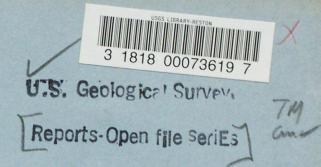
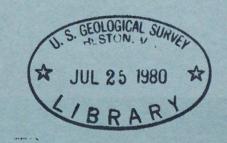
(200) R290 no. 80-447

UNITED STATES
DEPARTMENT OF THE INTERIOR
GEOLOGICAL SURVEY



FRESHWATER RESOURCES OF BIG PINE KEY, FLORIDA

OPEN-FILE REPORT 80-447



Prepared in cooperation with
MONROE COUNTY
and the
RIDGE AND LOWER GULF COAST WATER MANAGEMENT DISTRICT



CONVERSION FACTORS

The conversion factors and abbreviations for the terms used in this report are listed below:

Multip	ly inch-pound	unit	By	To obtain m	etric (SI) un	it
		1,2	2.54 0.3048 2.590 234.0	millimeter meter (m) square kilo cubic meter	meter (km ²)	
*	*	*	*	*	*	*
mean se	ea level (msl)				odetic Vertic 1929 (Datum o	-

UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY

FRESHWATER RESOURCES OF BIG PINE KEY, FLORIDA

By C. E. Hanson

Open-File Report 80-447

309904

Prepared in cooperation with

MONROE COUNTY
and the
RIDGE AND LOWER GULF COAST
WATER MANAGEMENT DISTRICT



Tallahassee, Florida

UNITED STATES DEPARTMENT OF THE INTERIOR

CECIL D. ANDRUS, Secretary

GEOLOGICAL SURVEY

H. William Menard, Director

ocha = 6535046

For additional information write to:

U.S. Geological Survey 325 John Knox Road Suite F-240 Tallahassee, Florida 32303

CONTENTS

			Page
Abstrac	ct		1
Introdu	icti	on	1
		se and scope	3
		wledgments	3
		ction	4
			4
			8
Company	onfi	guration of freshwater lenses	19
Potenta	i o l	ter quality freshwater supplies	
Summary	Iai		35
		references	
		ILLUSTRATIONS	
		ILLUSTRATIONS	
Figure	1.	Map showing location of Big Pine Key, Florida	2
	2.	Map showing location of monitor wells and ground- water and tidal-stage recording gages	5
	3.	Map showing depth below land surface of base of the Miami Oolite	9
	4.	Hydrographs of daily high water levels in wells 2 and 6, June to September 1976	13
	5.	Hydrograph of water levels at tidal-stage recorder on Pine Channel, October 1976	14
	6.	Profile of conductivity and chloride concentration in relation to geology at well 2A, March 1977	17
	7.	Profile of conductivity and chloride concentration in relation to geology at well 6A, March 1977	18
	8.	Graphs showing fluctuations of chloride concentrations 10 feet below the water table in selected wells in the northern part of Big Pine Key and rainfall, June 1976 to April 1977	20
	9.	Graphs showing fluctuations of chloride concentrations 10 feet below the water table in selected wells in the southern part of Big Pine Key and rainfall, June 1976 to April 1977	21

ILLUSTRATIONS (Continued)

			Page
Figure	10.	Map showing chloride concentrations 5 feet below the water table, September 14, 1976	22
	11.	Map showing chloride concentrations 10 feet below the water table, September 14, 1976	23
	12.	Map showing chloride concentrations 15 feet below the water table, September 14, 1976	30
	13.	Map showing chloride concentrations 5 feet below the water table, March 29, 1977	31
	14.	Map showing chloride concentrations 10 feet below the water table, March 29, 1977	32
	15.	Map showing chloride concentrations 15 feet below the water table, March 29, 1977	33
	16.	Map showing chloride concentrations at surface—water sites———————————————————————————————————	34
		TABLES	
Table	1.	Major inorganic ions in water from wells 2 and 6	6
	2.	Nutrient and total organic carbon analyses of water from wells 2 and 6	7
	3.	Depths to bottom of Miami Oolite	10
	4.	Seasonal rainfall distribution since 1969	11
	5.	Tidal levels, ground-water levels, and freshwater head data, May 1976 to April 1977	15
	6.	Chloride concentrations in wells, July 1976 to April 1977	24

FRESHWATER RESOURCES OF BIG PINE KEY, FLORIDA

By C. E. Hanson

ABSTRACT

The principal freshwater-bearing unit underlying Big Pine Key is a layer of oolitic limestone averaging 19 feet in thickness. The freshwater exists in two separate lenses, one in the northern half of the island and one in the southern half. The slightly larger north lens is separated from the south lens by a low-lying area 1 to 3 feet above NGVD of 1929. The lenses float on saltwater in the aquifer and are affected by tidal fluctuations. The areal and depth configuration of the lenses fluctuate in response to rainfall, evapotranspiration, lateral and vertical losses, and pumpage from local wells. The lenses are not a major source of freshwater. Only a small amount of the freshwater in the lenses can be removed before saltwater intrusion will occur.

INTRODUCTION

Big Pine Key, a reef limestone island in the lower Florida Keys (fig. 1), has an area of approximately 14 mi^2 . The northern part is primarily undeveloped pinelands and a buttonwood and mangrove fringe; it includes the Key Deer National Wildlife Refuge. The southern part has similar vegetation but is developed for residential and commercial use.

Big Pine Key lies within a tropical climatic zone. The average temperature is 29°C in summer and 22°C in winter. Temperatures seldom fall below 4°C due to the moderating effect of the Florida current and Florida Bay. Average annual rainfall is 39 inches, of which 75-85 percent falls from May through October (U.S. Weather Service, written commun., February 1977).

Development of Big Pine Key and other Florida Keys has accelerated since the middle 1960's. To protect the natural resources and minimize further environmental changes, the Florida Legislature declared the Florida Keys an "area of critical state concern."

The permanent population of Big Pine Key is about 800. During the winter the influx of tourists increases the population by about 2,000. This increase corresponds with the dry season when the island freshwater resources are lowest, and any increased pumpage stresses the available water supply.

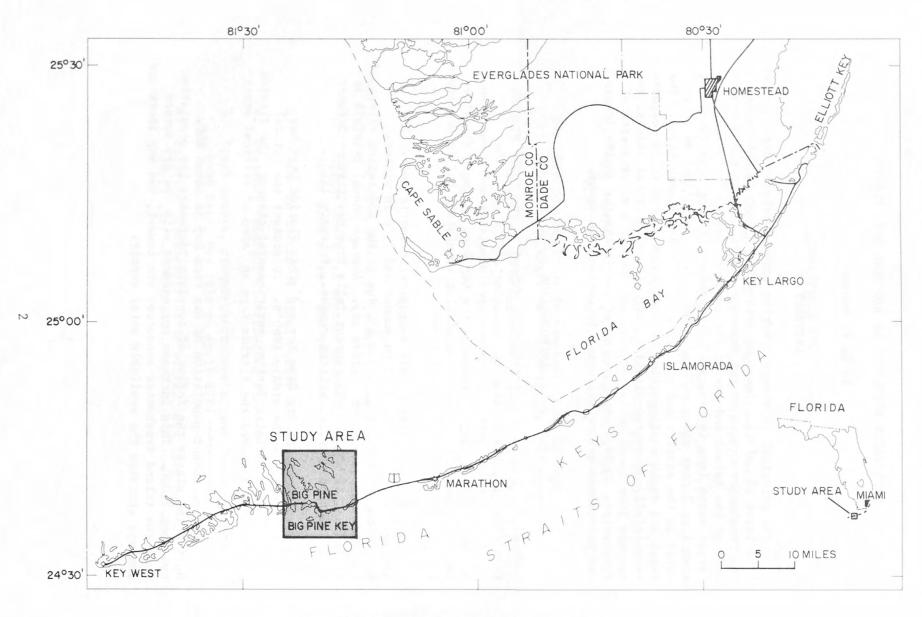


Figure 1.—Location of Big Pine Key, Florida

The main source of water for the island is the Florida Keys Aqueduct Authority pipeline from the mainland. Supplemental sources are shallow wells and rainfall cisterns. These sources have met the needs of residents and tourists, but concern for freshwater supplies has recently been generated because of increasing demands on the local ground-water resources by commercial and residential users. Proposed tropical plant nurseries would further increase freshwater demand.

Saltwater intrusion is the main threat to the fresh groundwater supply of Big Pine Key. Intrusion has occurred in some areas as a result of channelization for waterfront property, and in low parts of the island as a result of channelization for mosquito control.

In March 1976 the Ridge and Lower Gulf Coast Water Management District and Monroe County began a cooperative program with the U.S. Geological Survey to investigate the freshwater resources underlying Big Pine Key. This information will assist water managers in planning for the efficient use of water, as well as in short and long-range planning for the development of the Florida Keys freshwater resources.

