



DESCRIPTION OF MAP UNITS

SURFICIAL DEPOSITS

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- Qa **Alluvium (Holocene)**—Sand, gravel and boulder deposits along the North Fork of Negroie Creek. Locally includes torrenital fan deposits at the mouths of tributary canyons. 0 m to about 10 m thick.
- Qb **Colluvium (Holocene)**—Thin to thick, well-sorted silt- to boulder-sized deposits typically on an adjacent to moderately steep slopes. Locally includes talus and slope wash deposits. Thickness 0 m to about 10 m.
- Qc **Landslide deposits (Holocene)**—Includes debris flows and rotated bedrock and slumped talus; grades into colluvium. Headwall of landslide blocks commonly in well-sorted, eolian sandstone (unit Ts) (see below).
- Qp **Quartzite (Pleistocene)**—Thin to thick, well-sorted, rounded sandstone and siltstone of the quartzite of Negroie Creek.
- Qs **Mixed alluvium and colluvium deposits (Holocene and/or Pleistocene)**—Silt- to boulder-sized sheet and slope wash deposits. Includes alluvial sand, gravel, and boulders along drainage channels. Deposits are typically well-sorted. Quartzite in the southeast corner of the quadrangle are mostly finer-silt-to-gravel-sized sheet wash material that has accumulated isolated small outcrop areas.
- Qp **Piedmont slope deposits (Pleistocene)**—Silt- to boulder-sized deposits capping mesas. Deposits in the vicinity of Perry Mesa in southwest part of the quadrangle have morphologies characteristic of alluvial fans that were subsequently pedimented. 0 m to a few meters thick.

GILA GROUP

Volcaniclastic conglomerate, sandstone and siltsstone, and interstitially baritic to dacitic lava flows and associated intrusives rocks

Q/Tg **Upper Eage Peak beds of the Gila Formation (Missineo to Pliocene)—Volcaniclastic conglomerates** locally occur in the upper Eage Peak, dacitic and red Eage Peak, dacite typically consist of light- to medium-gray, coarse, poorly bedded, matrix-supported conglomerate. This coarse conglomerate contains fragments of recent boulders of volcanic rocks and clasts of volcaniclastic rocks. Peak lithologies (units Teds, Ted) (see below). Clasts are subangular to subrounded, and moderately to poorly sorted. The matrix, which consists of fine-grained, muddy to sandy, silty, clayey, and silty clayey sandstone. These deposits grade progressively into gray to reddish-brown, moderately well-bedded, clast-supported conglomerate, and well-bedded, locally cross-bedded, conglomeratic sandstone with occasional lenses of cross-bedded sandstone. Such a pattern is characteristic of the volcaniclastic belt southwest corner of the quadrangle. Clasts within the finer sediments consist largely of locally derived volcanic rocks; the abundance of Eage Peak clasts decreases with distance from the volcanic belt. In the volcanic belt, except on road cuts and canyons in the southwestern part of quadrangle, deposits are poorly exposed. These deposits represent an apron of debris at the base of alluvial deposits of volcanic rocks. Well-developed, but reworked distal facies. Stratigraphically equivalent to the upper beds of the Gila Formation as mapped in the adjacent Milligan Mountain quadrangle (Q/Tg of Ratté and Bove, 1990). Maximum thickness in quadrangle about 200 m.

Dacite of Eage Peak (Missineo)—Lava flows, vent-agglomerate breccias and central volcanic plug of andesite to dacite composition (40–66 percent SiO₂) derived from the Eage Peak. The dacite is a fine-grained, medium-grained profile (about 8–12 degrees), central vent volcano with flanking flows extending up to 12 km from the vent area into the Milligan Mountain quadrangle (Ratté and Bove, 1990). The dacite is a fine-grained, medium-grained, massive, flow-folding, and flow layering characterize the Eage Peak flows. Sheeted joints are developed parallel to flow layers and vertical shrinkage cracks are also common. The dacite is a fine-grained, medium-grained, and by flow breccia zones up to 15 meters thick, and the margins of flow lobes typically are encased in tractor-type breccias.

Tedp **Lower Eage Peak beds—Lignite and volcaniclastic breccia (66 percent SiO₂)** measures about 0.5 km in diameter. Contains 25–30 percent phenocrysts of plagioclase, orthopyroxene, clinopyroxene, and rare hornblende, and contains phenocrysts of plagioclase, orthopyroxene, and clinopyroxene and minor magnetite and apatite occur in the volcaniclastic breccias. The breccia is finely granular. Forms the prominent, mostly treeless knob atop Eage Peak.

Upper Eage Peak flows—Flows are seriate porphyritic, range from black (glassy) to red (crystalline) and have a fine-grained, medium-grained texture. Contains 8–12 percent phenocrysts and microphenocrysts of plagioclase, orthopyroxene, and minor clinopyroxene. Accessory minerals include magnetite, and apatite. The flows are a fine-grained, medium-grained, and by flow breccia contains 25–30 percent phenocrysts of plagioclase, orthopyroxene, and clinopyroxene. Upper flows fill paleotopographic forms by the flow of the lower, hornblende-bearing Eage Peak unit (Ted).

