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U.S. DEPARTMENT OF THE INTERIOR  
UNITED STATES GEOLOGICAL SURVEY

MINERAL RESOURCES OF THE MARBLE CANYON  
WILDERNESS STUDY AREA, WHITE PINE COUNTY,  
NEVADA, AND MILLARD COUNTY, UTAH

By

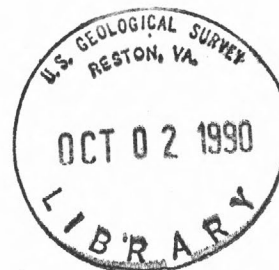
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Prepared by the U.S. Geological Survey and the U.S. Bureau of Mines



for the U.S. Bureau of Land Management

*Twined*

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## **STUDIES RELATED TO WILDERNESS**



### **Bureau of Land Management Wilderness Study Area**

The Federal Land Policy and Management Act (Public Law 94-579, October 21, 1976) requires the U.S. Geological Survey and U.S. Bureau of Mines to conduct mineral surveys on certain areas to determine the mineral values, if any, that may be present. Results must be made available to the public and be submitted to the President and the Congress. This report presents the results of a mineral survey of part of the Marble Canyon (NV-040-086) Wilderness Study Area, White Pine County, Nevada, and Millard County, Utah.

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## **SUMMARY**

### **Abstract**

The 19,150-acre Marble Canyon Wilderness Study Area (NV-040-086) was evaluated for mineral resources (known) and mineral resource potential (undiscovered), and field work was conducted in 1987. The acreage includes 6,435 acres that is now designated as part of the Mount Moriah Wilderness under the Nevada Wilderness Protection Act of 1989 (S. 974), most but not all of which is included in 8,300 acres for which the U.S. Bureau of Land Management requested a mineral survey. In this report, the "wilderness study area," or simply "the study area" refers to the entire 19,150-acre tract.

The area is underlain by quartzite shale and carbonate rocks. The northern Snake Range décollement is a detachment surface within the study area that separates rocks of similar age but different metamorphic grade. Large inferred subeconomic limestone and marble resources in the study area have no special or unique properties. The mineral resource potential for limestone and marble is high in two canyons and is moderate in the rest of the wilderness study area. Parts of the study area above and along the northern Snake Range décollement have low potential for undiscovered deposits of gold, silver, copper, lead, zinc, tungsten, molybdenum, beryllium, and fluorite. A zone around barite-bearing rock penetrated by adits inside the southeast boundary of the study area has moderate potential for barite, and the surrounding area has low potential for barite; both areas also have low potential for silver, copper, lead, zinc, and tungsten. The entire study area has moderate potential for oil and gas and low potential for geothermal energy resources.

### **Character and Setting**

The Marble Canyon Wilderness Study Area covers approximately 19,150 acres and is 10 mi west of Gandy, Utah (fig. 1). The terrain is rugged, and the elevation ranges from about 9,331 ft at Thunder Mountain to about 5,240 ft in the southeast corner of the study area. Access to and within the study area is by four-wheel-drive roads in Marble Wash, Coyote Canyon, Bars Canyon, and Smith Creek.

The study area is underlain by shales, quartzites, and carbonate rocks of early Paleozoic age (see "Appendixes" for geologic time chart). The oldest exposed rocks in the area are shale, quartzite, and marble of Early and Middle Cambrian age. Shale of Late Cambrian age has been faulted over the older metamorphosed rocks. Dolomite, shale, and quartzite of Ordovician and Silurian age are also in fault contact with the Upper Cambrian rocks. The structural history of the study area includes the development of a detachment fault that juxtaposes metamorphosed Paleozoic rocks beneath, and in flat-fault contact with, less metamorphosed to nonmetamorphosed rocks. This detachment surface is part of the northern Snake Range décollement.

### **Identified Mineral Resources**

Limestone has been prospected and quarried in and near the study area. Large inferred subeconomic limestone and marble resources inside the study area are without special or unique properties. The carbonate rocks are suitable for use in aggregate, but there are no nearby markets and similar quality resources are present elsewhere.

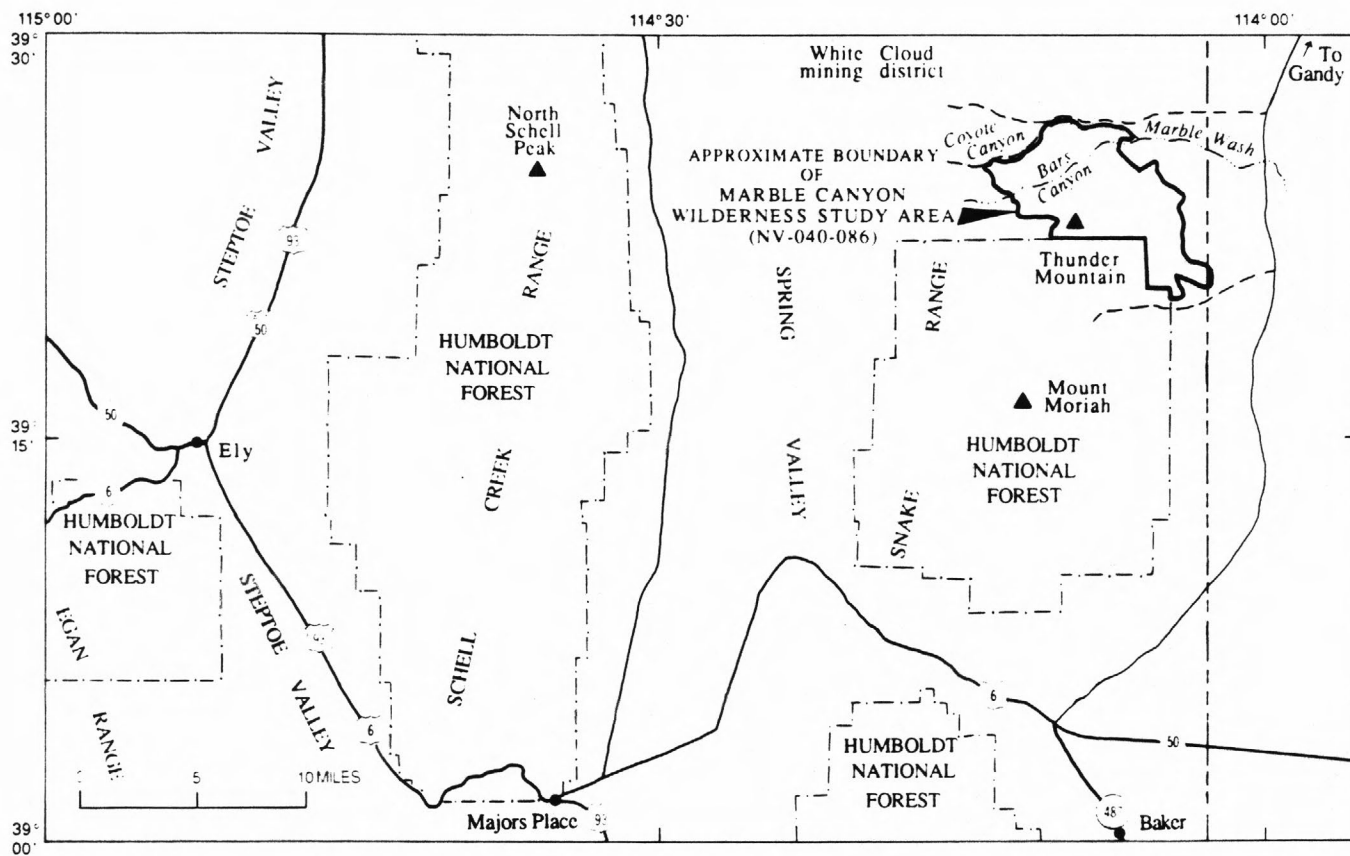


Figure 1.--Index map showing location of Marble Canyon Wilderness Study Area, White Pine County, Nevada, and Millard County, Utah.

The exposed northern Snake Range décollement within the study area is not mineralized, as indicated by a lack of surface expressions and by stream-sediment data. Samples taken from fault-breccia zones in nonmetamorphosed carbonate rocks above the northern Snake Range décollement contain some pathfinder elements typical of detachment-related deposits, but no mineral occurrences were observed.

### **Mineral Resource Potential**

In the lower parts of Bars Canyon and Marble Wash the mineral resource potential for limestone and marble is high, and in the rest of the study area it is moderate.

Areas of the Marble Canyon Wilderness Study Area underlain by upper-plate rocks and the zone of the northern Snake Range décollement have low mineral resource potential for gold, silver, copper, lead, zinc, tungsten, molybdenum, beryllium, and fluorite. The gold and silver are associated with low-angle faults within those areas. Geochemical evidence suggests the resource potential for copper, lead, zinc, tungsten, and molybdenum in those areas. The geophysical study suggests that no plutons are buried beneath the study area, but anomalous concentrations of beryllium in geochemical samples collected from the study area could be explained by the presence of an undetected buried pluton.

Adits inside the southeast boundary of the study area reveal exposures of barite, and anomalous concentrations of barium were measured in geochemical samples from upper-plate rocks there. A narrow zone around the adits has moderate mineral resource potential for barite, and the surrounding area has low mineral resource potential for barite. On the basis of geochemical data, the adits and a wider zone around them have low mineral resource potential for silver, copper, lead, zinc, and tungsten.

The metamorphosed lower plate rocks, the thin sequences of faulted upper plate rocks, and the extrusive volcanic rocks exposed in the study area are not conducive to the formation and accumulation of hydrocarbons. The resource potential for oil and gas in the entire study area is moderate, but this assessment may be too high due to the lack of likely source rocks. There are two thermal springs in Spring Valley to the west of the study area and other thermal springs just east of the study area. Range-front faults in the region may provide a conduit for the circulation of thermal water, and such a system could extend into the study area. Therefore, the entire Marble Canyon Wilderness Study Area has low potential for geothermal energy resources associated with low-temperature thermal springs.

### **INTRODUCTION**

This mineral survey was requested by the U.S. Bureau of Land Management and is the result of a cooperative effort by the U.S. Geological Survey and the U.S. Bureau of Mines. An introduction to the wilderness review process, mineral survey methods, and agency responsibilities was provided by Beikman and others (1983). The U.S. Bureau of Mines evaluates identified resources at individual mines and known mineralized areas by collecting data on current and past mining activities and through field examination of mines, prospects, claims, and mineralized areas. Identified resources are classified according to a system that is a modification of that described by McKelvey (1972) and U.S. Bureau of Mines and U.S. Geological Survey (1980). U.S. Geological Survey studies are designed to provide a scientific basis for assessing the potential for undiscovered mineral resources by determining geologic units and structures, possible environments of mineral deposition, presence of geochemical and geophysical anomalies, and applicable ore-deposit models. Goudarzi (1984) discussed mineral assessment methodology and terminology as they apply to these surveys. See "Appendixes" for the definition of levels of mineral resource potential and certainty of assessment and for the resource/reserve classification.

## Location and Physiography

The Marble Canyon Wilderness Study Area comprises approximately 19,150 acres in the Basin and Range physiographic province. It is situated in the northern Snake Range, principally in eastern White Pine County, Nev., but extending a short distance into Millard County, about 10 mi southwest of Gandy, Utah. The north and west boundaries follow Marble Wash and Coyote Canyon, respectively. The south boundary follows a four-wheel-drive road from Coyote Canyon to the base of Thunder Mountain and the north boundary of the Mount Moriah Division of Humboldt National Forest. The east boundary follows a four-wheel-drive road along the alluvial apron near the 6,000-ft contour. Elevations in the study area range from about 9,331 ft at Thunder Mountain to about 5,240 ft in the southeast corner of the study area. The Nevada Wilderness Protection Act of 1989 designates 6,435 acres of the study area as part of the Mount Moriah Wilderness (fig. 1).