Purpose and Scope

Increasing pressures on the water available from the Florida Keys Aqueduct Authority (FKAA) pipeline require water managers to look for new and alternate sources of freshwater. Knowledge of the availability of freshwater on Big Pine Key is needed for long-range planning and managing the freshwater resources of the island. This report includes information on the areal and vertical configuration of the freshwater lenses beneath the island, and describes how the lenses change seasonably.

Acknowledgments

The author wishes to thank Monroe County and the Ridge and Lower Gulf Coast Water Management District for support in this water-resources investigation. Acknowledgment is extended to the U.S. Fish and Wildlife Service, Key Deer National Wildlife Refuge; Florida State Division of Forestry; and the residents of Big Pine Key for valuable cooperation and assistance.

DATA COLLECTION

Between May 1976 and February 1977, 22 uncased wells were drilled to obtain geologic data and to monitor changes in chloride concentrations in the ground water. Ten wells were drilled along the north-south centerline of the island (fig. 2). Other wells were drilled perpendicular to the centerline in the northern and southern parts of the island. Four additional wells were drilled near the center of the island in an area of low elevation. Wells 1 through 20 range in depth from 20 to 26 feet (fig. 2). Wells 2A and 6A are 36.9 and 37.6 feet deep and were drilled to define the top of the saltwater zone. Chloride concentrations and specific conductance of ground water in the wells were determined by sampling at the water table, at successive 5-foot depth intervals, and at the bottom of each well. Specific conductance measurements were made in place with a conductance probe. Chloride concentrations were determined from specific conductance and by titration of the samples. Samples were collected and analyzed at monthly intervals from June 1976 through April 1977. Samples were also collected and analyzed for chloride concentrations at 14 surfacewater sites. Continuous water levels were obtained at two wells and one tidal gage.

Geological data were obtained from rock core samples which were taken during drilling of the monitor wells. Relative changes in porosity and degree of cementation or solution were noted at various depths.

Water samples were taken from wells 2 and 6 at a depth of 1 foot below the water table in October 1976 and April 1977. These samples were analyzed for major inorganic ions and macronutrients to determine differences during wet and dry seasons (tables 1 and 2).

Rainfall data from 1969 through June 1976 were obtained from records kept by the late Ralph Higgs at the meteorological station in the northern part of the island. Rainfall data were also supplied by the U.S. Weather Service in Key West, Florida.

GEOLOGY

The Florida Keys can be classified into two groups, distinguished by rock characteristics, orientation, shape, and origin. The Upper Keys (islands between Big Pine Key and the mainland) are oriented along a northeast-southwest arc. The Lower Keys (Big Pine Key to Key West) are almost at right angles to the arc and are separated by narrow channels that run northwest-southeast.

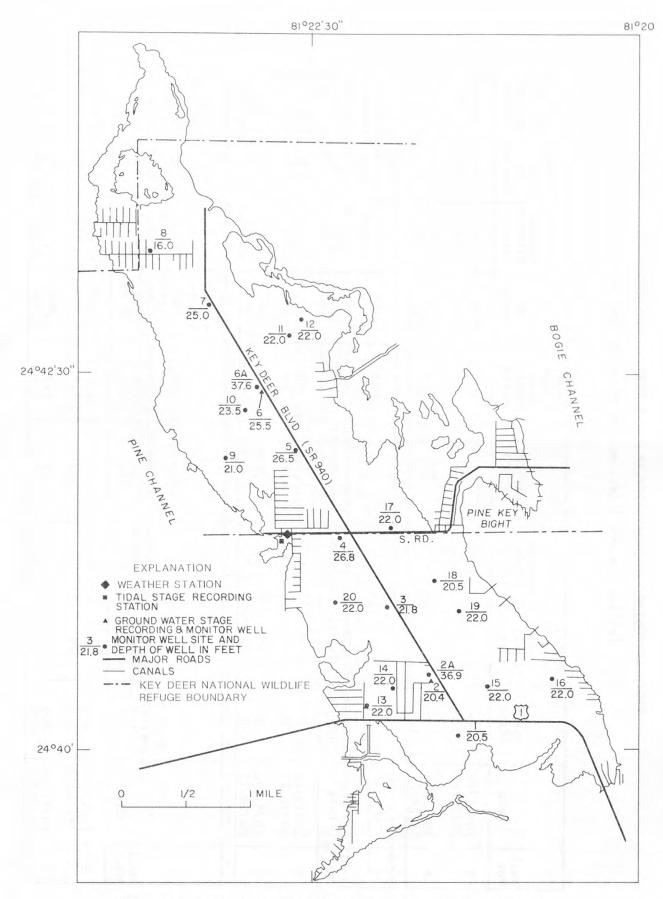


Figure 2.—Location of monitor wells and ground-water and tidal-stage recording gages

Table 1.—Major inorganic ions in water from wells 2 and 6 [milligrams per liter, unless otherwise indicated]

Well No.	Date	Depth (feet)	Total depth of well (feet)	Temper- ature (°C)	Turbidity (JTU)	Color (plat- inum cobalt units)	pH (units)	Carbon dioxide (CO ₂)	Alka- linity as (CaCO ₃)	Bicar- bonate (HCO ₃)
2	10/13/76 04/28/77	1.0	21 21	28.0 26.0	100 450	0	7.2 7.5	28 13	230 220	280 264
6	10/13/76 04/28/77	1.0	25 25	27.0 26.0	220 25	0	7.4 7.4	20 16	259 200	316 248

Well No.	Date	Carbonate (CO3)	Hardness (Ca,Mg)	Carbonate hardness	Specific conductance (micromhos)	Dis- solved chloride (C1)	Dis- solved sulfate (SO ₄)	Dis- solved calcium (Ca)	Dis- solved mag- nesium (Mg)	Dis- solved sodium (Na)
2	10/13/76 04/28/77	0 0	290 330	64 120	1,350 1,675	290 380	52 63	77 84	24 29	170 220
6	10/31/76 04/28/77	0 0	330 450	76 240	1,830 2,540	470 740	60 86	79 100	33 47	250 380

Well No.	Date	Sodium adsorp- tion ratio	Percent sodium	Dissolved potassium (K)	Dissolved fluoride (F)	Dissolved silica (SiO ₂)	Dis- solved stron- tium (micro- grams per liter)	Dis- solved solids (resi- due at 180°C)	Dis- solved solids (sum of con- stit- uents)	Dis- solved solids (tons per acre- feet)
2	10/13/76 04/28/77	4.3 5.3	55 58	10 13	.5	1.7 1.6	2,100 2,200	742 977	765 923	1.01
6	10/13/76 04/28/77	6.0 7.9	61 64	11 15	• 2	1.3 1.4	1,500 2,000	1,010 1,580	1,060 1,490	1.37 2.15

Table 2.—Nutrient and total organic carbon analyses of water from wells 2 and 6

[milligrams per liter]

Well No.	Date	Total nitrogen (N)	Total organic nitrogen (N)	Total kjeldal nitrogen (N)	Total ammonia nitrogen (N)	Total ni- trite (N)	Total ni- trate (N)	Total ni- trite plus ni- trate (N)	Total phos- phorus (P)	Total ortho- phos- phorus (P)	Total organic carbon (C)
2	10/13/76 04/28/77	.18	.08	•17 •44	.09 .10	.01	.00	.01	.01	.00	47 4
6	10/13/76 04/28/77	.19	.09	.18	.09	.01	.00	.01	.01	.00	10 4

The Upper Keys are of coral lime origin, known as the Key Largo Limestone, (Parker and others, 1955, p. 99-100), formed during the Pleistocene age as part of a coral reef. The islands are elongate parallel to the coastline along which the reef grew (Hoffmeister, 1974). The Key Largo Limestone is composed primarily of corals and some amorphous calcium carbonate, coral sand, breccia, and limey muds. The formation is very permeable, contains cavities, and has a porosity coefficient of about 0.2 (Parker and others, 1955).

The Lower Keys are an offshore extension of the Miami Oolite of Pleistocene age present in southeast Florida (Parker and others, 1955). Porosity of the Miami Oolite is moderately high, but the permeability is low. The porosity coefficient is approximately 0.4, but the interconnection of the voids between the oolids is very small, causing low permeability (oral commun., Robert Halley, 1977). The Miami Oolite was deposited in a warm, shallow sea as a shoal deposit (Parker and others, 1955).

The Miami Oolite extends from land surface to depths of 14 to 19 feet at the southern part of Big Pine Key and 16 to 23 feet at the northern part (table 3 and fig. 3); the thickness generally increases toward the centerline of the island where land surface altitude is highest. Key Largo Limestone underlies the Miami Oolite.

Three layers of hard, dense limestone were identified in test wells penetrating the Miami Oolite to the Key Largo Limestone. These layers are 2 to 8 inches thick and have low permeability. The first layer is near land surface, the second is 6 to 7 feet below land surface, and the third is 16 to 20 feet below land surface. The third layer is approximately at the Miami Oolite and Key Largo Limestone contact. These layers are probably discontinuous.

