



METHODS FOR ESTIMATING MONTHLY STREAMFLOW CHARACTERISTICS AT UNGAGED SITES IN WESTERN MONTANA

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U.S. GEOLOGICAL SURVEY

Open-File Report 89-40



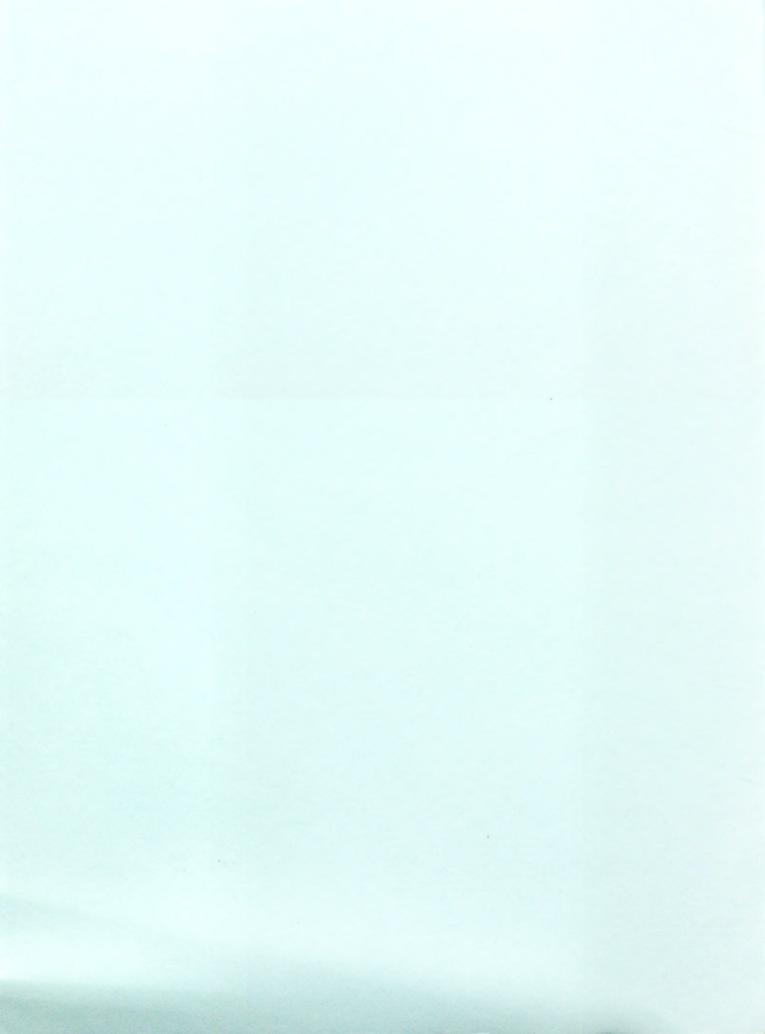
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U.S. BUREAU OF INDIAN AFFAIRS and the

CONFEDERATED SALISH AND KOOTENAI TRIBES

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AT UNGAGED SITES IN WESTERN MONTANA

By Charles Parrett, U.S. Geological Survey, and Kenn D. Cartier, Confederated Salish and Kootenai Tribes

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CONVERSION FACTORS

The following factors can be used to convert inch-pound units in this report to metric (International System) units.

Multiply inch-pound unit	By	To obtain metric unit
cubic foot per second (ft ³ /s) foot (ft) inch (in.) mile (mi) square mile (mi ²)	0.028317 0.3048 25.4 1.609 2.590	cubic meter per second (m ³ /s) meter (m) millimeter (mm) kilometer (km) square kilometer (km ²)

 $\underline{\text{Sea level}}$: In this report "sea level" refers to the National Geodetic Vertical Datum of 1929 (NGVD of 1929)—A geodetic datum derived from a general adjustment of the first-order level nets of both the United States and Canada, formerly called "Sea Level Datum of 1929."

METHODS FOR ESTIMATING MONTHLY STREAMFLOW CHARACTERISTICS AT UNGAGED SITES IN WESTERN MONTANA

by

Charles Parrett and Kenn D. Cartier

ABSTRACT

Three methods for estimating mean monthly discharge and various points on the daily mean flow-duration curve for each month (daily mean discharges exceeded 90, 70, 50, and 10 percent of the time each month) were developed for western Montana. A procedure for weighting two or more individual estimates to provide a minimum-variance-weighted-average estimate also was developed. This report describes the estimation methods developed and their reliability and limitations.

The first method is based on multiple-regression equations relating the monthly streamflow characteristics to various basin and climatic variables. Standard errors of the basin-characteristics equations range from 43 to 107 percent. The basin-characteristics equations are generally not applicable to streams that receive or lose water as a result of localized geologic features or to stream sites with appreciable upstream storage or diversions.

The second method is based on regression equations relating the monthly streamflow characteristics to channel width. Standard errors of the channel-width estimating equations range from 41 to 111 percent. The channel-width equations are generally not applicable to stream sites with exposed bedrock, with braided or sand channels, or with recent alterations.

The third method requires 12 once-monthly streamflow measurements at the ungaged site of interest. The 12 measured flows are then correlated with concurrent flows at some nearby gaged site using the curve-fitting technique MOVE.1 (Maintenance of Variance Extension, Type 1), and the relation defined is used to estimate the required monthly streamflow characteristic at the ungaged site from the streamflow characteristic at the gaged site. Standard errors, which are estimated by applying the method to 20 other gaged sites, range from 19 to 92 percent. Although generally substantially more reliable than either the basin-characteristics method or the channel-width method, this method may yield unreliable results if the measurement site and the index gaged site are not hydrologically similar.

The procedure for weighting individual estimates is based on the variance and degree of independence of the individual estimating methods. Standard errors for the weighted estimates of the monthly flow characteristics range from 15 to 43 percent when all three methods are used. The weighted-average estimates from all three methods are generally substantially more reliable than any of the individual estimates.

INTRODUCTION

Although western Montana generally has abundant surface water, shortages are common because of the large areal and seasonal variability of runoff. Making sound management decisions to relieve periodic shortages and to most efficiently allocate the supply among competing users thus requires reliable information about the variability of streamflow. In particular, the distribution of daily mean discharge by month is of interest to fish and wildlife managers, water-rights administrators, and other land- and water-use planners and managers. Unfortunately, techniques for estimating monthly streamflow characteristics are not as readily available as techniques for estimating annual and peak streamflow characteristics. For example, the only U.S. Geological Survey report containing estimating equations for mean monthly discharge in Montana is one by Boner and Buswell (1970); that report is based on a relatively small number of streamflow-gaging stations having at least 10 years of record then available. A more recent report by Parrett and Hull (1985, p. 8,9) indicates that mean monthly discharge can be estimated at an ungaged site by using existing techniques to estimate a mean annual discharge and then assuming that the monthly distribution of the annual discharge follows the same distribution as some nearby gaged site. The accuracy of the estimated monthly mean discharge using this technique, however, is not completely satisfactory in western Montana.

Because of the dearth of techniques available for estimating monthly stream-flow characteristics at ungaged sites in western Montana, the present study was undertaken in 1985 in cooperation with the U.S. Bureau of Indian Affairs and the Confederated Salish and Kootenai Tribes of the Flathead Indian Reservation. The objective of the project was to develop techniques for estimating long-term mean monthly discharge and various points on the daily mean flow-duration curve for each month (daily mean discharges exceeded 90, 70, 50, and 10 percent of the time each month) that would be applicable within the boundaries of the Flathead Indian Reservation in western Montana.

Purpose and Scope

The purposes of this report are to describe the estimation methods that were developed and to discuss their reliability and limitations. Three methods for estimating the required discharges were developed. One method is based on the relation between streamflow and various basin and climatic variables. The second method is similar to the first and is based on the relation between discharge and channel width. The third method requires once-monthly measurements of discharge at the ungaged site of interest and is based on the relation between the measured discharges and concurrent daily mean discharges at a similar, nearby gaged site. A procedure also is presented for weighting the individual estimates of discharge made from two or more of the three separate methods. The weighted-average estimate is based on the variance and degree of independence of the individual estimating methods. Calculated standard errors of prediction are used as a measure of reliability of each estimating method, and experience gained in the development and application of the methods is used to describe the major limitations.

Description of Study Area

Because of the small number of streamflow-gaging stations with monthly discharge data within the Flathead Indian Reservation, the study area was expanded to

include the entire part of the State within the upper Columbia River basin as well as the adjacent eastern side of the Rocky Mountains (fig. 1). This area, termed "western Montana" for the purposes of this report, is composed largely of north-to northwest-trending mountain ranges separated by long, fairly narrow valleys. Except for the valley-floor areas, the study area is generally rugged and forested. The flatter valleys are mostly cultivated or grazed. The Flathead Indian Reservation, like the larger study area, is composed of both mountains and valleys. The reservation is bounded on the east by the rugged Mission Mountains and on the south and west by less rugged and prominent mountains. Much of the interior part of the reservation includes broad intermontane valleys and gently rolling prairies.

Annual precipitation varies widely in the study area, primarily because of orographic effects. Annual precipitation tends to be greatest in the mountains, where it is as much as 100 in. in the northeastern corner of the study area and in the Mission Mountains on the eastern edge of the Flathead Indian Reservation (U.S. Soil Conservation Service, 1981, p. 1-2). In the drier valley areas, including the Little Bitterroot River valley within the Flathead Indian Reservation, annual precipitation is as little as 12 in.

Annual runoff generally follows the precipitation pattern, with greater quantities occurring in the higher elevation areas. Streamflows vary greatly on a seasonal basis, as snowmelt provides the bulk of annual runoff in May, June, and July for the mountain streams and in March, April, and May for the streams draining the lower foothills and valley-floor areas. The smallest streamflows generally occur in late fall and winter when streamflows are almost entirely the result of ground-water inflow. Smaller streams draining the valleys may become dry during this period.

Streamflow Data Used

Monthly streamflow characteristics were computed from data at 59 streamflow-gaging stations within the study area, including 12 stations within the Flathead Indian Reservation. All stations used in the analysis had at least 5 years of record through water year 1986, although some stations did not have complete record for all months. Streamflow-gaging stations where flows are substantially regulated or where large diversions substantially affect most flows were not used in the analyses. The location of the streamflow-gaging stations used is shown in figure 1. The monthly streamflow characteristics computed for each station are listed in table 11 in the Supplemental Data section at the back of the report.

METHODS FOR ESTIMATING MONTHLY STREAMFLOW

Basin-Characteristics Method

One method for estimating streamflow characteristics at ungaged sites uses multiple-regression equations that relate streamflow characteristics at gaged sites to various measured basin and climatic variables. This method, termed the "basin-characteristics method" in this report, has commonly been used in Montana to estimate flood flows and mean annual flows (Omang and others, 1986; Parrett and Omang, 1981; Parrett and Hull, 1985).

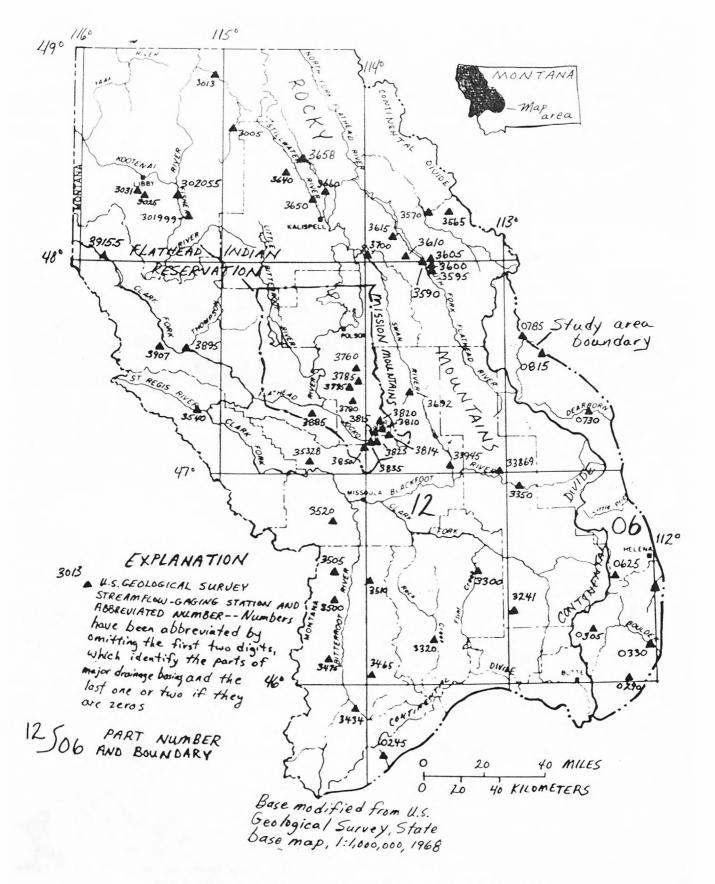


Figure 1.--Location of streamflow-gaging stations.

Because the basin-characteristics method has been widely used, several basin and climatic variables have been previously measured at virtually every U.S. Geological Survey streamflow-gaging station in Montana. These measurement data are stored in the Basin Characteristics File of the U.S. Geological Survey's Water Data Storage and Retrieval System (WATSTORE).

Boner and Buswell (1970) used basin characteristics to develop estimating equations for mean monthly flow in Montana, but the reported accuracy was generally unacceptable. According to Riggs (1972, p. 13-14), the basin-characteristics method is not well suited for the estimation of low flows, because low flows are largely affected by localized geology that cannot be quantified easily. For this study, several previously unmeasured basin characteristics that might be indicative of basin geology were investigated. Eighteen streamflow-gaging stations (table 1) in the study area were randomly selected, and the following geomorphic variables were measured at each site on U.S. Geological Survey topographic maps: basin perimeter, basin slope, circularity ratio, maximum basin relief, drainage density, stream frequency, and aspect.

Table 1.--Streamflow-gaging stations used to investigate new basin characteristics

Formal	Abbreviated station No.	
station No.	(fig. 1)	Stream name
06062500	0625	Tenmile Creek
06078500	0785	North Fork Sun River
12300500	3005	Fortine Creek
12301999	301999	Wolf Creek
12302055	302055	Fisher River
12302500	3025	Granite Creek
12303100	3031	Flower Creek
12324100	3241	Racetrack Creek
12330000	3300	Boulder Creek
12338690	33869	Monture Creek
12343400	3434	East Fork Bitterroot River
12347500	3475	Blodgett Creek
12350500	3505	Kootenai Creek
12356500	3565	Bear Creek
12357000	3570	North Fork Flathead River
12359500	3595	Spotted Bear River
12361500	3615	Graves Creek
12364000	3640	Logan Creek

Basin perimeter, expressed in miles, was determined by measuring the basin drainage area outline on the best-scale topographic map available. Basin slope, which is dimensionless, was determined by measuring the lengths of all contours at a fixed contour interval within the basin, multiplying by the contour interval, and dividing by the basin drainage area. Because the number of contours is largely dependent on the map scale, a single scale (1:24,000 U.S. Geological Survey 7.5-minute quadrangle maps) was used for determining basin slope at all sites; the contour interval selected was 400 ft. The map scale used at any ungaged site needs to be the same to ensure that the equations are applicable. Circularity ratio, which is also dimensionless, was determined by dividing the basin drainage area by the area of a circle with the same basin perimeter. Maximum basin relief, expressed in thousands of feet, was determined by subtracting the elevation of the stream at the basin outlet from the maximum elevation contour within the basin boundary shown on the contour map. Drainage density, expressed in miles per square mile, was determined by measuring and totaling the lengths of all channel segments shown on the contour map and dividing the result by the basin drainage area. with basin slope, only 1:24,000 quadrangle maps were used to determine drainage density and the closely related variable, stream frequency. Stream frequency, expressed as a number per square mile, was determined by dividing the total number of stream segments by the basin drainage area. Aspect, expressed in degrees, was determined by measuring the angle from north to the line connecting the basin centroid to the basin outlet. Measurements of aspect were made either clockwise or counterclockwise from north so that the maximum possible aspect was 180°. Thus, a line from the centroid to the outlet oriented due west would result in an aspect of 90°, as would a line from the centroid to the outlet oriented due east.

