

UNITED STATES DEPARTMENT OF THE INTERIOR  
GEOLOGICAL SURVEY

Recognition and geologic implications of potassium metasomatism  
in upper-plate volcanic rocks at the detachment fault at the  
Harcuvar Mountains, Yavapai County, Arizona

by

William E. Brooks<sup>1</sup>

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This report is preliminary and has not been reviewed  
for conformity with U.S. Geological Survey editorial  
standards and stratigraphic nomenclature.

<sup>1</sup>Denver, Colorado

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## INTRODUCTION

Potassium metasomatism of diverse origins has affected volcanic and sedimentary rocks from California to Colorado (Rattee and Steven, 1967; Wones and others, 1967; Chapin and Glazner, 1983; Chapin and Lindley, 1986). In southwestern Arizona, mid-Tertiary volcanic rocks that range in composition from basalt to rhyolite may have low  $\text{Na}_2\text{O}$  and anomalously high  $\text{K}_2\text{O}$  content which is indicative of metasomatism (Brooks, 1985a, 1986; Kerrich and others, 1986). Based on their anomalously high  $\text{K}_2\text{O}$  content alone, ash-flow tuffs in the upper-plate of the detachment fault at the Harcuvar Mountains have been referred to as as high-potassium trachytes (Scarborough and Wilt, 1979) and trachytes (Reynolds and Spencer, 1984).

Analytical data are important in resolving the controversy over the magmatic (Shafiqullah and others, 1976) or metasomatic (Brooks, 1985a, 1986) origin of some of the high potassium rocks. Because of the limited number of previous rock analyses (Brooks, 1984) and discordant potassium-argon (K-Ar) and fission-track (FT) dates (some by as much as 7 m.y.) (Scarborough and Wilt, 1979; Brooks, 1985b; Brooks and Marvin, 1985), samples of the upper-plate volcanic rocks (fig. 1) were submitted for analysis.

## DISCUSSION

The metasomatic nature of high-potassium volcanic rocks is more easily recognized in the laboratory than in the field, where samples of metasomatised lavas may appear no different than their unaltered counterparts. In thin-section, metasomatic alteration is confirmed by the chalky appearance of plagioclase, alteration of groundmass pyroxene and magnetite to hematite, and patchy replacement of plagioclase by adularia (Chapin and Lindley, 1985). Rock chemistry, particularly a  $\text{K}_2\text{O}:\text{Na}_2\text{O} > 2$ , may indicate metasomatic

