List layers starting in column 61; the first zero entry terminates the printing of head maps.
	71-78	A8	MESUR	Name of map length unit.
3	1-20	G20.10	SUM	Parameters in which elapsed times and cumulative volumes for mass balance are stored. For the start of a simulation insert three blank cards. <u>For continuation</u> of a previous run using cards as input, replace the three blank cards with the first three cards of punched output from the previous run. Using data from disk for input, leave the three blank cards in the data deck.
	21-40	G20.10	SUMP	
	41-60	G20.10	PUMPT	
	61-80	G20.10	CFLUXT	
4	1-20	G20.10	QRET	Parameters in which elapsed times and cumulative volumes for mass balance are stored. For the start of a simulation insert three blank cards. <u>For continuation</u> of a previous run using cards as input, replace the three blank cards with the first three cards of punched output from the previous run. Using data from disk for input, leave the three blank cards in the data deck.
	21-40	G20.10	CHST	
	41-60	G20.10	CHDT	
	61-80	G20.10	FLUXT	
5	1-20	G20.10	STORT	Parameters in which elapsed times and cumulative volumes for mass balance are stored. For the start of a simulation insert three blank cards. <u>For continuation</u> of a previous run using cards as input, replace the three blank cards with the first three cards of punched output from the previous run. Using data from disk for input, leave the three blank cards in the data deck.
	21-40	G20.10	ETFLXT	
	41-60	G20.10	FLXNT	

*Value of drawdown or head	FACT 1 FACT 2	or Printed value
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0.01	1	
	0.1	5
52.57	1.0	53
	10.0	526
	100.0	***

Group III: Array Data.--Each of the following data sets (except data set 1) consists of a parameter card and, if the data set contains variable data, a set of data cards for each layer in the model. Each parameter card contains at least five variables:

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
Every Parameter Card	1-10	G10.0	FAC	If IVAR = 0, FAC is the value assigned to every element of the matrix for this layer. If IVAR = 1, FAC is the multiplication factor for the following set of data cards for this layer.
	11-20	G10.0	IVAR	= 0 if no data cards are to be read for this year. = 1 if data cards for this layer follow.
	21-30	G10.0	IPRN	= 0 if input data for this layer are to be printed; = 1 if input data for the layer are <u>not</u> to be printed.
Trans- missivity Parameter cards also have these Variables	31-40	G10.0	FACT(K,1)	multiplication factor for transmissivity in x direction
	41-50	G10.0	FACT(K,2)	multiplication factor for transmissivity in the y direction

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
	51-60	G10.0	FACT(K,3)	multiplication factor for hydraulic conductivity in the z direction. (Not used when confining bed nodes are eliminated and TK values are read)
Every Parameter Card	61-70	G10.0	IRECS	= 0 if the matrix is being read from cards or if each element is being set equal to FAC. = 1 if the matrix is to be read from disk (unit 2)
	71-80	G10.0	IRECD	= 0 if the matrix is <u>not</u> to be stored on disk. = 1 if the matrix being read from cards or set equal to FAC <u>is</u> to be stored on disk (unit 2) for later retrieval.

<u>DATA SET</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
1	1-80	8F 10.4	PHI(I,J,K)	Head values for continuation of a previous run (L)
2	1-80	8F 10.4	STRT(I,J,K)	Starting head matrix (L)
3	1-80	20F 4.0	S (I,J,K)	Storage coefficient (dimensionless)

Note: This matrix is also used to locate constant head boundaries by coding a negative number at constant head nodes. At these nodes T must be greater than zero.

4	1-80	20F 4.0	T(I,J,K)	Transmissivity (L^2/t)
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Note 1) Zero values are required around the perimeter of the T matrix for each layer for reasons inherent in the computational scheme. This is done automatically by the program.

2) See the previous page for the additional requirements on the parameter cards for this data set.