The newly measured basin characteristics were combined with 10 standard basin and climatic characteristics previously measured at the 18 stations and treated as independent variables in a multiple-regression analysis. The 10 standard basin and climatic characteristics used were the following: drainage area, percentage of basin above 6,000 ft elevation, main-channel length, mean annual precipitation, mean basin elevation, main-channel slope, percentage of basin covered by forest, percentage of basin composed of lakes and ponds, precipitation, intensity of a storm of 24 hours duration with a recurrence interval of 2 years, and mean January minimum temperature. Individual equations for five monthly flow characteristics for each month (60 equations) were developed using a computerized step-wise regression procedure. Based on this initial analysis, the only new basin characteristics that were significant were basin perimeter, basin slope, circularity ratio, and maximum basin relief. Accordingly, these four new basin characteristics were considered to be worthy of inclusion in a regression analysis using all available streamflow-gaging-station data in the study area, and they were subsequently measured at 54 gaged sites. Suitable topographic maps were not available for four gaged sites (stations 06030500, 06033000, 06061500, and 06081500), so these sites were excluded from the regression analysis. In addition, station 12359000 was excluded from the regression analysis because total streamflows at this site are substantially greater than at any other site used in the analysis.

In the multiple-regression analysis using the 54 gaged sites, the following basin and climatic variables were significant in at least one regression equation:

A drainage area,

E6 percentage of basin above 6,000 ft elevation, plus 1,

PE basin perimeter,

BSL basin slope,

L main-channel length,

P mean annual precipitation,

E mean basin elevation,

BR maximum basin relief.

The most significant variable in almost all instances was main-channel length. Main-channel length is more susceptible to human change and measurement error than is drainage area, however, so drainage area was substituted for main-channel length and the regressions were repeated. Because main-channel length and drainage area are highly correlated, the substitution produced no substantial change in regression reliability. Although circularity ratio was determined to be significant in the initial regression analysis using 18 test sites, it was not significant in the analysis using all 54 gaged sites.

Drainage area, expressed in square miles, was determined by planimetering on the best-scale topographic map available. Percentage of basin above 6,000 ft elevation above sea level was determined by planimetering the drainage area above the 6,000-ft contour on the best available topographic map, dividing by the total drainage area, multiplying by 100, and adding 1 to ensure that 0 values did not occur. Mean annual precipitation, expressed in inches, was the basin average precipitation as determined from maps published by the U.S. Soil Conservation Service (1981). Mean basin elevation, expressed in thousands of feet, was determined by overlaying a transparent grid on the basin outline on a topographic map, reading the elevation at the grid intersections, and averaging the readings. The basin and climatic characteristics measured at each streamflow-gaging station used in the regression analysis are listed in table 12 at the back of the report.

Monthly streamflow data and basin and climatic characteristics at the 54 gaged sites in the study area were converted to logarithms and used in a multiple-regression analysis to derive estimating equations of the following linear form:

$$\log Q = \log a + b1 \log B + b2 \log C + \dots$$
 bn $\log N$, (1)

where

Q (dependent variable) is the desired monthly streamflow characteristic in cubic feet per second (daily mean discharge exceeded 90, 70, 50, or 10 percent of the time during the given month, or mean discharge for the month);

a is the multiple-regression constant;

bl, b2, ... bn are the regression coefficients; and

B, C, ... N are values of the significant basin characteristics (independent variables).

Taking antilogarithms yields the following nonlinear form of the regression equation:

$$Q = aB^{b1} \quad C^{b2} \dots N^{bn}, \tag{2}$$

The regressions were performed using a computerized step-wise regression procedure that adds independent variables to the equation one at a time until all significant variables are included. In this study, a variable was included in the model if the F statistic was greater than 5. The computerized procedure also provided statistical measures of the applicability of the derived equations such as standard errors of estimate and coefficients of determination. In general, the smaller the standard error and the larger the coefficient of determination, the more reliable is the estimating equation.

To ensure that estimates from the regression equations for any month would be consistent, the initial equations for some streamflow characteristics were modified. In these instances, variables that were significant in most of the equations for any given month were selected as key variables, and the regressions were repeated using the key variables as the only independent variables. For any given month, the equations for all streamflow characteristics thus have the same independent variables. Complete results of the regression analysis using basin characteristics are given in table 2, along with the coefficients of determination and standard errors associated with each estimating equation.

As indicated by the results in table 2, the basin-characteristics equations generally are more reliable for estimating the higher-flow monthly characteristics (for example, Q.50, Q.10, and QM) than the lower-flow characteristics (Q.90 and Q.70) in any given month. The basin-characteristics equations also generally are more reliable for estimating flow characteristics for the months of high runoff (May and June) than for the months of generally low runoff (July through April).

Channel-Width Method

The second method used in this study for estimating monthly streamflow characteristics at ungaged sites also uses multiple-regression equations developed from gaged data. In this instance, however, monthly streamflow characteristics at gaged sites are related to measured-channel widths at the gaged sites rather than to measured-basin characteristics. This method, termed the "channel-width method" in this report, has been used with generally good success in Montana and elsewhere for the estimation of flood flows and mean annual flows (Cartier, 1984; Parrett and others, 1983; Omang and others, 1983; Hedman and Osterkamp, 1982; Wahl, 1984). Because channel size is presumed to be largely the result of bankfull or near-bankfull flows, the channel-width method generally has not been used for monthly or low-flow characteristics. Nevertheless, the method was investigated for this study because the channel width had previously been measured at most of the gaged sites and because the relation between monthly flow characteristics and bankfull flows is fairly consistent for most perennial streams in the study area.

Channel features previously measured at gaged sites were active-channel width and bankfull width. At most sites the two features were about equally prominent and identifiable.

Osterkamp and Hedman (1977, p. 256) described the active channel as "...a short-term geomorphic feature subject to change by prevailing discharges. The upper limit is defined by a break in the relatively steep bank slope of the active channel to a more gently sloping surface beyond the channel edge. The break in slope normally coincides with the lower limit of permanent vegetation so that the two features, individually or in combination, define the active channel reference

Table 2.--Results of regression analysis based on basin characteristics

[R², coefficient of determination; Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; A, drainage area, in square miles; BR, maximum basin relief, in thousands of feet; BSL, basin slope, dimensionless; QM, mean monthly discharge, in cubic feet per second; P, mean annual precipitation, in inches; E, mean basin elevation, in thousands of feet; PE, basin perimeter, in miles; E6, percentage of basin above 6,000 feet elevation, plus 1]

Month and number of sites	Stream flow charac teristi	-		Equat	ion		\mathbb{R}^2	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
October	Q.90	=	0.123	A ^{0.84}	BR1.31	BSL ^{0.70}	0.69	0.31	82
(50)	Q.70	=	0.246	A ^{0.84}	BR1.21	BSL ⁰ .92	.76	.26	66
	Q.50	=	0.521	A0.80	BR1.07	BSL1.06	.77	.24	60
	Q.10	=	4.68	A ^{0.73}	BR ^{0.40}	BSL1.19	.69	.25	63
	QM	=	1.69	A ^{0.78}	BR ⁰ .59	BSL1.16	.77	.22	54
November	Q.90	=	0.140	A ^{0.89}	BR1.21	BSL0.86	.76	.27	69
(49)	Q.70	=	0.294	A0.84	BR1.22	BSL1.05	.76	.25	63
	Q.50	=	0.711	A ^{0.82}	BR1.00	BSL1.28	.77	• 24	60
	Q.10	=	3.45	A ^{0.85}	BR ^{0.59}	BSL1.74	.76	.24	60
	QM	=	1.19	A ^{0.84}	BR ^{0.83}	BSL1.48	.79	.23	57
December	Q.90	=	0.132	A ^{0.92}	BR1.18	BSL1.01	.76	•27	69
(49)	Q.70	=	0.258	A ^{0.87}	BR1.17	BSL1.11	.79	.24	60
	Q.50	=	0.552	A0.88	BR ^{0.95}	BSL1.33	.79	.24	60
	Q.10	=	2.00	A ^{0.94}	BR ^{0.64}	BSL 1.76	.79	.25	63
	QM -	=	0.874	A ^{0.91}	BR ^{0.84}	BSL1.57	.79	.24	60

Table 2.--Results of regression analysis based on basin characteristics--Continued

Month and number of sites	Stream flow charac terist:	2-	Eq	uation		R^2	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
January	Q.90	=	0.117 A ⁰	.96 _{BR} 1.2	3 _{BSL} 1.25	.77	.29	75
(47)	Q.70	=_	0.276 A ⁰	.93 _{BR} 0.9	6 _{BSL} 1.26	.80	• 24	60
	Q.50	=	0.431 A ⁰	.94 _{BR} 0.8	2 _{BSL} 1.27	.81	• 24	60
	Q.10	=	0.855 A ⁰	.96 _{BR} 0.7	9 _{BSL} 1.36	.79	.25	63
	QM	=	0.424 A ⁰	.96 _{BR} 0.8	8 _{BSL} 1.30	.82	• 24	60
February	Q.90	=	0.176 A ⁰	.98 _{BR} 0.9	9 _{BSL} 1.32	.81	.25	63
(47)	Q.70	=	0.301 A ⁰	.97 _{BR} 0.8	4 BSL1.28	.82	•23	57
	Q.50	=	0.405 A ⁰	.99 _{BR} 0.7	5 _{BSL} 1.35	.84	.22	54
	Q.10	=	1.34 A ¹	.07 _{BR} 0.3	7 _{BSL} 1.66	.80	.26	66
	QM	=	0.590 A ¹	.03 _{BR} 0.6	3 BSL1.53	.83	.23	57
March	Q.90	=	0.174 A ⁰	.99 _{BR} 1.0	3 BSL1.24	.82	.24	60
(48)	Q.70	==	0.307 A ¹	.00 BR0.8	7 _{BSL} 1.31	.84	.23	57
	Q.50	=	0.369 A ¹	.01 _{BR} 0.8	6 BSL1.32	.84	.23	57
	Q.10	=	0.629 A ¹	.05 _{BR} 0.7	7 _{BSL} 1.23	.79	.27	69
	QM	=	0.366 A ¹	.03 _{BR} 0.8	6 BSL1.22	.83	.24	60

Table 2.--Results of regression analysis based on basin characteristics--Continued

Month and number of sites	Stream flow charac terist	c-	Equation	\mathbb{R}^2	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
April	Q.90	=	0.0103 A ^{0.97} p ^{1.43} E ^{-0.86}	.82	.23	57
(49)	Q.70	= .	0.0271 A ^{0.98} P ^{1.50} E ^{-1.29}	.85	.22	54
	Q.50	=	0.0758 A ^{0.96} P ^{1.48} E ^{-1.55}	.83	. 24	60
	Q.10	=	0.119 A ^{1.01} P ^{1.48} E ^{-1.36}	.80	.28	72
	QM	=	0.0708 A ^{1.00} p ^{1.46} E ^{-1.38}	.83	.25	63
May	Q.90	=	0.00100 A ^{1.00} P ^{1.96}	.80	.27	69
(52)	Q.70	=	0.00321 A ^{1.04} P ^{1.75}	.83	.25	63
	Q.50	=	0.00802 Al.01 pl.64	.82	.24	60
	Q.10	=	0.106 A ^{0.91} p ^{1.25}	.84	.21	51
	QM	=	0.0249 A ^{0.96} P ^{1.43}	.84	.22	54
June	Q.90	=	0.122 A ^{0.87} BSL ^{1.06} P ^{1.00} E6 ^{0.17}	.77	.25	63
(53)	Q.70	=	0.144 A ^{0.92} BSL ^{0.98} P ^{1.00} E6 ^{0.18}	.85	.20	49
	Q.50	=	0.245 A ^{0.91} BSL ^{0.95} P ^{0.95} E6 ^{0.19}	.86	.19	46
	Q.10	=	0.511 A ^{0.90} BSL ^{0.79} P ^{0.89} E6 ^{0.19}	.87	.18	43
	QM	=	0.284 A ^{0.90} BSL ^{0.87} P ^{0.92} E6 ^{0.19}	.87	.18	43

Table 2.--Results of regression analysis based on basin characteristics--Continued

Month and number of sites	Stream flow characterists	2-	Equation	R ²	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
July	Q.90	=	0.192 PE ^{1.37} BR ^{0.96} BSL ^{1.31}	.62	.33	88
(53)	Q.70	=	0.173 PE1.33 BR1.28 BSL1.06	.72	.27	69
	Q.50	=	0.296 PE1.33 BR1.18 BSL1.10	.75	.25	63
	Q.10	=	0.871 PE1.35 BR1.01 BSL1.20	.80	.21	51
	QM	=	0.485 PE1.33 BR1.03 BSL1.18	.78	.22	54
August	Q.90	=	0.105 PE1.43 BR0.65 BSL1.11	.57	.38	107
(53)	Q.70	=	0.0931 PE ^{1.39} BR ^{0.92} BSL ^{0.90}	.63	.34	92
	Q.50	=	0.0978 PE ^{1.33} BR ^{1.12} BSL ^{0.75}	.66	.31	82
	Q.10	=	0.209 PE1.26 BR1.07 BSL0.64	.72	.26	66
	QM	=	0.136 PE1.32 BR0.97 BSL0.77	.69	.28	72
September	Q.90	=	0.0420 PE1.46 BR0.90 BSL0.92	.60	.37	103
(53)	Q.70	=	0.0522 PE ^{1.42} BR ^{0.93} BSL ^{0.72}	.65	.33	88
	Q.50	=	0.0604 PE1.35 BR1.12 BSL0.67	.66	.30	78
	Q.10	=	0.202 PE ^{1.24} BR ^{0.98} BSL ^{0.70}	.74	.23	57
	QM	=	0.102 PE ^{1.33} BR ^{0.97} BSL ^{0.78}	.73	.25	63

level. The section beneath the reference level is that portion of the stream entrenchment in which the channel is actively, if not totally, sculpted by the normal process of water and sediment discharge."

The bankfull-channel section (also referred to as the main-channel or whole-channel section) was described by Riggs (1974, p. 53) as "...variously defined by breaks in bank slope, by the edges of the flood plain, or by the lower limits of permanent vegetation." On perennial streams, the upper extent of the bankfull-channel section corresponds to the bankfull stage at a narrow stream section described by Leopold and others (1964). For most sites in the study area, the bankfull width was only slightly larger than the active-channel width. The lower limit of permanent vegetation was most commonly the recognizable reference feature for active-channel width, whereas the prominent break in slope was most commonly used to define bankfull width.