HYDROLOGY

The areal and depth configuration of the lenses of freshwater underlying Big Pine Key are influenced by five factors: (1) freshwater recharge (rainfall); (2) freshwater discharge; (3) response to tidal fluctuations; (4) proximity to saltwater bodies; and (5) permeability of subsurface materials.

Rainfall, which averages 39 inches per year, is the only source of freshwater recharge. During the study annual rainfall was 46.03 inches, with 86 percent occurring between May and October (table 4).

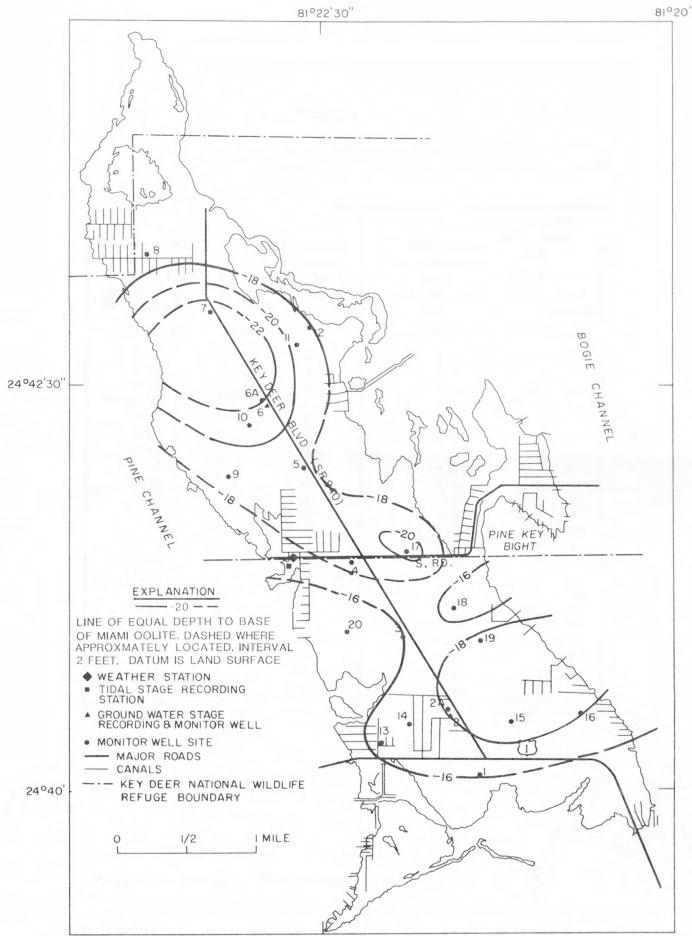


Figure 3. -- Depth below land surface of base of the Miami Oolite

Table 3.—Depths to bottom of Miami Oolite

Well Number	Depth from land surface to bottom of Miami Oolite (feet)	Approximate elevation of land surface (feet above NGVD of 1929)	Approximate elevation of bottom of Miami Oolite (feet below NGVD of 1929)
1	16.0	3.0	13.0
2	18.0	4.1	13.9
2A	19.7	4.2	15.5
3	16.0	1.75	14.25
4	18.5	4.5	14.0
5	18.5	5.0	13.5
6	22.0	5.3	16.7
6A	22.6	5.1	17.5
7	23.5	2.5	21.0
8	16.0	3.0	. 13.0
9	19.0	1.7	17.3
10	21.0	4.5	16.5
11	19.0	3.2	15.8
12	18.0	4.0	14.0
13	16.5	2.75	13.75
14	17.0	3.5	13.5
15	19.5	3.5	16.0
16	18.0	2.5	15.5
17	21.0	3.0	18.0
18	14.0	3.0	11.0
19	18.5	2.6	15.9
20	14.5	2.3	12.2

Table 4.—Seasonal rainfall distribution since 1969

Year	Total precipitation (inches)	Total precipitation for wet season (May-October) (inches)	Percent of total precipitation in wet season (%)
1969	61.63	49.12	79.7
1970	41.06	27.13	66.0
1971	41.11	33.45	81.3
1972	45.58	35.13	76.9
1973	55.98	43.02	76.8
1974		INCOMPLETE DATA	
1975	32.90	25.86	78.6
1976	47.01	39.68	84.4
Period of study May 1976 April 1977	46.03	39.68	86.2
Average since 1969*	46.48	36.19	77.8

^{*}Long-term average - 39.0 inches

Freshwater discharge is by surface runoff and losses from the lenses by evapotranspiration, lateral seepage, and pumpage. Average evapotranspiration in south Florida is about 70 percent of the rainfall (Parker and others, 1955). Much of the surface of Big Pine Key is riddled with solution cavities which allow rainfall to rapidly infiltrate, thereby, reducing evapotranspiration. In some parts of the island hard, dense limestone at or near the surface locally causes temporary ponding. This retards but does not prevent infiltration. Heavy vegetation in some areas near the lenses also increases evapotranspiration.

Freshwater losses by surface runoff have not been determined, but low relief on Big Pine Key suggests that this loss is small. Canals dug from the coast to inland parts of the island to obtain fill for housing construction and to provide waterfront properties tend to accelerate runoff of freshwater. The canals penetrate the water-bearing material, increasing natural discharge from the freshwater lenses and allowing saltwater to move farther inland. The largest canal networks are on the west side in the northern and central parts of the island.

Between May 1976 and April 1977, monthly average water levels in well 6 in the northern part of the island and well 2 in the southern part varied by up to 0.98 feet. The highest monthly average water level in well 6 was 1.44 feet above NGVD of 1929 in September 1976, and the lowest was 0.81 feet in February 1977 (table 5). In well 2 the highest monthly average water was 1.61 feet in October 1976, and the lowest was 0.63 feet in February 1977. Monthly average water levels at the tide recorder on Big Pine Channel were above NGVD of 1929 in all months except February when the average level was 0.10 feet below NGVD of 1929.

Water-table fluctuations on Big Pine Key are influenced primarily by rainfall and tides. During the wet season rainfall is probably the overriding influence causing the peaks in the hydrographs of wells 2 and 6 (fig. 4). During the dry season tides are the dominant influence. Fluctuations caused by evapotranspiration are subdued and are usually masked by those caused by tides and rainfall.

The tides in the Lower Keys result from the phasing together of the Atlantic Ocean and Gulf of Mexico tides. Tides from the Atlantic are a predictable two-cycle (diurnal) tidal movement, while the tides of the Gulf are normally one-cycle daily. The phasing together of these tide systems results in mixed tides. Varying amounts of this phasing cause large differences in the amplitudes and times of the tides in the Lower Keys. Tides at the recorder on Pine Channel on the west side of Big Pine Key (fig. 2) are mixed tides (fig. 5). Tide charts prepared by the National

