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CHEMISTRY OF ZEOLITES AND ZEOLITIC TUFFS FROM A PLIOCENE LACUSTRINE DEPOSIT NEAR DURKEE, BAKER COUNTY, OREGON

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### INTRODUCTION

Zeolites in an unnamed Pliocene lacustrine and fluviatile deposit near Durkee were first mentioned by Eakle (1898). He described a new zeolite, erionite, from a small opal quarry in a "rhyolite" (welded ash-flow tuff) about 6 km southeast of Durkee. Eakle, unfortunately, provided only a vague description of the type locality of erionite; and more than half a century passed before the type locality was rediscovered by L. W. Staples, and additional specimens were compared with Eakle's original material (Staples and Gard, 1959). Prior to the late 1950's, the Durkee locality was the only confirmed occurrence of erionite in the world.

Our investigation of the zeolites and the diagenesis of tuffaceous rocks near Durkee began late in 1970 after a brief reconnaissance early in that year showed an abundance and variety of authigenic zeolites in silicic tuffs of the lacustrine deposit. Field studies continued intermittently until 1972, and laboratory studies continued intermittently until 1976. Preliminary findings have been published in two brief reports (Sheppard and Gude, 1975; Gude and Sheppard, 1978). This present report is concerned only with the chemistry of certain zeolitic tuffs from the Durkee deposit and with the chemistry of chabazite, clinoptilolite, and erionite that were separated from six of the tuffs.

### GEOLOGIC SETTING

The deposit that contains the zeolitic tuffs was mapped as an unnamed unit of "tuffaceous lake and stream sediments" by Prostka (1967). This Pliocene deposit of lacustrine and fluviatile rocks and a welded ash-flow tuff occupies an irregularly shaped, northwesterly elongated basin. The deposit covers an area of about 118 km<sup>2</sup>, but the zeolitic part of the deposit seems restricted to an area of about 27 km<sup>2</sup> in the southeastern part of the basin (Gude and Sheppard, 1978). The total exposed thickness of the deposit is about 350 m. Tuffs are interbedded with mudstone, siltstone, sandstone, and diatomite. Zeolitic tuffs are chiefly in the lower 135 m of the deposit where diatomite is absent. Those tuffs interbedded with diatomite consist chiefly of fresh glass and occur stratigraphically higher in the section.

SEMIQUANTITATIVE SPECTROGRAPHIC ANALYSES OF ZEOLITE-RICH TUFFS Semiquantitative spectrographic analyses of 11 zeolitic tuffs (localities, fig. 1) rich in chabazite, clinoptilolite, and erionite are given in table 1. These analyses were performed on bulk samples that contained at least 90 percent of the stated zeolite. The chief impurities were trace amounts of other zeolites and smectite. Erionite-rich tuffs from the Durkee deposit have a consistently high potassium content and a highly variable calcium content relative to the other zeolitic tuffs. Clinoptilolite-rich tuffs seem to have relatively high contents of barium and strontium, although one sample (no. 8) of an erionite-rich tuffs seem to be especially low in strontium relative to the clinoptilolite-rich and erionite-rich tuffs.

### CHEMICAL ANALYSIS AND UNIT-CELL COMPOSITION OF CHABAZITE, CLINOPTILOLITE, AND ERIONITE

Relatively pure zeolite separates were prepared for chemical analysis from nearly monomineralic tuffs. The zeolites were separated by crushing the tuff and then disaggregating it in an ultrasonic bath. The zeolites were then concentrated by repeated centrifuging in a heavy-liquid mixture of bromoform and acetone, utilizing the technique described by Schoen and Lee (1964).

Conventional rock analyses were performed on zeolite separates of chabazite, clinoptilolite, and erionite, and the analyses are given in table 2. The analyses of those separates that contained minor calcite impurities were corrected by subtracting the analyzed  $CO_2$  content plus the equivalent CaO content to make calcite and then recalculating the analyses so that they total 100 percent. These corrected analyses are also given in table 2.