In this study, the monthly streamflow characteristics and measured-channel widths were converted to logarithms, and multiple-regression techniques were used to derive estimating equations relating monthly streamflow to either active-channel or bankfull width:

$$\log Q = \log a + b \log W, \tag{3}$$

where

Q is a monthly streamflow characteristic as previously defined,

a is the regression constant,

b is the regression coefficient, and

 ${\tt W}$ is the significant independent variable, either active-channel width (W_{AC}) or bankfull width (W_{BF}).

The nonlinear form of equation 3, obtained by taking antilogarithms, is the following:

$$Q = a W^b. (4)$$

The final regression equations derived using channel widths and their coefficients of determination and standard errors are given in table 3. As with the basin-characteristics equations, the channel-width equations are generally more reliable for the higher-flow characteristics (Q.50, Q.10, and QM) than for the lower-flow characteristics (Q.90 and Q.70). Likewise, the channel-width equations are more reliable for the months of high runoff than for the months of low runoff and base flow. Ignoring measurement error, comparison of results in tables 2 and 3 indicates that the basin-characteristics equations and channel-width equations are about equally reliable for most flows for most months.

Concurrent-Measurement Method

The third method for estimating monthly streamflow characteristics at an ungaged site requires a series of discharge measurements at the site. The measured discharges at the ungaged site are correlated with concurrent discharges at some nearby, hydrologically similar gaged site, and the relation between the discharges at the two sites is used to transfer the desired long-term streamflow characteristic

Table 3.--Results of regression analysis based on channel width

 $[R^2$, coefficient of determination; Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; W_{AC} , active-channel width, in feet; QM, mean monthly discharge, in cubic feet per second; W_{BF} , bankfull width, in feet]

Month and number of sites	Stream flow charac terists	2-	Equation	R^2	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
October	Q.90	=	0.0521 W _{AC} ^{1.58}	0.59	0.35	96
(44)	Q.70	=	0.0774 WAC 1.56	.67	.29	75
	Q.50	=	0.116 W _{AC} ^{1.51}	.72	.25	63
	Q.10	=	0.383 W _{AC} 1.38	.73	.22	54
	QM	=	0.186 W _{AC} ^{1.44}	.78	.20	49
November	Q.90	=	0.0508 W _{AC} 1.60	.66	.30	78
(43)	Q.70	=	0.0875 W _{AC} 1.55	.69	.27	69
	Q.50	=	0.124 W _{AC} 1.52	.74	.24	60
	Q.10	=	$0.215 \text{ W}_{AC}^{1.55}$.78	.21	51
	QM	=	0.138 W _{AC} ^{1.53}	.77	.22	54
December	Q.90	=	0.0356 W _{AC} 1.66	.67	.31	82
(43)	Q.70	=	0.0695 W _{AC} ^{1.58}	.71	.26	66
	Q.50	=	0.0896 WAC 1.57	.74	.25	63
	Q.10	=	0.118 W _{AC} 1.68	.79	.23	57
	QM	=	0.0875 W _{AC} ^{1.63}	.76	.24	60

Table 3.--Results of regression analysis based on channel width--Continued

Month and number of sites	Stream flow characterists	2-	Equation	\mathbb{R}^2	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
January	Q.90	=	0.0270 WAC 1.71	.69	.31	82
(41)	Q.70	=	0.0398 W _{AC} 1.69	.75	.26	66
	Q.50	=	0.0557 W _{AC} ^{1.66}	.76	• 24	60
	Q.10	=	0.0735 W _{AC} ^{1.76}	.80	.23	57
	QM	=	0.0509 W _{AC} 1.73	.78	• 24	60
February	Q.90	=	0.0265 W _{AC} 1.73	.71	.29	75
(41)	Q.70	=	0.0389 WAC 1.71	.75	.26	66
	Q.50	=	0.0444 W _{AC} 1.72	.77	•25	63
	Q.10	=	0.0659 W _{AC} 1.80	.77	.26	66
	QM	=	0.0476 W _{AC} 1.75	.77	.26	66
March	Q.90	=	0.0320 W _{AC} 1.74	.74	.27	69
(42)	Q.70	=	0.0416 WAC 1.74	.75	.26	66
	Q.50	=	0.0529 WAC 1.74	.75	.26	66
	Q.10	=	0.0633 W _{AC} 1.85	.75	.28	72
	QM	=	0.0519 W _{AC} ^{1.79}	.75	•27	69

Table 3.--Results of regression analysis based on channel width--Continued

Month and number of sites	Stream flow characterists	2-	Equation	R ²	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
April	Q.90	=	0.0535 W _{AC} 1.75	.76	.26	66
(43)	Q.70	=	0.0695 W _{AC} 1.82	.79	.25	63
	Q.50	=	0.115 W _{AC} 1.81	.76	.27	69
	Q.10	=	0.271 W _{AC} 1.85	.76	.27	69
	QM	=	0.144 W _{AC} ^{1.83}	.77	.25	63
May	Q.90	=	0.0548 W _{BF} ^{1.95}	.80	.25	63
(46)	Q.70	=	0.0698 W _{BF} ^{2.03}	.86	.21	51
	Q.50	=	0.128 W _{BF} 1.96	.86	.20	49
	Q.10	=	0.392 W _{BF} 1.85	.88	.18	43
	QM	=	0.175 W _{BF} ^{1.91}	.87	.19	46
June	Q.90	=	0.265 W _{BF} ^{1.58}	.68	.27	69
(47)	Q.70	=	0.300 W _{BF} 1.67	.78	.22	55
	Q.50	=	0.423 W _{BF} 1.67	.81	.21	50
	Q.10	=	0.657 W _{BF} 1.72	.86	.17	41
	QM	=	0.445 W _{BF} 1.68	.84	.19	46

Table 3.--Results of regression analysis based on channel width--Continued

Month and number of sites	Stream flow charac terists	2-	Equation	R ²	Standard error (loga- rithm, base 10)	Stand- ard error (percent)
July	Q.90	=	0.162 W _{AC} ^{1.51}	.58	.34	92
(47)	Q.70	=	0.258 W _{AC} 1.51	.67	.29	75
	Q.50	=	0.372 W _{AC} 1.51	.72	.25	63
	Q.10	=	0.857 W _{AC} ^{1.51}	.80	•20	49
	QM	=	0.498 W _{AC} ^{1.49}	.77	•22	54
August	Q.90	=	0.0746 W _{AC} 1.54	•55	.37	103
(47)	Q.70	=	0.107 W _{AC} 1.54	.59	.34	92
	Q.50	=	0.163 W _{AC} 1.49	.60	•33	88
	Q.10	=	0.347 W _{AC} 1.44	.68	.26	66
	QM	=	0.191 W _{AC} ^{1.47}	.65	.29	75
September	Q.90	=	0.0545 WAC 1.57	•54	.39	111
(47)	Q.70		0.0741 WAC 1.56	.60	.34	92
		=	0.112 W _{AC} 1.51	.62	.32	85
	Q.10	=	0.278 W _{AC} 1.42	.72	.24	60
	QM	=	0.142 W _{AC} ^{1.48}	.69	.27	69

at the gaged site to the ungaged site. This estimation method, referred to in this report as the "concurrent-measurement method," has been used previously in Montana to estimate mean annual streamflow (Parrett and Hull, 1985; Parrett, 1985) and selected flows on a duration curve of monthly mean streamflow (Parrett and Hull, 1986). According to Searcy (1959, p. 17) and Riggs (1972, p. 15), the concurrent-measurement method generally provides more reliable estimates of low-flow characteristics than other methods not using discharge measurements.

The concurrent-measurement method investigated in this study requires 12 measurements (1 per month) at the ungaged site of interest. The measurements are paired with concurrent daily mean discharges obtained from a similar, nearby gaged site, and a straight line is plotted through the logarithms of the data points. The curve-fitting technique used (MOVE.1) is described by Hirsch (1982). The MOVE.1 technique is similar to an ordinary least-squares regression, except that ordinary regression minimizes the squared vertical deviations of the dependent variable from the regression line, whereas the MOVE.1 technique minimizes the areas of the right triangles formed by the horizontal and vertical deviations from the regression line (Hirsch and Gilroy, 1984, p. 707). The equation describing the ordinary least-squares regression line is the following:

$$y = \overline{y} + r \left(S_{y} / S_{x} \right) \left(x - \overline{x} \right), \tag{5}$$

where

y is the dependent variable,

y is the sample mean of the dependent variable,

r is the sample correlation coefficient between the dependent and independent variables,

 S_n is the sample variance of the dependent variable,

 $S_{\mathbf{v}}$ is the sample variance of the independent variable,

x is the independent variable, and

x is the sample mean of the independent variable.

The following equation describing the MOVE.1 best-fit line is identical to equation 5 except that r is not included:

$$y = \overline{y} + (S_y/S_x) (x - \overline{x}), \tag{6}$$

where all terms are as defined above. An example of an ordinary regression line and a MOVE.1 line fit to concurrent daily-mean discharges at two gaged sites is shown in figure 2. Although the two best-fit lines in figure 2 are similar, Stedinger and Thomas (1985) have shown that the MOVE.1 line is an unbiased estimator of low flows, whereas the ordinary regression line is a biased estimator of low flows. An alternative approach to the MOVE.1 or ordinary least-squares

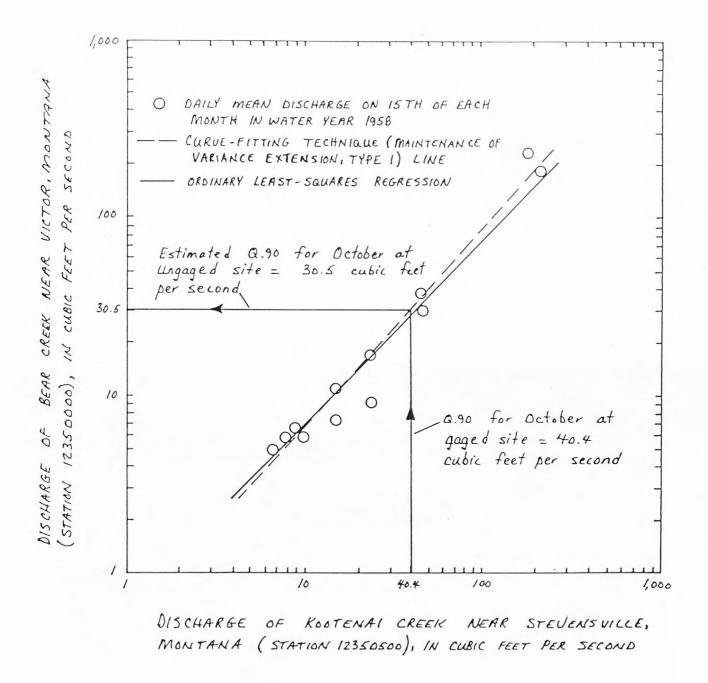


Figure 2.--Comparison between lines for the curve-fitting technique and ordinary least-squares regression lines.

regression would be a visual fit to the 12 data points. Although a visual fit would be subjective, it would allow the fitting of curves or multiple straight-line segments rather than a simple straight line.

To obtain an estimate of a particular monthly flow characteristic at the ungaged site, the value of the flow characteristic at the gaged site is located along the horizontal axis and projected to the MOVE.l line. The horizontal projection from the MOVE.l line to the vertical axis yields the estimate at the ungaged site

as shown in figure 2. As indicated by Searcy (1959, p. 20), the relation between concurrent high flows may be different from the relation between concurrent base flows so that a single straight line may not provide a good fit to the data. Riggs (1969) also showed that a difference in timing of runoff at two sites will result in a concurrent discharge plot that resembles a loop. Nevertheless, an examination of concurrent discharges from pairs of streamflow-gaging stations within the study area indicated that, in most instances, either the deviation from a single straight-line fit was not significant or the scatter about the line was great enough to mask any deviations. Accordingly, the reliability tests of the concurrent-measurement method are all based on a single MOVE.1 fit to the concurrent-measurement data. In applying the method at any particular site, however, the reader needs to be aware that a single straight line may not fit the data as well as two straight-line segments or that a timing-effects loop may exist. Using more complicated curve-fitting procedures in those instances will probably yield more accurate estimates than using the single MOVE.1 line.

To estimate the standard error of estimate of the concurrent-discharge method, the 20 pairs of streamflow-gaging stations listed in table 4 were tested. One

Table 4.--Streamflow-gaging stations used in the test of the curve-fitting technique

Station used as pseudo-ungaged site	Station used as index gaged site	Year of record used in test
06024500	06061500	1951
06030500	06033000	1947
06062500	06061500	1969
06073000	06078500	1951
06081500	06061500	1912
12301300	12302055	1980
12301999	12302055	1970
12303100	12302500	1968
12324100	12330000	1966
12346500	12343400	1967
12350000	12350500	1958
12351000	12350500	1958
12356500	12359000	1952
12360000	12359500	1953
12360500	12359500	1956
12361000	12359500	1956
12361500	12359500	1956
12365800	12366000	1979
12369200	12370000	1976
12390700	12389500	1983

¹Maintenance of Variance Extension, Type 1 (MOVE.1).

station of each pair was selected to be the test site (herein called the pseudo-ungaged site) for which estimates of monthly streamflow were required, and the other station served as the nearby, hydrologically similar index site. The stations were chosen such that the degree of similarity between the pseudo-ungaged and gaged sites was about the same as would be expected in actual practice. Thus, in some instances both sites were located in adjacent drainages and were very similar, and in other instances the sites were many miles apart and probably not so similar. One year from the concurrent period of record at each pair of stations was randomly selected, and the recorded daily mean discharge on the 15th of each month was used as the measured discharge at the pseudo-ungaged site and as the concurrent discharge at the gaged site. The MOVE.1 technique was then used to fit a line to the 12 data points, and the fitted line was used to estimate the monthly flow characteristics at the pseudo-ungaged site from the known monthly flow characteristics at the gaged site as described above.

The standard deviation of the differences (residuals) between the actual monthly flow characteristics at the 20 pseudo-ungaged sites and the estimated monthly flow characteristics from the MOVE.1 line was considered to be analogous to the standard error of estimate computed for the basin-characteristics method and the channel-width method. The resultant calculated "standard errors" for the monthly flow characteristics as determined from the 20 pairs of stations are presumed to be a reasonable approximation of the expected reliability of the concurrent-measurement method and are listed in table 5. Comparison of the standard errors in table 5 with the standard errors for the basin-characteristics method

Table 5.--Standard errors for concurrent-measurement method based on 12 measurements

[Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; QM, mean monthly discharge, in cubic feet per second]

	Stan	dard error, monthly	in percent flow charact		fied
Month	Q.90	Q.70	Q.50	Q.10	QM
Oct.	69	38	31	36	26
Nov.	46	28	21	31	21
Dec.	41	21	19	31	26
Jan.	38	21	21	28	26
Feb.	28	21	21	33	26
Mar.	26	23	26	38	28
Apr.	33	38	36	38	41
May	51	43	43	41	38
June	66	46	38	46	38
July	85	51	43	49	43
Aug.	92	66	54	38	46
Sept.	85	54	46	33	33

in table 2 and with the standard errors for the channel-width method in table 3 indicates that the concurrent-measurement method is substantially more reliable than the other methods for all months and nearly all monthly flow characteristics.