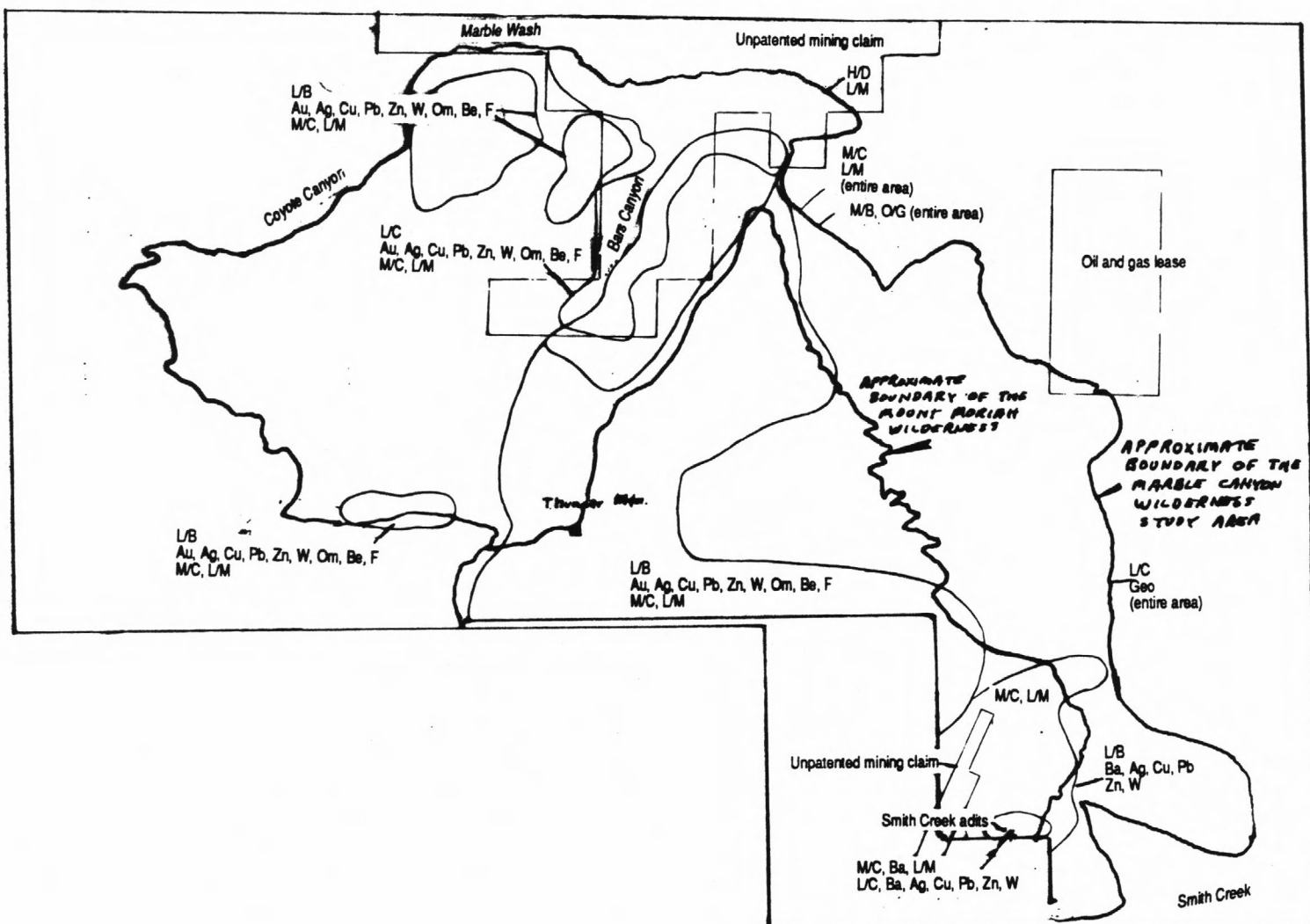
Climate in the study area is classified as arid; precipitation averages 10 in. per year and supports vegetation of the Sonoran and Transition Zones. The higher elevations are in the Transition Zone where vegetation includes piñon and bristlecone pine. Bristlecone pines grow primarily on limestone and are conspicuously absent from hillsides underlain by quartzite. The middle elevations are in the Upper Sonoran Zone and support white fir and incense cedar. On the dryer lower slopes, the vegetation is typically Lower Sonoran sagebrush and mountain mahogany.

## Procedures and Sources of Data

The U.S. Geological Survey and the U.S. Bureau of Mines conducted detailed field investigations of the Marble Canyon Wilderness Study Area in the summer of 1987. This work included geologic mapping at a scale of 1:24,000, field checks of existing geologic maps, geochemical sampling, and the examination of outcrops for evidence of mineralization.

A detailed literature search was made for geologic and mining information pertinent to the study area, and U.S. Bureau of Land Management records were examined for information on mining claims and oil and gas leases. Two U.S. Bureau of Mines geologists spent 6 days in the field for ground reconnaissance and an examination of prospects within and near the study area, and they collected 21 rock and 19 stream-sediment samples. Those samples were analyzed either by inductively coupled plasma-atomic emission spectroscopy or fire assay-atomic absorption methods by Chemex Labs Inc., Sparks, Nev. Sample data were discussed by Kness (1989) and are summarized in this report. Complete sample data are available for public inspection at the U.S. Bureau of Mines, Intermountain Field Operations Center, Building 20, Denver Federal Center, Denver, CO.

General stratigraphic studies in the region include a correlation of stratigraphic units in the Great Basin (Langenheim and Larson, 1973), a description of the stratigraphic section near Eureka, Nev. (Nolan and others, 1956), and paleogeographic interpretations (Stewart and others 1977). Detailed stratigraphic studies include descriptions of upper Precambrian and Lower Cambrian strata (Stewart, 1970), Lower and Middle Cambrian strata (Robison, 1960), and Upper Cambrian strata (Palmer, 1960). Whitebread and Lee (1961) and Whitebread (1969) described the geology of the Wheeler Peak area in what is now Great Basin National Park, 15 mi south of the study area. Hose and Blake (1970, 1976) compiled existing and original geologic mapping of White Pine County, and Hose (1981) presented the geology of the Mount Moriah Division of Humboldt National Forest (formerly the Mount Moriah Roadless Area and now included in the Mount Moriah Wilderness). Christiansen and others (1987) did detailed mapping just south of the study area. The geology in the western part of figure 2 is generalized from detailed mapping by Jeffrey Lee (1990).



### EXPLANATION

- H Area having high mineral resource potential  
M Area having moderate mineral resource potential  
L Area having low mineral resource potential

#### Certainty of Classification

- B Data only suggest level of potential  
C Data give good indication of level of potential  
D Data clearly define level of potential

#### Commodities

- |     |                   |     |                      |
|-----|-------------------|-----|----------------------|
| Ag  | Silver            | L-M | Limestone and marble |
| Au  | Gold              | Mo  | Molybdenum           |
| Ba  | Barite            | O-G | Oil and gas          |
| Be  | Beryllium         | Pb  | Lead                 |
| F   | Fluorite          | W   | Tungsten             |
| Geo | Geothermal energy | Zn  | Zinc                 |
| Cu  | Copper            |     |                      |

Figure 2.--Mineral resource potential and generalized geology of Marble Canyon Wilderness Study Area, White Pine County, Nevada, and Millard County, Utah. Geology generalized from Lee (1987) and unpublished mapping of Jeffrey Lee (written commun., 1989).



General treatments of the structural geology of the region include studies of central-northeastern Nevada (Misch, 1960; Misch and Hazzard, 1962) and descriptions of the structural evolution of the eastern Great Basin (Hose and Daneš, 1968, 1973; Hose and Blake, 1969). Hazzard and others (1953) and Nelson (1966) described the structural development of the region, including the large-scale thrust faulting in the northern Snake Range, a theory that has been superseded by the interpretation of the northern Snake Range décollement (Wernicke, 1981; Miller and others, 1983). Whitebread (1966), Coney (1974), Rowles (1982), and McGrew (1986) specifically addressed the northern Snake Range décollement. Gering (1987), Miller and others (1988b), and Jeffrey Lee (1990) discussed the metamorphic history of the Snake Range.

General treatments of mineral resources include a bibliography of Nevada resources (Weimer-McMillion and others, 1983), a map of resource areas in Nevada (Wong, 1983), and a description of computer-accessible Nevada mineral-resource data (Sherlock and Tingley, 1985). Cox and Singer (1986) presented mineral-deposit and grade-tonnage models. Lincoln (1923) provided an early treatment of mining districts in Nevada. Morton and others (1977) provided age data on ore deposits in Nevada. Darton (1908) reported on marble in White Pine County. Smith (1976) reported on the mineral resources of White Pine County, and Carlson and others (1984) presented the mineral resource potential of the Mount Moriah Division of Humboldt National Forest. In 1983, this study area was included in a U.S. Bureau of Land Management Geology, Energy, and Minerals (G.E.M.) inventory program (Great Basin GEM Joint Venture, 1983). Bullock and others (1989) presented the data from the U.S. Geological Survey stream-sediment sampling program for this wilderness study area. Kness (1989) presented the results of the U.S. Bureau of Mines mineral resource appraisal for this wilderness study area.

## **Acknowledgments**

The authors greatly appreciate the cooperation of Shaaron Netherton, William D. Robison, and Brian Amme of the U.S. Bureau of Land Management in Ely, Nev. Carl and Sheriton Baker made the Horse Canyon spring on their ranch available for a field camp site. Jeffrey Lee shared unpublished mapping that is part of his Ph.D. dissertation at Stanford University. Elizabeth Miller and her students from Stanford University shared detailed geologic mapping information on the northern Snake Range. R.A. Schreiner assisted the U.S. Bureau of Mines author with his field studies.

## **APPRAISAL OF IDENTIFIED RESOURCES**

By Richard F. Kness

*U.S. Bureau of Mines*

### **Mining Activity**

No mining has occurred, but prospecting and limestone quarrying have taken place in the study area. Unpatented mining claims for marble are along parts of Bars Canyon and Marble Wash within the study area and north of it. The following information on nearby mining districts and a mining area is summarized from Hose and Blake (1976) and Smith (1976).

The Marble Canyon mining district encompasses the Marble Wash area and the region north of it (fig 1). Marble claims were first located in 1891. No production records are available, but small quantities of marble were quarried for grave headstones and transported to Garrison, Utah, 30 mi to the south. Darton (1908) investigated the marble as a building stone source. Pink-colored marble was tested for crushed stone from 1966 to 1969. Results of these tests have not been located.

The White Cloud mining district is just west of the Marble Canyon district (fig. 1; Kness, 1989). The Lead King mine, the only one in the district, is near the summit of White Cloud Mountain, 7 to 8 mi northwest of the study area. Recorded district production from 1949 to 1952 totalled 92 tons of ore containing a trace of gold, 197 oz silver, 225 lb copper, 25,985 lb lead, and 1,553 lb zinc valued at \$4,481 at the time it was mined. No other information on the district was found.

The Mount Moriah mining area, in the Humboldt National Forest, is named for mines and prospects on and around Mount Moriah (fig. 1). This mining area has produced quartzite building stone, garnet abrasives, gold, silver, copper, lead, and zinc (Wood, 1983). The Grand View tungsten prospect is on the west slope of Mount Moriah. An assay of coarse-grained scheelite enclosed in 3 ft of silicified limestone shows as much as 7 percent tungsten oxide ( $\text{WO}_3$ ).

