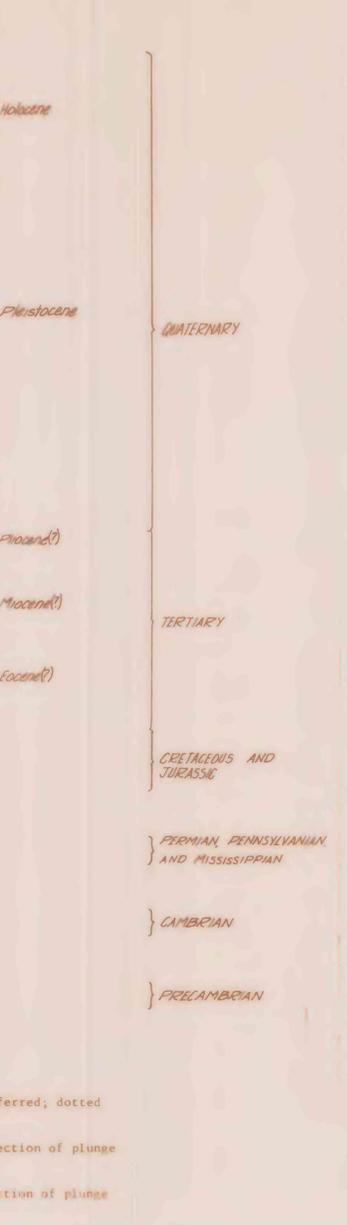
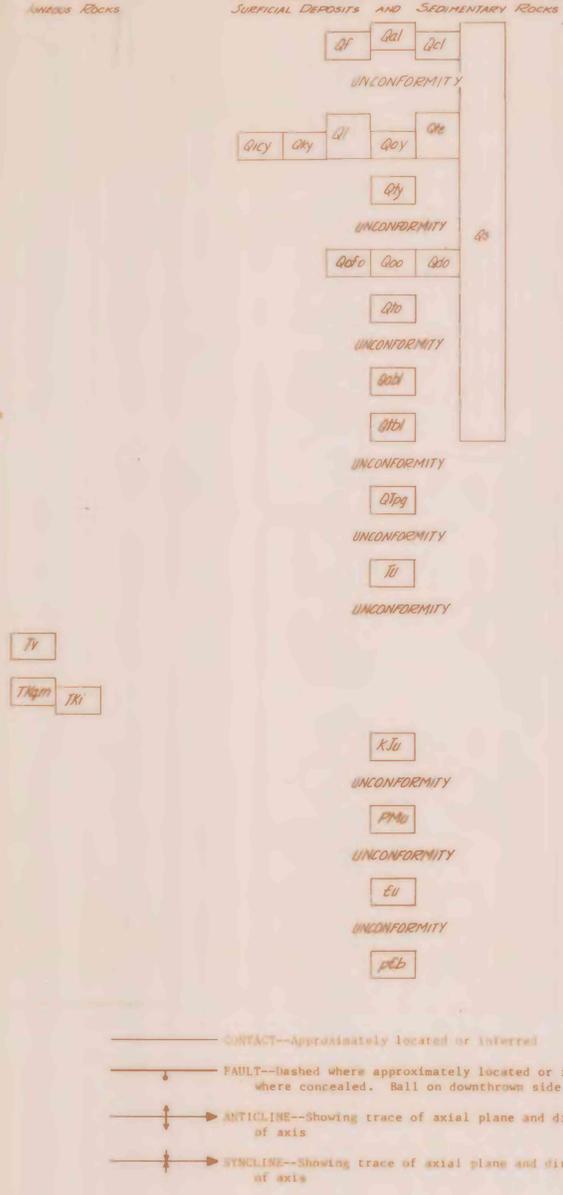


CORRELATION OF MAP UNITS



DESCRIPTION OF MAP UNITS

[This is a combined map description for quadrangles in the southern part of the Big Fork-Avon environmental study area (Open-File Reports 79-437 to 79-447). Not all units will appear on all maps.]

