

(Tvst), and the tuffaceous conglomerate (Ttc). Thickness not known with certainty because of poor exposures, but probably from 0 to about 50 m Ta Hornblende andesite (Miocene) -- Purplish-brown to medium gray, andesite lava flow, flow breccia, and intrusive breccia. Phenocrysts (5-6 percent of rock volume), typically about 1 mm long but some as much as 5 mm long, consist of 5 percent hornblende needles, and 1 percent plagioclase. Groundmass locally partly glassy. Phenocrysts and groundmass generally are altered, but are relatively fresh locally. Rock overlies and intrudes Precambrian Signal Granite (pes) and is overlain by basin-fill deposits (Tbf). Locally, flow and flow breccia of this unit are associated in unclear contact with sandstone that interfingers with breccia (Tbx) and boulder conglomerate (Tcg). The andesite thus may be as old as these units. Thickness from 0 to about 60 m To Conglomerate (Miocene) -- Pebble-, cobble-, and boulder conglomerate containing sandstone interlayers. Moderately to poorly consolidated, typically brown to reddish-orange, weathering into light-toned, boulder-strewn slopes with few outcrops. Coarsegrained porphyritic granite (pes) is a conspicuous clast type that in places unit can be mistaken for spheroidally weathered granite. Typically occurs as isolated patches mantling the granite. May correlate with Tbfc, Ttc, and Tcg. Thickness unknown Tvr Rhyodacite (Miocene?) -- Intrusive mass of light-gray to pale-pink porphyritic hornblende-biotite rhyodacite. Phenocrysts, which comprise about 30 percent of rock, consist of euhedral uncorroded plagioclase typically 1 mm long, with thin rims of iron oxide; hornblende as much as 4 mm long, fresh to completely altered to calcite and brown clays; biotite, about 1 mm in size, fresh to moderately altered; and rounded and embayed quartz as much as 2 mm in diameter. Groundmass consists of plagioclase laths, minor biotite and hornblende in a brown, dusty, relatively altered microcrystalline matrix. Unit intrudes Precambrian Signal Granite (pes)

Mesa 2900 (Tmbo); the more tilted rocks probably correlate with the tuffaceous sandstone (Tss), the pyroclastic and epiclastic rocks

and is overlain by basin-fill deposits (Tbf) Mzp6i Silicic and intermediate intrusive rocks (Mesozoic of Precambrian)-Pinkish-gray, grayish-orange-pink and very pale-orange, fine- to coarse-grained, equigranular leucogranite; fine-grained aplite; medium- to coarse-grained pegmatite; and leucocratic to mesocratic medium- to fine-grained biotite granodiorite. Typically fresh Occurs as a subequant stock measuring about 250 by 450 m; as thick dike-like bodies 50-70 m wide and about 500 m long; as numerous dikes a few meters thick and as much as 900 m long; and as abundant dikes and veins too small to map. Tabular bodies typically trend

ped Diabase (Precambrian?) -- Medium-gray to dark-greenish-gray, fine- to medium-grained diabase dikes, mostly with well-developed ophitic texture. Diabase in many places forms tabular bodies that intrude granite and gneiss and trend northwest to north-northwest. dically weathers to rubble and forms few good exposures per Hetarhyolite porphyry (Precambrian?) -- Light-gray to pinkish-gray rhyolite porphyry dikes. Contains l to 1.5-mm-long rounded quartz phenocrysts in altered felsic groundmass. No flow banding visible. Very hard. Weathers into angular fragments that litter

pes Signal Granite (Precambrian) -- Light-gray to medium-brownish-gray leucocratic to mesocratic, unfoliated to weakly foliated, typically massive, locally jointed and spheroidally-weathered, medium- to coarse-grained porphyritic biotite granite or monzogranite. Phenocrysts are subequant, typically 2-3 cm long, maximum 5 cm, composed of twinned potassium feldspar. A coarse-grained equigranular phase occurs locally. In places, contains sparse xenoliths of melanocratic diorite and biotite schist, and xenoliths of leucocratic to mesocratic gneiss. Also contains local zones of closely spaced subparallel joints in areas of structural disturbance. Intruded by quartz, aplite, pegmatite, leucogranite, metarhyolite porphyry, and diabase. Leucocratic and mesocratic phases form coherent bodies measuring hundreds of meters. Cut by numerous shear zones, locally mineralized, showing evidence of weak recrystallization. The age assignment is based on lithologic and structural similarity with Precambrian granite on Colorado Plateau to east and Hualapai Mountains to north. Best exposed in the Artillery Mountains, in the northern part of the Artillery Peak quadrangle, between the ghost town of Signal (SW 44 sec. 9, T. 1 I., R. 13 W.), and Eagle Point (E 1/2 sec. 7, T. 12 N., R. 14 W.). New name, proposed by Lucchitta and Suneson (1982)