## **Oil and Gas**

Drilling has occurred near the north boundary of the study area. A dry-hole marker was located in Marble Wash (fig. 2) but no records were located. Oil and gas leases have been filed near and on the east boundary of the study area (fig. 2).

## **Results of Field Investigation**

### **Smith Creek Adits**

The two short Smith Creek adits are just inside the south boundary of the study area (fig. 2). Barite, calcite, and manganese oxide are in northeast-striking detachment-fault breccia zones in upper-plate limestone. The breccia zones could not be traced beyond the workings due to surface talus. Christiansen and others (1987) mapped this upper-plate limestone as an unnamed Middle Cambrian carbonate unit. Samples of fault breccia contain anomalous amounts of arsenic, beryllium, bismuth, mercury, molybdenum, tungsten, and especially barium and manganese (Kness, 1989, fig. 2). Gold and uranium were not detected, and silver was at detection limits (0.2 ppm). No resources could be defined. The unpatented mining claim 0.5 mi west of the adits is shown in U.S. Bureau of Land Management records but no evidence of workings was found.

### **Limestone and Marble**

Large inferred limestone and marble resources are present in the study area but must meet industry standards or specifications in order to be marketable. Limestone and marble are carbonate rocks of essentially the same composition and are used in the chemical and construction industries. They have the same end uses and are cut, crushed, or ground to produce a wide variety of stone products. Other types of rocks may be utilized or substituted for carbonate rocks: granite, sandstone, and slate may be used as dimension or cut stone, and sand and gravel is frequently used as an alternative to crushed limestone or marble.

Impurities such as chert, clay, iron, organic matter, and silica may affect utilization of carbonate rock. Pure calcitic or dolomitic marble is white; any coloration results from impurities. Gray, pink, and white marble were noted during the field investigation. The quartz content of the upper marble units in the northern Snake Range is high; it averages 15 to 20 percent but may be as high as 40 percent (Nelson, 1969, p. 335). Therefore the marble is too impure to be marketable.

Most chemical uses require high-calcium limestone containing more than 95 percent calcium carbonate ( $\text{CaCO}_3$ ); some uses specify more than 97 percent  $\text{CaCO}_3$  (Carr and Rooney,

1983). Limestone and marble samples in and near the study area contain calcium carbonate ranging from 75.73 to 94.83 percent (Kness, 1989, table 2); thus, the best material in the study area is of marginal quality. In general, marble samples contain a higher percentage of magnesium carbonate than limestone samples and thus are unlikely to be suitable for chemical uses.

Limestone and marble may be cut and used as dimension stone. Principal uses are in building construction and monuments. Most dimension marble is cut into thin slabs (0.875 to 1.25 in.) for use as veneer or curtain wall construction. Quarry blocks typically weigh 15 to 30 tons and may be rejected for unacceptable color or, more commonly, structural unsoundness (cracks, joints, or fractures). Because most bedded deposits are sawed parallel to bedding, small-scale folds may make it impossible to saw satisfactory slabs. Some marble units in the study area are contorted, folded, and fractured. Dimension-stone quarries in general produce a large amount of waste because only about 25 percent of the quarried stone is sent to the mill for sawing, and 50 percent of that may be lost during manufacturing (see Power, 1972, 1978). A quarry in this area is likely to have higher waste percentages and therefore is less likely to be a viable commercial venture.

Marble dimension stone marketability is governed largely by stone availability and aesthetic qualities (uniform color, patterns, and texture) determined by architects and architectural fashions. A new prospect or deposit may be successful if it matches or closely matches a currently fashionable stone available in large quantities (see Power, 1972). Marble production continues in traditional areas that have favorable geology and market proximity such as the Pittsford district in Vermont and the Tate district in Georgia. Significant tonnage of marble dimension stone is imported from Italy. Further testing is necessary to determine if marble in and near the study area is suitable for dimension stone. Soundness and aesthetic qualities of the stone are largely unknown. Large inferred resources are present, but impurities, small-scale folds, and fractures are present in some marble units. It is not known whether large, uniform, sound blocks are present. Selective mining might be required if suitable marble units are identified. Transportation costs are high because of weight and the need for special handling to prevent finished-stone damage.

Limestone and marble may be crushed and used as aggregate. A private mining consultant's report (L.C. Armstrong, E.J. Longyear and Co., written commun., 1970) estimated that the Marble Wash deposit contains marble resources totaling 300 million tons, but aggregate is a high-bulk, low-unit-value commodity. In general, hauling crushed stone more than 30 mi is not economical (Schenck and Torries, 1983). The distance from the Marble Wash marble deposits to U.S. Highway 6 and 50 is about 37 mi and Ely, Nev. is the closest market, another 75 mi away. Transportation costs exceed the value of the commodity.

## Conclusions

Large inferred subeconomic limestone and marble resources are present in the study area; however, these rocks were determined to have no special or unique properties. Low calcium carbonate content may preclude utilization for chemical purposes. Aesthetic qualities of the marble are unknown, and the presence of impurities, folds, and fractures may make it unsuitable for dimension stone. Even though the carbonate rocks are suitable for use as aggregate, the transportation costs to the nearest markets exceed the unit value, and market areas may utilize closer carbonate rocks or substitute other rocks. The remoteness of the area would limit development of the carbonate rocks for all but local uses. Local demand at present is nonexistent. No nearby markets were identified, and similar quality resources are present elsewhere.



## ASSESSMENT OF MINERAL RESOURCE POTENTIAL

By Michael F. Diggles, Gary Nowlan, H. Richard Blank, Jr., and Susan M. Marcus

*U.S. Geological Survey*

### Geology

Rocks not exposed at the surface within the Marble Canyon Wilderness Study Area presumably present at depth, include the Proterozoic and Lower Cambrian Prospect Mountain Quartzite, a thick sequence of fine-to medium-bedded quartzite that contains interbeds of argillite. Exposed rocks in the study area consist of a lower plate of Middle and Upper Cambrian metamorphosed sedimentary rocks in detachment-fault contact with an upper plate of Lower Cambrian through Devonian limestone, dolomite, shale, and quartzite that have mostly undergone only brittle deformation. From Late Proterozoic through Early Triassic, the region of the study area was the site of relatively continuous continental shelf sedimentation (Stewart and others, 1977). Lower-plate rocks were metamorphosed to amphibolite facies during the Cretaceous and underwent retrograde metamorphism to greenschist facies in the Tertiary (J. Lee and others, 1988; Huggins and Wright, 1989; Miller and others, 1988b).

### Lower-Plate Rocks

The location of the northern Snake Range décollement and the stratigraphy of the upper and lower plates were modified from the early work of Hose and Blake (1970, 1976) by Jeffrey Lee (1990). The lower plate of the décollement within the study area includes Middle and Upper Cambrian rocks (fig. 2). Elsewhere in the northern Snake Range, the décollement cuts through the Middle Cambrian Pole Canyon Limestone (Huggins and Wright, 1989).

Middle Cambrian rocks consist of the Raiff Limestone of Young (1960) and the Monte Neva Formation of Young (1960). These light-gray to dark-gray metamorphosed limestones are combined on figure 2 (map unit *Grm*). The Middle Cambrian rocks are interrupted by faulting or erosion everywhere in the study area, and their total thickness is not known.

The Upper Cambrian Notch Peak Limestone and Dunderberg Shale are also combined on figure 2 (map unit *End*). The Dunderberg consists mostly of light-olive-gray to medium-olive-gray silty shale intercalated with thin beds of limestone. It is only about 70 ft thick in the study area. The Notch Peak Limestone, which underlies most of the study area, is medium-gray, massive, locally dolomitic, and contains lenses of chert. The basal part is thin bedded and silty. Metamorphosed parts of the Notch Peak unit include the marble for which the area has the highest resource potential. Whitebread (1969) shows a thickness of 1,600 to 1,800 ft for the Notch Peak Limestone south of the study area.

### Upper-Plate Rocks

The oldest rocks exposed in the study area are Lower Cambrian shale and quartzite (map unit *Gsq*) that include the Pioche Shale. The Pioche consists of 150 to 600 ft of dark-greenish-gray to olive-gray silty shale and clay shale intercalated with siltstone and includes some sandstone near its base and minor limestone in its upper part. It is exposed in an upper-plate block in the southwest part of the study area. Where the Pioche is metamorphosed within the study area, metamorphic minerals include biotite, muscovite, and garnet. The Middle Cambrian Monte Neva Formation and Raiff Limestone (map unit *Grm*) comprise the rocks in many small exposures of the upper plate in the northern part of the study area. The Upper Cambrian Notch Peak Limestone is the oldest unit in much of the upper plate, particularly in the southern part of the study area, where the northern Snake Range décollement lies within it in most places.

Ordovician rocks in the study area crop out in the upper plate and consist of the undivided Pogonip Group and Eureka Quartzite (map unit Oep). The Pogonip Group in the study area is limited to small outcrop areas near the décollement in the south-central part. It is composed of platy to thin-bedded, fine- to coarse-textured, gray detrital limestone. The Eureka Quartzite consists of about 310 ft of white fine-grained equigranular quartzite with minor dolomitic quartzite.

The undivided Ordovician Fish Haven Dolomite and Silurian Laketown Dolomite (map unit SOf) crop out in the south-central part of the study area. These formations consist of light-gray to dark-gray dolomite. South of the study area, chert is locally abundant and the Fish Haven has a dolomitic sandstone at its base.

Outcrops of Devonian rocks in the eastern part of the study area north of Petes Knoll are the undivided Guilmette Formation, Simonson Dolomite, and Sevy Dolomite (map unit Dgs). The Guilmette Formation consists of alternating dolomite and limestone in roughly equal amounts. The Sevy Dolomite is medium-gray to medium-light-gray, microcrystalline, well-bedded dolomite with sandy dolomite and quartzite near the top. The Simonson Dolomite is darker gray and commonly more coarsely crystalline.

Permian and Pennsylvanian rocks also crop out north of Petes Knoll (map unit PPre). The undivided Pennsylvanian Ely Limestone and Permian Riepe Springs Limestone of Steele (1960) are medium-gray limestones of organic detrital material.

### Tertiary Rocks and Quaternary Surficial Deposits

A thin layer of Tertiary volcanic rocks unconformably overlies the Paleozoic rocks on the east side of Petes Knoll. These rocks consist of alkali olivine basalt (map unit Tv). Quaternary surficial deposits include Holocene poorly sorted gravel, sand, and silt in stream-channel and flood-plain deposits and landslides (map unit Qal). Older Quaternary stream-channel deposits form terraces that are dissected by streams.

### Structure and Metamorphism

The presence of metamorphosed Paleozoic rocks beneath, and in flat-fault contact with, less-metamorphosed to nonmetamorphosed rocks in the Snake Range has been discussed since the early 1950's (Hazzard and others, 1953) when the nearly flat fault was interpreted as a thrust fault. The fault was interpreted later to be a detachment fault (Whitebread, 1966; Coney, 1974) that has juxtaposed younger rocks eastward over older rocks. Detachment faults have been described in many areas of the Western United States on the margins of metamorphic core complexes (Crittenden and others, 1980). Miller and others (1983) interpreted the fault to have undergone little or no horizontal displacement. An early deformation resulted in a northwest trend to the metamorphic fabric in the lower-plate rocks. A later lower-plate metamorphic fabric has a southerly trend that resulted from the Tertiary movement on the northern Snake Range décollement (Coney, 1974).

Rocks that have been metamorphosed to amphibolite-facies in the southern Snake Range have been cut by 160-Ma plutonic rocks, thereby constraining the age of the metamorphism to no younger than Jurassic (Miller and others, 1988b). Those rocks that were metamorphosed during the Jurassic have undergone retrograde metamorphism during more recent events.