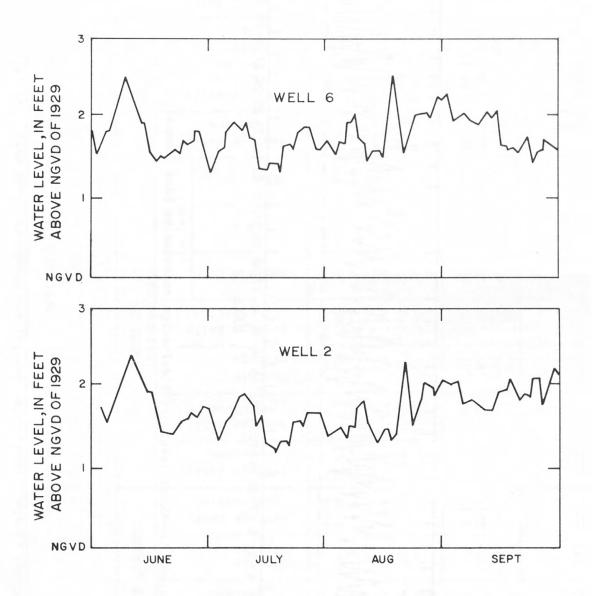


Figure 4.—Daily high water levels in wells 2 and 6, June to September 1976

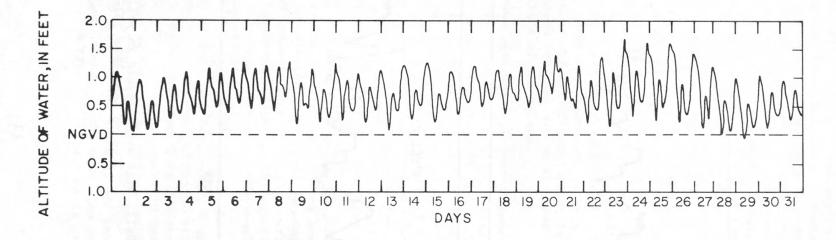


Figure 5.—Water levels at tidal-stage recorder on Pine Channel, October 1976

		BIO	G PINE	CHANNE	EL			W	ELL 2							WELL 6			
		y	7		Α.	7	y		8	Freshwater head in feet		8	>.		8. 1		water i	head	
Year	Month	Average dail maximum	Average dail minimum	Monthly average	Average dail fluctuation	Average daily maximum	Average daily minimm	Monthly average	Average daily fluctuation	Maximum	Minimum	Mean	Average daily maximum	Average daily minimum	Monthly average	Average dail fluctuation	Maximum	Minimum	Mean
	May	1.30	0.18	0.72	1.12	1.53	0.67	1.07	0.86	0.49	0.23	0.35	1.63	0.63	1.14	1.00	0.45	0.33	0.42
	June	1.25	.16	.67	1.09	1.70	.82	1.22	.88	.66	. 45	.55	1.80	.84	1.31	.96	.68	.55	. 64
	July	.98	10	.40	1.08	1.47	.69	1.03	.78	.79	.49	.63	1.62	.77	1.19	.85	.87	.64	.79
1976	Aug. Sept.	1.21	.27	.70	.94	2.02	1.10	1.49	.92	.83	.79	.81	1.85	1.08	1.44	.77	.81	.64	.86
1370	Oct.	1.24	.26	.73	.99	2.04	1.18	1.61	.86	.92	.80	.88	1.77	1.03	1.37	.73	.77	.53	.64
	Nov.	1.11	.08	.58	1.03	1.94	1.05	1.48	.90	.97	.83	.90	1.77	.93	1.31	.83	.85	.66	.73
	Dec.	1.08	.02	.54	1.06	1.81	.91	1.32	.90	.89	.73	.78	1.83	.92	1.34	.91	.90	.75	.80
	Jan.	.79	40	.16	1.19	1.35	.50	.87	.85	.90	.56	.71	1.76	.69	1.14	1.08	1.09	.97	.98
	Feb.	.50	49	10	1.09	1.07	.29	.63	.79	.88	.57	.73	1.17	.46	.81	.72	1.05	.67	.91
1977	March	.64	33	.12	.97	1.20	. 47	.81	.73	.80	.56	.69	1.38	.57	1.01	.81	.90	.74	.89
	April	.79	21	.28	1.00	1.35	.58	.95	.77	.79	.56	.67	1.36	.57	1.00	.79	.76	.57	.72
Stu	idy																		
ave	erage	.00	05	.44	1.05	1.59	.76	1.14	.84	.81	.60	.70	1.65	.79	1.20	.86	.84	.65	.76

Oceanographic and Atmospheric Administration for 1976 indicate that tide schedules in Bogie Channel on the eastern side of the island vary from those in Pine Channel by as much as several hours. Due to the shallow waters in the Lower Keys, local winds and weather systems offshore affect both the amplitude and times of tides in the area.

The extreme variability of the tides in the vicinity of Big Pine Key make it impossible to correlate water-table fluctuations with the tides or determine tidal efficiency at the monitor wells with data available.

The amount and distribution of freshwater in an island lens is determined by several factors which include rainfall, permeability of the sediments, and land-surface elevation. The depth at which saltwater underlies the freshwater is governed by the Ghyben-Herzberg relation (Chow, 1964) where conditions are homogeneous and under a steady state. These conditions do not exist in nature, but the relation is a useful tool for estimating the approximate depth of the fresh-saltwater interface. The relation states that for every foot of freshwater above NGVD of 1929, 40 feet of saltwater will be displaced below NGVD of 1929 due to the density difference between saltwater and freshwater. Freshwater has a density of 1.0 gram per cubic centimeter (g/cm³) and seawater has a density of 1.025 g/cm³; therefore, the freshwater is lighter and floats on the saltwater. Land-surface elevations on Big Pine Key are highest in the center of the island, and the freshwater head is greater there than at the coast. resulting body of freshwater is lens-shaped. The transition between freshwater and saltwater is a zone of dispersion maintained by fluctuations of the water levels and tidal action (Cooper and others, 1964). This zone grades from freshwater to saltwater. The freshwater lens expands during the wet season and contracts during the dry season. Parker (1955, p. 177) stated that Big Pine Key was one of the islands containing the largest amounts of freshwater in the Florida Keys. He also states pumpage of several thousand gallons per day (gal/d) would soon exhaust the island's freshwater even during the wet season.

On Big Pine Key the average freshwater head for March 1977 was 0.69 foot in well 2 and 0.89 foot in well 6 (table 5). According to the Ghyben-Herzberg relation, the freshwater lens theoretically should extend about 27 feet below NGVD of 1929 at well 2 and about 35 feet below NGVD of 1929 at well 6. Depth traverses of specific conductance profiles of water in well 2A (southern part of island) and well 6A (northern part of island) indicate that the relatively low salinity of the water is constant to a depth of about 20 feet below NGVD of 1929 (figs. 6 and 7). At that point the salinity increases rapidly with depth. The apparent zone of dispersion is about 10 feet thick. Samples taken 29 feet below NGVD of 1929 in well 6A had chloride concentrations of 6,600 mg/L.

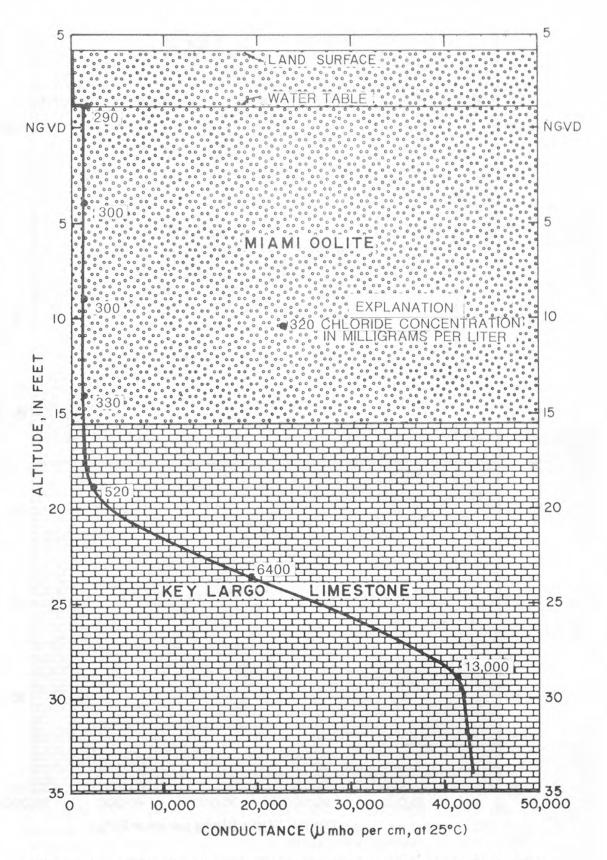


Figure 6.—Conductivity and chloride concentration in relation to geology at well 2A, March 1977

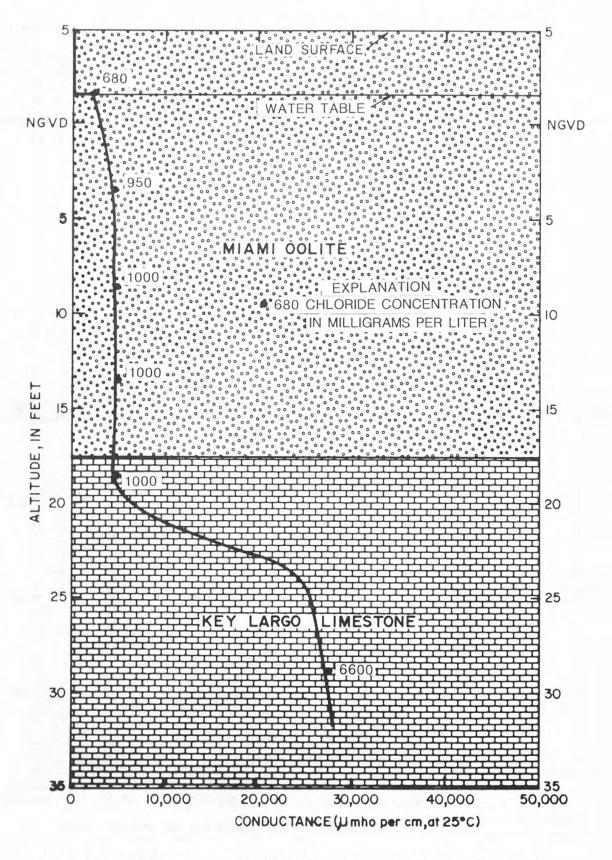


Figure 7.—Conductivity and chloride concentration in relation to geology at well 6A, March 1977

Configuration of Freshwater Lenses

During most of the year two separate freshwater lenses persist on Big Pine Key, one near well 6 in the north and one near well 2 in the south. These general areas are where land elevation above NGVD of 1929 is greatest. The lenses are separated by an area of low land-surface elevation, 1 to 3 feet above NGVD of 1929, extending from near Pine Key Bight southwest to Pine Channel. Monthly chloride determinations of ground water in the uncased monitor wells were made at the water table and at 5-foot intervals below the water table (table 6). Chloride concentrations increased rapidly with depth in some wells and more gradually in others.