Qal ALLUVIUM (HOLOCENE)—Light to dark gray and brownish gray; stream-deposited, unconsolidated, moderately to well-bedded, interstratified assemblage of silt, sand, and gravel. Typically clay-rich and cohesive where incised into claystones, siltstones, and shales of older basin-fill deposits. Gold-bearing sand and gravel occur locally in the northward-draining valleys of the Garnet Range, and in the westward-draining valleys of the mountains east of Helmsville and Avon. Characterized by shallow depths to ground water. Soil drainage is poor and shrink-swell potentials are inferred to be relatively low.

Qa1 ALLUVIUM (HOLOCENE)—Small, moderately sloping, poorly sorted, crudely stratified silt, sand, gravel, cobbles, and boulders at the mouths of small gullies and high-gradient streams. Ground-water conditions are quite variable, and flash-flood hazard potential is high, especially near the upper apex of the fan. Shrink-swell potential is inferred to be generally low.

Qa2 COLLUVIUM (HOLOCENE)—Unsorted debris that has slid or been washed downslope to form small, thin, and hummocky deposits. Best developed in terrane underlain by Tertiary volcanic rocks. Shrink-swell potentials are variable, and are inferred to range from moderate to high.

Qa3 LANDSLIDE DEPOSITS (QUATERNARY)—Small, fan- and lobe-shaped hummocky masses of semiconsolidated soil, till, Tertiary basin fill, and volcanic debris. Clay and soil-moisture content are generally high. Small springs or seeps are common. Many of these landslides are active and all should be considered unstable. Shrink-swell potentials are variable, and are inferred to range from moderate to high.

SEDIMENTS DEPOSITED BY MELTWATERS OF PINEDALE GLACIATION (QUATERNARY)—YOUNGER ICE

Qa4 Lacustrine Deposits—Pink; rhythmically bedded silt, sandy silt, and clay with included lenses of sand and gravel. Forms gently sloping, smooth surfaces at altitudes generally below 1,158 m (3,800 ft). Soil drainage ranges from fair to poor, but lateral permeability may be high within the interbedded sandy beds and lenses. High frost-heave susceptibility. Clay-rich beds have a moderate shrink-swell potential.

Qa5 Ice-Contact Deposit—Dark brown and brown; composed of moderately well sorted silt, sand, and gravel. Forms an elongate, gently sloping bench. Clasts range in shape from angular to rounded; most are rounded. Sizes range from 3 mm to 20 cm (1/8-8 in.); dominant sizes range from 13 mm to 5 cm (1/2-2 in.). Cobbles and small angular to rounded boulders 0.6-1 m (2-3 ft) are scattered across the surface. Formed in contact with wasting ice.

Qa6 EASE DEPOSITS—Dark brown and brown; consists of moderately to poorly sorted silt, sand, and gravel. Forms small, conical to ellipsoidal, steep-sided hillocks. Few well-rounded cobbles scattered irregularly through the unit. Formed by a glacial stream that flowed down into a stagnant ice mass.

Qa7 Outwash—Light brown, and grayish brown; moderately well sorted; composed of silt, sand, and subrounded to rounded pebbly and cobbly gravel. Commonly the upper 1 m (3 ft) is composed of light-brown, fine- to medium-grained sand. Generally forms broad, smooth-surfaced plains which grade to the main valley of the Blackfoot River, and lie 1.5-6 m (5-20 ft) above the modern stream channel. Ground water commonly occurs at shallow depths.

Terrace Deposits—Gray, tan, brown, and pink; unconsolidated, moderately well sorted, well-bedded silt, sand, gravel, and cobbles in smooth-surfaced, dissected deposits flanking modern stream channels, and in abandoned channels 3-8 m (15-20 ft) above modern stream channels. These deposits are typically moderately to well-drained, and are inferred to have relatively low shrink-swell potential.

SEDIMENTS DEPOSITED BY MELTWATERS OF PINEDALE GLACIATION (UPPER PLEISTOCENE)—OLDER ICE

Qa8 Outwash—Dark brown, reddish brown, and brown; consists of moderately well sorted silt, sand, and subangular to rounded gravel and cobbly gravel. Forms smooth, locally dissected plains and terraces which lie 6-30 m (20-100 ft) above the younger outwash deposits. Ground water occurs at moderate depths.