Cretaceous metamorphism was to upper greenschist or amphibolite facies. The age of metamorphism is inferred to be  $78 \pm 9$  Ma on the basis of uranium-lead zircon and monazite data (D.E. Lee and others, 1981). Jeffrey Lee and others (1988) and Jeffrey Lee (1990) also obtained a Cretaceous age from argon determinations ( $^{40}\text{Ar}/^{39}\text{Ar}$ ) on hornblende. This metamorphism

coincides with the time of movement on the Sevier thrust belt east of the study area (Miller and others, 1988a; Miller and Gans, 1989a, 1990).

Tertiary metamorphism is shown by ductile strain that produced penetrative foliation in the Paleozoic rocks of the lower plate of the northern Snake Range décollement (Miller and others, 1988b). Potassium-argon ages on micas from lower-plate rocks yield Tertiary ages (D.E. Lee and others, 1970) and  $^{40}\text{Ar}/^{39}\text{Ar}$  data show that the rocks were still being deformed as recently as the Miocene (J. Lee and others, 1988; J. Lee, 1990).

The northern Snake Range décollement is interpreted by Miller and others (1983) and Gans and others (1985, 1986) to have developed as a ductile-brittle transition zone at a depth of 3.8 to 4.4 mi. The detachment took place during the doming of the décollement (Miller and others, 1983; McCarthy, 1986) and the extension of this part of the Basin and Range province. The uplift and detachment represent an Oligocene to Miocene event (Miller and others, 1987). The rocks above the décollement underwent two generations of east-dipping high-angle normal faulting as the extension progressed. These high-angle faults do not extend below the décollement (Gans and Miller, 1983).

Basin-and-Range-type high-angle normal faulting in the Snake Range started in the early to middle Oligocene (Gans and others, 1989) and has occurred continuously since then. It is responsible for the relief in the study area and for the formation of Snake Valley to the east. Normal faults cut volcanic rocks 35 m.y. in age (Gans, 1987). Miller and Gans (1989b) put age constraints on the relative uplift of the range and downdropping of the adjacent basins from about 17 to 15 Ma and 11 to 13 Ma.

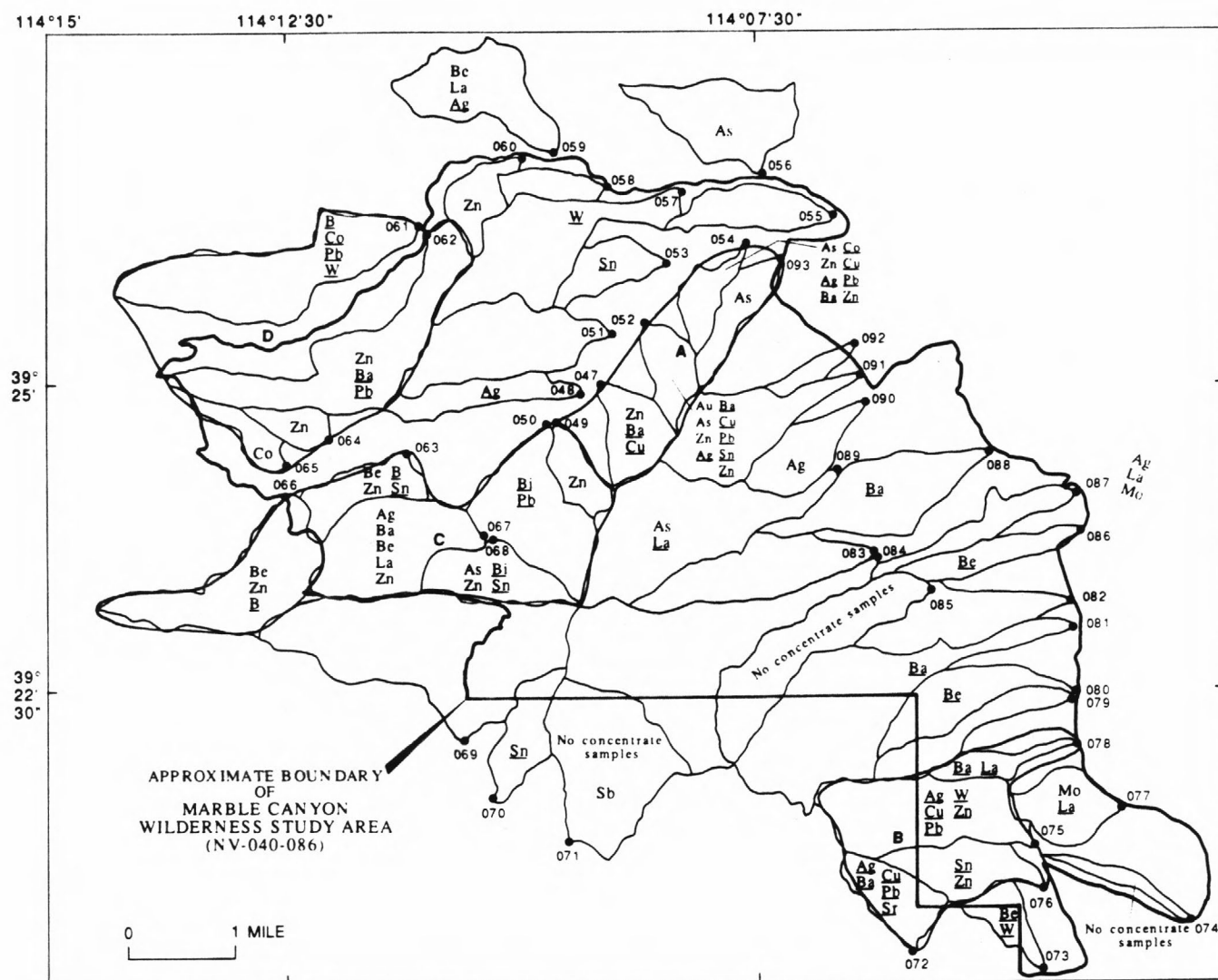
## Geochemistry

A reconnaissance geochemical survey of the Marble Canyon Wilderness Study Area was conducted in May 1987. Stream-sediment samples that were collected from all major streams and tributaries represent eroded bedrock that underlies the drainage basin whence the sediment came. Panned concentrates of the sediment were prepared and analyzed separately. The panned-concentrate fraction may contain minerals related to metallization processes.

## Methods and Background

Samples of sediment were collected at 47 sites on streams draining the wilderness study area and vicinity (fig. 3). Stream-sediment samples represent a composite of material eroded from the drainage basin of the stream sampled. Panned-concentrate samples derived from stream sediment contain selectively concentrated minerals that may be ore related and may include elements not easily detected in stream-sediment samples.

A stream-sediment sample and a panned-concentrate sample were collected at each site. The stream-sediment sample was air dried and then sieved through an 80-mesh stainless-steel sieve. The portion that passed through the screen was later pulverized to minus-100 mesh size prior to analysis. For the panned-concentrate sample at each site, enough stream sediment was screened through a 10-mesh sieve to obtain about 20 lb. The minus-10-mesh sample was panned to remove most of the quartz, feldspar, carbonate-rock material, clay-sized material, and organic matter. Most of the panned-concentrate samples were further concentrated by a series of steps that utilized bromoform (specific gravity 2.8) and magnetic separations to produce nonmagnetic heavy-mineral-concentrate samples that include most nonmagnetic ore minerals and accessory minerals such as sphene, zircon, apatite, and rutile. Prior to analysis, the 44 nonmagnetic heavy-mineral-concentrate samples were pulverized to minus-100 mesh size.



### Explanation

- Outline of drainage basin  
 — Outline of anomalous area  
 o Sample-collection site

Ag Silver  
 As Arsenic  
 Au Gold  
 B Boron  
 Be Beryllium  
 Bi Bismuth  
 Co Cobalt  
 Cu Copper

La Lanthanum  
 Mo Molybdenum  
 Pb Lead  
 Sb Antimony  
 Sn Tin  
 Sr Strontium  
 W Tungsten  
 Zn Zinc

**Figure 3.--Geochemical anomaly map of Marble Canyon Wilderness Study Area, Nevada and Utah.** Concentrations of As-i, Sb-i, and Zn-i determined by inductively coupled plasma spectroscopy, Au-a determined by atomic absorption, and other concentrations determined by emission spectrography. Underlined elements anomalous in nonmagnetic heavy-mineral-concentrate sample; other elements anomalous in minus-80-mesh stream-sediment sample.



The stream-sediment and nonmagnetic heavy-mineral-concentrate samples were analyzed by emission spectrography for antimony, arsenic, barium, beryllium, bismuth, boron, cadmium, calcium, chromium, cobalt, copper, gold, iron, lanthanum, lead, magnesium, manganese, molybdenum, nickel, niobium, scandium, silver, strontium, thorium, tin, titanium, tungsten, vanadium, yttrium, zinc, and zirconium. In addition, the nonmagnetic heavy-mineral-concentrate samples were analyzed for gallium, germanium, palladium, phosphorus, platinum, and sodium by emission spectrography. The stream-sediment samples were also analyzed for gold by graphite-furnace atomic absorption, for uranium by ultraviolet fluorimetry, and for antimony, arsenic, bismuth, cadmium, and zinc by inductively coupled plasma spectroscopy (ICP). Analytical data, sampling sites, and references to analytical methods are presented by Bullock and others (1989).

## Results and Interpretation

Table 1 lists selected elements determined in each sample type and for each of them the lower and upper limits of determination, the range of concentrations, the 50th and 90th percentile concentrations, and the threshold (highest background) concentrations. Threshold concentrations were established by statistical and subjective examination of the data. The number of samples that have anomalous concentrations of each selected element is also given.

Elements in table 1 that are present in anomalous concentrations in samples from various sampling sites are shown on figure 3. The elements selected for inclusion in table 1 and on figure 3 are elements that may reflect the presence of mineral deposits. The anomalous concentrations of beryllium and lanthanum are unusual in the absence of igneous rocks and the predominance of carbonate rocks in the sampled area; these two elements are usually associated with granitic rocks or pegmatites. The average concentration of beryllium in limestone is less than 1 part per million (ppm), and the average concentration of lanthanum in limestone is 4 ppm (Rose and others, 1979, p. 552, 567). Barium commonly accompanies base metals (copper, lead, zinc) in mineral deposits but also commonly forms bedded barite deposits in sedimentary rocks that have no significant amounts of base metals. Therefore, it is not unusual that anomalous concentrations of barium in samples sometimes accompany anomalous concentrations of other elements and sometimes do not.

Many samples from throughout the study area contain anomalous concentrations of only a single element or *none*. Samples from a few drainage basins contain anomalous concentrations of two to eight elements. Several geochemically anomalous areas, delineated A-D on figure 3, have similar geochemical signatures that are centered on anomalous concentrations of lead and zinc. The anomalous concentrations of lead and zinc are accompanied by anomalous concentrations of silver, gold, arsenic, barium, beryllium, bismuth, boron, cobalt, copper, lanthanum, tin, strontium, or tungsten; silver and gold are considered the most important after lead and zinc. The intensity of the geochemical anomaly of area A is the greatest, and that of area D is the least.

Anomalous areas A-D include a few adjoining drainage basins that have single-element anomalies. Beryllium and lanthanum are not included as contributors to multiple-element anomalies. When they are omitted, a few basins having them become single-element basins and are not included in the anomalous areas. However, an anomalous concentration of any element may be significant, and figure 3 only portrays relative significance.

Table 1. Statistics for selected elements in drainage samples collected in and near the Marble Canyon Wilderness Study Area, White Pine County, Nevada, and Millard County, Utah.