Using the concurrent-measurement method with 12 once-monthly measurements requires a large investment of time and money. Therefore, it is of some interest to investigate whether a program of fewer measurements might provide estimates of acceptable accuracy. Accordingly, the concurrent-measurement method was tested for the situation where only five once-monthly discharge measurements were available. For the same randomly selected year of record used in the 12-measurement test, the mid-monthly recorded discharge for the base-flow months November through March were used as data points for the 20 gage pairs, and the test described above was repeated. The five base-flow months were chosen for testing because many ungaged sites on the Flathead Indian Reservation had discharge measurements available for only those months. The computed standard errors for the concurrent-measurement method based on the five base-flow measurements are given in table 6. In this instance, the computed standard errors are substantially larger than the computed standard errors for the 12-measurement situation for most months when flow measurements were not available. The computed standard errors for the five-measurement situation are particularly large, substantially larger even than the standard

Table 6.--Standard errors for concurrent-measurement method based on five measurements

[Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; QM, mean monthly discharge, in cubic feet per second]

	Stan		in percent flow charact		fied	
Month	Q.90	Q.70	Q.50	Q.10	QM	
Oct.	82	49	38	75	49	
Nov. 1	57	36	36	60	38	
Dec.1	43	28	28	60	36	
Jan. 1	36	26	31	51	31	
Feb. 1	31	23	26	46	31	
Mar. 1	28	28	31	54	33	
Apr.	38	54	103	258	149	
May	179	326	471	772	471	
June	214	353	415	737	451	
July	120	129	160	339	214	
Aug.	99	92	96	116	92	
Sept.	92	69	63	69	60	

¹ Months when measurements were made.

errors for the basin-characteristics method or the channel-width method, April through July. Thus, the concurrent-measurement method based on fewer than 12 measurements may provide monthly flow estimates with an acceptable accuracy only for those months when measurements were made.

Weighted-Average Estimate

When different methods are available for estimating streamflow characteristics, it seems reasonable to assume that a weighted average of the individual estimates might provide a better answer than any of the individual estimates. When the individual estimates are independent, E.J. Gilroy (as cited by the U.S. Water Resources Council, 1981, p. 8-1) showed that the individual estimates could be weighted inversely proportional to their variances, and the resultant weighted average would have a smaller variance than any of the individual estimates.

To test whether the three estimating methods yield independent estimates, the cross-correlation coefficient between the residuals from the different methods was computed for 18 of the gaged sites used as pseudo-ungaged sites (table 4) in the concurrent-measurement method test. Two sites used in the concurrent-measurement test (stations 06030500 and 06081500) could not be used in this test because not all required basin-characteristics data were available. The equation used to compute the cross-correlation coefficient is the following:

$$r_{xy} = \underbrace{\sum_{i=1}^{N} x_i y_i - N \overline{x} \overline{y}}_{(N-1) S_x S_y}, \qquad (7)$$

where

 r_{xy} is the correlation coefficient between the residuals from method x and method y (ranges from -1.0 to 1.0),

N is the total number of sample residuals (18 in this computation),

 x_i and y_i are the *i-th* residuals from methods x and y,

 \overline{x} and \overline{y} are the mean values of the residuals from methods x and y, and

 $\mathbf{S}_{\mathbf{X}}$ and $\mathbf{S}_{\mathbf{y}}$ are the standard deviations of the residuals from methods \mathbf{x} and \mathbf{y}_{\bullet}

If the computed correlation coefficients between the residuals from any two estimating methods are zero or near-zero, the two methods may be considered to be independent. The results of the correlation-coefficient computations for all methods are listed in tables 7-9.

Table 7.--Correlation between residuals from basin-characteristics method and channel-width method

[Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; QM, mean monthly discharge, in cubic feet per second]

Flow		Corre	lation	coeff	icient	betwe	en res	iduals	for s	pecifi	ed mon	th
charac- teristic	Oct.	No v.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
Q.90	0.73	0.63	0.63	0.60	0.57	0.50	0.35	0.36	0.62	0.79	0.85	0.84
Q.70	.63	.58	•54	.52	.51	.46	.10	.18	.40	.68	.80	.79
Q.50	.57	.52	.51	.49	.43	.45	.18	.18	.31	.60	.77	.76
Q.10	.68	.45	.41	.42	.49	.52	.30	.19	.16	.45	.68	.65
QM	.54	.44	.45	.43	.45	.46	.17	.17	.23	.52	.74	.69

Table 8.--Correlation between residuals from basin-characteristics method and concurrent-measurement method

[Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; QM, mean monthly discharge, in cubic feet per second]

Flow	Correlation coefficient between residuals for spec									ecifie	ied month		
charac- teristic	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept	
Q.90	-0.34	-0.22	-0.11	-0.17	-0.16	-0.05	-0.18	-0.52	-0.46	-0.49	-0.42	-0.44	
Q.70	47	56	27	.02	08	15	38	41	29	42	42	44	
Q.50	 53	51	34	.01	03	12	36	33	24	32	38	44	
Q.10	46	47	34	21	37	26	56	09	.03	05	22	30	
QM	18	34	20	.19	09	09	36	20	16	06	24	28	

Table 9.--Correlation between residuals from channel-width method and concurrent-measurement method

[Q.xx, daily mean discharge exceeded xx percent of the time during the specified month, in cubic feet per second; QM, mean monthly discharge, in cubic feet per second]

Flow		Corre	lation	coeffi	icient	betwee	en resi	iduals	for s	pecifie	ed mont	h
charac- teristic	Oct.	Nov.	Dec.	Jan.	Feb.	Mar.	Apr.	May	June	July	Aug.	Sept.
Q.90	-0.45	-0.30	-0.21	-0.19	-0.05	0.18	-0.12	-0.42	-0.50	-0.47	-0.47	-0.51
Q.70	48	37	09	01	06	13	10	39	52	49	54	59
Q.50	43	19	03	.10	07	03	11	38	53	50	58	63
Q.10	46	44	26	.01	02	.08	23	17	23	57	65	55
QM	05	02	05	.34	.06	.11	03	30	46	- . 36	 50	49

As indicated by the results in table 7, the basin-characteristics method and the channel-width method yield monthly flow estimates that generally are not independent from each other. The results in tables 8 and 9 indicate that the concurrent-measurement method provides monthly flow estimates that are independent from either of the other two methods for some monthly flow characteristics for some months. For other flow characteristics and months, however, the concurrent-measurement method estimates are not independent from estimates made from the other two methods. Results in tables 8 and 9 also indicate that the correlation between the concurrent-measurement method and the other two methods commonly is negative. The negative correlations are an indication that the two methods being compared are providing estimates on either side of the true value, and that the errors of the individual estimates might be compensating when the estimates are combined.

If the individual estimates are not independent, the following equations (E.J. Gilroy, U.S. Geological Survey, written commun., 1987) can be used to weight the individual estimates so as to yield the weighted average estimate with the smallest variance:

$$Z = a1 \cdot X1 + a2 \cdot X2 + a3 \cdot X3,$$
 (8)

where

Z is the unbiased, weighted estimate of some flow characteristic,

al, a2, and a3 are weights which result in a minimum-variance, unbiased, linear combination of X1, X2, and X3, and

X1, X2, and X3 are estimates of the flow characteristic from three different methods.

Equations for the weights are as follows:

$$a1 = [C (SE_3^2 - S_{1.3}) - B (SE_3^2 - S_{2.3})]/(A C - B^2),$$
 (9)

$$a2 = [A (SE_3^2 - S_{2,3}) - B (SE_3^2 - S_{1,3})]/(A C - B^2),$$
 (10)

$$a3 = 1 - a1 - a2, \tag{11}$$

where

$$C = SE_2^2 + SE_3^2 - 2 S_{2,3},$$

 SE_1 , SE_2 , SE_3 are the standard errors of the three different estimating methods,

 $S_{1,2} = r_{1,2}$ ($SE_1 \cdot SE_2$) and is the covariance of methods 1 and 2,

 $S_{1,3} = r_{1,3}$ ($SE_1 \cdot SE_3$) and is the covariance of methods 1 and 3,

 $S_{2,3} = r_{2,3}$ (SE₂ • SE₃) and is the covariance of methods 2 and 3,

 $r_{i,j}$ is the cross-correlation coefficient between estimates from methods i and j.

$$A = SE_1^2 + SE_3^2 - 2S_{1,3}$$
, and

$$B = SE_3^2 + S_{1,2} - S_{1,3} - S_{2,3}$$

The estimated standard error of the weighted estimate, SE_Z , is determined as follows:

$$SE_z = [(a1 \cdot SE_1)^2 + (a2 \cdot SE_2)^2 + (1 - a1 - a2)^2 SE_3^2 + 2 a1 \cdot a2 \cdot S_{1,2} + 2 a1 (1 - a1 - a2) S_{1,3} + 2 a2 (1 - a1 - a2) S_{2,3}]^{0.5},$$
 (12)

where all terms are as previously defined.

If only two of the estimating methods are used, the following equations for computing weights and standard error are applicable:

$$Z = a1 \cdot X1 + a2 \cdot X2$$
, and (13)

$$SE_{z} = \sqrt{(SE_{1}^{2} SE_{2}^{2} - S_{1,2}^{2})/(SE_{1}^{2} + SE_{2}^{2} - 2 S_{1,2}^{2})}$$
 (14)

where

$$a1 = (SE_2^2 - S_{1,2})/(SE_1^2 + SE_2^2 - 2 S_{1,2}),$$
 and $a2 = (SE_1^2 - S_{1,2})/(SE_1^2 + SE_2^2 - 2 S_{1,2}).$

The above equations were used to calculate weights and standard errors for all combinations of the three estimating methods. For the basin-characteristics method and the channel-width method, the standard errors are based on the regression data from 54 gaged sites. The standard errors for the concurrent-measurement method are based on data from 20 gaged sites (table 4). The results, listed in table 13 at the back of the report, indicate that considerably more weight is given to the concurrent-measurement method estimates than either the basin-characteristics method or channel-width method estimates for all monthly streamflow characteristics for all months. Likewise, the weighted standard errors are substantially less when the concurrent-measurement estimates are included in the weighting procedure than when only estimates from the basin-characteristics method and channel-width method are used.

RELIABILITY AND LIMITATIONS OF ESTIMATING METHODS

Graphical comparisons of the standard errors for the individual methods of estimation and for the weighted-average estimates using all three methods are shown in figures 3-7. The standard errors, expressed in percent, range from 43 to 107 for the basin-characteristics method, from 41 to 111 for the channel-width method, from 19 to 92 for the concurrent-measurement method, and from 15 to 43 for the weighted-average estimates using all three methods. As indicated, the weighted-average estimates have the smallest standard errors for all monthly flow characteristics for all months. The weighted-average estimates thus are considered to be generally substantially more reliable than estimates from any of the three individual methods.

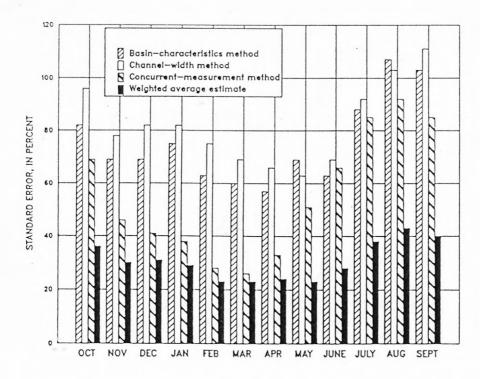


Figure 3.--Standard error for daily mean discharge exceeded 90 percent of the time based on different methods of estimation.

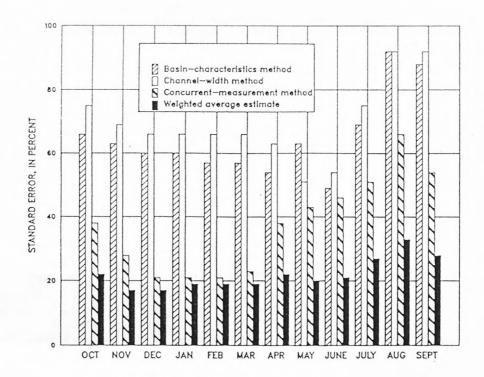


Figure 4.--Standard error for daily mean discharge exceeded 70 percent of the time based on different methods of estimation.

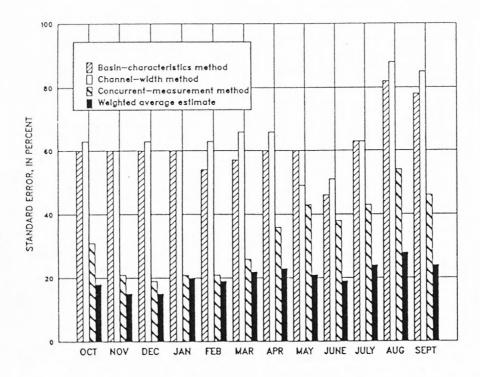


Figure 5.--Standard error for daily mean discharge exceeded 50 percent of the time based on different methods of estimation.

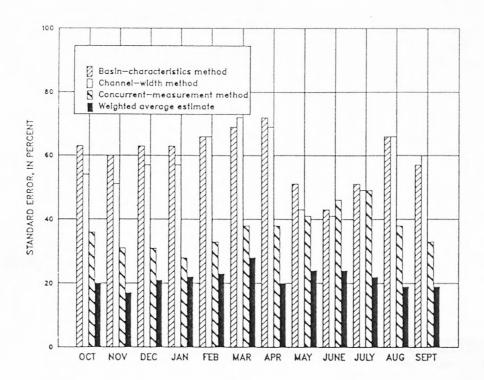


Figure 6.--Standard error for daily mean discharge exceeded 10 percent of the time based on different methods of estimation.

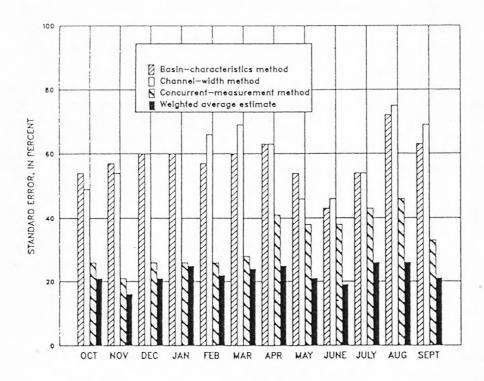


Figure 7.--Standard error for mean monthly discharge based on different methods of estimation.

Although figures 3-7 indicate the general reliability of the different estimating methods, the reader needs to be aware of certain limitations associated with the individual methods that may limit their applicability. Both the basin-characteristics method and the channel-width method, for example, are based on regression analyses, and the resultant regression equations may not be applicable beyond the range of variable values used to derive the equations. The ranges of basin and climatic characteristics and channel widths used in this study are given in table 10. Extrapolation beyond the values listed may yield erroneous estimates. Regression equations based on basin characteristics are also generally not applicable to streams that receive their water from springs or that lose substantial flows because of permeable streambeds or other localized geologic features. The equations also may not be applicable to stream sites with appreciable upstream lake storage or diversions.

Table 10.--Range of basin and climatic characteristics and channel widths used in the regression analyses

asin or width characteristic	Range of values
Prainage area (A), in square miles	3.59 - 838
Percentage of basin above 6,000 feet elevation, plus 1 (E6),	1.00 - 101
asin perimeter (PE), in miles	11.4 - 172
asin slope (BSL), dimensionless	0.19 - 0.64
ean annual precipitation (P), in inches	15 - 69
ean basin elevation (E), in thousands of feet	4.10 - 7.60
aximum basin relief (BR), in thousands of feet	2.04 - 7.09
ctive-channel width (WAC), in feet	12 - 172
ankfull width (W _{BF}), in feet	16 - 192

Regression equations based on channel width are probably more reliable than equations based on basin characteristics in such instances, because channel width is formed by the recent flow regime, no matter how anomalous the regime may be. Conversely, however, the channel-width method is generally not applicable where exposed bedrock occurs in either the streambed or banks, on braided or sand-channel streams, or on streams that have recently flooded or been altered by human activities.