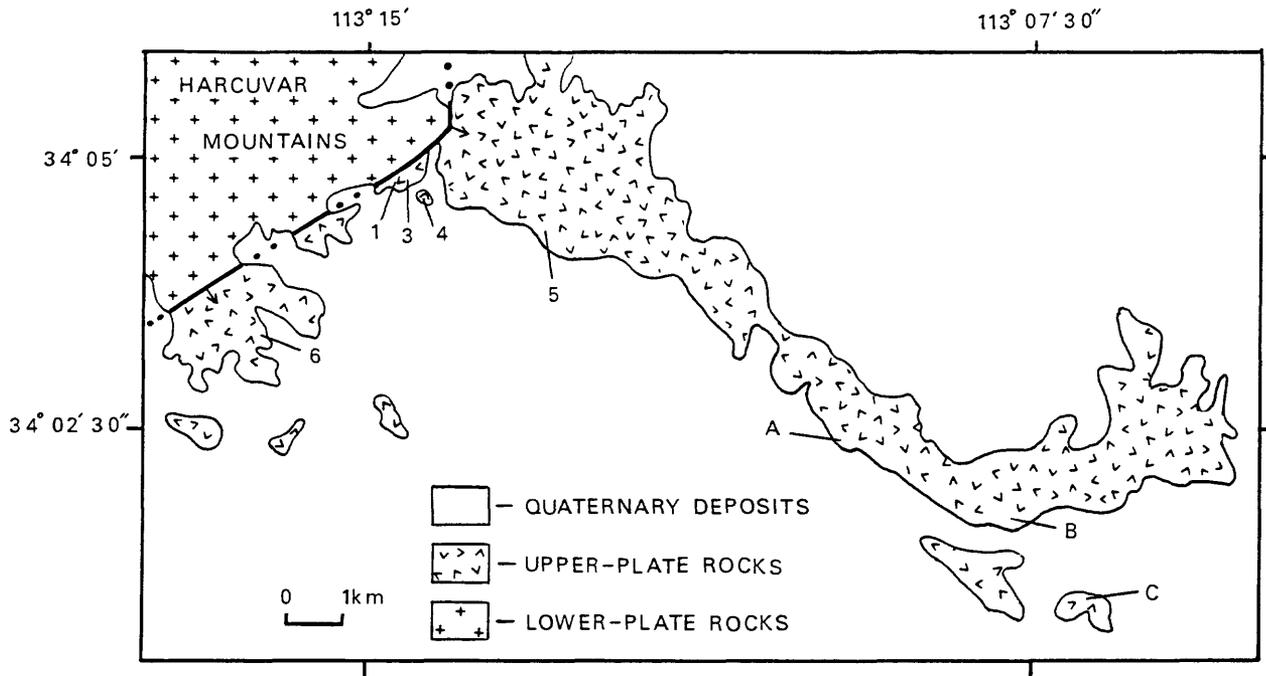


Figure 1.--Locations of analyzed samples of upper-plate volcanic rocks given in Table 1(HM prefix deleted for brevity). Heavy line indicates trace of detachment fault and arrows indicate direction to upper-plate. A=K-Ar site (Scarborough and Wilt, 1979); B=rock analysis site (Brooks, 1984); and C=K-Ar, fission-track, and rock analysis site (Brooks, 1984; 1985b).

alteration. Other indications of metasomatic alteration include the presence of 2-10 $\mu$ m orthoclase as adularia(?), secondary quartz and calcite, propylitic alteration, XRF analyses with a high or erratic LOI, and the absence of leucite as a host phase for the excess potassium.

Potassium content, alone, does not indicate metasomatism. For example, the K<sub>2</sub>O content of the ash-flow tuffs from the upper-plate of the detachment fault at the Harcuvar Mountains (HM-1, 3, 4, 5A, 5B; Table 1) is comparable to the K<sub>2</sub>O content of the unmetasomatised Topopah Springs tuff of southern Nevada (Lipman and others, 1966). However, metasomatic alteration has reduced the Na<sub>2</sub>O content and thereby increased the K<sub>2</sub>O:Na<sub>2</sub>O of the Harcuvar ash-flow tuffs to as much as 38:

|                                    | Harcuvar Mtns. tuff, AZ | Topopah Spring tuff, NV |
|------------------------------------|-------------------------|-------------------------|
| K <sub>2</sub> O wt%               | 4.3 - 8.6               | 4.7 - 6.0               |
| Na <sub>2</sub> O wt%              | 0.2 - 1.6               | 3.7 - 4.4               |
| K <sub>2</sub> O:Na <sub>2</sub> O | 2.7 - 38.0              | 1.3 - 1.5               |

An analysis of the Bullard andesite, mapped by Reynolds and Spencer (1984), shows a K<sub>2</sub>O content of 4.2 weight percent (HM-6; Table 1). This is only somewhat higher than K<sub>2</sub>O contents of 3.2 and 2.1 weight percent given by Taylor (1969) and Gill (1981), respectively, for high-potassium andesites. The low Na<sub>2</sub>O of 0.2 weight percent, high K<sub>2</sub>O:Na<sub>2</sub>O of 21, and high SiO<sub>2</sub> of 70.3 weight percent of this "hand-sample andesite" demonstrate the effects of metasomatism on rock chemistry.

In metasomatised upper-plate volcanic rocks, it is common to find discordant FT and K-Ar dates. Resetting of temperature-sensitive FT dates to 17-18 Ma with respect to K-Ar dates of 24 Ma has occurred in the upper-plate rocks at the Harcuvar Mountains and elsewhere in southwestern Arizona (Brooks and Marvin, 1985; Brooks, 1985c). Therefore, it is unlikely that a single

date, whether FT or K-Ar, from a metasomatised lava gives the true age of eruption.

Because the potassium content of volcanic rocks from southwestern Arizona has been used to establish depth-to-plate reconstructions for the Tertiary (Dickinson and Hatherton, 1967; Lipman and others, 1971; Rowell and Edgar, 1983) it is important to discriminate between analyses of metasomatised lavas and primary potassic lavas. More than just high  $K_2O$  content must be considered if analyses of the rocks are to be used in petrology, geochronology, and tectonic modeling. Evaluation of major oxides,  $K_2O:Na_2O$ , and petrography indicates that potassium metasomatism has affected the upper-plate volcanic rocks at the Harcuvar Mountains. Therefore, magmatic chemistry should not be inferred from major oxide analyses of these rocks.

Table 1.--Analyses of upper-plate volcanic rocks at the  
Harcuvar Mountains, Yavapai County, Arizona