[Results based on analyses of 47 stream-sediment samples and 44 nonmagnetic heavy-mineral-concentrate samples. Determined by inductively coupled plasma spectroscopy for As-i, Sb-i, and Zn-i, by atomic absorption for Au-a, and by emission spectrography for other elements. N, not detected at lower limit of determination; L, detected below lower limit of determination; G, greater than upper limit of determination; <, less than lower limit of determination; ---, upper limit is open ended]

Element	Limits of determination, ppm		Range, ppm		Percentiles, ppm		Threshold concentration, ppm	Number of samples having anomalous concentrations
	Lower	Upper	Minimum	Maximum	50th	90th		
Minus-80-mesh stream-sediment samples								
Ag	0.5	5,000	N	L	N	N	N	3
As-i	5	20,000	<	36	8	16	14	6
Au-a	0.001	—	<	0.007	0.001	0.001	0.001	1
Ba	20	5,000	150	1,500	300	700	700	1
Be	1	1,000	N	5	2	3	3	4
La	20	1,000	N	100	20	50	50	3
Mo	5	2,000	N	L	N	N	N	2
Sb-i	2	20,000	<	2	<	<	<	1
Zn	200	10,000	N	L	N	N	N	4
Zn-i	2	18,000	11	60	39	53	49	10
Nonmagnetic heavy-mineral-concentrate samples								
Ag	1	10,000	N	7	N	L	N	6
B	20	5,000	N	300	N	100	100	3
Ba	50	10,000	L	G	600	G	3,000	8
Be	2	2,000	N	15	N	4	7	3
Bi	20	2,000	N	70	N	N	N	2
Co	20	5,000	N	50	N	L	L	3
Cu	10	50,000	N	30	N	13	10	5
La	100	2,000	N	300	L	100	150	3
Pb	20	50,000	N	G	70	1,000	700	7
Sn	20	2,000	N	100	N	L	N	6
Sr	200	10,000	N	2,000	N	200	300	1
W	50	20,000	N	50	N	N	N	4
Zn	500	20,000	N	10,000	N	N	N	4

Samples from some drainage basins contain anomalous concentrations of from one to many elements (fig. 3). Except for beryllium and lanthanum, the elements present in anomalous concentrations are compatible with the possible existence of at least two models of mineral deposition. The first model is the polymetallic replacement deposit (Morris, 1986; Mosier and others, 1986), and the second is deposits associated with detachment faults (Bouley, 1986; Spencer and Welty, 1986).

A carbonate replacement deposit, the Silver Peak mine, is about 5 mi southwest of the wilderness study area. Recorded production from the mine is 275 oz silver, 0.06 oz gold, and about 38,000 lb of lead (Smith, 1976). Polymetallic replacement deposits characteristically occur in limestone, dolomite, and shale that have been intruded by plutons (Morris, 1986). Igneous bedrock is not present in the sampled area but intrusions are less than 2 mi from the southwest side of the wilderness study area, and volcanic rocks are less than 1 mi from the east side (Hose and Blake, 1976). Igneous rocks that were not detected in the geophysics study may be present beneath the upper-plate terrain that covers extensive areas of the wilderness study area. The presence of concealed granitic rocks could explain the anomalous concentrations of beryllium and lanthanum as well as the elements that are included in the geochemical signature interpreted here as possibly representing carbonate replacement deposits related to igneous intrusions. Barton (1987) observed that two-mica granites that intrude carbonate host rocks can give rise to a characteristic lithophile-element suite that includes beryllium, fluorine, tungsten, molybdenum, and zinc.

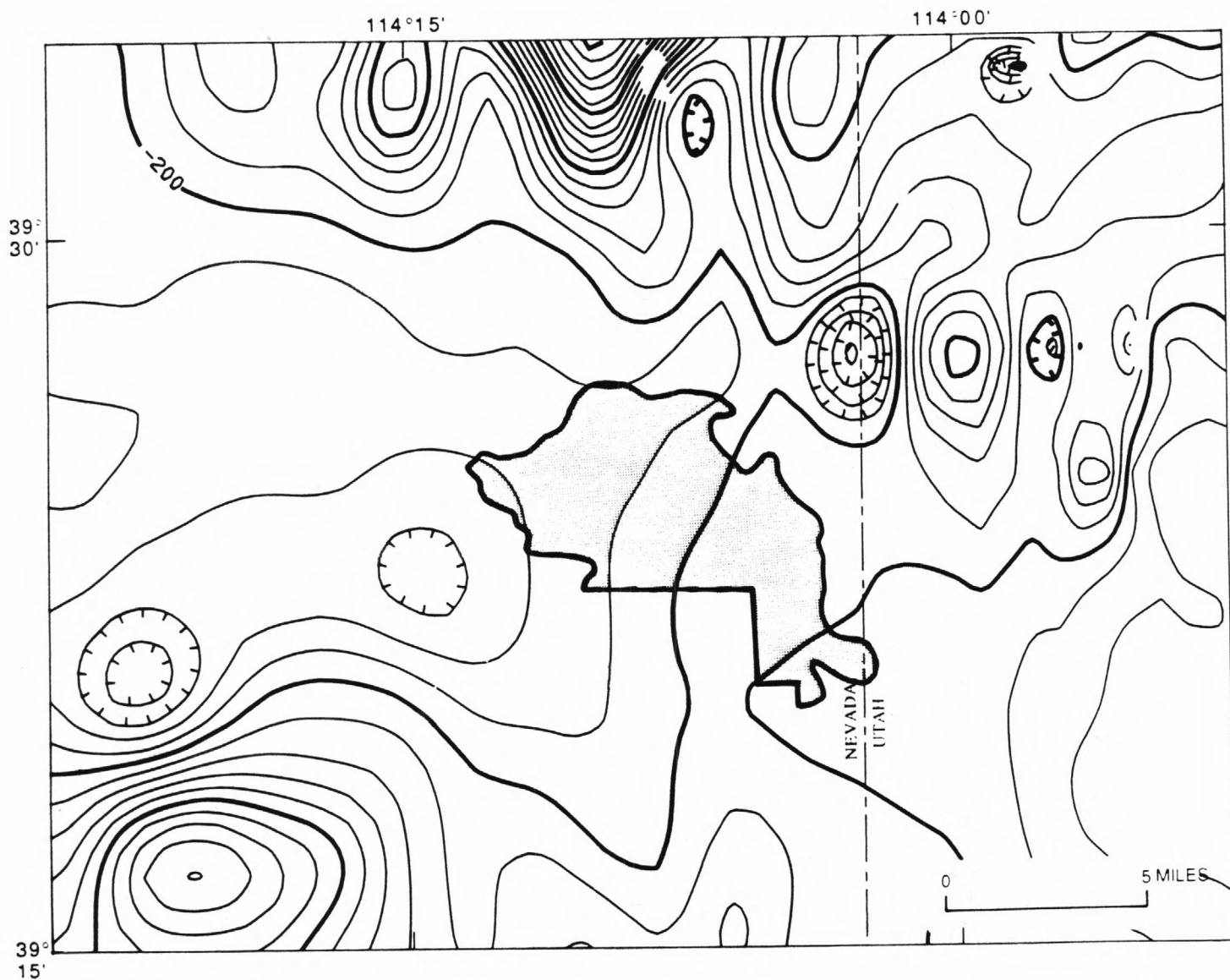
The detachment-fault model of mineral deposition (Spencer and Welty, 1986) does not require the presence of igneous rocks. The most intense geochemical anomalies in the sampled area are spatially associated with detachment faults. Alteration associated with detachment faults is probably the most reasonable explanation for the geochemical anomalies in the sampled area.

## Geophysics

Regional aeromagnetic and gravity maps for the Marble Canyon Wilderness Study Area and vicinity have been examined for indications of mineral resources. Two sources of aeromagnetic data have been utilized: surveys of the Ely, Nev., and Delta, Utah, 1° x 2° quadrangles flown for the National Uranium Resource Evaluation (NURE) program, and a survey of east-central Nevada flown for the U.S. Geological Survey (U.S. Geological Survey, 1978). The NURE traverses were flown east-west at a nominal 400 ft above ground and 3-mi spacings, and the U.S. Geological Survey traverses were flown on east-west headings at 12,000 ft barometric elevation and 4-mi spacings (U.S. Geological Survey, 1978). The main source of gravity data was the National Solar-Terrestrial and Geophysical Data Center in Boulder, Colo., which supplied principal facts for about 295 stations distributed mainly in low-lying sections of the area of interest. These data were supplemented by thirteen additional stations established during July 1988 specifically to improve coverage in remote portions. It should be noted that the area considered in the aeromagnetic interpretation is necessarily much larger than the area encompassed by the wilderness study area boundaries; only seven stations lie within the wilderness study area.

## Aeromagnetic data

The anomalous aeromagnetic field after removal of the International Geomagnetic Reference Field (IGRF) was computed from data of the NURE surveys (fig. 4). Because the traverse spacing is wide compared to the flight elevation, the shapes of short-wavelength features of the field are not accurately depicted. Sources of most of the anomalies are unknown. Volcanic rocks, which crop out in the northern part of the map area, have aeromagnetic signatures ranging from negligible to very intense and of relatively short wavelength.



**Figure 4.--Aeromagnetic anomaly map of region including Marble Canyon Wilderness Study Area, Nevada and Utah (shaded area). Contour interval 20 nanoteslas, hachures indicate closed areas of lower gravity.**