The analyses were calculated into atoms per unit cell on the basis of 72 oxygen atoms, and the composition of the unit cell is given in table 2 and plotted on figures 2 and 3. Figure 2 shows that the plots for chabazite and erionite overlap. The Durkee clinoptilolites are more siliceous than the chabazites or the erionites. The Si:  $Al+Fe^{+3}$  ratio of one clinoptilolite (no. 3) is anomalously low, inasmuch as most clinoptilolites have a Si:  $Al+Fe^{+3}$ 

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Figure 1.--Part of the U.S. Geological Survey Durkee, Oregon 15-minute quadrangle topographic map showing sample localities indicated by dots. Analyses of the samples are given in tables 1 and 2.

#### Table 1.--Semiquantitative spectrographic analyses of zeolite-rich tuffs

[Analyst: Harriet G. Neiman, U.S. Geological Survey. Results are to be identified with geometric brackets whose boundaries are 1.2, 0.83, 0.56, 0.38, 0.26, 0.18, 0.12, and so forth, but are reported arbitararily as midpoints of these brackets, 1, 0.7, 0.5, 0.3, 0.2, 0.15, 0.1, and so forth. The precision of a reported value is approximately plus or minus one bracket at 68 percent confidence, or two brackets at 95 percent confidence. G. greater than 10 percent; N, not detected at limit of detection; L, detected but below limit of determination]

		Chabazite-rich tuff			Clino	ptilolite-ri	Erionite-rich tuff				
	1	2	3	4	5	6	7	8	9	10	11
					Weight perc	ent					
S1	- G	G	G	G	G	G	G	G	G	G	G
A1	- 7	7	5	5	7	7	10	7	7	7	5
Fe	5	.7	.7	.7	2	3	3	2	3	3	. 5
Mg	2	.1	.7	.7	.5	.7	.5	.3	.5	.7	1.5
Ca	- 2	5	1.5	2	3	1.5	2	5	2	2	.0
Na	- 1.5	1	.5	.7	.7	2	3	.7	.3	1	1.5
K	7	2	2	2	2	3	3	5	3	3	5
Ti	07	.05	.1	.1	.07	.15	.2	.05	.03	.2	.0
					Parts per mil	lion					
8	- N	L	L	L	N	L	L	L	L	N	N
Ba	- 200	300	200	200	1,500	1,000	1,000	2,000	150	150	70
Be	- 2	2	N	N	2	N	N	5	2	3	2
Ce	- N	N	N	N	N	N	N	N	N	150	L
Cr	- L	1	L	L	1	2	1	1	L	L	L
Cu	- L	10	10	7	3	30	15	15	30	20	3
Ga	- 20	15	10	15	20	10	15	30	30	30	30
La	- 50	30	30	30	30	N	N	50	N	100	50
Mn	- 30	100	50	30	500	70	100	70	30	150	50
Mo	- N	N	N	N	N	50	30	7	15	3	N
Nb	- 10	30	L	L	10	N	L	10	10	50	15
Nd	- N	N	N	N	N	N	N	N	N	100	N
Pb	- 10	20	15	10	20	20	10	20	20	30	20
Sc	- L	N	L	N	L	5	7	N	N	L	N
Sr	- 30	100	70	70	200	1,000	1,000	700	150	150	70
Y	- N	15	5	- 7	20	15	10	20	30	70	N
Y	- 30	30	15	15	70	N	N	10	10	70	10
Yb	- 5	3	1.5	1.5	7	N	N	L	1	7	L
Zr	- 200	100	150	100	200	200	100	150	200	300	150
nalysis	Lab. No.	Local	ity No.	Thickness of tuff		Analysis	Lab No.	Locality No.	Thi	ckness of	tuff
		(fig	. 1)	(in meters)				(fig. 1)		(in meters	)
1	D152183		4	0.91		7	D152188	8		0.05	
2	D152184		4	0.55		8	0152179	21		0.04	
3	0152185		4	1.68		9	D152180	28		0.10	
4	D152186		4	1.52		10	D152182	29		0.46	
5	D152181	2	9	1.22		11	D152191	8		0.20	
6	D152187		8	0.10							

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### Table 2.--Chemical analyses and composition of unit cell of chabazite, clinoptilolite, and erionite