Qa9 Outwash Fan Deposit—Even-surfaced, gently sloping, cone-shaped deposit of unconsolidated, moderately sorted silt, sand, gravel, and cobbles. Probably formed at some time after the Blackfoot River was diverted by lobes of ice which lay in the Monture, North Fork of the Blackfoot, and Nevada Creek valleys.

Qa10 Deltaic Gravels—Tan to brown; locally crossbedded, unconsolidated, well-sorted, medium- to coarse-grained sand, and gravel. Clasts range in shape from angular to round. Deposited in a small glacial lake formed where southward advancing ice impinged upon the northern flank of the Garnet Range blocking the northward flowing ancestral Pearson Creek (see Chamberlain Mountain quadrangle).

TILL OF PINEDALE GLACIATION (UPPER PLEISTOCENE)

Qa11 Till deposited by Younger Ice—Characterized by a striking knob-and-kettle topography with many lakes and swamps. Consists of a heterogeneous assemblage of unsorted gravel, cobbles, and boulders in a light reddish-brown to tan silty to clayey matrix. Many large boulders are scattered across both the lateral and ground moraines, and the terminal moraine near Eleinschmidt Lake; locally parts of the moraine surface slope as much as 65 percent. Clasts range in shape from angular to well rounded, and are predominantly quartzites, argillites, and limestones of the Belt Supergroup. Soil drainage and permeability is variable, but generally is poor.

Qa12 Till deposited by Older Ice—Characterized by somewhat subdued knob-and-kettle topography. Consists of a heterogeneous assemblage of gravel, cobbles, and boulders in a reddish-brown, dense, granular, calcareous, clayey to silty, locally sandy matrix. Generally less than 15 percent of the till is gravel size or larger. Sparse boulders are scattered across both the lateral and ground moraines; surface slopes are commonly less than 25 percent. Clasts range in shape from angular to well rounded, and are predominantly quartzites, argillites, and limestones of the Belt Supergroup. Soil drainage and permeability generally is poor.

Qa13 OUTWASH OF HULL LAKE(?) ICE (PLEISTOCENE)—Brown and reddish brown; unconsolidated, mildly calcareous, well-bedded and moderately well sorted silt, sand, gravel, and cobbles. Forms even-surfaced, gently sloping, highly dissected remnants which are about 12 m (40 ft) above the outwash of the older ice. Soil drainage is moderate to good.

Qa14 TILL DEPOSITED BY HULL LAKE(?) ICE (PLEISTOCENE)—A generally formless, gently undulating deposit of low relief characterized by sparse kettle depressions, and a few boulders scattered across the surface. Consists of an unconsolidated, unsorted assemblage of sand, gravel, cobbles, and boulders in a brown to reddish-brown, compact, clayey matrix. Locally large glacial boulders, 3-4 m (10-13 ft) across, are scattered through the till; they are especially common near Ninemile Prairie (see Greenough quadrangle). Clay content is higher and this till consequently is more susceptible to mass movement than the younger tills. Where deeply dissected yields a discontinuous rubble veneer of cobbles and boulders.

PIEDMONT ALLUVIUM (QUATERNARY AND TERTIARY)—Reddish brown and tan; unconsolidated, moderately well sorted, stratified silty sand and silt, sand, and gravel which mantle broad, smooth-surfaced benches that flank the mountains in the Helmsville-Avon area. Deposits thin basinward, and typically are coarser-grained near the mountains. Overlies truncated pre-Tertiary and Tertiary sedimentary and volcanic rocks, and is segmented by modern stream valleys and gullies. Locally the upper part of the deposit is weakly cemented by calcium carbonate. Soundness and durability of the gravel is quite variable but typically is poor if the gravels contain high percentages of volcanic rock types. Those deposits rich in volcanic rock types are most common in basins which are adjacent to extensive outcrops of volcanic rocks.

BASIN DEPOSITS, UNDIVIDED (TERTIARY)—Brown, grayish brown, and yellowish brown; semiconsolidated to consolidated clays, shales, siltstones, and sandstones with minor amounts of conglomerate. Also included are white, gray, brown, yellowish brown, and red beds of altered volcanic ash, and ash-rich shale and siltstone. Competence is fair to good in the brown siltstones and sandstones but poor in the light-colored shales, clays, and ash beds. "Popcorn" surface on the shales, clays, and ash beds indicate a high-swell potential. Soil drainage is fair to poor. In many localities this unit underlies till, outwash, or alluvium and restricts and controls the downward and lateral migration of the local ground water.