pegn Gneiss (Precambrian) -- Includes a variety of gneissic rocks that are predominantly orthogneisses, are affected by open to isoclinal folding, and locally are migmatitic. Chief lithologies are: (a) Light- to medium-gray, leucocratic to mesocratic equigranular gneissic granite, locally porphyritic. Where porphyritic, includes 2-5 cm subequant crystals of potassium feldspar. Probably a marginal facies of the Signal Granite (pes). (b) Coarse-grained, equigranular leucocratic to mesocratic, weakly to moderately foliated biotite granite gneiss. (c) Dark-gray, medium- to finegrained, equigranular, mesocratic to melanocratic, well-foliated biotite granite gneiss. Locally contains bands of well-defined augen composed of potassium feldspar, and has a mylonitic fabric. Locally contains garnet.

The unit is intruded by quartz, pegmatite, and aplite veins, and by diabase, metarhyolite and leucogranite dikes. Locally it is cut by shear zones that are marked by crushing, mylonitization, and alteration, and are intruded by pegmatite, aplite, and quartz veins. Similar to and continuous with Precambrian rocks in Hualapai Mountains to north and Colorado Plateau to east. \_\_\_\_ Contact-Dashed where approximately located or gradational. In silicic

volcanic rocks (Tvs), contacts separate intrusions of slightly different ages and y/o indicates younger/older where determined ?...- Fault-Dashed where approximately located, dotted where concealed, queried where inferred. Arrow shows direction and amount of dip. Crossbar denotes vertical fault. Bar and ball on downthrown side, where determined

~ ~ ~ ~ Shear zone-Zone of crushing, brecciation, mineralization, and bleaching, for which no displacement can be determined Mineralized vein

ΔΔΔΔ Area of intense brecciation \* Vent area

Anticline-Showing crestline and direction and amount of plunge Syncline-Showing troughline and direction of plunge. Dashed where approximately located

Strike and dip Dot at end of leader shows location of measurement

Apparent (in areas of low dip and poor exposure) Generalized (crumpled, plicated, or indulating beds)

quadrangle.

Inclined showing range of dips

## GEOLOGIC SETTING

The Artillery Peak NW quadrangle is one of a series of quadrangles being mapped in a northeast-trending belt extending from the metamorphic-corecomplex terrane of the Rawhide, Buckskin, and Whipple Mountains to the Colorado Plateau near Wikieup, Arizona.

Four major sequences of rocks are present. From oldest to youngest they (1) Basement rocks, which consist of: (a) Banded gneiss, augen gneiss, granite gneiss, and medium-grained equigranular granitic rocks, all indistinguishable from similar rocks of Precambrian age in the Hualapai Mountains to the north and the Colorado Plateau to the east. (b) Coarsely porphyritic granite or monzogranite forming a large pluton. Age unknown, but thought to be Precambrian because of lithologic similarity with known Precambrian granites (the Prescott Granodiorite, Oracle (of Peterson, 1938), and Ruin Granites, all in Arizona). (c) Metarhyolite and diabase dikes trending northeast and dipping southwest. Age unknown, but probably Precambrian. (d) Aplite pegmatite, and leucogranite dikes trending northeast, and one small leucogranite stock. Age unknown, most likely Precambrian or Laramide.

(2) Fluvio-lacustrine rocks, with interlayered basaltic and rhyolitic rocks, that were deposited in basins not related to the present topography and that now are highly deformed. The lower part of the sequence reflects drainage chiefly from the northeast, whereas the upper part reflects drainage chiefly from the southwest. Because of the several basins of deposition, marked lateral changes in facies and thickness are the norm. These rocks are analogous to the Miocene Horse Spring Formation and related deposits in the Lake Mead area (Longwell, 1921, 1922; Bohannon, 1980). The Artillery Peak NW quadrangle is near the northeastern limit of deposition of fluvio-lacustrine rocks of the Bill Williams River area.

(3) Interior-basin deposits filling the present basins, which were formed by basin-range tectonism. Basaltic lava flows associated with these deposits typically form conspicuous mesas or gently sloping cuestas. All these basin-filling deposits are analogous to the Muddy Creek Formation, the Big Sandy Formation of Sheppard and Gude (1972), the fanglomerate of Metzger and others (1973), and the fanglomerate of Osborne Wash of Dickey and others (1980).

(4) Deposits related to through-flowing drainage. Chiefly piedmont-slope deposits, alluvial deposits, and basaltic lava flows that have followed present drainages. Probably younger than 5.5 m.y. (Lucchitta, 1972,

Rocks of the second, or fluvio-lacustrine, sequence are tilted southwest and cut by faults trending northwest and dipping northeast. Based on to be listric and associated with an episode of low-angle detachment faulting and intense extension.