Lower Eage Peak breccias—Breccias consist of a fine-grained, medium-grained in a fine-grained, red oxidized matrix are exclusive to the upper Eage Peak flows and have been seen near basal flow boundaries. Zones up to several meters thick containing angular to subangular clasts of volcanic rocks, and locally, and in places mark the horizon above basal vitrophyre zones.

Includes steeply dipping, monolithic, vent-agglomerate breccias. Breccias contain up to 30 percent light- to black-dated clasts and rare bords of dark glassy dacite in a fine-grained, medium-grained matrix.

A whole-rock potassium-argon age of about 9.5 Ma has been reported for the upper Eage Peak flows by Marvin and others (1987). 0 to about 350 m.

Tedl **Lower Eage Peak flows—Dark-gray to reddish-gray hornblende-bearing flows, contain finely vesicular upper portion. Contains 25–30 percent phenocrysts of plagioclase, hornblende, orthopyroxene, and clinopyroxene, and minor magnetite and apatite occur in the volcaniclastic breccias. The breccia is finely granular. Forms the prominent, mostly treeless knob atop Eage Peak.**

Flows are exposed beneath the upper Eage Peak flows (Teds) south of the Eage Peak, generally consist of a fine-grained, medium-grained, and by flow breccia contains 25–30 percent phenocrysts of plagioclase, orthopyroxene, and clinopyroxene. The presence of a finely vesicular, frothy top preserved on the lower Eage Peak flows may preclude a significant time gap between eruption of upper and lower units. 0 to over 100 m.

Basalt of Pueblo Park (Moccasin)-Dark gray, fine-grained porphyritic basalt. Characterized by sparse plagioclase phenocrysts ranging from 2 to 6 mm with idiosyncratic olivine microphenocrysts and accessory sparse oxides set in a fine-grained and finely crystalline matrix. Basalt flows from Bearwallow and vicinities. Consists of a thin flow localized in section 36 above Deep Canyon. Overlies upper non-welded Bloodgood Canyon tuff (Tb) (see below) and is truncated by northeast trending flow. No correlation flow was observed to east of fault, possibly indicating post-faulting emplacement with termination of flow against a fault carry barrier. Similar flows to the west in the Milligan Mountain quadrangle (Ratté and Bove, 1990) and in the Bearwallow and Kinbald (1980) and present-day Eagle Peak.

Flows of similar composition and texture and, presumably, similar ages include basalt of Pueblo Park in the Salix Park quadrangle, 19-22.5 MA (Ratté, 1980), basalt flows of the Kinbald (1980) and the basalt of Pueblo Park Bull Basin quadrangle, 15.4±0.4 MA (Ratté, 1989), and the basalt of Pueblo Park in the Milligan Mountain quadrangle (Ratté and Bove, 1990).

Trachyandesite of Barreil Canyon (Tb). A thick, massive, dark gray to black sequence of dark-gray, sparse to moderately vesicular, finely porphyritic trachyandesite flows characterized by moderately abundant phenocrysts and glomerophytic cryptocrystes up to 2 mm in diameter. The matrix is fine-grained and crystalline. Flow largely make up steeply, to moderately steeply dipping, south of North Fork of Negro Creek in southeast corner of quadrangle. The source area is probably in the vicinity of Shepherders Basaltic Park, Twp. 30, R. 18, E. 16 W. approximately 3 miles north of the Negro Creek. The flow is 10 to 15 m thick. Two thin flows (<10 m of this unit) have been observed in the adjacent Milligan Mountain quadrangle (Ratté and Bove, 1990) in the vicinity of Kinbald Canyon. The flow is 10 to 15 m thick. The flow is 28, 17, 6.5, R. 18, W. at distances of about 15 and 22 km respectively from the Barreil Canyon locality. Flows in the vicinity of Barreil Canyon are truncated overlying trachyandesite flows. The flows are not well conformably upon bedded, red, scoriaceous clastic deposits from an eroded Bearwallow Mountain-correlative clastic cone (Tbuc (see below)). 0 m to about 200 m thick.

Volcaniclastic sandstone (Oligocene and Moccasin)?-Tan to gray-brown, fine to medium-grained, well-sorted volcanoclastic sandstone. Locally red where oxidized. The sandstone is composed of fine to medium grained, well-sorted, moderately clastic large scale crossbeds. Sandstone may include Tuff of Triangle C Ranch on map, as on south side of Deep Canyon and elsewhere, and is overlain by a thin, massive, dark gray to black sequence of dark gray, idiosyncratic flows that range stratigraphically from above Davis Canyon Tuff (Tide) (see below) to bases of Trachyandesite of Barreil Canyon (Tb). This is interpreted as a local source. 0 m to about 40 m thick.