VOLCANIC ROCKS, UNDIVIDED (TERTIARY)—Small outcrops of trachyte, andesite, lamprophyre, and extrusive volcanic breccia are in the western part of the area near Greenough (see Greenough quadrangle). Extensive basalt and andesite flows, volcanic breccias, and minor amounts of rhyolite, dacite, diorite, gabbro, latite, and welded tuff are in the eastern part of the area near Helmsville and Avon (see Helmsville and Avon quadrangles). The volcanic breccia are clay rich and very susceptible to mass movement.

QUARTZ MONZONITE INTRUSIVE ROCKS (TERTIARY AND CRETACEOUS)—Consists of the intrusive Garnet stock and other smaller igneous bodies. Generally shows well-developed blocky jointing; weathers to spheroidal boulders. Gold minerals are in quartz veins that fill fractures in the quartz monzonite in a zone within about half a mile of the surrounding marble and limestone country rock.

ACIDIC AND INTERMEDIATE INTRUSIVE ROCKS (TERTIARY AND CRETACEOUS)—Includes small intrusive igneous bodies ranging in composition from quartz monzonite to dacite.

CRETACEOUS AND JURASSIC BEDROCK, UNDIVIDED—Folded and faulted outcrops of (in descending order) the Cretaceous Blackleaf and Eastern Formations and the Jurassic Swift, Klerdon, and Sawtooth Formations.

PERMIAN TO MISSISSIPPIAN BEDROCK, UNDIVIDED—Folded and faulted outcrops of the Phosphoria and Quadant Formations, and the Madison Group. Locally some limestone and dolomite beds of the Jefferson Formation (Devonian) may be included with the Madison Group.

CAMBRIAN STRATA, UNDIVIDED—Folded, faulted and locally intruded and metamorphosed outcrops of the (in descending order) Harnack and Silver Hill Formations and Flathead Quartzite. Near the Garnet stock the limestones of the Harnack(?) and Silver Hill Formations have been metamorphosed to marble.

BELT SUPERGROUP ROCKS, UNDIVIDED (PRECAMBRIAN)—Consists of various units of the Belt Supergroup, chiefly the Mount Shields (argillite and sandstone), Shepard (argillite and dolomite), and Snowslip (argillite and sandstone) Formations in the northern part of the area; and the quartzite and claystone of Black Mountain, Bonner Quartzite and the Mount Shields, Shepard, and Snowslip (calcareous sandstone and quartzite) Formations in the southern part of the area.

Introduction

The eleven maps in the set (79-437 through 79-447) are part of a series prepared as a result of a cooperative venture between Missoula and Powell Counties, Montana and the U.S. Geological Survey. Weber, the senior author, at one time served as geologic consultant to the counties, and he is primarily responsible for the geologic mapping of the southern half of the Big Fork-Avon area. Witkind mapped the northern half of the area and his maps are available as Open-File Reports (listed below). Weber is now a geologist with the U.S. Forest Service in Great Falls, Montana; Witkind is a geologist with the U.S. Geological Survey in Denver, Colorado.