In addition, accurate measurements of channel width require training and experience, and, even among experienced individuals, the variability in measured widths can be large. On the basis of a test in Wyoming, Wahl (1977) reported that the standard error in estimated flood discharge that could be attributed solely to measurement error might be as large as 30 percent. The total standard error of estimate for discharge using the channel—width method thus is comprised of both regression error and some unknown measurement error.

Because the concurrent-measurement method is based only on measured streamflow, the method is generally applicable where a suitable flow-measurement section can be found and where a suitable, nearby, concurrent streamflow-gaging station is available. Thus, the method can be used for sites where neither the

basin-characteristics method nor the channel-width method provides reliable estimates, but the reliability of the estimates made using the concurrent-measurement method is dependent on the degree of correlation between the measurement site and the index gaged site. If the concurrent measurements at the two sites are poorly correlated and show a large amount of scatter about the best-fit MOVE.1 line, the estimates made using the concurrent-measurement method may be unreliable. Extension of the MOVE.1 line beyond the range of discharge measurements may also result in errors in the long-term estimates. Additional limitations on the use of the concurrent-measurement method are the expense and time required to make the required 12 monthly flow measurements. Alternative measuring programs using fewer measurements can be devised, but the standard errors of the method may increase substantially.

APPLICATION OF ESTIMATING METHODS

The general procedures for using all methods to make estimates of monthly flow characteristics and for weighting the individual estimates are illustrated in the following examples. The examples are varied to illustrate typical applications of the various methods.

Example 1.

Estimates of the daily mean discharges exceeded 90 and 10 percent of the time (Q.90 and Q.10) during July are required for a stream located within the study area. The basin perimeter (PE), maximum basin relief (BR), and basin slope (BSL) were measured on suitable topographic maps and determined to be 13.1 mi, 5.22 thousands of feet, and 0.55, respectively. The site was visited and the active-channel width ($W_{\rm AC}$) was determined to be 16 ft. Using the applicable basin-characteristics equations from table 2, the required monthly streamflow characteristics are calculated as follows:

```
Q.90 = 0.192 PE<sup>1.37</sup> BR<sup>0.96</sup> BSL<sup>1.31</sup>

Q.90 = 0.192 (13.1)<sup>1.37</sup> (5.22)<sup>0.96</sup> (0.55)<sup>1.31</sup>

Q.90 = 14.5 ft<sup>3</sup>/s

Q.10 = 0.871 PE<sup>1.35</sup> BR<sup>1.01</sup> BSL<sup>1.20</sup>

Q.10 = 0.871 (13.1)<sup>1.35</sup> (5.22)<sup>1.01</sup> (0.55)<sup>1.20</sup>

Q.10 = 72.7 ft<sup>3</sup>/s
```

Similarly, the required monthly streamflow characteristics are calculated from the applicable channel-width equations in table 3:

$$Q.90 = 0.162 \text{ W}_{AC}^{1.51}$$

 $Q.90 = 0.162 \text{ (16)}^{1.51}$
 $Q.90 = 10.7 \text{ ft}^{3}/\text{s}$

$$Q.10 = 0.857 W_{AC}^{1.51}$$

 $Q.10 = 0.857 (16)^{1.51}$
 $Q.10 = 56.4 ft^3/s$

A program of once-monthly streamflow measurements was also instituted, and the measured flows were correlated with concurrent flows at a nearby, gaged correlating site as previously described. The MOVE.1 line through the plotted concurrent flows yielded the following estimates at the ungaged site:

$$Q.90 = 13.1 \text{ ft}^3/\text{s}$$

 $Q.10 = 53.0 \text{ ft}^3/\text{s}$

Weights for July were determined from table 13 for all three methods. Weighted estimates were calculated as follows:

Example 2.

Estimates of the daily mean discharge exceeded 50 percent of the time (Q.50) and the mean monthly discharge (QM) for June are required for a site in the study area. Insufficient time was available to use the concurrent-measurement method. The following basin and climatic characteristics were measured from topographic and precipitation maps:

```
Drainage area (A) = 22.6 \text{ mi}^2,
Basin slope (BSL) = 0.62,
Mean annual precipitation (P) = 40 \text{ in.}, and
Percentage of basin above 6,000 \text{ ft} elevation, plus 1 (E6) = 61.0.
```

On a site visit, the bankfull width $(W_{\rm RF})$ was measured as 35 ft.

Using the applicable basin-characteristics equations in table 2, the required monthly flow characteristics were calculated as follows:

```
Q.50 = 0.245 A<sup>0.91</sup> BSL<sup>0.95</sup> P<sup>0.95</sup> E6<sup>0.19</sup>
Q.50 = 0.245 (22.6)<sup>0.91</sup> (0.62)<sup>0.95</sup> (40)<sup>0.95</sup> (61.0)<sup>0.19</sup>
Q.50 = 193 ft<sup>3</sup>/s

QM = 0.284 A<sup>0.90</sup> BSL<sup>0.87</sup> P<sup>0.92</sup> E6<sup>0.19</sup>
QM = 0.284 (22.6)<sup>0.90</sup> (0.62)<sup>0.87</sup> (40)<sup>0.92</sup> (61.0)<sup>0.19</sup>
QM = 202 ft<sup>3</sup>/s
```

Using the appropriate channel-width equations in table 3, the monthly flow characteristics were calculated as follows:

$$Q.50 = 0.423 W_{BF}^{1.67}$$

 $Q.50 = 0.423 (35)^{1.67}$
 $Q.50 = 160 \text{ ft}^3/\text{s}$

$$QM = 0.445 W_{BF}^{1.68}$$

 $QM = 0.445 (35)^{1.68}$
 $QM = 175 \text{ ft}^{3}/\text{s}$

Using the appropriate weights from table 13 for June, the weighted estimates based on the basin-characteristics method and the channel-width method were calculated as follows:

Example 3.

Estimates of mean monthly discharge for January and February are required for a site in the study area. The following basin characteristics were measured on available topographic and precipitation maps:

```
Drainage area (A) = 21.0 \text{ mi}^2,
Maximum basin relief (BR) = 4.01 \text{ thousands of feet, and}
Basin slope (BSL) = 0.37.
```

On a site visit, the active-channel width (W_{AC}) was measured as 30 ft. During the site visit, the stream appeared to receive its water from a spring because streamflow was greater than at nearby, similar streams in the area. A concurrent-measurement program was instituted, and the 12 visits for measurements also confirmed that the site had greater flows than nearby, similar streams. Based on the concurrent-measurement program, estimates of the required monthly flow characteristics were as follows:

```
QM for January = 22.5 \text{ ft}^3/\text{s}
QM for February = 24.2 \text{ ft}^3/\text{s}
```

Using the appropriate basin-characteristics equations in table 2, mean monthly flow estimates were calculated as follows:

```
QM for January = 0.424 \text{ A}^{0.96} \text{ BR}^{0.88} \text{ BSL}^{1.30}

QM for January = 0.424 (21.0)^{0.96} (4.01)^{0.88} (0.37)^{1.30}

QM for January = 7.35 \text{ ft}^3/\text{s}
```

```
QM for February = 0.590 \text{ A}^{1.03} \text{ BR}^{0.63} \text{ BSL}^{1.53}
QM for February = 0.590 (21.0)^{1.03} (4.01)^{0.63} (0.37)^{1.53}
QM for February = 7.11 \text{ ft}^3/\text{s}
```

Using the appropriate channel-width equations in table 3, the estimates of mean monthly flow were calculated as follows:

```
QM for January = 0.0509 W<sub>AC</sub> 1.73

QM for January = 0.0509 (30)1.73

QM for January = 18.3 ft 3/s

QM for February = 0.0476 W<sub>AC</sub> 1.75

QM for February = 0.0476 (30)1.75

QM for February = 18.3 ft 3/s
```

Because the flow estimates made from the basin-characteristics equations were substantially smaller than the estimates made from the other two methods, and because the site appeared to receive its water from a spring during the site visits, the basin-characteristics estimates were considered to be erroneous. The final weighted estimates of mean monthly flow thus were made using only the concurrent-measurement method estimates and the channel-width method estimates from table 13 for January and February as follows:

```
QM for January = 18.3 (0.060) + 22.5 (0.940)

QM for January = 22.2 \text{ ft}^3/\text{s}

QM for February = 18.3 (0.136) + 24.2 (0.864)

QM for February = 23.4 \text{ ft}^3/\text{s}
```

The above examples were selected to illustrate how the various methods for estimating monthly streamflow characteristics could be used and combined in typical applications to provide the most reliable estimates. Considerable judgment is required to decide which methods may be appropriate or cost-and-time effective, however. Situations requiring the most accurate and reliable estimates will almost always require use of the concurrent-measurement method, but the additional time and cost required may be prohibitive.

SUMMARY AND CONCLUSIONS

Three methods for estimating mean monthly discharge and various points on the daily mean flow-duration curve for each month (daily mean discharges exceeded 90, 70, 50, and 10 percent of the time each month) were developed for western Montana. The first method was based on a multiple-regression analysis that related the streamflow characteristics to various basin and climatic variables. Several new basin characteristics were measured and tested to determine whether the regression equations might be improved. New characteristics that were found to be significant were basin perimeter, basin slope, and maximum basin relief. The estimating equations based on basin characteristics had standard errors ranging from 43 to 107. The standard error was smallest in the estimating equations for daily mean discharge that was exceeded 10 percent of the time (Q.10) for June and for mean monthly discharge for June. The standard error was largest in the estimating equations for daily mean discharge that was exceeded 90 percent of the time (Q.90) for August. Regression equations based on basin and climatic characteristics are generally not

applicable to streams that receive or lose water as a result of localized geologic features. They also may not be applicable to stream sites with appreciable upstream storage or diversions.

The second method for estimating monthly streamflow characteristics was based on a regression analysis relating the streamflow characteristics to channel width. The channel-width features used were active-channel width (W_{AC}) and bankfull width (W_{BF}). Most of the derived regression equations were based on active-channel width, but the equations for May and June were based on bankfull width. The standard errors for the estimating equations based on channel width ranged from 41 to 111. The standard error was smallest in the estimating equation for daily mean discharge that was exceeded 10 percent of the time (Q.10) during June, and was largest in the estimating equation for daily mean discharge that was exceeded 90 percent of the time (Q.90) in September. Regression equations based on channel width are generally not applicable where bedrock is exposed in the channel, on braided or sand-channel streams, or on streams that have recently been altered by floods or human activities. Proper application of the channel-width method also requires training and experience.

The third method for estimating monthly streamflow characteristics, termed the "concurrent-measurement method," required 12 once-monthly measurements of streamflow at the ungaged site of interest. The streamflow measurements at the ungaged site were correlated with concurrent discharges at a nearby gaged site using a MOVE.1 curve-fitting technique. The relation between flows at the two sites defined by the MOVE.1 curve then was used to compute the required monthly flow characteristics at the ungaged site from the monthly flow characteristics at the gaged site. Standard errors for the concurrent-measurement method were estimated by applying the method to 20 gaged sites and computing the standard deviation of the differences between the monthly flow characteristics determined from the estimation method and the monthly flow characteristics determined from the actual flow record. On this basis, the standard errors of the concurrent-measurement method ranged from 19 to 92 percent. The standard error was smallest in the estimate for the daily mean discharge that was exceeded 50 percent of the time (Q.50) during December, and was largest in the estimate for the daily mean discharge that was exceeded 90 percent of the time (Q.90) during August. Although the concurrentmeasurement method is generally substantially more accurate than either the basincharacteristics method or the channel-width method, it may yield unreliable results if there is poor correlation between the measurement site and the index gaged site. In addition, the monthly flow measurement method may be too expensive and timeconsuming for some applications.

A procedure for weighting individual estimates from any combination of the three different estimating methods to provide a minimum-variance-weighted-average estimate also was developed. The standard errors for the weighted estimates of monthly flow characteristics when all three methods were used ranged from 15 to 43 percent. The standard error was smallest for the weighted estimates for the daily mean discharge that was exceeded 50 percent of the time (Q.50) during November and December, and was largest for the daily mean discharge that was exceeded 90 percent of the time (Q.90) during August.

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SUPPLEMENTAL DATA

Table 11. -- Monthly streamflow characteristics for selected streamflow-gaging stations

[Monthly streamflow characteristic: Q.90, daily mean discharge exceeded 90 percent of the time; Q.70, daily mean discharge exceeded 70 percent of the time; Q.50, daily mean discharge exceeded 50 percent of the time; Q.10, daily mean monthly discharge; --, no data]