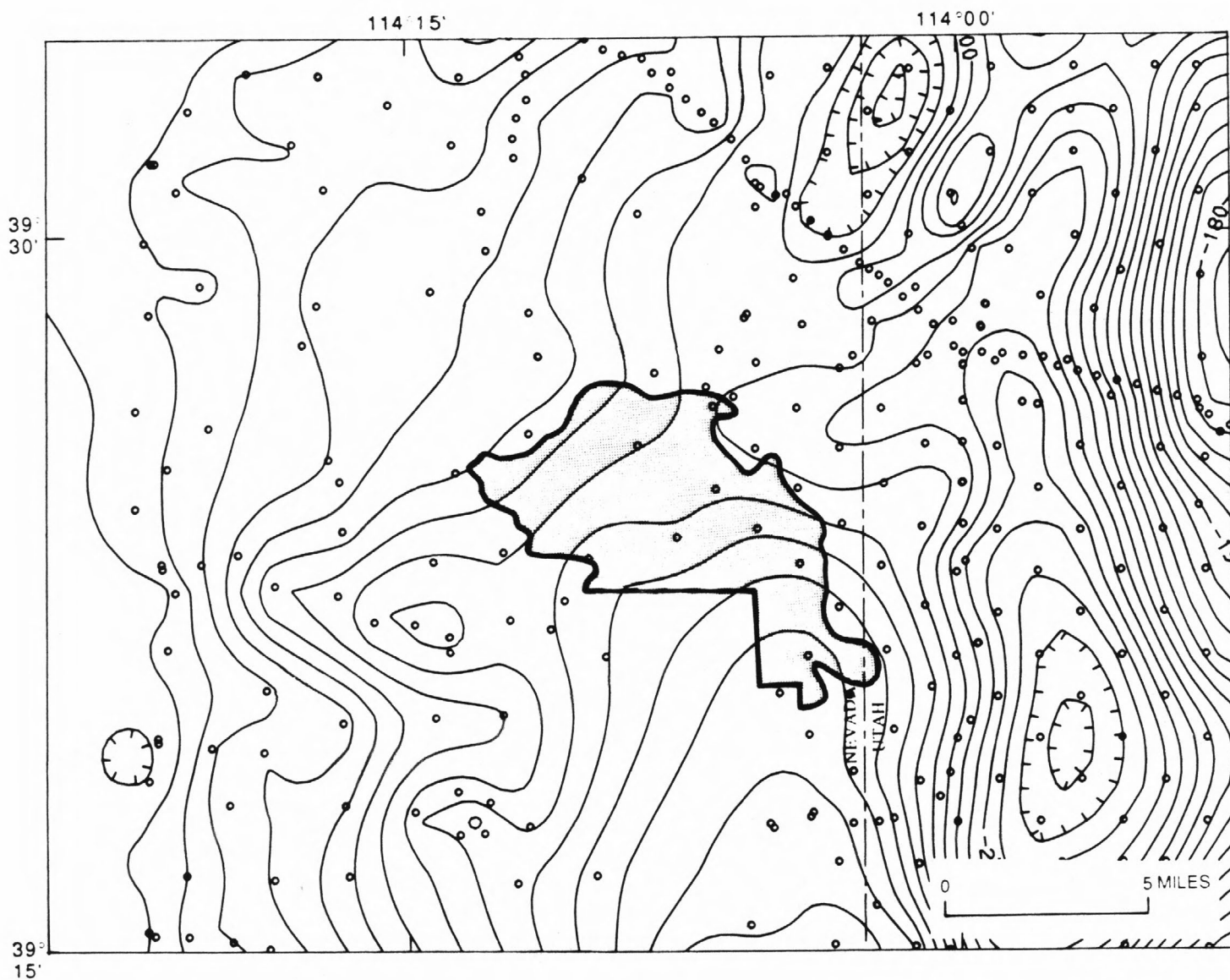


Outcrops of intrusive rock south of the study area have no detectable aeromagnetic expression. Long-wavelength positive anomalies of unknown origin north and southwest of the wilderness study area cover areas underlain by mostly Paleozoic rock or unconsolidated alluvium. The wilderness study area occupies a broad, west-northwest-trending anomaly trough between relatively positive areas where the field is more disturbed; the origin of this trough is speculative, but it may reflect a structural depression in the crystalline basement. No aeromagnetic evidence suggests that the wilderness study area is underlain by intrusive rocks or contains concealed structural discontinuities. However, the Jurassic and Cretaceous two-mica granites of the region are notable for their absence of a strong aeromagnetic signature (Grauch and others, 1988).

### Gravity data

The complete-Bouguer anomaly field, terrain-corrected to a distance of 42 mi (Hayford-Bowie zones A-O), is shown in figure 5. All data reduction and terrain corrections employed a standard density of 2.67 g/cm<sup>3</sup> and followed conventional U.S. Geological Survey procedures (see, for example, Cordell and others, 1982). The anomalous gravity field bears little resemblance to the aeromagnetic field. The anomalies have no direct correlation of anomalies with outcrops of volcanic or intrusive rocks; the only apparent consistency is that the most positive areas are largely associated with exposed bedrock. The main sources of positive anomalies are expected to be thick sections of Paleozoic miogeoclinal carbonate rocks or possibly carbonate rocks that have a core of elevated Proterozoic crystalline basement. A significant maximum, possibly reflecting a basement rise (Hose, 1981), occurs 5 to 6 mi southeast of Mount Moriah. Minima are produced mostly by thick accumulations of unconsolidated material and occur in the Snake Valley to the east or in Spring Valley and its adjuncts to the west of the wilderness study area. A major structural discontinuity on the east side of the Snake Range is indicated by the steep anomaly gradient east-southeast of the wilderness study area, several miles east of the nearest outcrops; terrain between this feature and the range front is probably an extensive pediment. No structural discontinuities can be identified from the gravity field within the boundaries of the wilderness study area. Absence of geophysical evidence of buried plutons reduces the likelihood that there may be pluton-related resources at depth. Such a pluton could be present, but only if it lacked a notable aeromagnetic signature; the data are permissive for the presence of a buried two-mica granite such as those described by Grauch and others (1988).

The gravity field southwest of figure 5 is not well constrained, but anomaly values that appear to fall off steadily to the southwest toward Sacramento Pass suggest an increase in thickness of the relatively low-density Proterozoic and Lower Cambrian quartzite section in that direction.



**Figure 5.--Bouguer gravity anomaly map of region including Marble Canyon Wilderness Study Area, Nevada and Utah (shaded area). Contour interval, 2 milligals; hachures indicate closed areas of lower gravity. Open circles represent gravity stations.**

## Mineral Resource Potential

### Limestone and Marble

The lower parts of Bars Canyon and Marble Wash have high mineral resource potential, certainty level D, for medium-purity limestone and marble for use as aggregate. The rest of the wilderness study area may have pockets of medium-purity limestone and marble, although none that have the purity of that observed in the canyons were. The rest of the wilderness study area, therefore, has moderate mineral resource potential, certainty C, for limestone and marble for use as aggregate.

### Gold and Silver

Parts of the Marble Canyon Wilderness Study Area underlain by rocks of the upper plate and the zone of the northern Snake Range décollement have low mineral resource potential for gold and silver associated with detachment faults. Bouley (1986) has described detachment-fault-type deposits. As so few such deposits are being mined, a grade-tonnage model has yet to be developed (D.A. Singer, oral commun., 1989). Gans and others (1988) suggests that the northern Snake Range décollement may be a rotated normal fault. If this is the case, the appropriate model would be carbonate-hosted gold-silver deposit described by Berger (1986) and Bagby and others (1986). With either model, the mineral resource potential for gold and silver is low, certainty level C, in the southeast side of lower Bars Canyon, where the geochemical data more strongly support this assessment. Other areas near the décollement, including both the upper and lower plates, have low mineral resource potential, certainty level B, for gold and silver.

### Copper, Lead, and Zinc

Geochemical evidence suggests that parts of the study area underlain by the upper-plate rocks and the zone of the northern Snake Range décollement or projections of that structure extending east and south to the Mount Moriah Division of Humboldt National Forest (Carlson and others, 1984), have low resource potential, certainty level B, for copper, lead, and zinc.

### Tungsten and Molybdenum

The geophysical study does not suggest that plutons are concealed beneath the study area, but the two-mica granites of the region are notable for their lack of geophysical expression (Grauch and others, 1988). Anomalous concentrations of tungsten and (or) molybdenum in stream-sediment and fault-breccia samples collected from scattered parts of the study area could be explained if a buried pluton were present. The detachment fault could also have provided a locus for concentration of the metals. Parts of the study area underlain by the upper-plate rocks and the zone of the northern Snake Range décollement have low mineral resource potential, certainty level B, for tungsten and molybdenum.

### Beryllium and Fluorite

Although there are beryllium and fluorite deposits in the Pioche Shale about 30 mi to the south in the Wheeler Peak region in the southern Snake Range (Whitebread and Lee, 1961), those deposits are confined to an (informal) limestone unit locally known as the "Wheeler limestone" in the lower part of the Pioche. A possible source of the beryllium is a quartz monzonite stock that intrudes the Wheeler Peak area (Whitebread and Lee, 1961). The Pioche Shale (part of map unit Csq) crops out in the southwestern part of the study area, but the so-called Wheeler limestone does not extend into the northern Snake Range; mineralization due to the intrusion of the stock is confined to the Wheeler Peak area. The geophysical study did not

show evidence that plutons are buried beneath the study area; however, the Snake Range is known for the presence of two-mica plutons that have no geophysical signature (Grauch and others, 1988). Anomalous concentrations of beryllium in geochemical samples collected from the study area could be explained by a buried two-mica pluton (Barton, 1987). The detachment fault could also have provided a locus for concentration of the beryllium. Parts of the study area underlain by the upper-plate rocks and the zone of the northern Snake Range décollement have low mineral resource potential, certainty B, for beryllium and fluorite.

## **Barite**

Barite is present in the walls of two adits that are just inside the southeast boundary of the study area (Kness, 1989), and anomalous concentrations of barium (greater than 10,000 ppm, Bullock and others, 1989) were measured in the panned concentrate from the basin containing the adits. Therefore, a zone around the adits has moderate mineral resource potential, certainty level C, for barite, and the surrounding area has low mineral resource potential, certainty level B, for barite. On the basis of geochemical data, the zone around the adits and the surrounding area also have low mineral resource potential, certainty level B, for silver, copper, lead, zinc, and tungsten.

## **Oil and Gas**

Sandberg (1982, 1983) evaluated the petroleum potential of wilderness lands in Nevada using the following four major parameters that govern oil and gas accumulation: presence of source rocks, hydrocarbon maturation, reservoir rocks, and traps. He rated the Marble Canyon Wilderness Study Area as having medium potential because of optimum maturity of source rocks in the Devonian and Mississippian Pilot Shale and the Mississippian Chainman Shale. However, metamorphosed lower-plate rocks, thin sequences of faulted upper-plate rocks, and extrusive volcanic rocks exposed in the study area may preclude the presence of hydrocarbons. Therefore, the oil and gas resource potential is moderate in the entire study area, but, on the basis of the degree of metamorphism of the possible lower-plate source rocks and the faulting and tilting of possible upper-plate trap rocks, this assessment could be too high. The certainty level of assessment, therefore, is B.

## **Geothermal Energy**

Bliss (1983) lists geothermal springs and wells for the Ely 1° by 2° quadrangle. He shows two thermal springs in Spring Valley to west of the study area. We observed thermal springs just east of the study area, 1 mi across the Nevada-Utah State line on the road to Gandy. Range-front faulting may be providing a conduit for thermal water, and that system could extend into the study area. The entire Marble Canyon Wilderness Study Area has low geothermal energy potential, certainty level B, for low-temperature thermal springs.

## **REFERENCES CITED**

- Bagby, W.C., Menzie, W.D., Mosier, D.L., and Singer, D.A., 1986, Grade and tonnage model of carbonate-hosted Au-Ag, in Cox, D.P., and Singer, D.A., eds., *Mineral deposit models*: U.S. Geological Survey Bulletin 1693, p. 175-177.
- Barton, M.D., 1987, Lithophile-element mineralization associated with Late Cretaceous two-mica granites in the Great Basin: *Geology*, v. 15, p. 337-340.
- Beikman, H.M., Hinkle, M.E., Frieders, Twila, Marcus, S.M., and Edward, J.R., 1983, Mineral surveys by the Geological Survey and the Bureau of Mines of Bureau of Land Management wilderness study areas: U.S. Geological Survey Circular 901, 28 p.