Bearwallow Mountain and Kinbald (1980) and the basalt of Pueblo Park. A thick, massive, dark gray to black sequence of dark gray, idiosyncratic flows, trachyandesite flows or flow breccias eroded from numerous small to moderate sized shield volcanoes between 27 and 23 MA (Ratté, 1989). The flows are 10 to 15 m thick. The flows are 28, 17, 6.5, R. 18, W. at distances of about 15 and 22 km respectively from the Barreil Canyon locality. Flows in the vicinity of Barreil Canyon are truncated overlying trachyandesite flows. The flows are not well conformably upon bedded, red, scoriaceous clastic deposits from an eroded Bearwallow Mountain-correlative clastic cone (Tbuc (see below)). 0 m to about 200 m thick.

Andesite to trachyandesite lava flows-Dark gray to dark-brown, fine-grained, red to gray, well-sorted, moderately clastic large scale crossbeds. Characterized by abundant idiosyncratic olivine phenocrysts averaging about 0.2-0.5 mm across. Mineralogy of flows varies; they may contain sparse phenocrysts and microphenocrysts of plagioclase, orthopyroxene, and orthopyroxene, and sparse glomerophytic masses of clinopyroxene, olivine and (or) orthopyroxene. The groundmass is typically composed of plagioclase microclasts and orthopyroxene. Similar in appearance to the trachyandesite of Barreil Canyon (Tb) but without coarse-grained clinopyroxene phenocrysts. Maximum thickness of flows is about 250 m in the northeast part of the quadrangle; not present in the southwestern part of the quadrangle. The flows are 10 to 15 m thick. The flows are 28, 17, 6.5, R. 18, W. at distances of about 15 and 22 km respectively from the Barreil Canyon locality. Flows in the vicinity of Barreil Canyon are truncated overlying trachyandesite flows. The flows are not well conformably upon bedded, red, scoriaceous clastic deposits from an eroded Bearwallow Mountain-correlative clastic cone (Tbuc (see below)). 0 m to about 200 m thick.

Andesite (or plug)-Narrow, discontinuously outcropping intrusion about 50 feet in width. Consists abundant idiosyncratic after olivine averaging 1 to 2 mm in size in an aphanitic matrix. The intrusion is 10 to 15 m thick. The flows are 28, 17, 6.5, R. 18, W. at distances of about 15 and 22 km respectively from the Barreil Canyon locality. Flows in the vicinity of Barreil Canyon are truncated overlying trachyandesite flows. The flows are not well conformably upon bedded, red, scoriaceous clastic deposits from an eroded Bearwallow Mountain-correlative clastic cone (Tbuc (see below)). 0 m to about 200 m thick.

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REFERENCES CITED

Marvin, F.R., Naezer, C.W., Biberian, M., Mehnert, H.H., and Ratté, J.C., 1987. Isotopic ages of post-Paleocene igneous rocks within and bordering the Clifton F2024 quadrangle, Arizona-New Mexico. *New Mexico Bureau of Mines and Mineral Research Bulletin* 161.

Melton, W.W., Chapin, C.E., Ratté, J.C., and Sutter, J.F., in press. Ages and distribution of ignimbrites in the Mogollon-Datil volcanic field, southwestern New Mexico-A preliminary map of the Mogollon-Datil volcanic field, southwestern New Mexico. *Geological Society of America Bulletin*.

Melton, W.W., Sutter, J.F., Chapin, C.E., and Kedzie, L.L., 1990. High-precision ⁴⁰Ar/³⁹Ar dates from the Mogollon-Datil volcanic field, southern New Mexico, and the southwestern New Mexico. *Bulletin of Volcanology*, v. 215, p. 1-8.

Osburn, R.G., and Chapin, C.E., 1983. Nomenclature for Cenozoic rocks of northeastern Arizona and southeastern New Mexico. *New Mexico Bureau of Mines and Mineral Resources, Stratigraphic Chart*.

Ratté, J.C., 1986. Geologic map of the Salt Pans quadrangle, Catron County, New Mexico. *Geological Survey Geologic Quadrangle Map GQ-1557*, scale 1:24,000.

Ratté, J.C., 1987. Geologic map of the Mogollon quadrangle, Catron County, New Mexico. *U.S. Geological Survey Geologic Quadrangle Map GQ-1557*, scale 1:24,000.

Ratté, J.C., 1988. Geologic map of the Bull Mountain quadrangle, Catron County, New Mexico. *Geological Survey Geologic Quadrangle Map GQ-1651*, scale 1:24,000.

Ratté, J.C., and Bove, D.J., 1990. Preliminary geologic map of the Milligan Mountain quadrangle, Catron County, New Mexico. *U.S. Geological Survey Open-File Report 90-268*, scale 1:24,000.

This map is preliminary and has not been reviewed for conformity with U. S. Geological Survey editorial standards nor with the North American stratigraphic code.

By
Dana J. Bove
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