Monthly changes in chloride concentrations between June 1976 and April 1977 at a depth of 10 feet below the water table in selected wells in the north and south freshwater lenses are shown in figures 8 and 9. The graphs indicate that chloride concentrations respond to rainfall on Big Pine Key. The greatest and most rapid response of chloride concentrations to rainfall variations occurs in well 7 near the northern end of the island and well 4 near the center. Both wells are located short distances from tidally affected canals. The canals provide channels for inland movement of seawater in addition to accelerating discharge of freshwater from storage in the lenses.

Chloride concentration fluctuations in wells other than 7 and 4 were more subdued, with concentrations decreasing as recharge from rainfall caused the lenses to expand and increasing during dry periods when evapotranspiration and pumpage reduced the lenses size. The somewhat larger chloride fluctuations in wells in the south lens probably are due to withdrawal of water for irrigation and domestic use. The increase in chloride concentration at well 6 in the north lens is probably due to pumping from a nearby rock quarry. The pumping temporarily lowers water levels and permits upward migration of saltwater from deeper zones. The data show that none of the monitor wells yields water whose chloride content meets drinking water standards (U.S. Environmental Protection Agency, 1976, p. 205) on a year-round basis.

Maps of equal chloride concentration at various depths below the water table were prepared to obtain synoptic configurations of the freshwater lenses during the investigation. The maps for September 14, 1976, when the freshwater resources were near maximum, show that water to a depth of 5 feet contains less than 500 mg/L chloride in a longitudinal strip extending nearly the entire length of the island (fig. 10). Water containing 250 mg/L chloride or less occurs in three areas, the largest in the north lens. At a depth of 10 feet below the water table (fig. 11), the

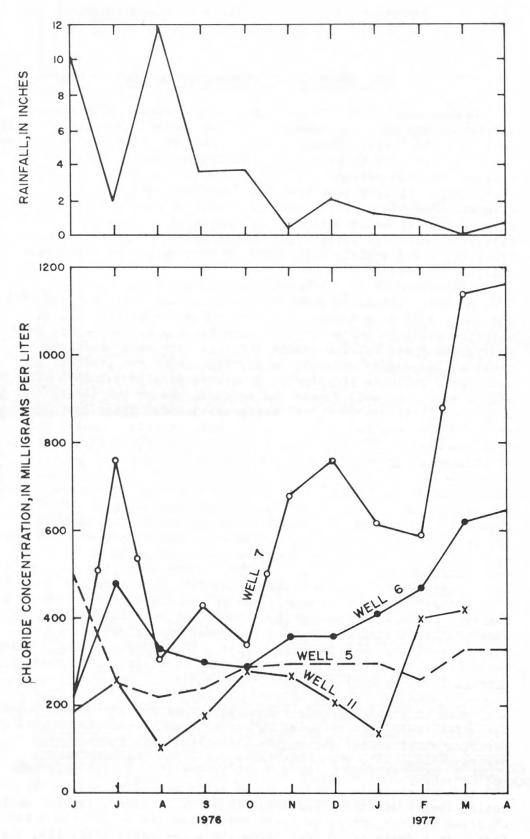


Figure 8.—Fluctuations of chloride concentrations 10 feet below the water table in selected wells in the northern part of Big Pine Key and rainfall, June 1976 to April 1977

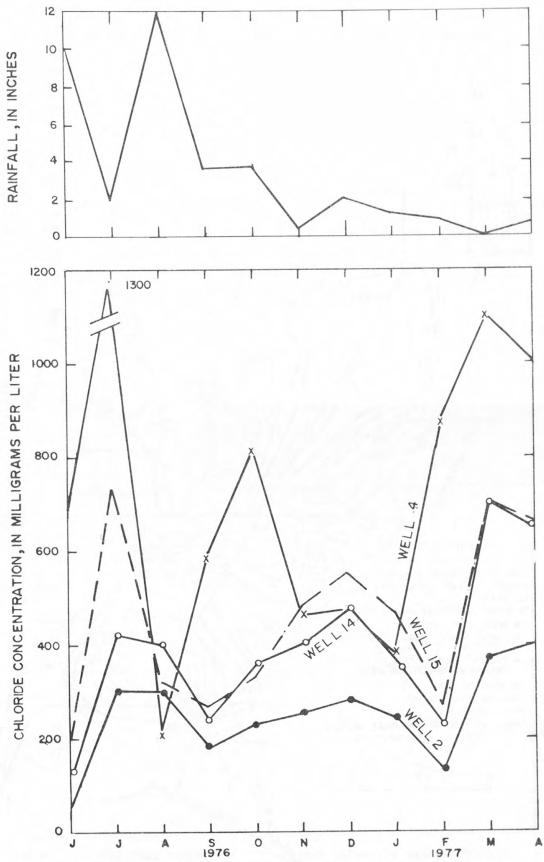


Figure 9.—Fluctuations of chloride concentrations 10 feet below the water table in selected wells in the southern part of Big Pine Key and rainfall, June 1976 to April 1977

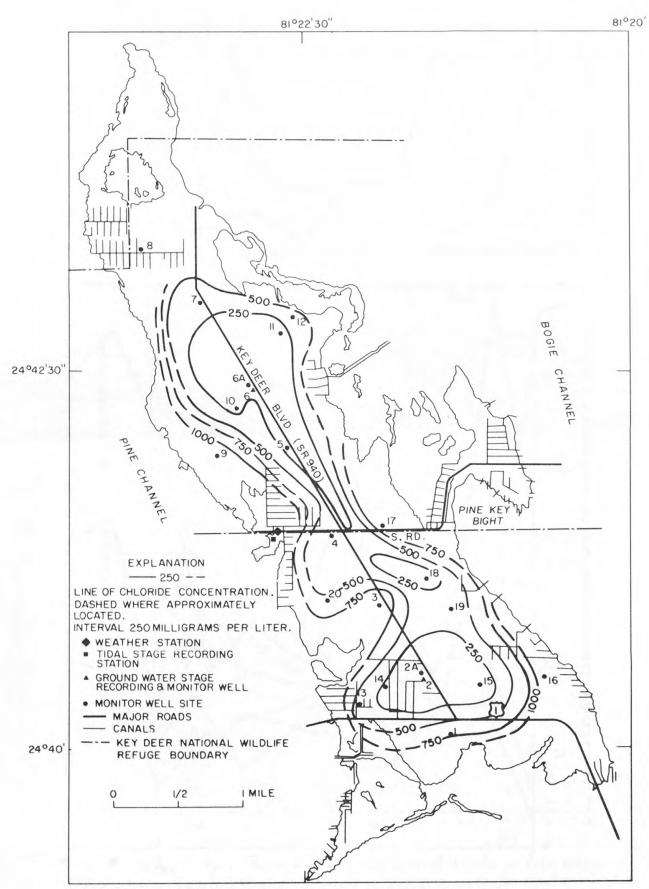


Figure 10.—Chloride concentrations 5 feet below the water table, September 14, 1976

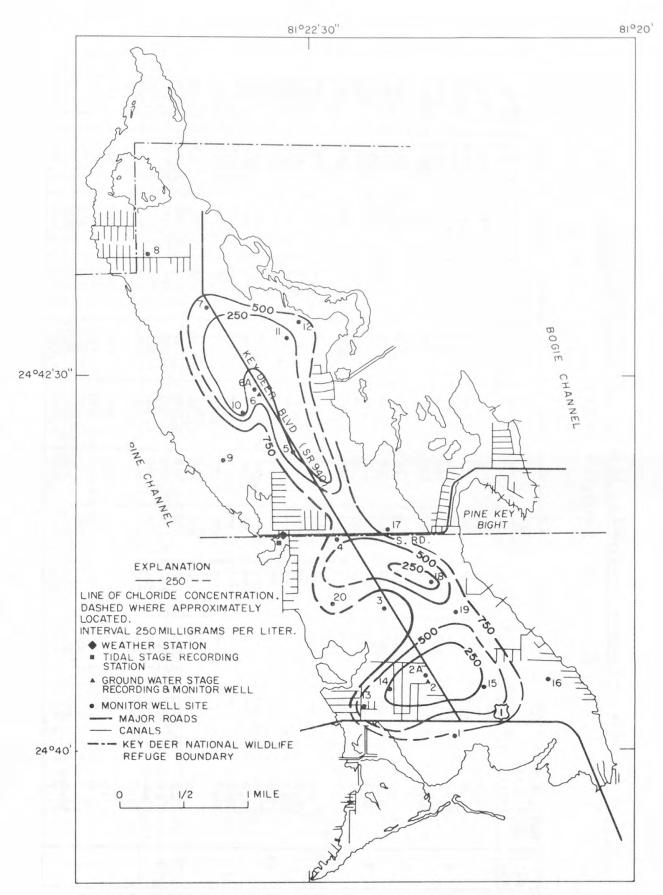


Figure 11.—Chloride concentrations 10 feet below the water table, September 14, 1976

Table 6.—Chloride concentrations in wells, July 1976 to April 1977
[milligrams per liter]