Interior-basin deposits are much less deformed than older rocks, but nevertheless are affected by basin-range faulting and mild warping. Faulting and deposition were largely synchronous, but faulting ceased before the end of deposition so that fault traces are buried by the youngest interior-basin The fourth, or youngest, group of deposits is not cut by faults in this

None of the rocks of sequence 2 are Eocene(?) or Oligocene in age, as had previously been believed (Lasky and Webber, 1949; Gassaway, 1977, 1978; Otton 1977: Eberly and Stanley, 1978; Shackelford, 1980). Our work in this and four quadrangles adjacent to the west indicates that these rocks are Miocene because: (a) K-Ar age determinations on basalts of the older basalt sequence range from about 16.5 to about 18.7 m.y. (Suneson and Lucchitta, 1979); (b) These basalts are interbedded with the Miocene Peach Springs Tuff of Young and Brennan (1974); (c) K-Ar age determinations on the silicic volcanic rocks are

from the lower part of the sequence have yielded Miocene microfossils (R. F. Hevly, written communication, 1979). Basalts that are part of sequence 3 (interior-basin deposits) have yielded K-Ar ages ranging from 9 to 13 m.y. (Suneson and Lucchitta, 1979).

mostly in the 12 to 13 m.y. range (Suneson and Lucchitta, 1979); and (d) Rocks

## REFERENCES CITED

Bohannon, R.G., 1980, Nonmarine sedimentary rocks of Tertiary age in the Lake Mead region, southeastern Nevada and northwestern Arizona: U. S. Geological Survey Professional Paper 1259, 72 p. Dickey, D.D., Carr, W.J., and Bull, W.B., 1980, Geologic map of the Parker NW, Parker, and parts of the Whipple Mountains SW and Whipple Wash quadrangles, California and Arizona: U. S. Geological Survey Miscellaneous Geologic Investigations Map I-1124, scale 1:24,000. Eberly, L.D., and Stanley, T.B., Jr., 1978, Cenozoic stratigraphy and geologic history of southwestern Arizona: Geological Society of America Bulletin, Gassaway, J.S., 1977, A reconnaissance study of Cenozoic geology in westcentral Arizona: San Diego State University Master's Thesis, 117 p. , 1978, Gravity tectonics and stratigraphic relations in the Transition Zone, Yuma and Mohave Counties, Arizona: Geological Society of America Abstracts with Programs, v. 10, no. 3, p. 106. Lasky, S.G., and Webber, B.N., 1949, Manganese resources of the Artillery Mountains region, Mohave County, Arizona: U.S. Geological Survey Bulletin Longwell, C. R., 1921, Geology of the Muddy Mountains, Nevada, with a section to the Grand Wash Cliffs in western Arizona: American Journal of Science, 5th Series, v. 1, p. 39-62. , 1922, The Muddy Mountain overthrust in southeastern Nevada: Journal of Geology, v. 30, p. 63-72. Lucchitta, Ivo, 1972, Early history of the Colorado River in the Basin and Range Province: Geological Society of America Bulletin, v. 83, p. 1933-, 1979, Late Cenozoic uplift of the southwestern Colorado Plateau and adjacent lower Colorado River region: Tectonophysics, v. 61, p. 63-95. Lucchitta, Ivo and Suneson, N.H., 1982, Signal Granite (Precambrian), Westcentral Arizona: U.S. Geological Survey Bulletin 1529-H, p. H87-H90. Metzger, D.G., Loeltz, O.J., and Irelan, Burdge, 1973, Geohydrology of the Parker-Blythe-Cibola area, Arizona and California: U.S. Geological Survey

Professional Paper 486-G, 130 p. Otton, J.K., 1977, Geology of uraniferous Tertiary rocks in the Artillery Peak - Date Creek basin, west-central Arizona: U.S. Geological Survey Circular 753, p. 35-36. Peterson, N.P., 1938, Geology and ore deposits of the Mammoth mining camp area, Pinal County Arizona: Arizona Bureau of Mines Bulletin 144, Geologic Series II, 63 p. Shackelford, T.J., 1980, Tertiary tectonic denudation of a Mesozoic-early Tertiary (?) gneiss complex, Rawhide Mountains, western Arizona: Geology, v. 8. p. 190-194. Sheppard R.A., and Gude, A.J., III, 1972, Big Sandy Formation near Wikieup,

Mohave County, Arizona: U.S. Geological Survey Bulletin 1354-C, 10 p.

Suneson, Neil, H., 1980, The origin of bimodal volcanism, west-central Arizona: University of California, Santa Barbara, Ph.D. dissertation, 282 Suneson, Neil H., and Lucchitta, Ivo, 1979, K/Ar ages of Cenozoic volcanic rocks, west-central Arizona: Isochron/West, no. 24, p. 25-29. Young, R.A., and Brennan, W.J., 1974, Peach Springs tuff: its bearing on structural evolution of the Colorado Plateau and development of Cenozoic drainage in Mohave County, Arizona: Geological Society of America

Bulletin, v. 85, p. 83-90.

(Tbft), and the tuffaceous sediments associated with the basalt of