DATA SET	COLUMNS	FORMAT	VARIABLE	DEFINITION
<p>Note 3) If the upper active layer is unconfined and PERM and BOTTOM are to be read for this layer, insert a parameter card <u>for this layer</u> with only the values for FACT on it.</p>				
5	1-80	20F 4.0	TK(I,J,K)	K_{zz}/b
<p>Note: This data set is read only if specified in the options. The number of layers of TK values = $K'=1$. See the discussion of the treatment of confining layers.</p>				
6	1-80	20F 4.0	PERM(I,J)	Hydraulic conductivity (L/T) (see note 1 for data set 4)
7	1-80	20F 4.0	BOTTOM(I,J)	Elevation of bottom of water-table unit (L)
<p>Note: Data sets 6 and 7 are required only for simulating unconfined conditions in the upper hydrologic unit.</p>				
8	1-80	20F 4.0	QRE(I,J)	Recharge rate (L/T) If recharge rate is read for each recharge period then this card is ignored and recharge rate is read from the Recharge File
<p>Note: Omit if not used.</p>				
9	1-80	8G10.0	DELX(J)	Grid spacing in x direction (L)
10	1-80	8G10.0	DELY(I)	Grid spacing in y direction (L)
11	1-80	8G10.0	DELZ(K)	Grid spacing in z direction (L)

Group IV: Parameters that change with the recharge period.--The Program has two options for the simulation period:

1. To simulate a given number of time steps, set TMAX to a value larger than the expected simulation period. The program will use NUMT, CDLT, and DELT as coded.

2. To simulate a given recharge period, set NUMT larger than the number required for the simulation period (for example, 100). The program will compute the exact DELT (which will be \leq DELT coded) and NUMT to arrive exactly at TMAX on the last time step.

<u>CARD</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
1	1-10	G10.0	<u>KP</u>	Number of the recharge period
	11-20	G10.0	<u>KPM1</u>	Number of the previous recharge period
Note: KPM1 is currently not used.				
	21-30	G10.0	<u>NWEL</u>	Number of wells for this recharge period
	31-40	G10.0	<u>TMAX</u>	Number of days in this recharge period
	41-50	G10.0	<u>NUMT</u>	Number of time steps
	51-60	G10.0	<u>CDLT</u>	Multiplying factor for DELT
Note: 1.5 is commonly used				
	61-70	G10.0	<u>DELT</u>	Initial time step in hours
	71-80	G10.0		Flag to indicate that recharge to the top layer is to be read from recharge file for each recharge period.

If NWEL: 0 the following set of cards is omitted

<u>DATA SET 1</u>		(NEWL cards)	
<u>COLUMNS</u>	<u>FORMAT</u>	<u>VARIABLE</u>	<u>DEFINITION</u>
1-10	G10.0	K	Layer in which well is located
11-20	G10.0	I	Row location of well
21-30	G10.0	J	Column location of well
31-40	G10.0	WELL(I,J,K)	Pumping rate (L ³ /t), negative for a pumping well.

For each recharge period only wells pumping at a different rate than the previous recharge period need to be entered.

Recharge File Instructions

The recharge file contains one record for each recharge period in the simulation. Each record contains a recharge rate per unit area for each node in the top layer of the grid. The values are in internal floating point notation in column major order.

River Data Deck

The river data deck consists of four arrays and three parameter cards.

<u>ARRAY</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>DEFINITION</u>
IDR(I,J)	1-80	80I1	Indicator array for cells containing a river, 1 for river 0 for no river.
RH(NRIV)	1-80	20F4.0	River water level for each of the NRIV cells containing a river. Entered in row major order.
RB(NRIV)	1-80	20F4.0	Bottom of streambed layer in each of the NRIV cells containing a river. Entered in row major order.
RC(NRIV)	1-80	10F8.0	Leakage factor for each of the NRIV cells containing a river. The factor for each cell is the percent area of the cell occupied by the hydraulic conductivity of the river bed and directed by the thickness of the river bed.

The RH, RB and RC arrays must each be preceded by a parameter card similar to those described for array data in the basic data deck.

Drainage Tile Deck

This file consists of three arrays and two parameter cards.

<u>ARRAY</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>DEFINITION</u>
ID(I,J)	1-80	80I1	Indicator array for cells containing drainage tile, 1 for tile, 0 for no tile.

<u>ARRAY</u>	<u>COLUMNS</u>	<u>FORMAT</u>	<u>DEFINITION</u>
LD(ND)	1-80	40F2.0	Length of drainage tile (in hundreds of feet) in the ND cells containing tile. Entered in row major order.

The LD and ELD arrays must each be preceded by a parameter card similar to those described for array data in the basic data deck.

* Several additional files are available for data input. They are described in detail in the model program documentation (Trescott 1975).

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