| Field no.   | HM-1 | HM-3 | HM-4 | HM-5A | HM-5B | HM-6  |
|---|------|------|------|-------|-------|-------|
| Major elements (weight percent, uncorrected) determined by x-ray spectroscopy, analysts, A. J. Bartel, K. Stewart, J. Taggart;<br>FeTO <sub>3</sub> indicates total iron reported as Fe <sub>2</sub> O <sub>3</sub> |      |      |      |       |       |       |
| SiO <sub>2</sub>  | 63.0 | 74.8 | 77.5 | 77.2  | 73.5  | 70.3  |
| Al <sub>2</sub> O <sub>3</sub>  | 9.3  | 11.6 | 9.3  | 10.1  | 12.8  | 6.6   |
| FeTO <sub>3</sub>   | 1.3  | 1.2  | 1.3  | 1.3   | 2.2   | 3.3   |
| MgO   | 0.3  | 0.2  | 0.1  | 0.3   | 1.0   | 0.3   |
| CaO   | 9.6  | 0.7  | 0.4  | 0.4   | 0.4   | 5.6   |
| Na <sub>2</sub> O   | 0.2  | 0.3  | 0.3  | 0.4   | 1.6   | 0.2   |
| K <sub>2</sub> O  | 7.6  | 8.6  | 6.4  | 8.4   | 4.3   | 4.2   |
| TiO <sub>2</sub>  | 0.1  | 0.1  | 0.2  | 0.1   | 0.2   | 0.5   |
| P <sub>2</sub> O <sub>5</sub>   | 0.1  | <0.1 | 0.1  | <0.1  | 0.1   | 0.2   |
| MnO   | 0.1  | 0.1  | 0.7  | 0.1   | 0.1   | 0.2   |
| LOI 900°C   | 8.0  | 1.7  | 1.6  | 1.3   | 3.0   | 5.8   |
| K <sub>2</sub> O:Na <sub>2</sub> O  | 38.0 | 28.7 | 21.3 | 21.0  | 2.7   | 21.0  |
| Trace elements (ppm) determined by ICP method,<br>analyst P. H. Briggs; --, not reported  |      |      |      |       |       |       |
| Mn  | 910  | 720  | 5300 | 670   | 320   | 1700  |
| Ag  | 6    | 3    | <2   | <2    | <2    | 2     |
| As  | <10  | 20   | <10  | <10   | <10   | <10   |
| Au  | <8   | <8   | <8   | <8    | <8    | <8    |
| B   | --   | --   | --   | --    | --    | --    |
| Ba  | 2000 | 2300 | 960  | 1300  | 380   | 550   |
| Be  | 2    | 2    | 1    | 2     | 3     | <1    |
| Bi  | <10  | <10  | <10  | <10   | <10   | <10   |
| Cd  | <2   | <2   | <2   | <2    | <2    | 13    |
| Ce  | 70   | 55   | 170  | 67    | 72    | 28    |
| Co  | 3    | 2    | 5    | 2     | 3     | 58    |
| Cr  | 2    | <1   | 2    | 5     | 7     | 150   |
| Cu  | 9    | 8    | 7    | 2     | 2     | 14000 |
| Eu  | <2   | <2   | <2   | <2    | <2    | <2    |
| Ga  | 11   | 20   | 15   | 16    | 16    | 8     |
| Ge  | --   | --   | --   | --    | --    | --    |
| Ho  | <4   | <4   | <4   | <4    | <4    | <4    |
| La  | 61   | 37   | 91   | 34    | 47    | 16    |
| Li  | 22   | 25   | 110  | 37    | 41    | 55    |
| Mo  | <2   | <2   | <2   | <2    | <2    | 6     |
| Nd  | 25   | 27   | 45   | 26    | 28    | 16    |
| Ni  | 4    | <2   | 7    | 3     | 4     | 26    |
| Pb  | 9    | 39   | 59   | 8     | <4    | 180   |
| Sc  | <2   | <2   | <2   | <2    | 3     | 10    |
| Sn  | <20  | <20  | <20  | <20   | <20   | <20   |
| Ta  | <40  | <40  | <40  | <40   | <40   | <40   |
| Th  | 9    | 18   | 10   | 15    | 8     | <4    |
| U   | <100 | <100 | <100 | <100  | <100  | <100  |
| V   | 8    | 17   | 23   | 9     | 12    | 62    |
| W   | --   | --   | --   | --    | --    | --    |
| Yb  | 2    | 3    | 2    | 3     | <1    | <1    |
| Zn  | 120  | 200  | 270  | 24    | 17    | 1000  |
| Trace elements (ppm) determined by an energy dispersive analyzer Cd <sup>109</sup> source<br>for Rb, Sr, Y, Zr, and Nb; analyst M. J. Blaskowski  |      |      |      |       |       |       |
| Rb  | 193  | 252  | 160  | 210   | 224   | 31    |
| Sr  | 67   | 154  | 90   | 15    | 27    | 56    |
| Y   | 13   | 32   | 17   | 23    | 17    | 7     |
| Zr  | 151  | 147  | 143  | 164   | 203   | 32    |
| Nb  | 12   | 27   | 11   | 29    | 16    | 2     |

Appendix 1.--Locations of samples for analytical data given in Table 1.

| Field no. | Latitude N. | Longitude W. |
|-----------|-------------|--------------|
| HM-1      | 34°04'44"   | 113°14'39"   |
| HM-3      | 34°04'44"   | 113°14'33"   |
| HM-4      | 34°04'40"   | 113°14'25"   |
| HM-5A     | 34°04'22"   | 113°13'04"   |
| HM-5B     | 34°04'22"   | 113°13'04"   |
| HM-6      | 34°03'23"   | 113°16'12"   |

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