- Berger, B.R., 1986, Descriptive model of carbonate-hosted Au-Ag, *in* Cox, D.P., and Singer, D.A., eds., Mineral deposit models: U.S. Geological Survey Bulletin 1693, p. 175.
- Bliss, J.D., 1983, Nevada-Basic data for thermal springs and wells as recorded in GEOTHERM, part A: U.S. Geological Survey Open-File Report 83-433-A, 101 p.
- Bouley, B.A., 1986, Descriptive model of gold on flat faults, *in* Cox, D.P., and Singer, D.A., eds., Mineral deposit models: U.S. Geological Survey Bulletin 1693, p. 251.
- Bullock, J.H., Jr., McHugh, J.B., Briggs, P.H., and Nowlan, G.A., 1989, Analytical results and sample locality map for stream-sediment and nonmagnetic heavy-mineral-concentrate samples from the Marble Canyon Wilderness Study Area (NV-040-086), White Pine County, Nevada: U.S. Geological Survey Open-File Report 89-0024, 11 p.
- Carlson, R.R., Martin, R.A., and Wood, R.H., II, 1984, Mineral resource potential map of the Mount Moriah Roadless Area, White Pine County, Nevada: U.S. Geological Survey Miscellaneous Field Studies Map MF-1244-B, scale 1:62,500.
- Carr, D.D., and Rooney, L.F., 1983, Limestone and dolomite, *in* Lefond, S.J., ed., Industrial minerals and rocks, 5th ed., V. 2: New York, Society of Mining, Metallurgical, and Petroleum Engineers, Inc., p. 833-868.
- Coney, P.J., 1974, Structural analysis of the Snake Range "décollement," east-central Nevada: Geological Society of America Bulletin, v. 85, no. 6, p. 973-978.
- Cordell, Lindrith, Keller, G.R., and Hildenbrand, T.G., 1982, Bouguer gravity map of the Rio Grande rift, Colorado, New Mexico, and Texas: U.S. Geological Survey Geophysical Investigations Map GP-949, scale 1:1,000,000.
- Christiansen, P., Cladouhos, T., DeBari, S., El-Shazly, A., Gans, E., Gans, P., Gehrke, I., Grissom, G., Huggins, C., Jones, C., Lee, J., Lomas, M., Miller, E., Minobe, W., Pinto, M., Riley, T., Stahl, D., and Thomas, A., 1987, Geologic map of the Horse Canyon/Smith Creek area of the northern Snake Range, White Pine County, Nevada: Stanford, Calif., Stanford University, Department of Earth Sciences, geological survey map, scale 1:12,000 [available at the main library].
- Cox, D.P., and Singer, D.A., eds., 1986, Mineral deposit models: U.S. Geological Survey Bulletin 1693, 379 p.
- Crittenden, M.D., Jr., Coney, P.J., and Davis, G.H., eds., 1980, Cordilleran metamorphic core complexes: Geological Society of America Memoir 153, 490 p.
- Darton, N.H., 1908, Building stones-Marble of White Pine County, Nev., near Gandy, Utah: U.S. Geological Survey Bulletin 340-G. pt. 1, p. 377-380.
- Gans, P.B., 1987, Cenozoic extension and magmatism in the eastern Great Basin: Stanford, Calif., Stanford University Ph.D. dissertation, 174 p.
- Gans, P.B., Clark, D.H., Miller, E.L., Wright, M.E., and Sutter, J.F., 1986, Structural development of the Kern Mountains and northern Snake Range [abs.]: Geological Society of America Abstracts with Programs, v. 18, no. 2, p. 108.

- Gans, P.B., Mahood, G.A., and Schermer, Elizabeth, 1989, Synextensional magmatism in the Basin and Range Province; a case study from the eastern Great Basin: Geological Society of America Special Paper 233, 53 p.
- Gans, P.B., and Miller, E.L., 1983, Style of mid-Tertiary extension in east-central Nevada, in Gurgel, K.D., ed., Geologic excursions in the overthrust belt and metamorphic core complexes of the intermountain region, Guidebook-Part 1, GSA Rocky Mountain and Cordilleran Sections Meeting: Utah Geological and Mineral Survey Special Studies 59, p. 107-160.
- Gans, P.B., Miller, E.L., McCarthy, Jill, and Ouldcott, M.L., 1985, Tertiary extensional faulting and evolving ductile-brittle transition zones in the northern Snake Range and vicinity: new insights from seismic data: *Geology*, v. 13, p. 189-193.
- Gans, P.B., Lee, Jeffrey, Miller, E.L., Kunk, Mick, and Sutter, J.F., 1988, Uplift history of mid-crustal rocks in the eastern Great Basin [abs.]: Geological Society of America Abstracts with Programs, v. 20, no. 7, p. A17.
- Gering, R.L., 1987, A study of the metamorphic petrology of the northern Snake Range, east-central Nevada: Dallas, Tex., Southern Methodist University M.S. thesis, 75 p.
- Goudarzi, G.H., 1984, Guide to preparation of mineral survey reports on public lands: U.S. Geological Survey Open-File Report 84-787, 51 p.
- Grauch, V.J.S., Blakely, R.J., Blank, H.R., Oliver, H.W., Plouff, Donald, and Ponce, D.A., 1988, Geophysical delineation of granitic plutons in Nevada: U.S. Geological Survey Open-File Report 88-11, scale 1:1,000,000, 2 plates.
- Great Basin GEM Joint Venture, 1983, Granite Spring G-E-M resources area (GRA No. NV-09) technical report (WSA NV 040-086): U.S. Bureau of Land Management Contract YA-554-RFP2-1054, 49 p.
- Hazzard, J.C., Misch, Peter, Wiese, J.H., and Bishop, W.C., 1953, Large-scale thrusting in northern Snake Range, White Pine County, Nevada [abs.]: Geological Society of America Bulletin v. 64, no. 12, pt. 2, p. 1507-1508.
- Hose, R.K., 1981, Geological map of the Mount Moriah Further Planning (RARE II) Area, eastern Nevada: U.S. Geological Survey Miscellaneous Field Studies Map MF-1244-A, scale 1:62,500.
- Hose, R.K., and Blake, M.C., Jr., 1969, Structural development of the eastern Great Basin during the Mesozoic [abs.]: Geological Society of America Abstracts with Programs for 1969, part 5, p. 34.
- \_\_\_\_\_, 1970, Geologic map of White Pine County, Nevada: U.S. Geological Survey open-file map, scale 1:150,000.
- \_\_\_\_\_, 1976, Part I, Geology, in Hose, R.K., Blake, M.C., Jr., and Smith, R.M., Geology and mineral resources of White Pine County, Nevada: Nevada Bureau of Mines and Geology Bulletin 85, p. 1-35.
- Hose, R.K., and Daneš, Z.F., 1968, Late Mesozoic structural evolution of the eastern Great Basin [abs.]: Geological Society of America Special Paper 115, p. 102.

- 1973, Development of late Mesozoic to early Cenozoic structures of the eastern Great Basin, in DeJong, K.A., and Scholten, Robert, eds., Gravity and tectonics: New York, Interscience, John Wiley and Sons, p. 429-441.
- Huggins, C.C., and Wright, J.E., 1989, Superimposed Cretaceous and Tertiary metamorphism in the northern Snake Range, Nevada [abs.]: Geological Society of America Abstracts with Programs, v. 21 no. 5, p. 95.
- Kness, R.F., 1989, Mineral investigation of the Marble Canyon Wilderness Study Area (NV-040-086), White Pine County, Nevada, and Millard County, Utah: U.S. Bureau of Mines Open-File Report MLA 17-89, 22 p.
- Langenheim, R.L., Jr., and Larson, E.R., co-chairmen, 1973, Correlation of Great Basin stratigraphic units: Nevada Bureau of Mines and Geology Bulletin 72, 36 p.
- Lee, D.E., Marvin, R.F., Stern, T.W., and Peterman, Z.E., 1970, Modification of potassium-argon ages by Tertiary thrusting in the Snake Range, White Pine County, Nevada, in Geological Survey research 1970: U.S. Geological Survey Professional Paper 700-D, p. D92-D102.
- Lee, D.E., Stern, T.W., and Marvin, R.F., 1981, Uranium-thorium-lead isotopic ages of metamorphic monazite from the northern Snake Range, Nevada: Isochron/West, no. 31, p. 23.
- Lee, Jeffrey, 1990, Structural geology and  $^{40}\text{Ar}/^{39}\text{Ar}$  thermochronology of the northern Snake Range, Nevada: Stanford, Calif., Stanford University Ph.D. dissertation, 186 p. [Includes a geologic map of the Mormon Jack Pass quadrangle, White Pine County, Nevada, scale 1:24:000].
- Lee, Jeffrey, Huggins, C.C., Miller, E.L., Gans, P.B., Wright, J.E., Geving, R.L., and Sutter, J.F., 1988, Polymetamorphism in the northern Snake Range, Nevada [abs.]: Geological Society of America Abstracts with Programs, v. 20, no. 7, p. A17.
- Lee, Jeffrey, Miller, E.L., and Sutter, J.F., 1987, Ductile strain and metamorphism in an extensional tectonic setting—a case study from the northern Snake Range, Nevada, U.S.A., in Coward, M.P., Dewey, J., and Hancock, P., eds., Continental extensional tectonics: Geological Society of London Special Paper 28, p. 267-297.
- Lincoln, F.C., 1923, Mining districts and mineral resources of Nevada: Reno, Nev., Nevada Newsletter Publishing Company, 295 p.
- McCarthy, Jill, 1986, Reflection profiles from the Snake Range metamorphic core complex—a window into the mid-crust, in Barazangi, Muawia, and Brown, Larry, eds., Reflection seismology—the continental crust: Washington, D.C., American Geophysical Union Geodynamics Series v. 14, p. 281-292.
- McGrew, A.J., 1986, Deformation history of the southern Snake Range, Nevada, and the origin of the southern Snake Range décollement: Stanford, Calif., Stanford University M.S. thesis, 46 p.
- McKelvey, V.E., 1972, Mineral resource estimates and public policy: American Scientist, v. 60, p. 32-40.

- Miller, E.L., and Gans, P.B., 1989a, Cretaceous crustal structure and metamorphism in the hinterland of the Sevier thrust belt, western U.S. Cordillera: *Geology*, v. 17, p. 59-62.
- \_\_\_\_\_, 1989b, Uplift history of the Snake Range metamorphic core complex, Basin and Range province, USA, from apatite fission track data [abs.]: *Eos, Transactions, American Geophysical Union*, v. 70, p. 1309.
- \_\_\_\_\_, 1990, Comment and reply on "Cretaceous crustal structure and metamorphism in the hinterland of the Sevier thrust belt, western U.S. Cordillera": *Geology*, v. 18, p. 577-578.
- Miller, E.L., Gans, P.B., and Garing John, 1983, Snake Range décollement: an exhumed mid-Tertiary ductile-brittle transition: *Tectonics*, v. 2, p. 239-263.
- Miller, E.L., Gans, P.B., and Lee, Jeffrey, 1987, The Snake Range décollement, eastern Nevada: *Geological Society of America Centennial Field Guide, Cordilleran Section*, p. 77-82.
- Miller, E.L., Gans, P.B., and Wright, J.E., 1988a, Cretaceous crustal structure and metamorphism, hinterland of Sevier thrust belt, Nevada and Utah [abs.]: *Geological Society of America Abstracts with Programs*, v. 20, no. 7, p. A18.
- Miller, E.L., Gans, P.B., Wright, J.E., and Sutter, J.F., 1988b, Metamorphic history of the east-central Basin and Range province: tectonic setting and relationship to magmatism, in Ernst, W.G., ed, *Metamorphism and crustal evolution of the western United States*: Englewood Cliffs, N.J., Prentice-Hall, Rubey Volume VII, p. 649-682.
- Misch, Peter, 1960, Regional structural reconnaissance in central-northeast Nevada and some adjacent areas-observations and interpretations in Boettcher, J.W., and Sloan, W.W., eds., *Geology of east-central Nevada: Intermountain Association of Petroleum Geologists, 11th Annual Field Conference, Salt Lake City, Utah, 1960 Guidebook*, p. 17-42.
- Misch, Peter, and Hazzard, J.C., 1962, Stratigraphy and metamorphism of late Precambrian rocks in central northeastern Nevada and adjacent Utah: *American Association of Petroleum Geologists Bulletin*, v. 46, no. 3, p. 289-343.
- Morris, H.T., 1986, Descriptive model of polymetallic replacement deposits, in Cox, D.P., and Singer, D.A., eds, *Mineral deposit models: U.S. Geological Survey Bulletin 1693*, p. 99-100.
- Morton, J.L., Silberman, M.L., Bonham, H.F., Jr., Garside, L.J., and Noble, D.C., 1977, K-Ar ages of volcanic rocks, plutonic rocks, and ore deposits in Nevada and eastern California-determinations run under the USGS-NBMG cooperative program: *Isochron/West*, no. 20, p. 19-29.
- Mosier, D.L., Morris, H.T., and Singer, D.A., 1986, Grade and tonnage model of polymetallic replacement deposits, in Cox, D.P., and Singer, D.A., eds., *Mineral deposit models: U.S. Geological Survey Bulletin 1693*, p. 101-104.
- Nelson, R.B., 1966, Structural development of northernmost Snake Range, Kern Mountains, and Deep Creek Range, Nevada and Utah: *American Association of Petroleum Geologists Bulletin*, v. 50, p. 921-951.
- \_\_\_\_\_, 1969, Relation of history of structures in a sedimentary succession with deeper metamorphic structures, eastern Great Basin: *American Association of Petroleum Geologists Bulletin*, v. 53, p. 307-339.