Well	Depth below water table				1976					19	77	
No.	(feet)	6/14	7/14-15	8/18	9/14	10/12	11/16	12/14	1/11	2/18	3/29	1 4/29
1	0.0 5.0 10.0 15.0 20.5	900 900 850 1,300	650 850 850 1,300	280 300 330 350 560	200 750 760 740 1,000	130 730 740 740 950	480 760 760 850 1,200	760 760 760 800 1,300	760 760 740 770 1,200	720 730 740 750 790	860 820 860 940 1,100	800 800 780 900 1,300
2	.0 5.0 10.0 15.0 20.4	38 37 42 42	200 300 300 300 300	180 270 300 — 300	48 160 180 190 210	65 180 230 240 240	250 240 250 260	270 270 280 300 300	200 210 240 260 270	130 120 130	350 370 370 380 380	360 390 400 390 400
2A	5.0 10.0 15.0 20.0 25.0 30.0 33.7		= = = = = = = = = = = = = = = = = = = =								290 300 300 330 520 6,400 13,000 13,000	280 280 320 360 650 7,200 15,000 16,000
3	.0 5.0 10.0 15.0 21.2	800 800 1,200 3,600	1,100 1,100 1,800 3,800	820 860 860 860 3,200	950 880 1,600 1,600 3,000	890 1,000 1,000 1,900 3,400	1,100 1,300 1,300 1,300 3,100	1,100 1,700 1,600 1,600 3,000	1,200 1,400 1,400 1,300 2,700	1,200 1,200 1,200 1,200 2,700	960 1,000 1,000 1,600 2,800	1,200 1,300 1,200 1,700 3,200
4	.0 5.0 10.0 15.0 24.7	192 650 650 2,300	1,000 1,300 1,300 3,300	170 160 210 290 1,700	120 310 580 900 2,200	240 320 810 850 2,500	440 440 460 560 2,500	450 450 470 590 2,400	350 370 380 480 2,300	380 390 870 900 2,400	560 1,000 1,100 1,200 2,500	470 700 1,000 1,300 2,800

Table 6.—Chloride concentrations in wells, July 1976 to April 1977 (Continued)

[milligrams per liter]

Well	Depth below water table				1976					19	977	
No.	(feet)	6/14	7/14-15	8/18	9/14	10/12	11/16	12/14	1/11	2/18	3/29	4/29
5	0.0 5.0 10.0 15.0 25.0	62 500 500 1,100	230 250 300 1,300	200 200 220 260 260	56 140 240 280 880	39 160 290 330 980	300 300 300 310 1,300	300 300 300 300 1,600	310 300 300 290 1,600	260 260 260 300 1,600	320 330 330 320 1,600	340 330 330 350 2,200
6	.0 5.0 10.0 15.0 25.0	190 230 270 750	310 480 500 800	280 300 330 350 560	180 250 300 360 490	240 260 290 360 500	360 350 360 360 560	370 360 360 370 470	410 390 410 410 470	480 450 470 450 590	610 590 620 600 780	610 650 650 640 820
6A	5.0 10.0 15.0 20.0 25.0 30.0 32.8	= = = = = = = = = = = = = = = = = = = =					=======================================				680 950 1,000 1,000 1,000 1,200 6,600 8,000	700 1,200 1,300 1,400 1,400 2,500 ———————————————————————————————————
7	.0 5.0 10.0 15.0 24.4	180 180 180 180	650 760 750 850	180 270 300 300 290	110 450 430 450 450	100 330 340 330 370	380 680 680 710 800	780 760 760 760 760	640 640 620 650 620	620 620 590 640 640	780 960 940 940 1,000	840 940 960 940 960
8	.0 5.0 9.8	3,000	11,000 11,000	7,200 7,500	5,800 9,500	8,200 8,800	15,000	19,000	20,000	18,000	13,000 14,000	13,000

Table 6.—Chloride concentrations in wells, July 1976 to April 1977 (Continued)

[milligrams per liter]

Well	Depth below water table			1976					1977				
No.	(feet)	6/14	7/14-15	8/18	9/14	10/12	11/16	12/14	1/11	2/18	3/29	4/29	
9	0.0 5.0 10.0 15.0 21.0	1,300 1,400 1,400 1,600	1,400 2,000 2,700 2,900	1,200 1,300 1,200 1,200 1,300	1,400 1,400 1,400 1,400 1,300	1,400 1,400 1,300 1,400 1,400	1,400 1,400 1,400 1,500 1,600	1,500 1,500 1,500 1,600 1,600	1,500 1,500 1,500 1,500 1,600	1,500 1,400 1,400 1,400 1,900	1,400 1,600 1,600 2,100 2,200	1,500 1,600 1,600 1,600 2,400	
10	5.0 10.0 15.0 23.4	180 170 170 170	270 490 780 800	140 160 160 200 200	96 210 240 260 270	200 220 220 240 270	310 300 300 400 410	290 290 280 390 480	260 240 260 260 320	230 230 230 230 230	370 370 370 610 780	350 340 340 450 770	
11	.0 5.0 10.0 15.0 21.0	58 60 60 65	180 190 190 190	220 260 260 260 270	51 100 110 110 120	160 160 180 190 200	280 280 280 280 310	270 260 270 270 280	220 210 210 230 220	140 140 140 140 140	340 390 400 400 420	380 410 420 430 430	
12	.0 5.0 10.0 15.0 22.0	280 310 310 310	600 650 750 600	310 270 330 430 430	280 360 420 430 450	300 300 450 480 470	410 400 480 560 560	470 450 490 630 660	450 440 480 530 560	360 360 390 460 480	600 910 960 990 1,100	450 490 730 880 960	
13	.0 5.0 10.0 15.0 21.1	400 400 400 1,300	320 350 350 350 1,400	340 400 390 440 1,500	130 440 470 1,200 1,600	440 500 630 860 1,900	460 460 460 1,500 1,900	490 460 480 930 1,700	470 450 470 1,400 2,100	530 490 520 610 1,200	520 970 1,100 1,600 1,700	540 1,400 1,500 1,600 2,400	

Table 6.—Chloride concentrations in wells, July 1976 to April 1977 (Continued)

[milligrams per liter]

Well	Depth below water table	1976								1977			
No.	(feet)	6/14	7/14-15	8/18	9/14	10/12	11/16	12/14	1/11	2/18	3/29	4/29	
14	0.0 5.0 10.0 15.0 21.1	94 130 160 770	350 420 500 950	400 400 400 460 860	140 180 240 450 720	110 300 360 430 690	400 400 400 500 750	470 460 470 590 820	360 360 370 600 800	220 220 230 680 740	650 680 700 830 830	450 500 650 770 1,000	
15	5.0 10.0 15.0 21.2	130 170 170 170	650 730 730 700	250 330 320 440 420	87 190 270 340 390	100 230 330 400 420	470 460 480 490 500	550 540 550 570 580	430 460 470 490 550	240 240 270 290 290	700 700 700 700 700	650 680 660 670 690	
16	5.0 10.0 15.0 21.2	250 250 250 250 240	2,200 2,200 2,200 2,200 3,300	2,400 3,100 3,100 3,200 3,200	530 1,900 1,900 1,800 2,200	350 2,800 2,800 2,800 2,900	1,100 2,900 2,900 2,900 3,400	3,100 3,200 3,200 3,200 4,000	2,600 3,000 3,100 3,100 4,000	3,000 3,100 3,200 3,200 4,000	2,900 2,900 3,000 3,000 3,200	3,000 3,600 3,700 3,400 4,500	
17	.0 5.0 10.0 15.0 20.5	1,200 1,300 1,400 1,500	1,300 1,600 1,700 2,800	1,100 2,100 2,100 2,100 2,100 3,100	500 900 950 950 950	670 1,000 1,000 1,000 2,000	1,200 2,200 2,200 2,200 2,900	1,800 2,600 2,600 3,200 4,000	2,200 2,300 2,200 2,500 3,400	1,800 1,700 1,800 1,800 3,000	1,600 2,800 2,800 3,000 3,600	2,200 2,100 2,600 3,100 3,800	
18	.0 5.0 10.0 15.0 20.3	310 300 290 800	260 260 280 800	95 120 100 120 530	55 190 210 600 920	100 230 280 440 730	210 200 220 560 960	240 230 350 660 1,100	200 210 340 590 1,000	210 210 230 320 660	450 490 510 690 920	350 370 390 440	

Table 6.—Chloride concentrations in wells, July 1976 to April 1977 (Continued)

[milligrams per liter]

Well	Depth below water table (feet)				1977							
No.		6/14	7/14-15	8/18	9/14	10/12	11/16	12/14	1/11	2/18	3/29	4/29
19	0.0 5.0 10.0 15.0 21.4	300 330 340 350	600 750 750 800	220 320 360 420 420	150 440 560 660 690	160 440 520 570 630	520 520 530 580 640	570 580 600 690 810	500 490 540 610 700	440 400 440 420 540	740 870 880 870 1,600	640 760 800 820 960
20	.0 5.0 10.0 15.0 20.3	650 850 1,100 4,500	900 1,100 1,200 4,600	190 290 290 1,400 4,200	220 670 600 1,600 2,800	500 700 700 1,400 3,400	640 610 630 2,500 3,300	670 730 770 2,500 4,000	640 590 600 2,400 3,800	440 460 450 2,100 3,700	850 850 850 2,000 3,700	750 1,100 1,300 2,300 4,400

areas containing water with 250 and 500 mg/L chloride have decreased considerably. At a depth of 15 feet (fig. 12) these areas continue to decrease. Figure 12 shows that at 15 feet below the water table, the two freshwater lenses are separated by an area in which the water contains more than 750 mg/L chloride.