Monthly streamflow characteristic, in cubic feet per second, for specified month

		(October				No	ovember				De	cember		
Gaging station	Q.90	Q.70	Q.50	Q.10	QM	Q.90	Q.70	Q.50	Q.10	QM	Q.90	Q.70	Q.50	Q.10	QM
06024500 06029000 06030500 06033000 06061500	16.0 2.1 1.5 13.7 18.0	19.1 2.7 1.9 22.7 23.1	21.7 4.1 2.3 31.7 30.7	29.8 15.9 3.9 66.7 58.2	21.8 7.3 2.4 36.0 32.7	14.1 1.6 1.4 17.8 18.5	17.0 2.1 1.8 26.6 23.4	19.4 2.3 2.2 33.2 28.3	30.7 3.6 3.2 54.3 47.7	21.0 2.6 2.2 34.4 30.7	10.6 1.3 1.1 15.5 14.9	13.8 1.7 1.4 23.0 20.9	16.6 2.0 1.8 28.0 25.5	21.4 3.0 3.0 42.4 37.9	15.6 2.1 1.9 28.1 25.1
06062500 06073000 06078500 06081500	15.9 73.9 3.2	28.2 88.4 7.0	1.1 39.7 106 12.5	8.1 80.7 185 25.7	3.0 43.7 122 12.7	25.2 65.4 4.3	.7 30.4 79.3 8.8	1.3 37.6 93.3 12.0	5.6 82.7 148 21.5	2.2 44.8 102 12.1	.3 22.1 54.4 4.9	.7 27.5 66.5 7.0	1.1 33.7 75.1 9.3	4.4 67.3 124 17.3	1.8 39.2 83.0 10.2
12300500 12301300 12301999 12302055 12302500	8.8 69.0 7.3 94.0 6.7	10.8 87.1 8.8 112 10.5	13.9 103 10.3 125 18.8	80.3 163 14.3 182 60.7	27.3 114 .10.5 132 27.4	9.0 73.6 7.9 103 7.8	11.6 94.6 9.1 127 15.4	17.0 108 10.8 147 24.4	55.8 169 17.0 294 63.4	25.0 118 11.4 174 32.6	8.4 68.1 7.0 97.9 11.6	11.6 82.5 8.5 132 16.6	18.8 96.0 9.6 163 22.1	44.1 158 23.1 366 71.5	23.8 108 12.0 223 38.8
12303100 12324100 12330000 12332000 12335000	5.4 24.4 13.5 35.9	6.8 27.6 19.0 41.9	8.0 30.0 23.8 46.5 162	15.2 49.3 38.6 68.7 204	9.0 33.2 24.7 51.6 159	6.0 18.7 16.7 28.5	7.7 22.2 21.0 36.3 130	9.2 24.8 24.2 41.1	19.3 36.0 34.7 64.0 194	11.5 25.3 24.8 43.8 149	5.3 16.5 14.8 23.4 93.3	6.5 19.4 19.4 30.3	7.8 21.6 21.5 35.0 135	15.6 29.3 30.0 52.1 173	10.8 21.5 21.9 36.4 132
12338690 12339450 12343400 12346500 12347500	37.1 51.3 83.6 30.7 3.5	41.8 61.4 97.5 36.2 8.8	45.4 75.3 108 41.2 16.8	60.3 123 151 53.9 59.7	81.7 113 41.5 27.3	34.7 54.4 72.3 26.9 5.6	38.7 66.7 89.7 31.3 12.2	41.8 74.3 102 35.1 18.3	61.0 156 142 45.8 56.6	45.6 86.1 104 35.4 25.6	31.2 63.2 52.7 23.0 6.2	35.1 70.2 75.0 26.3 11.6	38.3 76.6 87.5 29.0 16.6	84.3 180 126 39.1 42.4	45.2 99.8 92.4 30.0 22.2
12350000 12350500 12351000 12352000 12353280	2.9 5.9 15.8 15.0 23.1	6.4 15.8 17.9 39.0 29.0	14.1 27.9 20.2 52.1 31.3	55.2 85.4 31.8 138 39.2	22.2 40.4 21.8 65.1 30.9	5.7 7.1 16.0 26.4 22.5	11.2 16.5 18.2 53.5 28.5	15.4 23.4 20.9 75.7 31.7	51.2 74.1 33.3 179 44.7	22.5 33.5 22.3 88.9 32.2	5.2 6.1 14.3 39.4 23.3	8.4 12.5 17.3 54.1 27.2	13.7 14.8 19.3 66.3 31.4	40.2 41.6 27.9 163 82.9	19.2 21.1 19.6 87.3 42.7
12354000 12356500 12357000 12359000 12359500		108 10.9 187 316 48.5	122 17.5 225 433 60.8	207 37.7 678 1090 179	143 21.9 347 623 93.7	99.2 8.7 132 253 41.3	118 14.4 177 323 50.9	153 18.1 280 444 83.2	519 27.5 702 880 168	235 18.4 355 529 93.9	86.9 7.3 121 214 36.9	105 10.8 166 288 49.3	143 12.9 246 369 69.4	431 26.4 643 832 145	214 16.1 333 452 81.6
12360000 12360500 12361000 12361500 12364000	8.7 6.7 22.5 10.1 5.7	12.0 8.0 31.5 20.9 10.8	18.9 10.9 55.2 39.9 14.6	59.4 48.2 199 147 62.0	29.2 21.6 89.3 63.4 38.3	8.7 5.7 25.5 12.1 11.8	15.0 8.0 51.5 34.6 14.8	31.5 23.1 74.5 54.0 18.7	69.9 56.7 190 116 36.2	34.9 26.1 96.5 60.3 20.9	10.4 4.2 23.1 14.0 12.5	18.3 9.0 37.8 27.8 15.6	27.5 19.8 58.7 41.8 19.5	57.3 43.1 158 103 41.6	36.3 23.4 80.4 54.1 22.5
12365000 12365800 12366000 12369200 12370000	67.1 25.1 22.1 37.2 350	89.8 31.5 50.6 44.4 416	110 37.0 68.6 51.2 507	142 48.9 127 85.9 860	110 36.2 70.7 56.4 560	71.4 18.5 26.8 35.0 358	95.8 26.4 57.1 45.2 444	113 32.2 70.3 52.1 529	180 69.8 122 91.0 846	121 37.3 73.7 58.4 582	65.4 11.5 40.1 34.7 350	81.8 21.3 51.0 41.6 425	94.6 26.5 59.1 46.4 484	167 58.7 123 102 862	107 34.3 71.8 59.5 567
12376000 12378000 12378500 12379500 12381000	15.8 30.1 31.7 28.0	22.6 38.3 35.7 33.1	27.9 44.7 42.3 41.0	53.3 57.7 59.7 78.4	31.9 43.5 43.1 47.8	11.5 25.2 28.9 32.0	22.2 28.5 36.6 50.0	29.0 330 45.5 56.3	50.8 47.6 71.4 70.0	31.4 34.0 47.7 54.1	7.4 13.8 26.6 24.7	18.3 20.7 30.0 43.7	22.9 24.9 35.4 49.5	35.2 37.1 51.5 55.5	22.4 24.3 37.1 45.0
12381400 12381500 12382000 12382500 12383500	21.0	24.0 9.1	26.0 9.7	68.4	34.4	19.0	23.0 7.7	26.0	38.8	27.6	6.0	13.0	17.0 7.3	22.0 8.7	15.5
12385000 12388500 12389500 12390700 12391550	4.9 7.2 140 43.3 76.8	5.9 7.9 173 51.4 91.8	6.8 8.4 189 56.9	8.9 14.3 241 68.0 135	7.6 9.6 191 59.7	7.6 146 41.1 80.2	8.1 168 47.9 95.6	11.1 185 52.8 116	14.7 247 122 267	10.6 194 69.7 157	6.6 128 41.4 79.0	7.5 157 48.9 109	8.0 179 57.8 182	12.7 305 228 551	10.0 203 113 282

Monthly streamflow characteristic, in cubic feet per second, for specified month January February March Gaging Q.50 Q.10 Q.90 Q.90 Q.70 Q.10 QM Q.90 0.70 Q.50 QM Q.70 0.50 Q.10 station QM 9.5 06024500 9.6 12.3 14.8 20.5 14.2 8.6 11.5 12.7 27.7 14.9 13.3 14.8 23.1 15.5 .5 2.0 1.9 1.8 06029000 1.4 1.6 1.4 1.5 1.8 2.3 1.7 1.8 2.6 1.8 .9 1.7 2.6 2.4 .9 06030500 1.4 1.7 1.0 1.4 1.8 1.7 1.5 2.0 4.1 2.1 27.4 24.2 06033000 14.7 21.6 38.8 26.1 16.2 24.1 29.1 41.3 29.7 30.5 36.1 73.3 06061500 13.5 19.8 23.2 32.6 22.2 15.1 20.8 24.0 37.0 25.0 17.4 24.2 29.2 50.1 32:0 06062500 3.0 1.0 1.6 06073000 20.1 23.9 28.2 52.1 32.6 19.9 23.6 27.8 49.5 31.1 22.2 26.4 30.8 51.2 36.4 06078500 46.6 56.7 64.6 89.6 66.6 47.3 56.3 65.0 87.4 66.3 46.8 55.3 61.5 92.0 67.7 6.0 7.2 18.3 10.1 7.9 30.8 06081500 4.0 14.1 9.3 12.8 19.1 20.2 20.6 32.7 12300500 8.0 15.0 52.0 27.9 13.1 19.3 28.9 45.0 61.7 75.4 90.6 144 101 64.5 80.2 96.4 150 103 94.7 114 203 138 12301300 75.3 26.0 22.2 13.4 68.8. 12301999 4.7 8.2 10.7 5.5 9.5 58.3 26.6 11.1 15.8 31.8 127 12302055 86.0 126 173 437 265 119 158 210 679 308 172 238 356 1120 529 9.9 15.4 19.1 33.1 20.0 8.6 12.3 14.7 50.5 24.6 11.7 15.8 22.0 55.6 29.8 12302500 8.3 19.7 12303100 4.6 4.6 17.8 10.1 4.8 6.2 19.1 18.3 16.4 18.6 20.8 26.0 20.4 16.6 18.5 20.0 24.5 19.8 16.6 24.4 19.8 12324100 17.4 19.7 25.5 19.2 13.9 17.4 19.2 24.6 19.0 17.3 18.8 24.6 18.9 12330000 12.3 15.1 20.0 26.0 27.4 12332000 30.1 45.5 31.5 19.2 31.9 45.0 32.0 22.4 28.9 33.3 49.0 82.7 108 125 -155 119 90.0 106 157 121 . 93.9 106 116 156 120 12335000 29.2 36.9 86.2 46.9 69.3 34.4 51.9 93.9 55.9 28.0 33.1 37.9 43.5 12338690 27.6 65.0 77.3 82.9 67.7 77.3 94.7 51.3 140 84.5 78.9 198 179 61.1 96.8 124 58.1 12339450 112 118 85.6 65.7 85.3 95.7 57.9 71.4 81.1 132 94.1 72.8 180 12343400 25.6 20.0 23.6 26.3 36.1 27.0 20.4 23.4 32.1 25.7 20.6 23.6 25.7 34.3 26.2 12346500 25.7 9.3 17.6 27.8 14.7 6.3 12.1 30.4 12347500 6.4 8.4 11.4 13.2 15.4 7.6 15.5 19.5 24.5 21.8 13.4 12350000 5.2 7.7 11.8 12.7 5.4 7.9 8.1 11.4 11.9 6.8 11.1 15.6 22.6 18.7 9.6 17.7 6.9 10.1 10.5 12350500 11.7 13.1 13.8 14.8 18.8 22.8 14.6 23.3 16.8 11.7 11.9 14.6 22.4 16.4 11.5 16.0 12351000 14.9 16.8 16.7 16.6 101 52.9 150 141 90.9 48.3 64.3 91.2 161 12352000 37.6 60.3 76.8 49.7 61.5 72.5 66.5 12353280 19.1 23.3 30.4 53.7 23.0 29.1 35.4 91.1 47.6 33.0 48.0 84.6 187 109 162 307 487 610 854 412 12354000 88.7 129 169 282 114 215 305 138 206 7.7 16.7 15.2 394 595 7.0 15.7 10.0 9.8 10.9 6.1 8.1 9.6 6.3 8.6 12356500 6.3 171 112 223 148 197 387 223 203 238 211 394 251 130 112 12357000 349 544 395 518 328 386 352 220 321 283 237 309 12359000 271 92.2 46.9 31.1 47.2 61.0 31.8 51.0 91.7 63.0 34.7 57.1 61.8 12359500 56.6 102 42.6 12.3 5.8 24.4 41.4 23.1 17.5 20.1 44.0 28.7 46.3 30.6 9.2 17.1 20.6 26.7 15.5 19.7 12360000 14.2 15.6 9.5 26.9 27.8 14.7 11.0 12.8 27.0 16.2 3.6 13.2 10.6 12360500 41.6 53.1 142 75.0 63.9 12361000 23.5 36.0 45.0 110 36.1 45.1 131 31.0 21.2 25.1 22.5 53.0 20.2 44.6 26.9 24.0 53.0 28.7 12361500 31.1 16.0 31.4 17.1 12364000 8.4 15.4 21.7 45.1 22.2 9.7 13.4 18.4 49.5 20.6 14.7 18.0 20.3 46.6 25.6 223 140 108 63.3 76.9 99.2 77.9 93.5 113 12365000 60.5 88.7 144 137 79.0 121 14.6 52.7 25.3 79.3 68.4 34.0 12365800 8.0 13.4 18.9 30.0 11.1 14.6 18.7 59.2 26.0 17.9 12366000 34.0 47.8 58.6 68.9 35.3 48.9 59.6 122 67.7 66.4 153 92.8 59.8 78.0 30.2 70.5 33.8 50.4 12369200 29.6 37.3 47.4 54.4 36.2 42.9 47.6 39.3 699 _ 390 453 681 499 311 376 436 491 335 431 500 935 588 12370000 70.7 43.4 12376000 20.1 25.5 39.0 13.7 13.9 19.4 14.4 10.3 15.1 16.5 28.4 17.4 9.9 12.3 13.9 19.4 11.1 12.5 12378000 12378500 19.9 22.8 26.6 43.6 27.6 20.5 23.1 26.2 43.5 28.1 22.8 28.1 32.4 39.9 32.1 .12379500 --------12381000 -1-12.0 17.0 30.0 18.0 10.7 14.0 15.0 17.0 14.3 14.0 15.0 17.0 20.3 17.6 12381400 10.0 12381500 ----------------12382000 12382500 --------4.2 4.6 5.0 5.7 5.0 4.9 5.3 6.0 6.9 5.9 4.3 4.5 4.6 5.1 4.6 12383500 12385000 7.4 5.6 6.6 3.3 7.4 4.0 7.1 7.9 34.3 13.0 33.0 15.2 6.1 8.6 6.6 12388500 148 303 202 262 535 334 123 179 213 161 196 389 242 155 12389500 134 41.7 63.0 12390700 53.7 78.2 215 128 43.4 68.2 104 312 154 117 165 389 215