- Nolan, T.B., Merriam, C.W., and Williams, J.S., 1956, The stratigraphic section in the vicinity of Eureka, Nevada: U.S. Geological Survey Professional Paper 276, 77 p.
- Palmer, A.R., 1960, Some aspects of the early Upper Cambrian stratigraphy of White Pine County, Nevada, and vicinity, *in* Boettcher, J.W., and Sloan, W.W., eds., *Geology of east-central Nevada: Intermountain Association of Petroleum Geologists, 11th Annual Field Conference, Salt Lake City, Utah, 1960 Guidebook*, p. 53-58.
- Power, W.R., 1972, An evaluation of building dimension stone deposits: *Mining Engineering*, vol. 24, no. 6, p. 42-44.
- \_\_\_\_\_, 1978, Economic geology of the Georgia marble district, *in* Twelfth forum on the geology of industrial minerals: Georgia Geologic Survey Information Circular 49, p. 59-68.
- Robison, R.A., 1960, Lower and Middle Cambrian stratigraphy of the eastern Great Basin, *in* Boettcher, J.W., and Sloan, W.W., eds., *Geology of east-central Nevada: Intermountain Association of Petroleum Geologists, 11th Annual Field Conference, Salt Lake City, Utah, 1960 Guidebook*, p. 43-52.
- Rose, A.W., Hawkes, H.E., and Webb, J.S., 1979, *Geochemistry in mineral exploration*, 2d edit.: London, Academic Press, 657 p.
- Rowles, L.D., 1982, Deformation history of the Hampton Creek area, northern Snake Range, Nevada: Stanford, Calif., Stanford University M.S. thesis, 80 p.
- Sandberg, C.A., 1982, Petroleum potential of wilderness lands, Nevada: U.S. Geological Survey Miscellaneous Geologic Investigations Map I-1542, scale 1:1,000,000
- \_\_\_\_\_, 1983, Petroleum potential of wilderness lands in Nevada, *in* Miller, B.M., ed., *Petroleum potential of wilderness lands in the Western United States*: U.S. Geological Survey Circular 902-H, p. H1-H11.
- Schenck, G.K., and Torries, T.F.K., 1983, Crushed stone, *in* Lefond, S.J., ed., *Industrial Minerals and rocks; 5th ed., vol. 1*: New York, Society of Mining, Metallurgical, and Petroleum Engineers, Inc., p. 60-80
- Sherlock, M.G., and Tingley, J.V., 1985, Nevada mineral-resource data: information available through the U.S. Geological Survey Mineral Resource Data System: U.S. Geological Survey Circular 966, 35 p.
- Smith, R.M., 1976, Part II, Mineral resources, *in* Hose, R.K., Blake, M.C., Jr., and Smith, R.M., *Geology and mineral resources of White Pine County, Nevada*: Nevada Bureau of Mines and Geology Bulletin 85, p. 36-105.
- Spencer, J.E., and Welty, J.W., 1986, Possible controls of base- and precious-metal mineralization associated with Tertiary detachment faults in the lower Colorado trough, Arizona and California: *Geology*, v. 14, p. 195-198.
- Steele, Grant, 1960, Pennsylvanian-Permian stratigraphy of east-central Nevada and adjacent Utah, *in* Boettcher, J.W., and Sloan, W.W., eds., *Geology of east-central Nevada: Intermountain Association of Petroleum Geologists, 11th Annual Field Conference, Salt Lake City, Utah, 1960 Guidebook*, p. 91-113.

- Stewart, J.H., 1970, Upper Precambrian and Lower Cambrian strata in the southern Great Basin, California and Nevada: U.S. Geological Survey Professional Paper 620, 206 p.
- Stewart, J.H., Stevens, C.H., and Fritsche, A.E., eds., 1977, Paleozoic paleogeography of the Western United States: Society of Economic Paleontologists and Mineralogists, Pacific Section, Pacific Coast Paleogeography Symposium 1, 502 p.
- U.S. Bureau of Mines and U.S. Geological Survey, 1980, Principles of a resource/reserve classification for minerals: U.S. Geological Survey Circular 831, 5 p.
- U.S. Geological Survey, 1978, Aeromagnetic map of east-central Nevada: U.S. Geological Survey Open-File Report 78-281, scale 1:250,000.
- Weimer-McMillion, Becky, Tingley, S.L., and Shilling, John, 1983, Bibliography of Nevada geology and mineral resources through 1980, an alphabetical listing by author: Nevada Bureau of Mines and Geology Special Publication 7, 184 p.
- Wernicke, Brian, 1981, Low-angle normal faults in the Basin and Range Province: nappe tectonics in an extended orogen: *Nature*, v. 291, p. 645-648.
- Whitebread, D.H., 1966, The Snake Range décollement and related structures [abs.]: Geological Society of America Program, 62nd Annual Meeting, Reno, Nev., p. 77.
- \_\_\_\_\_, 1969, Geologic map of the Wheeler Peak and Garrison quadrangles, Nevada and Utah: U.S. Geological Survey Miscellaneous Geologic Investigations Map I-578, scale 1:48,000.
- Whitebread, D.H., and Lee, D.E., 1961, Geology of the Mount Wheeler mine area, White Pine County, Nevada, *in* Geological Survey research 1961: U.S. Geological Survey Professional Paper 424-C, p. C120-C122.
- Wong, George, 1983, Preliminary map of the resource areas in the Basin and Range Province of Nevada: U.S. Geological Survey Open-File Report 83-721, 37 p.
- Wood, R.H., 1983, Mineral inventory of the Mt. Moriah Roadless Area, White Pine County: U.S. Bureau of Mines Open-File Report MLA 50-83, 27 p.
- Young, J.C., 1960, Structure and stratigraphy in north central Schell Creek Range *in* Boettcher, J.W., and Sloan, W.W., eds., *Geology of east-central Nevada*: Intermountain Association of Petroleum Geologists, 11th Annual Field Conference, Salt Lake City, Utah, 1960 Guidebook, p. 158-172.

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## APPENDIXES

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# DEFINITION OF LEVELS OF MINERAL RESOURCE POTENTIAL AND CERTAINTY OF ASSESSMENT

## LEVELS OF RESOURCE POTENTIAL

- H HIGH** mineral resource potential is assigned to areas where geologic, geochemical, and geophysical characteristics indicate a geologic environment favorable for resource occurrence, where interpretations of data indicate a high degree of likelihood for resource accumulation, where data support mineral-deposit models indicating presence of resources, and where evidence indicates that mineral concentration has taken place. Assignment of high resource potential to an area requires some positive knowledge that mineral-forming processes have been active in at least part of the area.
- M MODERATE** mineral resource potential is assigned to areas where geologic, geochemical, and geophysical characteristics indicate a geologic environment favorable for resource occurrence, where interpretations of data indicate reasonable likelihood for resource accumulation, and (or) where an application of mineral-deposit models indicates favorable ground for the specified type(s) of deposits.
- L LOW** mineral resource potential is assigned to areas where geologic, geochemical, and geophysical characteristics define a geologic environment in which the existence of resources is permissive. This broad category embraces areas with dispersed but insignificantly mineralized rock, as well as areas with little or no indication of having been mineralized.
- N NO** mineral resource potential is a category reserved for a specific type of resource in a well-defined area.
- U UNKNOWN** mineral resource potential is assigned to areas where information is inadequate to assign a low, moderate, or high level of resource potential.

## LEVELS OF CERTAINTY

- A** Available information is not adequate for determination of the level of mineral resource potential.
- B** Available information only suggests the level of mineral resource potential.
- C** Available information gives a good indication of the level of mineral resource potential.
- D** Available information clearly defines the level of mineral resource potential.

	A	B	C	D
LEVEL OF RESOURCE POTENTIAL ↑	U/A  UNKNOWN POTENTIAL	H/B HIGH POTENTIAL	H/C HIGH POTENTIAL	H/D HIGH POTENTIAL
		M/B MODERATE POTENTIAL	M/C MODERATE POTENTIAL	M/D MODERATE POTENTIAL
		L/B LOW POTENTIAL	L/C LOW POTENTIAL	L/D LOW POTENTIAL
				N/D NO POTENTIAL
	LEVEL OF CERTAINTY →			

Abstracted with minor modifications from:

Taylor, R.B., and Steven, T.A., 1983, Definition of mineral resource potential: *Economic Geology*, v. 78, no. 6, p. 1268-1270.

Taylor, R.B., Stoneman, R.J., and Marsh, S.P., 1984, An assessment of the mineral resource potential of the San Isabel National Forest, south-central Colorado: U.S. Geological Survey Bulletin 1638, p. 40-42.

Goudarzi, G.H., compiler, 1984, Guide to preparation of mineral survey reports on public lands: U.S. Geological Survey Open-File Report 84-0787, p. 7, 8.

## RESOURCE/RESERVE CLASSIFICATION

	IDENTIFIED RESOURCES		UNDISCOVERED RESOURCES	
	Demonstrated		Probability Range	
	Measured	Indicated	Hypothetical	Speculative
ECONOMIC	Reserves		Inferred Reserves	
MARGINALLY ECONOMIC	Marginal Reserves		Inferred Marginal Reserves	
SUB-ECONOMIC	Demonstrated Subeconomic Resources		Inferred Subeconomic Resources	

Major elements of mineral resource classification, excluding reserve base and inferred reserve base. Modified from McKelvey, V.E., 1972; Mineral resource estimates and public policy: American Scientist, v. 60, p. 32-40; and U.S. Bureau of Mines and U.S. Geological Survey, 1980, Principles of a resource/reserve classification for minerals: U.S. Geological Survey Circular 831, p. 5.



# GEOLOGIC TIME CHART

Terms and boundary ages used by the U.S. Geological Survey in this report

EON	ERA	PERIOD		EPOCH	AGE ESTIMATES OF BOUNDARIES IN MILLION YEARS (Ma)	
Phanerozoic	Cenozoic	Quaternary		Holocene	0.010	
				Pleistocene	1.7	
		Tertiary	Neogene Subperiod	Pliocene	5	
				Miocene	24	
			Paleogene Subperiod	Oligocene	38	
				Eocene	55	
				Paleocene	66	
	Mesozoic	Cretaceous		Late	96	
				Early		
					138	
		Jurassic		Late	205	
				Middle		
				Early		
	Triassic		Late	~240		
			Middle			
			Early			
	Paleozoic	Permian		Late	290	
				Early		
		Carboniferous Periods	Pennsylvanian	Late	~330	
				Middle		
			Mississippian	Early		
			Devonian		Late	360
					Middle	
					Early	
			Silurian		Late	410
					Middle	
					Early	
			Ordovician		Late	435
					Middle	
			Early			
	Cambrian		Late	500		
			Middle			
			Early			
	Proterozoic	Late Proterozoic			~570	
Middle Proterozoic			900			
Early Proterozoic			1600			
Archean	Late Archean			2500		
	Middle Archean			3000		
	Early Archean			3400		
pre-Archean <sup>2</sup>					(3800?)	
					4550	

<sup>1</sup>Rocks older than 570 Ma also called Precambrian, a time term without specific rank.

<sup>2</sup>Informal time term without specific rank.

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