Maps of equal chloride concentrations at depths of 5, 10, and 15 feet below the water table were prepared for samples collected on March 29, 1977, when freshwater resources were near minimum (figs. 13, 14, and 15). The maps show that none of the water on Big Pine Key had chloride concentrations less than or equal to the drinking water criteria of 250 mg/L (U.S. Environmental Protection Agency, 1976, p. 205). Water from several wells in both lenses was found to have chloride concentrations less than 500 mg/L (table 6).

SURFACE-WATER QUALITY

Water samples were taken from 15 surface-water sites in April 1976 and March 1977 and analyzed for chloride concentrations (fig. 16). These sites represent areas of freshwater available to wild-life, primarily deer. Chloride concentrations ranged from 102 to 5,600 mg/L and exceeded 250 mg/L in about 80 percent of the samples.

POTENTIAL FRESHWATER SUPPLIES

The amount of freshwater remaining in the shallow lenses near the end of the dry season (fig. 16) does not represent the potential supply or the perennial yield available from the lenses. The quantity of freshwater that remains in the lenses at the end of the dry season will fluctuate depending upon the amount and timing of rainfall. During years of above normal rainfall the lenses will contain more freshwater and cover a larger area at the end of the dry season than shown in figure 16. During extended droughts the freshwater lenses may totally dissipate. Also any substantial increase in pumping would probably reduce water levels and increase saltwater intrusion.

Water is presently being withdrawn from the lenses from about 150 private and commercial wells. In most cases these wells are a secondary source of water; the Florida Keys Aqueduct Authority pipeline is the primary source. The principal use of water from the lenses is irrigation of lawns and plant nurseries. Due to the sporadic nature of pumpage of these wells, no estimates of withdrawals from the lenses were made.

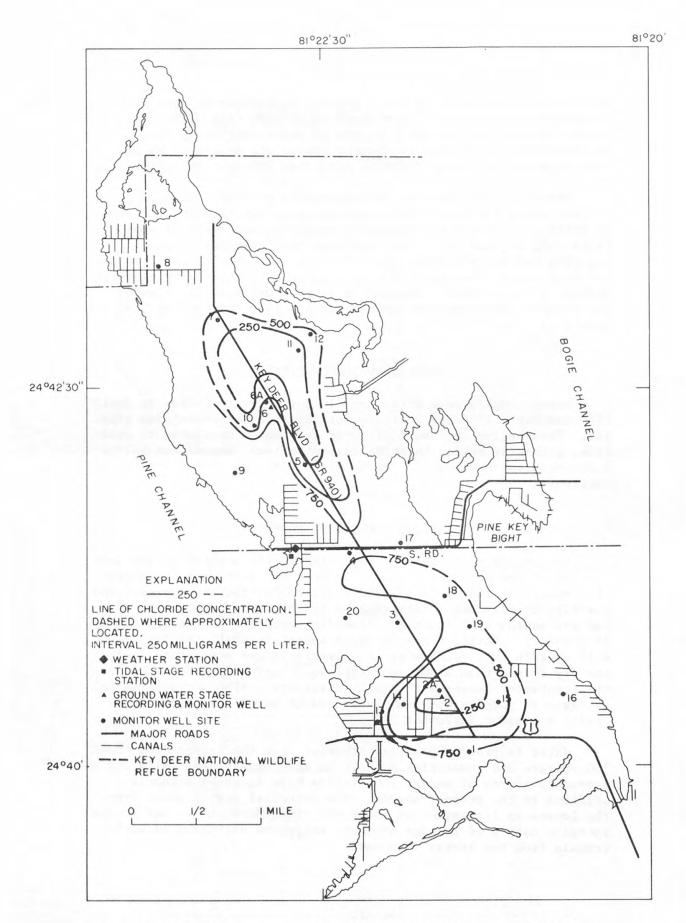


Figure 12.—Chloride concentrations 15 feet below the water table, September 14, 1976