154

489

246

152

201

128

12391550

73.0 126

170

420

261

100

257

510

300

Monthly streamflow characteristic, in cubic feet per second, for specified month

			Apri	1				May					June		
Gaging station	Q.90	Q.70		Q.10	QM	Q.90	Q.70	Q.50	Q.10	QM	Q.90	Q.70	Q.50	Q.10	QM
06024500 06029000 06030500 060333000 06061500	16.0 2.3 2.3 43.2 27.0	26.6 2.8 4.7 72.1 38.7	45.4 3.5 11.7 117 48.5	208 5.5 45.5 392 95.2	81.2 3.6 18.2 169 54.5	80.2 3.9 12.1 172 52.8	251 5.4 26.8 282 79.2	434 8.3 44.9 395 102	711 67.0 88.8 909 188	409 22.9 49.9 484 118	108 19.6 11.5 90.4 49.3	180 27.9 17.5 218 81.0	269 36.4 27.2 346 116	693 78.4 70.0 939 274	
06062500 06073000 06078500 06081500	1.9 26.8 67.1 8.7	4.8 42.9 93.1 13.1	9.4 71.2 121 22.0	46.5 328 505 43.7	17.5 129 225 23.8	22.2 106 351 10.8	44.8 201 749 31.1	71.0 318 1170 48.4	183 737 2400 142	88.7 375 1260 67.5	10.4 72.0 648 8.1	26.4 204 1080 30.4	54.9 345 1390 55.9	194 967 2590 239	79.9: 451 1543 98.7
12300500 12301300 12301999 12302055 12302500	28.9	88.3 205 70.4 553 61.9	223 302 130 818 85.8	611 760 656 2240 226	273 380 263 1110 116	130 437 50.0 646 110	208 607 164 1330 156	360 771 261 1450 215	754 1260 580 2690 399	412 808 296 1570 234	42.1 405 18.0 330 101	69.8 607 46.4 637 150	105 742 64.1 856 207	245 1230 168 1580 374	125 783 82:2 910 224
12303100 12324100 12330000 12332000 12335000	9.4 18.0 17.3 30.6 103	14.5 21.0 20.5 39.8 133	20.4 23.5 24.2 53.2 170	59.6 39.2 52.7 146 577	28.3 26.3 29.6 73.0 271	33.3 32.0 36.7 101 125	53.3 49.8 62.2 196 517	78.0 86.2 94.4 290 806	153 198 249 693 2070	86.0 101 121 347 970	41.6 112 75.6 229 458	67.9 174 126 357 756	91.4 222 178 463 1130	161 341 333 840 2370	96.3 221 191 509 1264
12338690 12339450 12343400 12346500 12347500	45.3 89.3 111 24.8 20.7	77.8 241 158 30.9 34.9	130 431 206 39.2 51.6	490 1240 469 98.3 165	211 579 248 51.3 75.3	250 562 355 64.0 97.5	404 847 568 120 157	592 1070 869 190 240	1250 1810 1740 .474 -457	668 1140 972 233 254	271 361 437 170 137	501 622 812 320 196	667 769 1080 397 257	1380 1250 1900 562 427	746 792 1160 386 269
12350000 12350500 12351000 12352000 12353280	17.1 25.0 16.6 96.9 70.5		62.4 63.2 31.3 314 231	171 184 85.5 663 551	81.1 83.6 42.1 372 271	85.7 87.2 46.9 412	151 148 76.0 650 328	229 213 106 809 453	468 454 277 1260 821	252 246 139 838 479	108 147 74.9 316 93.0	158 212 118 483 213	218 291 167 630 290	430 508 302 1030 633	243 309 180 652 339
12354000 12356500 12357000 12359000 12359500	9.0 285	24.0 600 771	1080 47.9 906 1290 253	2290 140 2940 3780 880	1240 63.9 1322 1760 394	929 90.8 1840 2520 504	1510 141 2880 4110 1040	1970 174 4130 6380 1480	3830 304 7440 12000 2830	2210 187 4350 6800 1590	468 44:0 1230 3400 487	833 64.3 2180 5680 1020	1320 113 3170 7730 1340	246 7580 12700	1542 127 3800 7950 1449
12360000 12360500 12361000 12361500 12364000	43.0 26.3 72.8 27.9 30.7	88.0 41.6 137 46.7 64.1	158 83.5 204 86.5 108	452 250 586 266 332	205 112 272 117 167	245 140 369 130 83.1	377 212 584 271 203	477 263 780 395 296	931 466 1440 784 610	545 284 837 429 341	121 76.3 369 264 38.0	247 141 538 391 84.1	358 202 713 504 114	696 395 1360 837 297	385 219 797 532 141
12365000 12365800 12366000 12369200 12370000	24.0 78.5 46.1	86.9	400 105 175 145 1330	1290 365 421 394 2690	586 152 216 186 1500	445 260 296 206 1580	847 387 416 311 2160	1170 481 517 397 2760	1990 796 848 739 4530	1213 507 540 435 2910	311 171 291 295 1890	577 321 474 387 2630	811 427 611 467 3240	1570 911 928 818 5120	880 499 621 520 3390
12376000 12378000 12378500 12379500 12381000	17.7 12.8 26.4 31.8	33.2 17.0 30.9 43.0	48.2 22.6 37.6 53.0	106 53.2 72.6 90.5	55.9 28.6 44.7 56.8	37.5 25.9 43.5 32.8	86.8 53.0 71.4 41.1	137 78.6 102 59.0	244 - 174 174 220 102	139 88.5 120 60.2	133 120 162 36.3	219 199 221 42.7	269 229 273 70.0	450 380 406 105	302 273 283 68.0
12381400 12381500 12382000 12382500 12383500	15.6	21.0	62.0	125	62.7	75.0 93.2 12.2 7.7	143 134 16.9 9.2	194 175 31.5 12.6	387 370 86.0 28.0	217 200 41.5 15.4	145 196 104 27.3 21.2	196 253 168 36.0 28.0	269 361 231 45.5 31.6	446 590 363 92.2 43.0	280 378 235 51.2 32.5
12385000 12388500 12389500 12390700 12391550	279 171	29.1 437 283 328	47.5 620 420 442	88.1 1400 924 991	49.9 769 499 531	11.2 52.1 629 415 562	16.7 71.9 1030 648 804	23.4 96.9 1360 825 1030	54.2 183 2430 1490 1680	28.4 110 1450 902 1080	27.4 34.0 496 228 424	36.8 52.2 788 377 659	44.3 81.0 1040 543 900	207 2070 1150	50.2 105 1160 615 1030

Monthly streamflow characteristic, in cubic feet per second, for specified month

			July					August	:			Sep	tember		
Gaging station	Q.90	Q.70	Q.50	Q.10	QM	Q.90	Q.70	Q.50	Q.10	QM	Q.90	Q.70	Q.50	Q.10	QM
06024500 06029000 06030500 06033000 06061500	32.8 13.4 2.5 13.4 20.9	42.8 19.2 3.8 33.0 34.8	55.8 24.9 5.5 60.3 50.5	102 40.2 17.1 211 125	61.6 25.1 7.8 93.9 61.1	17.4 17.9 1.3 7.6 13.9	21.4 20.7 1.8 13.5 21.2	24.8 28.0 2.2 19.8 28.3	39.2 40.3 4.6 56.8 56.5	26.2 27.7 2.5 26.5 31.2	14.9 7.9 1.4 7.5 15.0	17.2 15.2 1.7 15.2 20.7	19.7 18.0 2.0 21.7 26.5	30.9 26.7 3.3 47.9 56.1	21.0 17.4 2.2 26.7 30.5
06062500 06073000 06078500 06081500	.7 20.1 199 1.8	2.1 43.3 275 .11.3	5.2 84.0 373 25.2	34.7 287 935 82.1	12.7 123 483 39.0	12.6 114 1.6	20.0 138 7.4	1.2 31.3 156 15.5	6.1 95.6 241 39.8	2.3 47.5 167 17.9	.3 11.2 84.8 2.7	21.2 100 6.6	.9 27.9 115 11.0	5.3 62.9 169 30.0	2.0 34.6 124 13.2
12300500 12301300 12301999 12302055 12302500	15.8 133 6.9 126 27.0	22.3 208 15.8 208 42.8	30.8 281 21.8 276 56.4	69.5 563 57.1 586 136	36.3 319 29.6 313 71.6	7.9 69.4 4.7 81.4 9.5	9.8 104 10.0 113 14.6	14.1 127 12.2 145 20.4	27.3 195 17.8 213 37.7	16.4 130 11.4 144 21.6	7.3 71.4 4.9 90.1 7.2	9.0 92.5 8.0 112 10.1	10.2 112 9.1 126 13.0	20.7 158 13.0 162 35.8	12.2 115 9.2 130. 18:3
12303100 12324100 12330000 12332000 12335000	11.6 65.5 20.9 76.7 222	18.1 84.1 38.2 119 311	25.9 97.7 52.4 161 417	63.5 160 132 349 943	31.9 105 65.7 191 500	6.1 45.1 9.1 42.6 145	8.2 57.0 14.7 57.6 185	10.4 68.5 19.5 68.7 224	16.8 98.1 41.7 107 359	10.6 68.9 22.4 72.4 235	5.8 28.7 8.1 36.3	7.0 35.2 12.8 44.4 154	8.1 40.8 17.4 51.2	12.2 62.6 34.5 75.4 226	8.7 43.2 19.2 53.0 173
12338690 12339450 12343400 12346500 12347500	83.4 78.9 144 73.8 24.9	147 202	193 230 266 127 70.2	463 530 561 280 195	247 274 313 153 90.7	53.0 42.4 82.7 43.9 7.4	71.3 60.8 106 54.8 14.0	81.5 79.8 126 65.4 20.7	115 148 189 97.9 36.8	82.6 87.0 131 67.9 21.1	37.7 38.4 84.5 35.3 3.9	48.3 48.6 95.4 41.0 6.6	55.3 57.9 110 46.1 11.7	70.9 127 161 62.5 30.8	55.1 75.4 116 47.6 16.5
12350000 12350500 12351000 12352000 12353280	17.4 38.2 29.7 60.4 38.1	33.9 65.9 41.3 118 73.7	57.2 100 51.9 176 96.6	192 247 123 424 179	84.2 125 65.8 211 104	4.3 10.3 19.8 18.7 22.0	7.3 14.7 23.6 31.9 34.2	10.3 19.1 29.1 47.7 41.3	24.8 42.7 42.7 84.8 59.9	12.7 24.1 29.5 49.7 41.5	2.6 6.6 16.4 14.0 21.1	3.9 9.2 19.5 23.8 30.2	6.0 13.2 21.7 40.0 34.4	23.0 33.6 28.9 76.0 44.5	10.7 18.0 21.9 42.9 33.4
12354000 12356500 12357000 12359000 12359500	178 15.8 363 969 140	242 21.6 535 1450 227	313 28.4 747 2080 338	.746 88.2 2090 5290 960	400 39.1 1041 2720 470	106 10.2 194 431 66.6	132 12.8 254 546 92.5	155 15.6 306 648 115	240 25.1 495 1140 178	165 16.1 326 730	92.4 9.3 170 305 52.5	110 10.8 201 357 61.5	127 12.0 227 412 68.3	177 19.7 339 825 104	132 13.4 244 500 75.7
12360000 12360500 12361000 12361500 12364000	31.3 24.4 75.0 55.4 8.3	44.7 35.7 107 88.4 23.9	65.0 56.9 151 149 40.1	169 130 382 373 126	86.7 68.9 197 193 55.6	14.1 12.3 35.4 22.1 3.9	18.2 15.2 44.5 27.9 7.1	22.5 17.8 54.2 34.7 9.6	35.2 27.9 91.2 71.0 34.2	23.1 18.9 60.0 41.1 15.7	10.3 8.1 27.0 13.5 2.4	12.3 9.3 32.7 17.7 6.3	13.9 10.7 38.3 23.2 8.8	25.0 16.5 104 63.7 29.1	17.9 11.4 59.2 33.2 12.5
12365000 12365800 12366000 12369200 12370000	127 64.7 100 103 734	230 99.2 160 185	333 133 230 263 1470	665 338 491 523 2890	370 173 270 296 1660	68.6 34.6 39.5 53.0 416	119 51.9 74.2 74.7 545	166 69.8 105 98.8 663	278 114 184 182 1060	170 70.5 107 109 705	64.9 27.5 28.4 38.8 353	90.6 36.9 62.8 56.5 435	123 46.4 82.3 69.2 513	186 72.7 133 124 769	124 47.6 83.7 78.7 550
12376000 12378000 12378500 12379500 12381000	31.5 78.6 73.5 30.4	173	103 201 214 50.2	243 344 370 71.2	122 204 222 50.7	22.7 34.0 25.6 21.8	31.5 69.1 41.0 26.3	36.3 83.4 101 32.4	66.0 155 172 55.8	41.2 87.7 93.8 34.6	19.2 27.0 24.8 18.8	23.2 43.0 31.8 20.8	27.6 52.9 63.0 23.3	44.3 85.1 118 37.9	31.6 55.3 64.2 25.5
12381400 12381500 12382000 12382500 12383500	70.0 112 37.8 7.7 14.7	89.0 139 49.0 12.2 19.4	104 156 72.8 16.7 23.3	183 273 145 32.0 41.2	124 190 89.2 20.4 25.9	34.0 55.9 14.1 3.5 9.8	45.0 80.0 21.1 4.9 12.9	57.5 109 26.5 6.2 15.1	88.1 140 37.9 12.3 20.9	60.5 103 26.0 6.8 15.4	24.0 42.0 11.0 2.6 9.0	30.0 52.5 17.0 4.4 10.6	35.0 61.9 21.9 6.8 11.9	54.0 98.0 31.9 12.5 15.7	37.2 65.0 21.1 6.7 11.7
12385000 12388500 12389500 12390700 12391550	11.0 12.1 246 100 145	15.8 17.2 337 126 244	22.3 21.2 405 149 323	47.5 48.9 696 276 740	26.9 29.3 443 171 396	5.9 5.5 171 69.1 84.7	7.8 9.0 225 80.3	9.8 11.3 256 88.1 149	14.5 16.2 338 112 237	9.8 10.8 255 88.4 155	4.8 4.6 147 52.0 78.8	6.0 6.1 187 61.7	7.6 8.2 210 67.6	10.4 ⁻ 13.6 269 78.7 160	7.6 8.4 208 66.0

Table 12.--Basin and climatic characteristics and channel widths for selected streamflow-gaging stations

[Basin/climatic characteristic: A, drainage area, in square miles; E6, percentage of basin above 6,000 feet elevation, plus 1; PE, basin perimeter, in miles; BSL, basin slope, dimensionless; L, main-channel length in miles; P, mean annual precipitation, in inches; E, mean basin elevation, in thousands of feet; BR, maximum basin relief, in thousands of feet. Channel width: $W_{\rm AC}$, active-channel width, in feet; $W_{\rm BF}$, bankfull width, in feet; --, no data]

Gaging station	. A	E6	PE	BSL	L	P	Е	BR	W_{AC}	W_{BF}
06024500	71.4	101	44.6	0.25	19.3	30	7.11	2.04	39	48
06029000	30.8	98.2	29.9	.22	8.80	21	7.33	2.91		
06062500	32.7	87.2	27.9	.30	8.80	24	6.58	3.41	16	25
06073000	123	77.0	66.3	.41	21.8	37	6.23	4.66	54	83
06078500	258	62.0	84.5	.35	27.4	42	6.15	4.62		
12300500	110	2.00	47.1	.19	18.7	28	4.55	3.27	22	29
12301300	440	6.00	126	.24	40.2	32	4.17	4.99	48	58
12301999	216	1.00	96.2	.23	29.4	27	4.10	4.03	40	47
12302055	838	1.00	169	.24	66.4	32	4.10	5.13	111	130
12302500	23.6	32.0	25.0	.56	9.40	67	5.26	5.97		
12303100	11.1	44.9	16.8	.52	5.90	67	5.24	4.83	17	2.4
12324100	39.5	94.0	33.9	.37	12.7	35	7.60	4.31	22	28
12330000	71.3	84.0	41.1	.33	13.5	31	6.98	4.77	28	32
12332000	123	90.0	61.3	.29	20.2	35 ·	7.18	5.08	56	71
12335000	481	48.0	114	.34	47.7	15	5.89	5.10	100	120
12338690	140	55.0	72.3	.36	26.0	35	5.91	4.05	52	68
12339450	345	27.0	112	.25	35.7	37	5.28	5.27	90	110
12343400	381	62.0	116	.40	35.3	32	6.45	5.10	70	90
12346500	87.8	83.0	52.4	.43	12.9	36	6.80	4.61	34	44
12347500	26.4	68.0	30.5	.62	12.3	64	6.73	4.96	30	38
12350000	26.8	69.0	27.2	.54	11.8	63	6.43	5.01	41	47
12350500	28.9	66.0	25.6	.62	10.5	64	6.35	5.69	38	46
12351000	73.2	71.0	44.8	.37	16.7	32	6.57	4.41	20	28
12352000	250	34.3	93.6	.42	30.9	52	5.43	5.82	51	60
12353280	170	25.0	66.6	.35	25.1	38	4.92	4.96	48	60
12354000	303	2.00	99.3	.43	37.1	52	4.52	4.67	130	136
12356500	20.7	33.2	23.8	.27	8.10	47	5.77	4.13	25	36
12357000	510	47.0	126	.44	60.0	52	5.90	4.83	172	192
12359500	184	56.5	75.3	.40	29.8	56	6.00	5.17	65	105
12360000	47	58.0	36.3	.48	15.4	53	5.30	4.16	41	59

Table 12.--Basin and climatic characteristics and channel widths for selected streamflow-gaging stations--Continued