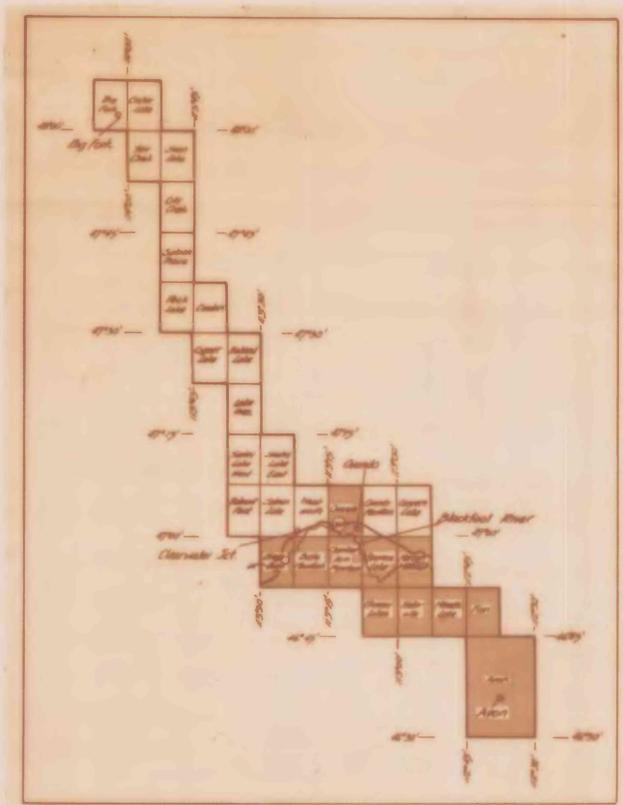
In addition to these maps, another map, by Witkind which discusses the seismicity of the Big Fork-Avon area, has been published as U.S. Geological Survey Miscellaneous Field Studies Map MF-923 and is titled "Major active faults and seismicity in and near the Big Fork-Avon area, northwestern Montana."

Surficial deposits

Most of the surficial deposits in the southern part of the Big Fork-Avon area were formed primarily during several advances and subsequent melts of large glaciers. The first advance recognized locally probably occurred about 150,000 years ago, during the Hull Lake Glaciation of the Pleistocene, when ice from the mountains to the north spread southward across the Blackfoot River. Deposits of this glacier are shown on the maps by the letters "hl" added to the symbol, thus "Qahl"—Till deposited by Hull Lake(?) ice. Subsequently, some 20,000 to 30,000 years ago during the Pinedale Glaciation of the Pleistocene, another glacier again moved southward into the area. It apparently reached almost as far south as Helmsville, before it withdrew and then readvanced again. This readvance, however, did not extend as far to the south, probably reaching only to Eleinschmidt Lake. Deposits of these two advances have also been differentiated on the map. The deposits of the first, or older, advance are shown on the map by the letter "o" added to the symbol, thus "Qao"—Till deposited by older ice. The deposits of the second, or younger advance, are shown on the map by the letter "y" added to the symbol, thus, "Qay"—Till deposited by younger ice.

In addition to the debris, mainly till, deposited by the ice, other material, chiefly sand and gravel, was carried far to the south by the meltwaters of the wasting glaciers. These deposits, termed outwash, are similarly identified on the map by the letters "bl" or "ol" or "yl" depending upon which glacier furnished them.

In some localities deposits of more than one ice advance may have been inadvertently grouped and thus mapped as the deposits of a single advance.



Index map showing quadrangles in the Big Fork-Avon area. Those quadrangles in Open-File Reports 79-437 to 79-447 are shaded.

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|----------------------------|--------|-------------------------------|--------|
| 1. Avon (S. 2/3) | 79-437 | 16. Lake Inez | 77-200 |
| 2. Beta Mountain (W. 1/3) | 79-438 | 17. Maroon Mountain (W. half) | 79-445 |
| 3. Big Fork | 78-174 | 18. Nevada Lake (S. half) | 79-446 |
| 4. Browns Lake | 79-439 | 19. Ovando | 77-196 |
| 5. Chamberlain Mountain | 79-440 | 20. Ovanda (S. 1/3) | 79-447 |
| 6. Chimney Lakes (NE 1/4) | 79-441 | 21. Ovando Mountain (S. half) | 77-465 |
| 7. Cilly Creek | 77-860 | 22. Peck Lake (E. half) | 77-539 |
| 8. Condon (W. half) | 77-866 | 23. Salmon Lake | 77-197 |
| 9. Coopers Lake (S. half) | 77-486 | 24. Salmon Prairie | 77-861 |
| 10. Crater Lake (W. half) | 78-173 | 25. Seeley Lake East | 77-201 |
| 11. Cygnet Lake | 77-198 | 26. Seeley Lake West | 77-201 |
| 12. Finn (SW 1/4) | 79-442 | 27. Swan Lake (SW 1/4) | 78-135 |
| 13. Greenough | 79-443 | 28. Woodward | 77-201 |
| 14. Helmsville (S. half) | 79-444 | 29. Yew Creek (NE 1/4) | 78-136 |
| 15. Holland Lake (W. half) | 77-199 | | |