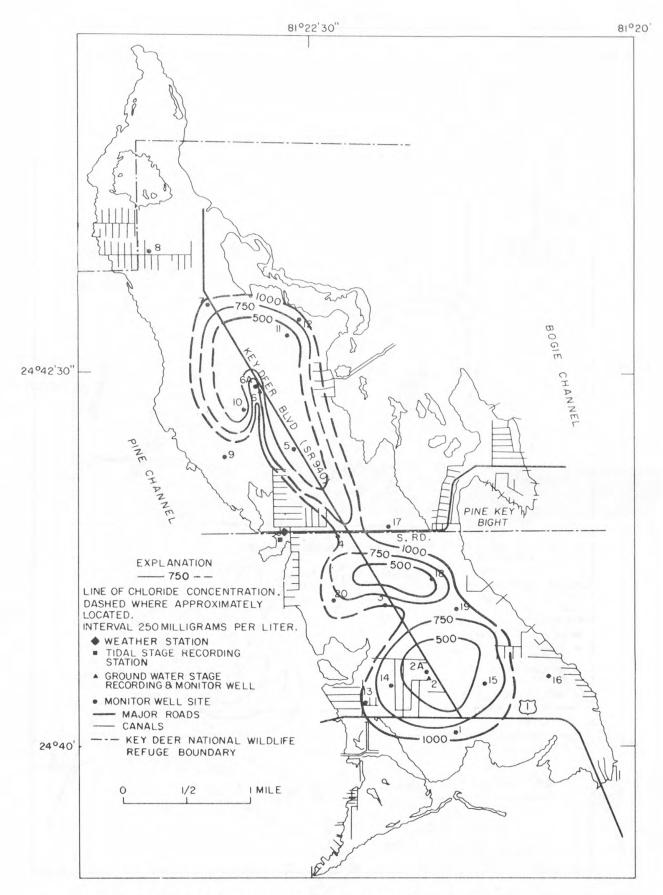


Figure 13.—Chloride concentrations 5 feet below the water table, March 29, 1977

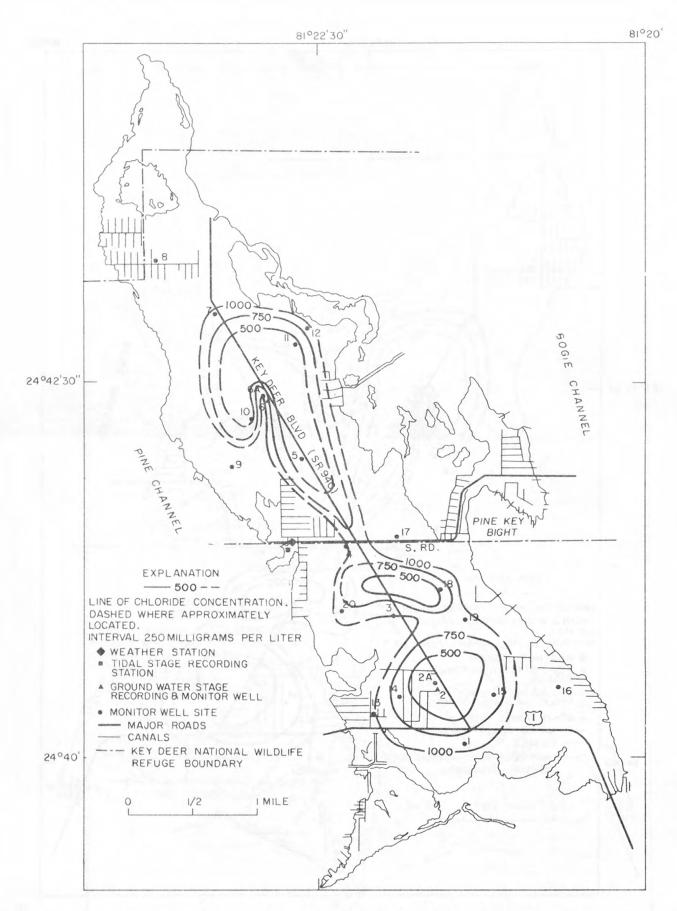


Figure 14.—Chloride concentrations 10 feet below the water table, March 29, 1977

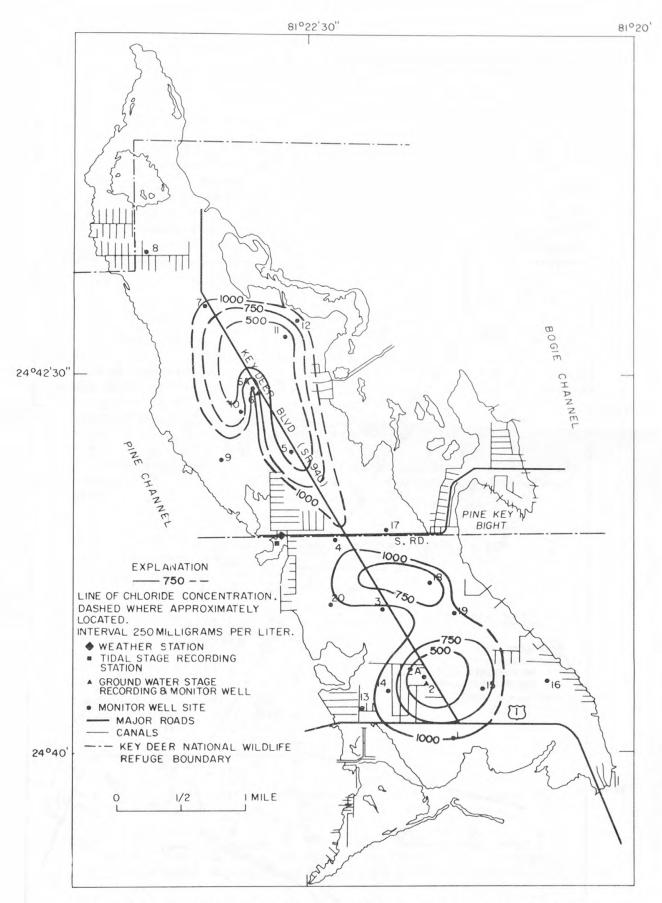


Figure 15.—Chloride concentrations 15 feet below the water table, March 29, 1977

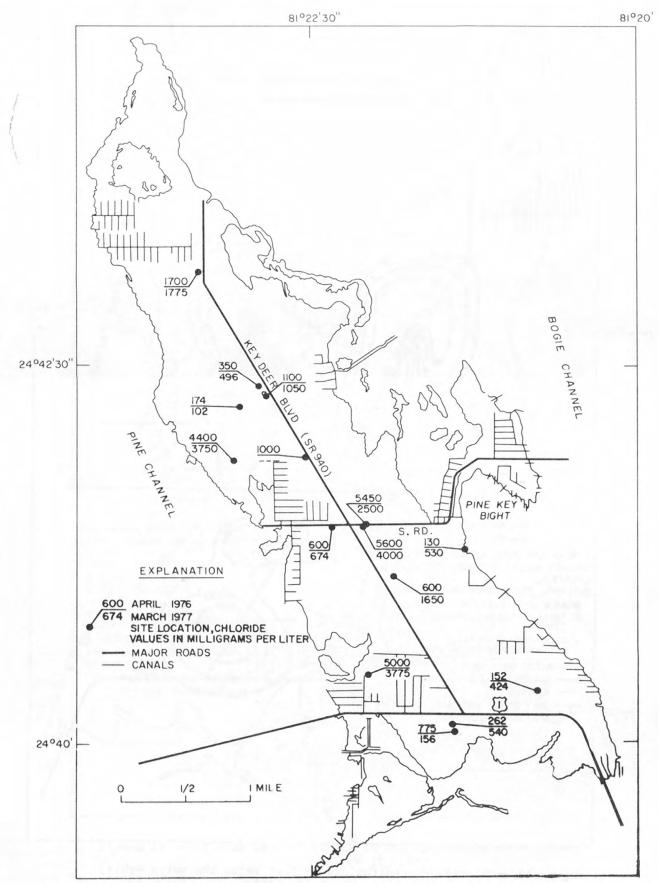


Figure 16. -- Chloride concentrations at surface-water sites

Much of the south end of Big Pine Key has been zoned for residential use. This use, along with new or enlarged plant nurseries, will increase the stress on the freshwater lenses which can only supply moderate amounts without detrimental effects during most years. Most of the area north of the east—west State Road is in the Key Deer National Wildlife Refuge and is not subject to private development. Large increases in water pumpage are not anticipated there.

SUMMARY

Miami Oolite, the primary freshwater bearing unit underlying Big Pine Key, averages 19 feet thick. Most of the freshwater is in two separate lenses, one in the northern half of the island and the other in the southern half. The slightly larger north lens is separated from the south lens by a low-lying area 1 to 3 feet above NGVD of 1929. Ground-water levels fluctuate in response to tides.

The areal and depth configuration of the lenses fluctuate in response to recharge by rainfall and to discharge by evapotranspiration, pumpage, and ground-water discharge to the sea.

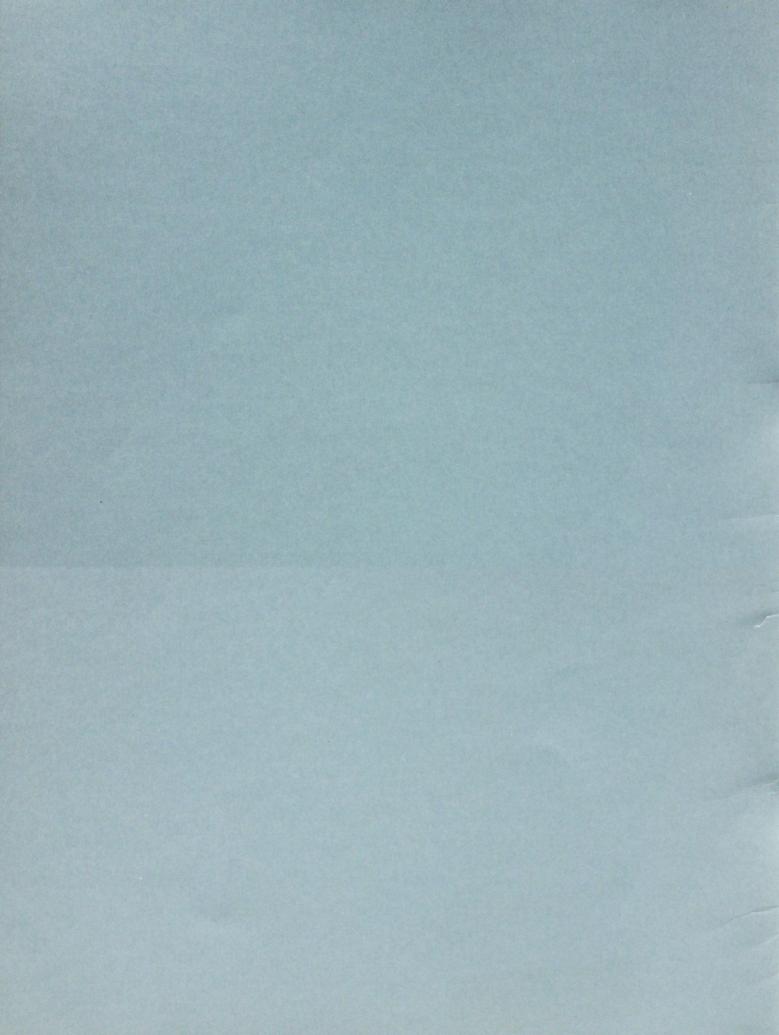
The freshwater lenses under Big Pine Key do not constitute a major source of freshwater. Continued pumping of freshwater from shallow wells will probably not damage the system. However, a continued point withdrawal or large scale general withdrawal would induce saltwater intrusion into freshwater lenses.

It is doubtful that the north lens will be subjected to excessive withdrawals, because most of this area is part of the Key Deer National Wildlife Refuge and is not subject to development. However, the area which overlies the south freshwater lens is platted for residential and commercial development. Future increased withdrawals will probably be from this area and will increase the threat of saltwater intrusion.

SELECTED REFERENCES

- Chow, V. T., 1964, Handbook of applied hydrology: McGraw-Hill Book Company, New York, N.Y., chap. 13.
- Cooper, H. H., Jr., Kohout, F. A., Henry, H. H., and Glover, R. E., 1964, Seawater in coastal aquifers: U.S. Geological Survey Water Supply Paper 1613-C, 84 p.

- Hem, J. D., 1970, Study and interpretation of the chemical characteristics of natural water: U.S. Geological Survey Water Supply Paper 1473, 363 p.
- Hoffmeister, J. E., 1974, Land from the sea: University of Miami Press, 143 p.
- National Academy of Sciences and National Academy of Engineering, 1973, Water quality criteria, 1972: U.S. Environmental Protection Agency Report EPA R3 73 033, 594 p.
- Parker, G. G., Ferguson, G. E., Love, S. K., and others, 1955, Water resources of southeastern Florida, with special reference to the geology and ground water of the Miami area: U.S. Geological Survey Water Supply Paper 1255, 965 p.
- U.S. Environmental Protection Agency, 1976, Quality criteria for water: U.S. Environmental Protection Agency, 256 p.





UNITED STATES DEPARTMENT OF THE INTERIOR GEOLOGICAL SURVEY 325 John Knox Rd--Suite F240

POSTAGE AND FEES PAID U.S. DEPARTMENT OF THE INTERIOR INT. 413

FIRST CLASS