Gaging station	A	E6	PE	BSL	L	Р	Е	BR	W _{AC}	W_{BF}
12360500	22.4	31.4	25.2	.49	8.90	56	5.49	4.16	32	44
12361000	71.3	39.0	50.3	. 47	13.0	35	5.51	4.03	63	78
12361500	27.0	43.0	24.8	•50	9.00	67	5.43	3.96	40	60
12364000	183	7.00	71.3	. 24	34.9	28	4.91	2.95	47	67
12365000	524	4.00	158	.25	50.5	31	4.32	4.39	70	85
12365800	78.0	30.0	54.8	.35	26.3	51	5.20	4.44	44	56
12366000	170	12.0	97.4	.23	36.9	37	4.17	4.48	64	82
12369200	73.3	40.0	49.8	.33	19.0	54	5.83	5.36	72	84
12370000	671	27.0	172	.33	84.5	23	5.02	6.33	165	185
12376000	50.6	17.7	37.9	.38	16.9	45	4.85	5.60	30	42
12378000	74.8	31.0	40.9	.33	14.8	48	4.84	6.63		
12378500	22.6	15.4	26.8	.64	12.7	66	6.12	6.76	35	40
12379500	67.1	22.0	39.0	.40	17.3	45	4.75	7.09	28	33
12381000	15.9	5.74	21.4	.46	10.9	38	5.65	3.69	22	25
12381400	58.3	32.5	42.3	.35	18.6	39	6.06	4.23	32	36
12381500	74.2	37.2	49.4	.38	18.7	39	5.97	4.26	35	42
12382000	20.0	14.3	23.1	.41	9.10	64	6.15	4.08	32	38
12382500	3.59	3.76	11.6	.48	5.46	69	6.63	4.88	17	23
12383500	6.90	5.57	12.8	.55	5.92	40	6.32	4.29	12	16
12385000	6.51	5.44	11.4	.58	5.27	40	6.41	4.66	18	22
12388500	26.3	37.0	21.8	.29	9.35	33	5.56	4.21		
12389500	642	6.00	145	.36	48.7	41	4.71	5.03	95	
12390700	182	4.00	72.2	.47	21.5	54	4.41	4.39	44	67
12391550	139	10.0	58.3	.51	26.4	65	4.47	6.53	64	70

Table 13.--Weights and standard errors for various combinations of methods of estimation

[Q.xx, daily mean discharge exceeded xx percent of time during specified month, in cubic feet per second; QM, mean monthly discharge, in cubic feet per second; log, logarithm, base 10; pct, percent]

			s for spec		
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM
	00	TOBER			
Basin-characteristics method	0.197	0.203	0.219	0.153	0.170
Channel-width method	.255	.159	.121	.233	.114
Concurrent-measurement method	.548	.638	.660	.614	.716
Weighted standard error (log)	.153	.094	.076	.087	.088
Weighted standard error (pct)	36	22	18	20	21
Basin-characteristics method	.719	.646	.547	.305	.397
Channel-width method	.281	.354	.453	.695	.603
Weighted standard error (log)	.302	.246	.217	.212	.183
Weighted standard error (pct)	79	61	53	52	44
.,					
Concurrent-measurement method	.551	.659	.689	.667	.762
Basin-characteristics method	.449	.341	.311	.333	.238
Weighted standard error (log)	.166	.101	.081	.096	.090
Weighted standard error (pct)	40	24	19	22	21
Concurrent-measurement method	.588	.690	.712	.628	.757
Channel-width method	.412	.310	.288	.372	.243
Weighted standard error (log)	.159	.104	.090	.092	.094
Weighted standard error (pct)	38	24	21	21	22
	NOVEM	IBER .			
Basin-characteristics method	0.181	0.235	0.209	0.176	0.190
Channel-width method	.203	.063	.015	.202	.023
Concurrent-measurement weight	.616	.702	.776	.622	.787
Weighted standard error (log)	.129	.073	.063	.075	.071
Weighted standard error (pct)	30	17	15	17	16
Basin-characteristics method	.641	•591	.500	.380	.460
Channel-width method	.359	.409	.500	.620	.540
Weighted standard error (log)	.255	.230	.209	.190	.191
Weighted standard error (pct)	64	57	51	46	46

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

				ified mont haracteris	
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM
	NOVEMBER	Continue	d		
Concurrent-measurement method Basin-characteristics method Weighted standard error (log) Weighted standard error (pct)	.640 .360 .138	.718 .282 .075	.782 .218 .063	.696 .304 .085	.798 .202 .071
Concurrent-measurement method Channel-width method Weighted standard error (log) Weighted standard error (pct)	.667 .333 .135	.763 .237 .090	.835 .165 .078	.661 .339 .085	.852 .148 .083
	DECEM	BER			
Basin-characteristics method Channel-width method Concurrent-measurement weight Weighted standard error (log) Weighted standard error (pct)	0.158 .167 .675 .130	0.163 .026 .811 .075	0.168 .000 .832 .065	0.182 .161 .657 .089	0.197 .037 .766 .091
Basin-characteristics method Channel-width method Weighted standard error (log) Weighted standard error (pct)	.683 .317 .258	.586 .414 .218	.542 .458 .213	.430 .570 .201	.500 .500 .204
Concurrent-measurement method Basin-characteristics method Weighted standard error (log) Weighted standard error (pct)	.697 .303 .136	.820 .180 .075	.832 .168 .065	.725 .275 .096	.783 .217 .091
Concurrent-measurement method Channel-width method Weighted standard error (log) Weighted standard error (pct)	.728 .272 .134	.872 .128 .082	.900 .100 .076	.712 .288 .099	.839 .161 .102

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

				ified mont haracteris	
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM
	JANUA	RY			
asin-characteristics method	0.159	0.091	0.099	0.189	0.115
Channel-width method	.139	.060	.042	.086	.009
Concurrent-measurement method	.701	.849	.859	.726	.875
Weighted standard error (log)	.124	.083	.084	.096	.106
eighted standard error (pct)	29	19	20	22	25
asin-characteristics method	.583	.583	.500	.428	.500
Channel-width method	.417	.417	.500	.572	.500
Weighted standard error (log)	.267	.217	.207	.201	.203
reighted standard error (pct)	68	53	51	49	49
Concurrent-measurement method	.733	.872	.879	.769	.88
asin-characteristics method	.267	.128	.121	.231	.119
Weighted standard error (log)	.129	.084	.085	.098	.106
reighted standard error (pct)	30	19	20	23	25
Concurrent-measurement method	.751	.891	.903	.788	.940
Channel-width method	.249	.109	.097	.212	.060
Weighted standard error (log)	.130	.085	.087	.107	.109
leighted standard error (pct)	31	20	20	25	26
	FEBRU	ARY			
Basin-characteristics method	0.196	0.114	0.101	0.276	0.183
Channel-width method	.036	.058	.085	.024	.04
Concurrent-measurement method	.768	.828	.814	.700	.77
Veighted standard error (log)	.100	.080	.080	.100	.09
Weighted standard error (pct)	23	19	19	23	22
Basin-characteristics method	.669	.624	.611	.500	.61
Channel-width method	.331	.376	.389	.500	.39
Weighted standard error (log)	.236	.211	.197	.224	.20
Weighted standard error (pct)	58	52	48	55	50
Concurrent-measurement method	.778	.848	.849	.710	.79
Basin-characteristics method	.222	.152	.151	.290	.20
Weighted standard error (log)	.101	.081	.082	.100	.09
erbited beamant of the (B)			19	23	22

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

				ified mont	
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM
	FEBRUARYC	ontinued			
Concurrent-measurement method	.844	.881	.869	.769	.864
Channel-width method	.156	.119	.131	.231	.136
Weighted standard error (log)	.109	.084	.083	.122	.103
Weighted standard error (pct)	26	19	19	29	24
	MARC	H			
	0.106	0.120	0 171	0.20/	0.000
Basin-characteristics method	0.186	0.138 .086	0.171 .069	0.304 .000	0.208
Channel-width method	.814	.776	.760	.696	.767
Concurrent-measurement method Veighted standard error (log)	.098	.084	.093	.120	.103
Weighted standard error (pct)	23	19	22	28	24
reignied scandard error (pec)	23	17		20	2 '
Basin-characteristics method	.616	.612	.610	.538	.608
Channel-width method	.384	.388	.390	.462	.392
Weighted standard error (log)	.219	.207	.207	.240	.216
Weighted standard error (pct)	54	51	50	60	53
Concurrent-measurement method	.814	.807	.787	.696	.778
Basin-characteristics method	.186	.193	.213	.304	.222
Weighted standard error (log)	.098	.086	.094	.120	.103
Weighted standard error (pct)	23	20	22	28	24
Concurrent-measurement method	.909	.841	.841	.773	.862
Channel-width method	.091	.159	.159	.227	.138
Weighted standard error (log)	.107	.089	.100	.144	.113
Weighted standard error (pct)	25	21	23	34	27
	ADDT	T			
	APRI	<u>L</u>			
Basin-characteristics method	0.222	0.315	0.271	0.256	0.286
Channel-width method	.132	.147	.131	.131	.159
Concurrent-measurement method	.645	.538	.597	.613	.555
Weighted standard error (log)	.104	.095	.097	.087	.105
Weighted standard error (pct)	24	22	23	20	25

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

	Weights for specified month and monthly flow characteristic								
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM				
	APRILCo	ntinued							
Basin-characteristics method	.593	.571	.549	.474	.500				
Channel-width method	.407	.429	.451	.526	.500				
Weighted standard error (log)	.200	.174	.192	.222	.191				
Weighted standard error (pct)	49	42	46	55	46				
Concurrent-measurement method	.698	.613	.666	.671	.638				
Basin-characteristics method	.302	.387	.334	.329	.362				
Weighted standard error (log)	.109	.103	.103	.095	.114				
Weighted standard error (pct)	26	24	24	22	27				
Concurrent-measurement method	.750	.692	.729	.691	.671				
Channel-width method	.250	.308	.271	.309	.329				
Weighted standard error (log)	.117	.128	.123	.118	.135				
Weighted standard error (pct)	27	30	29	28	32				
	MAY								
Danier share to sinting mother	0.257	0.220	0.207	0.209	0.194				
Basin-characteristics method Channel-width method	.235	.290	.320	.346	.316				
Concurrent-measurement method	.508	.490	.474	.445	.490				
	.097	.086	.089	.101	.090				
Weighted standard error (log) Weighted standard error (pct)	23	20	21	24	21				
Basin-characteristics method	.440	.395	.390	.406	.413				
Channel-width method	.560	.605	.610	.594	.587				
Weighted standard error (log)	.214	.174	.167	.149	.155				
Weighted standard error (pct)	52	42	40	35	37				
Concurrent-measurement method	.582	.614	.606	.596	.629				
Basin-characteristics method	.418	.386	.394	.404	.371				
Weighted standard error (log)	.115	.113	.118	.126	.116				
Weighted standard error (pct)	27	26	28	30	27				
Concurrent-measurement method	.561	•555	.538	.524	.565				
Channel-width method	.439	.445	.462	.476	.435				
Weighted standard error (log)	.122	.106	.105	.112	.102				
Weighted standard error (pct)	29	25	24	26	24				

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

				ified mont haracteris	
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM
	JUN	E			
Basin-characteristics method	0.274	0.203	0.179	0.241	0.192
Channel-width method	.260	.315	.305	.402	.331
Concurrent-measurement method	.465	.482	.516	.357	.478
Weighted standard error (log)	.119	.091	.080	.101	.081
Weighted standard error (pct)	28	21	19	24	19
Basin-characteristics method	.601	.579	.572	.466	.535
Channel-width method	.399	.421	.428	•534	.465
Weighted standard error (log)	.233	.175	.161	.133	.145
Weighted standard error (pct)	58	42	38	31	34
Concurrent-measurement method	.487	.520	.569	.472	.551
Basin-characteristics method	.513	.480	.431	.528	.449
Weighted standard error (log)	.132	.116	.107	.133	.110
Weighted standard error (pct)	31	27	25	31	26
Concurrent-measurement method	.513	.548	.588	.455	.558
Channel-width method	.487	.452	.412	• 545	.442
Veighted standard error (log)	.132	.100	.088	.111	.089
Weighted standard error (pct)	31	23	20	26	21
	JUL	<u>.Y</u>			
Basin-characteristics method	0.294	0.196	0.116	0.006	0.088
Channel-width method	.210	.236	.304	.496	.367
Concurrent-measurement method	.496	.568	.580	.498	.545
Weighted standard error (log)	.158	.115	.101	.093	.111
Weighted standard error (pct)	38	27	24	22	26
Basin-characteristics method	.571	.611	.500	.456	.500
Channel-width method	.429	.389	.500	.544	.500
Weighted standard error (log)	.317	.255	.224	.174	.192
Weighted standard error (pct)	84	64	55	42	46

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

				ified mont	
Combinations of methods of estimation	Q.90	Q.70	Q.50	Q.10	QM
	JULYCon	tinued			
Concurrent-measurement method	.510	.587	.622	.523	.594
Basin-characteristics method	.490	.413	.378	.477	.406
Weighted standard error (log)	.164	.127	.121	.141	.135
Weighted standard error (pct)	39	30	28	33	32
Concurrent-measurement method	.521	.606	.608	.500	.573
Channel-width method	.479	.394	.392	.500	.427
Weighted standard error (log)	.170	.122	.104	.093	.112
Weighted standard error (pct)	41	29	24	22	26
	AUGU	ST			
Basin-characteristics method	0.137	0.079	0.020	0.000	0.000
Channel-width method	.342	.344	.360	.358	.364
Concurrent-measurement method	.520	.577	.621	.642	.636
Weighted standard error (log)	.180	.140	.119	.081	.113
Weighted standard error (pct)	43	33	28	19	26
Basin-characteristics method	.411	.500	.635	.500	.567
Channel-width method	.589	.500	.365	.500	.433
Weighted standard error (log)	.360	.323	.300	.238	.265
Weighted standard error (pct)	99	86	78	59	67
Concurrent-measurement method	.539	•593	.621	.688	.651
Basin-characteristics method	.461	.407	.379	.312	.349
Weighted standard error (log)	.193	.158	.142	.122	.138
Weighted standard error (pct)	47	38	34	29	33
Concurrent-measurement method	.529	.586	.625	.642	.636
Channel-width method	.471	.414	.375	.358	.364
Weighted standard error (log)	.182	.141	.119	.081	.113
Weighted standard error (pct)	44	33	28	19	26

Table 13.--Weights and standard errors for various combinations of methods of estimation--Continued

Combinations of methods of estimation	Weights for specified month and monthly flow characteristic				
	Q.90	Q.70	Q.50	Q.10	QM
	SEPTEM	BER			
Basin-characteristics method Channel-width method Concurrent-measurement method Weighted standard error (log) Weighted standard error (pct)	0.149 .303 .548 .169	0.053 .322 .625 .120	0.040 .314 .646 .102	0.045 .301 .654 .083	0.049 .259 .693 .091
Basin-characteristics method Channel-width method Weighted standard error (log) Weighted standard error (pct)	.663 .337 .363	.571 .429 .317	.633 .367 .289	.561 .439 .213	.623 .377 .238
Concurrent-measurement method Basin-characteristics method Weighted standard error (log) Weighted standard error (pct)	.550 .450 .181	.637 .363 .139	.653 .347 .122	.681 .319 .101	.711 .289 .105
Concurrent-measurement method Channel-width method Weighted standard error (log) Weighted standard error (pct)	.565 .435 .172	.633 .367 .120	.654 .346 .102	.666 .334 .083	.706 .294 .092