[a, uncorrected analysis;	b, analysis	corrected fo	or CU2 plus	equivalent	CaO to	make calcite]
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Cha	bazite		Clinoptilol	ite		Erior	nite		
1	2		3	4		5		1	
۵	a	a	b	a	a	b	å	b	

Chemical analyses, in weight percent

[Analysts: Edythe E. Engleman and Vertie C. Smith, U.S. Geological Survey]

\$10 <sub>2</sub> 57.91	58.65	61.59	62.87	64.72	59.40	59.81	59.53	60.67
A1203 14.25	13.51	12.57	12.83	12.10	13.51	13.61	13.59	13.85
Fe <sub>2</sub> 0 <sub>3</sub> 42	.61	2.28	2.33	1.13	1.49	1.50	.22	.22
Fe002	.00	.03	.03	.06	.05	.05	.07	.07
Mg040	1.27	1.13	1.15	.74	1.09	1.10	1.67	1.70
Ca0 3.85	3.78	2.84	1.92	1.52	3.73	3.43	2.04	1.09
Ba001	.00			.11			.00	.00
Na20 2.46	.77	2.79	2.85	3.09	.38	.38	2.30	2.34
K20 1.33	1.76	1.84	1.88	2.41	3.53	3.55	3.35	3.41
H <sub>2</sub> 0 <sup>+</sup> 10.58	11.77	8.93	9.12	6.81	10.63	10.70	8.33	8.49
H <sub>2</sub> 0 <sup>-</sup> 7.80	7.19	4.72	4.82	6.51	5.74	5.78	7.95	8.11
.15	.02	.13	.13	.06	.07	.07	.03	.03
.03	.01	.02	.02	.01	.01	.01	.02	.02
Mn001	.00	.06	.06	.00	.01	.01	.00	.00
.04	.02	.75		.04	.25		.76	
Total 99.26	99.36	99.68	100.00	99.31	99.89	100.00	99.86	100.00

#### Composition of unit cell, in atoms per unit cell

 $[Fe^{+2}, Ti, P, and Mn were omitted in the calculation of the unit cell]$ 

\$1 27.90	28.21		28.51	29.23		28.08		28.32
A1 8.09	7.66		6.85	6.44		7.53		7.62
Fe <sup>+3</sup> 15	.22		.80	.38		.53		.08
Mg28	.91		.78	.50		.77		1.18
Ca 1.98	1.95		.93	.74		1.73		.54
Ba				.02				
Na 2.30	.72		2.51	2.70		.34		2.12
.82 K	1.08		1.09	1.39		2.13		2.03
H <sub>2</sub> 0 <sup>+</sup> 17.00	18.88		13.79	10.26		16.76		13.22
H20 12.54	11.54		7.29	9.81		9.05		12.63
0 72.00	72.00		72.00	72.00		72.00		72.00
S1: A1+Fe <sup>+3</sup> 3.38	3.58		3.73	4.28		3.48		3.68
Analysis	Lab No.	Locality No. (fig	g. 1)	Analysis	Lab No.	Lo	ocality No	o. (fig. 1)
1	D103736	4		4	D103738			54
2	D103737	4		5	D103020			28

5

D103739

6

81

3

D103021

7







Figure 3.--Atomic percentages of Na, K, and Ca+Mg+Ba for chabazite (dots), clinoptilolite (triangles), and erionite (squares). Numbers refer to analyses given in table 2.

ratio greater than 4 (Boles, 1972). The Al+Fe<sup>+3</sup> :Na+K+2(Ca+Mg) ratio for zeolites should be unity, but this ratio is about 1.1 for analysis no. 3. The Fe<sub>2</sub>O<sub>3</sub> content of this clinoptilolite is much higher than that for any of the other Durkee zeolites (table 2). Thus, some of the iron in this particular clinoptilolite is probably an impurity, and the true Si: Al+Fe<sup>+3</sup> ratio may be higher than that shown on figure 2. Table 2 and figure 3 show that, relative to each other, the Durkee erionites are potassic, the clinoptilolites are sodic, and the chabazites are calcic. Figure 3 shows that four of the six zeolites are rich in alkalis, that one chabazite (no. 2) is rich in alkaline-earth elements.

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